1 2	Diverging trends in aerosol sulfate and nitrate measured in the remote North Atlantic on Barbados are attributed to clean air policies, African smoke, and anthropogenic emissions
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## 11 ABSTRACT

Sulfate and nitrate aerosols degrade air quality, modulate radiative forcing and the hydrological 12 cycle, and affect biogeochemical cycles, yet their global cycles are poorly understood. Here, 13 we examined trends in 21 years of aerosol measurements made at Ragged Point, Barbados, the 14 easternmost promontory on the island located in the eastern Caribbean Basin. Though the site 15 has historically been used to characterize African dust transport, here we focused on changes 16 in nitrate and non-sea salt (nss) sulfate aerosol from 1990-2011. Nitrate aerosol concentrations 17 averaged over the entire period were stable at 0.59  $\mu$ g/m<sup>3</sup> ± 0.04  $\mu$ g/m<sup>3</sup> except for elevated 18 nitrate concentrations in the spring of 2010 as well as during the summer and fall of 2008 due 19 to the transport of biomass burning emissions from both northern and southern Africa to our 20 site. In contrast, from 1990 to 2000, nss-sulfate decreased 30% at a rate of 0.023  $\mu$ g/m<sup>3</sup>/yr, a 21 trend which we attribute to air quality policies enacted in the U.S. and Europe. From 2000-22 2011, sulfate gradually increased at a rate of 0.021  $\mu$ g/m<sup>3</sup>/yr to pre-1990s levels of 0.90  $\mu$ g/m<sup>3</sup>. 23 We used the Community Multiscale Air Quality (CMAQ) model simulations from the EPA's 24 Air OUAlity TimE Series (EQUATES) to better understand the changes in nss-sulfate after 25 2000. The model simulations estimate that increases in anthropogenic emissions from Africa 26 explain the increase in nss-sulfate observed in Barbados. Our results highlight the need to better 27 constrain emissions from developing countries and to assess their impact on aerosol burdens 28 in remote source regions. 29

30

#### 31 **1. INTRODUCTION**

Sulfate and nitrate aerosol, formed from gaseous sulfur dioxide (SO<sub>2</sub>) and nitrogen oxides (e.g., 32 NO<sub>x</sub>=NO+NO<sub>2</sub>) as well as reactive nitrogen (e.g., NO<sub>y</sub>), contribute to aerosol direct and indirect 33 34 radiative forcing, impact biogeochemical cycles (Jickells et al., 2017), and degrade air quality (Adams et al., 1999; Appel et al., 1978; Charlson et al., 1992; Lelieveld & Heintzenberg, 1992; 35 Murphy et al., 2006). An outstanding question is, how have sulfate and nitrate aerosol burdens in 36 remote regions changed in response to air quality policies, economic growth, and changing 37 38 frequency of wildfires—all of which have affected SO<sub>2</sub> and NO<sub>x</sub> emissions? Answering this question is important as remote regions are an important barometer of changing global emission 39 inventories, contain ecosystems sensitive to changing chemical inputs (Galloway et al., 2008; 40 41 Mahowald et al., 2017), and are most sensitive to fluctuations in aerosol burdens that alter aerosolcloud interactions (Carslaw et al., 2013). Long-term measurement records in remote regions can 42 provide insight into this question as well as further advance current chemical transport and global 43 climate models. However, there are few long-term measurement records in remote regions. In this 44 work, we leverage 21 years of nitrate and sulfate aerosol concentrations measured at Ragged Point, 45 Barbados, a remote site in the eastern North Atlantic marine boundary layer, and use simulations 46 from a hemispheric chemical transport model—the Community Multiscale Air Quality (CMAQ) 47 model within the EPA's Air QUAlity TimE Series (EQUATES) (Foley et al., 2023)-to link our 48 observed changes in nitrate and sulfate to changing emissions inventories and meteorological 49 conditions. In turn, comparing the EQUATES model output to our time series provides guidance 50 51 on where in-situ measurements are needed to improve emissions inventories and measurementmodel agreement. 52

53 Ragged Point, Barbados provides a unique opportunity to understand changes in the nitrate and sulfate aerosol burden in the remote North Atlantic marine boundary layer. Aerosol sampling 54 began in 1971 and continues to this day generating an~50-year measurement record-the longest 55 modern speciated aerosol record, to the best of our knowledge (Prospero et al., 2021; Prospero & 56 Mayol-Bracero, 2013). The site serves as a lynchpin for understanding the impact of long-range 57 aerosol transport on the remote North Atlantic marine boundary layer and Caribbean. The site's 58 primary objective has been to understand the factors affecting the long-range transport of African 59 60 dust to the Caribbean and North America, which peaks in boreal summer in association with the seasonal northward shift in the intertropical convergence zone (ITCZ). Summer dust events are 61 caused by the strong heating of northern Africa, which causes hot, dry dust-laden desert air to be 62 carried to high altitudes, e.g., 4 - 6 km. African Easterly Waves propagate dust westward within 63 an elevated air layer known as the Saharan Air Layer (SAL) that overrides the cool, moist marine 64 boundary layer (Adams et al., 2012; Carlson & Prospero, 1972; Goudie & Middleton, 2001; 65 Tsamalis et al., 2013). Background emissions at the site are dominated by sea spray and marine 66 biogenic emissions of dimethyl sulfide (DMS) that contribute non-sea salt (nss)-sulfate (Andreae 67 et al., 1985; Barnes et al., 2006; Savoie et al., 2002). Along with dust, anthropogenic emissions 68 from Europe (Lelieveld et al., 2002), North America, and northern Africa are also transported to 69 Barbados (Savoie et al., 2002). Transport from Africa takes ~5-7 days to reach our site while 70 transport from the U.S. and Europe takes longer, typically 7-10 days. 71

Anthropogenic emissions of SO<sub>2</sub> and NO<sub>x</sub> that impact Ragged Point have changed in recent 72 73 decades due to the opposing effects of decreasing emissions mandated by national air quality 74 policies, implemented mostly in developed countries, and increasing emissions linked to rapid 75 economic growth in developing countries. The United States (U.S.) curbed emissions of  $NO_x$  and SO<sub>2</sub> with the implementation of the Clean Air Act (amended in 1990), resulting in a 92% reduction 76 77 in SO<sub>2</sub> and a 71% reduction in NO<sub>x</sub> emissions from 1990-2022 (Aas et al., 2019; Hand et al., 2012; https://gispub.epa.gov/air/trendsreport/2023/; Smith et al., 2011; Zhao et al., 2017). Countries 78 within the European Union (EU) passed similar policies resulting in analogous reductions (Aas et 79 al., 2019; Rafaj et al., 2015; Smith et al., 2011; Yang et al., 2020). Notably, reductions in SO<sub>2</sub> can 80 reduce aerosol acidity resulting in increased nitrate aerosol. Further, reductions in pollutant gases 81 can relieve oxidant limitations leading to more efficient oxidation and, therefore, reductions in SO<sub>2</sub> 82 and NO<sub>x</sub> may not reduce sulfate and nitrate aerosol as much as expected (Shah et al., 2018). In 83 84 contrast to the U.S. and EU, emissions in regions such as the Middle East, India, and Africa are continuing to increase due to rapid economic growth with emissions from India predicted to 85 overtake China as the world's largest emitter of SO<sub>2</sub> (Lelieveld et al., 2009; Li et al., 2017; 86 McDuffie et al., 2020). Due to a lack of in-situ measurements in many of these regions, chemical 87 transport and emissions inventory models combined with remote sensing have been key tools to 88 understand changing pollutants. 89

In addition to fossil fuel emissions, biomass burning is also a major source of  $SO_2$  and  $NO_x$ 90 that can impact the Atlantic (Andreae, 2019; Andreae & Merlet, 2001; Rickly et al., 2022; Roberts 91 et al., 2009; Zuidema et al., 2018). Wildfire activity has a distinct seasonality linked to the dry 92 seasons in Africa. Burning is most intense in Sub-Saharan Africa from the equator to 20°N from 93 November through May while from May through October, the savannah regions of Sub-Saharan 94 Africa from the equator to 25°S are the most active fire sources (Giglio et al., 2006; Kganyago & 95 96 Shikwambana, 2019; Roberts et al., 2009; Van der Werf et al., 2003). African smoke can be transported to Barbados from Sub-Saharan Africa north of the equator in winter and spring (Quinn 97 et al., 2022; Royer et al., 2023; Wex et al., 2016) and, less frequently, from southern hemispheric 98 Africa in fall (Trapp et al., 2010). Conditions thought to be related to African climate (e.g., the 99 North Atlantic Oscillation and the position of the Azores High) can cause large quantities of 100 northern African dust (and smoke) to be transported during the winter and spring in elevated, 101 mixed aerosol layers (Chiapello et al., 2005; Doherty et al., 2008, 2012; Gutleben et al., 2022) 102 when dust is also carried to northeastern South America (Barkley et al., 2019; Prospero et al., 1981, 103 2014). Prescribed burns in the southeastern U.S. in winter and spring may also contribute biomass 104 burning emissions to the aerosol burden observed in Barbados (Nowell et al., 2018). 105

Here we highlight different trends in nitrate and sulfate aerosol over the remote North Atlantic marine boundary layer and relate them to changing emissions. We then compare our observations to simulated concentrations of nitrate and sulfate aerosol using the CMAQ model from EQUATES, which was chosen due to its skill at modeling changes in nitrate and sulfate chemistry within the U.S. (Benish et al., 2022) as well as its ability to simulate constituents and sources of air pollution in remote regions, such as Dhaka, Bangladesh (Sarwar et al., 2023). Our results highlight the importance of long-term atmospheric observations to understand the net outcome of changing 113 global  $SO_2$  and  $NO_x$  emissions on both the aerosol burden as well as air quality in distant 114 populations.

115

# **116 2. METHODS**

2.1 Aerosol Collection at the Barbados Atmospheric Chemistry Observatory: Aerosols were 117 collected daily at the University of Miami's Barbados Atmospheric Chemistry Observatory (UM 118 BACO: https://baco.rsmas.miami.edu/) located at Ragged Point, Barbados-the easternmost 119 promontory on the east coast of the island (13.16504N, 59.43207W). The site has been operated 120 by UM since 1971, and aerosol data has been used to document the long-range transport of African 121 dust to the Caribbean and the Americas carried by the easterly trade winds as it is the first landmass 122 encountered by African emissions transported across the Atlantic Ocean (Prospero et al., 2021). 123 The site is approximately 4500 km from the west coast of northern Africa, 3000 km from the east 124 coast of the U.S. and 6000 km from the west coast of the EU. Since 1989, aerosols have been 125 126 collected at the top of a 17-m sampling tower that stands atop a 30-m bluff (see Fig 1).



127

- 128 Figure 1: Photo of the Barbados Atmospheric Chemistry Observatory (BACO) including the 17-
- m sampling tower, two shipping container laboratories, and an Advanced Global AtmosphericGases Experiment (AGAGE) laboratory.
- 131 A high-volume sampler pulls air at a nominal rate of 0.75 m<sup>3</sup>/minute across a 20cm x 25cm
- 132 Whatman-41 filter. The upper particle size limit for our filter collection method is approximately
- 133 80-100µm or greater based on the geometry of our sampling rain hat (Barkley et al., 2021). Filters
- 134 were collected daily (e.g., every 24 hours); however, a few multiday samples were also collected

that typically span 2 days. Mass collection efficiencies are 90% for sulfate, 95% for nitrate, and 135 99% for dust (Savoie & Prospero, 1982). Filters are then folded into quarters under a laminar flow 136 137 hood, placed in a clean Ziploc bag, and periodically shipped to UM for processing. To ensure that local island emissions are not sampled, the sampling pump is only operational when the wind 138 blows directly from the ocean (from 335 degrees through N and E to 130 degrees) with speeds 139 greater than 1 m/s. A timer is used to record the "run time", the total amount of time that the 140 sampling pumps are on during the sampling interval between filter changes. Data with a run time 141 less than 10% of the sampling interval are discarded to minimize the impact of low-speed and 142 flukey wind conditions that might carry aerosols from local sources. These deletions account for 143 less than 10% of all data collected over the record highlighting the steady easterly winds measured 144 at the site year-round. Procedural blanks are collected weekly by placing a Whatman-41 filter in 145 the filter cassette with the sampling pump off for 15 minutes then placing the filter in a clean Ziploc 146 147 bag; the blank is subsequently processed along with the daily filter samples.

This work focuses on aerosol measurements conducted over 1990-2011, a period that overlaps with the implementation of more stringent air quality policies in the U.S. and Europe. Seasonal trends are also shown where winter is represented by December (from the previous year), January, February (DJF); spring is March, April, May (MAM); summer is June, July, August (JJA); fall is September, October, November (SON). African dust peaks annually at BACO in JJA with episodic transport in DJF and MAM in some years (Prospero & Mayol-Bracero, 2013; Quinn et al., 2022; Royer et al., 2023). Trends presented in this work are derived from Theil-Sen regression.

155 Significance is calculated using paired samples t-tests.

2.2 Quantification of Dust and Soluble Ion Mass Concentrations: A one-quarter filter section 156 is extracted three times with a total volume of 20 mL of Milli-Q water to remove soluble material. 157 The extracted filter is placed in a combustion oven (500°C) overnight. The resulting ash is weighed 158  $(m_{filter ash})$  and subsequently corrected for the blank by subtracting the ash obtained from 159 performing the same technique on the procedural blank ( $m_{procedural blank}$ ). The gross ash weight is 160 adjusted by a factor of 1.3 to account for losses during the extraction and combustion process 161 (Prospero, 1999; Zuidema et al., 2019). This corrected ash mass concentration is equated to 162 mineral dust concentrations present on the filter based on previous comparisons between filter ash 163 164 and concentrations of aluminum, a tracer for mineral dust (Prospero, 1999; Trapp et al., 2010).

165 
$$Dust = (m_{filter ash} - m_{procedural blank ash}) * 1.3$$
 (1)

The filtrate from the sample extraction process and procedural blanks are used to quantify 166 soluble ion concentrations. Anions (e.g., chloride (Cl<sup>-</sup>), nitrate (NO<sub>3</sub><sup>-</sup>), and sulfate (SO<sub>4</sub><sup>2-</sup>)) were 167 measured using ion chromatography (IC). Cations (sodium (Na<sup>+</sup>), potassium (K<sup>+</sup>), and calcium 168 (Ca<sup>2+</sup>)) were measured with a flame photometer after 2004 while flame atomic absorption 169 spectrophotometry was used prior to 2004 limiting cation analysis to sodium (Savoie, et al., 2002). 170 171 In addition to total soluble ion concentrations, we also report concentrations of non-sea salt (nss) sulfate, which is a useful tracer of sulfur from marine biogenic and pollution emissions, and nss-172 potassium, a tracer of biomass burning emissions (Andreae, 1983; Keene et al., 1986). 173 Concentrations of nss-SO<sub>4</sub><sup>2-</sup> and nss-K<sup>+</sup> are calculated using the following mass-based equations 174 and assuming that Na<sup>+</sup> is a conservative tracer of sea spray aerosol: 175

176 
$$nss - SO_4^{2-} = [SO_4^{2-}] - (0.2517 * [Na^+])$$
 (2)

(3)

177 
$$nss - K^+ = [K^+] - (0.03595 * [Na^+])$$

Filter samples with undetectable amounts of dust and soluble ions compared to procedural blanksare removed from our analysis.

#### **180 2.3 EQUATES Model Products**

EPA's Air QUAlity TimE Series (EQUATES) Project uses the Community Multiscale Air Quality 181 (CMAQ) model, a 3-D chemical transport air quality model, to simulate air quality over a 182 continuous 2002-2019 period (Foley et al., 2023). CMAQ accounts for gas, cloud, and aerosol 183 chemistry, including processes such as in-cloud sulfate formation from the oxidation of SO<sub>2</sub>. 184 EQUATES uses CMAQv5.3.2 (Appel et al., 2021) to model the Northern Hemisphere using 108-185 km resolution and 44 vertical layers (Mathur et al., 2017). Meteorological data are derived from 186 the Weather Research and Forecasting model (WRFv.4.1.1). Emissions from outside the 187 contiguous U.S. and China are generated using the Hemispheric Transport of Air Pollution version 188 2 inventory for the year 2010 (HTAPv2.2) and are scaled to other years using the Community 189 Emissions Data System (CEDS) for the years 2002-2019. The Fire INventory from NCAR (FINN) 190 is used to generate biomass burning emissions (Wiedinmyer et al., 2011), lightning NOx emissions 191 are derived from the Global Emissions IniAtive (GEIA), biogenic volatile organic compounds 192 (VOCs) are from MEGAN2, and soil NO<sub>x</sub> is from CAMSv2.1. 193

In this study, EQUATES was used to better understand observed trends, namely in nss-sulfate, 194 after 2000 when observed concentrations unexpectedly increased. EOUATES also simulated 195 concentrations of locally emitted gases, including  $SO_2(g)$  and  $NO_x(g)$ , which are not measured at 196 Ragged Point. The model product is not available for the years prior to 2000. The analysis focused 197 on dust, sea spray, nitrate, sulfate, and gaseous SO<sub>2</sub> and NO<sub>2</sub>. In addition to anthropogenic sources 198 of SO<sub>2</sub>, natural sources from the oxidation of DMS were included in model runs. Species 199 200 predictions were extracted from the lowest CMAQ model layer (~10m in thickness) for a source area over the Atlantic Ocean to the east of the island from 59.5627°W to 56.5448°W longitude and 201 14.3989°N to 11.4566°N latitude (equivalent to 16 grid cells with one cell over Ragged Point and 202 the others to the east of the site). Simulated concentrations of aerosol sulfate, nitrate, calcium, 203 potassium, and sodium were obtained for fine mode aerosol (e.g., Aitken and accumulation mode, 204 PMF model outputs) and coarse mode aerosol (Total PM - PM2.5, PMC model outputs). Previous 205 studies have shown that most of the aerosol mass at Ragged Point is below 10 µm diameter 206 207 (Prospero et al., 2001). Because aerosol filters collected at BACO capture total suspended particulate matter, model outputs of fine and coarse mode aerosol concentrations were combined 208 (e.g., PMF + PMC model outputs) to give total aerosol mass concentrations of sulfate, nitrate, 209 sodium, potassium, and calcium. Total concentrations of nss-sulfate and nss-potassium were 210 211 calculated using Equations 2 and 3, respectively, and model outputs of total sodium mass concentrations, total sulfate mass concentrations, and total potassium mass concentrations from 212 combined coarse and fine mode aerosol model outputs. We calculated concentrations of nss-213 214 calcium, which has been shown to be a good tracer for mineral dust in Barbados (Savoie & Prospero, 1980), and dust mass concentrations were then calculated using the average upper crustal 215

abundance of calcium in soil (an average of 4.1%) (Scheuvens et al., 2013; Taylor & McLennan,
1985) as shown in Equations 4 and 5.

218 
$$nss - Ca^{2+} = [Ca^{2+}] - (0.0376 * [Na^+])$$
 (4)

219 
$$Dust = [nss - Ca^{2+}] * 24.1$$
 (5)

220 We first assessed the ability of the EQUATES CMAQ simulations to capture trends in different aerosol types observed at Ragged Point. Simulations of sodium and dust from EQUATES capture 221 seasonal and monthly observed trends in dust (see Fig S1 and S2 of the Supporting Information 222 (SI)). However, the model overpredicts sodium (Na<sup>+</sup>, a proxy for sea spray) by a factor of 3-4 (Fig. 223 S1) and underpredicts dust by an average factor of ~7. This low bias for dust in CMAQv5.3.2 is 224 consistent with CMAQ development that occurred after the EQUATES simulations were 225 complete. A bugfix to the online dust emissions module in CMAQv5.4 increases dust emissions 226 by a factor of 3-7 over the Sahara Desert and parts of Asia (see the CMAQv5.4 release notes for 227 further information; https://www.epa.gov/cmaq/cmaq-documentation#release-notes). 228

In addition to simulating observed aerosol concentrations, EQUATES was also used to examine trends in gaseous indicators of anthropogenic and biomass burning emissions, oxidant concentrations and oxidant ratios important for investigating changes in the oxidation efficiency of locally emitted pollutant gases and the subsequent formation of nitrate and sulfate aerosol. Further, EQUATES was used to investigate whether the oxidation efficiency of SO<sub>2</sub>(g) changed during the 2002-2011 period. The oxidation ratio was calculated from Equation 6 (Shah et al., 2018):

236 Oxidation Ratio = 
$$\frac{nss - SO_4^{2^-}}{nss - SO_4^{2^-} + SO_2}$$
(6)

For this calculation, we used EQUATES model data for SO<sub>2</sub>(g) concentrations and filter-based observations of nss-sulfate.

### 239 2.4 HYSPLIT Back-trajectory Analysis

Air mass back-trajectory analysis was performed using NOAA's Hybrid Single-Particle 240 Lagrangian Integrated Trajectory (HYSPLIT) model (Draxler & Rolph, 2011; Rolph et al., 2017; 241 Stein et al., 2015). 13-day back-trajectories were initiated at heights of 500, 1000, and 2000-m to 242 capture both SAL and boundary layer transport. Frequency plots were also generated for the entire 243 sampling period (1990-2011) to illustrate seasonal differences in air mass transport and origin and 244 to explore any interannual variability. We used the global National Center for Environmental 245 Prediction (NCEP) reanalysis data that extends back to the beginning of our dataset in 1990 246 (Kramer et al., 2020). The finer resolution Global Data Assimilation System (GDAS) dataset (1° 247 resolution) was used to examine trends in air mass back trajectories in 2008, 2009, and 2010 when 248 transport conditions strongly impacted nitrate aerosol concentrations. We emphasize the 249 challenges in characterizing long-range transported aerosol due to the remoteness of the site, the 250 location of Barbados in the middle of the ocean, and the fact that air layers of smoke and dust 251 change in height during transport. As such, we caution against the over-reliance on HYSPLIT 252

analysis to extensively characterize long-range transport aerosol dynamics and instead use thisanalysis to quantitatively describe major changes in transport patterns.

#### 255 **3. RESULTS**

#### 256 **3.1 Measured Trends from 1990-2011 at Ragged Point.**

Figure 2 shows yearly averaged mass concentrations of non-sea salt (nss) sulfate and nitrate 257 from 1990 through 2011. Nitrate concentrations are remarkably stable from 1990 to 2011 258 (R<sup>2</sup>=0.006, p>0.05 (not significant)) with an average concentration of 0.59  $\mu$ g/m<sup>3</sup> ± 0.04  $\mu$ g/m<sup>3</sup>. 259 However, two anomalous peaks in nitrate are observed in 2008 and 2010 with annual average 260 nitrate concentrations of 0.73 and 0.81  $\mu$ g/m<sup>3</sup>, respectively. Similarly, dust mass concentrations 261 also show no trend over this period ( $R^2=0.06$ ). In contrast, nss-sulfate decreases by 30% from an 262 average concentration of 0.84  $\mu$ g/m<sup>3</sup> starting in 1990 to a minimum of 0.64  $\mu$ g/m<sup>3</sup> in 2000 at a rate 263 of -0.023 ug/m<sup>3</sup>/yr. Subsequently, sulfate gradually increased to 0.91 µg/m<sup>3</sup> in 2010 and 0.90 264  $\mu g/m^3$  in 2011, maximums across the entire record. The trends in the yearly average mass 265 concentrations of nitrate and nss-sulfate are significantly different (p-value < 0.005), which can be 266 explained by either different sources or different rates of change for precursor NO<sub>x</sub> and SO<sub>2</sub> 267 emissions or different responses in nitrate and sulfate aerosol production to changing emissions of 268 NO<sub>x</sub> and SO<sub>2</sub> as has been shown in North America (Shah et al., 2018; Vasilakos et al., 2018). 269



270

Figure 2: Yearly averages of non-sea salt sulfate (NSS Sulfate, red line, top panel), nitrate (middle panel,
blue line), and dust (bottom panel, brown line with squares) measured at Ragged Point, Barbados, from

273 1990-2011. Dashed lines show changes in NSS Sulfate pre-2000 (orange), post-2000 (gray line), nitrate

- (teal line) without the spikes in 2008 and 2010 considered, and dust (brown line) over the 1990-2011
  period.
- 276

The Barbados yearly concentrations differ from long-term observations of aerosol and 277 precipitation chemistry measured at Tudor Hill, a site on the west coast of Bermuda, from 1989-278 1997 and from 2006-2009 as part of the same program as that at Barbados (AEROCE) and using 279 the same protocols including sampling only when winds are over the ocean (Keene et al., 2014; 280 Savoie, et al., 2002). At the Bermuda site, the prevailing winds come from the West so that the 281 282 sampling sector extends from 180° through West to 330°. Aerosol nitrate is constant at both sites. However, the nitrate annual mean at Bermuda is  $\sim 1.05 \,\mu g/m^3$ , roughly double the nitrate observed 283 at Ragged Point (Keene et al., 2014; Savoie, et al., 2002). Also, the decline in nss-sulfate observed 284 in Bermuda, from ~2.59  $\mu$ g/m<sup>3</sup> in 1989 to ~1.63  $\mu$ g/m<sup>3</sup> in 2009, is greater than our observations at 285 Ragged Point (Keene et al., 2014). Furthermore, sulfate aerosol in Bermuda declines over the 286 entire record and does not exhibit the same reversal in the 2000s that we observe at Barbados. The 287 difference in concentrations and trend in nss-sulfate observed at Ragged Point compared to that at 288 Bermuda is not surprising given that the Bermuda site is located 1100 km from the east coast of 289 the U.S. and is more directly influenced by anthropogenic emissions compared to Barbados, which 290 is more remote and could be influenced by a multitude of emission sources (Savoie, et al., 2002). 291 Also, the trend in nss-sulfate at Barbados differs from long-term observations measured at an 292 IMPROVE site in the Virgin Islands. Sulfate shows no trend from 2004-2021 (Hand et al., 2024). 293 294 The difference in trend in nss-sulfate between Ragged Point and the IMPROVE site in the U.S. Virgin Islands is likely due to more influence from the U.S. and less influence from African 295 emissions. Further, the IMPROVE site does not follow a sector-controlled sampling protocol as 296 297 does the site in Barbados.

298 To assess whether annual trends in nss-sulfate and nitrate observed in Barbados are associated 299 with African aerosol transport conditions, annual average measured dust mass concentrations are also shown in Fig 2 and show no appreciable increase or decrease from 1990-2011 ( $\mathbb{R}^2=0.06$ ). 300 Therefore, annual trends in nss-sulfate aerosol, which show a decrease from 1990-2000 and an 301 increase from 2000-2011, are not correlated with African dust mass concentrations ( $R^2=0.001$ ). 302 This suggests that the increase in nss-sulfate after the year 2000 is not due to an increase in sulfate 303 associated with heterogeneous reactions between  $SO_2(g)$  and dust or to more favorable transport 304 from Africa. Nitrate aerosol, in contrast, is modestly correlated with dust ( $R^2=0.30$ ). Comparing 305 seasonal nitrate and dust mass concentrations by year reveals tighter correlations between nitrate 306 and dust for DJF and MAM (0.46 and 0.4, respectively, see Fig S3). DJF and MAM are not the 307 peak dust transport seasons to the Caribbean, but they are the seasons that favor co-transport of 308 309 dust and biomass burning emissions from Sub-Saharan Africa north of the equator (Royer et al., 2023). Transport of African smoke to Barbados has been shown to be associated with elevated 310 concentrations of nitrate, which likely explains the association between winter- and spring-time 311 dust and nitrate (Quinn et al., 2021; Savoie & Prospero, 1982). 312

## 313 **3.1.2** Seasonal Patterns in Air Mass Trajectories and Nitrate Concentrations

HYSPLIT air mass back-trajectories reveal similar seasonal patterns each year with 314 315 predominantly easterly transport to the site year-round (see Fig S4, which shows 5-day air mass 316 back-trajectory frequency plots for the entire 1990-2011 period and similar patterns year to year). As has been documented in prior work (Prospero et al., 2021), the dominant transport pathways 317 are over the ocean and traversing the African continent. Some trajectories do intercept North 318 319 America and come close to Europe while few (<10%) air masses come near South America, which is outside of our sampling sector. Air masses typically take 5-7 days to be transported from 320 northern Africa and longer than 7 days to intercept the E.U. and North America. From Figure S4, 321 DJF has some trajectories that intercept North America and some from the northern part of the 322 Atlantic Ocean toward Europe, MAM trajectories are easterly, JJA trajectories are from the west 323 coast of Africa, and SON trajectories are from the east with some from the southeast. We note that 324 the directional tendencies of the trajectories did not noticeably change in 2000, when the 325 326 concentration trends in nss-sulfate changed.

327 Figure 3 shows trends in nss-sulfate and nitrate during different seasons (e.g., DJF, MAM, JJA, and SON). Increases in nitrate in 2008 are driven by high nitrate concentrations in JJA and SON 328 in 2008, with most of the increase in September, while high nitrate levels in 2010 are primarily 329 observed during MAM. In 2010, a transition from an El Niño to the strongest La Niña event on 330 record occurred (Wolter & Timlin, 2011; Zhang et al., 2019), and long-range transport from Africa 331 was anomalously high during the springtime as evidenced by high mass concentrations of dust in 332 the spring of 2010 (Zuidema et al., 2019). Daily air mass back-trajectories passed more frequently 333 over the African continent in MAM of 2010 compared to MAM of 2009 (26 vs 6 days) (Fig 4a 334 and b), and nitrate levels exceeded 1  $\mu g/m^3$  on over half of those days when trajectories traversed 335 the northern African continent. Figure 4 focuses on trajectories initiated at 500-m height. If 336 trajectories initiated at 1000-m are also included, then transport over the North African continent 337 occurs on 36 days in MAM of 2010, i.e., 82% of days when nitrate levels exceeded 1 µg/m<sup>3</sup>. 338 339 However, during the summertime peak (JJA) in African dust transport in 2010, nitrate does not show the same increase as dust despite frequent transport from northern Africa (see Fig 4d). 340 Further, Fig 4c and 4d compare daily and monthly mean concentrations of nitrate, dust, and non-341 sea salt potassium (e.g., nss-K<sup>+</sup>), which is a marker for biomass burning emissions (Andreae & 342 Merlet, 2001). Nitrate and nss-K<sup>+</sup> clearly track each other and are both elevated in the spring of 343 2010. Although biomass burning peaks in December and January while MAM is the tail end of 344 the burn season in northern Sub-Saharan Africa (Giglio et al., 2013; Roberts et al., 2009), the high 345 nitrate loadings observed in 2010 in spring are likely due to strongly favorable transport conditions 346 that transport both dust and smoke to Barbados during this year. 347





349 Figure 3: Average concentrations of nitrate (blue markers), and nss-sulfate (red markers) for winter (DJF),

spring (MAM), summer (JJA), and fall (SON). The vertical black dashed line denotes the year 2000 when
 nss-sulfate trends shifted from decreasing to increasing.



Figure 4: 13-day air mass back-trajectories initiated at 500-m for MAM a) 2010 and b) 2009. Backtrajectories labeled in green are for days where nitrate concentrations measured at Ragged Point were < 1  $\mu g/m^3$  while trajectories labeled in red had nitrate > 1  $\mu g/m^3$ . Maps are from Google Earth. c) Daily comparison of nitrate and non-sea salt potassium (nss-K<sup>+</sup>), a biomass burning marker, show a tight correlation (R<sup>2</sup>=0.62). d) Monthly average concentrations of dust (brown), nitrate (blue), and nss-K<sup>+</sup> (green) for 2010.

Similar to 2010, air mass back-trajectories traversed the African continent on just 40% of the 359 days in JJA but 82% of days in September of 2008 with nitrate concentrations > 1  $\mu$ g/m<sup>3</sup> (Fig S5 360 and Fig 5a). If trajectories initiated at the 1000-m level are included, then African transport occurs 361 on 64% of days in JJA with nitrate concentrations > 1  $\mu$ g/m<sup>3</sup>, consistent with higher level transport 362 in JJA. In JJA, most of the air mass back-trajectories pass through Africa north of the equator. 363 However, in September, trajectories take a more southerly route traversing near the South 364 American coast and Sub-Saharan Africa south of the equator (Fig 5a). Figure 5b shows trajectories 365 for September but in 2009, when fewer trajectories take a southerly route and nitrate concentrations 366 rarely exceed 1  $\mu$ g/m<sup>3</sup>. September is during the peak of the burn season in Sub-Saharan Africa 367 south of the equator and also dovetails with strong smoke transport over the Atlantic Ocean due to 368 the increased intensity of the African Easterly Jet (Adams et al., 2012; Adebiyi & Zuidema, 2016; 369 Zuidema et al., 2018). Air mass back-trajectories in September from Sub-Saharan Africa south of 370 the equator to our measurement sites in Barbados and Cayenne have been linked with African 371 372 smoke transport (Barkley et al., 2019; Trapp et al., 2010) further suggesting a link between elevated concentrations of nitrate measured at Barbados with transported African smoke. Biomass 373 burning in the Amazon also peaks in SON (Adams et al., 2012) and could also explain some of 374 375 the increase in nitrate in this season in 2008.



**Figure 5:** 13-day air mass back-trajectories initiated at 500-m for September a) 2008 and b) 2009. Backtrajectories labeled in green are for days when nitrate concentrations measured at Ragged Point were < 1  $\mu$ g/m<sup>3</sup> while trajectories labeled in red had nitrate > 1  $\mu$ g/m<sup>3</sup>. The maps are from Google Earth.

380

### **381 3.1.3 Trends in Non-Sea Salt-Sulfate and Emissions of SO<sub>2</sub>**

Figure 3 reveals that the decrease in sulfate and subsequent increase are measured during most 382 seasons. Table 1 further provides correlation coefficients and the rate of change in either nitrate or 383 sulfate in µg m<sup>-3</sup>yr<sup>-1</sup> for data collected pre- and post-2000 when the trend in nss-sulfate changed. 384 Pre-2000, nss-sulfate shows a consistent decline in each season with the weakest decline in winter 385 and the strongest reduction measured in JJA at -0.036 µg m<sup>-3</sup>yr<sup>-1</sup> (R<sup>2</sup>=0.72, Table 1). Post-2000, 386 nss-sulfate increased during every season with comparable increases in every season. In JJA, nss-387 sulfate increased at  $+0.028 \ \mu g \ m^{-3} yr^{-1}$  (R<sup>2</sup>=0.61, Table 1). In contrast, nitrate shows no trend (e.g., 388  $R^{2}$ <0.2) in any season except a slight increase of +0.012 µg m<sup>-3</sup>yr<sup>-1</sup> post-2000 in JJA. Unlike nitrate 389 which showed intermittent spikes associated with increased biomass burning aerosol transport 390

- periods, the persistent increases in nss-sulfate observed during all seasons more likely reflect either
   increased emissions or more efficient oxidation of SO<sub>2</sub>.
- **Table 1:** Seasonal trends in nitrate and nss-sulfate shown pre- and post-2000. The rate of change in  $\mu g/m^3/yr$
- is shown and the correlation coefficient is also included. Values of  $R^2 < 0.2$  were denoted as having
- 395 "No Trend". DJF represents winter, MAM represents spring, JJA represents summer, and SON396 represents fall.

Season	Compound	Pre-2000 Rate of Change (μg/m <sup>3</sup> /yr)	Post-2000 Rate of Change (ug/m <sup>3</sup> /yr)
DJF	Nitrate	-0.0005 (R2=0.05, No Trend)	+0.013 (R2=0.07, No Trend)
	NSS-Sulfate	-0.0070 (R2=0.05, No Trend)	+0.035 (R2=0.48)
MAM	Nitrate	+0.0037 (R2=0.01, No Trend)	+0.028 (R2=0.08, No Trend)
	NSS-Sulfate	-0.034 (R2=0.64)	+0.033 (R2=0.09, No Trend)
JJA	Nitrate	-0.002 (R2=0.02, No Trend)	+0.012 (R2=0.21)
	NSS-Sulfate	-0.036 (R2=0.72)	+0.028 (R2=0.61)
SON	Nitrate	+0.0003 (R2=0.02, No Trend)	+0.013 (R2=0.13, No Trend)
	NSS-Sulfate	-0.023 (R2=0.29)	+0.026 (R2=0.44)

399 Figure 6 compares yearly trends in Ragged Point nss-sulfate along with SO<sub>2</sub> emissions reported from the Community Emissions Data System (CEDS) (McDuffie et al., 2020). We focus on the 400 most likely sources to impact Ragged Point: the U.S., EU, and Africa. A similar figure comparing 401 nitrate concentrations measured at Ragged Point and nitrogen dioxide (NO<sub>2</sub>) emissions from 402 CEDS can be found in the SI (Fig S6). Our near-constant nitrate mass concentrations do not match 403 the decline in NO<sub>2</sub> observed in the EU and U.S. and the increase in NO<sub>2</sub> in Africa. In contrast, 404 405 decreases in nss-sulfate observed from 1990-2000 at Ragged Point closely follow the 32% and 406 58% reductions of SO<sub>2</sub> emissions in the U.S. and Europe, respectively (Fig 6) (Aas et al., 2019; Hand et al., 2012; McDuffie et al., 2020; Rafaj et al., 2015; Yang et al., 2020). Our finding that 407 changing SO<sub>2</sub> emissions in the EU and U.S. is reflected in our record at Ragged Point agrees with 408 409 previous work examining both anthropogenic and biogenic sulfate (Savoie et al., 2002). Figure 6 also compares the trends of nss-sulfate observed at Ragged Point to increasing emissions of SO<sub>2</sub> 410 from Africa (McDuffie et al., 2020). Before 2000, SO<sub>2</sub> emissions from Africa oscillated around 411 4.44 $\pm$ 0.19 Tg S/yr but show no consistent trend (R<sup>2</sup>=0.027). However, after 2000, emissions of 412 SO<sub>2</sub> from Africa steadily increase by 37% from 2000-2011 (4.33 Tg S/yr in 2000 and 5.95 Tg S/yr 413 in 2011,  $R^2=0.88$ ). The rate of increase in SO<sub>2</sub> is on par with the rate of increase in nss-sulfate of 414 29% observed in Barbados suggesting that anthropogenic emissions from Africa are affecting the 415 nss-sulfate trends measured in Barbados. 416

We next utilized the CMAQ model results from EQUATES to gain further insight into the observed recovery of nss-sulfate (post-2000). We first note that EQUATES also predicts an increase in SO<sub>2</sub> emissions from northern hemispheric Africa after 2002 (see Figure S7). However, we use EQUATES to determine if other factors such as changes in the oxidation efficiency of locally emitted SO<sub>2</sub> and meteorological changes affected our observations.



Figure 6: Percent changes in NSS-Sulfate mass concentrations measured at Ragged Point from 1990-2011 423 are shown in red. NSS-Sulfate mass decreases from 0.87 µg/m<sup>3</sup> in 1990 to 0.64 µg/m<sup>3</sup> in 2000 and 424 subsequently increases to  $0.90 \ \mu g/m^3$  in 2011. The data are normalized to the maximum NSS-sulfate mass 425 concentration, 0.92 µg/m<sup>3</sup> measured in 2010. The percent changes in emissions of sulfur dioxide (SO<sub>2</sub>) 426 from the CEDS emissions inventory from McDuffie et al., 2020 are included for comparison. Decreasing 427 428 emissions of SO<sub>2</sub> from the U.S. and EU are shown in blue lines. U.S. SO<sub>2</sub> reduces from 21.12 to 5.85 Tg S/yr and EU SO<sub>2</sub> reduces from 28.06 to 7.74 Tg S/yr. Percentages are calculated by normalizing to the 429 430 maximum values in SO<sub>2</sub> emissions observed in 1990 for both the U.S. and EU. SO<sub>2</sub> emissions from Africa increased from 4.13 to 5.95 Tg S/yr; they are normalized to the maximum SO<sub>2</sub> emissions observed in 2010 431 432 at 6.16 Tg S/yr.

433

#### 434 **3.2** Comparison of Measured and Modeled Trends of Nitrate and Sulfate Aerosol

Monthly concentrations of simulated nss-sulfate and nitrate (for both the fine and coarse 435 mode combined) were compared with mass concentrations measured on filters collected at Ragged 436 Point (see Fig 7 and 8, respectively). The model simulates similar concentrations both at Ragged 437 Point and the area to the east of the site implying that emissions on Barbados are minimal. 438 EQUATES predicts nitrate concentrations in fine and coarse aerosol sizes. A greater proportion of 439 fine nitrate is predicted in DJF (40% of the total modeled nitrate) and MAM (26% of the total 440 modeled nitrate) during all years compared to JJA and SON (18% of the total modeled nitrate for 441 both seasons, see Fig S8). This seasonality is likely due to increased contributions of nitrate from 442 fine biomass-burning aerosol produced in Sub-Saharan Africa north of the equator in winter and 443 spring. This point is highlighted in MAM of 2010 when the amount of modeled fine nitrate was 444 elevated (35% of total modeled nitrate). Nss-sulfate is predicted to be almost exclusively in fine 445 mode aerosol. While previous observations have shown that nss-sulfate dominates the fine mode 446 in Barbados, some of the nss-sulfate is also present in the coarse mode likely due to heterogeneous 447 448 reactions between  $SO_2$  and coarse sea spray and mineral dust aerosols (Adams et al., 2005; Alexander et al., 2005; Li-Jones & Prospero, 1998). EQUATES does not seem to be capturing this 449 minor yet important budget of nss-sulfate. 450





Figure 7: Monthly averages of non-sea salt sulfate (nss-sulfate) mass concentrations measured at Barbados 453 on filters (solid red line with red circles) compared to monthly averages of NSS-sulfate calculated from 454 EQUATES model simulations of combined fine and coarse mode sodium (Na) and sulfate using equation 455 2 (dashed red line) for 2002-2011.



457 Figure 8: Monthly averages of nitrate mass concentrations measured in Barbados on filters (solid blue line
458 with blue circles) compared to monthly averages of combined fine and coarse mode nitrate calculated from
459 EQUATES model simulations (dashed blue line) for 2002-2011.

To assess the performance of the model compared to our observations, we calculated the 462 normalized mean bias (NMB) and Pearson correlation coefficient (r) for monthly averaged 463 464 concentrations of nss-sulfate and nitrate (see Table S1 (Boylan & Russell, 2006)). Additional calculations of the mean bias (MB) and root-mean-square error (RMSE) can also be found in Table 465 S1 of the SI. NMB for nitrate was generally within  $\pm 20\%$ , better than predictions of nitrate within 466 467 the U.S. (Kelly et al., 2019), except for 2005 (-24.18%) as well as 2008 (-35.06%) and 2010 (-28.72%) when the model underpredicted measurements likely due to the elevated African smoke 468 transport events. The model overpredicts nss-sulfate in the earlier years (2002-2007) then 469 converges closer to our measurements at Ragged Point after 2008 as shown in Fig 7. Because 470 trends in sea salt (e.g., sodium from EQUATES, Figure S1) show a constant high bias in all years, 471 the overprediction of nss-sulfate reflects biases in sources of sulfate (other than sea spray) or biases 472 in the conversion of SO<sub>2</sub> to sulfate. Further, the NMB for nss-sulfate also reflects the model 473 474 overprediction of nss-sulfate as high values are observed in 2002 (+81.45%) then the model begins to fall within  $\pm 20\%$  starting in 2008 (see Table S1). As such, trends in nss-sulfate simulated by 475 EQUATES show a decrease over time rather than an increase post-2000 (see Fig S9). The decrease 476 in EQUATES simulations of nss-sulfate post-2000 compared to the increase observed in Barbados 477 could be related to a changing bias over time. For example, recent predictions for 2019 indicate 478 CMAQ underpredicts sulfate by about 50% in the eastern U.S. (45% underestimate across entire 479 U.S.) (Vannucci et al., 2024). Previous simulations for 2002 and 2016 indicated more modest 480 normalized mean biases of 20% or less (Appel et al., 2021; Sarwar et al., 2011). As a result, 481 EQUATES may simulate a stronger decline in transported U.S. sulfate over 2002-2019 than 482 observations indicate. This overestimate in declining sulfate in the U.S. may mask trends 483 (including stronger increases in sulfate) in other regions. Pearson correlation coefficients range 484 from -0.22 to 0.88 by year for nitrate with a mean of 0.54. The poorest correlation (-0.22) occurs 485 for 2002 when Barbados filter measurements were unavailable from January until May. For nss-486 sulfate, r ranges from 0.23 to 0.82 by year with a mean of 0.59. The greatest variation between the 487 model and measurements is for 2009 (see Fig 7), which was also the year that had an erroneously 488 high model concentration of sea spray aerosol in winter. Overall, the model is capturing both the 489 magnitude and seasonal and interannual variation of nss-sulfate and nitrate at our remote location 490 in the tropical North Atlantic. 491

### 492 **3.2.1 Using EQUATES to Determine Contributions to Aerosol Sulfate**

Figure 9 shows a decrease in SO<sub>2</sub> simulated by EQUATES within the modeled grid space 493 over Ragged Point and to the east of the site. Consistent with a predicted decrease in SO<sub>2</sub> and the 494 observed increase in nss-sulfate, the oxidation ratio of SO<sub>2</sub> increases in the region near Barbados. 495 We note that because the lifetime of SO<sub>2</sub> is predicted to be  $\sim$ 20 hours and as much as  $\sim$ 40 hours 496 depending on the latitude of the source of SO<sub>2</sub> and the season (Green et al., 2019; Lee et al., 2011), 497 498 our oxidation ratio estimates here are most relevant for local emissions, including sulfur emitted from marine phytoplankton (e.g., DMS). However, prior measurements at Bermuda suggest a 499 longer lifetime for long-range transported SO<sub>2</sub> compared to local oceanic emissions from DMS, 500 which are subject to strong condensational losses to sea spray in the marine boundary layer (Keene 501 et al., 2014). As such, long-range transported sources from Africa, the U.S., and EU have likely 502 been converted to sulfate upwind of Barbados and SO<sub>2</sub> concentrations shown here most likely 503

reflect oceanic sources though some contribution from long-range transport is also possible. 504 Locally emitted hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) concentrations also increased over this period (see Fig 505 506 10) and is likely linked to decreases in locally emitted SO<sub>2</sub>, which is a major sink of H<sub>2</sub>O<sub>2</sub> during the aqueous phase formation of sulfate (Manktelow et al., 2007). We note that the predicted 507 oxidation ratios are likely an overestimate because EQUATES overpredicts nss-sulfate compared 508 509 to observations and the overprediction decreases with time. As such, the increase in the oxidation ratio, which is small from Figure 9, likely has a minimal influence on the observed trends in nss-510 sulfate observed at our site. 511



512

513 Figure 9: Annual trends in locally emitted sulfur dioxide (SO<sub>2</sub>) simulated by the EQUATES model and the

514 calculated oxidation ratio at Ragged Point. Linear fits and corresponding correlation coefficients are also

515 shown.

516



517

Figure 10: Annual trends in select gas phase species including benzene (pink trace, bottom panel), carbon
 monoxide (CO) (orange trace, middle panel), and hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) (purple trace, top panel)
 simulated by the EQUATES model at Ragged Point. Linear fits and corresponding correlation coefficients
 are also shown.

522

Changes in biomass burning, anthropogenic emissions, oxidant concentrations, and 523 meteorological parameters were also investigated using EQUATES. Fine mode non-sea salt 524 potassium (nss-K<sup>+</sup>) was used as our tracer for smoke emissions. We used EOUATES data rather 525 than our nss-K<sup>+</sup> observations, which were non-existent for most years from 2002-2011. Also, our 526 measurements included total (i.e., fine and coarse mode) nss-K<sup>+</sup> including contributions from 527 African dust (Savoie & Prospero, 1980). EQUATES simulations of nss-K+ were a factor of 2-3 528 higher than our measurements except for the dust and smoke transport event that occurred in MAM 529 of 2010. No significant increase in simulated nss-K<sup>+</sup> was estimated (Fig S10). Meteorological 530 parameters (temperature, relative humidity (RH), wind speed, wind direction) were constant over 531 time showing no shift in rainfall, winds, temperature, or RH (Fig S10). Concentrations of the 532 hydroxyl radical ('OH) did not show a significant change from 2002-2011 (Fig S10). EQUATES 533 concentrations of benzene (a tracer for fossil fuel combustion) and carbon monoxide (CO, a tracer 534 for combustion from fossil fuels and biomass burning) both increased modestly from 2002-2011 535 (Fig 10). Our results suggest that an increase in the oxidation efficiency of locally emitted SO<sub>2</sub> to 536 sulfate may contribute to the increase of nss-sulfate post-2000, but the contribution is likely small 537 compared to long-range transported sulfate. 538

## 539 4. DISCUSSION & CONCLUSIONS

Our 21-year record (1990-2011) of nitrate and nss-sulfate aerosol shows two distinct trends 540 over the tropical North Atlantic. Nitrate shows no significant change other than two spikes in JJA 541 and September of 2008 and MAM of 2010. Variations in winter- and spring-time dust transport 542 explained interannual oscillations in nitrate concentrations while increased transport of smoke 543 from Sub-Saharan Africa north of the equator in MAM 2010 and from south of the equator and, 544 possibly, South America in September 2008 caused the increased levels of nitrate observed in 545 those years. Nitrate has been shown to be enhanced in smoke by up to 5-fold over background due 546 to high emissions of NO<sub>x</sub> that are rapidly converted to nitrate (Adon et al., 2010; Hickman et al., 547 2021; Perron et al., 2022; Schlosser et al., 2017). Notably, nitrate was not enhanced in JJA of 2010 548 even though high quantities of dust were transported to Barbados. We speculate that the lack of 549 enhanced nitrate during summer of 2010 is due to the lack of nitrate contribution from African 550 smoke emissions during this season. This finding in addition to our observations of enhanced 551 nitrate associated with dust in DJF and MAM suggests that the nitrate associated with African 552 553 aerosol transport primarily originates from NO<sub>x</sub> emissions from African wildfires that are rapidly converted to nitrate prior to transport. 554

In contrast to the relatively flat trend in aerosol nitrate, nss-sulfate decreased by 30% from 1990-2000 then increased from 2000-2011, recovering to concentrations measured in the early 1990s. Reductions in nss-sulfate observed in Barbados are most likely due to decreased emissions of SO<sub>2</sub> in the U.S. and EU due to clean air policies implemented via technologies such as flue-gas desulfurization devices installed in power plants (Aas et al., 2019; Kharol et al., 2017; Smith et al., 2011). Thus, our results highlight that regulations aimed to improve national and regional air
quality also impact more distant locations such as the remote North Atlantic marine boundary layer
and Caribbean.

As shown in Fig 6, increases in SO<sub>2</sub> in Africa, namely from anthropogenic sources, are the 563 most probable cause for the increase in nss-sulfate levels from 2000-2011 in Barbados. 564 Simulations from both CEDS and EQUATES reveal an increase in anthropogenic emissions of 565 566 SO<sub>2</sub>. Further evidence for this speculation stems from the lack of change in air mass back 567 trajectories before and after 2000, suggesting that emissions rather than meteorological trends are 568 driving our observed patterns in nss-sulfate. Industrial contributions of sulfate from oil refineries, coal fired power plants, and fertilizer plants along the north and northwestern coast of Africa were 569 570 also observed in the Cape Verde Islands (Salvador et al., 2015). Anthropogenic sources of SO<sub>2</sub> in Africa include emissions from electricity generation, diesel combustion and transportation 571 (Assamoi & Liousse, 2010; Keita et al., 2021; Liousse et al., 2014), refineries, gas flaring, and 572 smelting (Doumbia et al., 2019; Osuji & Avwiri, 2005). Levels of pollution in African cities are 573 on par with Asian megacities with the largest rate of increases in SO<sub>2</sub> observed in western Africa 574 (Adon et al., 2016; Hopkins et al., 2009; Liousse et al., 2014; Val et al., 2013). The industrial sector 575 in the Highveld region of South Africa also shows some of the highest increases in SO<sub>2</sub> with 576 577 observed increases starting in 1980 due to an increase in the number of coal-fired power plants arising from an increased demand for electricity (Keita et al., 2021; Liousse et al., 2014; 578 Shikwambana et al., 2020). From 1990-2011, the CEDS inventory shows the largest increase in 579 580 SO<sub>2</sub> emissions in Africa starting in 2000, coincident with the increase in nss-sulfate observed at our site (see Figure 6). The emissions inventory for Africa is also likely underestimated due to a 581 lack of measurements (McDuffie et al., 2020). Therefore, we speculate that anthropogenic SO<sub>2</sub> 582 emissions are likely higher than shown from the CEDS model and are driving the observed 583 increases in nss-sulfate in Barbados. 584

In addition to sources of SO<sub>2</sub> within Africa, SO<sub>2</sub> emissions from other nearby countries and 585 regions have been shown to be exported to Africa (Koch et al., 2007). In particular, SO<sub>2</sub> emissions 586 in India have rapidly risen and overtaken China as the largest emitter of SO<sub>2</sub> (Li et al., 2017) while 587 remote sensing observations have highlighted that SO<sub>2</sub> emitted from oil and gas operations in the 588 589 Persian Gulf have been greatly underestimated in emissions inventories (McLinden et al., 2016). While it is possible that these two regions may also contribute to the increase in nss-sulfate 590 observed at our measurement site, we can only speculate of their importance to the remote North 591 592 Atlantic marine boundary layer.

593 While some of the increases in nss-sulfate after 2000 could be due, in part, to marine 594 biogenic, shipping, biomass burning, and volcanic emissions, their contributions are likely not the dominant cause of the observed trends. Marine biogenic sulfate is estimated to contribute up to 595 50% of nss-sulfate at Ragged Point during non-dust transport conditions (Li-Jones & Prospero, 596 1998; Royer et al., 2023; Savoie et al., 2002). However, our predictions of nss-sulfate and SO<sub>2</sub> 597 concentrations with EQUATES includes DMS chemistry that does not explain our observed 598 trends. In fact, locally emitted SO<sub>2</sub>, most likely from the ocean, is simulated to be decreasing from 599 2002-2011. Further, while recent studies have shown links between climate change and increased 600

DMS emissions at high latitudes, these trends have not been demonstrated at lower latitudes. 601 602 Instead, DMS is predicted to decrease with increasing temperature at low latitudes due to 603 stratification (Kloster et al., 2007) and increasing ocean acidification (Zhao et al., 2024) while a 604 recent modeling study, factoring in changes in phytoplankton dynamics, found that DMS emissions have not appreciably changed from the preindustrial to the present-day (Wang et al., 605 606 2018). As such, DMS emissions do contribute to the sulfate burden but likely do not explain the 607 recent increases in sulfate, which would also agree with findings in Bermuda where changes in the nss-sulfate burden were explained by anthropogenic rather than biogenic emissions (Keene et al., 608 2014). Shipping emissions likely do not explain our trends in nss-sulfate because Barbados is 609 somewhat isolated from proximal shipping impacts-heavy shipping is concentrated in the 610 Caribbean west of Barbados and along the north coast of South America (Czermański et al., 2021). 611 Biomass burning has declined in northern Sub-Saharan Africa starting in the early 2000s (Andela 612 613 et al., 2017; Andela & Van Der Werf, 2014; Zubkova et al., 2019). Further, increased nss-sulfate has been observed year-round rather than just during the main burn seasons, suggesting that this 614 source alone is likely not responsible for the major increases in nss-sulfate observed at our site. 615 The largest natural source of SO<sub>2</sub> in Africa is volcanic emissions from Mount Nyiragongo in the 616 Goma region of the Democratic Republic of Congo, which has been shown to impact sulfate 617 aerosol at the Amazon Tall Tower Observatory (ATTO) in Brazil (Opio et al., 2021; Saturno et 618 al., 2018). Volcanic emissions likely do impact nss-sulfate measured in Barbados but there is no 619 620 evidence that emissions from this source are increasing. These lines of evidence further support an anthropogenic source as the cause of the increase in nss-sulfate observed in Barbados. 621

In addition to increased sulfate transported from Africa, the oxidation ratio was simulated 622 to increase from 2002-2011, which would more efficiently convert locally emitted SO<sub>2</sub> to sulfate 623 and may represent a minor addition to the budget of nss-sulfate. Further, concentrations of H<sub>2</sub>O<sub>2</sub> 624 625 increased in the post-2000 period indicating a potential increase in the efficiency of aqueous phase oxidation. We would expect that even if the total burden of SO<sub>2</sub> has been reduced globally 626 (McDuffie et al., 2020; Smith et al., 2011), SO<sub>2</sub> emitted locally is more efficiently converted to 627 sulfate due to greater availability of oxidants at lower latitudes (Manktelow et al., 2007). It is 628 important to note, however, that the oxidation ratio calculated is most appropriate for accounting 629 for changing oxidation efficiencies of  $SO_2$  and sulfate formation near the site, and the oxidation 630 ratio does not account for changes in the oxidation efficiency of already formed sulfate aerosol 631 that has been transported to the site. For example, SO<sub>2</sub> emitted from Africa is likely already 632 oxidized to sulfate prior to being transported to our site. 633

One question that persists is why nitrate did not increase from 2000-2011 alongside the 634 increase in nss-sulfate? One possible explanation is a combination of reduced NO<sub>x</sub> from smoke 635 concurrent with increased dust and smoke transport that offset any changes in nitrate other than 636 the observed spikes in 2008 and 2010. In the 2000s, biomass burning emissions declined in 637 northern equatorial Africa due to a combination of increased precipitation in DJF associated with 638 a shift from more frequent El Niño events in the 1990s to more frequent La Niña events in the 639 2000s and land use practices converting tropical savanna to cropland (Andela et al., 2017; Andela 640 & Van Der Werf, 2014; Hickman et al., 2021; Zubkova et al., 2019). The recently updated 641 Barbados dust record highlights that dust is being transported to the Caribbean earlier in the year 642

and arriving more frequently in spring (Zuidema et al., 2019), which would increase the transport
of biomass burning emissions and associated nitrate to the Caribbean and remote North Atlantic,
which may effectively cancel out the impact of reduced smoke emissions.

The Ragged Point site in Barbados has historically been associated with research on 646 African dust transport (Prospero et al., 2021). However, this work highlights that the site is also 647 an excellent indicator of long-term and large-scale changes in emissions and the impact of air 648 quality policies or the lack of them or poor compliance to them. Looking forward, building upon 649 650 the existing time series of nitrate and sulfate aerosol while also expanding the measurement 651 capabilities at Ragged Point to incorporate measurements of metals, which will increase our ability to apportion aerosol sources, and carbonaceous aerosol will provide needed insight into the impact 652 653 of anthropogenic and biomass burning on sulfate and nitrate burdens over the remote North Atlantic that complement recent work performed in the South Atlantic (Zuidema et al., 2016, 654 655 2018).

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**Data Availability Statement:** Measured nitrate, nss-sulfate, and sea salt concentrations will be put on the University of Miami's repository in addition to EQUATES simulations of nitrate, nsssulfate, sea salt, gaseous tracers (SO<sub>2</sub>, benzene, CO, H<sub>2</sub>O<sub>2</sub>) and meteorological parameters. Dust mass concentrations from Barbados can be found in the data repository for Zuidema et al., 2019. EQUATES data is available via the Remote Sensing Information Gateway (RSIG): https://www.epa.gov/hesc/remote-sensing-information-gateway.

663

664 **Author Contribution:** CJG analyzed measurement and model data and wrote the manuscript. 665 JMP, LC, EB, PS collected data, operated the site, and analyzed filters for dust, sea salt, sulfate 666 and nitrate mass concentrations. KF and HOTP provided EQUATES model simulations and helped 667 with their interpretation. JAC performed HYSPLIT analysis. All authors read and edited the 668 manuscript.

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670 **Competing Interests:** The authors declare that they have no conflict of interest.

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## 684 **REFERENCES CITED**

- Aas, W., Mortier, A., Bowersox, V., Cherian, R., Faluvegi, G., Fagerli, H., Hand, J., Klimont, Z.,
  Galy-Lacaux, C., Lehmann, C. M. B., Myhre, C. L., Myhre, G., Olivié, D., Sato, K., Quaas,
- J., Rao, P. S. P., Schulz, M., Shindell, D., Skeie, R. B., Stein, A., Takemura, T., Tsyro, S.,
- 688 Vet, R., & Xu, X. (2019). Global and regional trends of atmospheric sulfur. *Scientific*
- 689 *Reports 2019 9:1, 9*(1), 1–11. https://doi.org/10.1038/s41598-018-37304-0
- Adams, A. M., Prospero, J. M., & Zhang, C. D. (2012). CALIPSO-derived three-dimensional
   structure of aerosol over the Atlantic Basin and adjacent continents. *Journal of Climate*,
   25(19), 6862–6879.
- Adams, J. W., Rodriguez, D., & Cox, R. A. (2005). The uptake of SO<sub>2</sub> on Saharan dust: A flow
  tube study. *Atmospheric Chemistry and Physics*, 5(10), 2679–2689.
  https://doi.org/10.5194/ACP-5-2679-2005
- Adams, P. J., Seinfeld, J. H., & Koch, D. M. (1999). Global concentrations of tropospheric
  sulfate, nitrate, and ammonium aerosol simulated in a general circulation model. *Journal of Geophysical Research: Atmospheres*, 104(D11), 13791–13823.
  https://doi.org/10.1029/1999JD900083
- Adebiyi, A. A., & Zuidema, P. (2016). The role of the southern African easterly jet in modifying
  the southeast Atlantic aerosol and cloud environments. *Quarterly Journal of the Royal Meteorological Society*, 142(697), 1574–1589.
- Adon, M., Galy-Lacaux, C., Yoboué, V., Delon, C., Lacaux, J. P., Castera, P., Gardrat, E.,
  Pienaar, J., Al Ourabi, H., Laouali, D., Diop, B., Sigha-Nkamdjou, L., Akpo, A., Tathy, J.
  P., Lavenu, F., & Mougin, E. (2010). Long term measurements of sulfur dioxide, nitrogen
  dioxide, ammonia, nitric acid and ozone in Africa using passive samplers. *Atmospheric Chemistry and Physics*, 10(15), 7467–7487. https://doi.org/10.5194/ACP-10-7467-2010
- Adon, M., Yobou, V., Galy-Lacaux, C., Liousse, C., Diop, B., Hadji, E., Doumbia, T., Gardrat,
  E., Ndiaye, S. A., & Jarnot, C. (2016). Measurements of NO<sub>2</sub>, SO<sub>2</sub>, NH<sub>3</sub>, HNO<sub>3</sub> and O<sub>3</sub> in
  West African urban environments. https://doi.org/10.1016/j.atmosenv.2016.03.050
- Alexander, B., Park, R. J., Jacob, D. J., Li, Q. B., Yantosca, R. M., Savarino, J., Lee, C. C. W., &
  Thiemens, M. H. (2005). Sulfate formation in sea-salt aerosols: Constraints from oxygen
  isotopes. *Journal of Geophysical Research: Atmospheres*, *110*(D10), 1–12.
  https://doi.org/10.1029/2004JD005659
- Andela, N., Morton, D. C., Giglio, L., Chen, Y., Van Der Werf, G. R., Kasibhatla, P. S., DeFries,
  R. S., Collatz, G. J., Hantson, S., Kloster, S., Bachelet, D., Forrest, M., Lasslop, G., Li, F.,

Mangeon, S., Melton, J. R., Yue, C., & Randerson, J. T. (2017). A human-driven decline in 717 global burned area. Science, 356(6345), 1356-1362. https://doi.org/10.1126/science.aal4108 718 Andela, N., & Van Der Werf, G. R. (2014). Recent trends in African fires driven by cropland 719 720 expansion and El Niño to La Niña transition. https://doi.org/10.1038/NCLIMATE2313 Andreae, M. O., Ferek, R. J., Bermond, F., Byrd, K. P., Engstrom, R. T., Hardin, S., Houmere, P. 721 D., LeMarrec, F., & Raemdonck, H. (1985). Dimethyl sulfide in the marine atmosphere. 722 Journal of Geophysical Research, 90(D7), 12891–12900. 723 Andreae, M. O. (1983). Soot Carbon and Excess Fine Potassium: Long-Range Transport of 724 Combustion-Derived Aerosols. Science, 220(4602), 1148-1151. 725 726 https://doi.org/10.1126/SCIENCE.220.4602.1148 Andreae, M. O. (2019). Emission of trace gases and aerosols from biomass burning - An updated 727 assessment. Atmospheric Chemistry and Physics, 19(13), 8523-8546. 728 https://doi.org/10.5194/ACP-19-8523-2019 729 Andreae, M.O., & Merlet, P. (2001). Emission of trace gases and aerosols from biomass burning. 730 Global Biogeochemical Cycles, 15(4), 955–966. 731 Appel, B. R., Kothny, E. L., Hoffer, E. M., Hidy, G. M., & Wesolowski, J. J. (1978). Sulfate and 732 nitrate data from California Aerosol Characterization Experiment (ACHEX). Environmental 733 Science & Technology, 12(4), 418–425. 734 Appel, K. W., Bash, J. O., Fahey, K. M., Foley, K. M., Gilliam, R. C., Hogrefe, C., Hutzell, W. 735 T., Kang, D., Mathur, R., Murphy, B. N., Napelenok, S. L., Nolte, C. G., Pleim, J. E., 736 Pouliot, G. A., Pye, H. O. T., Ran, L., Roselle, S. J., Sarwar, G., Schwede, D. B., Sidi, F. I., 737 Spero, T. L., & Wong, D. C. (2021). The Community Multiscale Air Quality (CMAQ) 738 model versions 5.3 and 5.3.1: system updates and evaluation. Geosci. Model Dev, 14, 2867-739 740 2897. https://doi.org/10.5194/gmd-14-2867-2021 Assamoi, E. M., & Liousse, C. (2010). A new inventory for two-wheel vehicle emissions in West 741 Africa for 2002. Atmospheric Environment, 44(32), 3985–3996. 742 https://doi.org/10.1016/J.ATMOSENV.2010.06.048 743 Barkley, A. E., Prospero, J. M., Mahowald, N., Hamilton, D. S., Popendorf, K. J., Oehlert, A. 744 M., Pourmand, A., Gatineau, A., Panechou-Pulcherie, K., Blackwelder, P., & Gaston, C. J. 745 (2019). African biomass burning is a substantial source of phosphorus deposition to the 746 Amazon, Tropical Atlantic Ocean, and Southern Ocean. Proceedings of the National 747 Academy of Sciences of the United States of America, 116(33), 16216–16221. 748 https://doi.org/10.1073/pnas.1906091116 749 Barkley, A. E., Olson, N. E., Prospero, J. M., Gatineau, A., Panechou, K., Maynard, N. G., 750 751 Blackwelder, P., China, S., Ault, A. P., & Gaston, C. J. (2021). Atmospheric transport of North African dust-bearing supermicron freshwater diatoms to South America: Implications 752 for iron transport to the equatorial North Atlantic Ocean. Geophysical Research Letters, 753 48(5), e2020GL090476. https://doi.org/10.1029/2020GL090476 754

- Barnes, I., Hjorth, J., & Mihalopoulos, N. (2006). Dimethyl sulfide and dimethyl sulfoxide and
  their oxidation in the atmosphere. *Chemical Reviews*, *106*(3), 940–975.
- Benish, S. E., Bash, J. O., Foley, K. M., Appel, K. W., Hogrefe, C., Gilliam, R., & Pouliot, G.
  (2022). Long-term regional trends of nitrogen and sulfur deposition in the United States
  from 2002 to 2017. *Atmospheric Chemistry and Physics*, 22(19), 12749–12767.
- 760 https://doi.org/10.5194/ACP-22-12749-2022
- Boylan, J. W., & Russell, A. G. (2006). PM and light extinction model performance metrics,
  goals, and criteria for three-dimensional air quality models. *Atmospheric Environment*, 40,
  4946–4959. https://doi.org/10.1016/j.atmosenv.2005.09.087
- Carlson, T. N., & Prospero, J. M. (1972). The large-scale movement of Saharan air outbreaks
   over the Northern Equatorial Atlantic. *Journal of Applied Meteorology, aa*(2), 283–297.
- Carslaw, K. S., Lee, L. A., Reddington, C. L., Pringle, K. J., Rap, A., Forster, P. M., Mann, G.
  W., Spracklen, D. V., Woodhouse, M. T., Regayre, L. A., & Pierce, J. R. (2013). Large
  contribution of natural aerosols to uncertainty in indirect forcing. *Nature*, *503*,
  doi:10.1038/nature12674.
- Charlson, R. J., Schwartz, S. E., Hales, J. M., Cess, R. D., Coakley, J. A., Hansen, J. E., &
  Hofmann, D. J. (1992). Climate forcing by anthropogenic aerosols. *Science*, 255(5043),
  423–430. https://doi.org/10.1126/SCIENCE.255.5043.423
- Chiapello, I., Moulin, C., & Prospero, J. M. (2005). Understanding the long-term variability of
   African dust transport across the Atlantic as recorded in both Barbados surface
- concentrations and large-scale Total Ozone Mapping Spectrometer (TOMS) optical
- thickness. *Journal of Geophysical Research: Atmospheres*, *110*(D18), 1–9.
- 777 https://doi.org/10.1029/2004JD005132
- Czermański, E., Cirella, G. T., Notteboom, T., Oniszczuk-Jastrzabek, A., & Pawłowska, B.
  (2021). An energy consumption approach to estimate air emission reductions in container
  shipping. *Energies 2021, Vol. 14, Page 278, 14*(2), 278.
- 781 https://doi.org/10.3390/EN14020278
- Doherty, O. M., Riemer, N., & Hameed, S. (2008). Saharan mineral dust transport into the
   Caribbean: Observed atmospheric controls and trends. *Journal of Geophysical Research: Atmospheres*, *113*(D7). https://doi.org/10.1029/2007JD009171
- Doherty, O. M., Riemer, N., & Hameed, S. (2012). Control of Saharan mineral dust transport to
  Barbados in winter by the Intertropical Convergence Zone over West Africa. *Journal of Geophysical Research: Atmospheres*, *117*(D19), 19117.
  https://doi.org/10.1029/2012JD017767
- Doumbia, E. H. T., Liousse, C., Keita, S., Granier, L., Granier, C., Elvidge, C. D., Elguindi, N.,
  & Law, K. (2019). Flaring emissions in Africa: Distribution, evolution and comparison with
  current inventories. *Atmospheric Environment*, *199*, 423–434.
- 792 https://doi.org/10.1016/J.ATMOSENV.2018.11.006

Draxler, R. R., & Rolph, G. D. (2011). HYSPLIT (HYbrid Single-Particle Lagrangian Integrated
Trajectory) Model access via NOAA ARL READY Website
(http://ready.arl.noaa.gov/HYSPLIT.php). NOAA Air Resources Laboratory, Silver Spring,
MD.

Foley, K. M., Pouliot, G. A., Eyth, A., Aldridge, M. F., Allen, C., Appel, K. W., Bash, J. O.,
Beardsley, M., Beidler, J., Choi, D., Farkas, C., Gilliam, R. C., Godfrey, J., Henderson, B.
H., Hogrefe, C., Koplitz, S. N., Mason, R., Mathur, R., Misenis, C., Possiel, N., Pye, H. O.
T., Reynolds, L., Roark, M., Roberts, S., Schwede, D. B., Seltzer, K. M., Sonntag, D.,

- Talgo, K., Toro, C., Vukovich, J., Xing, J., & Adams, E. (2023). 2002–2017 anthropogenic
  emissions data for air quality modeling over the United States. *Data in Brief*, 47, 109022.
- 803 https://doi.org/10.1016/J.DIB.2023.109022
- Galloway, J. N., Townsend, A. R., Erisman, J. W., Bekunda, M., Cai, Z., Freney, J. R.,
  Martinelli, L. A., Seitzinger, S. P., & Sutton, M. A. (2008). Transformation of the nitrogen
  cycle: Recent trends, questions, and potential solutions. *Science*, *320*(5878), 889–892.
  https://doi.org/10.1126/SCIENCE.1136674
- Giglio, L, Randerson, J. T., & van der Werf, G. R. (2013). Analysis of daily, monthly, and
  annual burned area using the fourth-generation global fire emissions database (GFED4). *Journal of Geophysical Research-Biogeosciences*, *118*(1), 317–328.
- Giglio, Louis, Csiszar, I., & Justice, C. O. (2006). Global distribution and seasonality of active
  fires as observed with the Terra and Aqua Moderate Resolution Imaging Spectroradiometer
  (MODIS) sensors. *Journal of Geophysical Research: Biogeosciences*, 111(2).
- 814 https://doi.org/10.1029/2005JG000142
- Goudie, A. S., & Middleton, N. J. (2001). Saharan dust storms: nature and consequences. *Earth-Science Reviews*, 56(1–4), 179–204. https://doi.org/10.1016/S0012-8252(01)00067-8
- Green, J. R., Fiddler, M. N., Holloway, J. S., Fibiger, D. L., McDuffie, E. E., Campuzano-Jost,
  P., Schroder, J. C., Jimenez, J. L., Weinheimer, A. J., Aquino, J., Montzka, D. D., Hall, S.
  R., Ullmann, K., Shah, V., Jaeglé, L., Thornton, J. A., Bililign, S., & Brown, S. S. (2019).
  Rates of wintertime atmospheric SO<sub>2</sub> oxidation based on aircraft observations during clearsky conditions over the Eastern United States. *Journal of Geophysical Research:*
- 822 *Atmospheres*, *124*(12), 6630–6649. <u>https://doi.org/10.1029/2018JD030086</u>.
- Gutleben, M., Groβ, S., Heske, C., & Wirth, M. (2022). Wintertime Saharan dust transport
   towards the Caribbean: an airborne lidar case study during EUREC<sup>4</sup>A. *Atmospheric Chemistry and Physics*, *22*, 7319-7330.
- Hand, J. L., Schichtel, B. A., Malm, W. C., & Pitchford, M. L. (2012). Particulate sulfate ion
  concentration and SO<sub>2</sub> emission trends in the United States from the early 1990s through
  2010. *Atmospheric Chemistry and Physics*, *12*(21), 10353–10365.

Hand, J. L., Prenni, A. J., & Schichtel, B. A. (2024). Trends in seasonal mean speciated aerosol
composition in remote areas of the United States From 2000 through 2021. *Journal of*

- Hickman, J. E., Andela, N., Tsigaridis, K., Galy-Lacaux, C., Ossohou, M., & Bauer, S. E.
- 834 (2021). Reductions in NO<sub>2</sub> burden over north equatorial Africa from decline in biomass
- burning in spite of growing fossil fuel use, 2005 to 2017. *Proceedings of the National*
- Academy of Sciences of the United States of America, 118(7).
- https://doi.org/10.1073/PNAS.2002579118/-/DCSUPPLEMENTAL
- Hopkins, J. R., Evans, M. J., Lee, J. D., Lewis, A. C., Marsham, J. H., McQuaid, J. B., Parker, D.
  J., Stewart, D. J., Reeves, C. E., & Purvis, R. M. (2009). Direct estimates of emissions from
  the megacity of Lagos. *Atmospheric Chemistry and Physics*, 9(21), 8471–8477.
  https://doi.org/10.5194/ACP-9-8471-2009
- 842 *https://gispub.epa.gov/air/trendsreport/2023/.* (2023).
- Jickells, T. D., Buitenhuis, E., Altieri, K., Baker, A. R., Capone, D., Duce, R. A., Dentener, F.,
- Fennel, K., Kanakidou, M., LaRoche, J., Lee, K., Liss, P., Middelburg, J. J., Moore, J. K.,
  Okin, G., Oschlies, A., Sarin, M., Seitzinger, S., Sharples, J., Singh, A., Suntharalingam, P.,
- Okin, G., Oschlies, A., Sarin, M., Seitzinger, S., Sharples, J., Singh, A., Suntharalingam, P.
  Uematsu, M., & Zamora, L. M. (2017). A reevaluation of the magnitude and impacts of
  anthropogenic atmospheric nitrogen inputs on the ocean. *Global Biogeochemical Cycles*,
  21(2), 280, 205, https://doi.org/10.1002/2016GP005586
- 848 *31*(2), 289–305. https://doi.org/10.1002/2016GB005586
- Keene, W. C., Pszenny, A. A. P., Galloway, J. N., & Hawley, M. E. (1986). Sea-salt corrections
  and interpretation of constituent ratios in marine precipitation. *Journal of Geophysical Research*, *91*(D6), 6647. https://doi.org/10.1029/JD091ID06P06647
- Keene, W. C., Moody, J. L., Galloway, J. N., Prospero, J. M., Cooper, O. R., Eckhardt, S., &
  Maben, J. R. (2014). Long-term trends in aerosol and precipitation composition over the
  western North Atlantic Ocean at Bermuda. *Atmospheric Chemistry and Physics*, *14*(15),
  8119–8135.
- Keita, S., Liousse, C., Assamoi, E. M., Doumbia, T., N'Datchoh, E. T., Gnamien, S., Elguindi,
  N., Granier, C., & Yoboué, V. (2021). African anthropogenic emissions inventory for gases
  and particles from 1990 to 2015. *Earth System Science Data*, *13*(7), 3691–3705.
  https://doi.org/10.5194/ESSD-13-3691-2021
- Kelly, J. T., Koplitz, S. N., Baker, K. R., Holder, A. L., Pye, H. O. T., Murphy, B. N., Bash, J.
  O., Henderson, B. H., Possiel, N. C., Simon, H., Eyth, A. M., Jang, C., Phillips, S., &
  Timin, B. (2019). Assessing PM<sub>2.5</sub> model performance for the conterminous U.S. with
  comparison to model performance statistics from 2007-2015. *Atmos Environ*, *214*,
  https://doi.org/10.1016/j.atmosenv.2019.116872.
- Kganyago, M., & Shikwambana, L. (2019). Assessing spatio-temporal variability of wildfires
  and their impact on Sub-Saharan ecosystems and air quality using multisource remotely
  sensed data and trend analysis. *Sustainability*, *11*, 6811, doi:10.3390/su11236811.

*Geophysical Research: Atmospheres*, *129*(2), e2023JD039902.
 https://doi.org/10.1029/2023JD039902

- Kharol, S. K., Mclinden, C. A., Sioris, C. E., Shephard, M. W., Fioletov, V., Van Donkelaar, A.,
  Philip, S., & Martin, R. V. (2017). OMI satellite observations of decadal changes in groundlevel sulfur dioxide over North America. *Atmos. Chem. Phys*, *17*, 5921–5929.
  https://doi.org/10.5194/acp-17-5921-2017
- Kloster, S., Six, K. D., Feichter, J., Maier-Reimer, E., Roeckner, E., Wetzel, P., Stier, P., & Esch,
- M. (2007). Response of dimethylsulfide (DMS) in the ocean and atmosphere to global
- 874 warming. Journal of Geophysical Research: Biogeosciences, 112(G3).
- 875 https://doi.org/10.1029/2006JG000224
- Koch, D., Bond, T. C., Streets, D., Unger, N., & van der Werf, G. R. (2007). Global impacts of
  aerosols from particular source regions and sectors. *Journal of Geophysical Research: Atmospheres*, *112*(D2), 2205. https://doi.org/10.1029/2005JD007024
- Kramer, S. J., Kirtman, B. P., Zuidema, P., & Ngan, F. (2020). Subseasonal variability of
  elevated dust concentrations over South Florida. *Journal of Geophysical Research: Atmospheres*, *125*(6). https://doi.org/10.1029/2019JD031874
- Lee, C., Martin, R. V., Van Donkelaar, A., Lee, H., Dickerson, R. R., Hains, J. C., Krotkov, N.,
  Richter, A., Vinnikov, K., & Schwab, J. J. (2011). SO<sub>2</sub> emissions and lifetimes: Estimates
  from inverse modeling using in situ and global, space-based (SCIAMACHY and OMI)
  observations. *Journal of Geophysical Research Atmospheres*, *116*(6).
  https://doi.org/10.1029/2010JD014758
- Lelieveld, J., & Heintzenberg, J. (1992). Sulfate cooling effect on climate through in-cloud
   processing of anthropogenic SO<sub>2</sub>. *Science*, *258*, 117–120.
- Lelieveld, J., Berresheim, H., Borrmann, S., Crutzen, P. J., Dentener, F. J., Fischer, H., Feichter,
  J., Flatau, P. J., Heland, J., Holzinger, R., Korrmann, R., Lawrence, M. G., Levin, Z.,
- J., Flatau, P. J., Heland, J., Holzinger, R., Korrmann, R., Lawrence, M. G., Levin, Z.,
  Markowicz, K. M., Mihalopoulos, N., Minikin, A., Ramanathan, V., De Reus, M., Roelofs,
- 6. J., Scheeren, H. A., Sciare, J., Schlager, H., Schultz, M., Siegmund, P., Steil, B.,
- 893 Stephanou, E. G., Stier, P., Traub, M., Warneke, C., Williams, J., & Ziereis, H. (2002).
- Global air pollution crossroads over the Mediterranean. *Science*, 298(5594), 794–799.
- 895 https://doi.org/10.1126/SCIENCE.1075457/ASSET/9D6C7086-E45D-4A4D-9892-
- 896 2054F591C047/ASSETS/GRAPHIC/SE4120969007.JPEG
- Lelieveld, J., Hoor, P., Jöckel, P., Pozzer, A., Hadjinicolaou, P., Cammas, J. P., & Beirle, S.
  (2009). Severe ozone air pollution in the Persian Gulf region. *Atmospheric Chemistry and Physics*, 9(4), 1393–1406. https://doi.org/10.5194/ACP-9-1393-2009
- Li, C., McLinden, C., Fioletov, V., Krotkov, N., Carn, S., Joiner, J., Streets, D., He, H., Ren, X.,
  Li, Z., & Dickerson, R. R. (2017). India Is overtaking China as the world's largest emitter
  of anthropogenic sulfur dioxide. *Scientific Reports 2017 7:1*, 7(1), 1–7.
- 903 https://doi.org/10.1038/s41598-017-14639-8

Li-Jones, X., & Prospero, J. M. (1998). Variations in the size distribution of non-sea-salt sulfate 904 905 aerosol in the marine boundary layer at Barbados: Impact of African dust. Journal of 906 Geophysical Research-Atmospheres, 103(D13), 16073–16084. 907 Liousse, C., Assamoi, E., Criqui, P., Granier, C., & Rosset, R. (2014). Explosive growth in African combustion emissions from 2005 to 2030. Environmental Research Letters, 9(3). 908 Mahowald, N., Scanza, R., Brahney, J., Goodale, C. L., Hess, P. G., Moore, J. K., & Neff, J. 909 (2017). Aerosol deposition impacts on land and ocean carbon cycles. Curr Clim Change 910 911 *Rep*, *3*, 16–31. Manktelow, P. T., Mann, G. W., Carslaw, K. S., Spracklen, D. V., & Chipperfield, M. P. (2007). 912 Regional and global trends in sulfate aerosol since the 1980s. Geophysical Research Letters, 913 34(14), 14803. https://doi.org/10.1029/2006GL028668 914 Mathur, R., Xing, J., Gilliam, R., Sarwar, G., Hogrefe, C., Pleim, J., Pouliot, G., Roselle, S., 915 Spero, T. L., Wong, D. C., & Young, J. (2017). Extending the Community Multiscale Air 916 Quality (CMAQ) modeling system to hemispheric scales: overview of process 917 considerations and initial applications. Atmos. Chem. Phys, 17, 12449–12474. 918 https://doi.org/10.5194/acp-17-12449-2017 919 McDuffie, E. E., Smith, S. J., O'Rourke, P., Tibrewal, K., Venkataraman, C., Marais, E. A., 920 Zheng, B., Crippa, M., Brauer, M., & Martin, R. V. (2020). A global anthropogenic 921 emission inventory of atmospheric pollutants from sector- And fuel-specific sources (1970-922 2017): An application of the Community Emissions Data System (CEDS). Earth System 923 Science Data, 12(4), 3413-3442. https://doi.org/10.5194/ESSD-12-3413-2020 924 McLinden, C. A., Fioletov, V., Shephard, M. W., Krotkov, N., Li, C., Martin, R. V, Moran, M. 925 D., & Joiner, J. (2016). Space-based detection of missing sulfur dioxide sources of global 926 air pollution. https://doi.org/10.1038/NGEO2724 927 Murphy, D. M., Cziczo, D. J., Froyd, K. D., Hudson, P. K., Matthew, B. M., Middlebrook, A. 928 M., Peltier, R. E., Sullivan, A., Thomson, D. S., & Weber, R. J. (2006). Single-particle mass 929 spectrometry of tropospheric aerosol particles. Journal of Geophysical Research-930 Atmospheres, 111(D23), D23S32, doi:10.1029/2006JD007340. 931 932 Nowell, H. K., Holmes, C. D., Robertson, K., Teske, C., & Hiers, J. K. (2018). A new picture of fire extent, variability, and drought interaction in prescribed fire landscapes: Insights from 933 Florida Government Records. Geophysical Research Letters, 45(15), 7874–7884. 934 https://doi.org/10.1029/2018GL078679 935 Opio, R., Mugume, I., & Nakatumba-Nabende, J. (2021). Understanding the Trend of NO<sub>2</sub>, SO<sub>2</sub> 936 937 and CO over East Africa from 2005 to 2020. Atmosphere 2021, Vol. 12, Page 1283, 12(10), 1283. https://doi.org/10.3390/ATMOS12101283 938 Osuji, L. C., & Avwiri, G. O. (2005). Flared gases and other pollutants associated with air 939 quality in industrial areas of Nigeria: An overview. Chemistry & Biodiversity, 2(10), 1277-940 1289. https://doi.org/10.1002/CBDV.200590099 941

- Perron, M. M. G., Meyerink, S., Corkill, M., Strzelec, M., Proemse, B. C., Gault-Ringold, M.,
  Sanz Rodriguez, E., Chase, Z., & Bowie, A. R. (2022). Trace elements and nutrients in
  wildfire plumes to the southeast of Australia. *Atmospheric Research*, *270*, 106084.
  https://doi.org/10.1016/J.ATMOSRES.2022.106084
- Prospero, J. M. (1999). Long-term measurements of the transport of African mineral dust to the
  southeastern United States: Implications for regional air quality. *Journal of Geophysical Research-Atmospheres*, *104*(D13), 15917–15927.
- Prospero, J. M., & Mayol-Bracero, O. L. (2013). Understanding the transport and impact of
  African dust on the Caribbean Basin. *Bulletin of the American Meteorological Society*,
  951 94(9), 1329–1337.
- Prospero, J. M., Glaccum, R. A., & Nees, R. T. (1981). Atmospheric transport of soil dust from
  Africa to South America. *Nature*, 289, 570–572.
- Prospero, J. M., Olmez, I., & Ames, M. (2001). Al and Fe in PM<sub>2.5</sub> and PM<sub>10</sub> suspended
  particles in South-Central Florida: The impact of the long range transport of African
  mineral dust.
- Prospero, J. M., Collard, F.-X., Molinie, J., & Jeannot, A. (2014). Characterizing the annual
  cycle of African dust transport to the Caribbean Basin and South America and its impact on
  the environment and air quality. *Global Biogeochemical Cycles*, 29, 757–773.
  https://doi.org/10.1111/1462-2920.13280
- Prospero, J. M., Delany, A. C., Delany, A. C., & Carlson, T. N. (2021). The discovery of African dust transport to the western hemisphere and the Saharan Air Layer: A history. *Bulletin of the American Meteorological Society*, *102*(6), E1239–E1260.
  https://doi.org/10.1175/BAMS-D-19-0309.1
- Quinn, P. K., Bates, T. S., Coffman, D. J., Upchurch, L. M., Johnson, J. E., Brewer, A., Baidar,
  S., McCoy, I. L., & Zuidema, P. (2022). Wintertime observations of tropical northwest
  Atlantic aerosol properties during ATOMIC: Varying mixtures of dust and biomass
  burning. *Journal of Geophysical Research: Atmospheres*, *127*(8), e2021JD036253.
  https://doi.org/10.1029/2021JD036253
- 970 Quinn, Patricia K., Thompson, E. J., Coffman, D. J., Baidar, S., Bariteau, L., Bates, T. S., Bigorre, S., Brewer, A., De Boer, G., De Szoeke, S. P., Drushka, K., Foltz, G. R., Intrieri, J., 971 Iyer, S., Fairall, C. W., Gaston, C. J., Jansen, F., Johnson, J. E., Krüger, O. O., Marchbanks, 972 R. D., Moran, K. P., Noone, D., Pezoa, S., Pincus, R., Plueddemann, A. J., Pöhlker, M. L., 973 Pöschl, U., Melendez, E. Q., Royer, H. M., Szczodrak, M., Thomson, J., Upchurch, L. M., 974 Zhang, C., Zhang, D., & Zuidema, P. (2021). Measurements from the RV Ronald H. Brown 975 and related platforms as part of the Atlantic Tradewind Ocean-Atmosphere Mesoscale 976 Interaction Campaign (ATOMIC). Earth System Science Data, 13(4), 1759–1790. 977 https://doi.org/10.5194/ESSD-13-1759-2021 978

- Rafaj, P., Amann, M., Siri, J., & Wuester, H. (2015). Changes in European greenhouse gas and
  air pollutant emissions 1960–2010: decomposition of determining factors. *Uncertainties in Greenhouse Gas Inventories*, 27–54. https://doi.org/10.1007/978-3-319-15901-0\_3
- Rickly, P. S., Guo, H., Campuzano-Jost, P., Jimenez, J. L., Wolfe, G. M., Bennett, R., Bourgeois,
  I., Crounse, J. D., Dibb, J. E., DiGangi, J. P., Diskin, G. S., Dollner, M., Gargulinski, E. M.,
  Hall, S. R., Halliday, H. S., Hanisco, T. F., Hannun, R. A., Liao, J., Moore, R., Nault, B. A.,
  Nowak, J. B., Peischl, J., Robinson, C. E., Ryerson, T., Sanchez, K. J., Schöberl, M., Soja,
  A. J., St Clair, J. M., Thornhill, K. L., Ullmann, K., Wennberg, P. O., Weinzierl, B.,
  Wiggins, E. B., Winstead, E. L., & Rollins, A. W. (2022). Emission factors and evolution of
- 988 SO<sub>2</sub> measured from biomass burning in wildfires and agricultural fires. *Atmos. Chem. Phys*,
- 989 22, 15603–15620. https://doi.org/10.5194/acp-22-15603-2022
- Roberts, G., Wooster, M. J., & Lagoudakis, E. (2009). Annual and diurnal African biomass
  burning temporal dynamics. *Biogeosciences*, 6(5), 849–866. https://doi.org/10.5194/BG-6849-2009
- Rolph, G., Stein, A., & Stunder, B. (2017). Real-time Environmental Applications and Display
  sYstem: READY. *Environmental Modelling & Software*, 95, 210–228.
- Royer, H. M., Pöhlker, M. L., Krüger, O., Blades, E., Sealy, P., Lata, N. N., Cheng, Z., China, S.,
  Ault, A. P., Quinn, P. K., Zuidema, P., Pöhlker, C., Pöschl, U., Andreae, M., & Gaston, C.
  J. (2023). African smoke particles act as cloud condensation nuclei in the wintertime
  tropical North Atlantic boundary layer over Barbados. *Atmospheric Chemistry and Physics*,
  23(2), 981–998. https://doi.org/10.5194/ACP-23-981-2023
- Salvador, P., Almeida, S.M., Cardoso, J., Almeida-Silva, M., Nunes, T., Cerqueira, M., Alves,
   C., Reis, M.A., Chaves, P.C., Artínano, B., & Pio, C. (2015). Composition and origin of
   PM<sub>10</sub> in Cape Verde: Characterization of long-range transport episodes. *Atmospheric Environment*, 127, 326-339.
- Sarwar, G., Fahey, K. M., Napelenok, S. L., Roselle, S. J., & Mathur, R. (2011). Examining the
   impact of CMAQ model updates on aerosol sulfate predictions. In *The 10th Annual CMAS Models-3 User's Conference* (p. vol 775). Chapel Hill, NC.
- Sarwar, Golam, Hogrefe, C., Henderson, B. H., Foley, K., Mathur, R., Murphy, B., & Ahmed, S.
   (2023). Characterizing variations in ambient PM<sub>2.5</sub> concentrations at the U.S. Embassy in
   Dhaka, Bangladesh using observations and the CMAQ modeling system. *Atmospheric Environment*, 296, 119587. https://doi.org/10.1016/J.ATMOSENV.2023.119587
- Saturno, J., Ditas, F., Penning De Vries, M., Holanda, B. A., Pöhlker, M. L., Carbone, S., Walter,
  D., Bobrowski, N., Brito, J., Chi, X., Gutmann, A., Hrabe De Angelis, I., Machado, L. A.
  T., Moran-Zuloaga, D., Rüdiger, J., Schneider, J., Schulz, C., Wang, Q., Wendisch, M.,
  Artaxo, P., Wagner, T., Pöschl, U., Andreae, M. O., & Pöhlker, C. (2018). African volcanic
  emissions influencing atmospheric aerosol particles over the Amazon rain forest.
- 1016 Atmospheric Chemistry and Physics, 18, 10391–10405. https://doi.org/10.5194/acp-2017-
- 1017 1152

Savoie, D. L., & Prospero, J. M.; (1980). Water-soluble potassium, calcium, and magnesium in
the aerosols over the tropical North Atlantic. *Journal of Geophysical Research: Oceans*,
85(C1), 385–392. https://doi.org/10.1029/JC085IC01P00385

- Savoie, D. L., & Prospero, J. M. (1982). Particle size distribution of nitrate and sulfate in the
   marine atmosphere. *Geophysical Research Letters*, 9(10), 1207–1210.
- Savoie, D.L., Arimoto, R., Keene, W. C., Prospero, J. M., Duce, R. A., & Galloway, J. N.
   (2002). Marine biogenic and anthropogenic contributions to non-sea-salt sulfate in the
   marine boundary layer over the North Atlantic Ocean. *Journal of Geophysical Research- Atmospheres*, 107(D18), 4356, doi:10.1029/2001JD000970.
- Scheuvens, D., Schutz, L., Kandler, K., Ebert, M., & Weinbruch, S. (2013). Bulk composition of
   northern African dust and its source sediments A compilation. *Earth-Science Reviews*,
   116, 170–194.
- Schlosser, J. S., Braun, R. A., Bradley, T., Dadashazar, H., MacDonald, A. B., Aldhaif, A. A.,
  Aghdam, M. A., Mardi, A. H., Xian, P., & Sorooshian, A. (2017). Analysis of aerosol
  composition data for western United States wildfires between 2005 and 2015: Dust
  emissions, chloride depletion, and most enhanced aerosol constituents. *Journal of Geophysical Research: Atmospheres*, *122*(16), 8951–8966.
  https://doi.org/10.1002/2017JD026547
- Shah, V., Jaeglé, L., Thornton, J. A., Lopez-Hilfiker, F. D., Lee, B. H., Schroder, J. C.,
  Campuzano-Jost, P., Jimenez, J. L., Guo, H., Sullivan, A. P., Weber, R. J., Green, J. R.,
  Fiddler, M. N., Bililign, S., Campos, T. L., Stell, M., Weinheimer, A. J., Montzka, D. D., &
  Brown, S. S. (2018). Chemical feedbacks weaken the wintertime response of particulate
  sulfate and nitrate to emissions reductions over the eastern United States. *Proceedings of the National Academy of Sciences of the United States of America*, *115*(32), 8110–8115.
  https://doi.org/10.1073/PNAS.1803295115
- Shikwambana, L., Mhangara, P., & Mbatha, N. (2020). Trend analysis and first time
  observations of sulphur dioxide and nitrogen dioxide in South Africa using
  TROPOMI/Sentinel-5 P data. *International Journal of Applied Earth Observation and Geoinformation*, 91, 102130. https://doi.org/10.1016/J.JAG.2020.102130
- Smith, S. J., Van Aardenne, J., Klimont, Z., Andres, R. J., Volke, A., & Delgado Arias, S.
  (2011). Anthropogenic sulfur dioxide emissions: 1850-2005. *Atmospheric Chemistry and Physics*, *11*(3), 1101–1116. https://doi.org/10.5194/ACP-11-1101-2011
- Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., & Ngan, F. (2015).
   NOAA's HYSPLIT Atmospheric Transport and Dispersion Modeling System. *Bulletin of the American Meteorological Society*, *96*(12), 2059–2077.

Taylor, S. R., & McLennan, S. M.; (1985). *The continental crust: Its composition and evolution*.
 Oxford: Blackwell Scientific Publications.

Trapp, J. M., Millero, F. J., & Prospero, J. M. (2010). Trends in the solubility of iron in dust-1055 1056 dominated aerosols in the equatorial Atlantic trade winds: Importance of iron speciation and 1057 sources. Geochemistry, Geophysics, Geosystems, 11(Q03014), doi:10.1029/2009GC002651. 1058 Tsamalis, C., Chédin, A., Pelon, J., & Capelle, V. (2013). The seasonal vertical distribution of 1059 the Saharan Air Layer and its modulation by the wind. Atmospheric Chemistry and Physics, 1060 13(22), 11235–11257. https://doi.org/10.5194/ACP-13-11235-2013 Val, S., Liousse, C., Doumbia, E. H. T., Galy-Lacaux, C., Cachier, H., Marchand, N., Badel, A., 1061 1062 Gardrat, E., Sylvestre, A., & Baeza-Squiban, A. (2013). Physico-chemical characterization 1063 of African urban aerosols (Bamako in Mali and Dakar in Senegal) and their toxic effects in human bronchial epithelial cells: Description of a worrying situation. Particle and Fibre 1064 1065 Toxicology, 10(1), 1–16. https://doi.org/10.1186/1743-8977-10-10/FIGURES/8 1066 Vannucci, P. F., Foley, K., Murphy, B. N., Hogrefe, C., Cohen, R. C., & Pye, H. O. T. (2024). Temperature-dependent composition of summertime PM<sub>2.5</sub> in observations and model 1067 1068 predictions across the Eastern U.S. ACS Earth and Space Chemistry, Submitted. Vasilakos, P., Russell, A., Weber, R., & Nenes, A. (2018). Understanding nitrate formation in a 1069 world with less sulfate. Atmospheric Chemistry and Physics, 18(17), 12765–12775. 1070 1071 https://doi.org/10.5194/ACP-18-12765-2018 Wang, S., Maltrud, M., Elliott, S., Cameron-Smith, P., & Jonko, A. (2018). Influence of 1072 dimethyl sulfide on the carbon cycle and biological production. Biogeochemistry, 138, 49-1073 1074 68. https://doi.org/10.1007/s10533-018-0430-5 Van der Werf, G. R., Randerson, J. T., Collatz, G. J., & Giglio, L. (2003). Carbon emissions 1075 from fires in tropical and subtropical ecosystems. Global Change Biology, 9(4), 547-562. 1076 https://doi.org/10.1046/J.1365-2486.2003.00604.X 1077 Wex, H., Dieckmann, K., Roberts, G. C., Conrath, T., Izaguirre, M. A., Hartmann, S., Herenz, P., 1078 Schafer, M., Ditas, F., Schmeissner, T., Henning, S., Wehner, B., Siebert, H., & Stratmann, 1079 F. (2016). Aerosol arriving on the Caribbean island of Barbados: physical properties and 1080 origin. Atmospheric Chemistry and Physics, 16(22), 14107–14130. 1081 Wiedinmyer, C., Akagi, S. K., Yokelson, R. J., Emmons, L. K., Al-Saadi, J. A., Orlando, J. J., & 1082 1083 Soja, A. J. (2011). The Fire INventory from NCAR (FINN): a high resolution global model to estimate the emissions from open burning. Geosci. Model Dev, 4, 625-641. 1084 https://doi.org/10.5194/gmd-4-625-2011 1085 Wolter, K., & Timlin, M. S. (2011). El Niño/Southern Oscillation behaviour since 1871 as 1086 diagnosed in an extended multivariate ENSO index (MEI.ext). International Journal of 1087 1088 *Climatology*, *31*(7), 1074–1087. https://doi.org/10.1002/JOC.2336 Yang, Y., Lou, S., Wang, H., Wang, P., & Liao, H. (2020). Trends and source apportionment of 1089 aerosols in Europe during 1980-2018. Atmospheric Chemistry and Physics, 20(4), 2579-1090 2590. https://doi.org/10.5194/ACP-20-2579-2020 1091

Zhang, T., Hoell, A., Perlwitz, J., Eischeid, J., Murray, D., Hoerling, M., & Hamill, T. M. 1092 1093 (2019). Towards probabilistic multivariate ENSO monitoring. Geophysical Research 1094 Letters, 46(17-18), 10532-10540. https://doi.org/10.1029/2019GL083946 1095 Zhao, B., Jiang, J. H., Gu, Y., Diner, D., Worden, J., Liou, K. N., Su, H., Xing, J., Garay, M., & 1096 Huang, L. (2017). Decadal-scale trends in regional aerosol particle properties and their 1097 linkage to emission changes. Environmental Research Letters, 12(5), 054021. https://doi.org/10.1088/1748-9326/AA6CB2 1098 1099 Zhao, J., Zhang, Y., Bie, S., Bilsback, K. R., Pierce, J. R., & Chen, Y. (2024). Changes in global 1100 DMS production driven by increased CO<sub>2</sub> levels and its impact on radiative forcing. Npj Climate and Atmospheric Science 2024 7:1, 7(1), 1-8. https://doi.org/10.1038/s41612-024-1101 1102 00563-y Zubkova, M., Boschetti, L., Abatzoglou, J. T., & Giglio, L. (2019). Changes in fire activity in 1103 1104 Africa from 2002 to 2016 and their potential drivers. Geophysical Research Letters, 46, 1105 7643-7653. https://doi.org/10.1029/2019GL083469 Zuidema, P., Redemann, J., Haywood, J., Wood, R., Piketh, S., Hipondoka, M., & Formenti, P. 1106 (2016). Smoke and clouds above the Southeast Atlantic: Upcoming field campaigns probe 1107 1108 absorbing aerosol's impact on climate. Bulletin of the American Meteorological Society, 97(7), 1131-1135. https://doi.org/10.1175/BAMS-D-15-00082.1 1109 Zuidema, P., Sedlacek, A. J., Flynn, C., Springston, S., Delgadillo, R., Zhang, J., Aiken, A. C., 1110 Koontz, A., & Muradyan, P. (2018). The Ascension Island boundary layer in the remote 1111 Southeast Atlantic is often smoky. Geophysical Research Letters, 45(9), 4456-4465. 1112 https://doi.org/10.1002/2017GL076926 1113 Zuidema, P., Alvarez, C., Kramer, S. J., Custals, L., Izaguirre, M., Sealy, P., Prospero, J. M., & 1114 Blades, E. (2019). Is summer African dust arriving earlier to Barbados? The updated long-1115 term in situ dust mass concentration time series from Ragged Point, Barbados, and Miami, 1116

- 1117 Florida. Bulletin of the American Meteorological Society, 100(10), 1981–1986.
- 1118 https://doi.org/10.1175/BAMS-D-18-0083.1
- 1119