

5 **Initial Author Response for “Multi-physics ensemble modelling of Arctic tundra snowpack properties”, Woolley et al.**

10 The authors would like to thank the editor and both reviewers for the time taken to read and review the original manuscript and for the great feedback provided which improves the manuscript. Our proposed initial responses are in blue, modified text that we will add to the revised document is in italics and reviewer comments are in black.

**Reply to comments from Reviewer 1:**

The font size on the figures (legends, axis labels, etc) throughout the paper should be increased to make them readable in print (including Figure 9).

We will increase the font size.

15 The author report that the vegetation effect does not allow for improved simulations compared to the standard crocus, and that this highlights the need to account for water vapor transport in snow. However, looking at the density profiles (fig 4 and 6), simulations do show a drop of density at the base. This drop seems of the same order as the ones reported in measurements, when observed. As you mentioned, the definition of what is the wind slab and what is the depth hoar could impact the conclusion of the statistical analysis. The simulated density drop indeed seems to impact less than the 40 to 70 % of the profile, as indicated as the range of DHF reported from the snow pits. An idea could be to compare bulk density of only the first 10 cm (or any other relevant value that for sure describe only depth hoar, even if not part of the depth hoar layer is not included). **Essentially, my question is to what extent the error in defining the depth hoar boundaries could have an impact on the conclusion regarding the proposed vegetation parameterizations. Some more comments could be done on that in the discussion.**

20 We will follow the suggestion of the reviewer and evaluate the first 10 cm of each simulated and measured profile. As measurements are not available for the base of the snowpack, we will only compare the first 10 cm of each profile where measurements are available. We will identify the depth in which the measured values begin and remove any simulated values that lie below this. We will then extract the first 10 cm of values from this point for both the measured and simulated profiles and carry out our statistical analysis. Initial statistical scores from this analysis are presented below:

35 *Table 1: Mean, RMSE, SS and CRPS scores for measured and simulated snow density (kg m<sup>-3</sup>) for the bottom 10 cm (starting where measurement profiles begin) for the March 2018, March 2019, March 2022 and March 2023 snow seasons.*

		<b>Mean</b>	<b>RMSE</b>	<b>SS</b>	<b>CRPS</b>
<b>Density (kg m<sup>-3</sup>)</b>	Measured	234	-	-	-
	Default	277	69	0.6	51
	Arctic	262	79	1.1	35

40 These results suggest that Arctic SVS2-Crocus is able to simulate a mean value for snow density that better matches that of measurements (by 15 kg m<sup>-3</sup>) for the first 10 cm of the snowpack. However, we find the error in the simulation is higher than default SVS2-Crocus (default RMSE: 69 kg m<sup>-3</sup>; Arctic RMSE: 79 kg m<sup>-3</sup>). A better mean but higher RMSE indicates that while the average value from the Arctic SVS2-Crocus ensemble is closer to the average value from the measurements, the ensemble has larger deviations from the measurements than default SVS2-Crocus. This is reflected in the higher SS score for Arctic SVS2-Crocus (default SS: 0.6; Arctic: 1.1) and is due to the design of the *Basal Vegetation Effect*

50 modifications. We investigate R2V and R2D individually and then as a combination as R21 (described in section 3.2.3). This produces a large ensemble spread, as R2V and R21 are able to simulate basal densities that are representative of the measurements, whereas the modification of R2D is not as effective. As we highlight in section 5.2 of the manuscript, we determine that the use of modification R21 due to its high relative occurrence in both the mixed and Arctic ensembles, can reduce basal layer densities that better match measured results, in comparison to default SVS2-Crocus. We will carry out statistical analysis on the performance of the individual modifications (R21, R2V and R2D) to support this statement.

55 We will also add further information to section 3.0 Data and Methods for determining the selection of DHF boundaries. The information included will be as follows:

- 60 1. We plot each measured profile and identify transitions in the density and/or SSA. The SSA for different layers is often more distinct than density (Rutter et al., 2019) which provides a sharper transition between wind slab and depth hoar that can be visibly identified.
2. Where the transition occurs, we note the density or SSA value.
3. We cross reference the value with those presented in Fig. 9 of Rutter et al. (2019) to ensure we classify the area of the snowpack with the correct snow type.

### Specific Comments:

65 Introduction

The second paragraph of the introduction (lines 49 – 67) focuses on the limitations of arctic snow modeling due to misrepresentation of some physical processes, not suited for the Arctic. This part provides a state of the art which is not fully convincing because some information are lacking and/or disorderly. For example, it is not mentioned what is the current issues regarding wind parameterization (is it too weak or too strong?), how is modeled the vegetation (is it accounted at all?) or why should we consider a different thermal conductivity. The process of water vapor transport is well described but is actually not essential as it is not addressed in this paper (not modeled). Most of the missing information are provided latter in the paper but should be described already in the state of the art so that the reader understands the motivation for the work done. This is why I suggest that, for each of the processes addressed in this paper, which are the effect of wind, the effect of vegetation, and the snow thermal conductivity, you check that the following information are provided in a well-structured way:

- 70 - description of the physical processes and the consequences on snow and what is different in the arctic
- 80 - how it is modeled or not in “standard” snow model (crocus, snowpack) and what are the errors done by applying it to the arctic (quantify if possible)
- if any, what are the modifications for the arctic already presented and how do they improve the situation, what is left to be addressed (contributions of this paper).

85 We will modify the second paragraph of the introduction to include the relevant and requested information. We will split this into the following paragraphs:

90 *Detailed multi-layered snowpack models Crocus (Vionnet et al., 2012) and SNOWPACK (Bartelt and Lehning, 2002) suffer from several weaknesses when applied within Arctic environments due to the misrepresentation and lack of consideration for many Arctic processes (Domine et al., 2019; Fourteau et al., 2021; Barrere et al., 2017). Despite showing reasonable agreement in their simulation of snow depth and SWE of Arctic snowpacks*

95 *(Barrere et al., 2017; Gouttevin et al., 2018; Krinner et al., 2018; Domine et al., 2019; Royer et al., 2021; Krampe et al., 2021; Lackner et al., 2022), both models simulate profiles of increasing density with snow depth because both Crocus and SNOWPACK were originally developed for Alpine applications. Inaccurate simulation in snow density is further compounded by an underestimation of wind-induced compaction (Barrere et al., 2017; Royer et al., 2021; Lackner et al., 2022), misrepresentation of basal vegetation influencing compaction and metamorphism (Gouttevin et al., 2018; Royer et al., 2021), uncertainties in thermal conductivity formulations (Royer et al., 2021; Dutch et al., 2022), and omission of water vapour flux transport (Brondex et al., 2023).*

100 *Within the Arctic, high wind speeds compact the surface of the snowpack leading to the development of high-density wind slab snow layers (King et al., 2020; Derksen et al., 2014). The effect of wind is underestimated within current versions of Crocus leading to underestimations in simulated surface snow density (Barrere et al., 2017). Attempts to account for an underestimation in wind speed have been proposed by Barrere et al. (2017) where wind speed was increased within the falling snow density and snow compaction equations. Barrere et al. (2017) and Royer et al. (2021) also deemed the maximum density constraint ( $350 \text{ kg m}^{-3}$ ) to be too low for Arctic applications and was raised (to  $600 \text{ kg m}^{-3}$ ) to allow further compaction to occur.*

110 *Basal vegetation (shrubs and sedges) modifies the temperature gradient through trapping effects by reducing compaction and enhancing snow metamorphism, which promotes further depth hoar formation (Domine et al., 2016; Domine et al., 2022). The ability of basal vegetation to promote the development of depth hoar is currently not considered within Crocus or SNOWPACK where compaction due to the weight of the overlying snow is the dominant process in shaping density profiles (Vionnet et al., 2012; Bartelt and Lehning, 2002). To consider the vegetation trapping effect, Gouttevin et al. (2018) and Royer et al. (2021) proposed to switch of snowdrift and increase snow viscosity below a set vegetation height which contributed towards density reduction and enhanced grain growth in basal layers.*

120 *Thermal conductivity of snow is often computed as a function of density within many snowpack models (Gouttevin et al., 2018), with a number of different relationships proposed (Yen, 1981; Calonne et al., 2011; Sturm et al., 1997). The parameterisation of Sturm et al. (1997) has been found to produce better results for Arctic snow than that of the default parameterisation of Yen (1981), due to its development on Arctic and sub-Arctic snow, and has recently been implemented into Crocus (Royer et al., 2021; Calonne et al., 2011). The formulations of Calonne et al. (2011), who use 3D tomographic images of most snow types, and Fourteau et al. (2021), who propose a formulation suitable for temperatures within Arctic snowpacks have also been found to improve the simulation of snow thermal conductivity at an Arctic site (Dutch et al., 2022). Calonne et al. (2011) is available for use within the ensemble system version of Crocus (ES-CROC) however the parameterisation of Fourteau et al. (2021) is yet to be implemented within Crocus.*

130 *Strong temperature gradients within an Arctic snowpack generate water vapour flux transport that redistributes mass from the bottom to the top of the snowpack, leading to the formation of low-density basal depth hoar layers (Bouvet et al., 2023; Weise, 2017). Attempts have been made to implement water vapour diffusion into Crocus (Touzeau et al., 2018), SNOWPACK (Jafari et al., 2020) and SNTHERM (Jordan, 1991), however no approach was successful in accounting for all aspects of vapour diffusion or able to be numerically stable at the typical time steps of snowpack models, and is therefore currently not included within any of these models (Brondex et al., 2023).*

The third paragraph of the introduction (line 69 -83) focuses on another issue on arctic snow modeling, which, this time, concerns the model evaluation method. If this is correct, I suggest

140 to change the first sentence of this paragraph so that this new topic is introduced. The sentence in the current version of the paper refers to issues related to physical processes, so to the topic of to the previous paragraph. It could start as “Another limitation in Arctic snow modeling concerns the method for model evaluation. Indeed, previous evaluation of simulated arctic snow density (e.g. Gouttevin ...) neglect uncertainties that arise from ...”

145 We will modify the first sentence of the third paragraph to introduce uncertainties in snowpack modelling:

150 *Snowpack modelling contains parameter uncertainties that can be quantified using an ensemble approach (Lafaysse et al., 2017). Previous evaluations of simulated Arctic snow density profiles focus on individual modifications to existing snow physical processes that account for high wind speeds, the presence of basal vegetation, and/or better simulations of snow thermal conductivity (Barrere et al., 2017; Lackner et al., 2022; Royer et al., 2021; Gouttevin et al., 2018). The uncertainties that arise from the interaction between model components, site specific calibration of parameter choice and limited evaluation datasets (one site, few snow seasons) – see e.g. Gouttevin et al. (2018) and Barrere et al. (2017) are therefore not considered through this approach.*

Line 56: “strong temperature gradients generate water vapour flux transport that redistributes mass from the bottom to the top of the snowpack, leading to the formation of low-density basal depth hoar layers (Domine et al., 2016b; Fourteau et al., 2021)”. Citations here should refer to observation, Fourteau et al 2021 did modeling, if my not mistaking. Citation of the experimental work of Weise’s thesis (in Chap.5) and Bouvet et al. 2023 could be included.

160 We will change the references to include the experimental work of Bouvet et al. (2023) and Weise (2017):

165 *strong temperature gradients generate water vapour flux transport that redistributes mass from the bottom to the top of the snowpack, leading to the formation of low-density basal depth hoar layers (Bouvet et al., 2023; Weise, 2017).*

Line 65: Is “Domine et al 2016b and Domine et al. 2019” the reference papers to describe wind effect on arctic snow? Papers that describe this process should be given here (maybe the references provided line 210).

170 We will change the references to include the work of King et al. (2020) and Derksen et al. (2014):

*Furthermore, high Arctic wind speeds compact the surface of the snowpack leading to the development of high-density wind slab snow layers (King et al., 2020; Derksen et al., 2014).*

175 Line 65 “Attempts to account for missing processes that specifically impact Arctic snowpack properties have been made by implementing simplified adaptations to existing snow physical processes (Gouttevin et al., 2018; Royer et al., 2021; Barrere et al., 2017; Lackner et al., 2022).” This sentence should be more explicit or deleted here.

We will delete the sentence but will modify the second sentence of the third paragraph to provide the reader with the same information. This is shown in detail in our response to the review comment on line 145 of this document:

180 *Previous attempts to simulate Arctic snow density profiles focus on individual modifications to existing snow physical processes that account for high wind speeds, the presence of basal vegetation, and/or better simulations of snow thermal conductivity (Barrere et al., 2017; Lackner et al., 2022; Royer et al., 2021; Gouttevin et al., 2018).*

185 Line 87: It should be mentioned in this paragraph if the parameterizations implemented in crocus are new or from the literature.

We will modify the sentence on line 87 to state that the parameterizations implemented are from literature:

190 *This study uses the multi-physics ensemble version of Crocus (Lafaysse et al., 2017; Vionnet et al., 2012) embedded within the Soil, Vegetation and Snow version 2 (SVS2) land surface model (hereby referred to as SVS2-Crocus, Garnaud et al., 2019; Vionnet et al., 2022) to evaluate the impact on simulated Arctic snowpack properties by modifying parameterisations of falling snow density, snowdrift, compaction and thermal conductivity that have been proposed within previous literature.*

Line 103: how is the topography of the study site?

195 We will add the following sentence to describe the topography at Trail Valley Creek:

*The terrain consists of mineral earth hummocks that range in diameter between 0.4 to 1.0 m and inter-hummock areas of peat (Quinton and Marsh, 1999).*

200 Line 223: "Following the approach of Gouttevin et al. (2018), Royer et al. (2021) and Domine et al. (2016a) deactivated wind compaction and increased  $\eta$  under a set vegetation height." → What was the improvement of this modification.

We will modify the following sentence, to outline the improvement of the modification:

*Following the approach of Domine et al. (2016), Gouttevin et al. (2018) and Royer et al. (2021) deactivated wind compaction and increased  $\eta$  under a set vegetation height which reduced the rate of densification through compaction processes.*

205 Line 225: "Both options" → It is not clear which options are meant here.

We will change line 225 to:

*Modifications R2D and R2V are also investigated together as R21.*

Line 259: "... and are applied to the normalized profiles of simulated density and SSA" → describe what are the normalized profiles.

210 We will add the following sentence to explain the normalized profiles:

*Measured and simulated density and SSA profiles report different depths and vertical resolutions; therefore, we rescale each individual profile to a 0.005 m grid interpolated using layer thickness, where all profiles begin at 0 m and end at 1 m.*

215 Line 260: "Vegetation in the base of an Arctic snowpack makes density and IceCube measurements difficult meaning measurements do not always reach the base of the snowpack for evaluation of simulated basal layer density and SSA." → to be reformulated "which might impact the evaluation of simulated basal layer density and SSA."

We will reword line 260 to:

220 *Vegetation in the base of an Arctic snowpack makes density and IceCube measurements difficult meaning measurements do not always reach the base of the snowpack which may impact the evaluation of simulated basal layer density and SSA.*

Line 270: The start of the paragraph should introduce what do we look at now, instead of going directly into details. I would suggest something like "Over the years 1991-2023, different result

225 can be found in the evolution of snow depth and SWE over the course of the winter between  
the model and the measurements. Over estimation, good agreement and under estimation  
can be observed depending on the year considered, as illustrated in Figure 2. These biases  
can be explained ... ect”.

We will add a new sentence to introduce the paragraph:

230 Differences in the seasonal evolution of simulated and measured snow depth, SWE and bulk  
density can be found over the 1991-2023 period. Model over-estimation, good model  
agreement and model under-estimation in simulated snow depth and SWE are observed when  
compared to measurements, depending on the year considered, as illustrated in Fig. 2.

Line 277: “In this case, a snow drift in the SR50 footprint can lead to exaggerated differences  
235 between simulated and measured snow depth.” → Is this feature was observed during the  
field campaign or is it an hypothesis?

This is a feature observed during the field campaigns. We will add a sentence that explains  
why the surrounding topography, wind direction and timing of precipitation leads to a snowdrift  
in the footprint of the SR50 sensor.

Line 287: Here the data that we look at now should be introduced first, otherwise it is unclear.  
240 Such as “We look now at the statistical scores when comparing model and measurements  
over the entire 1991-2023 period at the time of the snow course measurement, i.e. around the  
peak SWE accumulation.”

We will add a new sentence to introduce the data:

245 *We now analyse the statistical scores computed when comparing default and Arctic SVS2-  
Crocus with measurements at the time of peak SWE accumulation over the entire 1991-2023  
period.*

Line 289: “Deeper snow depths are simulated by default SVS2-Crocus (default mean: 0.54 m;  
Arctic mean: 0.47 m) due to the Wind Effect modifications applied to Arctic SVS2-Crocus  
250 resulting in increased density in the surface layers of the snowpack, leading to higher bulk  
density (default mean: 239 kg m<sup>-3</sup>; Arctic mean: 278 kg m<sup>-3</sup>; Table 1, Appendix B3) and  
shallower snow depths.” → This sentence should be reformulated.

We will reformulate the sentence on line 289 to:

255 *Wind Effect modifications applied to Arctic SVS2-Crocus increase surface layer density  
leading to a higher bulk density (default mean: 239 kg m<sup>-3</sup>; Arctic mean: 278 kg m<sup>-3</sup>; Table 1,  
Appendix B3) and shallower snow depths (default mean: 0.54 m; Arctic mean: 0.47 m) than  
default SVS2-Crocus.*

The first paragraph is made of an overview of the snowpack structure at TVC. Then details on  
the dataset are provided and seem out of context. They seem to come to early in the result,  
260 before an overview or a general description of the dataset is provided (following paragraphs),  
which make it difficult to follow. I would suggest to move them lower down in the paper.

We will restructure this section to include an initial paragraph that outlines the dataset  
presented:

265 *Evaluation of the development of snow density and SSA over a winter season in the Arctic is  
a unique opportunity that differs from the traditional snow model evaluation methods of utilising  
measurements from March, April or May (Barrere et al., 2017; Gouttevin et al., 2018; Royer et  
al., 2021; Domine et al., 2019). Figures 4 and 5 compare measured and simulated density and*

SSA by default and Arctic SVS2-Crocus over the 2018/19 winter using in-situ measurements from November 2018, January 2019 and March 2019. Figures 6 and 7 show measured and simulated snow density and SSA across four winter seasons for a March snowpack.

270 We will also revise the leading sentences to introduce the following two paragraphs:

Line 356: *Over the course of the 2018/19 winter season, default SVS2-Crocus simulated a snowpack subject to consistent compaction, with basal layers increasing in density from ~200 kg m<sup>-3</sup> in November 2018 to ~300 kg m<sup>-3</sup> in March 2019 (Fig. 4).*

275 *Line 369: Across the four winter seasons for a March snowpack, the dominance of compaction is clear when using default SVS2-Crocus where the ensemble simulated high-density basal layers (default mean DHF: 268 kg m<sup>-3</sup>) overlain with lower density surface layers (default mean WS: 177 kg m<sup>-3</sup>) (Table 2, Fig. 6) across each year.*

280 Line 232: “November 2018 shows less variability and range in snow density than other snow seasons (Fig. 4) as the snowpack was shallow and metamorphism in basal layers and compaction in surface layers had little time to affect the density.” → unclear. Do you compare snow in November with snow in later months? in early season, snow had less time to evolve? Or does the comment refer to the year 2018 which was special? Again, this specific comment is hard to follow, as the reader is not yet familiar with the data / figures, which are actually described lower down.

285 We agree this sentence may lead to some confusion. We compare the structure of the density profile with other snow seasons to help describe how this evolves. Along with the restructure outlined in our response to the previous comment, we will add a new sentence to paragraph 2 of section 4.2 to give the reader a description of the data about to be discussed:

290 *We first analyse the measured profiles of density at TVC across the 2018/19 winter and the four winter seasons for a March snowpack.*

We will then elaborate on our description and comparison of the November 2018 profile with other snow seasons as follows:

295 *The density profile from November 2018 was measured early in the snow season and shows less variability and range than other snow seasons (Fig. 4) as the snowpack has had less time to evolve. The snowpack was shallow and metamorphism in basal layers and compaction in surface layers had little time to affect the density.*

300 Line 235: “Heightened variability in the density of the top 20% of the January 2019, March 2019 and March 2022 snowpacks was more pronounced than in other winter seasons due to the timing of sampling relative to a fresh snowfall event”. → unclear. Do you mean that the variability observed in snow density at the top 20% at these dates could have been introduced by a changing snowpack during the measurements, as they were performed during a snow fall.

305 *The heightened variability in density occurs in the top 20% of the January 2019, March 2019 and March 2022 simulated profiles and is not captured by measurements. We agree that this sentence is unclear and leads to confusion. Analysis of the meteorological forcing data shows that fresh snowfall occurred relative to this snapshot. It is likely that the measurements technique performed is not of a high enough resolution to capture small layers of fresh, low-density snowfall. We will modify this sentence for clarity.*

310 In the paragraph line 357 – 365, the rather sharp drop of density reported in the simulated profiles with the arctic version of crocus needs to be described in the text, as it seems an

important feature of the modeling. It differs strongly with the standard crocus. The fact that this sharp drop comes from the modified parameterizations of the vegetation effect (if I'm not mistaken) could also be pointed out.

We will modify the paragraph to include the following sentence:

315 *As the season progresses, and snow depth increases, the Basal Vegetation Effect modifications counteract the dominance of compaction found within default SVS2-Crocus and lead to a sharp drop in simulated density (reduction of  $\sim 50 \text{ kg m}^{-3}$  in November 2018). This decrease in density is retained within the snowpack over the entire winter season, with a greater reduction of  $\sim 150 \text{ kg m}^{-3}$  simulated by March 2019.*

320 Line 373: "density" → mean density

We will change this.

325 Line 398: "The Arctic SVS2-Crocus ensemble was however more skilled at capturing the variability in measurements" → It is not clear to me if this comment is relevant. In which way the arctic crocus would be more tuned / suited to capture spatial variability. Which introduced parameterizations would allow for that. If so, it should be described here.

We agree that this comment is not relevant and will remove from the manuscript.

330 Line 427 : "As these modifications were developed to consider Arctic processes, it is likely that they are better at simulating physical processes that occur in the Arctic environment over the default parameterisations, leading to lower CRPS scores." → to be deleted, it was shown and quantified in the sentence above.

We will delete this sentence.

Line 446: "Snowdrift parameterisations implemented into Arctic SVS2-Crocus modify the microstructure of snow grains during blowing snow events, which occur frequently at TVC." → should this sentence be in the discussion?

335 We agree that this sentence should be within the discussion, and we will move it to within paragraph 4 of section 5.2 as follows:

340 *Arctic modifications R21F, R21W and R21R are dominant parameterisations within the snowdrift scheme that lead to lower CRPS scores for the simulation of SSA as they work to modify the microstructure of snow grains during blowing snow events, which occur frequently at TVC.*

Line 504: "For the same reason, basal densities using default SVS2-Crocus may be underestimated." → please provide more explanation of this link.

345 Through this sentence we intended to state that the statistical scores using default SVS2-Crocus are underestimated, due to the lack of measurements in the base of the snowpack. As default SVS2-Crocus is subject to densification, leading to higher basal densities, the CRPS score for the DHF would in fact be higher than what is shown within this study. We will elaborate on the above sentence to provide more detail on why the basal densities using default SVS2-Crocus may be underestimated.

Line 520 : "The Basal vegetation effect is ..." → "In this case the Basal vegetation effect is ..."

350 We will change this sentence to:



*In this case, the Basal Vegetation Effect is activated immediately, causing compaction to occur at a very low rate where low basal densities are then retained within the snowpack throughout the entire winter.*

**Conclusion:**

355 I would suggest to include a comment on the potential to improve arctic model evaluation by using dataset that allow to better capture snow properties at and near the base, especially to evaluate the parameterizations of the vegetation effect. Your study pointed out the challenge (but the need) of having density and SSA data near the ground due to the measurement method used.

360 We will add the following sentence to the first paragraph of the conclusion outlining how an improved dataset would allow better evaluation of the parameterisations of the vegetation effect:

365 *The ability to evaluate the simulation of microstructure properties at the base of the snowpack and the performance of the Basal Vegetation Effect parameterisations would benefit from the use measurements from the snow micropenetrometer (SMP) (Johnson and Schneebeli, 1999), that is not hindered by the presence of basal vegetation and can reach the base of the snowpack.*

**Reply to comments from Reviewer 2:**

370 In the introduction, you mention the problems of site specific calibration of parameter choices, which you avoid here by working with a set of parameters calibrated for different sites. But that doesn't prevent your conclusions from being site specific. This is more of a suggestion than a comment, but other observations are available at other Arctic sites, for instance in Vargel et al. (2020), which would make it easy to elaborate on this point.

375 We understand our conclusions are currently specific to the site of Trail Valley Creek. The goal of this paper was to explore the effect and interaction of Arctic parameterisations and identify preferential combinations of parameters which we can now test elsewhere. We will look into the work of Vargel et al., 2020 to allow us to elaborate on this.

L109: How is the data filtered? Is the method known?

We will add the following sentence to include information about how the data is filtered:

380 *Depths below 0 m and above instrument sensor height (1.63 m) and abrupt jumps or spikes (negative or positive) that lie outside the reasonable range of values within the SR50A snow depth data were removed.*

L137: What is this most suitable option based on?

385 The temperature threshold was selected based upon qualitative assessment of the impact of temperature on precipitation partitioning at TVC (between 0°C and 5 °C) compared with observations of the precipitation type from the TVC Main Meteorological Station (TMM) and the adjacent Meteorological Service of Canada (MSC) weather station.

390 Figs. 4,5,6,7,8: Increase label size and legend and explain what is normalized depth in the text for better readability. Wouldn't it be more practical to display the observations and the model on the same graph.

We will increase the font size.

We will add a description of the normalized depth, this is outlined in our response to Reviewer 1 (Line 210 of this document).

395 Including both the model output and observations on the same graph made visual analysis of the shape and differences in profiles difficult. We decided to separate model output and observations to allow ease in the analysis by readers. For this reason, we will keep the figures as they appear in the preprint.

## References

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