

1 **Response to comments posted 11 August 2024 by Referee #1 to “DALROMS-NWA12 v1.0, a**  
2 **coupled circulation-ice-biogeochemistry modelling system for the northwest Atlantic Ocean:**  
3 **Development and validation” (28 August 2024)**  
4

5 We thank the referee for taking the time to review our manuscript and for providing supportive  
6 comments. Please find below our response to your comments.

7  
8 *(...) Although the short simulation period of 4 years may not be useful for climate research, I*  
9 *suggest the authors select specific locations to demonstrate whether the models can properly*  
10 *simulate the evolution of key variables (monthly mean or annual mean). Numerous observed sea*  
11 *surface temperature data, bottom temperature data (summertime), current meter data, and other*  
12 *types of data are available for this region, which could be used to further investigate model*  
13 *performance. This would enhance the confidence of readers and future users in the model.*  
14 *Additionally, winter convection events in the Labrador Sea are believed to be a significant driver of*  
15 *circulation in the North Atlantic Ocean, impacting water mass properties in this region. It is worth*  
16 *reporting whether the ROMS model can well simulate these winter convection events during the*  
17 *study period.*

18  
19 We agree with you that four-year simulations are too short for climate research. The main focus of  
20 this paper, however, is on the development and validation of an advanced coupled physical-  
21 biogeological modelling system for the northwest Atlantic Ocean. In our future studies, we will run  
22 this modelling system for much longer simulation periods to examine the effects of climate change  
23 on the marine conditions over the region. This sentence was added to the last paragraph of section  
24 5. We also examined the model performance in simulating winter convection in the new Figure 16  
25 (discussed in pages 30 and 32–33 of the revised manuscript), which shows daily-mean temperatures  
26 simulated in the prognostic simulation, vertically interpolated to 5-m depth intervals between the 0-  
27 and 2000-m depths, at two locations in the Labrador Sea: one near the centre of the “convective  
28 region” identified by Luo et al. (2014), and the other on the AR7W transect and within the area for  
29 which Yashayaev (2024) composited available observations to construct his time series of depth  
30 profiles. Although the horizontal grid size of our model ( $O(\text{km})$ ) is too coarse to resolve convective  
31 plumes ( $O(100 \text{ m})$ ) and the model does not include any parameterization of convection, the  
32 response of the model’s turbulent vertical mixing scheme to the conditions that trigger convection  
33 can be expected to generate conditions similar to those resulting from convection. At the location

34 within the “convective region”, the time series of simulated temperature profiles includes several  
35 features that appear in Yashayaev’s observation-based time series, such as the contrast between  
36 weak and strong convections in 2014 and 2015 respectively. In the future, we hope to develop a  
37 finer-resolution model focusing on the Labrador Sea which will be nested within our current model  
38 and will include the capability to simulate or parameterize convection.

39

40 *Below is a list of specific comments.*

- 41 • *Line 34 “up to 7”, could “up to 7 psu” be better”?*

42

43 We agree that “up to 7” sounds somewhat awkward. Modifications were made accordingly.

44

- 45 • *Lines 153-154. Why are the eddy viscosity and diffusivity set to zero? Given that you employ*  
46 *high-order accuracy schemes (3rd and 4th order), it’s unclear why these values would be zero.*  
47 *Could you provide a rationale for this choice?*

48

49 The reason for setting the horizontal eddy diffusivity to zero is that the third-order upstream  
50 advection scheme generates some numerical diffusion which is large enough to eliminate numerical  
51 noise in tracers. As we discuss in lines 525–535 (page 19) of the revised manuscript, this numerical  
52 diffusion associated with the advection scheme already appears to cause excessive mixing in areas  
53 such as the St. Lawrence Estuary-Gulf system. We set the horizontal eddy viscosity to zero in order  
54 to preserve as much as possible of the simulated eddy activity and because we have not noticed any  
55 numerical noise in the simulated circulation field that needs to be smoothed out with a non-zero  
56 horizontal eddy viscosity. Text was added on Page 7 accordingly.

57

- 58 • *Lines 172-174. Are the inter-annually variable GLORYS12 data used at open boundaries or the*  
59 *daily climatological ones? Please specify.*

60

61 It is the former. We use daily GLORYS values with interannual variability for the simulation period  
62 and not climatologies. Text was changed to “derived from the daily fields of Copernicus global  
63 1/12° oceanic and sea ice reanalysis (GLORYS12V1, hereafter GLORYS; Lellouche et al., 2021)  
64 for the simulation period” in lines 245–246 (page 8) of the revised manuscript.

65

66 • *Lines 182-183. I am curious why it is through the bottom. River discharge is through top layers.*  
67

68 In previous versions of ROMS, riverine freshwater input was specified as horizontal flow into the  
69 river mouth grid cell from a neighbouring dry grid cell. When the riverine input is specified through  
70 the bottom of a wet grid cell, the river scheme has several advantages including: 1) the model is  
71 more stable because there is no longer an addition of horizontal momentum and 2) the user no  
72 longer needs to worry about specifying the grid cell from which the fresh water enters the river  
73 head, which can be tedious when the domain includes a large number of rivers flowing in different  
74 directions. Details of this method are described by its developers in this ROMS user forum post:  
75 <https://www.myroms.org/forum/viewtopic.php?t=5156>. This link was added to line 254 on page 8.

76

77 • *Lines 298-299. Glorys12 uses data assimilation (4d var), and in theory, this can eliminate those*  
78 *biases from the non-inclusion of tides.*

79

80 We agree with you that, in theory, the use of data-assimilation in generating GLORY12 can  
81 eliminate biases associated with non-inclusion of tides. But in reality observations used in data  
82 assimilation are very sparse in time and space. Therefore it is highly unrealistic to expect GLORYS  
83 to be an exact reproduction of real ocean conditions at all locations. In areas and during periods for  
84 which the number of observations are relatively small, exclusion of tides can cause GLORYS  
85 values in the subsurface zone of relatively shallow waters to deviate from true solutions. Text was  
86 added accordingly in lines 378–382 (page 15) of the revised manuscript.

87

88 • *Lines 310-311. Same as above. You may want to find some literature to support this statement.*

89

90 Here we are merely speculating, based on the fact that error metrics tend to be large: 1) near the  
91 model's lateral open boundaries, 2) during the summer when temperatures are not constrained by  
92 sea ice, and 3) at the surface, where the metrics are calculated with respect to an independent  
93 dataset, that the errors in GLORYS (which exist even after the assimilation of available  
94 observations) might be propagating into the model domain via lateral open boundary inputs.

95

96

- 97 • Lines 369-378. This seems an indication of numerical scheme issue or the horizontal mixing  
98 issue (zero is used).

99

100 The third-order upstream and fourth-order Akima horizontal tracer advection schemes that we  
101 tested have their advantages and disadvantages. The third-order scheme is much less prone than  
102 fourth-order schemes to numerical under- and over-shootings, but it generates relatively large  
103 numerical diffusion (even with zero eddy diffusivity) that can result in excessively smooth  
104 simulated tracer fields. The fourth-order scheme performed better at reproducing the general three-  
105 dimensional structures of temperature and salinity in areas such as the St. Lawrence Estuary, which  
106 in turn led to more realistic simulations of sea ice. This scheme, however, is prone to numerical  
107 under- and over-shootings that produced patches of unrealistic tracer values, which is why we  
108 decided to use the third-order scheme. The topic of numerical schemes, including the possible  
109 development of a flux limiter for a fourth-order advection scheme, is something we hope to revisit  
110 in the future. Text in lines 448–458 (page 19) of the revised manuscript was modified accordingly.

111

- 112 • *Lines 409-425. Data from Drinkwater (1988) were for the year of 1982, and the data your*  
113 *model and GLORYS12 are for the recent years. You need to mention the probable existence of*  
114 *decadal or even longer variability for the current in this area.*

115

116 We agree. We added “Although we need to keep in mind the existence of interannual variability  
117 and long-term trends which limit the conclusions we can derive” to the beginning of the sentence in  
118 which we point to the possibility that the inclusion of tides in our model results in a more realistic  
119 vertical structure of currents (lines 522–524, page 23 of revised manuscript).

120

- 121 • *Figure 13. I cannot see any dots there. Could the quality of the figure need improving?*

122

123 The “dots” referred to the straight black line at the sea surface. We have revised the caption to  
124 clarify what the lines represent.