Improved Improving winter conditions simulations in SURFEX-TEB v9.0 with a multi-layer snow model and icefor road winter maintenance

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Abstract. Snow-covered or icy roads increase the risk of accidents for drivers, pedestrians, and cyclists. In cities or in remote areas, to prevent these slippery conditions, road winter maintenance decisions are weather-informed by simulations. Numerical In winter, snow and ice covered artificial surfaces are important aspects of the urban climate and trigger road maintenance operations. Urban climate and road weather models have been developed for this purpose, and mostly built to simulate the road conditions in open environments without shadowing and reflections effects specialised in simulating these conditions in cities or in the countryside, respectively. In this study, we intent-intend to bridge the gap between road weather models and urban climate models to improve in terms of cold regions urban modeling and road modelling and artificial surface condition predictions in any environment. We have refined the modelling of road surface processes related to winter conditions in the Town Energy Balance (surface externalisée; SURFEX-TEB v9.0), which is TEB), an urban climate model used for complex environment modeling. For icy conditions, we designed for complex environments. We have developed an ice content prediction to account for the freezing and melting of the water content on the surface. Additionally, we enhanced TEB's have enhanced the TEB representation of snow on road, previously relying on moving from a single-layer snow model (1-L), with to a more precise multi-layer snow model known as Explicit Snow (ES). We have conducted evaluations at two distinct locations: isolated the winter surface processes from other physical interactions by limiting the evaluation of the experiments to open environments. The experiments are carried out at two locations: the Col de Porte in the Alps and a road weather station in southern Finland. Our findings shows show that the enhanced TEB model (named TEB-ES) outperforms TEB, as well as two benchmark models benchmark model, ISBA-Route/CROCUS, and but with mixed results against a multiple linear regression in open environments. This results are promising for using TEB to inform road winter maintenance decisions in-sample algorithm. For roads with high-traffic and/or winter maintenance operations, future modelling work should focus on the representations of anthropogenic effects.

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1 Introduction

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For the past 20 years, meteorological offices in countries impacted by winter conditions have provided road and airport runway weather forecasts to operators during winter maintenance season. These forecasts are valuable to them, as they facilitate the planning of deiging activities, traffic optimization and snow removal. They also facilitate advance planning of preventive measures that should be conducted before the road surface conditions become too dangerous for users There are significant interconnections between urban climate and winter conditions. As in summer, high-density building distribution can induce a pronounced Urban Heat Island effect (UHI). Few studies tried to study its extent, and unveil its drivers. For instance, Maleyich and Klink (?) measured during winter 2008-2009, an average winter UHI of 1°C at Minneapolis. In Alaska, in a small settlement of 4600 residents, an average 2.2°C winter UHI was measured (?). In a typical Artic city, Varenstov (?) measured an average UHI of 1.9°C with extremes up to 11°C in Apatity city center. Hinkel et al. (?), Varenstov et al. (?), and Bohnenstengel et al. (?) agree to say that the anthropogenic heat, mainly released mainly from house heating, is strongly correlated with the UHI magnitude. But its relative impact with other potential drivers on the UHI magnitude is still unclear and need for further studies. The snow cover which can hold for months on artificial surfaces after some large snowfall, may play an important role in this phenomenon. A few studies based on simulations have shown that the presence of snow cover decreased the surface air temperature (?, ?; ?, ?), whereas on the contrary Malevich et al. (?) shown an increased UHI. It is clear that specific events of cold conditions such as snowfall and freezing temperatures have a considerable influence on the city climate (?). The reverse is true, as some evidence suggests that the urban heat island decreases the amount of snowfall and increases the amount of rainfall (?)

40 The presence of a layer of ice or snow insulates the artificial surfaces from the atmosphere by covering significant parts of the city surface. Therefore, covered artificial surfaces temperatures evolve by conduction. The soil atmosphere interaction is driven by the changing properties of the ageing snow, instead of the artificial surface materials. Lemonsu et al. (?) have shown that snow-covered urban surfaces contribute to changes in sensible and latent heat fluxes. The snow layer stores the incoming energy, modifies the surface output fluxes, and releases energy by melting snow or by sensible and latent heat fluxes (?). The 45 Net surface radiation seems to be greatly impacted due to the high albedo of the snow (?). Depending on the characteristics of the snow, the incoming energy from the atmosphere can be stored, released as heat flux, or as melting snow, as measured in the China case study of Taotao et al. (?). In addition, the ice and snow cover acts as a water reservoir. It delays precipitation runoff and drives urban hydrology (?). These conditions lead to concrete impacts on human activities because of the slippery nature of frozen water. They increase the risk of accidents for drivers, pedestrians, and bicycles. According to Michaelides et al. (?), the risk of accidents on icy and slippery road-roads is 2-3 times higher than on dry roads. In Sweden, Andersson and Chapman (?) showed that the accidents were most accidents were more frequent under winter conditions with a road surface temperature below -3°C and snow-covered or icy road surfaces roads. Remote areas are the most vulnerable, and even light snowfall can have serious consequences in countries not adapted accustomed to these conditions (?). Every year, a significant amount of deicing agents is applied to road surfaces. Pollutants can find their way into natural ecosystems and disrupt the delicate balance of nature. During 2022 in the French Alsace department, 24 000 tonnes of salt have been stored for deicing measures. Also, 55

500 agents and 121 machines have been rallied for the season (?). Proper utilization of this fleet is essential to maintain road safety across the extensive road network. Thus, the costs associated with dangerous weather conditions in the road sector are significantly higher than those in rail and aviation (?). Accurate simulations of the winter conditions on artificial surfaces can help to increase our knowledge of cold urban conditions and increase the safety of artificial surfaces used by citizens.

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Many factors affect the pavement conditions through various and sometimes complex processes (?). Numerical modeling tools are used to monitor surface conditions and plan road maintenance. Physical and statistical models with varying levels of refinement Simulation tools are needed to simultaneously represent the specific properties of urban environments as a coupling of human activities with the specific processes associated with cold conditions. Land Surface Models (LSM) based on the physical heat-balance equation are well suited as they can represent various surface types ranging from natural to artificial surfaces in urban environments. Mainly used to provide boundary conditions for atmospheric models, LSM are key for the prediction of soil-atmosphere fluxes. Several LSMs have been developed to accurately predict road conditions. Mostly, road weather forecasts are simulated by Land Surface Models (LSM), but in recent years many statistical modelshave also been created specifically for urban environments; the urban climate models.

Empirical relationships have been established between the various variables, which influence road conditions. These statistical models often produce satisfactory results with few input variables. They are also easy to implement. Different machine learning algorithms have been used for road surface temperature prediction like simple linear regression (?, ?; ?, ?), GAM (?) or more recently gradient boosting algorithms (?) and neural networks (?). Models that directly predict the road conditions, have also been developed with recurrent neural networks (?) or random forests methods (?). Statistical models inherently fail to capture the joint physical evolution of the variables that influence road conditions. They are also considered to have lower performance in complex environments such as urban areas with multiple shading and anthropic effects where physical models can perform better (?). For the reasons described previously, we will focus on heat balance models in this study.

Land Surface Models (LSM) simulate surface variables using the heat balance equation. They can represent various surface types from natural surfaces to urban environment with urban elimate models. Mainly used to provide boundary conditions for atmospheric models, LSM surfaceconditions are key for the prediction of soil atmosphere fluxes In urban environments, specific physical processes are needed to represent the town energetics (?). Simple building-averaged models are able to compute radiative trapping, surface energy budgets, and wind channelling (?). They have demonstrated the ability to simulate the urban heat island in summer driven by urban morphology and the capacity of artificial materials to store energy like the TEB model (?). In comparison, the modelling of urban winter conditions has been less studied in the urban climate community (?). Lemonsu et al. (?) showed that coupling the road with a simple one-layer snow model in an urban climate model (TEB) leads to improved fluxes in winter. The SUEWS (?), TEB (?) and Lodz-SUEB (?) models include a one-layer snow model to take the effect of snow into account during wintertime. CLMU (?) and JULES (?) go further and include a multi-layer snow model on the road. Karsisto et al. (?) compared TEB, SUEWS, and CLM at two Helsinki sites and showed that the snow-covered ground fraction plays a major role in winter and spring for the model fluxes performance. More studies are needed to model and compare the winter conditions with the observations. In particular, the key processes driving the urban response at the surface: snow cover evolution, ice layer evolution, and human activities. To that extent, road weather models are also well

suited to simulate the artificial surface response to winter conditions. Many national weather services run land surface models specifically designed to help road winter maintenance. These so-called road weather models focus on integrating various factors affecting that affect the evolution of road conditions (?), including the difficult winter road conditions related to snow and ice.

The Canadian road weather model METRo (?) and the Norwegian model NORTRIP $_{\tau}$ (?, ?; ?, ?) predict slippery road conditions with a single shared ice/snow storage content. In Finland, RoadSurf (?) computes two distinct snow and ice reservoirs with a simple approach to the melting of ice and snow on the road. The melting energy is taken into account by using the excess energy to melt the ice and snow instead of warming the road when temperature is above the melting point. In Netherlands, the The Dutch road weather model takes into account freezing and melting energy (?). Chen et al (?) developed a complex formulation for road ice prediction. It computes an explicit one layer one-layer water/ice energy equation with complex heat exchanges between the road and the atmosphere. In France, a modified version of the hydrological model ISBA-ISBA hydrological model coupled with two multiple multi-layer multilayer snow models (CROCUS and ES) was built for road maintenance purposes $_{\tau}$ (?, ?; ?, ?). The snow model computes prognostic heat content, water content and density and has been validated on ES and CROCUS compute prognostic heat contents, and densities and have been validated at many alpine sites. CROCUS utilization. The use of CROCUS within ISBA-Route leads to accurate road conditions simulations simulations of road conditions on snow-covered roads (?).

In urban environments, road weather models fail to accurately model the road surface conditions since specific physical processes are needed to represent the town energetics (?). Simple building-averaged models are able to compute the radiative trapping, surface energy budgets and wind channeling (?). The modeling of urban winter conditions has been little studied in the urban climate community in comparison with summer (?). Lemonsu et al. (?) showed that snow-covered urban surfaces contribute to changes in sensible and latent heat fluxes. In an urban climate model (TEB), coupling the road with a simple one-layer snow model leads to improved fluxes in winter (?). Thus, the models SUEWS (?), TEB (?) or Lodz-SUEB (?) include a one-layer snow model to take the effect of snow into account during wintertime. CLMU (?) and JULES (?) go further and includes a multi-layer snow model over the road.

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This study attempts to bridge the a gap between urban climate models and road weather models used for road maintenance. On the one hand, as they focus on soil-atmosphere heat exchange, urban climate models do not include processes relevant to winter road maintenance. On the other hand, road weather models fail—with physics comparable to one-tile urban models (?) struggle to compute accurate road conditions in an urban environmenturban conditions. Some urban climate models bridge part of the gap by including snow (?) and ice accumulation in the road component (?). There are also road weather models that take into account sky-view factors and radiation trapping (?, ?; ?, ?). Thus, the The aim of this study is to improve the representation of winter processes in the TEB urban climate model to make it suitable for winter road operations. A new version of TEB (TEB-ES) has been developed. It models the challenging winter road artificial surfaces conditions associated with snow and ice. Our work presents a new ice storage term and an improved snow model with a multi-layer parametrization (ES) on of snow over the road surface.

Section 2 of the paper describes first the TEB initial version and (simply TEB) then the new processes added in the model new version of the model (TEB-ES). Section 3 presents the experimental set up set up to evaluate the model performance models

performances in winter conditions. In Section 4 presents and 5, we present the performance of TEB initial version and the new TEB model against measurement and the TEB-ES model against measurements at two sites, Col de Porte in France and a road weather station in southern Finland. Section 6 comments on the findings of this study and the limitation of TEB-ES to represent artificial winter conditions. Finally, we conclude this study section 7.

130 2 Methods

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2.1 Initial TEB model

The Town Energy Balance Model (TEB) (?) is embedded in the SURFEX software (SURface EXternalisée). The idea of this system was to build a modular system disconnected from an atmospheric model. Rather than being tied to a single atmospheric model, SURFEX can be launched autonomously or coupled to any atmospheric model and it provides to provide surface state variables. It-SURFEX consists of four sub-models submodels that describe different surface types on the globe with TEB appropriate to simulate urban environments. Extensively used to study the Urban Heat Island (UHI) in summer, TEB has been validated and incorporated in-into the Meso-NH model (?). Pigeon et al. (?) performed a winter evaluation of the model on-in the 2004-2005 Capitoul campaign and showed that the model accurately simulates surface temperature in cases without snow. Lemonsu et al. (?) evaluated the one-layer snow model coupled with the road and the roof snow during the Montreal campaign. TEB is an heat balance heat-balance model with a local canyon geometry that represents a simplified urban environment. It models two facing walls separated by a road, as first proposed by Oke et al. (?), which leads to a fast computation. The TEB model solves distinct heat equations for each surface (roof, wall, and road). The radiative trapping inside the canyon geometry leads to specific shortwave and longwave energy balance equations forced by the atmosphere for each surface. The net longwave radiation absorbed by each surface is computed between in interaction with each TEB component interaction. The direct solar flux received by the road or the walls is computed according to shadowing effects and road direction the direction of the road. The diffuse solar flux is processed using a sky-view factor and a geometric system for an infinite number of reflections. The following description of TEB is restricted to the road component its road component, schematically displayed Fig 1, as it is the focus of our study. Depending on the TEB configuration during simulations, the road component can represent simultaneously roads, sidewalks, and car park.

The ground is discretized discretised with layers of artificial ground representing the road structure and layers of natural soil beneath them. The temperature evolution across all layers is driven by a heat equation that computes the energy stored or emitted depending on the weather conditions. The road is assumed to be impermeable, so there is no water drainage within the vertical road layers. Snow The snow and rain intercepted by the soil is confined to the road surface. The snow cover defined as a fraction of the total road surface p_{nc} divides f_{sn} partitions the road surface. The fraction p_{nc} is computed with snow-covered fraction f_{sn} is computed from the total snowpack water equivalent W_{snow} (kg m⁻²) and the parameter $W_{snowmax}$ set to 1kg m⁻² (?) as:

$$p_{nc}f_{sn} = W_{snow}/(W_{snow} + W_{snowmax}) \tag{1}$$

Snow cover fraction p_{nc} The snow cover fraction f_{sn} that depends on the snow water content, is included in the heat balance heat-balance equation as:

$$160 \quad C_{R1} \frac{\partial T_{road}}{\partial t} = (1 - \underline{\underline{p_{nc}}} \underline{f_{sn}})) \frac{1}{d_{R1}} (S_R^* + L_R^* - H_R - LE_R - G_{R1,2}) + \underline{\underline{p_{nc}}} \underline{f_{sn}} \frac{1}{d_{R1}} (G_{Rsnow} - G_{R1,2})$$

$$(2)$$

With Where T_{road} is the road surface temperature driven by the snow-road conduction flux G_{Rsnow} , the conduction flux between the first and second road layers $G_{R1,2}$, net radiation fluxes S_R^* and L_R^* shown in Fig. 1, and sensible and latent heat fluxes H_R and LE_R . C_{R1} is the road surface heat capacity of the road surface depending on the dry surface heat capacity and of the mass contents amount, and d_{R1} the depth of the first road layer with values as described in Table 1.

According to Eq. (2) the TEB road surface energy budget is split divided according to the snow fraction. Indeed, snow cover insulates the road surface from the first canyon air layer and vice versa. The snow-covered road surface budget is only driven by the snow-road conduction term (second right-hand term in Eq. (2)). The energy budget on the snow-free fraction of the road is driven by the latent and sensible turbulent fluxes between the road and the interface canyon air layer, by the radiation absorbed by the road and by the heat conduction from the road sub-layers (first right-hand term Eq. (2)).

The rain is intercepted by the snow-free fraction of the road, and transferred into the available water reservoir $\overline{W_s}$ at the road surface in $\overline{W_s}$ (kg m⁻²). Its maximum capacity W_{smax} in the snow-free fraction is the maximum possible content set to 1 kg m⁻² (?). Thus, the evolution equation water for water volumetric water the water equivalent content W_s is:

$$\frac{\partial W_s}{\partial t} = R + R_{melt} - (\underline{1 - p_{nc}} \underline{1 - f_{sn}}) LE/L_v \qquad W_s \le W_{smax}$$
(3)

where R is the rain rate, R_{melt} represents the snow melting rate (kg m⁻² s⁻¹), L_v the latent heat of vaporization and LE is the latent heat flux between the road and the lower air layer of the urban canyon. If W_s reaches the maximum capacity available W_{smax} , the excess liquid water leaves the system as runoff.

The TEB road surface is coupled with a one-layer snow model (1-L) with an albedo parameterization, albedo and density parameters adjusted for urban environments (?). Temperature, water content, density, and albedo are solved prognostically and represent the evolution of the snow layer stateat any time. A simple formulation is used for the snow density with an exponential law. Simple formulations are used for snow density and snow albedo with exponential evolution laws to represent snow aging ageing. Liquid water melted from snow, R_{melt} , is transferred in into the available water reservoir, W_s , or it goes directly into runoff into runoff if W_s reaches W_{smax} .

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The road component of the TEB initial version described previously and represented in Fig. 1 is then modified to improve the representation of winter processes. The following sections describe the modification within TEB represented in Fig. 2 with a new ice content W_i and an improved representation of the snow mantle with the Explicit Snow (ES) model.

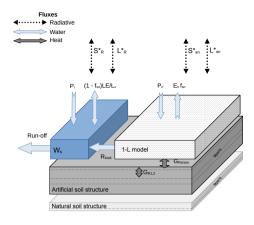


Figure 1. Schematic implementation of the road component in TEB initial version with the 1-L snow model in white, the water content W_s in blue, and their interaction with the road surface in grey. Heat, water and radiation effects are represented by arrows with radiative fluxes S_{sn}^* and L_{sn}^* (net shortwave and net longwave over snow, respectively).

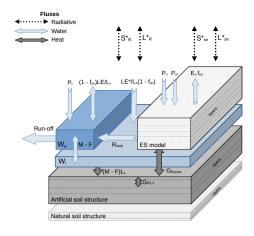


Figure 2. Schematic implementation of the new TEB model (TEB-ES) with the ice content W_i in light blue, the ES model in white and their interaction with the road surface in grey and the water content W_s in blue. Heat, water and radiation effects are represented by arrows with radiative fluxes S_{sn}^* and L_{sn}^* (net shortwave and net longwave over snow, respectively.)

2.2 **Ice content**Implementation of road ice

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To account for icy conditions on the road surface, we model road conditions, we introduce a prognostic evolution of the amount of ice on the road. It is described by the state variable W_i in $(kg m^{-2})$, which represents the liquid water equivalent ice content. W_i evolves by phase-induced changes. First, it interacts with the water content W_i , by melting and freezing. Second, it interacts with the atmosphere with by deposition as shown in Fig. 42.

Several hypotheses are made to best model reality and to To be in agreement with the TEB modeling choices. The fraction occupied by ice on the TEB road surface is set to 1 and does not depend on the snow-free fraction. Somodelling choices, the ice layer coexists with W_i energy, like the water reservoir To, To contrary to the liquid water content, ice content can grow without any limitation on the part that is not snow-covered. But, when the water content available for freezing is null, the ice content can no longer grow. Also, the ice layer can be snow covered. When the snow covers all the road surface To is not explicitly modelled. Their temperatures are considered indistinguishable from the road surface. The ice layer under the snow does not interact with the snow layer. As shown in Fig. 1, ice is transparent to the snow-road heat conduction flux To content To is intercepted by the snow-free road fraction and falls directly into the available water content To in the ice reservoir evolution equation is: Thus, the processes involved in the ice evolution, freezing, melting and sublimation impact directly the

road surface temperature as follows:

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$$\frac{\partial W_i}{\partial t} \frac{\partial T_{road}}{\partial t} = \underline{F - M} - \underbrace{\frac{1}{d_{R1}}}_{(1 - \underline{p_{nc}} f_{sn}) LE^* - (F - M) L_v)} \underbrace{\frac{LE^*}{L_s}}_{(4)}$$

With with F representing the freezing rate $(kg \cdot m^{-2} s^{-1})$, M the melting rate $\frac{in (kg \cdot (kg m^{-2} s^{-1}))}{p_{nc} f_{sn}}$, the snow fraction on road, L_s the sublimation heat constant and LE^* the solid-gas latent heat flux $\frac{in \cdot (W m^{-2})}{m^{-2}}$. The ice content W_i changes according to:

$$\frac{\partial W_i}{\partial t} = F - M - (1 - f_{sn}) \frac{LE^*}{L_s} \qquad W_i \ge 0$$
 (5)

Freezing of water is an exothermic reaction while melting is endothermic. This will affect the energy balance at the surface of the road as shown in Fig. 2. Ice on the roads also changes the exchange coefficient, based on the aerodynamical resistance, resulting in modified turbulent exchange between the road and the first air layer in the canyon. For this process, we assume that the ice is at the first road layer temperature and the aerodynamical resistance is the same as the one for water.

Several hypotheses arise from this parametrization. The fraction occupied by ice on the TEB road surface is set to 1. Thus, the ice layer can be partially or completely snow-covered depending on the snow fraction value f_{sn} . In addition, the ice content can grow without limitation as long as it is supplied by the freezing rate F of the available water content W_{s} , or by the ice deposition by sublimation on the snow-free fraction $(1-f_{sn})\frac{LE^*}{L_s}$. As the ice-layer energy is considered indistinguishable from the road surface, the snow-covered ice layer is transparent to the snow-road heat conduction flux G_{Rsnow} as shown in Fig. 2. Therefore, the ice-layer part that is snow-covered is insulated from the atmosphere and interacts only with the road surface by the melting process M.

The freezing rate, the melting rate and the solid-gas latent heat flux are defined by the following equations adapted equations for the ice evolution on the natural surface layer from Boone et al. (?) :- are adapted to impermeable artificial surfaces. The insulating effects of the vegetation are removed as well as the phase change coefficients for subgrid-scales effects. The adapted water mass rates for impermeable artificial surfaces are therefore defined as:

$$F = \frac{1}{\tau} \min_{\infty} (W_s, \frac{\max(O, T_f - T_{road})}{C_l L_f})$$
(6)

$$M = \frac{1}{\tau} \min_{\infty} \left(W_i, \frac{\max(O, T_{road} - T_f)}{C_l L_f} \right) \tag{7}$$

$$LE^* = \underline{\gamma_{ice}} \rho_a \frac{1}{R_a L_s} [Q_{sati}(T) - Q_a] \underline{1.inLE}^* \le \underline{0}$$
(8)

where the triple point temperature represented by with T_f , the freezing temperature, L_f is the latent heat of fusion of water, R_a the air aerodynamic resistance, ρ_a the air density, τ the characteristic timescale for phase change (3300/0.05 s used in this study), and C_l represents the ice heat capacity thermal inertia coefficient described in Boone et al. (?). The parameter τ is set to 25000s rather than the 3300s value used for natural soil in order to get realistic simulations for this study experiments. Ice sublimation is assumed to be negligible because its evolution is small compared to F or M. Thus, when the road surface reaches the saturation specific humidity with $Q_{sati}(T) \geq Q_a$, LE^* should be greater than 0, but the term γ_{ice} LE^* is set to 0.

Otherwise when $S_{\infty} LE^* \leq 0$, γ_{ice} is set to 1 and deposition as frost on the road can occur. The melting and freezing process couples the evolution of the ice and water contents. Thus, the water-reservoir evolution equation becomes:

$$\frac{\partial W_s}{\partial t} = R + R_{melt} - (1 - \underline{p_{nc}} \underline{f_{sn}}) LE/L_v - F + M \tag{9}$$

With F, M the freezing and melting rates for ice.

Freezing of water is an exothermic reaction while melting is endothermic. This will affect the energy balance at the surface of the road as shown in Fig. 1. Ice on the roads also changes the exchange coefficient, based on the aerodynamical resistance, resulting in modified turbulent exchange between the road and the first air layer in the canyon. For this process, we suppose that the ice is at the first road layer temperature and the aerodynamical resistance is the same as the one for water.

This process is added to the heat-balance equation as follows:

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$$\frac{\partial T_{road}}{\partial t} = \frac{1}{d_{r1}}((1 - p_{nc})LE^* - (F - M)L_v)$$

Schematic implementation of the ice content W_i in light blue, the ES model in white and their interaction with the road surface in grey and the water content W_s in blue. Heat, water and radiation effects are represented by arrows with radiative fluxes S_{sn}^* and L_{sn}^* respectively net shortwave and net longwave over snow.

2.3 Explicit Snow model coupling Improvement of the road snow processes with a multi-layer scheme

245 The snow mantle

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The thermal and liquid profiles of the snow mantle cannot be represented by averaged single-layer variables used by as in one-layer snow model schemes, but they require multi-layer models instead (?)(?). Cristea et al. (?) showed that numerous using several layers in snow models increase the performance improve the realism of the heat change and liquid transfer changes and liquid transfers between the snow mantle layers. Decharme and al. (?) have also shown that the same snow model with 5-layers rather 3-layers converting a snow model from 3 to 5 layers leads to a more accurate soil temperature evolution. Explicit Snow (ES) (?) is a multi-layer snow model that resolves explicitly explicitly resolves the heat-energy balance. The prognostic variables are snow density, heat content, thickness for each snow layer and albedo. To solve a proper snow thermal profile there should be Sun et al. (?) suggested that at least 3 layers in the model. Mass and heat are conserved to an high degree of accuracy in the models now layers are needed to represent a snow thermal profile.

In this work, ES simulates the snow mantle evolution on a road modeled modelled inside the local canyon geometry of TEB. The snow model is forced by the TEB variables. ES receives the computed shortwave radiation from the road sky-view factor and the trapped longwave radiation. It is also forced by the local atmospheric variables computed inside the canyon such as the specific humidity and air temperature. Finally, ES intercepts the snow precipitation and the liquid precipitation. Unlike the one-layer snow scheme (1-L), ES computes the impact of the liquid precipitation on the snow mantle. So, the The total liquid precipitation rate R (kg m⁻² s⁻¹), is split into a fraction that enters the snowpack with P_{rn} (kg m⁻² s⁻¹) and a fraction that is

intercepted by the water reservoir with P_r (kg m⁻² s⁻¹):

$$P_{rn} = p_{nc} f_{sn} R \tag{10}$$

$$P_r = R(1 - p_{nc}1 - f_{sn}) \tag{11}$$

Thus, the water-reservoir evolution equation receives P_r rather than the total liquid precipitation, R in Eq. (9). The snow fraction, p_{nc} , defined f_{sn} , defined by Eq. (1), is modified and is set to $p_{nc} = 1$ with respect to the TEB initial version is modified as follows:

$$f_{sn} = \min(1, D_s/0.01) \tag{12}$$

So $f_{sn} = 1$ when the total snow mantle depth D_s is higher than 0.01 mas:

$$p_{nc} = \min(1, D_s/0.01)$$

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The atmospheric variables in the TEB canyon are modified by this new snow scheme. For both the 1-L and ES options, the amount of radiation received by the snow-free fraction of the road is weighted by the snow cover. The snow/atmosphere interaction is modified by the ES scheme. The net heat flux, the sensible, latent, and radiative fluxes all depend on the local variables inside the snow mantle simulated by ES.

At the bottom of the snow mantle, ES is coupled with the impermeable road surface. Liquid water leaving ES is treated as in 1-L. It: it is transferred to the water reservoir and then taken into account in the road surface energy balance or, it leaves the system as runoff if W_s reaches W_{smax} . However, the heat conduction between the road surface and the lower snow layer (G_{Rsnow}) is not treated as in 1-L. ES scheme is implicitly coupled to the road surface following the procedure of Masson et al. (?), to improve stability. Heat conduction between 1-L and the road component heat equations is strictly explicit. It impacts the road surface energy balance as in Eq. (2).

The mass conservation equation for the total snowpack in TEB-ES is:

$$\frac{\partial W_{snow}}{\partial t} = P_n + P_{rn} - R_{melt} - E_n \tag{13}$$

 W_{snow} is the product of the average snowpack density and the total thickness. It corresponds to the total snowpack water equivalent (SWE) (kg m⁻²). P_{rn} is the liquid precipitation rate defined Eq. (10), P_n the snow-snowfall rate, R_{melt} (kg m⁻² s⁻¹) the melt rate and E_n (kg m⁻² s⁻¹) the total latent heat flux caused by evaporation and condensation.

Instead of Unlike in 1-L, each ES layer is characterized characterized by a liquid water content of the snow W_{li} . Index i refers to the layer. It is modeled modelled as a series of bucket-type reservoirs and the total layer liquid water content W_{li} stays < 10% of the total layer snow mantle mass represented by W_{limax} with:

$$\frac{\partial W_{li}}{\partial t} = R_{li-1} - R_{li} + \frac{F_{si}}{L_f} \tag{14}$$

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$$R_{l0} = P_{rn} - (1 - \chi_1)E_n \tag{15}$$

Where R_{li-1} and R_{li} is the water flow between the layers i-1 and i (kg m⁻² s⁻¹), F_{si} the phase change heat flux (W m⁻²) that represents the sum of two terms, the available energy for snow to melt and the available energy for the liquid water to freeze (considered as snow W_{snow}), R_{l0} the flux at the snow surface and χ_1 the fraction of the total mass of the surface layer which is frozen frozen surface layer defined as:

$$\chi_1 = 1 - \frac{W_{l1}}{W_{snow1}} \tag{16}$$

The snow-layer density prognostic variable ρ_{si} (kg m⁻³), changes because of few factors such as the weight of the overlying snow, the settling mainly due to fresh snowfall, the thermal metamorphism and the snow viscosity viscosity of the snow. Also, the fresh snowfall usually reduces the uppermost layer density and is defined as:

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$$\rho_{new} = a_{sn} + b_{sn}(T_a - T_f) + c_{sn}(V_a)^{1/2}$$
 (17)

Where T_a is the air temperature inside the canyon in Kelvin, V_a the wind speed, and coefficients $a_{sn} = 109 \text{ kg m}^{-3}$, $b_{sn} = 6 \text{ kg m}^{-3} \text{ K}^{-1}$ and $c_{sn} = 26 \text{ kg}$. Melting, infiltration of rainwater and retention of snow melt also affect the snow layer density as described in Boone et al.? (?).

The snow mantle is slightly transparent to the solar radiation flux. The snow mantle 's heat balance heat-balance equation is modified at each layer by this positive heat flux. The solar transmission heat flux is an a negative exponential of the snow depth and the extinction coefficient for shortwave radiation products. This flux is weighted by the snow surface albedo.

In ES, the snow surface albedo process is adjusted for natural environments and computed as in Decharme et al. (?). The impact of human activity, such as pollution sources, on the whiteness of snow is not considered. Thus, the albedo equation and parameters used in 1-L from Lemonsu et al. ?-(?) are used in replacement. It impacts directly impacts the solar radiation transmission heat flux.

3 Experimental set up and assessment

3.1 Model configuration

This study compares the TEB model initial version released in SURFEX v9.0 with the modified version called TEB-ES in-with the road component processes shown in Fig. 1 and in Fig. 2 respectively. The TEB-ES version used in this study is published in the repository (?), that includes the focussing on the processes related to snow and ice related processes described above. Both models are initialized with the same shared configuration configured in a similar way in order to evaluate the impact of the new processes at the locations selected for this study: Col de Porte in France and Hajala in Finland. Two benchmark road weather models are set up and compared to established and compared with TEB and TEB-ES performance. First, the heat-balance

model ISBA-Route/CROCUS described in Bouilloud and Martin (?)named ISBA-Route/CROCUS, which is in operation in operation at the French national meteorological officewith a downgraded version, is used in comparison at Col de Porte location. The model is not used at the Hajala location. Indeed, it lacks processes to model the impacts of human activities, such as salting or traffic heating, as in TEB and TEB-ES. CROCUS is a more complex snow model than ES and is coupled to the surface of ISBA-Route. CROCUS and ES share many similarities, but CROCUS can explicitly calculate snow metamorphism, including grain size and shape evolution, which impact the mechanical properties and albedo of the snow mantle (?). Secondly, at each experimental location, a simple statistical model as described in Kršmanc et al. (?) that is built with a multiple linear regression method (MLR) method, is used to predict road surface temperature. Simple empirical models are valuable for assessing the need to construct complex physical models for predicting surface variables (?). The best predictive variables for the MLR are found by using a stepwise regression procedure that minimizes the root mean squared error. They are chosen among the as explained in Appendix 1, using the same available forcing variables that are used as input for the heat-balance modeldescribed next section. Following Kršmanc et al. (?), lags of the variables are introduced in the model to help capture the time-varying nature of the physical variables. Thus, input variables have a lag of 3 hours, 2 hours and no lag before the prediction time.

TEB is designed for urban areas, but it needs to be adapted for validation sites located in open areas where the pavement is constructed without adjacent structures. The local canyon geometry configuration of the model cannot be completely removed. So, we flatten the canyon geometry to the limit. The canyon aspect ratio was set to 0.0001 which causes the sky-view factor of the road to be close to 1. This nullifies the radiative trapping of the canyon. The building fraction is set to 0.0001 to limit interactions between the air inside the canyon and the TEB building component. With these settings, TEB is considered to simulate a road with in open surroundings. The surface boundary layer options option is activated and computes explicit atmospheric variables inside the urban canyon (?) as well as the. The explicit calculation of the longwave exchanges is also activated. We set the pavement structure physical parameters described physical parameters of the pavement structure described in Table 1 to be in agreement with Bouilloud and Martin (?) for a French highway. The natural soil under the artificial structure is initialized initialized in TEB by the dry soil thermalthermal conductivity of the dry soil, whereas for TEB-ES it is initialized with the moist soil thermal conductivity (?). For the second thermal conductivity of the moist soil (?). Indeed, at our experiment site, water infiltrates beneath the road from surrounding natural soil. For the Finnish experiment in this paperarticle, the TEB-Hydro component is enabled to simulate water wear-off in order to simulate the water drainage on the roads (?), and W_{snowmax} W_{snowmax} is set to 0.6 kg m⁻² rather than 1 kg m⁻² to take into account the because of the different road properties.

A straightforward parameterisation of snow removal operations is implemented in TEB, TEB-ES, and ISBA-Route/CROCUS. Within these models, snow depth and ice content are set to zero when an operator clears the snow. This parameterisation is adapted for the snow removal procedures carried out at the experimental sites presented next section. In the Col de Porte experiment, snow and ice are reset to 0 on the known date of snow removal by a manual operator. In the Hajala experiment, the exact dates of winter maintenance activities are unknown. So, snow and ice are reset to 0 at 6h UTC in the models.

Soil material	Concrete	Natural soil
Number of layers	9	3
Discretization [m]	[0.001, 0.01, 0.05, 0.10, 0.15, 0.20, 0.30, 0.60, 1.00]	[1.50 ,2.00 ,3.00]
Soil thermal conductivity $[W.m^{-2}]$	2	1.5 (TEB-ES), 0.5 (TEB)
Dry specific capacity $[J.m^{-2}]$	1840000	1940000
Surface albedo	0.1	n/a
Surface rugosity [m]	0.005	n/a
Surface emissivity	1.	n/a

Table 1. Road configuration and parameters in TEB and the modified model TEB-ES

3.2 Experiments

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The models are mainly forced by the on-site measurements observations at the experiment set-up location. They are first assessed at the Col de Porte Météo France measurement site. Located measurement site of Météo France. It is located at an altitude of 1325m-1325 m in the Chartreuse mountain range in the Alps (45.30° N, 5.77° E), this site. This site is in a grassy meadow surrounded by a coniferous forest, it is covered by snow several months a year. In operation since 1959, the Col de Porte Météo-France site is equipped with standard meteorological and snow mantle sensors (?). It is a European reference for the study of snow-covered surfaces thanks to the meteorological conditions and its collection of sensors. Thus, data from this site has have been exploited to validate many snow models (?, ?; ?, ?; ?, ?) and even used for large snow model intercomparison projects (?). During 3 winters (1997/98-1999/2000), the Col de Porte hosted a large experiment to study the snow-road interface from within the GELCRO project (?). Six experimental artificial pavements. An artificial pavement (2m x 3m) equivalent to a French highway shown in Fig. 3 were was installed at the site (?). Road surface temperatureand The road surface temperature, with a probe inside the artificial structure, and the snow depth were monitoredon each artificial pavement. The snow cover was frequently cleared by an operator throughout the entire experiment. This experiment was used to evaluate the model operating in France (?) ISBA-Route/CROCUS model (?) used for road maintenance purposes, which is our heat-balance road weather model benchmark benchmark model for this study.

TEB-ES is assessed and compared with TEB, ISBA-Route/CROCUS and MLR. The heat-balance models are forced hourly by the local atmospheric measurements at the static nearby meteorological station. The 6min 6 min precipitation measurements are aggregated every full hour to give a precipitation intensity. The type of precipitation is divided by assuming (mm h⁻¹). The precipitation phase is selected by assuming that it is rainfall when the air temperature is < 1°C and snowfall when the air temperature of 1°C is the average temperature of separation of the precipitation phase in the Alps. But this assumption can often fail (?). For the other atmospheric atmospheric measurements, the value closest to the whole hour is considered. The hourly surface observations

provides provide validation data for the models. The models are evaluated with the experimental artificial pavement equivalent to a French highway surface observations only. Finally, snow depth and ice content are reset to 0 in the models after operator snow removal surface observations of the artificial pavement.

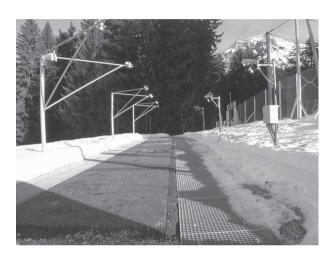
Next, a site with recurring snowy and ice road conditions outside of controlled experimental conditions was selected to assess the modelsTEB, TEB-ES and MLR. In southern Finland, these kinds of conditions are normal in winter and the temperature crosses zero degrees multiple times during the winter season, which makes the making surface condition forecasting a challenge. Fintraffic has installed numerous road weather stations to monitor atmospheric variables (wind speed, air temperature, humidity and precipitation), road surface temperature and road conditions. The stations road surface temperature sensors are manufactured by Vaisala and they usually measure road surface temperature with asphalt embedded sensors. Many stations also have optical instruments that measure water, ice, and snow layer thickness. Anthropic effects such as traffic and winter maintenance directly influence the physical variables. The surface temperature and layer thicknesses can vary greatly depending on which part of the road they are measured. For instance, snow compaction by the traffic can drop the snow depth and lead to measurements errors. In addition, the optical sensor might only see the top of the snow or ice layerand is unable to measure the actual thickness. For these reasons, the quantities measured by optical sensor should not be treated as absolute truths but as approximate measurements. They can still be used in the qualitative validation of the models the thickness of the water, ice, and snow layer. Among several stations with the most sensors, the Salo Hajala road weather station (60.435°N, 22.969°E) shown in Fig. 4 has been arbitrarily selected. It is ealled from now on From now on it is called just "Hajala" for sake of simplicity.

To force and validate the model, we used data from a study conducted by Karsisto and Loven? The used data were observed variables at Hajala road weather station. Wind and from measurements provided by the Finnish Meteorological Institute (?). Observed wind speed, air temperature, humidity and precipitation were from the Hajala road weather station are used as atmospheric forcing in the model and processed in the same way as the Col de Porte forcing. Since there is no radiation measurement at the Hajala road weather stations, shortwave and longwave radiation were extracted from ERA5 reanalysis at the closest grid point (?). Surface measurements including road surface temperature, water/ice contents, Snow Water Equivalent (W_{snow}) and road conditions were and snow water equivalent (SWE) are then used to validate the models. The hourly atmospheric forcing for the model consisted of a mix of observation data and ERA5 reanalysis. Shortwave and longwave radiation were extracted from ERA5 at the closest grid point (?). The studied period was from October 2017 to May studied period was from October 2017 to May 2018.

In many cases, optical instruments are considered unreliable for detecting road conditions with precision subject to anthropic effects and winter maintenance road operations. Indeed, optical sensors always failed to distinguish between ice and snow content on the road surface. In the Hajala experiment, on the 706 snow occurrences and 743 ice occurrences measured, the sensors recorded 706 occurrences of both ice and snow at the same time, and the other 37 hourly occurrences for ice are at the beginning or at the end of a snow event. Anthropic effects such as traffic and winter maintenance directly influence the physical variables. In addition, the optical sensor might only see the top of the snow or ice layer and is unable to measure the actual thickness. For these reasons, the ice and snow mass contents measured by the optical sensor should not be used to validate

the models quantitatively but qualitatively as occurrences. They are compared in the Hajala experiment, with the snow and ice
410 output variables from the models transformed as occurrences.

Statistical scores are calculated hourly at the Hajala site for the whole simulation, similarly to the Col de Porte experiment. The scores in Table 3, Table 4, Table 6, Table 7 and Table 8 are calculated from the confusion matrices that report the number of true positives (TP), false negatives (FN), false positives (FP), true negatives (TN). They are calculated as follow: detection rate = TP/(TP + PN), missed event rate = FN/(TP + PN), false positive rate = FP/(FP + TN) and false discovery rate = FP/(TP + FP). These metrics help to evaluate the models performances for important thresholds, in particular for decision making in the context of road weather forecasts.



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Figure 3. Col de porte experimental artificial soil during the GELCRO campaign, extracted from (?)



Figure 4. Salo Hajala road weather station in Finland, ©Google Steet-Street View 2024

4 Evaluation at the Col de Porte site

First, we compare the performance of TEB, TEB-ES, and the benchmarks at the Col de Porte meteorological site during the GELCRO campaign. They are forced by the in-situ in situ measurements and set up to compute the physical variables from 21st October 1998 6 UTC to 14th May 1999 4 UTC in a continuous simulation over the whole time span. Snow height In this reference experience, the new winter processes implemented are evaluated to see whether they have a positive impact on the model performance and the physics consistency. Snow depth and road surface temperature are studied on this reference experience for snow-road interface modeling simulations are compared with the observations.

4.1 November 7th to 18th period

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The time range extracted from the simulation from November 7th to November 18th in Fig. 5 shows typical snow conditions at the measurement site with 5 snow events. During this period, the road was cleared by hand three times three times by hand. Thus, in the models, on November 13rd at 12h UTC, on November 16th at 11h UTC and on November 17th at 17h UTC, the snow heights and the ice contents are reset to 0. TEB and TEB-ES behaviors are tested and described with 1-L and ES enabled, respectively.

On November 7th, 8th and 9th, a synoptic high pressure centered over Western centred over western Europe brought calm weather. Conditions were dry with positive air temperature and clear skies. The observed daily evolution of the road surface temperature is accurately simulated by TEB and TEB-ES. Both simulations are nearly identical, except that TEB-ES has a reduced cold bias during the evening and nighttimenight. Here, the road surface temperature is driven by road-atmosphere interaction and the pavement conduction. The soil-atmosphere interaction in the absence of ice or snow has not been changed for the new TEB version in the new version of TEB. But the moisture conductivity in the natural soil under the pavement added in TEB-ES as seen in Table 1 leads to improved pavement heat restitution and reduces the cold bias by 0.5 °C.

Several weather perturbations occurred during the following days. The first low pressure system reached the station on November 10th. Rain fell in the afternoon, followed by snow in the evening. TEB-ES simulates lower snow depths than TEB (around 2.7 cm lower for the episode). ES simulates more accurately the heat transfer between the positive road temperature and the fresh snow layers. ES melts almost all snowpack and is closer to the observations as shown in Fig. 5. Therefore, the TEB-ES road temperature follows the observations that are driven by the report a negative air temperature; while whereas the TEB road surface temperature is insulated from the atmosphere by the snow mantle. Then, its road surface heat change is driven by pavement conduction and snow-road heat transfer. TEB-ES road surface simulation is better than TEB on November 11st and 12nd.

A new shallow low pressure system formed off the coast of Ireland during the November 11st, and narrowed rapidly with weak activity over France on November the 12th. It brought a small amount of snowfall with a snow depth increase during After small observed snowfall in the afternoon of November the 12th. The snow mantle height, the snow depth evolution is well computed by TEB-ES with less than a 2 cm difference with the observations. TEB adds fresh snow to the previous snow mantle on the road and leads to a snow cover 6 cm higher than the measured value.

From the 14th to the 17th of November, a low pressure system persisted over Eastern Europe with several rainfall and snowfall events before a strong ridge brought back high pressure and clear skies. At the beginning of this event, the precipitation forcing is wrong; it was rain rather than snowfall that affected the location. This explains the excessive snow cover in both models. The following snow event is well modeled modelled by both models. ES simulates the fresh snow accumulation more accurately due to the multi-layer parameterization. The mixed composition density of fresh and old snow layer is better represented than in 1-L as described in section 2.3. The temporal evolution of the density of the old snow layer is also better represented with ES (?). Also, (?). In addition, the road surface temperature is better modeled modelled by ES with a Mean mean absolute error (MAE) of 0.3 °Cand for 1.4 °C during this event. This was a typical

isothermal event with a snow-pavement interface layer at constant freezing temperature. This effect is poorly represented by the 1-L snow modelthat underestimates snow-road heat transfer.

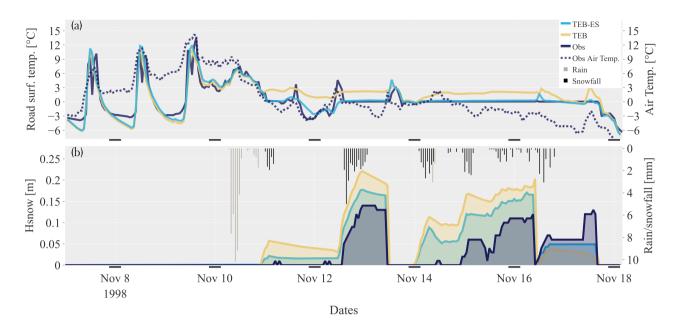


Figure 5. Comparison between the models and the observations at the Col de Porte location. Road surface temperature (modeled modelled and observed) and observed air temperature (a), snow height (modeled modelled and observed) and observed rain/snowfall with a reversed y-axis (b)

460 4.2 Statistical results

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Twenty-three snow events occurred during the date range from the October 21st, 1998 at 6 UTC to May 14, 1999 at 4 UTC at the Col de Porte, which is a large enough sample to show statistical differences between the models. Scores—The scores displayed in Table 2 show notable differences in the performance of the road surface temperature and snow height simulations. The heat balance heat-balance models outperform the statistical benchmark. In Table 2, the absence of bias in the TEB road surface temperatures is explained by several seasonal biased scores that compensates compensate. Indeed, Fig. 6 shows significant seasonal temperature differences with the observations in TEB. For TEB-ES and ISBA-Route/CROCUS simulations, the seasonal temperature differences with the observations are much lower. TEB-ES's and ISBA-Route's road surface temperatures are more consistent with the observations than the TEB simulation as shown in Table 2 and Fig. 6. Snow height is better simulated by TEB-ES than TEB in terms of RMSE, MAE and R² shown in Table 2. Multi-layer snow model coupling greatly improves the snow height and the road surface temperature performance. Because the TEB model is more adapted for man-made structure simulations, validated and calibrated at many locations, it is more accurate than ISBA-Route/CROCUS.

Scores	ores Snow height [m])		Road surface temp. [°C]				
Models	TEB	TEB-ES	ISBA-Route/CROCUS	TEB TEB-ES ISBA-Route/CR		ISBA-Route/CROCUS	MLR
RMSE	0.19	0.13	0.14	2.82	2.05 <u>2.27</u>	2.53	3.64
MAE	0.12	0.08	0.09	2.10	1.33 _1.39	1.40	2.45
\mathbb{R}^2	0.54	0.84	0.80	0.82	0.89_0.86	0.83	0.57
Bias	-0.04	-0.02	-0.02	0.02	-0.62 -0.75	-0.92	0.00

Table 2. Scores for TEB, TEB-ES and the benchmarks (ISBA-Route/CROCUS, MLR) at the Col de Porte location during winter 1998-1999

In addition, it is essential to we evaluate the ability of the models to capture the occurrence of significant events that could compromise road safety. Table 3 and Table 4 evaluate respectively the capacity the ability of the models to predict potential dangerous conditions and snow road condition occurrences. Similar detection rates and missed event rates respectively. Similar confusion matrices are found for the models for snow height > 0.5 cm. However, TEB-ES outperforms TEB and ISBA-Route/CROCUS reliability of positive detection with lower false detection and false alarm rates. Surprisingly, TEB outperforms slightly ISBA-Route/CROCUS despite the simple 1-L snow model embedded. During conducive conditions for slippery roads, ISBA-Route/CROCUS has a better detection rate as shown in Table 3 but is less reliable to predict a true negative among all the simulated snow occurrences as shown by the similar rates Table 4.

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Larger differences are observed in road surface temperature simulations between the heat-balance models are observed in snow-covered isothermal situations. This is is-particularly visible during spring 1999 shown in Fig. 6, panel (c), with snow-covered isothermal situations only. In those these situations, the TEB road surface temperature is strongly biased, while TEB-ES and ISBA-Route/CROCUS show good performance. ISBA-Route/CROCUS has a slightly better performance than TEB-ES on in these situations, due to the complexity of the CROCUS snow model. However, during one particular event in early spring, road surface temperature was very poorly simulated by ISBA-Route/CROCUS. These outliers are not shown in Fig. 6. The TEB-ES detection rate is much higher than TEB for road surface temperature < 0.5 °C as shown in Table 3 which can be attributed to snow-covered isothermal situations.

Models	Detection rate %	Missed event rate %	False detection positive rate %	False alarm discovery rate %
TEB	74	26	4	2
TEB-ES	90. 92	9 .8.	5. 7	2. 3 _~
ISBA-Route/CROCUS	97	3	21	9
MLR	70	30	6	4

Table 3. Performance of TEB, TEB-ES and the benchmarks (ISBA-Route/CROCUS, MLR) surface temperature occurrence, below 0.5 C, for at 1 hour time step, at the Col de Porte location during the winter 1998-1999

Models	Detection rate %	Missed event rate %	False detection positive rate %	False alarm discovery rate %
TEB	99	1	18	26
TEB-ES	98	2	16 . <u>17</u>	24 -25
ISBA-Route/CROCUS	98	2	19	27

Table 4. Performance of TEB, TEB-ES and ISBA-Route/CROCUS show snow depth occurrence greater than 0.5 cm, for at 1 hour time step, at the Col de Porte location during the winter 1998-1999

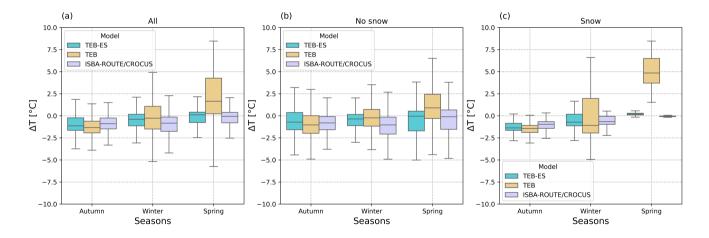


Figure 6. Seasonal road surface temperature differences comparison with the observations between TEB, TEB-ES and the benchmark ISBA-Route/CROCUS. Figures with road conditions partition of the situations: all cases (a), no-snow observed (b), non zero snow observed (c). The boxes extend from the first quartile (Q_1) to the third quartile (Q_3) , with whiskers up to the farthest point lying within $1.5\times$ the interquartile range (Q_3-Q_1) .

5 Evaluation at the Hajala site

In this section, we take advantage of the detailed observation observations at the Finnish Salo Hajala road weather station to further validate the physics of the model, in particular, the different mass contents in terms of mass content at the road surface. The model is set up to calculate the evolution of the TEB variables for about 6 months in a continuous run, from October 23rd, 2017 at 15h UTC to May the 1st, 2018 at 19h UTC.

5.1 January 17th to 26th 2018 period

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From January 17th, 2018 to January 23rd, 2018, the Salo Hajala site was affected by a cold air mass with road temperatures below -3°C (Fig. 7). Three small light synoptic scale snow events impacted traffic conditions by causing snow cover on the roads. Then, on January 24th the weather regime changed with a low pressure system which brought snowfall, then warmer air with rain.

Small persistent snowfalls on January 17th and January 18th in the morning were induced by a small low surface pressure associated with a weakly active quasi-stationary frontare captured simultaneously as ice and snow measurements on the road until January 19th in the morning. This can be attributed to the solid content detection issue with the sensor, that usually does not discriminate between ice or snow. TEB and TEB-ES simulates the appearance of a thin snow cover several simulate the snow cover 8 hours before the sensor measurement, as shown in Fig. 7. This difference The models fail to match the time span of the event. It could be explained by the high traffic intensity during these hours. The blowing of snow by the traffic strong morning traffic commuting pattern that blew the snow away, removed the thin flake layer on the roadand delayed the buildup, and then delayed the accumulation of the snow cover (?). Then, the sensors measured ice and snow simultaneously as the road began to be affected by solid precipitation. This is a common spurious ice detection issue with the sensor. There could be not ice, as the conditions were snowy and there was no liquid water to freeze on the road surface. Both models correctly capture the snow layer evolution. The snow cover is removed by hand at 6 h UTC in the models. The modeled However, the modelled road surface temperatures are consistent with the observed increasing trend during the afternoon of January 18 and show good performance.

On January 19th and 20th, during the nights, the sensor detected liquid water on the road surface. The sky was clear, and the raingauge did not capture any rain. Thus, it is may be a spurious water detection by the sensor. TEB-ES simulates a possible small hoarfrost event that could have been eaptured by this sensor classified as water.

Both snow simulations match the models snow mantle evolutions match the beginning of the observed SWE in late evening on January 21st. But in January 22nd in the afternoon, the observed contents dropped to 0 mm whereas the simulated SWE's increase. Snow ploughing weekday operations occurred during the day until noon. Afterward, the TEB and TEB-ES simulations snow mantle evolutions are consistent with the observed snowfalland the surface state.

Then, in the night between January 23rd and January 24th, the upstream warm front of the weakening low pressure system brought moderate snowfall. The snow water equivalent evolution by the modelsmatches the observed snowfall. In the morning, snow ploughing and salting removed the snow deposition on the road which we modeled by a simple manual removal of snow

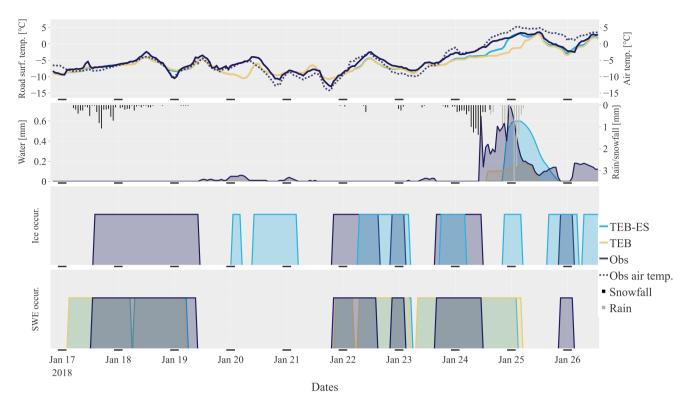
cover at 6 h UTC. These operations are largely responsible for the observed water increase. In the afternoon of January 24th, the air temperature rises due to . For this episode, the simulated snow occurrence is precise enough to return an accurate road surface temperature for both models.

In this 9-day period, the warmer air mass brought in by the low pressure system. TEB-ES snow melts accordingly whereas TEB the snowpack melts slowly before it is set to zero by the manual removal operation parameterization road surface temperature is well simulated for both models with similar results. But TEB-ES significantly outperforms TEB on January 25th moderate snowfall followed by a rainfall episode. It shows an improved modelling of the snow mantle variables with ES for positive air temperatures. In both models, the simulated SWE occurrence is consistent with the observed snowfall but fails to match the observed mass content on the road. The snow removal parameterisation at 6 h UTC. Finally, the road was wetted by rainfall in the warmer air mass brought by the low pressure system. In the night of January 25th, in clear sky conditions, TEB-ES model correctly captures the freezing of water on the road.

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the models appears to be an oversimplification of road maintenance operations such as salting or snow ploughings.



Comparison between TEB, TEB-ES and the observations at the road surface on a road weather station at Hajala. Road surface temperature (modeled and observed) (a), snow water equivalent (modeled and observed) and observed rain/snowfall (b), ice on road (modeled and observed) (c), water on road (modeled and observed) (d)

Figure 7. Comparison between TEB, TEB-ES and the observations at the road surface on a road weather station at Hajala. Road surface temperature (modelled and observed) (a), snow water equivalent (modelled and observed) and observed rain/snowfall (b), ice on road (modelled and observed) (c), water on road (modelled and observed) (d)

5.2 April 1st to 7th period

In early April, a low pressure system traveled fast from the Baltic States to Hajala and hit the station with the first snowfall on

535 5.2 April 1st to 7th period

Several cold conditions struck Hajala during this 6 days of this mid-spring period (Fig. 8). The road surface temperature is well simulated by both models, especially the diurnal cycle of surface temperature in clear sky conditions on April 3rd, 4th, and 7th. For the road conditions, the models have varying degrees of success in representing the different mass contents evolutions. In snow and ice-free conditions, the TEB and TEB-ES simulated water content evolutions are similar.

On April 2nd, at 2 h UTC. Uninterrupted, uninterrupted moderate snowfall occurred at the station until 14 h UTC from the warm front and then the occlusion of the low pressure under a low-pressure system. This standard snow event had probably been anticipated with brine injection during the night, since the water content observed increases as shown in Fig.8. Since the salting effect is not modeledmodelled, both models poorly simulated the snow water equivalent content occurrences captured by the Vaisala sensorsas shown in Fig.8. However, the overall timing of snow cover is well captured by the models because they use observed precipitations as input.

The road surface temperature is well simulated by both model. The models capture the diurnal cycle of surface temperature in clear sky conditions. They also capture the positive surface temperature with rainy conditions and an overcast sky on April 4th. In snow and ice-free conditions, the simulated water content evolutions are similar. When the road was salted on the morning of April 2th, the more sophisticated modeling in TEB-ES fails to improve the SWE since the salting process is missing to represent these impacts. On April 6that noon, our precipitation type procedure diagnoses small snowfall instead of rainfall, leading to a misdiagnoses twice the precipitation types that fall into the rain gauge. It leads to two false positive light snow event. The sensors detected a small amount of water on the road. Then, in the early afternoon, the roadsurface temperature simulated by both models are positive with negative air temperature and snowfall observed. TEB-ES melts the simulated snowpack quickly, whereas the SWE simulated by TEB remains stable until the parametrized snow removal at 6 h UTC, events on April 6th at noon, and in the afternoon instead of wet road.

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Clear sky conditions during the same night were cold enough to freeze the water on the road. Later, the road was salted, which melted the ice. In TEB-ESwater content from, the water produced by the snow mantle melting is frozen and the model accurately reproduces the observed content on the road which was presumably ice.

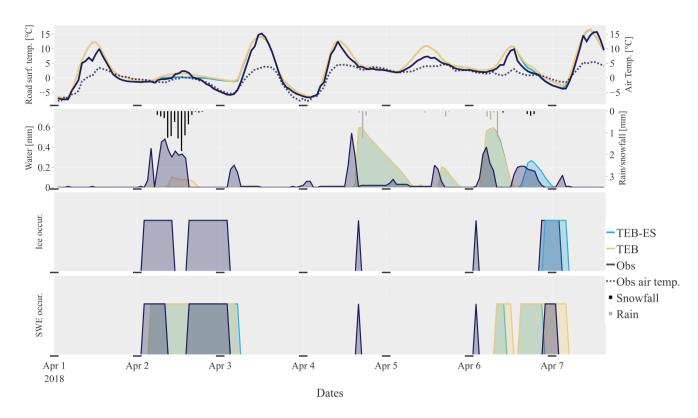


Figure 8. Comparison between TEB, TEB-ES and the observations at the road surface on a road weather station at Hajala. Road surface temperature (modeled modelled and observed) (a), snow water equivalent (modeled modelled and observed) and observed rain/snowfall (b), ice on road (modeled modelled and observed) (c), water on road (modeled modelled and observed) (d)

5.3 Statistical results

Statistical scores are calculated hourly at the Hajala site for the whole simulation. The approach is similar to that used for the Col de Porte site. Overall TEB and TEB-ES road surface temperature performance shown in Table 5 are almost similar. TEB-ES is slightly better on RMSE and MAE but does not improve the R²-R² and bias. The For TEB-ES, the slight increase in performance is due to better simulation of the snow cover in the snow-covered situation conditions, as shown in Fig. 9. On the panel (c), TEB-ES road surface temperature differences are less spread than TEB and almost centered are almost centred over 0 °C. Figure 9 reveals, revealing a seasonal performance trend in snow-free situations with all models. The overall temperature differences in spring is higher than in autumn and in winter. In addition, snow-covered situations with the heat-balance models. It is consistent with the behaviour of the Col de Porte experiment typical of snow isothermal events. However, the statistical benchmark significantly outperforms the heat-balance models with better RMSE, MAE, R² and bias. It is able to retrieve an accurate road surface temperature for each season as shown Fig. 9. Even if the MLR predictions are mainly

570 unbiased, Fig. 9 shows that in snow-covered conditions the road surface temperature difference spread is lower in winter in both models simulated is slightly biased.

The snow water equivalent scores are difficult to interpret due to salting. The road lanes are frequently treated with salt and the snow is removed by winter service vehicles. Thus, the observed occurrences of snow cover are much lower than modeled As previously seen, the snow removal parameterisation at 6 h UTC in both TEB and TEB-ES is not complex enough to represent road maintenance operations such as salting or snow ploughings which greatly influences the mass content evolution. For instance, freezing is simulated more often than observed with 743 observed ice occurrences and 1368 modelled ice occurrences. As said in the sensor descriptions, the optical sensors are not able to distinguish between snow-covered or ice-covered road conditions on busy lanes. This explains the high missed event rate shown in Table 8. The scores are shown in Table 8 and Table 7, but only a raw analysis can be extracted from these values. Therefore, the mass contents performances can not be compared between the models on this road, because of human activities.

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However, the models simulations are consistent with physics, and tend to accurately represent the road conditions without human activities. The physical consistency of the models with the observed precipitation leads to an high detection rates for SWE but a snow cover that is lower in observations than modelled (706 hours observed, 1394 modeled modelled by TEB and 1417 modeled modelled by TEB-ES) and lead which leads to a > 60 % false alarm-discovery rate on both models as shown Table 7. Also In addition, the snow occurrence detection by the model shows < 25 % false detection rate and > 70 % detection rate.

Several stacked effects explain the low performance for the ice content variable as shown in Table 8. The optical ice sensor detects ice whenever snow is detected. The measurement recorded 706 occurrences of both ice and snow at the same time. In total, 706 snow occurrences and 743 ice occurrences are measured. Thus, the optical sensor is not able to distinguish between snow-covered or ice-covered road conditions on busy lanes. The other 37 hourly occurrences for ice are at the beginning or at the end of a snow event. This explains the high missed event rate shown in Table 8. In Hajala experiment, around 28 % of the events modeled as ice are observed as water events in non freezing conditions (Road surface temperature measured is > 0C) as illustrated in Fig. ??. Around 96 % of these 28 % events were actually observed as ice-free wet road. This explains the There are a high number of false alarms shown Table 8, which are explained by the cold bias in the model. Freezing occurs more often than observed with 743 ice occurrences observed and 1368 iceoccurrences modeled events without snow and ice.

	TEB	TEB-ES	<u>MLR</u>
RMSE	2.51 -1.58	2.27 _1.51	1.19
MAE	1.19	1.12 _1.13	0.83
\mathbb{R}^2	0.95	0.95	0.97
Bias	-0.31	-0.31 -0.32	<u>Q</u>

Table 5. Road surface temperature scores for the Hajala site during winter 2017-2018.

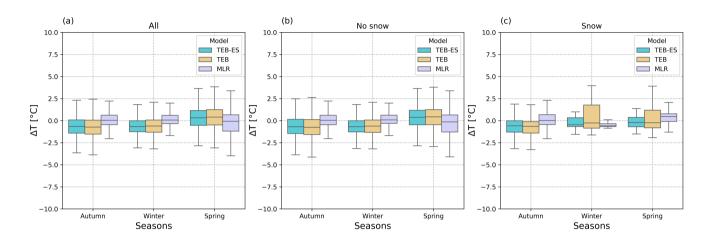


Figure 9. Seasonal road surface temperature differences comparison with the observations between TEBand, TEB-ES and MLR. Figures with road conditions partition of the situations: all cases (a), no-snow observed (b), non zero snow observed (c). The boxes extend from the first quartile (Q_1) to the third quartile (Q_3) , with whiskers with whiskers up to the farthest point lying within $1.5\times$ the interquartile range (Q_3-Q_1) and without outliers.

Hourly ice depositions in mm modeled on the road by TEB-ES (with ice depth occurrence greater than 0.001mm); function of the observed and simulated surface temperatures differences $\Delta T(obs-mod)$ when $T_{obs}>0C$ and $T_{mod}<0C$. The colors shares the points by: observed water content on the road $(ObsW_s>0)$ or dry road $(ObsW_s=0)$.

6 Discussion

In this study, the performance evaluation of TEB and its modified version TEB-ES has been carried out in open surrounding areas to investigate whether this urban climate model can accurately reproduce artificial surface conditions predicted by road weather models in winter. We isolated the winter surface processes from other physical interactions by limiting the current experiments to open environments. The surroundings are limited to few trees and no buildings. Although urban climate and road weather models mainly focus on simulating summer and winter conditions, respectively, we had attempted to bridge the

Models	Detection rate %	Missed event rate %	False detection positive rate %	False alarm discovery rate %
TEB	95	5	18	12
TEB-ES	97	3	18	12
<u>MLR</u> €	93	7_	6_	4

Table 6. Performance of TEBand, TEB-ES and MLR surface temperature occurrence below 0.5 °C, for 1 hour time step, at the Hajala site during the Winter 2017-2018

Models	Detection rate %	Missed event rate %	False detection positive rate %	False alarm discovery rate %
TEB	71	29	22	63
TEB-ES	72	28	20	60

Table 7. Performance of TEB and TEB-ES SWE occurrence greater than 0.01 mm, for 1 hour time step, at the Hajala site during the winter 2017-2018

gap between both model types by improving the TEB winter conditions modelling by including new processes from road weather models.

TEB and TEB-ES simulations have demonstrated good performance on the demonstrated physical consistency with the reality. The new developments in TEB-ES improve the performance of the road surface simulation in winter conditions as seen on the 1998-1999 Col de Porte and Hajala experiments, and the model physics are consistent with reality. The urban elimate model TEB is well-suited winter experiment, while beating the ISBA-Route/CROCUS road weather model in operation at Météo-France and even outperforming the in-sample statistical benchmark. TEB-ES appears well suited to simulate the surface response to atmospheric variables in man-made structures. When using snow models with similar characteristics, TEB outperforms ISBA-Route in winter conditions. The heat balance models outperform the statistical benchmark. The surface temperature depends on physical variables that can be difficult to observe, such as subsurface temperatures or snow layer properties. Simple relationships with atmospheric variables are insufficient to describe this complexity. Therefore, it is useful to develop heat balance models, such as urban models, land surface models, and snow models, for land surface modeling, artificial environments. The urban climate model is less effective for roads with human activities with snow ploughing, salting, and traffic as shown in Hajala experiments. The prediction of the road surface temperature variable is better for TEB-ES than for TEB but they are both outperformed by the in-sample MLR benchmark.

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Overall Our study suggests that new developments within TEB are interesting for artificial surface predictions but are flawed for roads impacted by human activities. Indeed, overall model performance for the Finland experiment is poorer than for the Col de Porte experiment, as shown by the experiment's missed event, false detection and false alarm rates shown in Table 7, Table 6, Table 4 and Table 3 analysis and scores. This inferior performance is caused by several factors, including errors in

Models	Detection rate %	Missed event rate %	False detection positive rate %	False alarm discovery rate %
TEB-ES	56. 57	44 .43	25 <u>24</u>	70 - <u>69</u>

Table 8. Performance of TEB and TEB-ES ice depth occurrence greater than 0.001mm, for 1hour time step, at the Hajala site during the winter 2017-2018

modeling caused by human activities: errors in modelling snow removal, anthropic effects not modeled salting not modelled, or traffic effects not modelled (snow compaction and heating effects by traffic), precipitation not detected by the raingauge, errors in distinguishing between snow and rain, and sensor detection errors). In fact, the traffic has a large effect on snow compaction. It drops: it reduces the snow depth and leads to measurements errors. Also measurement errors. In addition, Finland's winter road maintenance operator salt salts major roads whenever any a slippery road condition is observed or forecasted forecast. Snow ploughing and salting is roughly simulated in the models by mechanical snow and ice removal every morning at 6h UTC in the Hajala experiment. The actual effects and timings of the winter service vehicles are more complicated and impact the water contents and the surface heat energy. Salting indirectly affects road surface temperature by melting the snow cover that insulates the road from the atmosphere. Indeed, Fig.8 shows that while the large modelled snow depth keeps the road surface temperature steady on April 2nd, the variation of the measured road surface temperature follows the variation of the air temperature.

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The seasonal differences in atmospheric forcing impact various processes that influence road surface temperatures and lead to a seasonal road surface temperature performance trend (Fig. 6). During autumn, the road surface temperature increase by the heat conduction from the sub-layers, while in winter, the road surface temperature is driven by the road-atmosphere conduction. In spring, radiative forcing is higher and exhibits pronounced diurnal cycles. Road surface temperature performance differences between TEB and TEB-ES are significant due to the natural soil moisture conductivity initialization. The cold bias on road surface temperatures is reduced in TEB-ES by this initialization modification. It increases the heat restitution from the natural soil in winter. That situation is reflected on the November 7th, 8th Other measurements errors and sources of uncertainty may decrease the reliability of numerical experiments: lack of precipitation detection by the raingauge, errors in distinguishing between snow and rain, sensor detection errors, and 9th 1998 as described in the section above. It also reduces the radiation forcing errors from the ERA5 reanalysis.

The Col de Porte simulations have better performance since no human activity impacts the different variables. It allows us to evaluate in detail the snow/road coupled behaviour in the models with the 1-L and ES snow models. The performance of the TEB-ES road surface temperature cold bias when the snow cover insulates the road surface from the atmosphere.

The Col de Porte experiment is more reliable to validate the snowpack evolution simulations appears similar to the heat-balance road weather models at other locations (?, ?; ?, ?; ?, ?). Some of these models have been tested in open environments, while others have been tested in urban areas. The main differences in snow height between the TEB and TEB-ES models can be summarized summarized by three processes. First, the TEB-ES snow mantle depth tends to be lower than TEB snow mantle

in TEB-ES. Also In addition, fresh snow properties and accumulation on old snow cover is also better modeled are also better modelled in TEB-ES due to because of the modelling of a specific density for each layer in the model. Secondly, the TEB-ES snow mantle tends to be higher than the TEB a few hours after each snowfall. TEB snow density. The snow density of TEB follows a simple formulation with an exponential law to represent snow mantle aging whereas TEB-ES density the ageing of the snow mantle, while the density of TEB-ES is affected by weight compaction, melting, rainwater infiltration, and snowmelt retention. Thirdly, In Third, in snow-covered isothermal situations, there are large differences between TEB and TEB-ES. These isothermal situations are common during early winter and spring snowfalls, when the radiative forcing is high. The payement returns the energy stored in its structure to the snow cover. The lower layers of the snowpack melt, causing liquid water to drain. In ES and CROCUS as in CROCUS snow model, the lower snowpack reaches its maximum liquid water content and the snowpack temperatures are at the freezing point. Thus, in TEB with the simple 1-L snow model, snow-soil heat-transfer heat transfer is underestimated. The difficulty of snow models with one or few layers in representing the evolution of the SWE in the spring melt season has also been shown by Cristea et al. ((?)). Overall, the TEB-ES snow height follows the observed trend evolution more closely, as shown by the significantly higher R². There are some important differences between the snowpack evolution of TEB-ES and ISBA-Route/CROCUS but the overall snowpack height performance heights and road surface temperature in observed snow-covered situation as shown are close(Fig 2 - Fig. 6). However, the heat-balance model ISBA-Route is generally more cold biased than TEB and leads to larger false alarm and, false detection rates (Fig. 3, Fig. 4) are close. In one snow-covered event in early spring (not shown here), ISBA-Route/CROCUS has a very large error in simulated surface temperature unlike TEB and TEB-ES. This is caused by the different snow fraction formulation between TEB and ISBA-Route. In this study. Comparison of heat-balance models with statistical benchmarks provides interesting insight for further studies. The artificial surface is a low inertia and simple enough system with easily modelled behavior as shown by the good in-sample MLR performance in the Finland experiment. Although this behaviour is true in an open environment, more validation is needed with roadside components, trees, or buildings. The in-sample MLR Hajala simulation which has been trained using observed road surface temperature is also capable of correcting the forcing errors and captures the impacts of human activities. So, these components could be systematic and cyclical enough to be easily modelled. It means that there is potential for further studies to take into account these effects in the performance evaluation of TEB and its modified version TEB-ES has been carried out in open surrounding areas to investigate whether this urban climate model can accurately reproduce the road conditions predicted by road weather models. The TEB-ES heat-balance models. In both experiments, MLR models struggle to simulate the road surface temperatures when snow-covered. It leads to poor performance on Col de Porte with a mostly snow-covered road during the 6 month experiment. It suggests that the snow/road coupling is crucial for the heat-balance model performances. Indeed, it is difficult to capture surface physics when the road surface temperature is insulated from the atmosphere by the snow mantle. More complex statistical methods are needed, such as recurrent neural networks, to take into account the long-term system inertia and model the coevolution of road surface temperature performance shown in this

the TEB snow depth at the beginning of the events. Heat transfer between the payement and the snow is better represented

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of accurate mass content observations.

studyappears similar to heat-balance with road mass contents. However, training such models is likely to require the acquisition

Further research is needed to address modelling and evaluations limitations from this study. First, ice content modelling could be improved by finding a better estimate of the characteristic timescale for phase change τ set now at 25000s. This parameter value should be evaluated more rigorously in more experiments to get a better estimate. Then continuous effects from traffic 690 such as heating, snow compaction, turbulence, splash, and intermittent effects such as winter maintenance activities are not modelled in TEB and TEB-ES despite their major impact on artificial surface conditions (?, ?; ?, ?). This should be addressed in future work to match the mass contents observations on busy road lanes as in the Hajala experiment. In most road weather modelsat other locations (?, ?; ?, ?; , some of these effects are taken into accounts, with different levels of complexity (?, ?). Some of these models have been tested in open environments, while others have been tested in urban areas. TEB has 695 been created to simulate the urban climate and has been successful in modeling urban heat island effects (?). Further research is: ?, ?). It goes from simple linear modelling to the full parametrisation of the salting effects. Finally, in this study, we assessed the model on open environment only, to analyze the specific process at the surface. So, further research are needed to evaluate the road condition forecast these new processes in complex environment such as facing walls or roadside trees (?). TEB is used as well for surface boundary conditions, coupled with research (?) and operational (?) atmospheric models. Flux modeling assessment for modelcoupling should be performed as in Lipson et al. ?. There are some limitations in our winter road 700 condition modeling, in particular in ice content modeling and its evaluation, which could be upgraded with layer temperature evolution as in Chen et al. (?) or Fujimoto et al. (?). Missing real-world processes in the models may be a part of the explanation (?) for lower peformance in Hajala experiment. The influence of anthropic effects such as traffic and salting are not modeled in TEB and TEB-ES despite its major impact on the snow depth (?, ?; ?, ?) should behave well in such complex environments as 705 they have been the focus of the TEB developments for the past decades. To complete the evaluation of the model, particularly for using TEB coupled to an atmospheric model, winter fluxes should be extensively assessed at many locations. This could be performed following the Urban Plumber initiative, with extensive model comparisons (?). In relation with the former comment, we propose in Appendix 2, a snow removal parametrisation in urban environment to support further studies,

7 Conclusions

Bringing together the best of urban climate and road weather models benefits would benefit both communities. For In the urban climate community, better modeling of winter road conditions can lead to important improvements of heat fluxes in eities where winter conditions are frequentcold conditions have been largely understudied, allowing many unknowns about the urban climate response to harsh winter conditions. This study is a first step toward addressing the literature limitations on this topic by improving the modelling of artificial surface winter conditions, that have a major impact on cold cities climate.
715 Of particular relevance to the road weather community, improved modeling of the environment modelling of artificial surface conditions in an urban climate model could improve the accuracy of road condition prediction, especially in complex urban predictions in complex environments.

Thus, a A modified version of TEB from SURFEX-TEB v9.0 has been developed for winter road conditions. The road surface processes have been enhanced to model hazardous winter driving conditions to improve winter processes modelling. We

incorporated a basic ice content prognostic evolution to depict frost and water freezing on the road surface and a new, precise, 720 snow model that is coupled with the road. The model's new physics has TEB road component. The new physics have been verified at two different winter sites. One experiment was conducted carried out under controlled conditions at Col de Porte in the French Alps, while the other was based on a real-world scenario busy road with traffic and with recurrent winter maintenance operations in Haiala, Finland, TEB-ES significantly enhanced improved the surface condition prediction accuracy for the Col de Porte controlled experiment, outperforming both benchmarks the benchmarks provided by ISBA-Route/CROCUS and MLR. 725 Periods that are suitable for conducive to slippery conditions are well detected in TEB-ES. Thus, during During snowfall, the road's snow coverage snow coverage of the artificial surface is accurately simulated. However, Road surface temperatures are also more accurately predicted for both the Hajala and Col de Porte experiment with TEB-ES. However, the Hajala road weather station experiment has shown that further developments are needed to account for anthropic effects. Traffic heating anthropogenic effects. Future works on road heating by traffic, salting, water splashing, and snow compaction impact the road 730 surface conditions and the roadsurface temperature are expected to improve the model performance in terms of snow and ice content for high traffic managed road.

Appendix A: Statistical benchmark model

A Multiple Linear Regression (MLR) is developed in this study. In this study, two multiple linear regressions (MLRs) are developed for predictive purposes as a benchmark for the Col de Porte and Hajala experiment. The model is models are developed using the same knowledge used by the heat balance models that the heat balance models use to simulate the road surface temperaturetemperature of the road surface. Thus, the wind-speedwind speed, wind direction, solar radiation direct and scattered, longwave radiation, air temperature, specific humidity, and pressure forcings are used as input data for the MLR modelmodels. Following 2(?), the hourly forcing variables from a lag of 3 hours, 2 hours, and no lag are concatenated. In total, 24 explanatory variables are considered by the modelmodels. A backward feature selection procedure is performed to prevent avoid overfitting. The selection is made using the adjusted-R² criterion that, unlike the R² is not monotonically non-decreasing nondecreasing by the number of explanatory variables. The selection procedure results are drawn results of the selection procedure are drawn in Fig. A1 and Fig. A2 with cross-validation estimator performance uses. The maximum mean adjusted-R² calculated by the selection procedure shown in both experiments shown in Fig. A1 is and Fig. A2, are used to select the variables needed for the modelmodels. The maximum mean adjusted-R² is are receptively 0.678 for nine variables extracted shown and 0.967 for nine and seven variables extracted as shown in Table A1. In this paper, the model predictions are made in-sample. It means that the data used for inference are also used to learn the model.

Selected variables with lag	-3 hours	-2 hours	No lag
CDP experiment	Temperature Direct shortwave	Temperature Direct shortwave	Temperature Direct shortwave
	Specific humidity		Scattered Shortwave Longwave
Hajala experiment	Temperature	Wind	Temperature Direct shortwave Scattered Shortwave Specific humidity Longwave

Table A1. Variables used for the MLR model learning and inference for the Hajala and the Col de Porte experiments

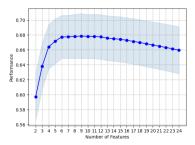


Figure A1. Sequential feature selection performances with 20 cross-validation steps, for the multiple linear regression at \overrightarrow{CDP} with the adjusted- \mathbb{R}^2 in function of the backward selected features number, with the 0.95 confidence interval

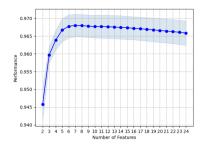


Figure A2. Sequential feature selection performances with 20 cross-validation steps, for the multiple linear regression at Hajala with the adjusted- R^2 in function of the backward selected features number, with the 0.95 confidence interval

Appendix B: Snow cover parametrisation in cities

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In cities, winter operations do not completely remove the snow cover. It remains for much longer on sideways or on car parks, and the excess snow on the road is gathered and packed at specific spots. This is particularly true in cold cities, subject to significant snowfall, with months of snow left over on the surfaces. While for road-focused simulations, the snow fraction is reset to 0 to model the complete removal of the snow cover on the road, this parametrisation is incorrect at a city scale. Thus, the snow cover occupation is adapted here for urban environments.

We developed an option that modifies how the snow cover occupation is handled in the model. When road maintenance operation occurs, snow cover is collected and stacked on a fraction of the total road surface. In the model, we call this fraction $f_{snstore}$. This parameter should then be representative of the remaining snow-covered surface fraction from the city snow removal operation procedure.

Between two snow removal operations, the fresh snow, from new snowfall, accumulates. The prognostic fresh snow height variable D_{fresh} evolves as:

$$\frac{\partial D_{fresh}}{\partial t} = \frac{P_n}{\rho_{new}} \tag{B1}$$

With P_n the snow rate in $(kg m^{-2} s^{-1})$ and ρ_{new} the fresh snow density in $(kg m^{-3})$ as defined in Sect. 2.3. D_{fresh} is set to zero when the snow removal operations occur in the model, as it is then fully considered as old snow.

This fresh snow is considered to fall first, into the snow-covered area of the road, sidewalks, or car park, then to gradually cover the previously cleared roads and sidewalks, as shown in Fig.C1. The fresh snow occupation fraction is modelled as:

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$$f_{snfresh} = \min(1, \max(D_{fresh}/0.01, f_{snstore}))$$
 (B2)

So, $f_{snfresh}$ increases when the fresh snow cover exceeds the "old snow" cover.

To parameterize the total occupation fraction of the snow cover, this fraction is compared to the total snow mantle depth D_s , which is the sum of the fresh snow, and the piled up snow from the previous snow removal operation, used in Eq. 12. Thus, the new fraction is computed as:

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$$f_{sn} = \min(\min(1.0, D_s/0.01), f_{snfresh})$$
 (B3)

$$\iff f_{sn} = \min(1, D_s/0.01, \max(D_{fresh}/0.01, f_{snstore}))$$
(B4)

This also means that during spring or even relatively warm situation in the cold season, when the total of the snow layer starts to completely melt, the fraction of snow will become smaller than the fraction of impervious surfaces assigned to store the old snow. This modelling is subject to an important assumption: for all computations linked to other processes than cover fraction on the impervious surfaces, the new fresh snow is considered to have the same properties and behaviour of the residual snow mantle. Thus, the evolution of the snow mantle is computed only once with the total snow height D_8 , although the properties of the snow layer are not really horizontally homogeneous in the model tile considered.

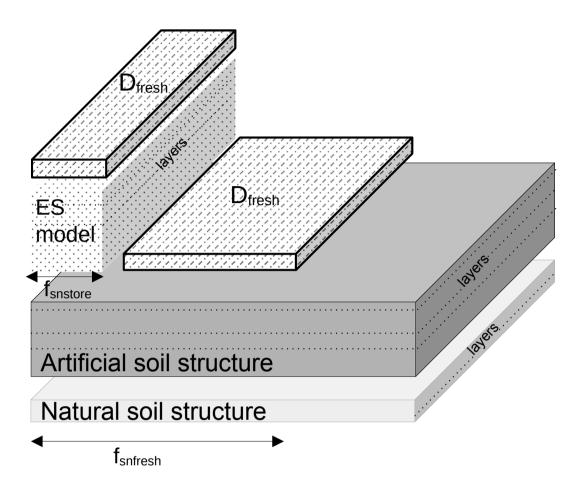


Figure C1. Schematic implementation of the new snow fraction procedure between two snow removal operations, as explained in the text when the fresh snow cover a fraction of the cleared road

Code and data availability. TEB is embedded in the software SURFEX available from the CNRM open source website: https://opensource. umr-cnrm.fr under the CeCILL Free Software License Agreement v1.0 license. The exact version of SURFEX v9.0 including the TEB model, the TEB-ES model, and the MLRs statistical models used to produce the results in this paper are available for public access on the Zenodo platform (?), as are the input data to run the models and the output data to evaluate all the simulations presented in this article. ISBA-Route/CROCUS code is not publicly available because it is not an open-source model.

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Author contributions. GC conducted the model improvements and benchmarks, investigation, methodology, formal analysis, validation, data curation and wrote the paper. VM, FB and LC conceptualized and supervised the project, participated for the methodology, validation, and reviewed the paper. VK has participated in the data curation, validation, and reviewed the paper. LP made a draft of the model improvements with investigations and validations and reviewed the paper. All authors discussed the performance of the models.

Competing interests. The authors declare that they have no conflict of interest