

1 **Nocturnal Atmospheric Synergistic Oxidation Reduces the Formation of Low-volatility**
2 **Organic Compounds from Biogenic Emissions**

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9

10 **Abstract**

11 Volatile organic compounds (VOCs) are often subject to synergistic oxidation by different oxidants
12 in the atmosphere. However, the exact synergistic oxidation mechanism of atmospheric VOCs and
13 its role in particle formation remain poorly understood. In particular, the reaction kinetics of the key
14 reactive intermediates, organic peroxy radicals (RO₂), during synergistic oxidation is rarely studied.
15 Here, we conducted a combined experimental and kinetic modelling study of the nocturnal
16 synergistic oxidation of α -pinene (the most abundant monoterpene) by O₃ and NO₃ radicals as well
17 as its influences on the formation of highly oxygenated organic molecules (HOMs) and particles.
18 We find that in the synergistic O₃ + NO₃ regime, where OH radicals are abundantly formed via
19 decomposition of ozonolysis-derived Criegee intermediates, the production of C_xH_yO_z-HOMs is
20 substantially suppressed compared to that in the O₃-only regime, mainly because of the depletion of
21 of α -pinene RO₂ derived from ozonolysis and OH oxidation by those arising from NO₃ oxidation
22 via cross reactions. Measurement-model comparisons further reveal that the cross-reaction rate
23 constants of NO₃-derived RO₂ with O₃-derived RO₂ are on average 10 – 100 times larger than those
24 of NO₃-derived RO₂ with OH-derived RO₂. Despite a strong production of organic nitrates in the
25 synergistic oxidation regime, the substantial decrease of C_xH_yO_z-HOM formation leads to a
26 significant reduction in ultralow- and extremely low-volatility organic compounds, which
27 significantly inhibits the formation of new particles. This work provides valuable mechanistic and
28 quantitative insights into the nocturnal synergistic oxidation chemistry of biogenic emissions and
29 will help to better understand the formation of low-volatility organic compounds and particles in
30 the atmosphere.

31

32 **1. Introduction**

33 The Earth's atmosphere is a complex oxidizing environment in which multiple oxidants coexist.
34 During the nighttime, NO₃ radicals (generated by the reaction of NO₂ and O₃) and O₃ contribute
35 significantly to the oxidation of volatile organic compounds (VOCs) (Huang et al., 2019), while
36 during the daytime, the fast photolysis of NO₃ radicals and rapid photochemical formation of OH
37 radicals and O₃ make the latter two the major oxidants for VOCs (Zhang et al., 2018). Therefore,
38 the degradation of ambient VOCs is subject to concurrent oxidation by different oxidants. Gas-phase
39 oxidation of VOCs from biogenic emissions (BVOCs) by these major atmospheric oxidants
40 produces a key type of reactive intermediates, organic peroxy radicals (RO₂), a portion of which can
41 undergo fast autoxidation forming a class of highly oxygenated organic molecules (HOMs) with
42 low volatilities (Jokinen et al., 2014; Mentel et al., 2015; Berndt et al., 2016; Zhao et al., 2018; Iyer
43 et al., 2021; Shen et al., 2022; Ehn et al., 2014). HOMs typically contain six or more oxygen atoms,
44 and play a key role in the formation of atmospheric new particles and secondary organic aerosol
45 (SOA) (Kirkby et al., 2016; Berndt et al., 2018; Zhao et al., 2018; Ehn et al., 2014; Bianchi et al.,
46 2019), which have important influences on air quality (Huang et al., 2014), public health (Pye et al.,
47 2021), and Earth's radiative forcing (Shrivastava et al., 2017).

48 Due to the complexity of oxidation mechanisms of BVOCs, previous laboratory studies typically
49 featured only one oxidant and a single SOA precursor (Berndt et al., 2016; Berndt, 2021; Clafin et
50 al., 2018; Iyer et al., 2021; Boyd et al., 2015). However, the synergistic oxidation by different
51 oxidants may significantly alter the fate of RO₂ intermediates, therefore influencing the formation
52 of HOMs and SOA (Bates et al., 2022). Recently, a field study at a boreal forest site in Finland
53 observed a series of nitrate-containing HOM-dimers from the coupled O₃ and NO₃ oxidation of
54 monoterpenes (Zhang et al., 2020). At the same site, Lee et al. (2020) found that the synergistic
55 oxidation of BVOCs by OH radicals and O₃ contributed to the largest fraction of SOA. These studies
56 suggest that the synergistic oxidation of BVOCs by different oxidants plays an important role in the
57 formation of HOMs and SOA in the atmosphere and highlight the needs to investigate the synergistic
58 oxidation mechanisms of BVOCs for a better representation of atmospheric particle formation.

59 Several laboratory studies have attempted to address the role of synergistic oxidation of BVOCs in
60 the formation of new particles and SOA (Kenseth et al., 2018; Inomata, 2021; Liu et al., 2022; Li et

61 al., 2024). Kenseth et al. (2018) identified a suite of dimer esters in flow tube experiments that can
62 be only formed from the OH and O₃ synergistic oxidation of β-pinene. These dimers exhibit
63 extremely low volatility and contributed 5.9 – 25.4% to the total β-pinene SOA. Similarly, Inomata
64 (2021) found that the presence of OH radicals during α-pinene ozonolysis is a key factor for the
65 production of low-volatility organic species and significantly promotes new particle formation
66 (NPF). On the other hand, the addition of O₃ in the monoterpene photooxidation system also
67 significantly increases the SOA mass yield (Liu et al., 2022). In addition, a recent chamber study
68 by Bates et al. (2022) showed that the synergistic oxidation of α-pinene by NO₃ radicals and O₃ can
69 significantly enhance the SOA yield compared to the NO₃ + α-pinene regime, which has nearly 0%
70 SOA yield (Fry et al., 2014; Hallquist et al., 1999; Mutzel et al., 2021), and they revealed that the
71 SOA yield in the NO₃ + O₃ oxidation system largely depends on the RO₂ fates. Most recently, Li et
72 al. (2024) found that during α-pinene ozonolysis, the presence of nitrooxy-RO₂ radicals formed from
73 NO₃ oxidation can significantly suppress the production of ultralow-volatility organic compounds
74 (ULVOCs) and thereby NPF. These laboratory studies together provide growing evidence that
75 synergistic oxidation of BVOCs by different oxidants have profound impacts on atmospheric
76 particle formation. However, the specific synergistic mechanisms of different oxidants and
77 oxidation pathways remain obscure. Although a few studies underscored the importance of the RO₂
78 fates (Bates et al., 2022; Li et al., 2024), the exact interactions between RO₂ species derived from
79 different oxidants are still unclear, and quantitative constraints on the reaction rate of different RO₂
80 species are quite limited.

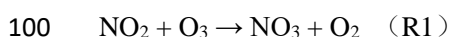
81 Here we conducted an investigation of the synergistic O₃ + NO₃ oxidation of α-pinene, one of the
82 most abundant monoterpenes in the atmosphere, using a combination of laboratory experiments and
83 detailed kinetic modelling, and focusing on the fate of RO₂ intermediates arising from different
84 oxidation pathways. The α-pinene oxidation experiments were conducted in a custom-built flow
85 reactor. The molecular composition of RO₂ species and HOMs in different oxidation regimes was
86 characterized using a chemical ionization atmospheric pressure interface time-of-flight mass
87 spectrometer (CI-APi-ToF) employing a nitrate ion source. The measured distributions of specific
88 RO₂ and HOMs across different oxidation regimes were fitted with a kinetic model using Master
89 Chemical Mechanisms (MCM v3.3.1) updated with recent advances of α-pinene RO₂ chemistry

90 (Wang et al., 2021; Iyer et al., 2021; Shen et al., 2022; Zang et al., 2023), which allows for
91 quantitative constraints on RO₂ kinetics and synergistic oxidation mechanisms. Atmospheric
92 relevance of the experimental results was evaluated by modelling the investigated oxidation
93 chemistry under typical nocturnal atmospheric conditions.

94 **2. Materials and Methods**

95 **2.1 Flow tube experiments**

96 Experiments of α -pinene oxidation in different regimes (i.e., synergistic O₃ + NO₃ oxidation vs. O₃-
97 only) were carried out under room temperature (298 K) and dry (relative humidity < 5%) conditions
98 in a custom-built flow tube reactor (FTR, Figure S1). O₃ and NO₂ were added into a glass tube
99 (Figure S1) to form NO₃ radical and its precursor N₂O₅:



102 O₃ was generated by passing a flow of ultra-high-purity (UHP) O₂ (Shanghai Maytor Special Gas
103 Co., Ltd.) through a quartz tube housing a pen-ray mercury lamp (UV-S2, UVP Inc.) and its
104 concentration was measured by an ozone analyzer (T400, API). NO₂ was obtained from a gas
105 cylinder (15.6 ppm, Shanghai Weichuang Standard Gas Co., Ltd.). The initial NO₂ concentration in
106 the flow tube was 4.5 – 6.4 ppb. To prevent the titration of NO₃ radicals by NO, all the experiments
107 were performed without the addition of NO. The total air flow in the NO₃ generation glass tube was
108 0.6 L min⁻¹ and 0.4 L min⁻¹ for the gas-phase HOM and SOA formation experiments, respectively.
109 The produced N₂O₅ and NO₃ radicals, as well as the excessive O₃ were added into the FTR to initiate
110 α -pinene oxidation. For the O₃-only experiments, only O₃ was added into FTR.

111 The α -pinene gas was generated by evaporating a defined volume of its liquid (99%, Sigma-Aldrich)
112 into a cleaned and evacuated canister (SilcoCan, RESTEK), and then added into FTR through a
113 movable injector at a flow rate of 22 – 108 mL min⁻¹. The initial concentration of α -pinene in the
114 flow reactor ranged from 100 – 500 ppb. In some experiments, the gas of cyclohexane (~ 100 ppm),
115 which was generated by bubbling a gentle flow of UHP N₂ through its liquid (LC-MS grade, CNW),
116 was added into the flow reactor as a scavenger of OH radicals formed from α -pinene ozonolysis.

117 For experiments characterizing the formation of HOMs, the total air flow in the FTR was 10.8 L

118 min⁻¹ and the residence time was 25 seconds. The short reaction time and the small amount of
119 reacted α -pinene (see Table S1) in these experiments prevented the formation of particles. For the
120 experiments characterizing the formation of SOA particles, a larger FTR was used, with a total air
121 flow of 5 L min⁻¹ and a residence time of 180 seconds. A summary of the conditions including the
122 simulated concentrations of NO₂, N₂O₅ and NO₃ radicals, as well as the concentration of α -pinene
123 oxidized by each oxidant in different experiments are shown in Table S1.

124 The gas-phase RO₂ radicals and closed-shell products were measured using a nitrate-based CI-API-
125 ToF (abbreviated as nitrate-CIMS; Aerodyne Research, Inc.), which has been described in detail
126 previously (Zang et al., 2023). A long ToF-MS with a mass resolution of ~ 10000 Th/Th was used
127 here. The mass spectra within the m/z range of 50 – 700 were analyzed using the tofTools package
128 developed by Junninen et al. (2010) based on Matlab. The total ion counts (TIC) with values of $(5.9$
129 $- 6.2) \times 10^4$ cps are similar under different reaction conditions. In this study, we assume that the
130 C_xH_yO_z-HOMs derived from ozonolysis and OH oxidation of α -pinene exhibit the same sensitivity
131 in nitrate-CIMS. However, the highly oxygenated organic nitrates (ONs) may have a significantly
132 lower sensitivity compared to the C_xH_yO_z-HOM counterparts, given that the substitution of -OOH
133 or -OH groups by -ONO₂ group in the molecule would reduce the number of H-bond donors, which
134 is a key factor determining the sensitivity of nitrate-CIMS (Shen et al., 2022; Hyttinen et al., 2015).
135 Recently, Li et al. (2024) used CI-Orbitrap with ammonium or nitrate reagent ions to detect
136 oxygenated organic molecules in the synergistic O₃ + NO₃ regime and found that both the ion
137 intensity of ONs and their signal contribution to the total dimers were much lower when using nitrate
138 as reagent ions.

139 A scanning mobility particle sizer (SMPS, TSI), consisting of an electrostatic classifier (model
140 3082), a condensation particle counter (model 3756), and a long or nano differential mobility
141 analyzer (model 3081 and 3085) with a measurable size range of 4.6 – 156.8 nm or 14.6 – 661.2
142 nm, respectively, was employed to monitor the formation of particles in the flow tube. During the
143 HOM formation experiments, even under conditions with the highest initial α -pinene concentration
144 (500 ppb), only a tiny amount of particles was formed, with mass concentrations of $(6.4 \pm 1.7) \times 10^{-3}$
145 and $(1.0 \pm 0.4) \times 10^{-2}$ $\mu\text{g m}^{-3}$ and number concentrations of 574 ± 148 and 256 ± 68 cm⁻³ in the O₃-
146 only (Exp 5) and O₃ + NO₃ regimes (Exp 11), respectively. These results suggest that the formation

147 of SOA particles in the HOM formation experiments is negligible and would have no significant
148 influence on the fate of RO₂ and closed-shell products.

149 **2.2 Estimation of HOM volatility**

150 A modified composition-activity method was used to estimate the saturation mass concentration
151 (C^*) of HOMs in this study according to the approach developed by Li et al. (2016):

$$152 \log_{10}C^* = (n_C^0 - n_C)b_C - n_O b_O - 2 \frac{n_C n_O}{n_C + n_O} b_{CO} - n_N b_N - n_S b_S$$

153 where n_C^0 is the reference carbon number; n_C , n_O , n_N , and n_S are the atom numbers of carbon,
154 oxygen, nitrogen, and sulfur, respectively; b_C , b_O , b_N , and b_S are the contribution of each atom
155 to $\log_{10}C^*$, respectively; b_{CO} is the carbon–oxygen nonideality (Donahue et al., 2011). These b -
156 values were provided by Li et al. (2016).

157 It should be noted that the CHON compounds used in the data set by Li et al. (2016) are mostly
158 amines, amides, and amino acids, and only contain a limited number of ONs (0.07%). Since different
159 types of CHON compounds have very different vapor pressures (Isaacman-Vanwertz and Aumont,
160 2021), this formula-based approach can be biased to estimate the C^* of ONs. Considering that the
161 –ONO₂ and –OH groups have similar impacts on vapor pressure and that the CHON species are
162 predominantly ONs in our study, all –ONO₂ groups are treated as –OH groups during the estimation
163 of vapor pressure (Daumit et al., 2013; Isaacman-Vanwertz and Aumont, 2021).

164 Gas-phase HOMs are grouped into five classes based on their $\log_{10}C^*$ (Donahue et al., 2012;
165 Bianchi et al., 2019; Schervish and Donahue, 2020), that is, ULVOCs ($\log_{10}C^* < -8.5$), extremely
166 low-volatility organic compounds (ELVOCs, $-8.5 < \log_{10}C^* < -4.5$), low-volatility organic
167 compounds (LVOCs, $-4.5 < \log_{10}C^* < -0.5$), semi-volatile organic compounds (SVOCs, $-0.5 <$
168 $\log_{10}C^* < 2.5$), and intermediate-volatility organic compounds (IVOCs, $2.5 < \log_{10}C^* < 6.5$).

169 **2.3 Kinetic model simulations**

170 Model simulations of specific RO₂ radicals and closed-shell HOMs formed in different oxidation
171 regimes were performed to constrain the reaction kinetics and mechanisms using the Framework
172 for 0-D Atmospheric Modeling (F0AM v4.1) (Wolfe et al., 2016), which employs MCM v3.3.1
173 (Jenkin et al., 2015). The α -pinene oxidation mechanism was updated with the state-of-the-art

174 knowledge on the chemistry of RO₂ autoxidation and cross reactions forming HOM monomers and
175 dimers, respectively (Zhao et al., 2018; Wang et al., 2021; Iyer et al., 2021; Shen et al., 2022). The
176 detailed updates have been described in our previous study (Zang et al., 2023). In particular, the
177 formation and subsequent reactions of the ring-opened primary C₁₀H₁₅O₄-RO₂, the highly
178 oxygenated acyl RO₂, as well as the C₁₀H₁₅O₂-RO₂ arising from H-abstraction by OH radicals
179 during α -pinene ozonolysis are included in the model according to recent studies (Iyer et al., 2021;
180 Zhao et al., 2022; Zang et al., 2023; Shen et al., 2022).

181 To investigate the synergistic reactions of RO₂ derived from the oxidation of α -pinene by different
182 oxidants, we added the cross reactions of the primary nitrooxy-RO₂ derived from NO₃ oxidation
183 (^{NO₃}RO₂), i.e., C₁₀H₁₆NO₅-RO₂, with RO₂ derived from ozonolysis (^{Cl}RO₂) and OH oxidation
184 (^{OH}RO₂). Recently, Zhao et al. (2018) revealed the bulk rate constant for ^{Cl}RO₂ and ^{OH}RO₂ self/cross
185 reactions to be $2 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$, and Bates et al. (2022) constrained the rate constant for
186 ^{NO₃}RO₂ self/cross reactions to be $1 \times 10^{-13} - 1 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$. In the present study, the
187 default rate constant for ^{NO₃}RO₂ + ^{Cl}RO₂ was set to $2 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$, the same to that for
188 self/cross reactions of ^{Cl}RO₂ and ^{OH}RO₂. The default rate constant for ^{NO₃}RO₂ + ^{NO₃}RO₂ was set to
189 $1 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$. The ratio of the cross-reaction rate constant of ^{NO₃}RO₂ + ^{Cl}RO₂ to that
190 of ^{NO₃}RO₂ + ^{OH}RO₂ was tuned to achieve a good measurement-model agreement for the distribution
191 of specific RO₂ and HOMs across different oxidation regimes. Recent studies suggested that the
192 ROOR' dimer formation rates from the highly oxygenated RO₂ are fast (Berndt et al., 2018; Molteni
193 et al., 2019). As a result, a relatively high dimer formation branching ratio of 50% was used for
194 different RO₂ (e.g., ^{Cl}RO₂, ^{OH}RO₂, ^{NO₃}RO₂) in the model, except for the reaction of ^{NO₃}RO₂ + ^{NO₃}RO₂,
195 for which ROOR' dimer formation was not considered, given the extremely low signals of CHON₂
196 dimers observed in the synergistic oxidation regime (see Section 3.1). With these default kinetic
197 parameters, the RO₂ bimolecular lifetimes were predicted to be 10.9 – 25.9 s in the O₃-only regime
198 and 8.4 – 11.8 s in the O₃ + NO₃ regime in the HOM formation experiments. Considering that the
199 RO₂ cross-reaction kinetics remain highly uncertain, sensitivity analyses were performed to evaluate
200 their influences on the results in this study (see Section 3.2). Previous studies indicated that the
201 primary ^{NO₃}RO₂ radicals arising from α -pinene are prone to lose the nitrate group and form
202 pinonaldehyde with high volatility (Kurtín et al., 2017; Fry et al., 2014). Therefore, we did not

203 consider the autoxidation of primary NO_3RO_2 in the model. Considering the presence of NO_2 in the
204 experiments, the reactions of $\text{RO}_2 + \text{NO}_2 \rightleftharpoons \text{ROONO}_2$ were also included in the model (Zang et al.,
205 2023).

206 3. Results and Discussion

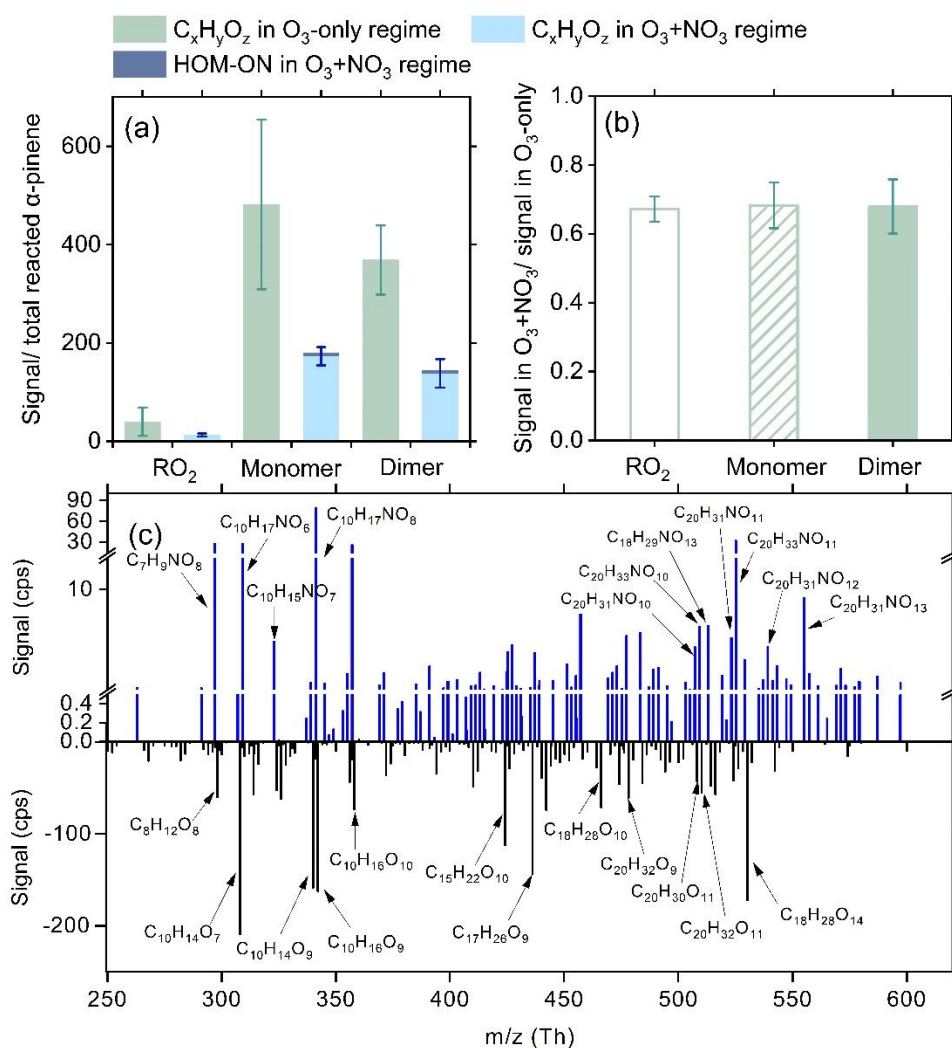
207 3.1 Molecular distribution of RO_2 and HOMs in the synergistic oxidation regime

208 The abundance of gas-phase RO_2 species and HOMs in different oxidation regimes is shown in
209 Figure 1a. The species signals are normalized by the total reacted α -pinene in each regime.
210 Compared to the O_3 -only regime, the normalized signals of total RO_2 and HOMs decrease by 63 –
211 68% in the synergistic $\text{O}_3 + \text{NO}_3$ regime. Although NO_3 oxidation accounts for a considerable
212 fraction of reacted α -pinene in the synergetic oxidation regime, the signal contributions of HOM-
213 ONs are not significant. This might be due to the low sensitivity of nitrate-CIMS to the ONs formed
214 involving NO_3 oxidation (Section 2.1). Although there remain considerable uncertainties in
215 instrument sensitivities to different compounds, sensitivity analyses suggest that varying the CIMS
216 sensitivities to RO_2 and HOMs by a factor of 10 would not significantly influence their relative
217 distribution across different oxidation regimes (see Section S1 for details).

218 Note that the initial concentrations of α -pinene and O_3 in the two oxidation regimes were the same.
219 In addition, model simulations show that in the synergistic $\text{O}_3 + \text{NO}_3$ regime, over 97% of OH
220 radicals react with α -pinene and the depletion of OH by NO_2 is minor (0.2 – 1.3%). Also, NO_3
221 radicals almost entirely (over 98.5%) react with α -pinene and their reaction with RO_2 has negligible
222 influence on the fate of RO_2 (Figure S2). Meanwhile, the depletion of acyl RO_2 by NO_2 only leads
223 to a small reduction (4 – 5% and 7 – 12%, respectively) in total $\text{C}_x\text{H}_y\text{O}_z$ -HOM monomers and dimers
224 in the synergistic regime compared to the O_3 -only regime. As a result, the strong reduction in HOM
225 formation due to the presence of NO_3 oxidation is likely mainly due to (i) the fast competitive
226 consumption of α -pinene by NO_3 radicals, which leads to a reduction in the reacted α -pinene by O_3
227 ($\Delta[\alpha\text{-pinene}]_{\text{O}_3}$, Figure S3) and thereby $\text{C}_x\text{H}_y\text{O}_z$ -HOM signals, and (ii) the cross reactions of $^{\text{Cl}}\text{RO}_2$
228 or $^{\text{OH}}\text{RO}_2$ with NO_3RO_2 , which suppress the autoxidation and self/cross reactions of $^{\text{Cl}}\text{RO}_2$ and $^{\text{OH}}\text{RO}_2$
229 to form $\text{C}_x\text{H}_y\text{O}_z$ -HOMs.

230 To quantify the contribution of cross reactions of NO_3RO_2 with $^{\text{Cl}}\text{RO}_2/^{\text{OH}}\text{RO}_2$ to the suppressed

231 formation of $C_xH_yO_z$ -HOMs in the synergistic oxidation regime, $C_xH_yO_z$ -HOM signals shown in
 232 Figure 1a are first normalized to $\Delta[\alpha\text{-pinene}]_{O_3}$ in each oxidation regime and then compared between
 233 different oxidation regimes (see Figure 1b). Notably, after excluding the influence of reduced $\Delta[\alpha\text{-}$
 234 pinene] $_{O_3}$, the $C_xH_yO_z$ -HOM signals still drop by 32 – 33% in the $O_3 + NO_3$ regime compared to
 235 those in the O_3 -only regime, indicating a significant contribution of the coupled reactions between
 236 NO_3RO_2 and $ClRO_2$ or $OHRO_2$ to suppressed $C_xH_yO_z$ -HOM formation.



237
 238 Figure 1 Distributions of RO_2 and HOMs in the O_3 -only and $O_3 + NO_3$ regimes. (a) Signals of total
 239 RO_2 , as well as HOM monomers and dimers normalized by the total reacted α -pinene in each
 240 oxidation regime (Exps 1-5, 7-11). (b) Relative changes in the normalized signals of $C_xH_yO_z$ -HOMs
 241 in the $O_3 + NO_3$ regime versus the O_3 -only regime. Ion signals are normalized to $\Delta[\alpha\text{-pinene}]_{O_3}$
 242 in each oxidation regime to highlight the suppression effect of the synergistic chemistry between
 243 NO_3RO_2 and $ClRO_2$ or $OHRO_2$ on $C_xH_yO_z$ -HOM formation. (c) Difference mass spectrum between
 244 the two oxidation regimes. The positive and negative peaks indicate the species with enhanced and
 245 decreased formation in the $O_3 + NO_3$ regime compared to the O_3 -only regime, respectively.

246 Figure 1c shows a difference mass spectrum highlighting the changes in species distribution
247 between the two oxidation regimes. Almost all $C_xH_yO_z$ -HOM species decrease significantly in the
248 $O_3 + NO_3$ regime compared to the O_3 -only regime. Besides, a large set of HOM-ON species are
249 formed, despite their relatively low signals. It should be noted that no obvious signals of highly
250 oxygenated $^{NO_3}RO_2$ ($C_{10}H_{16}NO_x$, $x \geq 6$) were observed by nitrate-CIMS in the $O_3 + NO_3$ oxidation
251 system. One possible reason is that nitrate-CIMS exhibits relatively low sensitivity to the ONs.
252 Secondly, the instrument's mass resolution is not high enough to differentiate the mass closure
253 between some of $^{NO_3}RO_2$ and $C_xH_yO_z$ -HOMs with strong peaks (Table S2), limiting the detection
254 of $^{NO_3}RO_2$ species. Furthermore, previous studies revealed that the primary $^{NO_3}RO_2$ radicals (i.e.,
255 $C_{10}H_{16}NO_5$ - RO_2) in the α -pinene + NO_3 system mainly react to form pinonaldehyde (Kurtén et al.,
256 2017; Perraud et al., 2010). It is likely that only a very small amount of $^{NO_3}RO_2$ can undergo
257 intramolecular H-shift/ O_2 addition to form highly oxygenated $^{NO_3}RO_2$. It should be pointed out that
258 although the primary $C_{10}H_{16}NO_5$ - RO_2 species arising from NO_3 oxidation may not undergo fast
259 autoxidation, they tend to efficiently terminate $^{Cl}RO_2$ and/or $^{OH}RO_2$ and suppress the formation of
260 $C_xH_yO_z$ -HOMs.

261 As shown in Figure 1c, although several closed-shell monomeric HOM-ONs have been observed in
262 the synergistic oxidation regime, only a few of them exhibit relatively high signals. Among them,
263 $C_{10}H_{17}NO_8$ may be formed by the autoxidation of $C_{10}H_{16}NO_6$ - RO_2 derived from the intramolecular
264 H-shift of primary ^{NO_3}RO radicals ($C_{10}H_{16}NO_4$ - RO). In addition, although CI is a soft ionization
265 method, the fragmentation of chemically labile species still occurs during the ionization in nitrate-
266 CIMS. It is possible that some of dimeric HOM-ONs are fragmented to $C_{10}H_{17}NO_8$ during nitrate-
267 CIMS measurements. In a recent study by Li et al. (2024), $C_{10}H_{17}NO_8$ was also identified during
268 the synergistic oxidation of α -pinene by O_3 and NO_3 . However, the exact origin of this species
269 remains to be clarified.

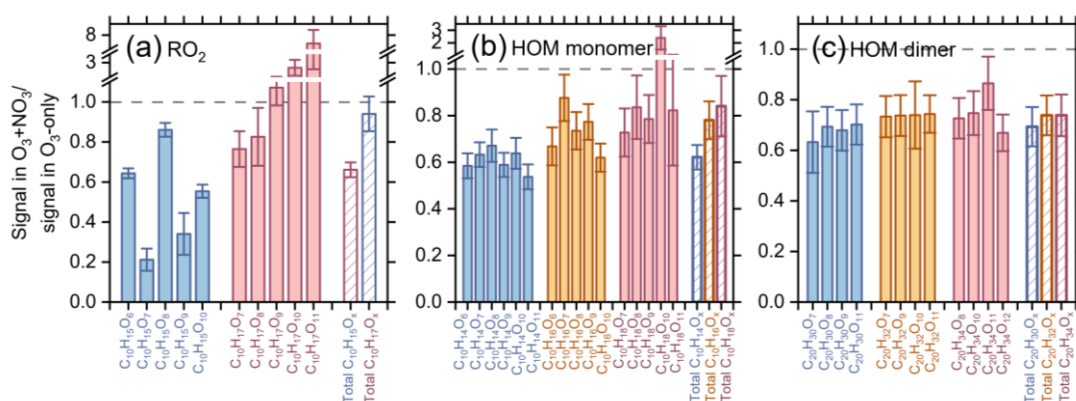
270 The C_{20} dimers with only one nitrogen atom are very likely to be formed from the cross reactions
271 of $^{Cl}RO_2$ or $^{OH}RO_2$ with $^{NO_3}RO_2$, which provides direct evidence for the synergistic RO_2 chemistry
272 in the $O_3 + NO_3$ regime. The $CHON_2$ dimers were also observed in the $O_3 + NO_3$ regime, despite
273 their much lower signals than $CHON$ dimers, which is different from the recent studies by Bates et
274 al. (2022) and Li et al. (2024), which found $CHON_2$ dimers account for an important fraction of the

275 total dimer signals in the synergistic oxidation regime. A potential explanation for this discrepancy
276 is the difference in the instrument sensitivity in these studies (Section 2.1). In general, the nitrate-
277 CIMS has lower sensitivities to ONs than to the $C_xH_yO_z$ -HOM counterparts (Shen et al., 2022;
278 Hyttinen et al., 2015). Bates et al. (2022) used CF_3O^- as the reagent ion of CIMS. Its sensitivity to
279 ONs might be significantly higher than the nitrate ion. In addition, Li et al. (2024) observed a
280 significantly lower signal contribution of $CHON_2$ dimers using CI-Orbitrap with nitrate reagent ions
281 than with ammonium ions. Despite both using nitrate reagent ions, the nitrate CI-Orbitrap in Li et
282 al. (2024) possibly exhibits higher sensitivities to ONs than the nitrate-CIMS in our study.

283 3.2 Synergistic reaction efficiencies of different RO_2 species

284 In the $O_3 + NO_3$ regime, synergistic reactions are likely to occur between $^{Cl}RO_2$, $^{OH}RO_2$ and $^{NO_3}RO_2$.
285 Figure 2 shows the $\Delta[\alpha\text{-pinene}]_{O_3}$ -normalized signal ratios of specific C_{10} RO_2 as well as their
286 related $C_xH_yO_z$ -HOM monomers and dimers in the synergistic $O_3 + NO_3$ regime vs. the O_3 -only
287 regime. It should be noted that the second-generation oxidation processes are strongly inhibited by
288 the excess of α -pinene in this study, thus the predominant type of RO_2 observed here is primary RO_2 .
289 Model simulations show that the H-abstraction of α -pinene by OH radicals contributes less than 2%
290 to the formation of $C_{10}H_{15}O_x$ - RO_2 and related HOMs under different experimental conditions
291 (Figure S5). Therefore, $C_{10}H_{15}O_x$ - RO_2 observed in this study are primarily $^{Cl}RO_2$. Notably, the
292 $^{Cl}RO_2$ ($C_{10}H_{15}O_x$) and related $C_{10}H_{14}O_x$ -HOMs decrease by $\sim 20 - 80\%$ in the $O_3 + NO_3$ regime
293 (Figures 2 a, b), while the decreasing extent of $^{OH}RO_2$ ($C_{10}H_{17}O_x$) and related $C_{10}H_{18}O_x$ -HOMs are
294 significantly smaller (0 – 30%). In particular, some of the most oxygenated $C_{10}H_{17}O_x$ - RO_2 and
295 $C_{10}H_{18}O_x$ -HOMs ($x \geq 9$) even increase unexpectedly in the synergistic oxidation regime. For the
296 $C_{10}H_{16}O_x$ -HOMs that can be derived from the self/cross reactions of both $^{Cl}RO_2$ and $^{OH}RO_2$, their
297 reductions are at a medium level. Because of the very small contribution of acyl RO_2 to the total C_{10}
298 RO_2 (0.4%) (Zang et al., 2023), their consumption by NO_2 leads to less than 2% reduction in the
299 C_{10} $^{Cl}RO_2$ signals. Therefore, the more significant decrease in signals of $^{Cl}RO_2$ and related HOMs
300 as compared to the OH-derived ones in the synergistic $O_3 + NO_3$ regime is primarily due to the more
301 efficient cross reactions of $^{NO_3}RO_2$ with $^{Cl}RO_2$ than with $^{OH}RO_2$. Because a large amount of $^{Cl}RO_2$
302 is terminated by $^{NO_3}RO_2$, fewer $^{Cl}RO_2$ are available to terminate $^{OH}RO_2$. As a result, more $^{OH}RO_2$
303 can undergo autoxidation to form highly oxygenated $C_{10}H_{17}O_x$ - RO_2 and $C_{10}H_{18}O_x$ -HOMs ($x \geq 9$),

304 leading to an increase in signals of these species. Consistently, the signals of C₂₀ HOM dimers
 305 decrease by ~20 – 40% in the O₃ + NO₃ regime compared to that in O₃-only regime, and the signal
 306 reduction of dimers (C₂₀H₃₀O_x) formed by ^CIRO₂ is slightly larger than that of the dimers (C₂₀H₃₄O_x)
 307 arising from ^{OH}RO₂ (Figure 2c). Note that the highly oxygenated C₂₀H₃₄O_x dimers (x ≥ 13) that can
 308 be formed from self/cross reactions of C₁₀H₁₇O_x-RO₂ (x ≥ 9) are not observed in this study, likely
 309 due to their low abundance and the limitation of instrument sensitivity.



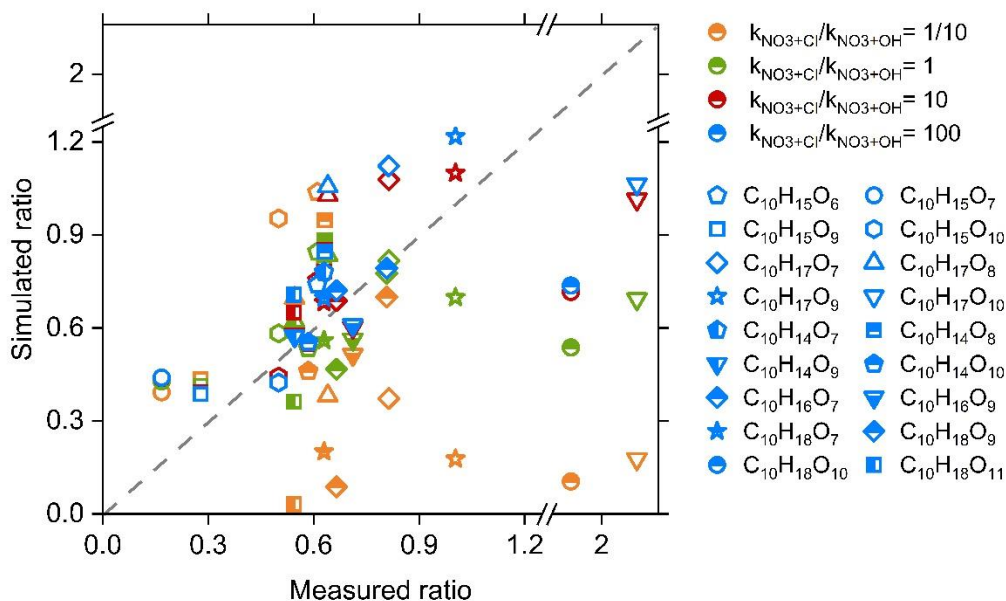
310
 311 Figure 2. Normalized signal ratios of (a) specific and total C₁₀H_{15,17}O_x-RO₂ radicals, as well as their
 312 related (b) C₁₀ HOM monomers and (c) C₂₀ HOM dimers in the O₃ + NO₃ regime vs. the O₃-only
 313 regime (Exps 1-5, 7-11). Ion signals observed in each oxidation regime are normalized to Δ[α-
 314 pinene]_{O₃}.

315 The above results are somewhat different from the most recent study by Li et al. (2024), which
 316 found that the measured C₁₀H₁₅O_x-RO₂ increased slightly with NO₃ radicals while C₁₀H₁₇O_{5,7}-RO₂
 317 from OH chemistry decreased by a factor of 9. Li et al. (2024) indicated that additional C₁₀H₁₅O_x
 318 could be produced from the H-abstraction pathway of NO₃ oxidation of α-pinene. However, in the
 319 monoterpene oxidation system, the rate constant for H-abstraction by NO₃ radicals is (4 – 10) × 10⁻
 320 ¹⁷ cm³ molecule⁻¹ s⁻¹, which is 10³ – 10⁴ times lower than that for the NO₃ addition channel (Martinez
 321 et al., 1998). Besides, the subsequent reactions of RO₂ species formed from H-abstraction by NO₃
 322 radicals should be very similar to those derived from H-abstraction by OH radicals, which was found
 323 not important for C_xH_yO_z-HOM formation in the absence of NO (Zang et al., 2023). Therefore, the
 324 H-abstraction of α-pinene by NO₃ radicals would have negligible influence on C₁₀H₁₅O_x formation.
 325 As Li et al. (2024) used a low α-pinene concentration and relatively high O₃ and NO₃ concentrations
 326 in their experiments, the secondary oxidation of aldehydes, such as the substantially formed
 327 pinonaldehyde, by NO₃ radicals might be important, which could contribute to the additional

328 formation of $C_{10}H_{15}O_x-RO_2$. However, as noted above, the second-generation oxidation processes
329 are strongly inhibited due to the excess of α -pinene in this study, therefore the formation of
330 secondary $C_{10}H_{15}O_x-RO_2$ is not important.

331 In addition, Li et al. (2024) reported that the fraction of α -pinene oxidized by OH radicals decreased
332 from 44% in the O_3 oxidation system to 6% in the $O_3 + NO_3$ system, mainly due to the depletion of
333 OH radicals by NO_2 and the competitive consumption of α -pinene by NO_3 radicals, which resulted
334 in a significant decrease in $C_{10}H_{17}O_{5,7}$ radicals from OH chemistry as observed in their experiments.
335 However, in the present study, because of the excess of α -pinene, over 97% of OH radicals react
336 with α -pinene and the depletion of OH by NO_2 is minor (0.2 – 1.3%) in the $O_3 + NO_3$ regime. The
337 reduction in the reacted α -pinene by OH radicals is less than 10% compared to the O_3 -only regime.
338 As a result, a smaller decrease in $C_{10}H_{17}O_{5,7}$ radicals was observed in our study.

339 To gain quantitative constraints on the relative reaction efficiency of $^{NO_3}RO_2 + ^{Cl}RO_2$ vs. $^{NO_3}RO_2 +$
340 $^{OH}RO_2$ (i.e., k_{NO_3+Cl}/k_{NO_3+OH}), the signal ratios of $C_{10}-^{Cl}RO_2$ and $^{OH}RO_2$ as well as their related C_{10}
341 HOMs in the synergistic oxidation regime vs. the O_3 -only regime were predicted using a kinetic
342 model (see Section 2.3) with different k_{NO_3+Cl}/k_{NO_3+OH} ratios. Figure 3 shows a measurement-model
343 comparison of those signal ratios. When the ratio of k_{NO_3+Cl}/k_{NO_3+OH} is smaller than or equal to 1,
344 the simulated signal ratios of many RO_2 and HOMs differ significantly from the measured ratios,
345 especially for some $C_{10}H_{17}O_x-RO_2$ and $C_{10}H_{18}O_x-HOMs$. When the ratio of k_{NO_3+Cl}/k_{NO_3+OH} is 10 –
346 100, there is a good measurement-model agreement for most of RO_2 and HOMs. Therefore, we
347 conclude that the cross-reaction rate constants of $^{NO_3}RO_2 + ^{Cl}RO_2$ are on average 10 – 100 times
348 larger than those for $^{NO_3}RO_2 + ^{OH}RO_2$. This different RO_2 cross-reaction efficiency is the main
349 reason for the significantly larger decrease in the abundance of $^{Cl}RO_2$ and related HOMs as
350 compared to the OH-derived ones (see Figure 2).



351

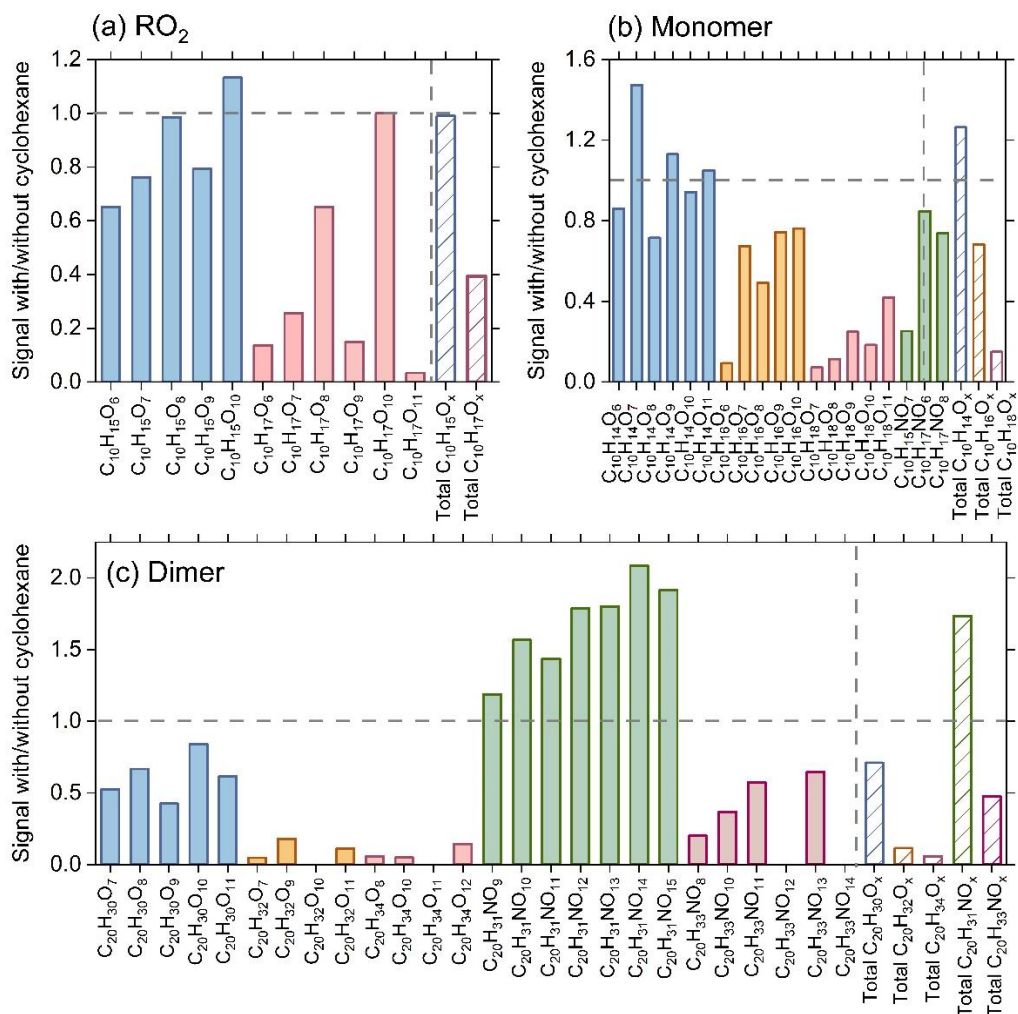
352 Figure 3. Measurement-model comparisons of the signal ratios of different C_{10} RO_2 and HOMs in
 353 the synergistic $O_3 + NO_3$ regime vs. the O_3 -only regime. The cross-reaction rate constant of $^{NO_3}RO_2$
 354 + $^{Cl}RO_2$ was set to $2 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ and the rate of $^{NO_3}RO_2 + ^{OH}RO_2$ was varied from $2 \times$
 355 $10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ to $2 \times 10^{-14} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ in the model.

356 As a competitive reaction pathway, the autoxidation rates of RO_2 can affect the extent to which RO_2
 357 cross reactions influence the RO_2 fate and HOM formation. Therefore, sensitivity analyses of the
 358 autoxidation rate of RO_2 were conducted to evaluate its influence on the changes of RO_2 and related
 359 HOM concentrations in the synergistic $O_3 + NO_3$ regime vs. the O_3 -only regime (Figure S6). In
 360 these analyses, a k_{NO_3+Cl}/k_{NO_3+OH} ratio of 10 was used according to the above discussions. As the
 361 autoxidation rate of $^{OH}RO_2$ increases from 0.28 to 10 s^{-1} , corresponding to the rate range reported in
 362 previous studies (Berndt et al., 2016; Zhao et al., 2018; Xu et al., 2019), the simulated reduction of
 363 highly oxygenated $^{OH}RO_2$ and related $C_{10}H_{18}O_x$ -HOMs in the synergistic $O_3 + NO_3$ regime exhibits
 364 a slight decrease ($< 10\%$) but still agrees reasonably well with the measured value (Figures S6 a-d).
 365 Considering that the autoxidation rates of $^{Cl}RO_2$ used in the model approach their upper limits
 366 reported in the literature, i.e., $\sim 1 \text{ s}^{-1}$ for the butyl ring-opened $C_{10}H_{15}O_4$ - RO_2 (Iyer et al., 2021) and
 367 relatively smaller rates for ring-retained $C_{10}H_{15}O_4$ - RO_2 ($0.02 - 0.29 \text{ s}^{-1}$, see Scheme S1) (Zhao et
 368 al., 2021), we also lowered the autoxidation rate constants of $^{Cl}RO_2$ by a factor of 10 to see its
 369 influence on RO_2 and HOM distribution in the $O_3 + NO_3$ regime. The simulated reduction of $^{Cl}RO_2$
 370 and $C_{10}H_{14}O_x$ -HOMs in this case decreases by 7 – 16% (Figures S6 e-h), while that of $C_{10}H_{16}O_x$ -
 371 HOMs increases by up to 31% (Figures S6 i, j). However, the simulated results are still close to the

372 measured values. These sensitivity analyses suggest that the uncertainty in the autoxidation rates of
373 $^{\text{OH}}\text{RO}_2$ and $^{\text{Cl}}\text{RO}_2$ could slightly affect the simulated distribution of RO_2 and HOMs across different
374 oxidation regimes but not significantly change the $k_{\text{NO}_3+\text{Cl}}/k_{\text{NO}_3+\text{OH}}$ ratio obtained in this study.
375 Further sensitivity analyses on the rate constant and dimer formation branching ratio of RO_2 cross
376 reactions indicate that the uncertainties in these reaction kinetics do not alter the conclusion
377 regarding the $k_{\text{NO}_3+\text{Cl}}/k_{\text{NO}_3+\text{OH}}$ ratio either (see details in Sections S2 and S3).

378 Cyclohexane was added in some experiments as an OH scavenger to elucidate the role of $^{\text{OH}}\text{RO}_2$
379 chemistry in HOM formation in the $\text{O}_3 + \text{NO}_3$ regime. In the presence of cyclohexane, most of
380 $^{\text{OH}}\text{RO}_2$ ($\text{C}_{10}\text{H}_{17}\text{O}_x$) and related HOM monomers ($\text{C}_{10}\text{H}_{18}\text{O}_x$) and dimers ($\text{C}_{20}\text{H}_{32}\text{O}_x$ and $\text{C}_{20}\text{H}_{34}\text{O}_x$)
381 decrease by more than 70% (Figure 4), while $^{\text{Cl}}\text{RO}_2$ ($\text{C}_{10}\text{H}_{15}\text{O}_x$) and related HOM monomers
382 ($\text{C}_{10}\text{H}_{14}\text{O}_x$) only decrease slightly. Accordingly, the reduction in $\text{C}_{20}\text{H}_{32}\text{O}_x$ and $\text{C}_{20}\text{H}_{34}\text{O}_x$ dimers is
383 significantly larger than that of $\text{C}_{20}\text{H}_{30}\text{O}_x$. These results are in a good agreement with previous
384 measurements (Zhao et al., 2018; Zang et al., 2023). The $\text{C}_{10}\text{H}_{16}\text{O}_x$ species, which can arise from
385 both $^{\text{Cl}}\text{RO}_2$ and $^{\text{OH}}\text{RO}_2$, exhibit a medium reduction (Figure 4b). It is interesting to note that with the
386 addition of cyclohexane, there is a significant increase in $\text{C}_{20}\text{H}_{31}\text{NO}_x$, which are formed from the
387 cross reactions of $^{\text{Cl}}\text{RO}_2$ with $^{\text{NO}_3}\text{RO}_2$. Such an enhanced production of $\text{C}_{20}\text{H}_{31}\text{NO}_x$ as compared to
388 the slightly decreased formation of $\text{C}_{20}\text{H}_{30}\text{O}_x$ indicates that the $^{\text{Cl}}\text{RO}_2 + ^{\text{NO}_3}\text{RO}_2$ reactions are
389 competitive compared to the $^{\text{Cl}}\text{RO}_2 + ^{\text{Cl}}\text{RO}_2$ and $^{\text{Cl}}\text{RO}_2 + ^{\text{OH}}\text{RO}_2$ reactions. As a result, when the
390 $^{\text{OH}}\text{RO}_2$ are depleted, the $^{\text{Cl}}\text{RO}_2$ that are supposed to react with $^{\text{OH}}\text{RO}_2$, efficiently react with $^{\text{NO}_3}\text{RO}_2$
391 to form $\text{C}_{20}\text{H}_{31}\text{NO}_x$, leading to the increase in $\text{C}_{20}\text{H}_{31}\text{NO}_x$ signals. Consistent with the experimental
392 measurements, the model simulations show that the concentrations of $\text{C}_{20}\text{H}_{31}\text{NO}_x$ in the $\text{O}_3 + \text{NO}_3$
393 regime increase with the addition of cyclohexane as an OH scavenger (Figure S9). However, the
394 simulated enhancement is slightly lower than the measurements, which might be due to the
395 uncertainties in the RO_2 cross-reaction kinetics in the model.

396



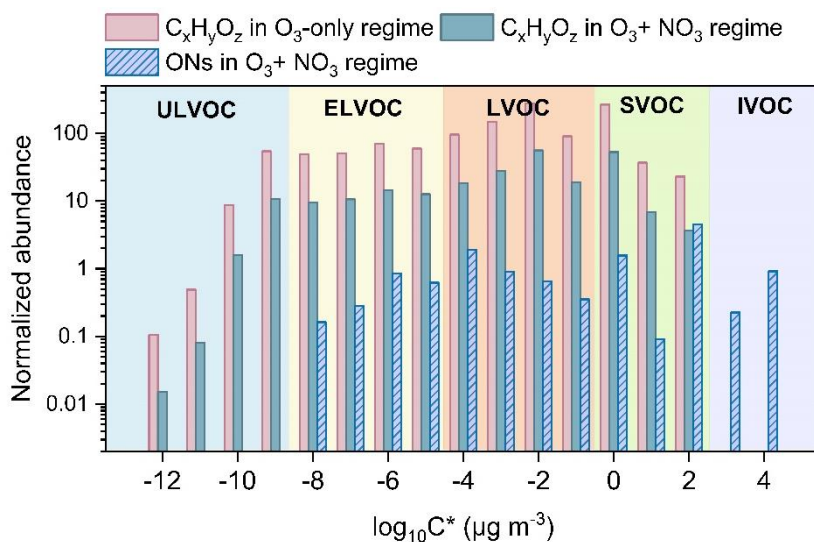
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398 Figure 4. Relative changes in signals of (a) C₁₀ RO₂, (b) C₁₀ HOMs, and (c) C₂₀ dimers due to the
 399 addition of 100 ppm cyclohexane as an OH scavenger derived in the synergistic O₃ + NO₃ regime
 400 (Exps 6 and 12).

401 3.3 Influence of synergistic oxidation on low-volatility organics and particle formation

402 Compared to the O₃-only regime, there are a remarkable reduction in C_xH_yO_z-HOMs and a strong
 403 formation of HOM-ONs due to the efficient cross reactions between ^{NO3}RO₂ and ^{Cl}RO₂ in the
 404 synergistic oxidation regime. This significant change in HOM composition and abundance would
 405 alter the volatility distribution of HOMs and influence the formation of particles. The volatilities of
 406 HOMs formed in the two oxidation regimes are estimated using a modified composition-activity
 407 method (see Section 2.2) and shown in Figure 5. The abundance of C_xH_yO_z-HOMs characterized as
 408 ULVOCs and ELVOCs decreases considerably in the synergistic O₃ + NO₃ regime compared to the
 409 O₃-only regime (Figure 5), in agreement with the very recent observations by Li et al. (2024) who
 410 found that the presence of NO₃ radicals during α-pinene ozonolysis significantly reduced the

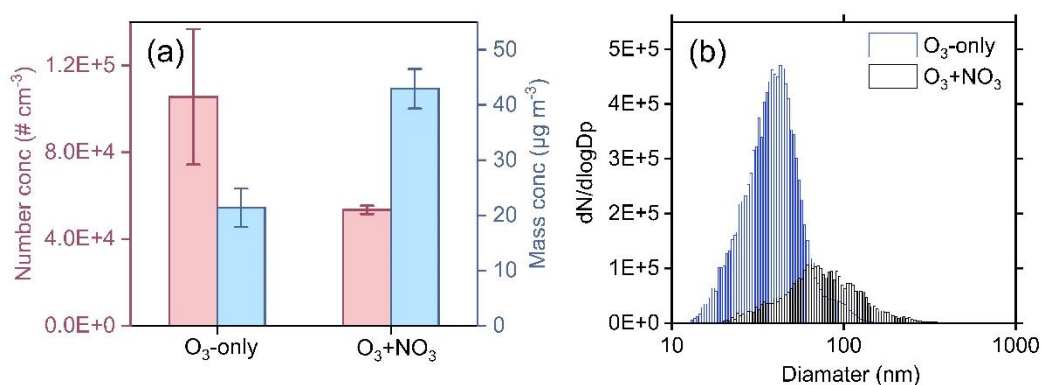
411 abundance of ULVOCs. Although substantial amounts of HOM-ONs are formed in the $O_3 + NO_3$
 412 regime, they generally have higher volatilities (i.e., characterized as ELVOCs to IVOCs) (Figure 5).
 413 Therefore, the synergistic $O_3 + NO_3$ oxidation of α -pinene significantly reduces the formation of
 414 ULVOCs and increases the overall volatility of total HOMs.



415
 416 Figure 5. Volatility distribution of $C_xH_yO_z$ -HOMs and HOM-ONs formed in the $O_3 + NO_3$ regime
 417 and O_3 -only regime (Exps 1, 7). Ion signals in each oxidation regime are normalized to the
 418 corresponding total reacted α -pinene.

419 Figure 6a shows the particle number and mass concentrations formed in the two oxidation regimes
 420 in SOA formation experiments (Exps 13, 14). The particle number concentration decreases by ~50%
 421 whereas the particle mass concentration increases by a factor of 2 in the synergistic $O_3 + NO_3$ regime,
 422 compared to that in the O_3 -only regime. The presence of NO_3 radicals during α -pinene ozonolysis
 423 reduces the abundance of ULVOCs, which are the key species driving particle nucleation, thereby
 424 leading to a reduction in the particle number concentration in the $O_3 + NO_3$ regime. On the other
 425 hand, substantial formation of HOM-ONs is expected from the cross reactions of $^{NO_3}RO_2$ with $^{Cl}RO_2$
 426 and $^{OH}RO_2$ in the synergistic oxidation regime (Li et al., 2024; Bates et al., 2022), although their
 427 signals are relatively low due to the low sensitivity of nitrate-CIMS to ONs in this study. The newly
 428 formed HOM-ONs have relatively higher volatilities and are inefficient in initiating particle
 429 nucleation, but they are able to partition into the formed particles and contribute to the particle mass
 430 growth. Meanwhile, as the particle number concentration decreases drastically in the synergistic
 431 oxidation regime, more condensable vapors are available for each particle to grow to larger sizes
 432 (Figure 6b), which would in turn favor the condensation of more volatile organic species including

433 ONs due to the reduced curvature effect of the larger particles, ultimately resulting in an increase in
434 SOA mass concentrations.



435
436 Figure 6. Number and mass concentrations (a), as well as the size distribution (b) of particles formed
437 from the ozonolysis and synergistic O₃ + NO₃ oxidation of α -pinene (Exps 13-14).

438 Recently, Bates et al. (2022) also found that in chamber experiments with seed particles, the SOA
439 mass yields were significantly higher during α -pinene oxidation by O₃ + NO₃ than during ozonolysis,
440 mainly due to the substantial formation and condensation of dimeric ONs. However, in the absence
441 of seed particles, synergistic O₃ + NO₃ oxidation of α -pinene does not nucleate in their study. This
442 phenomenon might be due to the high concentrations of NO₂ (72 ppb) and O₃ (102 ppb) as well as
443 the relatively low concentration of α -pinene (27 ppb) in their experiments. As indicated by Bates et
444 al. (2022), under this conditions NO₃ radicals were substantially formed and contributed to a
445 dominant fraction (75%) of α -pinene oxidation, which strongly inhibited the production of low-
446 volatility species and particle nucleation.

447 3.4 Atmospheric relevance of experimental results

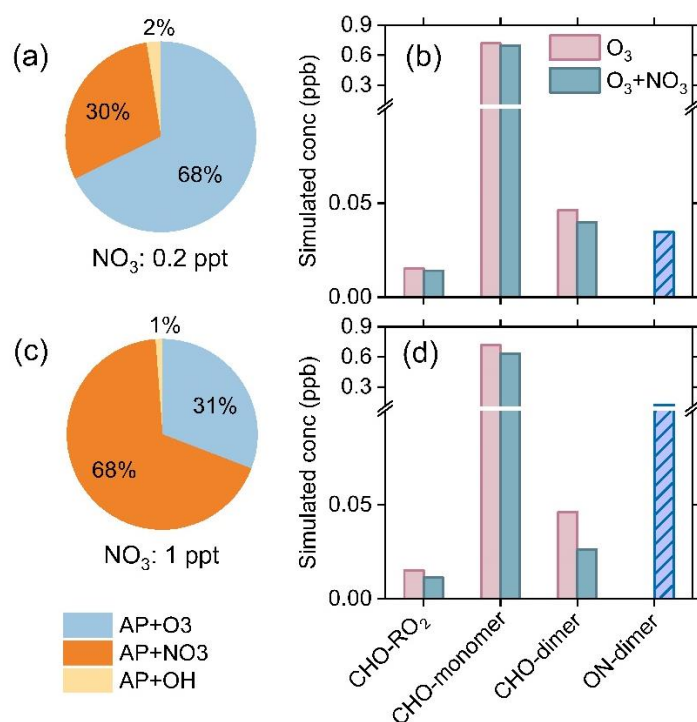
448 In the present study, the flow tube experiments were conducted under dry conditions. Although
449 water vapor may affect the fate of Criegee intermediates and RO₂ radicals and thereby HOM
450 formation during the oxidation of organics under humid conditions, there is growing evidence that
451 such effects in the α -pinene oxidation system are small. Kinetics studies have found that the
452 stabilized Criegee intermediates (SCIs) arising from α -pinene ozonolysis can undergo fast
453 unimolecular decay at a rate constant of 60 – 250 s⁻¹ (Vereecken et al., 2017; Newland et al., 2018),
454 which is rapid compared to their reaction with water vapor, in particular for syn-SCIs, under
455 atmospheric conditions (Vereecken et al., 2017; Newland et al., 2018). In addition, the yield of OH

456 radicals from Criegee decomposition is independent of RH (Atkinson et al., 1992; Aschmann et al.,
457 2002). Consistent with the fast unimolecular reaction kinetics revealed by these studies, recent
458 laboratory measurements have shown that the contribution of SCIs to the formation of gas-phase
459 and particle-phase dimers are small (<20%) during α -pinene ozonolysis (Zhao et al., 2018; Zhao et
460 al., 2022). Furthermore, the molecular composition and abundance of HOM monomers and dimers
461 (Li et al., 2019) and the formation of particle-phase dimers (Zhang et al., 2015; Kenseth et al., 2018)
462 do not change significantly with RH ranging from 3% to 92%. These studies suggest that the
463 humidity condition does not strongly affect the HOM formation chemistry in the α -pinene
464 ozonolysis system.

465 To evaluate the relevance of our experimental findings to the real atmosphere, we performed
466 chemical model simulations of HOM formation from nocturnal synergistic $O_3 + NO_3$ oxidation of
467 α -pinene under typical atmospheric conditions. In these simulations, constant concentrations of α -
468 pinene (1 ppb), O_3 (30 ppb), NO (5 ppt), NO_2 (1.8 ppb), NO_3 radicals (0.2 or 1 ppt), OH radicals (5
469 $- 50 \times 10^4$ molecules cm^{-3}), HO_2 radicals (4 ppt), as well as a constant RH of 50% and temperature
470 of 298 K were used as typical nocturnal conditions in the boreal forest according to the field studies
471 (Stone et al., 2012; Lee et al., 2016a; Brown and Stutz, 2012; Geyer et al., 2003b; Kristensen et al.,
472 2016; Hakola et al., 2012; Liebmann et al., 2018). Considering the rapid deposition of oxidized
473 biogenic compounds (Nguyen et al., 2015), a typical dilution lifetime of 5 h (i.e., $k_{dil} = 1/5 \text{ h}^{-1}$) was
474 assumed in the model. According to the above analysis, the cross-reaction rate constants for NO_3RO_2
475 $+ ClRO_2$ and $NO_3RO_2 + OHRO_2$ were set to $2 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ and $2 \times 10^{-13} \text{ cm}^3 \text{ molecule}^{-1}$
476 s^{-1} in the model, respectively. The formation of RO_2 with oxygen numbers higher than 11 was not
477 considered in the model, due to the large uncertainty in the autoxidation rate constants of the highly
478 oxygenated RO_2 . In fact, the autoxidation rate of the highly oxygenated RO_2 is expected to be small
479 given the significant decrease in the number of active sites for intramolecular H-abstraction in the
480 molecule. As a result, the contribution of the most oxygenated HOMs to the total HOM monomers
481 could be relatively small (Zhao et al., 2018; Clafin et al., 2018).

482 In the absence of NO_3 radicals (with NO_3 concentrations and formation rates set to zero), the amount
483 of α -pinene consumed during 4 hours of simulation is 1.04 ppb. When a relatively low NO_3
484 concentration (0.2 ppt) is considered (Figure 7a), the amount of α -pinene consumed is 1.48 ppb, and

485 the ozonolysis is the primary loss pathway of α -pinene (68%), followed by NO_3 (30%) and OH
486 oxidation (2%). The reactions of $\text{RO}_2 + \text{HO}_2$, $\text{RO}_2 + \text{NO}$, and $\text{RO}_2 + \text{RO}_2$ account for ~49%, ~27%,
487 and ~24% of the total RO_2 fate, respectively (Figure S10a). Compared to the ozonolysis of α -pinene,
488 the synergistic $\text{O}_3 + \text{NO}_3$ oxidation leads to a reduction of 3% and 13% in the formation of $\text{C}_x\text{H}_y\text{O}_z$ -
489 HOM monomers and dimers, respectively (Figure 7b). Given that the concentrations of α -pinene
490 and oxidants were held constant during the simulation, the consumptions of α -pinene by O_3 and OH
491 radicals are the same across different oxidation regimes. Therefore, the decreases in the
492 concentrations of $\text{C}_x\text{H}_y\text{O}_z$ -HOM monomers and dimers in the presence of NO_3 oxidation are mainly
493 due to the cross reactions of $^{\text{NO}_3}\text{RO}_2$ with other RO_2 . When the NO_3 concentration is as high as 1
494 ppt as reported in field studies (Liebmann et al., 2018), the consumption of α -pinene reaches 3.24
495 ppb, of which 68% is contributed by NO_3 oxidation (Figure 7c). Under this condition, the $\text{RO}_2 +$
496 RO_2 reactions account for ~34% of the total RO_2 fate (Figure S10b). As a result, the cross reactions
497 of $^{\text{NO}_3}\text{RO}_2$ with other RO_2 play a more important role in the HOM formation. The production of
498 $\text{C}_x\text{H}_y\text{O}_z$ -HOM monomers and dimers decreases by 12% and 43%, respectively, due to the presence
499 of NO_3 oxidation (Figure 7d). We note that the variation in RH from 0 – 90% in the model has
500 negligible influence on the relative changes in $\text{C}_x\text{H}_y\text{O}_z$ -HOMs under these nocturnal atmospheric
501 conditions (Figure S11). Considering that there are uncertainties in the dilution rate constant, a
502 sensitivity analysis was performed by varying the k_{dil} in the range of 0.04 – 0.2 h^{-1} . It is found that
503 the variation within these rate values does not significantly influence the response of $\text{C}_x\text{H}_y\text{O}_z$ -HOM
504 dimer formation to concurrent NO_3 oxidation (Figure S12).



505

506 Figure 7. Model simulations of α -pinene oxidation and HOM formation under typical nighttime
 507 conditions in the boreal forest. (a, c) Contributions of different loss pathways of α -pinene by
 508 different oxidants at NO_3 concentrations of 0.2 and 1 ppt, respectively; (b, d) Concentrations of
 509 C_xH_yO_z-HOMs and HOM-ONs formed by synergistic O₃ + NO₃ oxidation and ozonolysis of α -
 510 pinene under conditions corresponding to (a) and (c). The simulations were run for 4 h after an 8-h
 511 spin-up for intermediates and secondary species.

512 Field observations have shown that NO_3 radicals, O₃, and OH radicals all had important
 513 contributions to monoterpene oxidation during the early morning after sunrise and late afternoon
 514 before sunset in the southeastern United States (Zhang et al., 2018). In addition, relatively high
 515 nighttime OH concentrations of $(2 - 10) \times 10^5$ molecules cm⁻³ were measured in some areas such
 516 as Germany and New York City (Faloona et al., 2001; Geyer et al., 2003a). As a result, a model
 517 simulation was conducted using a 10 times higher OH concentration (5×10^5 molecules cm⁻³). The
 518 concentration of NO_3 radicals is 1 ppt and the concentrations of other species are the same as the
 519 values mentioned above. With a higher OH concentration, O₃, NO₃, and OH radicals account for
 520 28%, 61%, and 11% to the total α -pinene consumption, respectively (Figure S13 a). Compared to
 521 the results under low OH concentration, the formation of C_xH_yO_z-HOM monomers and dimers are
 522 all enhanced under high OH concentration (Figure S13 b). This is mainly due to the promoted
 523 self/cross reactions of ^{OH}RO₂, as well as the promoted formation of C₁₀H₁₅O_x-RO₂ derived from H-
 524 abstraction pathway by OH radicals. Nevertheless, the presence of NO_3 oxidation still reduces the

525 formation of $C_xH_yO_z$ -HOM dimers by 26% (Figure S13 b).

526 Furthermore, model simulations under typical conditions in the southeastern United States (see
527 details in Section S4) suggest that the coexistence of isoprene appears to exacerbate the suppression
528 effect of synergistic oxidation on HOM formation from monoterpenes. As shown in Figure S14, in
529 the absence of isoprene, the synergistic $O_3 + NO_3$ oxidation of α -pinene leads to a reduction of 13%
530 and 24% in the formation of $C_xH_yO_z$ -HOM monomers and dimers, respectively. When isoprene is
531 present, as the isoprene + NO_3 oxidation produces a significant amount of nitrooxy RO_2 that can
532 also scavenge α -pinene-derived $^{Cl}RO_2$ and $^{OH}RO_2$ via cross reactions, the synergistic oxidation leads
533 to a slightly larger reduction in $C_xH_yO_z$ -HOM monomers and dimers (15% and 31%, respectively).

534 The above model simulations suggest that under nocturnal atmospheric conditions with a very low
535 NO_3 concentration, the RO_2 radical pool is dominated by $^{Cl}RO_2$ and their self/cross reactions are a
536 major contributor to ULVOCs such as the highly oxygenated C_{20} dimers as observed in boreal forest
537 (Bianchi et al., 2017). When the NO_3 concentration is high, the production of $^{NO_3}RO_2$ becomes
538 significant and their cross reactions with $^{Cl}RO_2$ would suppress the formation of ULVOCs. Although
539 HOM-ON dimers are readily produced by cross reactions between $^{NO_3}RO_2$ and $^{Cl}RO_2$, they generally
540 have higher volatilities than $C_xH_yO_z$ -HOM dimers and therefore are less efficient in initiating
541 particle formation. However, these HOM-ONs can be an important contributor to the particle mass
542 growth. As suggested by the model simulations in Bates et al. (2022), the NO_3 oxidation of α -pinene
543 led to a particulate nitrate yield of 7% under nocturnal atmospheric conditions in rural Alabama
544 during the SOAS campaign. Our results offer mechanistic and quantitative insights on how the
545 synergistic oxidation of α -pinene by O_3 and NO_3 radicals can influence the formation of low-
546 volatility organic compounds and hence particle formation and growth. They also provide a potential
547 explanation for field observations that NPF events frequently occur in monoterpene-rich regions
548 during daytime but not at nighttime (Mohr et al., 2017; Kulmala et al., 2001; Junninen et al., 2017).

549 **4. Conclusions**

550 This study provides a comprehensive characterization of the nocturnal synergistic oxidation of α -
551 pinene by O_3 and NO_3 radicals and its influence on the formation of HOMs and low-volatility
552 organic compounds using a combination of flow reactor experiments and detailed kinetic model
553 simulations. It is found that the formation of $C_xH_yO_z$ -HOMs in the $O_3 + NO_3$ regime is significantly

554 suppressed compared to that in the O₃-only regime, mainly due to the depletion of ozonolysis-
555 derived RO₂ (i.e., ^{Cl}RO₂ and ^{OH}RO₂) by ^{NO₃}RO₂ via cross reactions. In addition, the decreases in the
556 abundance of ^{Cl}RO₂ and related HOMs are significantly larger than those of OH-derived ones,
557 indicating that the ^{NO₃}RO₂ species react more efficiently with ^{Cl}RO₂ than with ^{OH}RO₂. Detailed
558 measurement-model comparisons for the distribution of a suite of ^{Cl}RO₂, ^{OH}RO₂, and associated
559 HOMs across different oxidation regimes further reveal that the cross reactions between ^{Cl}RO₂ and
560 ^{NO₃}RO₂ are averagely 10 – 100 times more efficient than those of ^{OH}RO₂ and ^{NO₃}RO₂.

561 The suppressed formation of C_xH_yO_z-HOMs in the synergistic O₃ + NO₃ regime results in a
562 significant reduction in ULVOCs. Although substantial amounts of HOM-ONs are formed from the
563 cross reactions between ^{NO₃}RO₂ and ^{Cl}RO₂ or ^{OH}RO₂ in the synergistic oxidation regime, they have
564 higher volatilities and are less likely to participate in the formation and initial growth of new
565 particles. As a result, in our experiment the formation of new particles in the synergistic oxidation
566 regime is substantially inhibited compared to the O₃-only regime. Chemical model simulations
567 further confirm that the synergistic oxidation of α-pinene by O₃ and NO₃ radicals can significantly
568 inhibit the formation of C_xH_yO_z-HOMs, especially the ultra-low volatility C_xH_yO_z-HOM dimers
569 under typical nighttime atmospheric conditions. Our study sheds lights on the synergistic oxidation
570 mechanism of biogenic emissions and underscores the importance of considering this chemistry for
571 a better depiction of the formation of low-volatility organics and particles in the atmosphere.

572

573 *Data availability.* The data presented in this work are available upon request from the corresponding
574 author.

575

576 *Author contributions.* YZ and HZ designed the study, HZ and DH performed the experiments. YZ
577 and HZ analyzed the data, conducted model simulations, and wrote the paper. All other authors
578 contributed to discussion and writing.

579 *Competing interests.* The authors declare no conflict of interest.

580

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585 **References**

- 586 Aschmann, S. M., Arey, J., and Atkinson, R.: OH radical formation from the gas-phase reactions of O₃
587 with a series of terpenes, *Atmos. Environ.*, 36, 4347-4355, [https://doi.org/10.1016/S1352-](https://doi.org/10.1016/S1352-2310(02)00355-2)
588 2310(02)00355-2, 2002.
- 589 Atkinson, R., Aschmann, S. M., Arey, J., and Shorees, B.: Formation of OH radicals in the gas-phase
590 reactions of O₃ with a series of terpenes, *J. Geophys. Res.-Atmos.*, 97, 6065-6073,
591 <https://doi.org/10.1029/92JD00062>, 1992.
- 592 Ayres, B. R., Allen, H. M., Draper, D. C., Brown, S. S., Wild, R. J., Jimenez, J. L., Day, D. A.,
593 Campuzano-Jost, P., Hu, W., de Gouw, J., Koss, A., Cohen, R. C., Duffey, K. C., Romer, P.,
594 Baumann, K., Edgerton, E., Takahama, S., Thornton, J. A., Lee, B. H., Lopez-Hilfiker, F. D., Mohr,
595 C., Wennberg, P. O., Nguyen, T. B., Teng, A., Goldstein, A. H., Olson, K., and Fry, J. L.: Organic
596 nitrate aerosol formation via NO₃ + biogenic volatile organic compounds in the southeastern United
597 States, *Atmos. Chem. Phys.*, 15, 13377-13392, <https://doi.org/10.5194/acp-15-13377-2015>, 2015.
- 598 Bates, K. H., Burke, G. J. P., Cope, J. D., and Nguyen, T. B.: Secondary organic aerosol and organic
599 nitrogen yields from the nitrate radical (NO₃) oxidation of alpha-pinene from various RO₂ fates,
600 *Atmos. Chem. Phys.*, 22, 1467-1482, <https://doi.org/10.5194/acp-22-1467-2022>, 2022.
- 601 Berndt, T.: Peroxy radical processes and product formation in the OH radical-initiated oxidation of alpha-
602 pinene for near-atmospheric conditions, *J. Phys. Chem. A*, 125, 9151-9160,
603 <https://doi.org/10.1021/acs.jpca.1c05576>, 2021.
- 604 Berndt, T., Mentler, B., Scholz, W., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion
605 product formation from ozonolysis and OH radical reaction of alpha-pinene: mechanistic insight
606 and the influence of isoprene and ethylene, *Environ. Sci. Technol.*, 52, 11069-11077,
607 <https://doi.org/10.1021/acs.est.8b02210>, 2018.
- 608 Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G., Stratmann,
609 F., Herrmann, H., Sipilä, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of
610 highly oxidized organic compounds, *Nat. Commun.*, 7, <https://doi.org/10.1038/ncomms13677>,
611 2016.
- 612 Bianchi, F., Garmash, O., He, X. C., Yan, C., Iyer, S., Rosendahl, I., Xu, Z. N., Rissanen, M. P., Riva, M.,
613 Taipale, R., Sarnela, N., Petäjä, T., Worsnop, D. R., Kulmala, M., Ehn, M., and Junninen, H.: The
614 role of highly oxygenated molecules (HOMs) in determining the composition of ambient ions in the
615 boreal forest, *Atmos. Chem. Phys.*, 17, 13819-13831, <https://doi.org/10.5194/acp-17-13819-2017>,
616 2017.
- 617 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crouse, J. D.,
618 Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R.,
619 Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly oxygenated organic molecules
620 (HOM) from gas-phase autoxidation involving peroxy radicals: a key contributor to atmospheric
621 aerosol, *Chem. Rev.*, 119, 3472-3509, <https://doi.org/10.1021/acs.chemrev.8b00395>, 2019.
- 622 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.:
623 Secondary organic aerosol formation from the β-pinene+NO₃ system: effect of humidity and peroxy
624 radical fate, *Atmos. Chem. Phys.*, 15, 7497-7522, <https://doi.org/10.5194/acp-15-7497-2015>, 2015.
- 625 Brown, S. S. and Stutz, J.: Nighttime radical observations and chemistry, *Chem. Soc. Rev.*, 41,
626 <https://doi.org/10.1039/c2cs35181a>, 2012.
- 627 Claffin, M. S., Krechmer, J. E., Hu, W., Jimenez, J. L., and Ziemann, P. J.: Functional group composition

628 of secondary organic aerosol formed from ozonolysis of α -pinene under high VOC and autoxidation
629 conditions, ACS Earth Space Chem., 2, 1196-1210,
630 <https://doi.org/10.1021/acsearthspacechem.8b00117>, 2018.

631 Daumit, K. E., Kessler, S. H., and Kroll, J. H.: Average chemical properties and potential formation
632 pathways of highly oxidized organic aerosol, Faraday Discuss., 165,
633 <https://doi.org/10.1039/c3fd00045a>, 2013.

634 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis
635 set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11, 3303-3318,
636 <https://doi.org/10.5194/acp-11-3303-2011>, 2011.

637 Donahue, N. M., Henry, K. M., Mentel, T. F., Kiendler-Scharr, A., Spindler, C., Bohn, B., Brauers, T.,
638 Dorn, H. P., Fuchs, H., Tillmann, R., Wahner, A., Saathoff, H., Naumann, K.-H., Möhler, O., Leisner,
639 T., Müller, L., Reinnig, M.-C., Hoffmann, T., Salo, K., Hallquist, M., Frosch, M., Bilde, M.,
640 Tritscher, T., Barnet, P., Praplan, A. P., DeCarlo, P. F., Dommen, J., Prévôt, A. S. H., and
641 Baltensperger, U.: Aging of biogenic secondary organic aerosol via gas-phase OH radical reactions,
642 P. Natl. Acad. Sci. USA, 109, 13503-13508, <https://doi.org/10.1073/pnas.1115186109>, 2012.

643 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
644 Tillmann, R., and Lee, B.: A large source of low-volatility secondary organic aerosol, Nature, 506,
645 476-479, <https://doi.org/10.1038/nature13032>, 2014.

646 Faloon, I., Tan, D., Brune, W., Hurst, J., Barkot, D., Couch, T. L., Shepson, P., Apel, E., Riemer, D.,
647 Thornberry, T., Carroll, M. A., Sillman, S., Keeler, G. J., Sagady, J., Hooper, D., and Paterson, K.:
648 Nighttime observations of anomalously high levels of hydroxyl radicals above a deciduous forest
649 canopy, J. Geophys. Res.-Atmos., 106, 24315-24333, <https://doi.org/10.1029/2000JD900691>, 2001.

650 Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkler, P. M., Lawler, M. J., Brown,
651 S. S., Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary organic aerosol formation and organic
652 nitrate yield from NO₃ oxidation of biogenic hydrocarbons, Environ. Sci. Technol., 48, 11944-11953,
653 <https://doi.org/10.1021/es502204x>, 2014.

654 Geyer, A., Bächmann, K., Hofzumahaus, A., Holland, F., Konrad, S., Klüpfel, T., Pätz, H. W., Perner, D.,
655 Mihelcic, D., Schäfer, H. J., Volz-Thomas, A., and Platt, U.: Nighttime formation of peroxy and
656 hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies, J. Geophys.
657 Res.-Atmos., 108, <https://doi.org/10.1029/2001JD000656>, 2003a.

658 Geyer, A., Bächmann, K., Hofzumahaus, A., Holland, F., Konrad, S., Klüpfel, T., Pätz, H. W., Perner, D.,
659 Mihelcic, D., Schäfer, H. J., Volz-Thomas, A., and Platt, U.: Nighttime formation of peroxy and
660 hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies, J. Geophys.
661 Res.-Atmos., 108, <https://doi.org/10.1029/2001JD000656>, 2003b.

662 Hakola, H., Hellén, H., Hemmilä, M., Rinne, J., and Kulmala, M.: In situ measurements of volatile
663 organic compounds in a boreal forest, Atmos. Chem. Phys., 12, 11665-11678,
664 <https://doi.org/10.5194/acp-12-11665-2012>, 2012.

665 Hallquist, M., Wängberg, I., Ljungström, E., Barnes, I., and Becker, K. H.: Aerosol and product yields
666 from NO₃ radical-initiated oxidation of selected monoterpenes, Environ. Sci. Technol., 33, 553-559,
667 <https://doi.org/10.1021/es980292s>, 1999.

668 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G.,
669 Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G.,
670 Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z.,
671 Szidat, S., Baltensperger, U., El Haddad, I., and Prevot, A. S.: High secondary aerosol contribution

672 to particulate pollution during haze events in China, *Nature*, 514, 218-222, 10.1038/nature13774,
673 2014.

674 Huang, W., Saathoff, H., Shen, X., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical characterization
675 of highly functionalized organonitrates contributing to night-time organic aerosol mass loadings and
676 particle growth, *Environ. Sci. Technol.*, 53, 1165-1174, <https://doi.org/10.1021/acs.est.8b05826>,
677 2019.

678 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.: Modeling
679 the charging of highly oxidized cyclohexene ozonolysis products using nitrate-based chemical
680 ionization, *J. Phys. Chem. A*, 119, 6339-6345, <https://doi.org/10.1021/acs.jpca.5b01818>, 2015.

681 Inomata, S.: New particle formation promoted by OH reactions during α -pinene ozonolysis, *ACS Earth*
682 *Space Chem.*, 5, 1929-1933, <https://doi.org/10.1021/acsearthspacechem.1c00142>, 2021.

683 Isaacman-VanWertz, G. and Aumont, B.: Impact of organic molecular structure on the estimation of
684 atmospherically relevant physicochemical parameters, *Atmos. Chem. Phys.*, 21, 6541-6563,
685 <https://doi.org/10.5194/acp-21-6541-2021>, 2021.

686 Iyer, S., Rissanen, M. P., Valiev, R., Barua, S., Krechmer, J. E., Thornton, J., Ehn, M., and Kurten, T.:
687 Molecular mechanism for rapid autoxidation in alpha-pinene ozonolysis, *Nat. Commun.*, 12, 878,
688 <https://doi.org/10.1038/s41467-021-21172-w>, 2021.

689 Jenkin, M., Young, J., and Rickard, A.: The MCM v3.3.1 degradation scheme for isoprene, *Atmos. Chem.*
690 *Phys.*, 15, 11433-11459, <https://doi.org/10.5194/acp-15-11433-2015>, 2015.

691 Jokinen, T., Sipilä, M., Richters, S., Kerminen, V. M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala,
692 M., Ehn, M., and Herrmann, H.: Rapid autoxidation forms highly oxidized RO₂ radicals in the
693 atmosphere, *Angew. Chem. Int. Edit.*, 53, 14596-14600, <https://doi.org/10.1002/anie.201408566>,
694 2014.

695 Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M.,
696 Fuhrer, K., and Kulmala, M.: A high-resolution mass spectrometer to measure atmospheric ion
697 composition, *Atmos. Meas. Tech.*, 3, 1039-1053, <https://doi.org/10.5194/amt-3-1039-2010>, 2010.

698 Junninen, H., Hulkkonen, M., Riipinen, I., Nieminen, T., Hirsikko, A., Suni, T., Boy, M., Lee, S.-H., Vana,
699 M., Tammet, H., Kerminen, V.-M., and Kulmala, M.: Observations on nocturnal growth of
700 atmospheric clusters, *Tellus B: Chemical and Physical Meteorology*, 60, 365-371,
701 <https://doi.org/10.1111/j.1600-0889.2008.00356.x>, 2017.

702 Kenseth, C. M., Huang, Y., Zhao, R., Dalleska, N. F., Hethcox, J. C., Stoltz, B. M., and Seinfeld, J. H.:
703 Synergistic O₃ + OH oxidation pathway to extremely low-volatility dimers revealed in beta-pinene
704 secondary organic aerosol, *P. Natl. Acad. Sci. USA*, 115, 8301-8306,
705 <https://doi.org/10.1073/pnas.1804671115>, 2018.

706 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M.,
707 Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A.,
708 Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A.,
709 Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T.,
710 Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K.,
711 Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P.,
712 Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N.,
713 Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F.,
714 Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D.,
715 Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,

716 Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure
717 biogenic particles, *Nature*, 533, 521-526, <https://doi.org/10.1038/nature17953>, 2016.

718 Kristensen, K., Watne, Å. K., Hammes, J., Lutz, A., Petäjä, T., Hallquist, M., Bilde, M., and Glasius, M.:
719 High-molecular weight dimer esters are major products in aerosols from α -pinene ozonolysis and
720 the boreal forest, *Environ. Sci. Tech. Let.*, 3, 280-285, <https://doi.org/10.1021/acs.estlett.6b00152>,
721 2016.

722 Kulmala, M., Hämeri, K., Aalto, P. P., Mäkelä, J. M., Pirjola, L., Nilsson, E. D., Buzorius, G., Rannik,
723 Ü., Dal Maso, M., Seidl, W., Hoffman, T., Janson, R., Hansson, H. C., Viisanen, Y., Laaksonen, A.,
724 and O'Dowd, C. D.: Overview of the international project on biogenic aerosol formation in the
725 boreal forest (BIOFOR), *Tellus Series B-Chemical and Physical Meteorology*, 53, 324-343,
726 <https://doi.org/10.1034/j.1600-0889.2001.530402.x>, 2001.

727 Kurtén, T., Møller, K. H., Nguyen, T. B., Schwantes, R. H., Misztal, P. K., Su, L., Wennberg, P. O., Fry,
728 J. L., and Kjaergaard, H. G.: Alkoxy radical bond scissions explain the anomalously low secondary
729 organic aerosol and organonitrate yields from α -pinene + NO₃, *J. Phys. Chem. L*, 8, 2826-2834,
730 <https://doi.org/10.1021/acs.jpcllett.7b01038>, 2017.

731 Lee, B. H., D'Ambro, E. L., Lopez-Hilfiker, F. D., Schobesberger, S., Mohr, C., Zawadowicz, M. A., Liu,
732 J., Shilling, J. E., Hu, W., Palm, B. B., Jimenez, J. L., Hao, L., Virtanen, A., Zhang, H., Goldstein,
733 A. H., Pye, H. O. T., and Thornton, J. A.: Resolving ambient organic aerosol formation and aging
734 pathways with simultaneous molecular composition and volatility observations, *ACS Earth Space*
735 *Chem.*, 4, 391-402, <https://doi.org/10.1021/acsearthspacechem.9b00302>, 2020.

736 Lee, S. H., Uin, J., Guenther, A. B., de Gouw, J. A., Yu, F., Nadykto, A. B., Herb, J., Ng, N. L., Koss, A.,
737 Brune, W. H., Baumann, K., Kanawade, V. P., Keutsch, F. N., Nenes, A., Olsen, K., Goldstein, A.,
738 and Ouyang, Q.: Isoprene suppression of new particle formation: Potential mechanisms and
739 implications, *J. Geophys. Res.-Atmos.*, 121, <https://doi.org/10.1002/2016jd024844>, 2016a.

740 Lee, S. H., Uin, J., Guenther, A. B., de Gouw, J. A., Yu, F. Q., Nadykto, A. B., Herb, J., Ng, N. L., Koss,
741 A., Brune, W. H., Baumann, K., Kanawade, V. P., Keutsch, F. N., Nenes, A., Olsen, K., Goldstein,
742 A., and Ouyang, Q.: Isoprene suppression of new particle formation: Potential mechanisms and
743 implications, *J. Geophys. Res.-Atmos.*, 121, 14621-14635, <https://doi.org/10.1002/2016jd024844>,
744 2016b.

745 Li, D., Huang, W., Wang, D., Wang, M., Thornton, J. A., Caudillo, L., Rörup, B., Marten, R., Scholz, W.,
746 Finkenzeller, H., Marie, G., Baltensperger, U., Bell, D. M., Brasseur, Z., Curtius, J., Dada, L.,
747 Duplissy, J., Gong, X., Hansel, A., He, X.-C., Hofbauer, V., Junninen, H., Krechmer, J. E., Kürten,
748 A., Lamkaddam, H., Lehtipalo, K., Lopez, B., Ma, Y., Mahfouz, N. G. A., Manninen, H. E., Mentler,
749 B., Perrier, S., Petäjä, T., Pfeifer, J., Philippov, M., Schervish, M., Schobesberger, S., Shen, J., Surdu,
750 M., Tomaz, S., Volkamer, R., Wang, X., Weber, S. K., Welti, A., Worsnop, D. R., Wu, Y., Yan, C.,
751 Zauner-Wieczorek, M., Kulmala, M., Kirkby, J., Donahue, N. M., George, C., El-Haddad, I.,
752 Bianchi, F., and Riva, M.: Nitrate radicals suppress biogenic new particle formation from
753 monoterpene oxidation, *Environ. Sci. Technol.*, 58, 1601-1614,
754 <https://doi.org/10.1021/acs.est.3c07958>, 2024.

755 Li, X. X., Chee, S., Hao, J. M., Abbatt, J. P. D., Jiang, J. K., and Smith, J. N.: Relative humidity effect
756 on the formation of highly oxidized molecules and new particles during monoterpene oxidation,
757 *Atmos. Chem. Phys.*, 19, 1555-1570, <https://doi.org/10.5194/acp-19-1555-2019>, 2019.

758 Li, Y., Pöschl, U., and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in the
759 chemical evolution of organic aerosols, *Atmos. Chem. Phys.*, 16, 3327-3344,

760 <https://doi.org/10.5194/acp-16-3327-2016>, 2016.

761 Liebmann, J., Karu, E., Sobanski, N., Schuladen, J., Ehn, M., Schallhart, S., Quéléver, L., Hellen, H.,
762 Hakola, H., Hoffmann, T., Williams, J., Fischer, H., Lelieveld, J., and Crowley, J. N.: Direct
763 measurement of NO₃ radical reactivity in a boreal forest, *Atmos. Chem. Phys.*, 18, 3799-3815,
764 <https://doi.org/10.5194/acp-18-3799-2018>, 2018.

765 Liu, J., D'Ambro, E. L., Lee, B. H., Schobesberger, S., Bell, D. M., Zaveri, R. A., Zelenyuk, A., Thornton,
766 J. A., and Shilling, J. E.: Monoterpene photooxidation in a continuous-flow chamber: SOA yields
767 and impacts of oxidants, NO(x), and VOC precursors, *Environ. Sci. Technol.*, 56, 12066-12076,
768 <https://doi.org/10.1021/acs.est.2c02630>, 2022.

769 Martinez, E., Cabanas, B., Aranda, A., and Martin, P.: Kinetics of the reactions of NO₃ radical with
770 selected monoterpenes: A temperature dependence study, *Environ. Sci. Technol.*, 32, 3730-3734,
771 <https://doi.org/10.1021/es970899t>, 1998.

772 Mentel, T., Springer, M., Ehn, M., Kleist, E., Pullinen, I., Kurtén, T., Rissanen, M., Wahner, A., and Wildt,
773 J.: Formation of highly oxidized multifunctional compounds: autoxidation of peroxy radicals
774 formed in the ozonolysis of alkenes—deduced from structure—product relationships, *Atmos. Chem.*
775 *Phys.*, 15, 6745-6765, <https://doi.org/10.5194/acp-15-6745-2015>, 2015.

776 Mohr, C., Lopez-Hilfiker, F. D., Yli-Juuti, T., Heitto, A., Lutz, A., Hallquist, M., D'Ambro, E. L.,
777 Rissanen, M. P., Hao, L., Schobesberger, S., Kulmala, M., Mauldin, R. L., Makkonen, U., Sipilä,
778 M., Petäjä, T., and Thornton, J. A.: Ambient observations of dimers from terpene oxidation in the
779 gas phase: Implications for new particle formation and growth, *Geophysical Research Letters*, 44,
780 2958-2966, [10.1002/2017gl072718](https://doi.org/10.1002/2017gl072718), 2017.

781 Molteni, U., Simon, M., Heinritzi, M., Hoyle, C. R., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M.,
782 Brilke, S., Dias, A., Duplissy, J., Frege, C., Gordon, H., Heyn, C., Jokinen, T., Kürten, A., Lehtipalo,
783 K., Makhmutov, V., Petäjä, T., Pieber, S. M., Praplan, A. P., Schobesberger, S., Steiner, G., Stozhkov,
784 Y., Tomé, A., Tröstl, J., Wagner, A. C., Wagner, R., Williamson, C., Yan, C., Baltensperger, U.,
785 Curtius, J., Donahue, N. M., Hansel, A., Kirkby, J., Kulmala, M., Worsnop, D. R., and Dommen, J.:
786 Formation of highly oxygenated organic molecules from α -pinene ozonolysis: chemical
787 characteristics, mechanism, and kinetic model development, *ACS Earth Space Chem.*, 3, 873-883,
788 <https://doi.org/10.1021/acsearthspacechem.9b00035>, 2019.

789 Mutzel, A., Zhang, Y., Böge, O., Rodigast, M., Kolodziejczyk, A., Wang, X., and Herrmann, H.:
790 Importance of secondary organic aerosol formation of α -pinene, limonene, and m-cresol comparing
791 day- and nighttime radical chemistry, *Atmos. Chem. Phys.*, 21, 8479-8498,
792 <https://doi.org/10.5194/acp-21-8479-2021>, 2021.

793 Newland, M. J., Rickard, A. R., Sherwen, T., Evans, M. J., Vereecken, L., Muñoz, A., Ródenas, M., and
794 Bloss, W. J.: The atmospheric impacts of monoterpene ozonolysis on global stabilised Criegee
795 intermediate budgets and SO₂ oxidation: experiment, theory and modelling, *Atmos. Chem. Phys.*,
796 18, 6095-6120, <https://doi.org/10.5194/acp-18-6095-2018>, 2018.

797 Nguyen, T. B., Crounse, J. D., Teng, A. P., St. Clair, J. M., Paulot, F., Wolfe, G. M., and Wennberg, P. O.:
798 Rapid deposition of oxidized biogenic compounds to a temperate forest, *P. Natl. Acad. Sci. USA*,
799 112, <https://doi.org/10.1073/pnas.1418702112>, 2015.

800 Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Greaves, J., and Finlayson-Pitts*, B. J.:
801 Identification of organic nitrates in the NO₃ radical initiated oxidation of α -pinene by atmospheric
802 pressure chemical ionization mass spectrometry, *Environ. Sci. Technol.*, 44, 5887-5893,
803 <https://doi.org/10.1021/es1005658>, 2010.

804 Pye, H. O. T., Ward-Caviness, C. K., Murphy, B. N., Appel, K. W., and Seltzer, K. M.: Secondary organic
805 aerosol association with cardiorespiratory disease mortality in the United States, *Nat. Commun.*, 12,
806 <https://doi.org/10.1038/s41467-021-27484-1>, 2021.

807 Schervish, M. and Donahue, N. M.: Peroxy radical chemistry and the volatility basis set, *Atmos. Chem.*
808 *Phys.*, 20, 1183-1199, <https://doi.org/10.5194/acp-20-1183-2020>, 2020.

809 Shen, H., Vereecken, L., Kang, S., Pullinen, I., Fuchs, H., Zhao, D., and Mentel, T. F.: Unexpected
810 significance of a minor reaction pathway in daytime formation of biogenic highly oxygenated
811 organic compounds, *Sci. Adv.*, 8, eabp8702, <https://doi.org/10.1126/sciadv.abp8702>, 2022.

812 Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L., Kuang, C.,
813 Laskin, A., Martin, S. T., Ng, N. L., Petaja, T., Pierce, J. R., Rasch, P. J., Roldin, P., Seinfeld, J. H.,
814 Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J., Worsnop, D. R., Zaveri, R. A.,
815 Zelenyuk, A., and Zhang, Q.: Recent advances in understanding secondary organic aerosol:
816 Implications for global climate forcing, *Reviews of Geophysics*, 55, 509-559,
817 [10.1002/2016rg000540](https://doi.org/10.1002/2016rg000540), 2017.

818 Simon, M., Dada, L., Heinritzi, M., Scholz, W., Stolzenburg, D., Fischer, L., Wagner, A. C., Kürten, A.,
819 Rörup, B., He, X.-C., Almeida, J., Baalbaki, R., Baccarini, A., Bauer, P. S., Beck, L., Bergen, A.,
820 Bianchi, F., Bräkling, S., Brilke, S., Caudillo, L., Chen, D., Chu, B., Dias, A., Draper, D. C., Duplissy,
821 J., El-Haddad, I., Finkenzeller, H., Frege, C., Gonzalez-Carracedo, L., Gordon, H., Granzin, M.,
822 Hakala, J., Hofbauer, V., Hoyle, C. R., Kim, C., Kong, W., Lamkaddam, H., Lee, C. P., Lehtipalo,
823 K., Leiminger, M., Mai, H., Manninen, H. E., Marie, G., Marten, R., Mentler, B., Molteni, U.,
824 Nichman, L., Nie, W., Ojdanic, A., Onnela, A., Partoll, E., Petäjä, T., Pfeifer, J., Philippov, M.,
825 Quéléver, L. L. J., Ranjithkumar, A., Rissanen, M. P., Schallhart, S., Schobesberger, S., Schuchmann,
826 S., Shen, J., Sipilä, M., Steiner, G., Stozhkov, Y., Tauber, C., Tham, Y. J., Tomé, A. R., Vazquez-
827 Pufleau, M., Vogel, A. L., Wagner, R., Wang, M., Wang, D. S., Wang, Y., Weber, S. K., Wu, Y., Xiao,
828 M., Yan, C., Ye, P., Ye, Q., Zauner-Wieczorek, M., Zhou, X., Baltensperger, U., Dommen, J., Flagan,
829 R. C., Hansel, A., Kulmala, M., Volkamer, R., Winkler, P. M., Worsnop, D. R., Donahue, N. M.,
830 Kirkby, J., and Curtius, J.: Molecular understanding of new-particle formation from α -pinene
831 between -50 and $+25$ °C, *Atmos. Chem. Phys.*, 20, 9183-9207, [https://doi.org/10.5194/acp-20-](https://doi.org/10.5194/acp-20-9183-2020)
832 [9183-2020](https://doi.org/10.5194/acp-20-9183-2020), 2020.

833 Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and HO₂ radicals: field measurements and
834 model comparisons, *Chem. Soc. Rev.*, 41, <https://doi.org/10.1039/c2cs35140d>, 2012.

835 Vereecken, L., Novelli, A., and Taraborrelli, D.: Unimolecular decay strongly limits the atmospheric
836 impact of Criegee intermediates, *Phys. Chem. Chem. Phys.*, 19, 31599-31612,
837 <https://doi.org/10.1039/c7cp05541b>, 2017.

838 Wang, Y., Zhao, Y., Li, Z., Li, C., Yan, N., and Xiao, H.: Importance of hydroxyl radical chemistry in
839 isoprene suppression of particle formation from α -pinene ozonolysis, *ACS Earth Space Chem.*, 5,
840 487-499, <https://doi.org/10.1021/acsearthspacechem.0c00294>, 2021.

841 Wolfe, G. M., Marvin, M. R., Roberts, S. J., Travis, K. R., and Liao, J.: The framework for 0-D
842 atmospheric modeling (F0AM) v3. 1, *Geosci. Model Dev.*, 9, 3309-3319,
843 <https://doi.org/10.5194/gmd-9-3309-2016>, 2016.

844 Xu, L., Møller, K. H., Crounse, J. D., Otkjær, R. V., Kjaergaard, H. G., and Wennberg, P. O.:
845 Unimolecular reactions of peroxy radicals formed in the oxidation of α -pinene and β -pinene by
846 hydroxyl radicals, *J. Phys. Chem. A*, 123, 1661-1674, <https://doi.org/10.1021/acs.jpca.8b11726>,
847 2019.

848 Zang, H., Huang, D., Zhong, J., Li, Z., Li, C., Xiao, H., and Zhao, Y.: Direct probing of acylperoxy
849 radicals during ozonolysis of α -pinene: constraints on radical chemistry and production of highly
850 oxygenated organic molecules, *Atmos. Chem. Phys.*, 23, 12691-12705, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-23-12691-2023)
851 [23-12691-2023](https://doi.org/10.5194/acp-23-12691-2023), 2023.

852 Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J. H.,
853 Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K.
854 S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Weber,
855 R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz,
856 A., Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A., and Goldstein,
857 A. H.: Monoterpenes are the largest source of summertime organic aerosol in the southeastern
858 United States, *P. Natl. Acad. Sci. USA*, 115, 2038-2043, <https://doi.org/10.1073/pnas.1717513115>,
859 2018.

860 Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.:
861 Formation and evolution of molecular products in α -pinene secondary organic aerosol, *P. Natl. Acad.*
862 *Sci. USA*, 112, 14168-14173, <https://doi.org/10.1073/pnas.1517742112>, 2015.

863 Zhang, Y., Peräkylä, O., Yan, C., Heikkinen, L., Äijälä, M., Daellenbach, K. R., Zha, Q., Riva, M.,
864 Garmash, O., Junninen, H., Paatero, P., Worsnop, D., and Ehn, M.: Insights into atmospheric
865 oxidation processes by performing factor analyses on subranges of mass spectra, *Atmos. Chem.*
866 *Phys.*, 20, 5945-5961, <https://doi.org/10.5194/acp-20-5945-2020>, 2020.

867 Zhao, Y., Thornton, J. A., and Pye, H. O. T.: Quantitative constraints on autoxidation and dimer formation
868 from direct probing of monoterpene-derived peroxy radical chemistry, *P. Natl. Acad. Sci. USA*, 115,
869 12142-12147, <https://doi.org/10.1073/pnas.1812147115>, 2018.

870 Zhao, Y., Yao, M., Wang, Y., Li, Z., Wang, S., Li, C., and Xiao, H.: Acylperoxy radicals as key
871 intermediates in the formation of dimeric compounds in α -pinene secondary organic aerosol,
872 *Environ. Sci. Technol.*, 56, 14249-14261, <https://doi.org/10.1021/acs.est.2c02090>, 2022.

873 Zhao, Z., Zhang, W., Alexander, T., Zhang, X., Martin, D. B. C., and Zhang, H.: Isolating alpha-pinene
874 ozonolysis pathways reveals new insights into peroxy radical chemistry and secondary organic
875 aerosol formation, *Environ. Sci. Technol.*, 55, 6700-6709, <https://doi.org/10.1021/acs.est.1c02107>,
876 2021.

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