



1 **Increasing Area and Decreasing Depth: Climate Change**

2 **Influence on Snow Variations in the Qilian Mountains**

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16

17 **Abstract:** The Qilian Mountains serve as a critical water source for the Yellow River and various  
18 inland rivers, playing a pivotal role in regulating the regional climate. Given their significance as  
19 one of the foremost water resources in the area, the spatial and temporal dynamics of the snow are  
20 crucial for understanding potential impacts on regional hydrology and ecology. This study  
21 examines the characteristics of spatial and temporal variations in snow-covered extent (SCE),  
22 snow depth (SD), snow-covered days (SCD), snow onset date (SOD), and snow end date (SED)  
23 within the Qilian Mountains region. We investigate the hydrological and ecological implications  
24 utilizing snow area and phenology data, alongside SD data. The findings indicate that: (1) the  
25 distribution of snow across the Qilian Mountains mainly splits between the central and western  
26 areas, with the central region showing deeper snow than both the eastern and western parts; (2) the  
27 area covered by snow in the Qilian Mountains is growing, but the depth of the snow is on a  
28 decline, especially in the central area; (3) in terms of snow phenology, most of the region is  
29 witnessing an earlier start of SOD, a longer SCD, and an earlier SED. An overall increase in



30 precipitation is identified as the key factor behind the expanded SCE in the Qilian Mountains,  
31 while rising temperatures are pinpointed as the primary cause for the reduction in SD. As global  
32 climate change intensifies, the observed alterations in the snow of the Qilian Mountains present  
33 emerging challenges for regional water security and ecological equilibrium.

34 **Keywords:** Qilian Mountains, snow, water resources, climate change

### 35 **1. Introduction**

36 Snow constitutes the predominant form of cryospheric moisture (Pulliainen, J., et al. 2020 ).It  
37 is estimated that over one-sixth of the global population relies on glaciers and seasonal snows as a  
38 critical water supply (Barnett, J., et al. 2005). Alpine regions play a crucial role in serving as  
39 reservoirs for ice and snowmelt recharge. Specifically, the Qilian Mountains capture airborne  
40 water vapor, leading to the formation of both permanent glaciers and seasonal snows. This process  
41 underpins the watershed of numerous streams that sustain the Hexi Corridor and the Qaidam Basin  
42 (Zhu, et al. 2022). Consequently, a comprehensive understanding of the spatial and temporal  
43 dynamics of mountain snow, alongside its evolving patterns, is imperative for safeguarding the  
44 ecological equilibrium of oasis environments. Furthermore, such knowledge is vital for ensuring  
45 the continued sustainable development of both the economy and society within arid regions.

46 Globally, approximately 78% of mountainous snow exhibits a declining trend, characterized  
47 by a reduction in snow duration by up to 43 days in extreme cases and a decrease in snow-covered  
48 area by up to 13% (Notarnicola, et al. 2020). Within the highland rangelands of Central Asia, a  
49 positive correlation has been observed between the Peak Height of Vegetation (PH) and  
50 snow-covered days (SCD), alongside a negative correlation between Thermal Time to Peak (TTP)  
51 and SCD. Topographical attributes, particularly slope and orientation, have been identified to  
52 significantly influence the snow end date (SED) (Tomaszewska, et al. 2020). Snow phenology  
53 parameters such as SCD, snow onset date (SOD), and SED have undergone notable changes in the  
54 arid regions of Asia (Tang, et al. 2022). An increasing trend in snow has been documented on the  
55 southwestern edge and southeastern part of the Tibetan Plateau (Huang, et al. 2016). Since the turn  
56 of the millennium, the consequences of global warming have grown increasingly pronounced,  
57 making snow changes in the Northwest a subject of considerable interest in climate change  
58 research.

59 The effects of climate change on the cryosphere have significantly influenced hydrological



60 processes and water resources in mountainous areas, leading to increased runoff due to accelerated  
61 glaciation and snowmelt (Dibesh, et al. 2014). Therefore, analyzing the snow depth(SD), along  
62 with spatial and temporal variation characteristics of snow in the Qilian Mountains, holds  
63 paramount importance for comprehending climate change and ecological evolution in  
64 mountainous regions. While previous research has predominantly concentrated on changes in  
65 snow-covered extent (SCE) and SD, a comprehensive analysis of the spatio-temporal patterns of  
66 snow phenology across the Qilian Mountains is critical for appraising and forecasting future  
67 climatic conditions. Furthermore, to gain a more profound understanding of the alterations in  
68 snow and snow phenology, along with their driving mechanisms and impacts on the global climate  
69 system, it is imperative to conduct analyses on the scale of the entire study area.

70 Snowmelt water plays a critical role in supporting local agriculture and livestock irrigation;  
71 consequently, variations in the SCE can profoundly impact local environmental dynamics and  
72 human livelihoods. Situated in the arid region of Central Asia, the Qilian Mountains experience  
73 substantial snowfall in winter, whereas summers are characterized by high temperatures and scant  
74 rainfall, positioning snowmelt runoff as a pivotal water resource for urban, agricultural, and  
75 industrial development within the region (Wu, et al. 2021). Hence, sustained and precise  
76 monitoring of the temporal and spatial variations in the SCE of the Qilian Mountains is of both  
77 practical and theoretical importance. Such efforts are crucial for advancing our understanding of  
78 snowmelt runoff dynamics in mountainous areas, facilitating the effective management and  
79 utilization of water resources, and preparing for winter snowstorms and spring and summer floods  
80 in pastoral and agricultural regions, respectively.

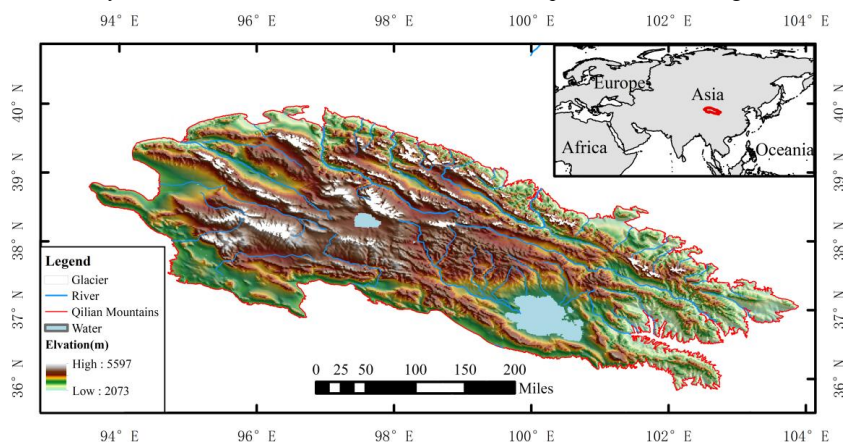
81 This study employs snow data spanning from 1980 to 2019 to analyze the spatial and  
82 temporal patterns of changes in snow depth and area within the Qilian Mountains, aiming to:(1)  
83 delineate the processes underlying spatial and temporal variations in snow accumulation; (2)  
84 identify the characteristics of snow volume and snow phenology; and (3) investigate the  
85 implications of snow accumulation changes for watershed water resources. The findings of this  
86 research offer a scientific foundation for elucidating the impacts of climate change on the Qilian  
87 Mountains' snow and for guiding the protective management, rational development, and utilization  
88 of snow resources in the area.

## 89 **2.Data and methods**



90 **2.1 Study area**

91 The Qilian Mountains, situated in the northeastern segment of the Qinghai-Tibet Plateau,  
92 span across Qinghai and Gansu Provinces. Characterized by its complex and varied topography,  
93 the region predominantly features altitudes ranging from 3,500 to 5,000 meters. Owing to its  
94 positioning within the north temperate zone and the consequential effects of elevation, the Qilian  
95 Mountains region records an average annual temperature that generally remains below 4 degrees  
96 Celsius. Precipitation patterns within this area are influenced by a myriad of factors, including  
97 altitude, geographic location, as well as the slope and orientation of the terrain, resulting in  
98 pronounced seasonal and inter-annual variabilities. The Qilian Mountains encompass a diverse  
99 snow classification, comprising permanent snow areas, stable seasonal snow areas, annually  
100 cyclical unstable seasonal snow areas, and non-annually cyclical unstable seasonal snow areas.  
101 This diversity culminates in a multifaceted snow distribution pattern across the region.



102  
103 **Fig. 1 Study area**

104 **2.2 Data description**

105 This research leveraged snow data products from the period 2000-2020, obtained  
106 from the National Cryosphere Desert Data Center (NCDC), including MODIS  
107 (Moderate Resolution Imaging Spectroradiometer) day-by-day clear-sky snow  
108 products and phenology dataset products. Furthermore, it utilized the "Big Earth  
109 Data" longitudinal snow depth dataset (referenced in Table 1). This dataset elucidates  
110 the snow accumulation characteristics across China, employing the MODIS  
111 reflectance product MOD/MYD09GA. A novel multi-index combined snow



112 accumulation discrimination algorithm was developed, tailored to different land cover  
 113 types. This algorithm notably enhances the precision of identifying snow  
 114 accumulation areas within forested and mountainous regions. Additionally, it  
 115 implements complete cloud removal using the Hidden Markov Model and integrates  
 116 multi-source data fusion methodologies.

117 The data fusion framework integrates various sources, including the Advanced  
 118 Microwave Scanning Radiometer-2 (AMSR2), Global Snow Monitoring for Climate  
 119 Research (GlobSnow), Northern Hemisphere Snow Depth (NHSD), ERA-Interim,  
 120 and the Modern-Era Retrospective Analysis for Research and Applications, version 2  
 121 (MERRA-2). It incorporates geographic (latitude and longitude) and topographical  
 122 (elevation) data as input independent variables. Utilizing over 30,000 terrestrial-based  
 123 observations as dependent variables, the models underwent training and validation  
 124 across distinct temporal scales. This comprehensive fusion framework yielded a  
 125 longitudinally continuous daily snow depth dataset for the Northern Hemisphere,  
 126 featuring a spatial resolution of 0.25°.

127

**Table 1 Data source**

Data name	Spatial resolution/m	Temporal resolution	Format	Number of files	Source
A long-term daily gridded snow depth dataset for the Northern Hemisphere from 1980 to 2019 based on machine learning	27000	1day	TIF	10,656	Big Earth Data ( <a href="https://doi.org/10.1080/20964471.2023.2177435">https://doi.org/10.1080/20964471.2023.2177435</a> )
China MODIS Daily Cloudless 500m Snow Area Product Dataset	500	1day	HDF5	7,614	National Cryosphere Desert Data Center( <a href="http://www.ncdc.ac.cn">www.ncdc.ac.cn</a> )
A dataset of snow phenology in China based on MODIS from 2000 to 2020	500	1year	TIF	60	National Cryosphere Desert Data Center( <a href="http://www.ncdc.ac.cn">www.ncdc.ac.cn</a> )
1-km monthly precipitation dataset for China (1901-2022)	1000	1month	NetCDF	1,464	National Tibetan Plateau Data Center ( <a href="https://data.tpdc.ac.cn/">https://data.tpdc.ac.cn/</a> )
CSR GRACE/GRACE-FO RL06.2 Mascon Solutions (RL0602)	27000	1month	NetCDF	255	Gravity Recovery and climate Experiment ( <a href="http://www2.csr.utexas.edu/grace">http://www2.csr.utexas.edu/grace</a> )

128 **2.2Data processing**

129 **2.2.1Snow cover**

130 The presence of extensive cloud cover in snow-covered regions presents a significant  
 131 challenge to the utilization of MODIS (Moderate Resolution Imaging Spectroradiometer) snow  
 132 products for snow monitoring. This study has addressed this issue by employing a day-by-day



133 cloud-free dataset, which classifies data as follows: 0 for land, 1 for image-recognized snow, 2 for  
134 de-clouded interpolated snow, 3 for snow-depth interpolated snow, 4 for water, and 255 for  
135 regions that are unrecognizable. Prior to analysis, the snow product undergoes a preprocessing  
136 phase. Given that the product is in the HDF5 format, a batch conversion to the GeoTIFF format,  
137 which incorporates a geographic coordinate system (GCS), is necessary to facilitate visualization.  
138 Subsequent steps involve cropping the data to match the study area and reclassifying snow raster  
139 values (where  $t=1,2,3$ ) to  $i=1$  (indicating snow presence), and no-snow raster values (where  
140  $t=0,4,255$ ) to  $i=0$  (indicating snow absence). The mean values of these raster datasets within the  
141 study area are then calculated using band set statistics to determine the snow-covered extent  
142 (SCE). To accurately delineate the snow-covered area, a conversion from the geographic  
143 coordinate system (GCS) to a projected coordinate system (PCS) is required. This process, along  
144 with operations such as image superimposition and geometric computation, facilitates the  
145 determination of the study area's snow-covered area. Data preprocessing is accomplished through  
146 the application of Python programming and MODIS-specific software tools, including the MODIS  
147 Reprojection Tools (MRT) and the Environment for Visualizing Images (ENVI).

$$SCE (\%) = \frac{Count (i = 1)}{Count (i = 1, 0)}$$

### 148 **2.2.2 Snow depth**

149 Research on SD predominantly concentrates on the snow period, which is characterized by  
150 the presence of snow with considerable depth. Defined as the interval during a year when snow is  
151 continuous, the snow period generally extends from the occurrence of the first widespread  
152 snowfall to the complete melting of spring snow, traditionally spanning from November 1 to  
153 March 31 of the subsequent year. In the context of this study, the time series of SD dataset  
154 products were reclassified to align with the snow accumulation period, utilizing an extensive data  
155 series from 1980 to 2019. This reclassification was further delineated into five-year intervals:  
156 1980-1984, 1985-1989, 1990-1994, 1995-1999, 2000-2004, 2005-2009, 2010-2014, and  
157 2015-2019. The objective of calculating the average SD during the snow period was to facilitate a  
158 more intuitive understanding of the overall snow dynamics and its temporal variations.

### 159 **2.2.3 Snow Phenology**

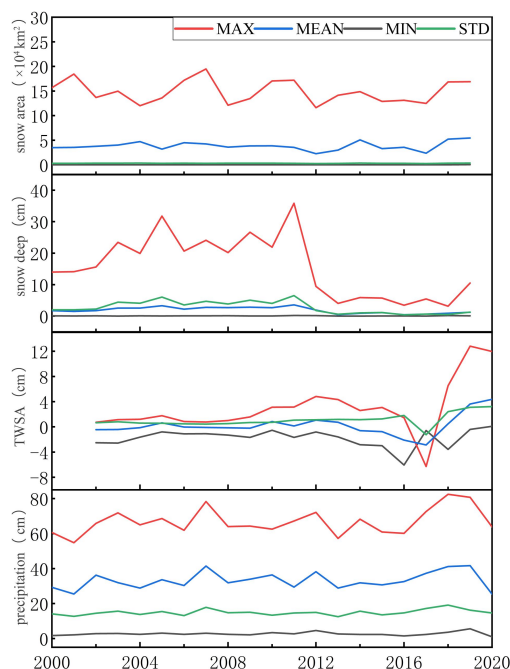
160 Similar to the methodology applied to the analysis of SCE, the study employed a



161 reassignment approach for raster data, subsequently reclassifying the snow phenology dataset in  
 162 accordance with the hydrological year by segmenting the product time series. The evaluation of  
 163 SCD、SOD and SED was conducted on an annual basis, per hydrological year and per image  
 164 element. These calculations derived from the predefined snow-covered climate parameters tailored  
 165 for China, spanning the years 2000 to 2020, facilitated a nuanced understanding of snow dynamics  
 166 within the specified period.

### 167 3 Results

#### 168 3.1 Spatial and temporal variation of snow area in the Qilian Mountains

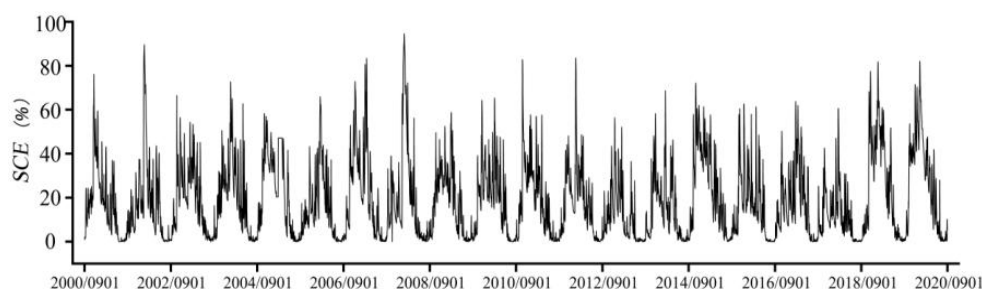


169  
 170 **Fig. 2 Interannual Changes in snow area, snow deep, Terrestrial Water Storage Anomaly (TWSA), and**  
 171 **precipitation in Qilian Mountains 2000-2020**

172 The analysis of the interannual mean snow in the Qilian Mountains from 2000 to 2020  
 173 reveals a generally fluctuating upward trend, characterized by periods of increase and decrease.  
 174 Specifically, increments were observed during the periods of 2000/2001-2004/2005,  
 175 2012/2013-2014/2015, and 2017/2018-2019/2020, while declines were noted in  
 176 2006/2007-2012/2013 and again in 2015-2017/2018. The peak of the interannual mean  
 177 snow-covered area was recorded in 2019/2020, reaching approximately  $5.45 \times 10^4 \text{ km}^2$ , whereas the  
 178 lowest extent was observed in 2012/2013, at nearly  $2.26 \times 10^4 \text{ km}^2$ . Over the entire study period



179 from 2000/2001 to 2019/2020, the trend of the mean snow-covered area exhibited a growth rate of  
180 approximately  $0.17 \times 10^4 \text{ km}^2/\text{a}$  per annum. A historical maximum snowage area was noted on  
181 January 27, 2008, with an extent of  $19.47 \times 10^4 \text{ km}^2$ , corresponding to a 94.52% snow. In contrast,  
182 the year 2012/2013 recorded the lowest maximum snow area within the same timeframe, with a  
183 coverage of  $11.6 \times 10^4 \text{ km}^2$  and snow constituting 78.35%. It is notable that during most years, there  
184 was at least one day with a complete absence of snow in the ablation period. Exceptions to this  
185 pattern were recorded in 2002/2003 and 2019/2020, during which snow persisted throughout the  
186 year, maintaining minimal areas of  $82.4 \text{ km}^2$  and  $206.01 \text{ km}^2$ , respectively. Peak snow areas in the  
187 Qilian Mountains typically manifested in November and January. The analysis identified four  
188 notable peaks in the inter-annual mean snow area, occurring in 2004/2005 ( $4.72 \times 10^4 \text{ km}^2$ ),  
189 2014/2015 ( $5.07 \times 10^4 \text{ km}^2$ ), 2018/2019 ( $5.21 \times 10^4 \text{ km}^2$ ), and 2019/2020 ( $5.45 \times 10^4 \text{ km}^2$ ). Additionally,  
190 the SCE within the Qilian Mountains exhibited significant seasonal variability. December and  
191 January witnessed the highest SCE levels, characterized by substantial fluctuations, whereas from  
192 June to September, the SCE's standard deviation decreased, often reaching a low value or even  
193 zero, as depicted in Figures 2 and 3.



194 **Fig. 3 Changes in SCE (%) of snow in the Qilian Mountains for multi-year days 2000-2020**

### 195 **3.2 Spatial and temporal variations of snow in the Qilian Mountains**

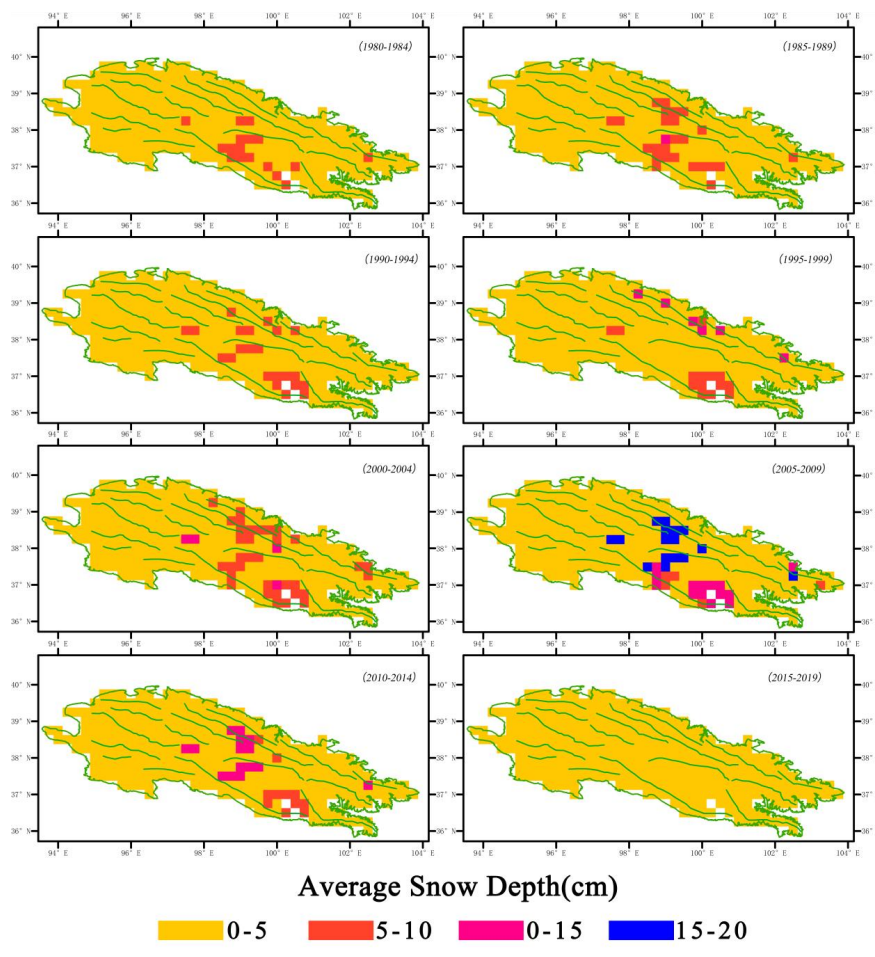
196 SD is widely recognized as a crucial metric for assessing snow conditions. It is defined as the  
197 vertical thickness of snow present on the ground. While SD does not provide direct information  
198 about the volume or mass of the snow, it serves as a reliable proxy due to the simplicity and  
199 immediacy of its measurement. In this research, the multi-year average SD within the Qilian  
200 Mountains snow was evaluated using data collected from 1980 to 2019. The analysis of this  
201 dataset reveals distinct spatial distribution patterns of snow across the Qilian Mountains, with the  
202 highest concentrations of snow predominantly located in the mountain range's central area. These





203 spatial characteristics highlight a notable trend: regions east of 98°E exhibit greater snow  
204 accumulations compared to those west of 98°E. Furthermore, regarding elevation distribution, the  
205 elevated areas west of 98°E, particularly those with altitudes ranging between 3,500 and 4,000  
206 meters and featuring SDs exceeding 5 cm, serve as the primary zones for substantial snow  
207 accumulation. This observation underscores a significant concentration of the densest snow within  
208 the central sector of the Qilian Mountains, establishing a demonstrable link between snow  
209 accumulation patterns and both geographical positioning (longitude) and elevation (Fig. 5a).

210       Between 2005 and 2009, the observed spatial distribution of SDs within the region was  
211 favorable, especially in zones where SDs ranged from 15 to 20 centimeters. Conversely, the period  
212 from 2015 to 2019 marked a significant reduction in the region's SD performance, with average  
213 SDs consistently registering between 0 and 5 centimeters across the area. This pattern denotes a  
214 marked, initial uptrend in snow within the Qilian Mountains region from 1980 to 1984,  
215 progressing through to 2005-2009. Post-2009, however, the trend in snow began to demonstrate a  
216 decline, with regions previously averaging more than 5 cm in SD gradually diminishing until such  
217 extents were no longer observed (Fig. 3). In synthesizing the SD distribution data across the Qilian  
218 Mountains over the past four decades, it becomes evident that the interval from 2005 to 2009  
219 represented a zenith in terms of snow levels within the region, succeeded by a substantial decline  
220 between 2015 and 2019. This observed trend underscores the susceptibility of the region's snow  
221 distribution to climatic shifts and anthropogenic influences. Consequently, it heralds potential  
222 future challenges for climate change mitigation efforts and strategies pertaining to regional water  
223 resource management.



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**Fig. 4 Changes in the distribution of mean SD in the Qilian mountainous area snow from 1980-2019**

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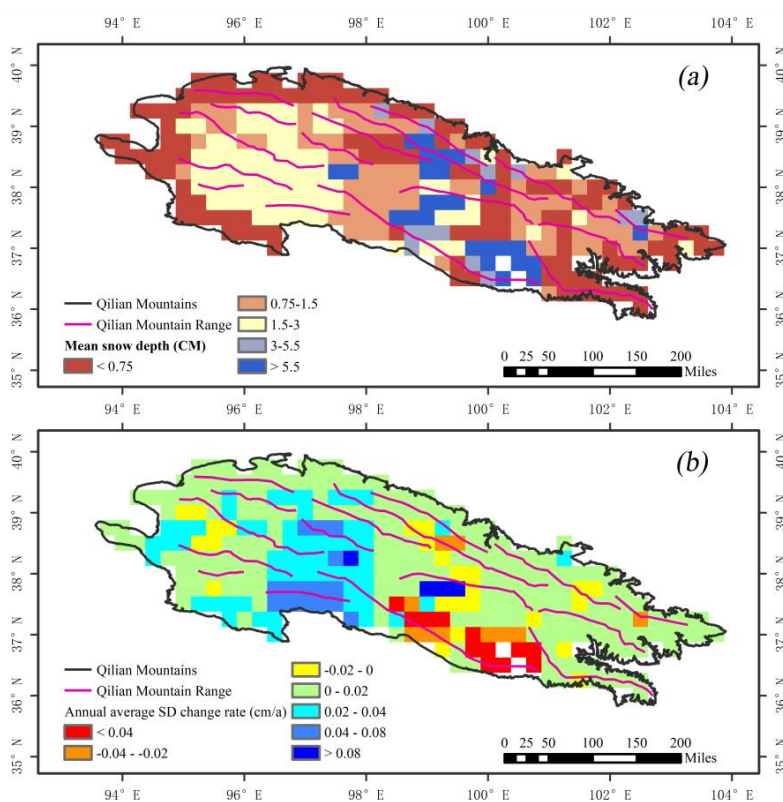
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The analysis of inter-annual variations in the mean SD in the Qilian Mountains from 1980 to 2019 reveals that the SD exhibited significant volatility up to the 1994/1995 period. Subsequently, a gradual yet steady upward trend in SD was observed until the 2005/2006 period. Despite experiencing some fluctuations within this time frame, the SD consistently increased, culminating in a peak depth of 3.577 cm in the 2011/2012 season. It is critical to underscore that, post-2011, the SDs have undergone a marked decline. Although this decline was interspersed with periodic fluctuations, the overall SD levels have persistently remained low up until 2019. This analysis highlights notable temporal fluctuations in SD within the region, underscoring a period of growth followed by a significant and sustained reduction in depths observed in recent years.

The analysis of the interannual rate of change in mean SD between 1980 and 2019 within the



236 Qilian Mountains demonstrates a significant decreasing trend in SD to the east of 98°E. This  
 237 decline is notably acute at elevations around 3,500 meters, where the inter annual rate of change  
 238 for the majority of this region registers at less than -0.04 cm/year. Historically, this area has been  
 239 characterized by relatively deeper SDs. In contrast, regions situated to the west of 98°E exhibit a  
 240 modest increase in SD, with inter annual rates of change varying between 0.04 and 0.08 cm/year  
 241 (Fig. 5b). These contrasting trends likely underscore the spatial heterogeneity of climate change  
 242 impacts across different regions and highlight how topographical and elevation gradients influence  
 243 precipitation and temperature distributions.



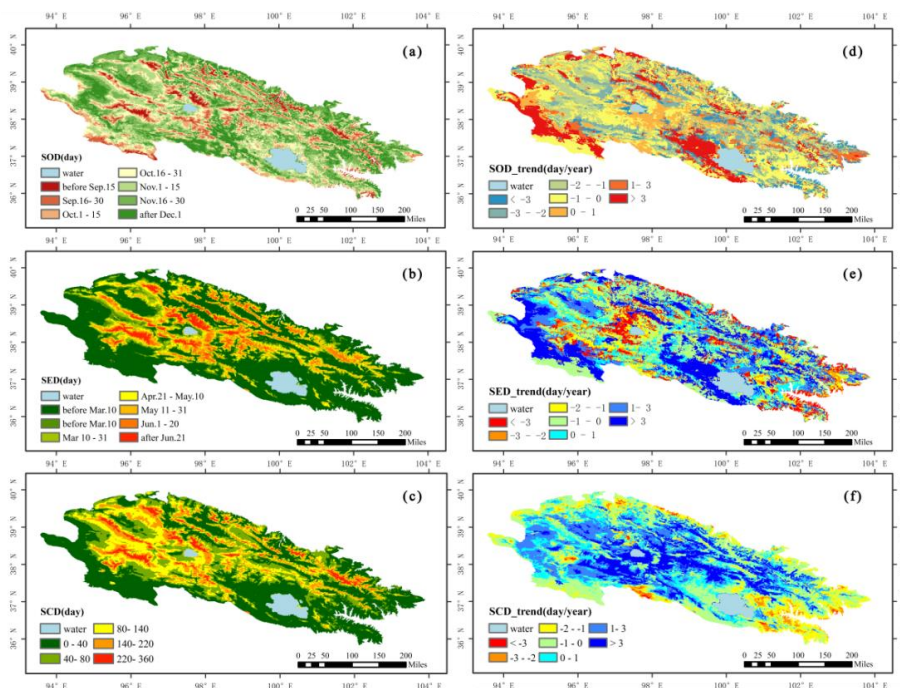
244  
 245 **Fig. 5 Mean SD in the snow of the Qilian Mountains, 1980-2019 (a) Distribution of mean SD; (b)**  
 246 **Distribution of inter-annual rate of change of mean SD**

247 **3.3 Alterations in snow Phenology within the Qilian Mountains**

248 The physical characteristics of the snow within the Qilian Mountains exhibit substantial  
 249 spatial heterogeneity. This study analyzed the spatial distribution characteristics of the SOD ,  
 250 SED and SCD utilizing observational data, which revealed pronounced disparities. Generally, the  
 251 SOD across the region predominantly occurs post-December, while the SED is chiefly observed



252 before May of the subsequent year, with the SCD primarily spanning 0 to 40 days. However, this  
 253 broad characterization markedly contrasts with the conditions observed in the High Altitude  
 254 Mountains. In the higher elevation zones, the climatic attributes of the snow display an earlier SOD  
 255 and a later SED. Specifically, the SOD typically commences before September 15, and the SED  
 256 extends beyond June 21st. Furthermore, the SCD in these areas significantly exceeds that of lower  
 257 elevations, ranging between 220 and 360 days. This pattern is predominantly observed along the  
 258 major mountain ridges of the Qilian Mountains, underscoring the profound impact of elevation on  
 259 snow phenology. Conversely, areas of lower altitude situated on the periphery of the Qilian  
 260 Mountains exhibit an antithetical trend, characterized by later SOD, earlier SED, and consequently,  
 261 a reduced SCD. These observations highlight the intricate relationship between altitude and snow  
 262 dynamics within the Qilian Mountains, reflecting the complex interplay of geographical and  
 263 climatic factors in shaping regional snow phenology.



264  
 265 **Fig. 6 Spatial distribution of multi-year average snow phenology in Qilian Mountains snow from 2000 to**  
 266 **2020: (a) SOD (b) SED (c) SCD; Spatial distribution of the trend of multi-year snow phenology in the Qilian**  
 267 **Mountains from 2000 to 2020: (d) SOD (e) SED (f) SCD**

268 The analysis of changes in snow phenology reveals significant spatial heterogeneity across  
 269 the Qilian Mountains. Examination of the interannual variation of the SOD indicates that the



270 majority of the area, approximately 68.94%, experienced an advancement in SOD. Conversely,  
271 approximately 12.12% of the snow areas within the study region exhibited a noticeable delay in  
272 SOD, with a change potential greater than 3 days per annum (d/a). This delay predominantly  
273 occurred in the low-altitude areas at the southwestern edge of the Qilian Mountains region and in  
274 central-southern mid-altitude areas. A smaller proportion, about 6.8%, of the snow area witnessed  
275 a significant advancement in SOD, with a change potential less than -3d/a, and this distribution  
276 was more scattered across the region (as shown in Fig. 6d).

277       Regarding the interannual variation of the SED, the data reveals a roughly equal division  
278 between areas experiencing delayed and advanced SED within the Qilian Mountains. Areas with  
279 advanced SED accounted for approximately 51.6% and were slightly more prevalent than those  
280 with delayed SED. Notably, the area with significantly delayed SED made up 23.15% of the  
281 region, exhibiting a trend greater than 3d/a and featuring a wide and sporadic distribution. In  
282 contrast, areas with significantly advanced SED constituted 9.09% of the snow, demonstrating a  
283 trend of less than -3d/a with a similarly sporadic distribution (as depicted in Fig. 6e). Analysis of  
284 the interannual variation of the SCD indicated that the area within the Qilian Mountains  
285 experiencing an extension of SCD comprised 54.12% of the total region, which was marginally  
286 higher than the area witnessing a reduction in SCD. Importantly, regions with significantly  
287 prolonged SCD accounted for 15.06%, featuring a trend of change greater than 3d/a. In stark  
288 contrast, regions with a significantly reduced SCD represented a mere 0.3%, with a trend of  
289 change less than -3d/a (as illustrated in Fig. 6f). This detailed analysis underscores the complex  
290 and varied impact of climate dynamics on snow phenology across different geographical and  
291 altitude gradients within the Qilian Mountains.

## 292 **4. Discussion**

### 293 **4.1 Mechanisms driving snow changes**

294       As a pivotal mountain range in northwestern China, the Qilian Mountains are influenced by  
295 an array of potential mechanisms driving changes in snow, shaped by both natural and  
296 anthropogenic factors. Observations from 2000 to 2019 indicate a dynamic trend in the  
297 snow-covered area within the region, characterized by an initial increase followed by a subsequent  
298 decline. Specifically, the period from 2000 to 2005 witnessed an increasing trend in snow-covered  
299 area, whereas a pronounced decreasing trend was observed from 2008 to 2013. Further analysis



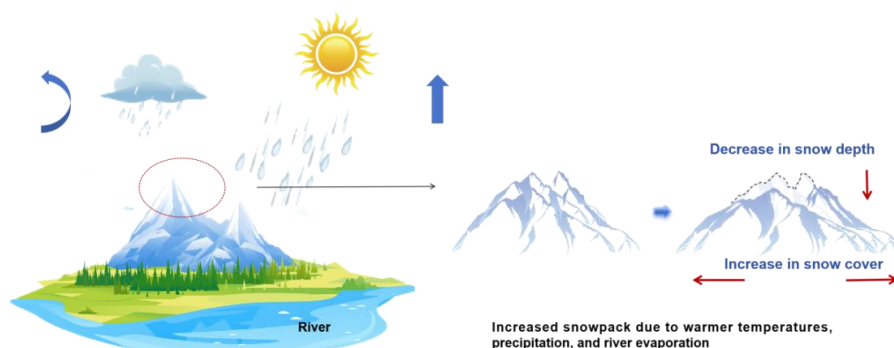
300 aligns the fluctuations in precipitation within the Qilian Mountains with the observed trends in  
301 SCE, strongly indicating that variations in precipitation exert a significant impact on the snow  
302 dynamics. This finding underscores the critical role of precipitation as a pivotal climatic  
303 determinant in the modulation of snow characteristics. The investigation into the effects of wind  
304 speed and temperature on snow variability posits that temperature, despite the global warming  
305 phenomenon, exerts a relatively minor influence on the interannual variability of snow area in the  
306 Qilian Mountains. This assertion is based on the analysis which demonstrates that the interannual  
307 variability in these climatic factors is relatively small, consequently having a limited direct impact  
308 on SCE. However, an analysis covering the years 2000 to 2020 revealed a generally fluctuating  
309 upward trend in the mean annual SCE within the Qilian Mountains. This observation suggests that  
310 the interannual variability of the SCE, even in a global warming context, remains complex and  
311 variable (as shown in Fig. 7).

312 The aforementioned insight does not negate the influence of climate change on the snow in  
313 the Qilian Mountains. Instead, it highlights the imperative for further research into the complexity  
314 and multifaceted nature of climate change impacts on snow dynamics, aiming for a more precise  
315 understanding and prediction of future snow trends. For example, this study's analysis of spatial  
316 and temporal changes in snow over the extensive period of 1980-2019 revealed a significant  
317 decline in mean SD values in the Qilian Mountains beginning after a peak in the 2011/2012 season,  
318 with no areas recording mean SD values exceeding 5 cm in the period of 2015-2019. This  
319 indicates that the trends in snow phenology could be critical for comprehending snow dynamics  
320 amid climate change, suggesting shifts towards higher winter temperatures, alterations in snowfall  
321 patterns, earlier snowmelt, and shortened snow season durations. Nevertheless, changes observed  
322 in the snow phenology of the Qilian Mountains, with a majority of the area experiencing an  
323 advance in the SOD (68.94%), a slight predominance of extended SCD over decreased areas  
324 (54.12%), and a minor advance in the SED (51.6%), indicate a complex interplay of factors.  
325 Furthermore, regions with significantly prolonged SCD and notably delayed SED substantially  
326 outnumber those with significantly reduced SCD and notably advanced SED, respectively.

327 Comparative studies in snow phenology, such as those by C. Notarnicola (2020), reveal that  
328 approximately 78% of mountainous regions globally experienced reductions in snow, with  
329 durations shortening by 43 days and area decreasing by 13%. Only a few regions exhibited



330 positive changes. Conversely, Wang et al. (2017) found no significant decrease in snow over the  
331 last 15 years on the Tibetan Plateau, utilizing MODIS data from 2000 to 2015. This aligns with the  
332 observations in the Qilian Mountains, a significant mountain system on the northeastern edge of  
333 the Tibetan Plateau, a generally positive trend in snow surface and phenology was identified, with  
334 high spatial heterogeneity in snow phenology trends. This reinforces the conclusions drawn in the  
335 current study, emphasizing the nuanced impacts of climate dynamics on snow.



336

337 **Fig. 7 Conceptual diagram of potential driving mechanisms of snow change in the Qilian Mountains**

#### 338 **4.2 Impacts of the Qilian Mountains snow on hydrology and ecosystems**

339 The hydrological dynamics within the Qilian Mountains are profoundly influenced by the  
340 snow's intra-annual accumulation and ablation processes, where snowmelt acts as a critical source  
341 of freshwater for tens of millions of people. Varied watershed studies within this region illustrate  
342 that the contribution of meltwater from snowmelt to total runoff significantly fluctuates across  
343 different watersheds. For instance, the Yarlung Tsangpo River and its tributaries derive  
344 approximately 9.7% of their runoff from snow and ice melt. Conversely, in the source area of the  
345 Yangtze River, characterized by high elevation and the prevalent development of glaciers and  
346 permafrost, precipitation stands as the predominant source of recharge. Here, snow, largely  
347 occurring as patchy distributions, along with ice meltwater, constitutes about 13.6% of the runoff.  
348 Furthermore, the proportion of snow and ice meltwater in the runoff of the upper reaches of the  
349 Black River in the Qilian Mountains accounts for roughly 16.1%, a figure that underscores the  
350 variance in snow development and hydrothermal conditions across different watersheds.

351 Although the overall contribution of snowmelt to total runoff might appear modest, its  
352 significance cannot be overstated. The peak of snowmelt typically aligns with late spring, a critical  
353 juncture for agricultural irrigation needs and the growth phase of natural vegetation. Thus,



354 snowmelt plays an indispensable role in the recharge of soil moisture and river runoff. Climate  
355 change introduces notable impacts on the hydrological processes associated with the snow in the  
356 Qilian Mountains, with responses varying across different regional watersheds. For example, an  
357 increase in snowfall within the eastern watersheds is anticipated to moderate the rate of runoff  
358 increase and delay the commencement of peak runoff periods. While snow presently constitutes the  
359 primary driver of river flow in these regions, future climate scenarios predict an increase in the  
360 frequency and intensity of rainfall events, potentially reducing the relative contribution of  
361 snowfall. This shift highlights the evolving dynamics of hydrological processes in the face of  
362 climate change, underscoring the need for adaptive water resource management strategies in the  
363 Qilian Mountains.

#### 364 **4.3 Impacts of the Qilian Mountains snow on ecosystems**

365 Snow significantly influences the climate system, water resource management, and  
366 ecological diversity. It plays a crucial role in regulating ground-level energy absorption,  
367 maintaining the water balance, influencing surface temperature, and facilitating gas exchange  
368 processes in vegetation. In the Qilian Mountains, variations in snow are directly associated with  
369 agricultural productivity and the preservation of biodiversity. snow critically impacts water cycle  
370 dynamics by regulating the availability of water resources across different seasons: it limits water  
371 resources during the cold season and ensures their abundance during the snowmelt period.  
372 Additionally, snow indirectly influences the energy balance at the ground level. The potential heat  
373 flow to the atmosphere is reduced, and the heat flow to the soil is altered, as snow impedes  
374 groundwater recharge by capturing precipitation and meltwater, and transports substantial volumes  
375 of water downstream during melting. Hence, future shifts in snow are poised to significantly  
376 transform the hydrology of the Qilian Mountains. Decreases in snow, earlier melting periods, and  
377 heightened rates of evapotranspiration and sublimation are likely to affect both seasonal and  
378 long-term water and ice storage. Consequently, many areas of temporary and semi-permanent  
379 snow in the Qilian Mountains may experience reduction or complete disappearance, impacting the  
380 SCE on glaciers and adversely influencing their mass balance. The diminution of snow is also  
381 expected to lead to the desiccation of numerous patchy wetlands and deterioration of conditions in  
382 other wetlands within the region, which rely on late-season snow to sustain their wet state.

383 The distribution of snow significantly influences the types of vegetation present in exposed





384 areas, where low-lying plant forms have evolved to withstand the dual stresses of wind erosion  
385 and summer drought, demonstrating remarkable adaptation to extreme environmental conditions.  
386 Moreover, species of vegetation that remain covered by snow during winter exhibit the capacity to  
387 sprout rapidly following snowmelt — capitalizing on the brief growing season in the Qilian  
388 Mountains to optimize their growth and reproductive success. Consequently, alterations in snow  
389 conditions play a pivotal role in determining vegetation distribution, biodiversity, and ecosystem  
390 productivity. In the Qilian Mountains, the intricate and vital interdependence between plant and  
391 animal communities and snow conditions cannot be overstated. Changes in snow influence not  
392 only the survival of specific species, such as the rock sheep and snow grouse, but also impact the  
393 migratory patterns and reproductive behaviors of species that migrate seasonally. For species  
394 embarking on long-distance migrations, snow conditions at breeding sites during spring are  
395 particularly critical. Therefore, modifications in snow dynamics could lead to significant  
396 repercussions for the ecosystem's structure and function, impacting the composition and  
397 distribution of species communities within the region.

#### 398 **4.4 Uncertainties and limitations of the study results**

399 The findings of this study are derived from the MODIS snow product and, thus, inherit the  
400 limitations associated with it. Despite its lower spatial resolution of 500 meters, inferior to the  
401 likes of Landsat and Sentinel-2 data, MODIS remains the most viable data source for monitoring  
402 the spatial and temporal dynamics of snow on a large scale over extended time series. Nonetheless,  
403 the utility of MODIS snow accumulation products is significantly impeded by cloud  
404 contamination. Over the past decades, numerous de-clouding techniques for snow accumulation  
405 products have been proposed (Li et al. 2019). These include: (1) spatial approaches such as spatial  
406 filtering, snowline mapping methods, and locally weighted logistic regression (Gafurov and  
407 Bardossy, 2009; Lopez-Burgos et al., 2013; Parajka et al., 2010); (2) temporal strategies,  
408 combining Terra and Aqua data, and implementing temporal filters that involve adjacent time  
409 inference, multi-day combinations, seasonal filters, and temporal interpolation using mathematical  
410 functions (Dozier et al., 2008; Gafurov and Bardossy, 2009; Parajka and Blöschl, 2008; Paudel  
411 and Andersen, 2011; Tang et al., 2013); (3) spatio-temporal combination methods (Dariane et al.,  
412 2017; Jing et al., 2019; Li et al., 2017); (4) multi-source fusion methods incorporating optical,  
413 microwave, and station observations (Brown et al., 2010; Gafurov et al., 2015; Huang et al., 2016;



414 Liang et al., 2008).

415 The de-clouded snow accumulation product employed in this study is developed through a  
416 methodology that utilizes high-resolution Landsat TM data as the baseline truth value. It is  
417 augmented with MODIS land cover classification products to calibrate index thresholds for  
418 discriminating snow accumulation under forested and non-forested categories. These are then  
419 integrated with MODIS snow accumulation inversion algorithms to generate the primary dataset,  
420 which undergoes further refinement through Hidden Markov de-clouding and snow-depth data  
421 interpolation methods to produce a cloud-free, daily snow area product for the study region.  
422 However, the employment of spatio-temporal interpolation algorithms and other void-filling  
423 techniques may introduce discrepancies due to challenges like prolonged cloud cover and the  
424 complexity of terrain and landcover. When quantifying snow area, variations arise from the use of  
425 different snow products. Consequently, the calculated snow area over the Qilian Mountains  
426 displays slight deviations from values reported by other researchers. However, the overall trend  
427 remains consistent, and the margin of error falls within an acceptable range.

## 428 **5. Conclusion**

429 This study systematically examined the spatial and temporal dynamics of snow accumulation  
430 in the Qilian Mountains, analyzing trends in snow area and snow phenology from high-resolution  
431 MODIS snow products and snow phenology products spanning 2000-2020, as well as multi-year  
432 SD trends derived from long-term SD data covering 1980-2019. The findings revealed several key  
433 trends:

434 (1) The overall snow in the Qilian Mountains exhibited a fluctuating upward trend from the  
435 1980/1981 season until a peak depth of 3.577 cm was reached in the 2011/2012 season, after  
436 which a significant decline was observed. The highest SD were located in the central and western  
437 regions of the Qilian Mountains, with the central region experiencing the most pronounced  
438 reduction in SD, exhibiting an interannual variability of less than -0.04 cm/year.

439 (2) The snow within the Qilian Mountains demonstrated an overall increasing trend, with  
440 peak snow typically occurring in November and January. The study also noted strong seasonal  
441 fluctuations in SCE. The seasons of 2002/2003 and 2019/2020 experienced no snow-free days,  
442 recording the lowest snow areas of 82.4km<sup>2</sup> and 206.01km<sup>2</sup>, respectively.

443 (3) In terms of snow phenology, the majority of the Qilian Mountains area experienced an



444 advancement in the SOD, which accounted for 68.94% of the total area studied. There was a  
445 marked increase in areas where the SCD was significantly prolonged compared to those where it  
446 was significantly reduced. Similarly, areas where the SED was significantly advanced  
447 outnumbered those where it was significantly delayed. The overall increase in precipitation was  
448 identified as the main driver of these trends, while rising temperatures were pinpointed as the  
449 primary cause for the decrease in SD across the Qilian Mountains region.

#### 450 **CRedit authorship contribution statement**

451 **Enwei Huang** : Conceptualization, Methodology, Formal analysis, Writing – review &  
452 editing, Supervision, Project administration, Funding acquisition. **Guofeng Zhu**: Methodology,  
453 Validation, Formal analysis, Investigation, Data curation, Software, Writing original draft. **Yuhao**  
454 **Wang**: Writing – review & editing. **Gaojia Meng**: Writing – review & editing. **Ling Zhao**:  
455 Writing – review & editing. **Xuan Zhang**: Writing – review & editing. **Xiaoyu Qi** : Writing –  
456 review & editing. **Qinqin Wang**: Writing – review & editing. **Yinying Jiao**: Writing – review &  
457 editing. **Jiawei Liu** : Writing – review & editing. **Siyu Lu** : Writing – review & editing. **Longhu**  
458 **Chen**: Writing – review & editing. **Rui Li**: Data curation, Software.

#### 459 **Declaration of Competing Interest**

460 The authors declare that they have no known competing financial interests or personal  
461 relationships that could have appeared to influence the work reported in this paper.

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