





## 18 **Abstract**

19 Non-methane hydrocarbons (NMHCs) in the marine atmosphere have been  
20 extensively studied due to their important roles in regulating the atmospheric chemistry  
21 and climate. However, very little is known about the distribution and sources of  
22 NMHCs in the lower atmosphere over the marginal seas of China. Herein, we  
23 characterized the atmospheric NMHCs (C2-C5) in both the coastal cities and marginal  
24 seas of China in spring 2021, with a focus on identifying the sources of NMHCs in the  
25 coastal atmosphere. The NMHCs in urban atmospheres, especially for alkanes, were  
26 significantly higher compared to that in marine atmosphere, suggesting that terrestrial  
27 NMHCs may serve as an important reservoir/source of the marine atmosphere. A  
28 significant correlation was observed between the alkane concentrations and the  
29 distances from sampling sites to the nearest land or retention of air mass over land,  
30 indicating that alkanes in the marine atmosphere are largely influenced by terrestrial  
31 inputs through air-mass transport. For alkenes, a greater impact from oceanic emissions  
32 was determined due to the lower terrestrial concentrations, short atmospheric lifetime,  
33 and substantial sea-to-air fluxes of alkenes compared to alkanes ( $489 \pm 454$  vs  $129 \pm$   
34  $106 \text{ nmol m}^{-2} \text{ d}^{-1}$ ). As suggested by the positive matrix factorization, terrestrial inputs  
35 contributed to 89 % of alkanes and 69.6 % of alkenes in Chinese marginal seas,  
36 subsequently contributing to 84 % of the ozone formation potential associated with C2-  
37 C5 NMHCs. These findings underscore the significance of terrestrial outflow in  
38 controlling the distribution and composition of atmospheric NMHCs in the marginal  
39 seas of China.

40 **Keywords:** non-methane hydrocarbons, oceanic ventilation, terrestrial outflow, source  
41 apportionment



## 42 **1 Introduction**

43 Non-methane hydrocarbons (NMHCs), a significant subset of volatile organic  
44 compounds (VOCs), are acknowledged as key precursors to tropospheric ozone  
45 formation (Houweling et al., 1998; Solomon et al., 2005) and second organic aerosol  
46 (SOA) generation (Hallquist et al., 2009; Wu and Xie, 2018), playing a pivotal role in  
47 atmospheric chemistry. The presence and activity of NMHCs in the troposphere have  
48 far-reaching implications, not only influencing the dynamics of ozone and SOA  
49 formation but also significantly impacting air quality. These compounds are intricately  
50 linked to heightened human health risks as well as possessing indirect yet profound  
51 effects on the broader climate system through their interactions with various  
52 atmospheric processes (Yuan et al., 2018).

53 The emission of NMHCs into the atmosphere stems from an array of natural and  
54 anthropogenic processes. Oceanic sources of NMHCs predominantly entail the  
55 biogenic production of phytoplankton and photochemical degradation of dissolved  
56 organic matter (DOM) (Bonsang et al., 1992; Li et al., 2019; Riemer et al., 2000; Sahu  
57 et al., 2010). However, they are minimal when compared to terrestrial inputs. Previous  
58 estimates driven from a global VOCs emissions model formulated by Guenther et al.  
59 (1995) assign terrestrial sources a cumulative emission of 1141 Tg C year<sup>-1</sup>, with  
60 oceanic emissions at merely 5 Tg C year<sup>-1</sup>. A substantial amount of NMHCs originating  
61 from terrestrial sources (e.g., vehicular emissions, biomass combustion, industrial  
62 activities, and continental vegetation emissions) can be transported into the offshore  
63 atmosphere via air mass conveyance (Wang et al., 2005; Kato et al., 2007; Song et al.,  
64 2020). Subsequently, these supplementary terrestrial NMHCs will play a pivotal role in  
65 shaping the chemical composition of the offshore atmosphere and influencing local



66 environmental dynamics. Hence, to further understand the characteristics, variation,  
67 and origins of NMHCs in the offshore atmosphere, it is imperative to scrutinize oceanic  
68 emissions and meanwhile, it is necessary to figure out the effect of terrestrial outflow  
69 on nearshore NMHCs.

70 The Yellow Sea and the East China Sea are important parts of Chinese marginal seas,  
71 situated along the eastern coast of China where it is densely populated and has intensive  
72 industries. The rapid pace of Chinese development has seen a notable escalation in  
73 anthropogenic NMHCs emissions over recent decades (He et al., 2019). Presently,  
74 excessive NMHCs emissions and severe ozone pollution have emerged as urgent  
75 environmental challenges in China, particularly in highly urbanized and industrialized  
76 areas along the eastern coast (Liu et al., 2016; Zhang et al., 2018). The seasonal cycle  
77 of the Asian monsoon and diurnal fluctuations of sea-land breezes can facilitate the  
78 transport of terrestrial pollution to the marine atmosphere (Ding et al., 2004; Wang et  
79 al., 2003; Talbot et al., 2003; Russo et al., 2003). Additionally, eutrophication in coastal  
80 regions fosters the proliferation of phytoplankton, potentially augmenting the natural  
81 emissions of NMHCs. Consequently, conducting atmospheric investigations in the  
82 coastal region of eastern China is effective in revealing the potential effects of land-sea  
83 interactions on offshore atmospheric NMHCs.

84 In the spring of 2021, atmospheric samples were systematically collected from both  
85 coastal cities and marginal seas of China, providing representative insights into the  
86 characteristics of NMHCs (C2-C5) and facilitating discussion on the interplay between  
87 ocean emission and terrestrial outflow concerning atmospheric NMHCs. Ultimately,  
88 the contributions of diverse sources to NMHCs were quantified using the positive  
89 matrix factorization (PMF) model, with assistance from indications provided by other  
90 typical gases, mainly dimethyl sulfur (DMS), volatile halogenated compounds (VHCs),



91 and monocyclic aromatics.

## 92 **2 Methods**

### 93 **2.1 Samples collection**

94 The urban samples were collected from eight coastal cities in China from March 27  
95 to April 1, 2021 (Fig. 1). Air samples were collected using fused silica-lined canisters  
96 (2.5 L), which were cleaned three times via a Canister Cleaning System (2101DS,  
97 Nutech) and were pumped into a negative pressure state before sampling. The sampling  
98 sites were selected at the top of high buildings to minimize contamination of particular  
99 point sources. Air samples were collected at 09:00 and 21:00 local time (UTC+8) aimed  
100 to represent the urban atmospheric conditions during the daytime and night, respectively.  
101 Note that night samples in Xiamen and Qinzhou were missing.

102 Oceanic air samples were collected aboard RV “*Dong Fang Hong 3*” during the  
103 voyage in the Yellow Sea and the East China Sea from April 17 to May 2, 2021.  
104 Nineteen oceanic air samples were collected on the top deck facing the wind when the  
105 ship was about to arrive at the station and started to slow down. Seawater samples were  
106 collected via prewashed Niskin bottles (12 L) incorporated into the Conductivity-  
107 Temperature-Depth Sensor Rosette (Seabird 911). Sampling details for urban and  
108 marine samples are shown in Table S4 and Table S5, respectively.

### 109 **2.2 Analysis of air samples**

110 All air samples were processed immediately after being brought back to the  
111 laboratory, using an Atmospheric Pre-concentrator System (8900DS, Nutech) coupled  
112 with a GC-MSD system (GC-7890A, MSD-5975, Agilent). The pretreatments of air



113 samples were as follows. First, the Atmospheric Pre-concentrator System was baked for  
114 10 min to clean the interior instrument. Then trap 1 was cooled to  $-170\text{ }^{\circ}\text{C}$  using liquid  
115  $\text{N}_2$  and a 300 mL air sample was pumped from the canister into trap 1 for the initial  
116 concentration of the target compounds, while  $\text{N}_2$  and  $\text{O}_2$  escaped due to their lower  
117 boiling points. After trap 2 was cooled to  $-50\text{ }^{\circ}\text{C}$ , trap 1 was heated to  $30\text{ }^{\circ}\text{C}$  to transfer  
118 the target compounds from trap 1 to trap 2. Moisture and  $\text{CO}_2$  were removed in the  
119 second concentration. Then, trap 2 was warmed up to transfer the target compounds  
120 into the last trap for cryofocusing ( $-175\text{ }^{\circ}\text{C}$ ). Finally, the last trap was instantaneously  
121 heated to  $200\text{ }^{\circ}\text{C}$  via gas bath heating, and the target compounds were delivered into  
122 the GC-MSD system by ultra-pure He.

123 For the parameter settings of GC-MSD, the temperature of the inlet, quadrupole, and  
124 ionization source was  $150\text{ }^{\circ}\text{C}$ ,  $150\text{ }^{\circ}\text{C}$ , and  $230\text{ }^{\circ}\text{C}$ , respectively. The inlet was set to  
125 split mode with a ratio of 10:1. The flow rate of carrier gas (He) was set to  $1.5\text{ mL min}^{-1}$   
126 in the instant flow mode. Specific columns were selected to separate the NMHCs (Rt-  
127 Alumina BOND/KCl, Restek), monocyclic aromatics (DB-624, Agilent), DMS  
128 (CP7529, Agilent), and VHCs (DB-624, Agilent). Gas standards of NMHCs in  $\text{N}_2$   
129 (Linde Gases, Germany) were diluted to 0.1-1 ppb for identification and calibration.  
130 Details of temperature programming and detector parameters can be seen in Zou et al.  
131 (2021) and Li et al. (2019). The precision and detection limits for the trace gases in the  
132 present study were 1-7 % and 0.03-20.0 ppt, respectively (Table S1). Note that DMS  
133 and VHCs data in marine atmospheric samples were graciously provided by colleagues  
134 in the same laboratory. These data were only used as supporting information in the  
135 interpretation (e.g., correlation analysis) of our core dataset in this paper.



### 136 **2.3 Analysis of seawater samples**

137 C2-C5 NMHCs in seawater were measured immediately on board using a purge and  
138 trap system coupled with the gas chromatography equipped with a flame ionization  
139 detector (GC-FID, 7890B, Agilent). The purge and trap system was improved based on  
140 a previously self-designed device described by Li et al. (2019). Briefly, seawater was  
141 collected using a customized glass sampler (500 mL) and was connected to the inlet of  
142 the system. Then seawater was transformed into the extraction cell under the pressure  
143 of pure N<sub>2</sub> and was purged with pure N<sub>2</sub> bubble flow (250 mL min<sup>-1</sup>). The moisture of  
144 the carrier gas condensed in a thin glass tube that was placed in a cold chamber (4-6 °C)  
145 and the carbon dioxide was absorbed by the glass tube filled with Ascarite II (Merck).  
146 The targets were concentrated in a passivated stainless-steel tube immersed in liquid  
147 nitrogen for 26 mins. Then, the steel tube was heated by boiling water and immediately,  
148 the six-way valve was turned for the inlet situation. The concentrated target compounds  
149 were transferred into the Rt-Alumina BOND/KCl capillary column for separation and  
150 were determined by the FID. The parameters of the inlet, oven, and detector are shown  
151 in Table S2. The gas standard (Linde Gases, Germany) was diluted with ultra-pure N<sub>2</sub>  
152 to 10 ppb for identification and quantification. The instrumental blank was made to  
153 guarantee data reliability. The precision and detection limits were 3-6 % and 0.5-1.0  
154 pmol L<sup>-1</sup> (Table S3).

### 155 **2.4 Calculation of sea-to-air flux**

156 The sea-to-air flux of each NMHCs ( $F$ , nmol m<sup>-2</sup> d<sup>-1</sup>) was calculated using Eq. (1):

$$157 \quad F = k \times (C_w - C_a \times H) \quad (1)$$

158 where  $k$  (m s<sup>-1</sup>) is the gas transfer velocity described by Eq. (2);  $H$  is Henry's law  
159 constant;  $C_w$  (pmol L<sup>-1</sup>) and  $C_a$  (ppb) are concentrations of each NMHCs in the



160 surface seawater (5 m depth) and atmosphere, respectively.

$$161 \quad k = 0.31 \times u^2 \times \left(\frac{Sc}{660}\right)^{-0.5} \quad (2)$$

162 where  $u$  ( $\text{m s}^{-1}$ ) is the wind velocity at 10 m.  $Sc$  is the Schmidt number and is defined

163 as  $Sc = \nu/D$ .  $\nu$  was the kinematic viscosity of seawater calculated by Eq. (3)

164 (Wanninkhof, 1992).  $D$  is the gas diffusion coefficient related to temperature described

165 by Eq. (4) (Wilke and Chang, 1955).

$$166 \quad \nu = 1.052 + 1.300 \times 10^{-3} \times t + 5.000 \times 10^{-6} \times t^2 + 5.000 \times 10^{-7} \times t^3 \quad (3)$$

$$167 \quad D = \frac{7.4 \times 10^{-8} (q \times M_b)^{0.5} \times T}{n_b \times V_a^{0.6}} \quad (4)$$

168 where  $t$  ( $^{\circ}\text{C}$ ) is the degree Celsius of seawater,  $q$  is the association factor of water,

169  $M_b$  ( $\text{g mol}^{-1}$ ) is the molar weight of water,  $T$  (K) is the degree Kelvin of seawater,  $n_b$

170 is the dynamic viscosity of seawater and  $V_a$  is the molar volume at the boiling point.

## 171 **2.5 Normalized concentrations and lifetime-weighted concentrations of NMHCs.**

172 To effectively compare the NMHCs variation with respect to the distance from the

173 sampling sites to the land (like Fig. 2d, f), we calculated the normalized concentration

174 for each NMHCs ( $C_{Nor-i}$ ) using Eq. (5).

$$175 \quad C_{Nor-i} = \frac{C_i}{C_{max-i}} \quad (5)$$

176 where  $C_i$  is the concentration of gas  $i$  and  $C_{max-i}$  is the maximum of gas  $i$ .

177 A novel approach was employed to analyzed the correlation between the

178 concentrations of various NMHCs and their sea-to-air fluxes. Concentrations were

179 weighted according to the respective atmospheric  $\bullet\text{OH}$  lifetime of each NMHCs. This

180 was achieved by dividing the concentration of each NMHCs by its corresponding

181 atmospheric  $\bullet\text{OH}$  lifetime, yield a “lifetime-weighted concentration” for each NMHCs

182 ( $C_{life-i}$ ) (Eq. 6). This method provides a more nuanced understanding of the impact of





183 oceanic emission on NMHCs, taking into account not only their abundance but also  
184 their residence in the atmosphere.

$$185 \quad C_{life-i} = \frac{C_i}{\tau_i} \quad (6)$$

186 where  $C_i$  is the atmospheric concentration of gas  $i$ ,  $\tau_i$  is the  $\bullet$ OH lifetime of gas  $i$ .  
187 Approximate atmospheric lifetime of each NMHCs was calculated assuming an  
188 average [ $\bullet$ OH] of  $6 \times 10^5$  molecules  $\text{cm}^{-3}$  within 24 h at 288 K (Jobson et al., 1999), with  
189 specific data listed in Table 1.

## 190 2.6 Calculation of retention of air mass over land

191 To identify whether an air mass was mainly from terrestrial or oceanic regions, the  
192 retention ratio of the air mass over land ( $R_L$ ) was calculated by Eq. (7).

$$193 \quad R_L = \frac{\sum_{n=1}^{N_{land}} e^{-\frac{t_n}{48}}}{\sum_{n=1}^{N_{total}} e^{-\frac{t_n}{48}}} \quad (7)$$

194 Where  $N_{total}$  is the total number of trajectory endpoints (downloaded from NOAA  
195 Air Resources Laboratory HYSPLIT trajectory model <https://www.arl.noaa.gov/>).  
196  $N_{land}$  is the total number of trajectory endpoints located over land, while  $t_n$  is the  
197 backward tracking time with the unit of hour and  $e^{-\frac{t_n}{48}}$  is the weighting factor related  
198 to tracking time as the diffusion of air mass takes place along the transport path than in  
199 the nearby regions. As a result, the larger  $R_L$  value indicates that the air mass is more  
200 influenced by terrestrial transport and its source is more likely to be on land. Similar  
201 methods have been used to calculate the average residence time of sampled air masses  
202 in the Arctic (Willis et al., 2017) and identify the percentage of time spent by trajectories  
203 over different surface types in the Antarctic (Decesari et al., 2020).  $R_L$  values were  
204 calculated by three different time-scale trajectories (48h, 72h, and 96h). The mean  $R_L$   
205 ( $n = 3$ ) was finally applied to analyze the terrestrial influence on oceanic NMHCs,



206 mitigating the uncertainty caused by the trajectory with different time-scales.

## 207 **2.7 Application of the PMF model**

208 PMF model introduced in detail in the study of (Paatero and Tapper, 1994) was  
209 applied to analyze the data of atmospheric NMHCs in the Yellow Sea and South China  
210 Sea. Based on a matrix consisting of the concentrations of diverse chemical species, the  
211 objective of PMF is to determine the number of NMHCs source factors, the chemical  
212 composition profile of each factor, and the contribution of each factor to species. In the  
213 application of the PMF model, the significance of missing data in the matrix was  
214 decreased by using the species median. The uncertainty for normal data was estimated  
215 as 20 % of the NMHCs concentrations because the analytical uncertainty was not  
216 available (Buzcu and Fraser, 2006). In this analysis, the model ran 20 times and we  
217 selected the result with the minimum “Q value”. Besides, approximately 94 % of the  
218 scaled residuals given by PMF ranged from -3 to 3 (Fig. S1), suggesting a reasonable  
219 fit of the model result.

## 220 **3 Results and discussion**

### 221 **3.1 Atmospheric concentrations of NMHCs in coastal cities and coastal** 222 **seas of China**

223 To clarify, NMHCs determined in this study were separated into two groups for  
224 further discussion based on their distinctly different atmospheric reactivity and lifetimes:  
225 alkanes (long lifetime, 8.2-78 d) and alkenes (short lifetime, 0.19-2.3 d). In urban  
226 atmosphere (n = 14), the mean (range) concentration of ethane, propane, i-butane, and  
227 n-butane was 2.26 (0.277-5.72), 2.95 (0.149-20.1), 2.57 (BD-27.6), and 3.29 (0.018-



228 30.2) ppb, respectively (Table 1). Alkanes combined accounted for ~76 %-99 % of total  
229 NMHCs measured in this study, which agrees with previous studies reporting alkanes  
230 as the dominant NMHCs in the urban atmosphere of China e.g., 43.7 % (Song et al.,  
231 2007), and > 50 % (Li et al., 2015). For alkene species in the urban atmosphere ( $n =$   
232 14), the mean (range) of ethylene, propylene, and isoprene was 0.180 (0.035-0.390),  
233 0.036 (BD-0.129), and 0.046 (0.006-0.250) ppb, respectively.

234 Similarly, alkanes were also dominant components in the marine atmosphere,  
235 accounting for ~86 %-95 % of NMHCs. In the marine atmosphere ( $n = 19$ ), the mean  
236 (range) concentration of ethane, propane, i-butane, n-butane, ethylene, propylene, and  
237 isoprene was 1.24 (0.686-1.72), 0.822 (0.226-1.79), 0.283 (BD-1.17), 0.256 (0.025-  
238 0.694), 0.151 (0.028-0.295), 0.033 (0.022-0.060), and 0.008 (BD-0.043) ppb,  
239 respectively. These values were comparable to those reported in the Bengal Bay (Sahu  
240 et al., 2011) and the Northwest Pacific Ocean (Li et al., 2019) (Table S6). Alkanes in  
241 the urban atmosphere were on average more than four times higher than those in the  
242 marine atmosphere, while no significant difference was observed for concentrations of  
243 alkenes between urban and marine air ( $t = 2.224$ ,  $p = 0.156$ ) (Fig. 2a, b). This suggests  
244 that the terrestrial alkanes may potentially serve as a reservoir/source of the alkanes in  
245 the marine atmosphere through the transport of terrestrial air mass.

### 246 **3.2 Atmospheric NMHCs variability vs. estimated lifetime**

247 The standard deviation of the natural logarithm of the NMHCs mixing ratios ( $S_{\ln x}$ )  
248 was established to correlate to their •OH lifetime ( $\tau$ ) in the atmosphere following an  
249 exponential function of  $S_{\ln x} = A\tau^{-b}$  (Jobson et al., 1998), where  $A$  and  $b$  are fitting  
250 parameters. A  $b$  value approaching zero suggests that the NMHCs variability is  
251 primarily controlled by local emission fluctuations while a  $b$  value of 1 indicates the



252 minimal impact of local emissions, with the variability predominantly controlled by the  
253 extent of photochemical reactions.

254 Employing the analytical framework in Jobson et al. (1998), we analyzed our  
255 atmospheric NMHCs data from urban areas and the Chinese marginal seas. The derived  
256  $b$  value for urban areas was 0.05 (Fig. 3a), suggesting that atmospheric NMHCs in  
257 coastal cities were mainly controlled by local emissions. In the marine atmosphere, the  
258  $b$  value was 0.26 (Fig. 3b) which was comparable to values reported for Gosan (0.30)  
259 (Wong et al., 2007) and continental outflow from southern China (0.31) (Wang et al.,  
260 2005), but it was significantly lower than the values for Ogasawara (0.43) (Kato et al.,  
261 2004), the Northwest Indian Ocean (0.40) (Warneke and De Gouw, 2001), and the South  
262 China Sea (0.42) (Wang et al., 2005). The  $b$  value of 0.26 in the atmosphere over the  
263 Chinese marginal suggests that the NMHCs composition in nearshore atmosphere is  
264 influenced both by local oceanic emissions and the remote sources from the continent.  
265 As sites closer to the source position tend to have lower  $b$  values, the Yellow Sea and  
266 the East China Sea experience a more pronounced influence from terrestrial pollution  
267 sources compared to Ogasawara, the South China Sea, and the Northwest Indian Ocean.

### 268 **3.3 Terrestrial influence on marine atmospheric NMHCs variation**

269 Given the discernible impact of terrestrial input on the spatial distributions and  
270 variabilities of marine atmospheric NMHCs, we further elucidated the role of terrestrial  
271 outflow in shaping marine atmospheric NMHCs levels. This examination focused on  
272 three key factors: distance from the sampling site to the land, retention of air mass over  
273 land, and transport time of air mass.



274 **Distance from the sampling site to the land**

275 The distances from the oceanic sampling sites to the nearest land spanned from 13.9  
276 to 331 km, with an average of 123 km (Table S9). Significant correlations were  
277 observed between the distances and concentrations of ethane ( $r = -0.553$ ,  $n = 19$ ,  $p =$   
278  $0.014$ ), propane ( $r = -0.605$ ,  $n = 19$ ,  $p = 0.006$ ), i-butane ( $r = -0.513$ ,  $n = 19$ ,  $p = 0.025$ ),  
279 and n-butane ( $r = -0.573$ ,  $n = 19$ ,  $p = 0.010$ ). When plotted against the distances, the  
280 concentrations of alkanes combined decreased with the increasing distance (Fig. 2c),  
281 and different species exhibited distinctly specific decreasing rates (Fig. 2d). Since the  
282 concentrations between different NMHCs species varied considerably, the normalized  
283 concentrations were employed to fit an attenuation equation ( $y = Ae^{-tx} + y_0$ ) for each  
284 species. As evident in Fig. 2d, the attenuation coefficients for ethane, propane, i-butane,  
285 and n-butane were 0.003, 0.030, 0.031, and 0.022, respectively. These coefficients were  
286 correlated with their atmospheric reactivities. Species with lower reactivity and longer  
287 lifetimes, such as ethane (with a lifetime of 78 d), have the lowest attenuation  
288 coefficient. This implies that long-lifetime species could be affected by the terrestrial  
289 input even at a more remote marine site. Terrestrial influences on propane, i-butane,  
290 and n-butane were discernible only in areas much closer to land, as their concentrations  
291 stabilized at low values beyond a distance of around 100 km (Fig. 2d).

292 **Retention of air mass over land**

293 A larger retention of air mass over land ( $R_L$ ) has previously been suggested to serve  
294 as an indicator of a greater terrestrial influence (Zhou et al., 2021). To mitigate the  
295 uncertainty derived from varying time-scale trajectories, we calculated the  $R_{L-mean}$   
296 based on 48, 72, and 96-hour backward trajectories.  $R_{L-mean}$  ranged from 0.10 to 0.96  
297 (Table S9). When plotted against  $R_{L-mean}$ , a linear relationship was observed between



298 the concentrations of NMHCs combined and  $R_{L-mean}$ , with a slope of 2.51 (Fig. 4a).  
299 A statistically significant correlation ( $r = 0.599$ ,  $n = 19$ ,  $p = 0.007$ ) was observed when  
300 only plotting alkanes with  $R_{L-mean}$ . However, the correlation between alkenes and  
301  $R_{L-mean}$  was statistically insignificant ( $r = 0.248$ ,  $n = 19$ ,  $p = 0.306$ ).

### 302 **Transport time of air mass**

303 The transport time of air mass was estimated as the interval from the last point of the  
304 trajectory contacting the continent to the moment when the air mass reached the  
305 sampling location, as detailed by Kato et al. (2001). These times ranged from 4 to 81 h,  
306 with an average of 30 h (Table S9). A shorter air mass transport time signifies a stronger  
307 terrestrial influence, as NMHCs within the air mass undergo further oxidation and  
308 dispersion over time. Total NMHCs concentrations exhibited a significant decrease  
309 with the increase of air mass transport time, characterized by a slope of -0.04 (Fig. 4d).  
310 Alkanes displayed a steeper decline, indicated by a slope of -0.0079 (Fig. 4e) compared  
311 to alkenes (-0.0038, Fig. 4f). However, similar to the analysis of  $R_L$ , the correlation  
312 between the air mass transport time and alkenes was statistically insignificant ( $r = 0.248$ ,  
313  $n = 19$ ,  $p = 0.306$ ).

314 Overall, the analysis above suggests that the terrestrial input plays an important role  
315 in driving the variability observed for the atmospheric NMHCs over the marginal seas  
316 of China. In particular, a stronger terrestrial impact was determined for the alkanes  
317 based on the larger slopes from linear regression analysis and the significant  
318 correlations with terrestrial indicators. In contrast, no discernible trend was found for  
319 alkenes when plotting their concentrations against the distance from sampling sites to  
320 the coastline (Fig. 2e, f). There was no significant correlation between alkenes and  $R_L$   
321 or air mass transport time. Therefore, the variability of alkenes in the coastal



322 atmosphere seems to be weakly impacted by the terrestrial sources when compared to  
323 alkanes. We attribute this to two main factors. First, the mean concentration of alkenes  
324 in the urban air was only 1.4 times of that in marine air, whereas it was 5.4 times for  
325 alkanes. Alkenes undergo more rapid oxidation due to their higher reactivities compared  
326 to alkanes during air mass transport. Secondly, oceanic ventilation may play a more  
327 substantial role in affecting marine alkenes (discussed in section 3.4).

### 328 **3.4 Oceanic impact on marine atmospheric NMHCs composition**

#### 329 **Sea-to-air fluxes of NMHCs**

330 The mean (range) of sea-to-air fluxes of ethane, propane, i-butane, n-butane, ethylene,  
331 propylene, and isoprene was 44.6 (0.2-118), 41.5 (0.2-157), 31.7 (0.1-146), 10.9 (-0.8-  
332 96.1), 321 (1.7-775), 56.1 (0.2-212), and 112 (0.5-468)  $\text{nmol m}^{-2} \text{d}^{-1}$ , respectively, in  
333 the Yellow Sea and the East China Sea (Table 1). These values were comparable to  
334 those reported in Chinese marginal seas (Wu et al., 2021; Li et al., 2021) and 23-38°N  
335 Atlantic Ocean (Tran et al., 2013), but were larger than those reported values in the  
336 North Sea (Broadgate et al., 1997) and the Northwest Pacific Ocean (Li et al., 2019;  
337 Wu et al., 2023) (Table S10).

338 The mean of sea-to-air fluxes of the total observed NMHCs was  $698 \pm 607 \text{ nmol m}^{-2}$   
339  $\text{d}^{-1}$  in areas within 100 km from the coastline, which was relatively higher than that in  
340 the regions beyond 100 km ( $480 \pm 481 \text{ nmol m}^{-2} \text{d}^{-1}$ ). These elevated fluxes in the sea  
341 areas closer to land could be attributed to the influence of phytoplankton biomass and  
342 chromophoric dissolved organic matter (CDOM). Seawater NMHCs are not only  
343 directly synthesized by phytoplankton (Ratte et al., 1995), but they can also be emitted  
344 through the photochemical degradation of CDOM (Ratte et al., 1993; Lee and Baker,  
345 1992). To substantiate our findings, we analyzed the monthly Chl-*a* concentration and



346 the absorption coefficient at 443 nm of seawater in April 2021 from the remote sensing  
347 dataset from the NASA Ocean Color data service (<https://oceancolor.gsfc.nasa.gov/>)  
348 (Fig. S2). The mean ( $\pm$ SD) of Chl-*a* concentrations was  $2.83\pm 1.17$  and  $1.68\pm 1.44$   $\mu\text{g}$   
349  $\text{L}^{-1}$  in the areas within and beyond 100 km from coastline, respectively. Correspondingly,  
350 the mean ( $\pm$ SD) of seawater absorption coefficients at 443 nm was at  $0.124\pm 0.060$  and  
351  $0.069\pm 0.040$   $\text{m}^{-1}$ , respectively. Hence, the heightened phytoplankton biomass and  
352 enriched photoreaction substrate collectively enhanced both the biological production  
353 and abiotic formation of NMHCs, consequently resulting in a pronounced NMHCs  
354 emission in nearshore regions.

### 355 **Assessing the effect of oceanic emission on NMHCs**

356 Prior to delving into the correlation between oceanic emissions and NMHCs  
357 concentrations, it is imperative to acknowledge the influence of different gases'  
358 reactivity on this relationship. For instance, ethane possesses an atmospheric lifetime  
359 of approximately 78 d at 24 h  $[\bullet\text{OH}]$  concentration of  $6\times 10^5$  molecules  $\text{cm}^{-3}$ . This means  
360 all ethane emitted from the ocean within this period to potentially contribute to the  
361 accumulation of atmospheric ethane. Conversely, isoprene, with a much shorter lifetime  
362 of only 0.2 d, emitted within a very brief window can impact its atmospheric level. Thus,  
363 to mitigate the impact of varying reactivity among the different gas species, we  
364 calculated the life-weighted concentrations of each NMHCs according to their  
365 atmospheric lifetime (introduced in section 2.5). This novel method is more nuanced to  
366 assess the impact of oceanic emission on atmospheric NMHCs, as it acknowledging not  
367 only their abundance but also their residence in the atmosphere.

368 In spite of the elevated oceanic emission of NMHCs within the 100 km from land,  
369 its impact on atmospheric NMHCs composition was comparatively weak displaying a





370 slope of 0.0187 (Fig. 5c), which was lower than the fitted result of the dataset in areas  
371 beyond 100 km from land with a slope of 0.0415 (Fig. 5d). This could be attributed to  
372 the disturbance of terrestrial outflow in nearshore areas, mitigating the direct impact of  
373 oceanic emission on NMHCs. As it extended further from the land, the terrestrial  
374 influence diminished. This, in turn, strengthens the regulatory impact of oceanic  
375 emission on atmospheric NMHCs levels.

376 In addition, the average flux of total alkenes across the entire region was  $163 \pm 221$   
377  $\text{nmol m}^{-2} \text{d}^{-1}$ , which was approximately 5 times higher than that of alkanes ( $32.2 \pm 37.5$   
378  $\text{nmol m}^{-2} \text{d}^{-1}$ ). This substantial discrepancy indicates that alkanes and alkenes are  
379 certainly influenced differently by oceanic emissions. The correlation between the  
380 lifetime-weighted concentrations of alkenes and their fluxes was statistically significant  
381 ( $r = 0.548, n = 57, p < 0.001$ ), while it was insignificant for alkanes ( $r = 0.113, n = 76,$   
382  $p = 0.329$ ). When specific species of alkanes (Fig. 5e) and alkenes (Fig. 5f) were  
383 separately plotted against their sea-to-air fluxes, alkenes exhibited a steeper slope of  
384 0.0072 compared to the slope of 0.0044 for alkanes. This signifies that oceanic emission  
385 has a more significant impact on atmospheric alkenes compared to alkanes, which  
386 verifies our hypothesis as stated at the end of section 3.3.

### 387 **3.5 Identification and apportionment of the sources of marine** 388 **atmospheric NMHCs**

#### 389 **Source identification**

390 Since the chemical compositions are largely controlled by the sources of emissions,  
391 specific ratios of hydrocarbons have been widely employed to identify the sources of  
392 NMHCs (Gilman et al., 2013; Rossabi and Helmig, 2018). For instance, elevated iso-  
393 pentane/n-pentane ratios are indicative of the heavy influence of vehicular emissions



394 (2.2-3.8) and gasoline fuel evaporation (1.8-4.6) (Gentner et al., 2009; Jobson et al.,  
395 2004; Liu et al., 2008; Russo et al., 2010). Conversely, the lower ratios indicate the  
396 importance of tropical forest fires (0.43-0.57) (Andreae and Merlet, 2001; Rossabi and  
397 Helmig, 2018), natural and oil gas operations (0.81-1.1) (Gilman et al., 2013; Swarthout  
398 et al., 2013), and marine vessel exhaust (1.59-1.71) (Bourtsoukidis et al., 2019) in  
399 controlling the chemical composition of NHMCs. In this study, a significant correlation  
400 was observed between i-pentane and n-pentane ( $r = 0.67$ ,  $p < 0.01$ ) (Fig. 6), and the i-  
401 pentane/n-pentane ratio spans a wider range from 0.89 to 2.46, suggesting that the  
402 composition of NMHCs in the marginal seas of China is controlled by multiple sources  
403 e.g., natural and oil gas operations, marine vessel exhaust, vehicular emissions, and  
404 gasoline evaporation.

405 Furthermore, propane, i-butane, and n-butane exhibited strong intercorrelations ( $r =$   
406  $0.52-0.95$ ,  $p < 0.05$ ). They also displayed strong correlations with ethane, i-pentane,  
407 and n-pentane ( $r = 0.55-0.98$ ,  $p < 0.05$ ). These alkanes were recognized as the primary  
408 components of liquid petroleum gases (Blake and Rowland, 1995), extensively utilized  
409 as fuel in taxis, private cars, and public buses in China (Guo et al., 2017; Zhang et al.,  
410 2015). Notably, C3-C5 alkanes also exhibited significant correlations with ethane ( $r =$   
411  $0.55-0.72$ ,  $p < 0.05$ ) and carbon monoxide ( $r = 0.59-0.81$ ,  $p < 0.05$ ), while ethane and  
412 carbon monoxide are acknowledged tracers for fossil fuel or biomass/biofuel  
413 combustion and incomplete combustion, respectively (Lai et al., 2010; Tang et al., 2009;  
414 Parrish et al., 2009). This indicated the contribution of vehicular emissions of liquid  
415 petroleum gases and combustion of fossil fuel or biomass to light alkanes. Additionally,  
416 strong correlations were observed among monocyclic aromatics (benzene, toluene,  
417 ethylbenzene) ( $r = 0.67-0.83$ ,  $p < 0.05$ ). This finding was consistent with recent  
418 emission inventory research identifying monocyclic aromatics as significant



419 constituents of ship exhaust (Xiao et al., 2018b; Wu et al., 2019). As for oceanic  
420 emissions, we have presented the sea-to-air fluxes of NMHCs and discussed the  
421 significant effect of oceanic emissions on NMHCs in Section 3.4. Multiple studies  
422 highlighted that the ocean is one of the important sources of these gases (Kato et al.,  
423 2007; Li et al., 2019; Mallik et al., 2013; Sahu et al., 2010; Rudolph and Johnen, 1990).

#### 424 **Source apportionment**

425 The potential sources of the atmospheric NMHCs and their respective contributions  
426 to each category were determined using the PMF model. Four isolate factors were  
427 extracted according to their composition profiles depicted in Fig. 7a. These factors,  
428 including industrial production, exhaust emission, terrestrial vegetation, and oceanic  
429 ventilation, were identified based on chemical profiles in literature.

430 Propane, i-butane, n-butane, i-pentane, n-pentane, and CFC-11 showed strong  
431 loadings (> 70 %) on factor 1. The presence of propane, butanes, and pentanes suggests  
432 the influence of the refinery activities (Buzcu and Fraser, 2006). Additionally, propane  
433 has been recognized as a characteristic NMHCs derived from natural gas emissions and  
434 butane is indicative of liquefied petroleum gas (LPG) (Guo et al., 2011; Tsai et al., 2006;  
435 Hui et al., 2018; Ho et al., 2009). Moreover, CFC-11 is a typical artificial industrial  
436 product. Subsequently, factor 1 was identified as a factor relating to industrial activities.

437 The profile of factor 2 showed strong loadings of benzene (72 %), toluene (57 %),  
438 and ethylbenzene (64 %), along with moderate impacts of ethylene (34 %) and  
439 propylene (32 %). Benzene emissions are notably associated with vehicle exhaust  
440 (Zhang et al., 2013; Zhang et al., 2016) and considerable fractions of aromatics can be  
441 emitted from ship exhaust during both berthing and cruising (Cooper, 2005; Xiao et al.,  
442 2018a). C2-C4 alkenes could stem from ship emissions in the open ocean (Eyring et al.,



443 2005). Therefore, factor 2 can be potentially assigned as a source of the exhaust  
444 emissions of vehicles and ships.

445 Factor 3 was assigned as oceanic ventilation due to elevated percentages of DMS  
446 (74 %) and  $\text{CHBr}_3$  (53 %), considering the dominant contributions of ocean emission  
447 to DMS (Lana et al., 2011; Lee and Brimblecombe, 2016) and  $\text{CHBr}_3$  (Quack and  
448 Wallace, 2003; Ashfold et al., 2014). Factor 4 was mainly characterized by a high  
449 percentage of isoprene (68 %), an indicator of biogenic emission from terrestrial  
450 vegetation (Guenther et al., 2006; Wu et al., 2016). However, given isoprene's high  
451 reactivity, this factor should be treated cautiously and regarded as a lower limit (Fujita,  
452 2001). Although its short atmospheric lifetime hinders long-range transport, the  
453 minimum air mass transport time from land to the oceanic station was four hours in this  
454 study, implying the potential for terrestrial isoprene to reach the nearshore atmosphere.

455 According to the results of the PMF model analysis, the dominant source of  
456 atmospheric alkanes in the Chinese marginal seas was industrial activities (0.253 ppb,  
457 60.8 %), followed by exhaust emissions (0.095 ppb, 23 %). Contributions from  
458 terrestrial vegetation emission (0.049 ppb, 11 %) and oceanic ventilation (0.021 ppb,  
459 5.2 %) were relatively smaller. Furthermore, exhaust emissions (0.017 ppb, 32.5 %),  
460 industrial activities (0.017 ppb, 31 %), and ocean ventilation (0.016 ppb, 30.4 %)  
461 contribute almost equally to atmospheric alkenes. Collectively, these three factors  
462 constitute the main sources of alkenes (93.8 %), whereas the contribution from  
463 terrestrial vegetation is minimal, at merely 6.2 %. Particularly, the contribution of  
464 terrestrial sources to alkanes (89 %) is greater than that to alkenes (69.6 %), while the  
465 contribution of ocean emission to alkenes (30.4 %) is greater than that to alkanes  
466 (5.2 %). This is consistent with the conclusions in section 3.3 and section 3.4.

467 To assess the environmental implications of different sources, the ozone formation



468 potential (OFP) of NMHCs was calculated using  $OFP = MIR \times C$ , where *MIR* depicts  
469 the maximum incremental reactivity and *C* represents the concentration of NMHCs  
470 (Carter, 1994). Specific data can be seen in supplementary Table S11. The contributions  
471 of different factors to the OFP are as follows: industrial activities ( $2.30 \mu\text{g m}^{-3}$ , 56 %),  
472 exhaust emissions ( $0.87 \mu\text{g m}^{-3}$ , 21 %), oceanic ventilation ( $0.64 \mu\text{g m}^{-3}$ , 16 %), and  
473 terrestrial vegetation emissions ( $0.27 \mu\text{g m}^{-3}$ , 7 %). Notably, terrestrial sources  
474 collectively accounted for 89 % of alkanes and 69.6 % of alkenes within the coastal  
475 marine atmosphere of China. Furthermore, these terrestrial factors contributed 84 % of  
476 the OFP associated with C2-C5 NMHCs. These findings highlight that terrestrial  
477 outflow substantially constitutes the atmospheric NMHCs and plays a significant role  
478 in regulating air quality in nearshore environments.

#### 479 **4 Conclusions**

480 Our study characterized the atmospheric NMHCs in both coastal cities and Chinese  
481 marginal seas, and determined that both oceanic ventilation and terrestrial inputs play  
482 important roles in controlling the distribution and chemical composition of NMHCs in  
483 the coastal atmosphere of China.

484 Alkanes were the dominant NMHCs both in urban and nearshore atmosphere, and  
485 the atmospheric concentrations of alkanes were significantly higher in coastal cities  
486 compared to coastal seas, showing the potential of terrestrial alkanes as a source of  
487 alkanes in the marine atmosphere through transport. Generally, alkane concentrations  
488 tended to be higher in cases: sampling sites closer to land, longer retention of air mass  
489 over land, and shorter air mass transport time from land to sampling site. However,  
490 these effects could not apply to alkenes due to their higher reactivities and the  
491 substantial sea-to-air fluxes. Additionally, the impact of oceanic emissions on NMHCs



492 composition was more pronounced in areas beyond 100 km from land compared to  
493 areas within 100 km, because the terrestrial input gradually diminishes along the  
494 direction towards the open ocean.

495 Combining the outcomes of the PMF model and chemical profiles of diverse sources  
496 in the literature, we extracted four isolated sources of NMHCs in the nearshore  
497 atmosphere. Terrestrial sources (including industrial activities, vehicular exhaust, and  
498 vegetation emission) primarily constitute the NMHCs in the nearshore atmospheres,  
499 further contributing 84 % to the OFP associated with C2-C5 NMHCs. This highlights  
500 the significant influence of terrestrial outflow on the distribution and composition of  
501 NMHCs in the nearshore atmosphere of China, emphasizing the necessity for a  
502 comprehensive understanding of both natural and anthropogenic emissions of NMHCs.

### 503 **Code and data availability**

504 Data presented in this paper are publicly available at Figshare via  
505 <https://doi.org/10.6084/m9.figshare.24722286>. The remote-sensing datasets of Chl-*a*  
506 and total absorption at 443 nm are available at <https://oceancolor.gsfc.nasa.gov>. Code  
507 to calculate retention of air mass over land can be download from  
508 <https://doi.org/10.1029/2021JD034960> (Zhou et al., 2021).

### 509 **Competing interests**

510 The authors declare that they have no conflict of interest.



511 **Author contributions**

512 Honghai Zhang and Jian Wang designed the investigation and experiments. Jian  
513 Wang, Qian Yao Ma, Feng Xu, Gaobin Xu, Shibo Yan, Jiawei Zhang, and Jianlong Li  
514 collected and determined the samples. Jian Wang analyzed the data and wrote the  
515 manuscript. Honghai Zhang, Lei Xue, Zhaohui Chen, and Guiling Zhang reviewed and  
516 revised the manuscript.

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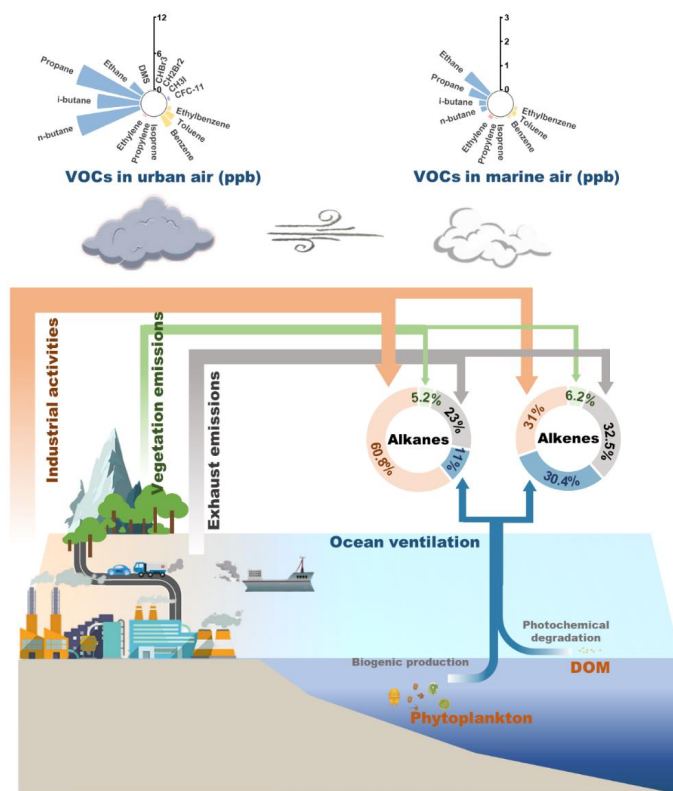
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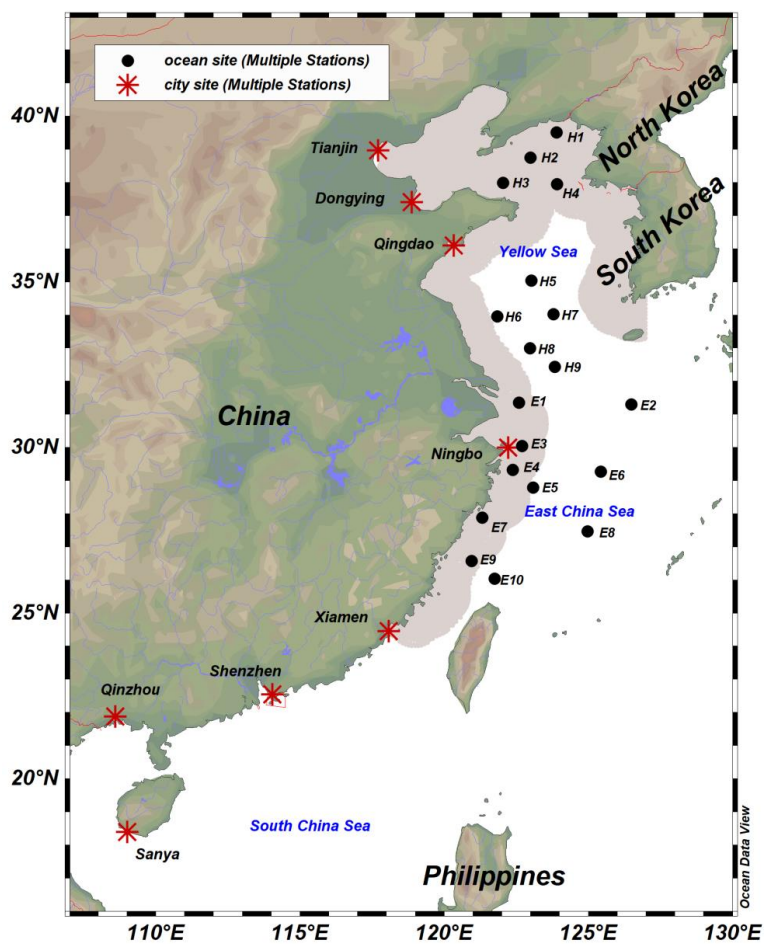


816 **Figure Captions**



817

818 **Graphical abstract** Schematic diagram showing the main sources and their relative contributions  
 819 to the non-methane hydrocarbons (NMHCs) budget in the nearshore atmospheres of China. The ring  
 820 bar chart above the land or the ocean shows the composition of urban or marine atmospheric trace  
 821 gases determined in this study. The axes with unit of ppb indicate the atmospheric concentrations of  
 822 gases. The distinct colored wedges indicate the alkanes (skyblue), alkenes (pink), monocyclic  
 823 aromatics (yellow), volatile halogenated compounds (VHCs, lilac), and dimethyl sulfur (DMS,  
 824 palegreen). Note that only alkanes, alkenes, and monocyclic aromatics are shown in marine  
 825 atmosphere. The colored arrows or annuli indicate the main sources of NMHCs in offshore  
 826 atmosphere: industrial activities (sandybrown), exhaust emissions (darkgray), oceanic ventilation  
 827 (steelblue), and vegetation emissions (lightgreen). The numbers on the annuli are their respective  
 828 relative contributions.

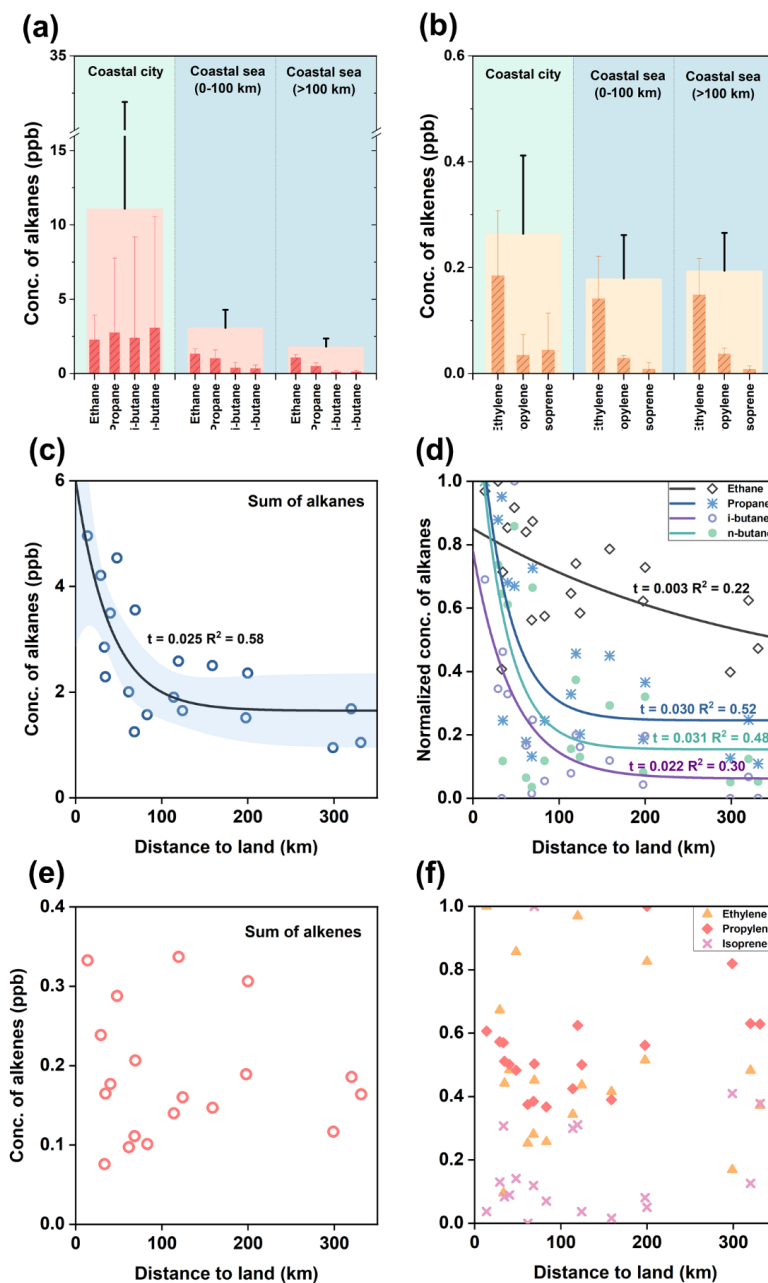


829

830 **Figure 1** Map showing the sampling stations in the coastal cities (red asterisks) and marginal seas

831 (black dots) of China from March to May 2021. The gray shaded area represents the inshore region

832 within 100 km from the coastline.



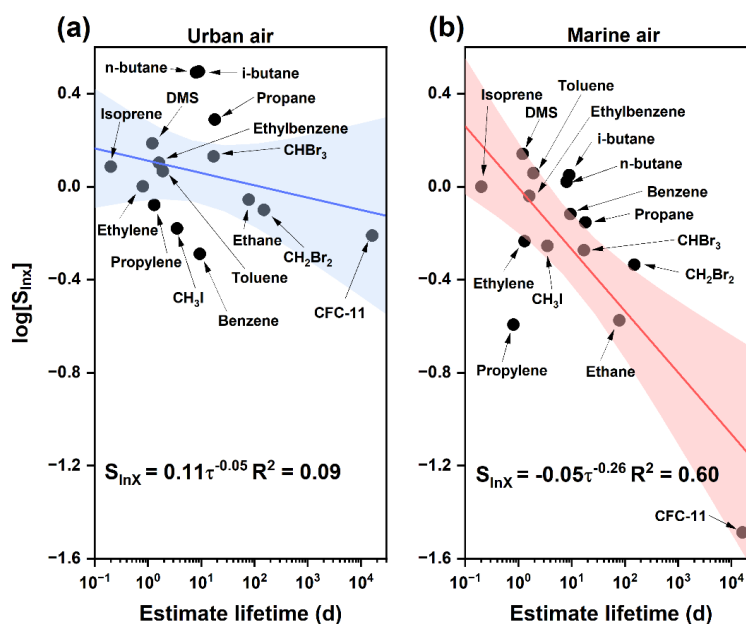
833

834 **Figure 2** Means of the concentrations of alkanes (panel a) and alkenes (panel b) in the atmosphere

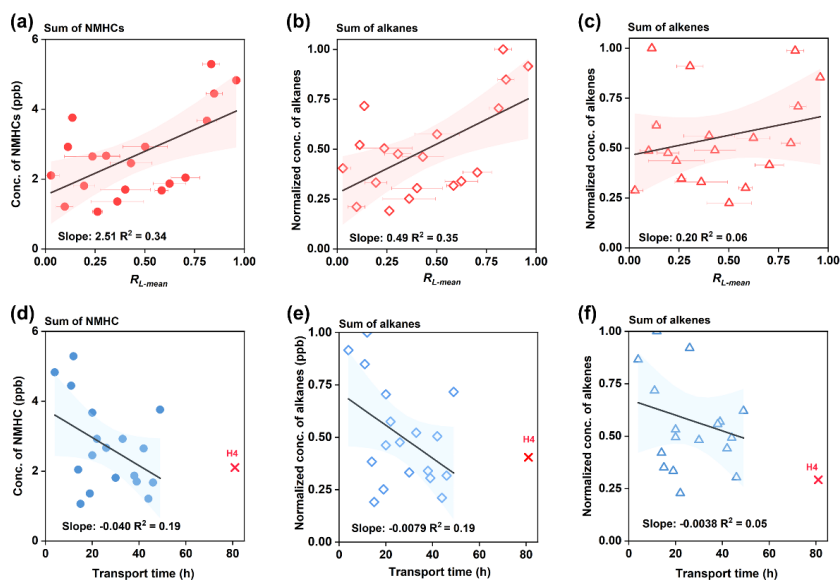
835 over coastal cities (n = 14) and nearshore (0-100 km, n = 10) and offshore (>100 km, n = 9) coastal



836 seas of China. The wider columns in panel a or b represent the sums of individual alkanes or alkenes  
 837 with error bars depicting the propagated errors from each NMHCs. Summed alkane (panel c) or  
 838 alkene (panel e) and normalized concentrations of specific alkane (panel d) or alkene (panel f)  
 839 plotted as a function ( $y = Ae^{-tx} + y_0$ ) of the distance from sampling sites to the nearest land.  
 840

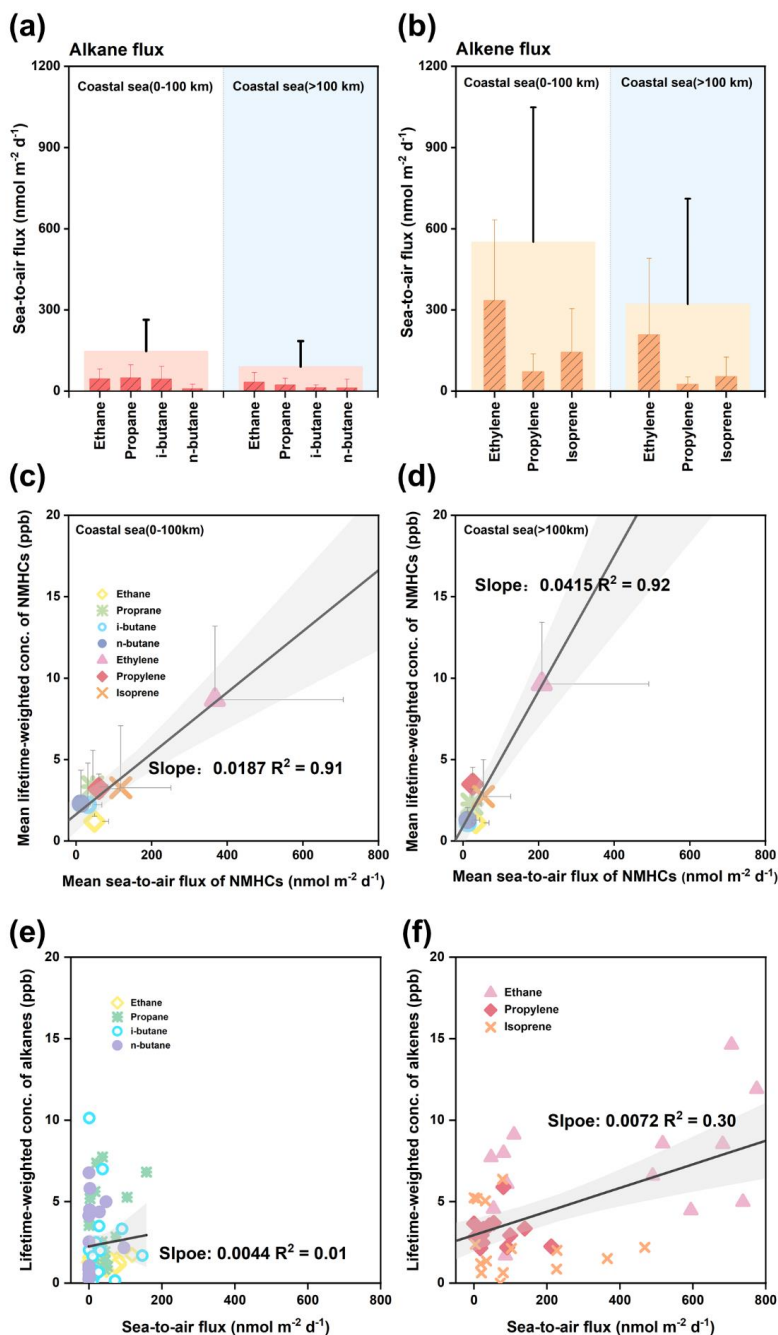


841  
 842 **Figure 3** Atmospheric variability ( $\log[S_{inx}]$ ) plotted as a function of the estimated  $\bullet$ OH lifetime for  
 843 each non-methane hydrocarbons (NMHCs) from the coastal cities (panel a) and marginal seas of  
 844 China (panel b). The blue or red line is the best linear fitting. Shaded area represents the  
 845 confidence band at a 95 % confidence level.



846

847 **Figure 4** Concentrations of non-methane hydrocarbons (NMHCs) combined (panel a or d), alkanes  
 848 (panel b or e), and alkenes (panel c or f) at each site plotted against the mean retentions of air mass  
 849 over land ( $R_{L-mean}$ ,  $n = 3$ ) or the transport time of air mass, respectively. The error bars for  
 850  $R_{L-mean}$  indicate the standard deviation from three different time-scale trajectories (48h, 72h, and  
 851 96h). The black line is the best fitting of liner function and shadowed area represents the confidence  
 852 band at a 95 % confidence level. H4 (marked with red "x") is treated as an outlier since it alone  
 853 deviates from the main dataset.



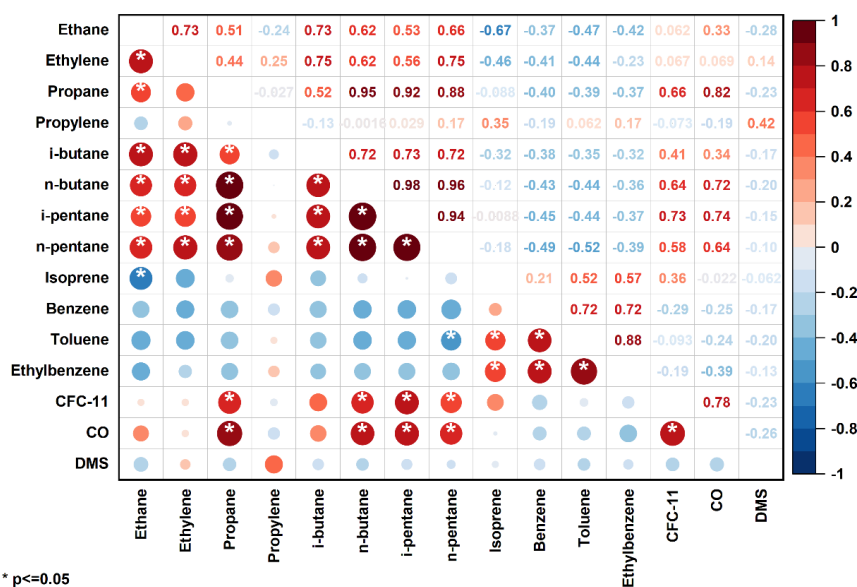
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855 **Figure 5** Means of sea-to-air fluxes of alkanes (panel a) and alkenes (panel b) in sea areas within

856 100 km (n = 10) and beyond 100 km (n = 9) from the nearshore land. The wider columns represent



857 the sum of alkanes or alkenes. Panel c or d shows the means of lifetime-weighted concentrations of  
 858 NMHCs plotted against the means of their mean sea-to-air fluxes in the area within 100 km or  
 859 beyond 100 km from the coastline. Specific lifetime-weighted concentrations of alkanes (panel e)  
 860 and alkenes (panel f) plotted against sea-to-air fluxes in the whole coastal sea region. The black,  
 861 blue or red line is the best linear fitting for each dataset and shadowed area represents the confidence  
 862 band at a 95 % confidence level.  
 863

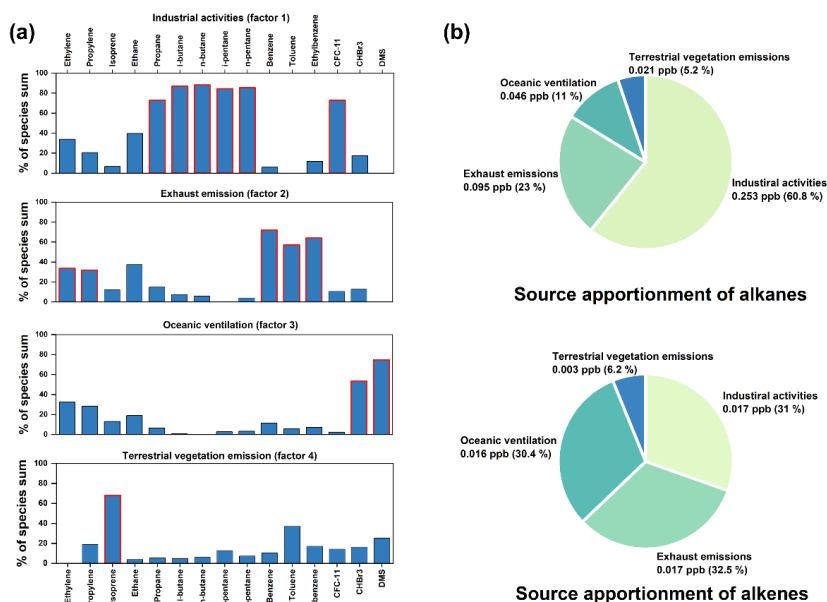


864

\*  $p < 0.05$

865 **Figure 6** Correlation coefficients (r) between the various trace gases determined in the atmosphere  
 866 over the Yellow Sea and the East China Sea. The white asterisk means the correlation is significant  
 867 at the  $p < 0.05$  level. The color of dots, red or blue, indicates the positive or negative correlation and  
 868 the size of the dots indicates the absolute value of r.





869

870 **Figure 7** Representative factor profiles from the positive matrix factorization (PMF) model (panel

871 a). Non-methane hydrocarbons (NMHCs) marked with red rim are selected as indicators for the

872 specific factors. Relative contributions of different factors/sources to the alkanes and alkenes in the

873 oceanic atmosphere (panel b).



## Table

**Table 1** Atmospheric and seawater concentrations, sea-to-air fluxes, and the calculated atmospheric lifetime of each non-methane hydrocarbons (NMHCs)

based on the reaction with hydroxyl radicals ( $\bullet\text{OH}$ ).

Species	Conc. in urban air (ppb)	Conc. in oceanic air (ppb)	Conc. in seawater ( $\mu\text{mol L}^{-1}$ )	Sea-to-air flux ( $\text{mmol m}^{-2} \text{d}^{-1}$ )	Atmospheric lifetime <sup>b</sup> (d)
Ethane	2.26 (0.277-5.72)	1.24 (0.686-1.72)	11.6 (4.70-22.8)	44.6 (0.2-118)	78
Propane	2.95 (0.149-20.1)	0.822 (0.226-1.79)	12.6 (3.68-136)	41.5 (0.2-157)	18
i-butane	2.57 (BD <sup>a</sup> -27.6)	0.283 (BD-1.17)	9.46 (1.54-35.3)	31.7 (0.1-146)	9.1
n-butane	3.29 (0.018-30.2)	0.256 (0.025-0.694)	4.95 (BD-32.9)	10.9 (-0.8-96.1)	8.2
Ethylene	0.180 (0.035-0.390)	0.151 (0.028-0.295)	70.4 (8.40-136)	321 (1.7-775)	2.3
Propylene	0.036 (BD-0.129)	0.033 (0.022-0.060)	15.2 (2.42-27.6)	56.1 (0.2-212)	0.73
Isoprene	0.046 (0.006-0.250)	0.008 (BD-0.043)	31.0 (3.43-105)	112 (0.5-468)	0.19

<sup>a</sup>: Below the detection limit.

<sup>b</sup>: Assuming an average  $[\bullet\text{OH}]$  of  $6 \times 10^5$  molecules  $\text{cm}^{-3}$  within 24 h (Jobson et al., 1999), and using the rate constant with  $\bullet\text{OH}$  at 288 K taken from Atkinson et al. (1997).