

Response to egusphere-2024-123 RC1

Below is our response to RC1. Referee #1's comments are in **bold**. Our responses are in black. Changes to the manuscript are indicated in blue. The line numbers refer to the previously submitted manuscript without track changes.

In this manuscript, Li et al have implemented benthic denitrification using the empirical function derived by Bohlen et al (2012) on Uni-Vic ESM, and parameterized using chl, O₂, NO₃, PO₄, N* observations. Relatively speaking they found considering N* to constrain their model has given better model performance when compared to the observations from WOA13. They found that this implementation has resulted in better simulation of N₂ fixation patterns in the world ocean, and more realistic simulation of ODZs, suggesting that inclusion of benthic denitrification in ESMs are a key piece to improve model performance.

I found the procedure and inferences are sound. This work also highlights the importance of consideration of benthic processes in ESMs which are currently underrepresented, and its potential in improving model predictions. I have no major concerns with the present study; however, I feel the following comments will increase the readability of the paper.

We thank the reviewer for these very positive and helpful comments.

Minor comments:

Line 23: replace “however, even though” with in spite of.

NEW: “Research on benthic denitrification has been more limited, **in spite of** the greater part of global ocean denitrification ($\approx 60 - 75\%$) occurring in the sediments (Somes et al., 2013; DeVries et al., 2013; Eugster and Gruber, 2012; Brandes and Devol, 2002). “

Line 33: “..thus will be..”

NEW: “Reproducing realistic patterns and rates of denitrification **thus will be** an important aspect for simulating N₂ fixation.”

Line 35-39: maybe expand a bit on this, exactly what you are going to test?

OLD: “Contrasting these geochemical inversion estimations, we aim to test the ability of N* to constrain our prognostic model that includes the representation of large-scale stoichiometric diversity of phytoplankton and diazotrophs (Pahlow et al., 2020).”

NEW: “**We compare the ability of N* with that of NO₃⁻ and PO₄³⁻ to constrain parameters of our prognostic model**, that includes the representation of large-scale stoichiometric diversity of phytoplankton and diazotrophs (Pahlow et al., 2020).”

Line 55: “..description of the..”

NEW: “Below we provide a brief **description of** the original OPEM implementation in UVic 2.9 (Pahlow et al., 2020), followed by the newly-implemented benthic denitrification.”

Line 75: Please define what "H" version is meant for, and briefly mention how it is different from OPEM.

OLD: “The temperature dependence is from the OPEM-H configuration of Pahlow et al. (2020), where growth and nutrient uptake of phytoplankton and diazotrophs follow the same exponential Eppley (1972) temperature function but N₂ fixation is driven by the unimodal function of Houlton et al. (2008) with an optimal temperature of 26°C.”

NEW: “The temperature dependence is from the OPEM-H configuration of Pahlow et al. (2020), which applies the Houlton et al. (2008) unimodal temperature function to N₂ fixation by diazotrophs. The growth and nutrient uptake of phytoplankton and diazotrophs follow the same exponential (Eppley, 1972) temperature function. The temperature dependence in the OPEM configuration is the same as in the original UVic-ESCM, where diazotrophs grow only when temperature is above 15°C. Thus, the only difference between the OPEM and OPEM-H configurations is the temperature dependence of diazotrophs.”

Line 170-171: it would be good to expand a bit more on this that why this might be the case. Give some perspective.

OLD: “This suggests that N* provides a more powerful constraint on water-column denitrification rates, and hence N fixation, than simply combining NO₃⁻ and PO₄³⁻.”

NEW: “This suggests that N* provides a better, more independent constraint on water-column denitrification rates, and hence N fixation, than simply combining NO₃⁻ and PO₄³⁻. NO₃⁻ and PO₄³⁻ are highly correlated, largely following the Redfield N:P ratio. The correlations between different data (Chl, NO₃⁻, PO₄³⁻ and O₂) are already accounted for by our misfit metric (Krishna et al., 2019). However, N* not only accounts for the correlation between NO₃⁻ and PO₄³⁻, but can be related directly to non-Redfield processes, such as N₂ fixation, denitrification, or variations in particulate C:N:P stoichiometry.”

Line 180-181: Expand a bit more on the export production. What percentage of NPP is exported to the deep ocean and how does it compare to the estimates elsewhere.

OLD: “The new estimates of export production at 130m (model euphotic depth) are relatively close to and slightly lower than no_bdeni (Table 2), yielding higher export efficiencies (export production / NPP) than no_bdeni.”

NEW: “Export production at 130m (model euphotic depth) ranges from 6.9 to 8.1 Pg C yr⁻¹, slightly lower than that of no_bdeni (8.7 Pg C yr⁻¹, Table 2), yielding higher export efficiencies (export production / NPP) than no_bdeni. Approximately 13% of NPP is exported as sinking particles to the deep ocean among our new solutions (Table 2). Simulated export production at 100m in the Coupled Model Intercomparison Project Phase 5 (CMIP5) and newer CMIP6 models ranges from about 4.5 to 7.5 Pg C yr⁻¹ (Bopp et al., 2013; Laufkötter et al., 2016; Fu et al., 2016) and 5 to 12 Pg C yr⁻¹ (Séférián et al., 2020; Henson et al., 2022), respectively. Thus, our estimates fall within the CMIP6 spread. Other data-assimilated global ecosystem and biogeochemistry models yield particulate organic carbon export across the 100m depth horizon of 10.64 (Wang et al., 2023), 6.4 (Nowicki et al., 2022), and 6.7 Pg C yr⁻¹ (DeVries and Weber, 2017), yielding export ratios between 12% and 20%.”

Correspondingly, we will revise Table 2 with export ratios.

variable	model configuration				units
	no_bdeni	calib_NO ₃ ⁻	calib_N*_only	calib_N*	
N ₂ fixation	69.5	95.5	136.1	137.7	Tg N yr ⁻¹
Water-column Deni. (W)	69.5	2.8	42.4	30.3	Tg N yr ⁻¹
Benthic Deni. (B)	0	91.4	94.4	105.5	Tg N yr ⁻¹
B/W	–	32.6	2.2	3.5	1
B _{continental_shelf} /B	–	0.5	0.6	0.5	1
NPP	88	52.8	63.0	59.3	Pg C yr ⁻¹
Export production (at 130m)	8.7	6.9	8.1	7.9	Pg C yr ⁻¹
EP/NPP	9.9	13.1	12.9	13.3	%

Line 197: “fixed-N nitrogen” only one nitrogen.

OLD: “The global oceanic fixed-N nitrogen inventory is maintained by balanced supply of N by N₂ fixation at the surface ocean and removal by water-column and benthic denitrification.”

NEW: “The global oceanic **fixed-N** inventory is maintained by the balance of N₂ fixation **in** the surface ocean and water-column and benthic denitrification. ”

Line 279-280: Not clear to me. Please explain. In figure 8, Calib_N* - no_bdeni values are mostly positive, that means Calib_N* has more O₂ in the ocean, I see that in figure 6 as well. But how does the prediction of ODZs in no_bden case. If is a bit confusing, please clarify in the text.

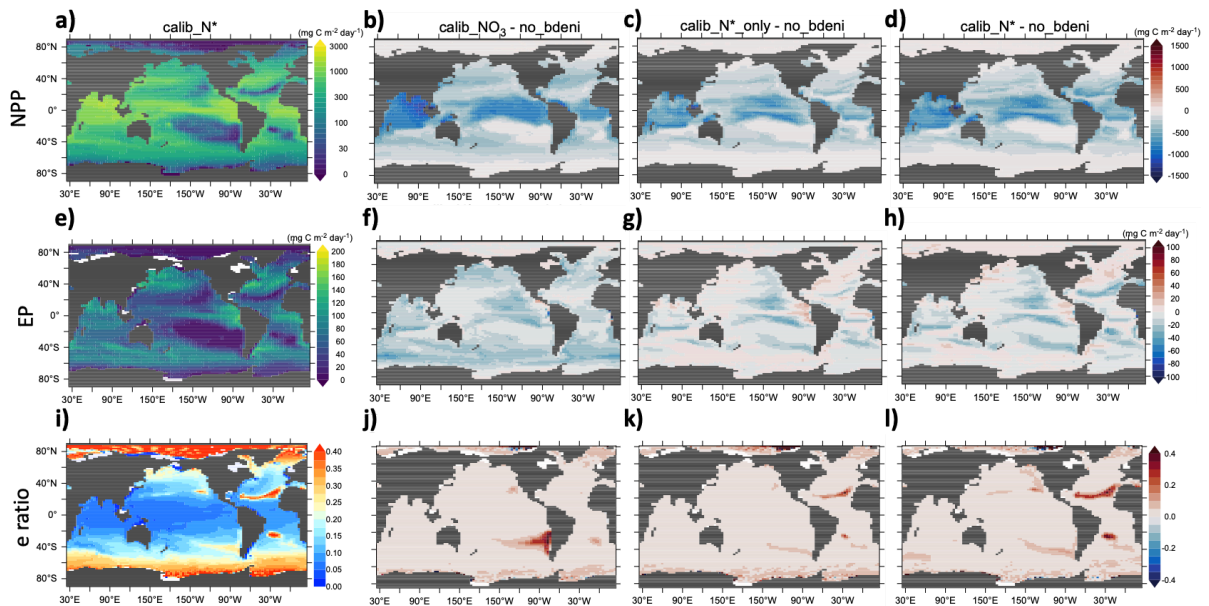
To make it clear, we added a row of ODZ volumes of all simulations in Table 2. If we define ODZ as the region with O₂ concentration < 5 mmol m⁻³, the ODZ volume in no_bdeni is >10¹⁵m³, whereas that of calib_N* is 6.58*10¹⁴m³.

OLD: “This implies that including benthic denitrification may improve the representation of ODZs in global ocean biogeochemical models that typically overestimate their volume (Cabr e et al., 2015).”

NEW: “**The lower ODZ volume when including benthic denitrification (Table 2)** implies that including benthic denitrification may improve the representation of ODZs in global ocean biogeochemical models that typically overestimate their volume (Cabr e et al., 2015).”

287-298: It will be nice to add a third row of figures on figure 9 plotting the EP ratio. Also, can you comment on the EP ratio how it compares to any other literature (observation or modeling)? It is good to put this information as the rain rate to the seafloor will ultimately drive the denitrification in the model.

Thank you for the suggestion. We added e-ratios in Figure 9.



We added the description of spatial distribution of e-ratio and comparison with other literature in L291.

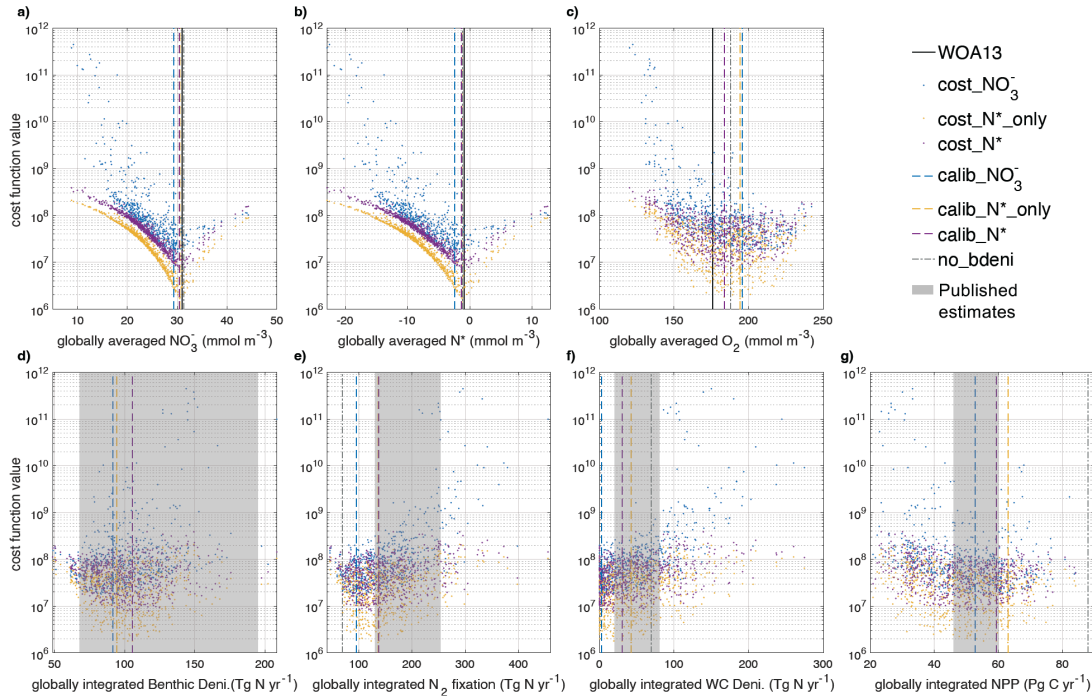
L291: “The global distribution of export efficiency (EP/NPP, hereafter e-ratio) exhibits a negative relationship between NPP and e-ratio. Low e-ratios (<0.2) occur in low latitudes and high e-ratios (>0.2) in high latitudes, with a few exceptions in the Atlantic and Pacific subtropical gyres (Fig. 9i). Its general pattern is similar to the observational estimate of Dunne et al. (2005) and the model estimate of Henson et al. (2015).”

305: I think you have not put any number of observational estimates of N fixation, if not please provide some number in relevant places.

OLD: “The absence of a persistent ODZ in the Arabian in our model is likely a reason why our global water-column denitrification and N₂ fixation rates are on the low-end of observational estimates.”

NEW: “The absence of a persistent ODZ in the Arabian Sea in our model is likely a reason why our global water-column denitrification and N₂ fixation rates are on the low-end of observational estimates range, between 131 and 253 Tg N yr⁻¹ (Großkopf et al., 2012; Luo et al., 2012; Landolfi et al., 2018; Shao et al., 2023).”

To make it easier to compare our estimates with other observational estimates, we indicated the observational estimates with an uncertainty range of 131 to 253 Tg N yr⁻¹ as a shaded grey area in Fig.2e.



Define what is biome when it first occurs.

Line 119: We added “Ocean biomes are geographical regions characterized by coherent large-scale patterns in physical and biogeochemical functions (Fay and McKinley, 2014), providing a representation of global ocean biogeography.”

If the Arabian sea has one, how does it compare with other modelled areas/seas?

Thank you for the question. In the cost function, each biome is characterised by two numbers for each depth layer (and month for the upper 5 layers of UVic): the spatial average and variance within the biome. In this study, we consider the Indian Ocean as a single biome including the Arabian Sea, thus the spatial average and variance of physical and biogeochemical variables within the whole Indian Ocean is taken into account in the cost function. If we consider the Arabian Sea to be one biome, the model calibration will explicitly reflect the spatial average and variance of its physical and biogeochemical variables, providing the same importance in cost function as other biomes.

We clarified further the calculation of the cost function in Section 2.4.2 (Line 116) as below.

OLD: “To assess the model performance with respect to the spatial distributions of dissolved tracers and surface chlorophyll a, we apply global misfit metrics J based on a maximum-likelihood estimation (ML) method for parameters, assuming log-normally distributed errors, with observations grouped by distinct biogeochemical biomes (Fay and McKinley, 2014), as described in Chien et al. (2020). Briefly, the calculation of our cost function comprises two terms for every depth level in our model,

$$J_k = \mathbf{d}_k^T \mathbf{R}_k^{-1} \mathbf{d}_k + \mathbf{v}_k^T \mathbf{V}_k^{-1} \mathbf{v}_k, k \in \{1, 2, \dots, 19\} \quad (1)$$

where the residual (\mathbf{d}) between observations (\mathbf{o}) and the model (\mathbf{m}) within each biome is defined as $\mathbf{d} = \mathbf{o} - \mathbf{m}$, and the discrepancy in the spatial variance (\mathbf{v}) as $\mathbf{v} = \mathbf{v}(\mathbf{o}) - \mathbf{v}(\mathbf{m})$. For each variable included in the cost function, we normalize the concentrations to threshold values and log-transform to achieve approximately Gaussian error distributions for \mathbf{d} and \mathbf{v} with zero mean and covariance matrices \mathbf{R} and \mathbf{V} , respectively. We integrate the results

across the 17 biomes, applying Eq. (1) to monthly means for the upper 550m (top 5 layers in UVic) and annual means below (bottom 14 layers).”

NEW: “To assess the model performance with respect to the spatial distributions of dissolved tracers and surface chlorophyll a, we apply global misfit metrics J based on a maximum-likelihood estimation (ML) method for parameters, assuming that the errors for the residuals of log-transformed variables between model simulations and observations follow normal distributions (Chien et al., 2020). Minimizing J ensures the best parameter estimates for the given model configuration.

Instead of calculating residuals between model simulations and observations for each model grid cell, we categorize model simulations and observations into 17 biomes (Fay and McKinley 2014). Ocean biomes are geographical regions characterized by coherent large-scale patterns in physical and biogeochemical functions (Fay and McKinley 2014), providing a representation of global ocean biogeography. We represent each variable by two statistical measures per biome: spatial average and variance. Therefore, the residuals comprise the discrepancies in spatial averages and variances for all biomes. In the vertical spatial dimension, we do not make any simplifications and residuals are calculated at each depth layer (k). We calculate the residuals of variables between monthly-averaged simulations and observations to resolve seasonal variations in the upper ocean (0-550m), and between annually-averaged simulations and observations below 550m.

Thus, the calculation of our cost function comprises two components,

$$J_k = \mathbf{A}_k^T \boldsymbol{\Sigma}_k^{-1} \mathbf{A}_k + \mathbf{V}_k^T \mathbf{Q}_k^{-1} \mathbf{V}_k, \quad k \in \{1, 2, \dots, 19\} \quad (1)$$

where \mathbf{A} is the residual between the spatial means of observations (\mathbf{o}) and the model (\mathbf{m}) defined as $\mathbf{A} = \text{mean}(\mathbf{o}) - \text{mean}(\mathbf{m})$ and \mathbf{V} is the residual in the spatial variance (var) as $\mathbf{V} = \text{var}(\mathbf{o}) - \text{var}(\mathbf{m})$. The covariance matrices of \mathbf{A} and \mathbf{V} are denoted by $\boldsymbol{\Sigma}$ and \mathbf{Q} , respectively.”

325-326: Define what is w_{d0} and w_{dd} , when it appears first.

w_{dd} is defined at Line 96 and we modified Line 325–326 as below.

OLD: “Improvements to the upper ocean may also have the potential to improve the deep-ocean performance, such as including w_{d0} into the calibrated parameters. By taking both w_{d0} (particle sinking velocity at the surface ocean) and w_{dd} into account for the calibration, ...”

NEW: “Improvements to the upper ocean may also have the potential to improve the deep-ocean performance, such as including the sinking speed of particles at the ocean surface (w_{d0}) in addition to its increase with depth (w_{dd}) into the set of calibrated parameters. By taking both w_{d0} and w_{dd} into account for the calibration, ...”

339-340: first part of this sentence is not clear.

OLD: “In order to explore the sensitivity to prior assumptions, we applied our cost function to calibrate model solutions against three different combinations of observations.”

NEW: “In order to explore the sensitivity to prior assumptions about which data are used for calibration, we applied our cost function to three different combinations of observations.”

345: “demonstrates”

OLD: “The greater constraining capacity of N* in comparison to considering nitrate and phosphate separately highlights the importance of accounting for correlations among variables within the cost function (Krishna et al., 2019) and demonstrates the power of diagnostic tracers such as N* for diagnostic studies of the ocean nitrogen cycle (DeVries et al., 2013; Eugster and Gruber, 2012; Deutsch et al., 2007).”

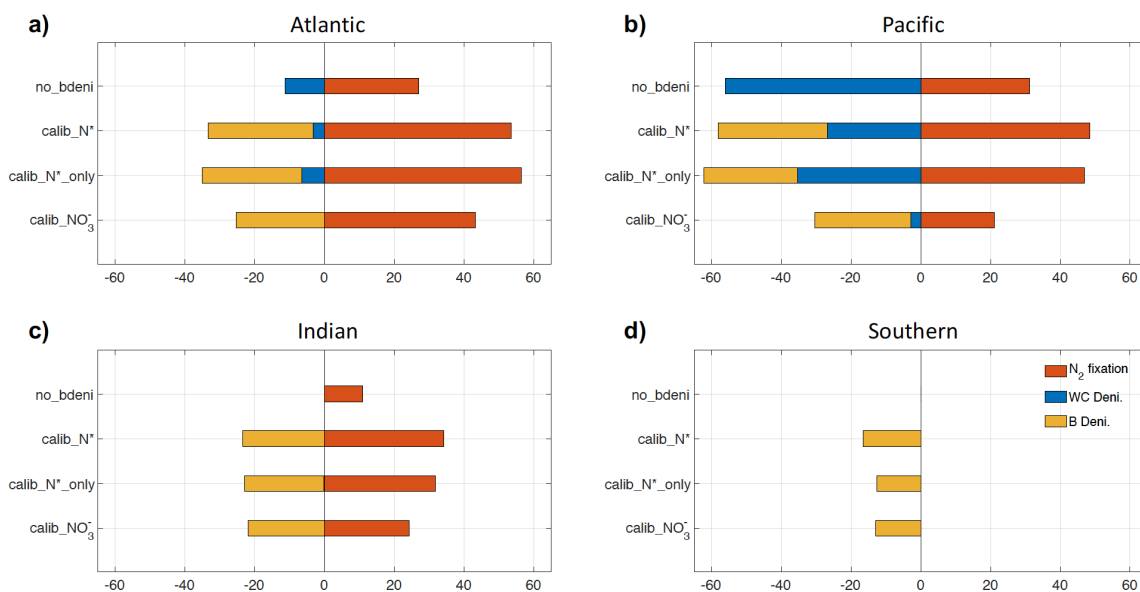
NEW: “The greater constraining capacity of N* in comparison to considering nitrate and phosphate separately highlights the importance of accounting for correlations among variables within the cost function (Krishna et al., 2019) and **demonstrates** the power of diagnostic tracers such as N* for diagnostic studies of the ocean nitrogen cycle (DeVries et al., 2013; Eugster and Gruber, 2012; Deutsch et al., 2007).”

Figure 2: Any reference for NPP to compare with?

Figure 2g: we marked published estimates of NPP from 46 to 60 Pg C yr⁻¹ (Johnson and Bif, 2021; Silsbe et al., 2016) as grey area.

Figure 5: explain why you did not include Indian ocean and southern ocean.

Figure 5: we added the Indian and Southern Ocean.



Line 213: we added the description of the simulated N fluxes in the Indian and Southern Ocean as below.

“Since our solutions lack water-column denitrification in the Indian Ocean, this area supplies extra nitrogen to the global ocean, even when benthic denitrification is taken into account. In contrast, the Southern Ocean becomes a nitrogen sink when benthic denitrification is included, as N₂ fixation does not occur in this area.”

Figure 6 caption: Typo.

OLD: “Vertical distributions of tracers (NO₃⁻, PO₃⁴⁻, N*, O₂) in global and respect basins (Atlantic, Pacific, Indian, Southern Oceans).”

NEW: “Vertical distributions of tracers (NO_3^- , PO_3^{4-} , N^* , O_2) in the global ocean and the Atlantic, Pacific, Indian, Southern Oceans.”

Figure 8 caption: three solutions or N^* ?

yes, it should be solution calib_ N^* . We removed “of three new solutions”.

Table 1 caption: units. Typo.

Table 1 caption: Parameter names, ranges, units and descriptions.

Table 1: did you get the reference range from somewhere? please clarify.

yes, the reference range comes from Pahlow et al., 2020 and Chien et al., 2020. We added the clarification “The reference range refers to the calibration range in Pahlow et al. (2020); Chien et al. (2020)” below the table.

Table A1: Why do you provide bar on top of TN^* ? It is not defined in the table.

The bar on top of TN^* means the average of TN^* of each basin. We added the definition in the table title.

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Response to egosphere-2024-123 RC2

Below is our response to RC2. Referee #2’s comments are in **bold**. Our responses are in black. Changes to the manuscript are indicated in blue. The line numbers refer to the previously submitted manuscript without track changes.

This manuscript describes the improvement of an Earth system model by incorporating benthic denitrification, highlighting the importance of benthic denitrification in shaping the global distributions of NPP, N₂ fixation, oxygen, etc. In addition, the authors conducted a large ensemble of simulations and applied the Latin-hypercube sampling method to choose the model parameters. The simulation identified that the N* is an essential parameter to calibrate the model, which has an important implication for future model calibration. Overall, the manuscript was well-written and organized, and the references cited are up-to-date and appropriate. I recommend publication after making the following clarifications and modifications.

We thank Referee #2 for the positive feedback.

I'd like to see more discussion on the pattern of N₂ fixation. Currently, the distribution patterns in the S. Pacific and N. Indian Ocean (Fig. 4a) seem not correct to my eyes. What causes such distribution, is there anything to do with benthic denitrification?

We agree that our model may underestimate N₂ fixation rates in the South Pacific and N. Indian Ocean (Fig. 4a). However, it is unlikely that the underestimation stems from benthic denitrification, since the differences between new simulations with no_bdeni (Fig. 4b-d) indicate there is no strong pattern shift in these basins compared to the changes in other subtropical regions. We suspect there are two different dominant processes that contribute to the underestimation of the N₂ fixation rates at these two basins separately.

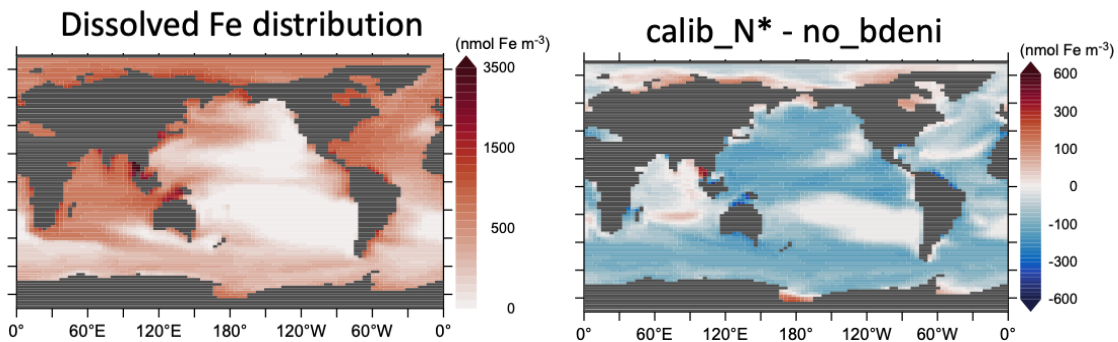
In our model, diazotroph growth and N₂ fixation in the South Pacific are subjected to the low Fe availability indicated by the low dissolved Fe concentration (see the left panel of Figure below), especially at the South Pacific Subtropical Gyre. Our simulated surface Fe concentration in the South Pacific Subtropical Gyre is lowest among all subtropics, whereas observations and other models (Huang et al., 2022) suggest that the surface dissolved Fe in the South Pacific is comparable to that in the North Pacific. We further assess the effect of benthic denitrification on dissolved Fe concentration (see the right panel of Figure below). In the South Pacific, the differences of surface dissolved Fe concentrations between calib_N*

and no_bdeni shows only minor changes. This minimal influence of benthic denitrification on dissolved Fe suggests that benthic denitrification does not affect N_2 fixation through Fe concentration in this region.

The absence of nitrogen fixation in the Northern Indian Ocean may result from the absence of an oxygen-deficient zone (ODZ) and water-column denitrification in this area in our model. A proper representation of the ODZ in the Arabian Sea is one of our goals of ongoing model development.

We added on Line 213: “The very low N_2 fixation rates in the South Pacific (Fig. 4a) can be attributed to the underestimated surface dissolved iron concentration in this region.”

We added on line 207: “This is also the main reason for the very low rates of N_2 fixation predicted for the Northern Indian Ocean, which could be unrealistic. Yet, there is considerable uncertainty about the regional pattern of N_2 fixation in the Northern Indian Ocean due to the sparsity of available observations. For example, Shao et al. (2020) found strong N_2 fixation rates at only a few of places along the southwest coast of India in the eastern Arabian Sea, Löscher et al. (2020) could find no evidence for N_2 fixation in the Bay of Bengal, and vast areas in the Northern and Western Indian ocean remain unsampled.”



In section 4 "Model and calibration limitations", the authors discussed multiple ways to improve the model, including applying different parameterization schemes for upper and deeper ocean sinking speed and resolving the dependency of remineralization rate on O₂. The methods described are all reasonable. But why not try them, this is part of model development.

This research concentrates on the impact of benthic denitrification on global ocean biogeochemistry in our model. A full investigation of the possible development discussed in Section 4 requires calibrating the model configurations for each change, which is very time-consuming and goes beyond the scope of this work. We plan to do some of these in the future.

Some minor suggestions

Line 83 (Somes and Oschlies, 2015) --> Somes and Oschlies, (2015)

ok

Line 92 (Pahlow et al., 2020; Chien et al., 2020) --> Pahlow et al., (2020) and Chien et al., (2020)

ok

Line 149 Best models choices --> Best model choice.

ok

Line 305 Arabian --> Arabian Sea

ok

References

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