

1 **Measurement Report: Elevated atmospheric ammonia**
2 **may promote the particle pH and HONO formation:**
3 **Insights from the COVID-19 pandemic**

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18

19 **Abstract**

20 HONO plays a crucial role as a precursor to OH radicals in the tropospheric atmosphere.
21 The incongruity between HONO concentration and NO_x emissions during the COVID-
22 19 pandemic remains puzzling. Here, we show evidence from field observations of ten
23 sites in China that there was a noticeable increase in NH₃ concentrations during the
24 COVID-19 pandemic. In addition to the meteorological conditions, the significant
25 decrease in sulfate and nitrate concentrations enhanced the partitioning of NH₄⁺ to NH₃.
26 Sensitivity analysis indicated that the decrease in anion concentrations (especially
27 sulfate and nitrate) and the increase in cation concentrations during the COVID-19
28 pandemic led to an increase in particle pH. In other words, changes in the excess
29 ammonia drove changes in particle pH that may consequently have impacted the rate
30 of HONO formation. The calculation of reaction rates indicates that during the epidemic,
31 the increase in pH may promote the generation of HONO by facilitating redox reactions,
32 which highlights the importance of coordinating the control of SO₂, NO_x, and NH₃
33 emissions.

34 **Keywords:** Ammonia, HONO, Gas-particle partitioning, Acidity, COVID-19 pandemic

35

36 **1. Introduction**

37 Nitrous acid (HONO) is a critical precursor of hydroxyl radical (OH), contributing
38 to more than 60% of OH production (Alicke, 2003;Platt et al., 1980;Kleffmann et al.,
39 2005). The OH can react with carbon monoxide, nitrogen oxides (NO_x), sulfur dioxide
40 (SO₂), and volatile organic compounds to produce secondary pollutants such as ozone
41 (O₃) and PM_{2.5} (particulate matter with an aerodynamic diameter less than or equal to
42 2.5 μm), thereby affecting air quality, human health, and global climate change (Li et
43 al., 2021a;Wang et al., 2023b;Lu et al., 2018)

44 High concentrations of HONO are present in urban daytime atmospheres, and
45 exploring its sources has become a hot and challenging topic in the field of atmospheric
46 chemistry (Jiang et al., 2022;Xu et al., 2019). Various sources of atmospheric HONO
47 have been identified, including combustion processes (e.g., vehicle emissions) (Kramer
48 et al., 2020;Liao et al., 2021a;Li et al., 2021b), direct emissions from soil (Su and Zhang,
49 2011;Oswald et al., 2013;Meusel et al., 2018), homogeneous reactions between NO and
50 OH radicals (Pagsberg, 1997;Atkinson and Rossi, 2004), heterogeneous reactions of
51 NO₂ on aerosols and ground surfaces (Zhang et al., 2020a;McFall et al., 2018;Liu et al.,
52 2014;Liu et al., 2020a), and photolysis of nitrate (Spataro and Ianniello, 2014;Scharko
53 et al., 2014;Romer et al., 2018;Ye et al., 2017;Shi et al., 2021). During the pandemic
54 control periods, there was a substantial reduction in vehicle traffic flow and industrial
55 emissions, leading to a decrease of more than 60% in NO_x emissions in eastern China
56 (Huang et al., 2021a). It was initially expected that the concentration of HONO would
57 also decrease proportionally. However, Liu et al. (2020b) observed that the decrease in

58 HONO concentration during the pandemic period was only 31% (from 1.5 ppb to 0.9
59 ppb), which was significantly lower than the reductions in NO (62%, from 26.3 to 4.2
60 ppb) and NO₂ (36%, from 15.5 to 6.2 ppb). Furthermore, the observed concentrations
61 of HONO during the COVID-19 pandemic in 2020 were higher than those during the
62 corresponding period in 2021 in Beijing (Luo et al., 2023). These findings suggest the
63 existence of a considerable unknown source of HONO during the COVID-19 pandemic
64 period.

65 Ammonia (NH₃) is a primary alkaline gas in the atmosphere, capable of influencing
66 the pH level of particulate matter and plays a crucial role in the atmospheric nitrogen
67 cycle (Gu et al., 2022; Xu et al., 2020; Gong et al., 2011). Several studies have indicated
68 that NH₃ can promote the formation of HONO by promoting the hydrolysis of NO₂ (Xu
69 et al., 2019) or the redox reaction of NO₂ with SO₂ (Liu et al., 2023). Moreover,
70 previous studies have reported that NH₃ concentrations in the atmosphere, particularly
71 in rural areas, significantly increased during the pandemic (Xu et al., 2022).
72 Consequently, the rise in NH₃ may contribute to the enhanced formation of HONO
73 (Huang et al., 2021a). Unfortunately, there is currently a lack of research on the
74 relationship between enhanced NH₃ and HONO during the COVID-19 pandemic period.

75 To address this, online observational data on the chemical composition of PM_{2.5},
76 gaseous pollutants, and meteorological conditions at ten sites in China before and
77 during the COVID-19 pandemic period were analyzed to investigate the variation in
78 NH₃ concentrations and particle pH, and explore the promoting effect of increased pH
79 values on HONO formation.

80 **2. Materials and methods**

81 **2.1 Observation sites**

82 Online measurements were conducted at four urban and six rural sites in Henan
83 Province, China from January 1 to February 29, 2020, including Sanmenxia (U-SMX),
84 Zhoukou (U-ZK), Zhumadian (U-ZMD), and Xinyang (U-XY), as well as rural
85 locations including Anyang (R-AY), Xinxiang (R-XX), Jiaozuo (R-JZ), Shangqiu (R-
86 SQ), Nanyang (R-NY), and Puyang (R-PY). Descriptions and the spatial distribution
87 of these ten sites can be found in Table S1 and Fig. S1.

88 **2.2 Measurements**

89 The aerosol and gas monitor (MARGA, Metrohm, Switzerland) was used to analyze
90 the hourly water-soluble ions (Na^+ , NH_4^+ , K^+ , Mg^{2+} , Ca^{2+} , Cl^- , NO_3^- , and SO_4^{2-}) in $\text{PM}_{2.5}$,
91 as well as gaseous species (NH_3 , HNO_3 , HCl , and HONO) at ten sampling sites. The
92 MARGA instrument is widely used (Chen et al., 2017; Stieger et al., 2019; Twigg et al.,
93 2022). A detailed description of the instrument and QA/QC can be found in Text S1. In
94 brief, the atmospheric sample passes through a $\text{PM}_{2.5}$ cut-off head, and both particles
95 and gases enter a wet rotating dissolution device for diffusion. Subsequently, the
96 particles in the sample undergo hygroscopic growth and condensation in an aerosol
97 supersaturated vapor generator, followed by collection and ion chromatographic
98 analysis. The gases in the sample are oxidized by H_2O_2 in the dissolution device,
99 absorbed into a liquid solvent, and then entered the gas sample collection chamber for

100 ion chromatographic quantification. The range of minimum detection limits for water-
101 soluble ions was between $0.002 \mu\text{g}/\text{m}^3$ (Cl^-) to $0.081 \mu\text{g}/\text{m}^3$ (NH_4^+). Uncertainties of 20%
102 are assumed for the detection of NH_3 and NH_4^+ , while uncertainties of 10% are assumed
103 for other components (Wang et al., 2020; Wang et al., 2022). In addition, a detailed
104 description of HONO measurement using this system can be found in Text S2. Overall,
105 the limit of detection for HONO was 4 pptv and the uncertainty was estimated to be \pm
106 20%.

107 The data for NO_2 and SO_2 were obtained from a series of instruments provided by
108 Thermo Fisher Scientific (USA). The hourly concentrations of organic carbon (OC) in
109 $\text{PM}_{2.5}$ were analyzed using a carbon analyzer (Model 4, Sunset Laboratory., USA).
110 Detailed descriptions of the NO_2 , SO_2 , and carbon analyzers can be found in Text S3.
111 The smart weather stations (LUFFTWS500, Sutron, Germany) were utilized for
112 synchronized observation of meteorological parameters including pressure,
113 temperature (T), and relative humidity (RH).

114 **2.3 Data analysis.**

115 **2.3.1 pH prediction.**

116 The thermodynamic model ISORROPIA-II was used to estimate the pH value of the
117 particles (Fountoukis, 2007) by inputting RH, T, K^+ , Ca^{2+} , Mg^{2+} , total ammonia
118 ($\text{TNH}_x = 17 \times (\frac{[\text{NH}_4^+]}{18} + \frac{[\text{NH}_3]}{17})$), total sulfuric acid (TH_2SO_4 , SO_4^{2-}), total sodium
119 (TNa , Na^+), total chlorine (TCl , Cl^-), and total nitrate ($\text{TNO}_3 = \text{NO}_3^- + \text{HNO}_3$). The

120 model has two calculation modes: the forward mode and reverse mode, and the aerosol
 121 dissolution systems can be set to simulate a metastable state (aqueous phase) or stable
 122 state (aqueous and solid phase). Studies have shown that the forward mode is less
 123 affected by instrument measurement errors than the reverse mode (Ding et al.,
 124 2019; Song et al., 2018). Additionally, the minimum average RH of about 55% was
 125 recorded during the sampling period at the ten sites. Thus, ISORROPIA-II was run in
 126 the forward model for the aerosol system in the metastable condition and only used data
 127 with $RH \geq 30\%$ for simulation accuracy (Ding et al., 2019; Song et al., 2018). The
 128 ISORROPIA model calculated the particle hydrate ion concentration per volume of air
 129 (H_{air}^+) and aerosol water associated with inorganic matter (AWC_{inorg}). The pH value
 130 was calculated using the following equation (Bougiatioti et al., 2016):

$$131 \quad pH = -\log_{10} H_{aq}^+ = -\log_{10} \frac{1000 H_{air}^+}{AWC_{inorg} + AWC_{org}} \quad (2.1)$$

132 where the modeled concentrations for AWC_{inorg} and H_{air}^+ are $\mu\text{g}/\text{m}^3$, and AWC_{org} is the
 133 particle water associated with the organic matters predicted using the following method:

$$134 \quad AWC_{org} = \frac{m_s}{\rho_s} \frac{k_{org}}{\left(\frac{1}{RH} - 1\right)} \quad (2.2)$$

135 where m_s is the mass concentration of organic matter ($OM = OC \times f$). The f is the
 136 conversion factor of OC, which is dependent on the extent of OM oxidation and
 137 secondary organic aerosol formation (Chow et al., 2015). Studies on the ratio of
 138 OM/OC in fourteen cities in China suggested that the mean value of f was 1.59 ± 0.18
 139 during the winter season in Northern China (Xing et al., 2013), and thus we adopted 1.6
 140 as the f in this study. k_{org} is the organic hygroscopicity parameter and depends on organic

141 functionality, water solubility, molecular weight, and oxidation level. Han et al. (2022)
142 have reported that the k_{org} generally increased with O: C ratios, with a range of 0–0.3
143 for 23 organics, including carboxylic acids, amino acids, sugars, and alcohols. Gunthe
144 et al. (2011) estimated a $k_{org} = 0.06 \pm 0.01$ for the effective average hygroscopicity of
145 the non-refractory organic particulate matter in the aerosols in Beijing. Our previous
146 study has estimated that the uncertainties of k_{org} value (0.06) for pH in the range of 0–
147 0.3 only lead to –1–3% errors, which can be neglected (Wang et al., 2023a). Therefore,
148 the value of 0.06 was selected in this paper. ρ_s is the organic density, which was chosen
149 to be 1.35 g/cm³ following previous studies (Table S2).

150 **2.3.2 The sources of HONO**

151 The sources of HONO include direct emission (P_{emi}), the homogeneous reaction of
152 NO and •OH (P_{OH+NO}), the heterogeneous reaction of NO₂ on the ground (P_{ground}) and
153 aerosol ($P_{aerosol}$), the photo-enhanced heterogeneous reaction of NO₂ on the ground
154 ($P_{ground+hv}$) and aerosol ($P_{aerosol+hv}$), and nitrate photolysis ($P_{nitrate}$). The detailed
155 calculation method is described in the Supplementary Material (Text S4, Table S3, Figs.
156 S2 and S3).

157 Soil emission has been demonstrated to be a major source of HONO, which is affected
158 by temperature to some extent (Liu et al., 2020c; Liu et al., 2020b). However, during the
159 sampling periods, there was no significant positive correlation between HONO
160 concentration and temperature (Fig. S4). In addition, temperatures did not exceed 10°C,
161 under which the soil HONO emission rate is generally considered to be zero (Zhang et

162 al., 2023). Furthermore, the equilibrium gas-phase concentration over an aqueous
163 solution of nitrous acid, [HONO]*, a key parameter controlling the exchange of HONO
164 between the gas and aqueous phase in soil, is calculated according to Su et al.(2011).
165 The results indicate that the temperature difference between PC and DC periods only
166 led to approximately a 0.01% concentration change. On the other hand, studies on the
167 sources of HONO in the North China Plain of China during winter consistently showed
168 that soil HONO emissions contribute around 1% (Liu et al., 2020c;Liu et al.,
169 2020b;Zhang et al., 2023). Therefore, this study does not consider soil HONO
170 emissions.

171 **2.3.3 Redox reaction of NO₂ with SO₂.**

172 The redox reaction of NO₂ with SO₂ (R₁) is considered a crucial potential source of
173 high concentrations of HONO in Northern China (Cheng et al., 2019;Wang et al., 2016):



175 the rate expression for reaction (R₁) was estimated to:

$$176 \quad d[\text{S(VI)}] / dt = k_1[\text{NO}_2][\text{S(VI)}], \quad (2.3)$$

177 the rate constant k₁ value is pH dependent, e.g., for pH, 5, k₁ = (1.4×10⁵+1.24×10⁷)/2
178 M⁻¹ s⁻¹. For k₁ values under other pH conditions and other related information, please
179 refer to Text S5, Table S4, and Table S5.

180 **3. Results and discussion**

181 **3.1 Variations of NH₃, NH₄⁺ and TNH_x.**

182 The temporal variations of NH₃, NH₄⁺, and TNH_x at 10 sampling sites in the pre-
183 COVID-19 pandemic period (PC, January 1 to 23, 2020) and during the COVID-19
184 pandemic period (DC, January 24 to February 29, 2020) are presented in Fig. 1, with
185 their average concentration listed in Table 1. In general, rural sites exhibited higher
186 concentrations of NH₃, NH₄⁺, and TNH_x compared to urban sites, except for the R-NY
187 site. This finding is consistent with previous studies conducted in Zhengzhou (Wang et
188 al., 2020), Shanghai (Chang et al., 2019), and Quzhou (Feng et al., 2022), owing to the
189 intense agricultural ammonia emissions. The highest concentrations of NH₃ and TNH_x
190 were recorded at site R-JZ, with average values of 25.3 ± 11.5 and 40.8 ± 20.1 $\mu\text{g}/\text{m}^3$,
191 respectively. Site R-AY had the highest NH₄⁺ concentration, measuring 19.3 ± 12.9
192 $\mu\text{g}/\text{m}^3$. Note that the current study area exhibited higher NH₃ levels compared to other
193 regions (Table S6), which probably was attributed to the highest NH₃ emissions of
194 Henan Province in China, primarily from nitrogen fertilizer application and livestock
195 farming (Wang et al., 2018; Ma, 2020).

196 Compared to the PC, NH₃ concentrations increased in the DC at all sites. Notably,
197 rural sites experienced more significant increases in NH₃ concentrations than urban
198 sites, which was similar to the trend in Shanghai (Xu et al., 2022). The largest increases
199 in NH₃ concentrations were observed at R-SQ (71%, 7.3 $\mu\text{g}/\text{m}^3$) and U-ZK (37%, 4.8
200 $\mu\text{g}/\text{m}^3$) for rural and urban sites, respectively. In contrast, the concentrations of NH₄⁺

201 and TNH_x decreased in the DC with the largest reduction at rural site R-PY (51%, 12.9
202 $\mu\text{g}/\text{m}^3$) and urban site U-ZMD (48%, 9.3 $\mu\text{g}/\text{m}^3$). Regarding TNH_x , rural sites exhibited
203 larger reductions, with site R-SQ experiencing the largest decrease of 37% (4.7 $\mu\text{g}/\text{m}^3$).

204 Figure 2 illustrates the spatial distribution and the diurnal variation of NH_3
205 concentrations at the ten sites before and during the pandemic. NH_3 concentrations in
206 most sites exhibited an unimodal trend in PC that NH_3 concentrations gradually
207 increased after sunrise, reaching a peak around noon (11:00–12:00), and then decreased
208 to a valley around 16:00–17:00. This diurnal pattern is similar to NH_3 variations
209 observed in urban areas of Houston, USA, as a result of the natural emissions from
210 vegetation and soil during photosynthesis (Gong et al., 2011). However, other studies
211 have recorded a significant NH_3 peak during the early morning of 8:00–10:00 (Ellis et
212 al., 2011; Meng et al., 2018; Gu et al., 2022), suggesting the influence of vehicle
213 emissions (Gong et al., 2011; Gu et al., 2022), residual NH_3 mixing, soil or plant
214 emissions (Ellis et al., 2011), and dew volatilization (Wentworth et al., 2016; Huang et
215 al., 2021b). Therefore, the NH_3 in urban and rural areas of this study was probably less
216 affected by NH_3 emissions from vehicles, different from the recent studies in megacities
217 of China (e.g., Beijing and Shanghai) (Gu et al., 2022; Wu et al., 2023; Zhang et al.,
218 2020b). In addition to the transport from agricultural emissions, urban NH_3 in this
219 region might also originate from other non-agricultural sources, such as wastewater
220 treatment, coal combustion, household waste, urban green spaces, and human
221 excrement (Chang et al., 2019).

222 During the COVID-19 pandemic, the diurnal variation of NH_3 in both urban and rural

223 sites still maintained an unimodal distribution. The peak values in urban sites remained
224 consistent with PC levels, further demonstrating that the influence of vehicles on NH₃
225 in urban areas was limited. Notably, the peak time of NH₃ in rural sites shifted 1–2 hours
226 earlier compared to the trend in PC. Ammonia in rural areas primarily originates from
227 nitrogen fertilizer application, livestock, and poultry breeding (Feng et al., 2022; Meng
228 et al., 2018), which are significantly influenced by T and RH (Liu et al., 2023). Table
229 S7 and Fig. S5 reveal that there was an increased T and a decreased RH at rural sites in
230 the DC than the PC, which could accelerate the evaporation of NH₃ and thus potentially
231 lead to earlier peak NH₃ concentrations.

232 **3.2 Gas-to-particle conversion of NH₃**

233 The increased NH₃ accompanying decreased NH₄⁺ in the DC suggests that the gas-
234 particle partition of NH₃/ NH₄⁺ may determine the elevated NH₃ concentrations.
235 Meteorological parameters, including RH and T, play a crucial role in the gas-particle
236 partitioning of NH₃ (Liu et al., 2023; Xu et al., 2020). Therefore, the higher T and lower
237 RH in the DC (Table S7 and Fig. S5) favored the conversion of NH₄⁺ to NH₃, resulting
238 in a decrease in $\epsilon(\text{NH}_4^+)$ ($[\text{NH}_4^+]/([\text{NH}_3] + [\text{NH}_4^+])$) compared to those in the PC (Table
239 S7).

240 NH₃ primarily enters particles to neutralize acidic ions (Wang et al., 2020; Xu et al.,
241 2020; Liu et al., 2017; Ye et al., 2011; Wells, 1998). Accordingly, the concentrations of
242 required ammonia (Required-NH_x) and excess ammonia (Excess-NH_x) were calculated
243 based on the acidic substances as follows (Wang et al., 2020):

$$\begin{aligned}
 \text{Required-NH}_x &= 17 \times \left(\frac{[\text{SO}_4^{2-}]}{48} + \frac{[\text{NO}_3^-]}{63} + \frac{[\text{Cl}^-]}{35.5} + \frac{[\text{HNO}_3]}{64} + \frac{[\text{HCl}]}{36.5} \right) \\
 &- 17 \times \left(\frac{[\text{Na}^+]}{23} + \frac{[\text{K}^+]}{39} + \frac{[\text{Ca}^{2+}]}{20} + \frac{[\text{Mg}^{2+}]}{12} \right)
 \end{aligned}
 \tag{3.1}$$

$$\text{Excess-NH}_x = \text{TNH}_x - \text{Required-NH}_x
 \tag{3.2}$$

where [W] represents the concentration of the substance ($\mu\text{g}/\text{m}^3$). The significant linear fitting (R^2 is greater than 0.96, and the slope is close to 1) in Fig. S6 demonstrates that the anions and cations at each site were close to the equilibrium state. Therefore, the organic acids in $\text{PM}_{2.5}$ may have less effect on NH_3 and NH_4^+ and were not considered in Formula 3.1.

As shown in Fig. 3 and Table S8, compared to those in the PC, the concentration of Required- NH_x in the DC significantly decreased (ranging from 37% at site R-JZ to 58% at site R-PY), while the concentration of Excess- NH_x increased (ranging from 9% at site R-AY to 78% at site R-SQ). The reduction in the concentrations of sulfate and nitrate (Fig. S7) was responsible for the decrease in the concentration of Required- NH_x . To sum up, in addition to meteorological conditions, the substantial reduction in anthropogenic emissions of SO_2 , NO_x , and other pollutants in the DC had led to a decrease in acidic substances (e.g., sulfate and nitrate) in particles, in turn, resulting in more gas-phase NH_3 concentration remaining in the atmosphere.

3.3 Particle pH before and during COVID-19

Diurnal patterns of particle pH in PC and DC at ten sites are summarized in Fig. 4 with their average values listed in Table S9. $\text{PM}_{2.5}$ shows consistent moderate acidity, with mean values in the range of 4.2–5.1, which were close to the values in previous

264 studies (Table S9). Compared to the PC, the particle pH at ten sites increased obviously
265 in the DC, with the highest increase of 0.5 (U-ZK) and 0.3 (R-PY) at urban and rural
266 sites, respectively, which were the subject of an in-depth discussion in the following
267 text.

268 To explore the dominant factors that determine the local particle pH level and result
269 in the high pH during the DC, sensitivity tests of pH to chemical species (i.e., TNH_x ,
270 TH_2SO_4 , TNO_3 , TCl , TNa , K^+ , Ca^{2+} , and Mg^{2+}) and meteorological parameters (i.e., T
271 and RH) were performed. A given range for a variable (i.e., TNH_x) with corresponding
272 average values of other parameters (i.e., TH_2SO_4 , TNO_3 , TCl , TNa , K^+ , Ca^{2+} , Mg^{2+} , T ,
273 and RH) was input into the model and simulated to compare its effects on pH. As shown
274 in Fig. S8, pH increases with the cation concentrations (i.e., TNH_x , Na^+ , K^+ , Ca^{2+} , and
275 Mg^{2+}) increasing as well as the anion concentrations (i.e., TH_2SO_4 , TNO_3 , and Cl^-), T
276 and RH decreasing. According to the average values of input data during PC (Blue line
277 in Fig. S8) and DC (Red line in Fig. S8) at U-ZK and R-PY sites respectively, the
278 changes in pH (ΔpH in Fig. 5) indicate that the decrease in TNH_x concentration and the
279 increase in T in DC led to a decrease in pH values (ΔpH : 0.09 at U-ZK and 0.08 at R-
280 PY sites) compared to PC. However, this effect was outweighed by the decrease in
281 TH_2SO_4 (ΔpH : 0.07 and 0.8 at U-ZK and R-PY sites, respectively) and TNO_3 (ΔpH :
282 0.05 and 0.4 at U-ZK and R-PY sites, respectively) concentrations as well as the
283 increase in K^+ (ΔpH : 0.03 at U-ZK and 0.2 at R-PY site) and Mg^{2+} (ΔpH : 0.01 at U-ZK
284 and 0.04 at R-PY site) concentrations in the DC, and resulting in an overall increase in
285 pH values in the DC. Furthermore, the relationship between particle pH with the

286 concentrations of Required-NH_x, and Excess-NH_x, which considers all chemical
287 components, is investigated to examine the dominant factor on the increasing pH in DC.
288 As shown in Fig. 6, the higher Excess-NH_x concentrations in the DC led to higher
289 increases in pH values (Δ pH: 1 at U-ZK and 0.5 at R-PY site) than those in PC (Δ pH:
290 0.3 at U-ZK and 0.2 at R-PY site), thus Excess-NH_x concentrations may be the key
291 factor in promoting the pH values.

292 **3.4 The influence of pH on HONO.**

293 The observed HONO concentrations decreased by 18% and 54% at U-ZK (0.8 ppb)
294 and R-PY (0.9 ppb) sites in the DC, respectively, compared to those (1.0 and 2.2 ppb)
295 in the PC. Moreover, all the known HONO production sources rates including P_{emi} , P_{OH}
296 + NO, P_{ground} , $P_{\text{ground+hv}}$, P_{aerosol} , $P_{\text{aerosol+hv}}$, and P_{nitrate} (Fig. 7, Fig S9 and S10) show a
297 decreasing trend from PC to DC, with the total reductions of 38% (from 30% to 45%
298 in the scenario with the minimum and maximum uncertainty, respectively) and 79%
299 (from 77% to 82% in the scenario with the minimum and maximum uncertainty,
300 respectively) for U-ZK and R-PY, respectively. At the U-ZK, $P_{\text{ground+hv}}$ decreased the
301 most (84%), while at the R-PY, P_{nitrate} had the largest decrease about 85%, which was
302 speculated to be related to the decrease of NO_x and NO₃⁻ concentration in DC. Note that
303 the reduction rates in the overall known source and almost individual sources were
304 greater than the reduction rates in HONO concentrations (Figs. 7 and 8), thus we
305 hypothesized that there should be other sources capable of promoting HONO
306 production.

307 There were positive correlations between HONO with SO₂, Excess-NH_x, SO₄²⁻, and
308 pH (Fig. S12) indicating that the R₁ reaction might form an amount of HONO and
309 contribute to less reduction in the observed HONO concentrations. Considering that R₁
310 mainly reacts in the liquid phase, the calculated reaction rates of R₁ under the conditions

311 of RH > 60% in the PC and DC periods are illustrated in Figs. 8 and S12. Despite the
312 decrease in NO₂ and SO₂ concentrations in the DC, the increase in particle pH,
313 increasing HSO₃⁻ concentration in the aqueous phase, promoted the R₁ reaction rates by
314 58% and 59% at U-ZK and R-PY (Figure 8), respectively. Consequently, the enhanced
315 R₁ reaction might prevent a large decrease in HONO (18% at U-ZK and 53% at R-PY)
316 under the conditions of a significant reduction in vehicle emissions and a decline of 66%
317 and 69% in NO₂ concentrations at U-ZK and R-PY, respectively.

318 **3.5 Uncertainty**

319 According to sensitivity tests of pH (Fig. S8) and R₁ (Fig. S12), pH increases with
320 the concentrations of cations (TNH_x, TNa, K⁺, Ca²⁺, and Mg²⁺) and OC increasing as
321 well as anions (TH₂SO₄, TNO₃, and Cl⁻) concentrations, T, and RH decreasing. R₁
322 reaction rate increases with the concentrations of AWC, NO₂, SO₂, pH, and pressure,
323 while increasing as well as T decreasing. Therefore, two extreme scenarios (i.e., the
324 maximum and minimum rate scenarios) were evaluated to estimate the uncertainty of
325 pH, and R₁ based on the measurement uncertainties at the U-ZK and R-PY sites. Figure
326 S13 suggests that the two extreme scenarios can lead to -10-7% and -71-125%
327 uncertainties at the U-ZK site and -10-7% and -78-123% uncertainties at the R-PY
328 site for pH and R₁, respectively. Even considering the above uncertainty in Fig. 8, it can
329 still be observed that during the DC period, the decrease in HONO was less than that
330 of NO₂, and the rate of the R₁ reaction increased.

331 Considering the conclusions of this study are based solely on observational data,

332 there are certain limitations. For example, only the changes in the R_1 reaction of $PM_{2.5}$
333 were calculated, without considering variations in components, pH values, and R_1
334 reaction rates of coarse particles. Additionally, although this study selected scenarios
335 with $RH > 60\%$ to calculate the R_1 reaction to ensure the presence of a liquid phase, it
336 is evident that this approach overlooks some R_1 reactions. Furthermore, due to
337 thermodynamic model calculations of pH values, changes in the mixed state of particle
338 components, and the omission of organic acids, alongside the absence of gaseous HNO_3
339 and HCl in this study, these factors may lead to inaccuracies in pH value simulations
340 and uncertainty in R_1 calculations (Pye et al., 2020; Haskins et al., 2018; Nah et al., 2018).
341 Therefore, there is a certain degree of uncertainty in the conclusions regarding the
342 growth of R_1 reactions in this paper. Nevertheless, by calculating the changes in R_1
343 reactions, this study provides a possible explanation for the relatively small decrease in
344 HONO during the epidemic period.

345 **4. Conclusions**

346 Elevated NH_3 concentration was observed during the COVID-19 pandemic at both
347 urban and rural sites in China. In addition to the rise in T and decrease in RH during the
348 COVID-19 pandemic, which favored the conversion of NH_4^+ to NH_3 , the significant
349 decrease in sulfate and nitrate concentrations led to the decline in Required- NH_x and
350 was beneficial to the particle-phase NH_4^+ partitioning to gas-phase NH_3 . Furthermore,
351 under the environmental conditions of increased anion concentrations (especially
352 sulfate and nitrate) and increased cation concentrations, the pH values increased by 0.5

353 and 0.3 at U-ZK and R-PY sites increased during the pandemic, respectively.
354 Consequently, the high pH values accelerated the formation rate of HONO through the
355 oxidation-reduction reaction of NO_2 with SO_2 (an increase of 58% at U-ZK and 59% at
356 R-PY, respectively), partially compensating for the decrease in HONO concentration
357 caused by the decline in vehicle emissions, NO_2 and NO_3 concentrations during the
358 COVID-19 pandemic.

359 **5. Implications**

360 HONO plays a crucial role as a precursor to OH radicals in the tropospheric
361 atmosphere (Xue, 2022). There have been significant observations of high HONO
362 concentrations in urban areas during the daytime, leading to a growing interest in
363 understanding its sources in atmospheric chemistry (Jiang et al., 2022; Xu et al., 2019).
364 The heterogeneous reaction mechanism of NO_2 on aerosol surfaces is currently the
365 focus of research on HONO sources, particularly in regions with elevated levels of
366 atmospheric particulate matter, where it could potentially become a major contributor
367 to HONO production (Zhang et al., 2022; Liao et al., 2021b). One of the pathways for
368 heterogeneous reactions on aerosol surfaces is the redox reaction of NO_2 with SO_2 .
369 However, the significance of this reaction in HONO production in the real atmosphere
370 is often overlooked, as it relies on the high pH of aerosols (Ge et al., 2019). In recent
371 years, there has been increasing attention on the enhancing effect of NH_3 on the redox
372 reaction, with laboratory experiments demonstrating its ability to generate substantial
373 amounts of HONO (Ge et al., 2019). This study highlights the importance of this

374 reaction based on actual atmospheric observations. Furthermore, numerous studies
375 have indicated that if control over NH₃ emissions continues to relax while SO₂ and NO₂
376 emissions decrease, the particle pH in future China is expected to rise steadily (Xie et
377 al., 2020; Song et al., 2019; Wang et al., 2020). Consequently, the redox reaction of NO₂
378 with SO₂ could become a significant source of HONO in China. Therefore, it is crucial
379 to coordinate the control of SO₂, NO_x, and NH₃ emissions to avoid a rapid increase in
380 the particle pH.

381

382 **Data availability:** All the data presented in this article can be accessed through
383 <https://zenodo.org/records/10273539>. (Zhang, 2023).

384

385 **Author contributions.** XZ Data Curation, Writing - Original Draft, Visualization.
386 LW, NW, SM, and DZ Investigation, Visualization, Data Curation. DZ, HZ, and MW
387 Investigation. SW Conceptualization, Data Curation, Supervision. RZ Data Curation,
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389

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391

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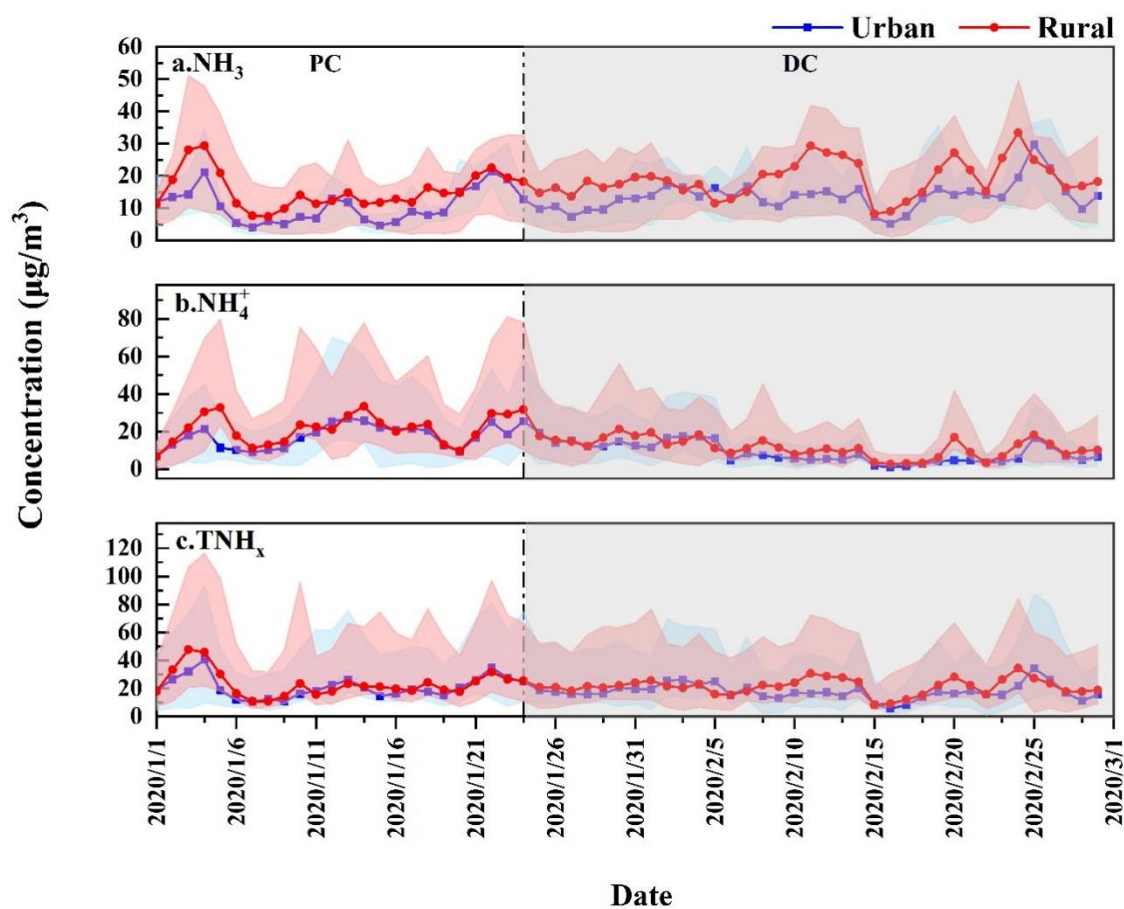
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706 **Figures:**

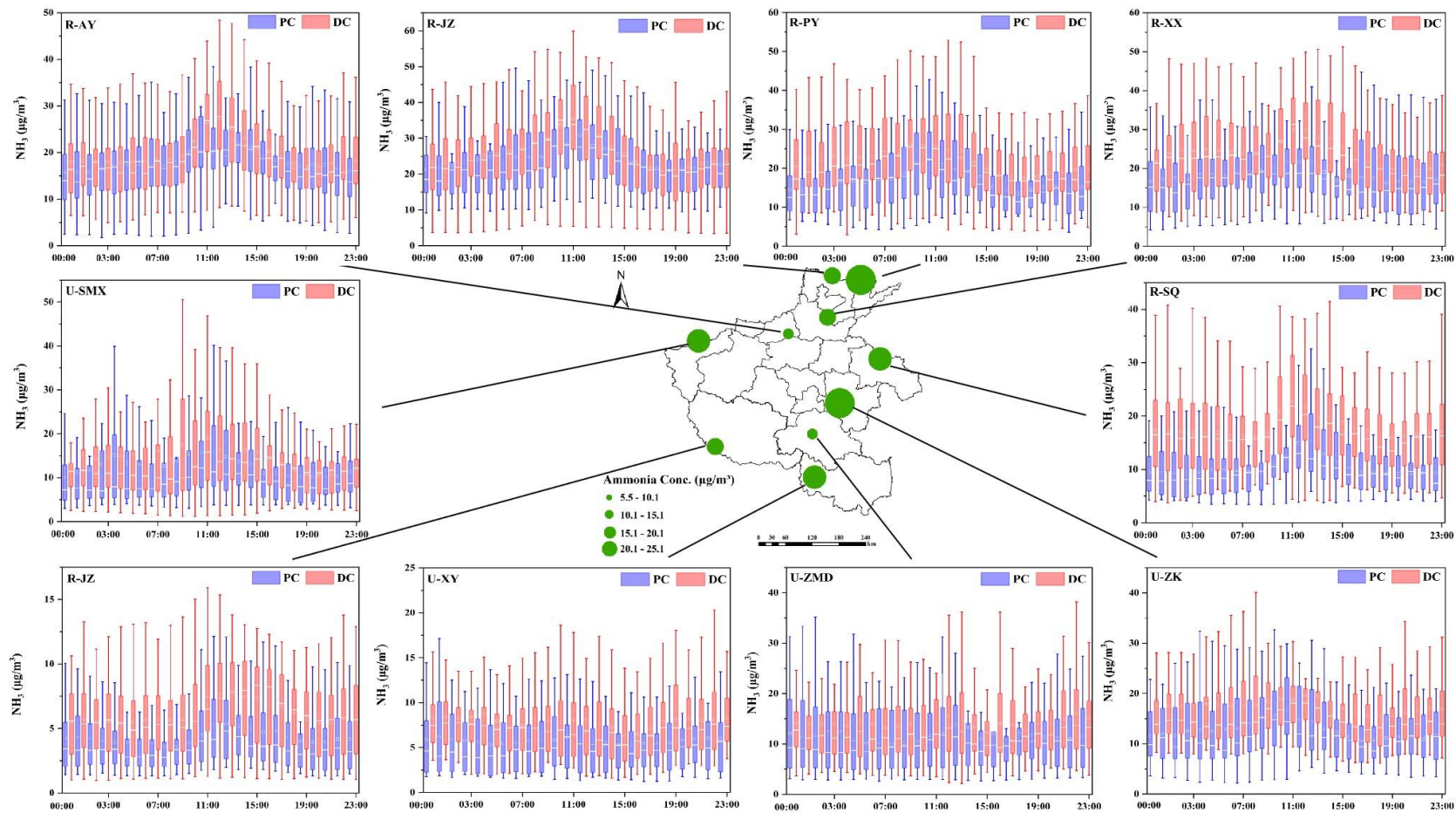


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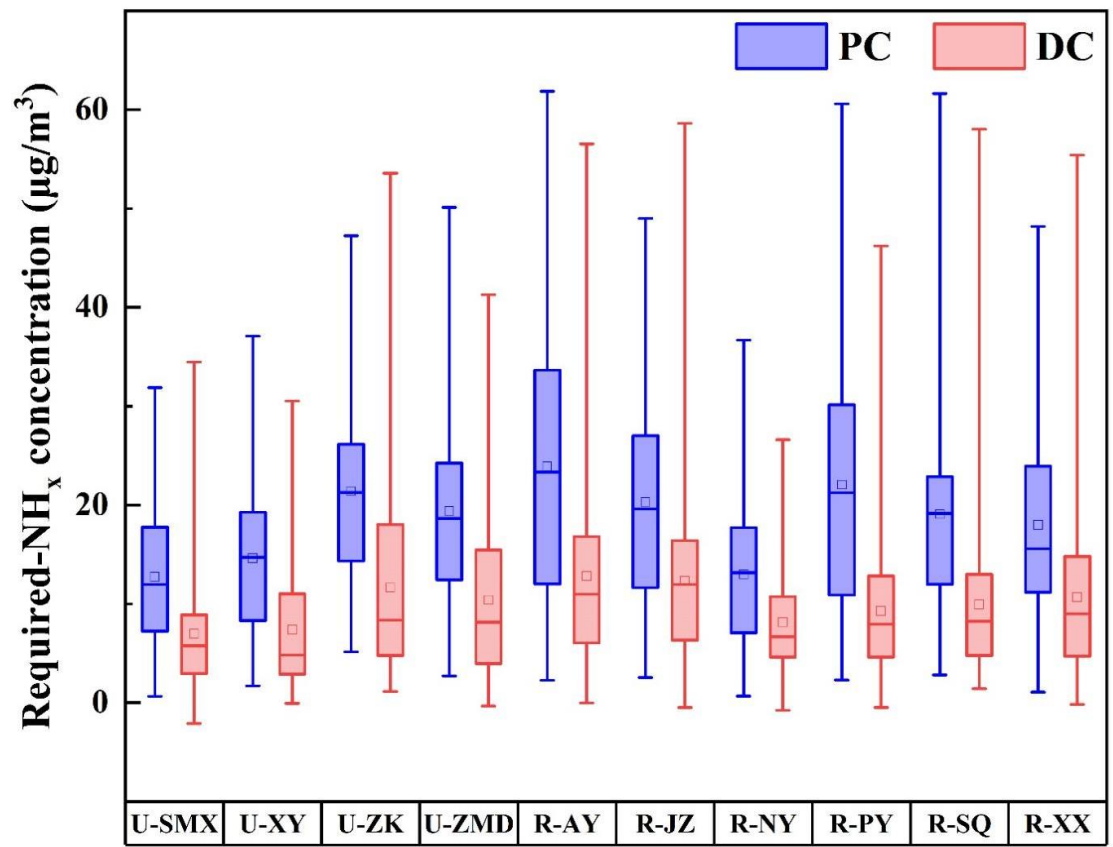
708 Figure 1. Temporal variations of a. NH₃, b. NH₄⁺, and c. TNH_x at the urban and rural

709 sites before (PC) and during (DC) the COVID-19 outbreak, respectively. The shaded

710 areas of the curve represent the maximum and minimum values.



711
 712 Figure 2. Daily variation of NH_3 concentrations at ten sites before (PC) and during (DC) the COVID-19 outbreak. The green dots represent the
 713 location of ten sites and their size represents the concentration of NH_3 ; In each box, the top, middle, and bottom lines represent the 75, 50, and 25
 714 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.



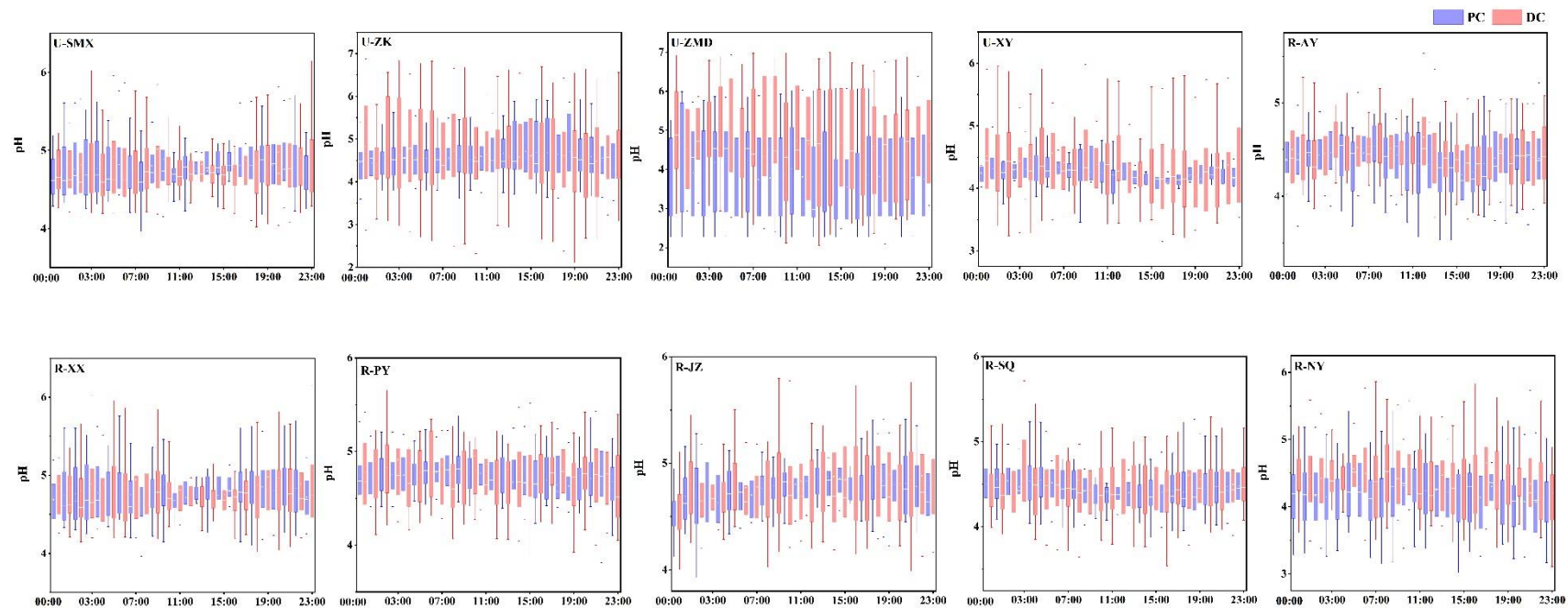
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716 Figure 3. Box diagram of changes in Required-NH_x at ten sites before (PC) and during

717 (DC) the COVID-19 outbreak. In each box, the top, middle, and bottom lines represent

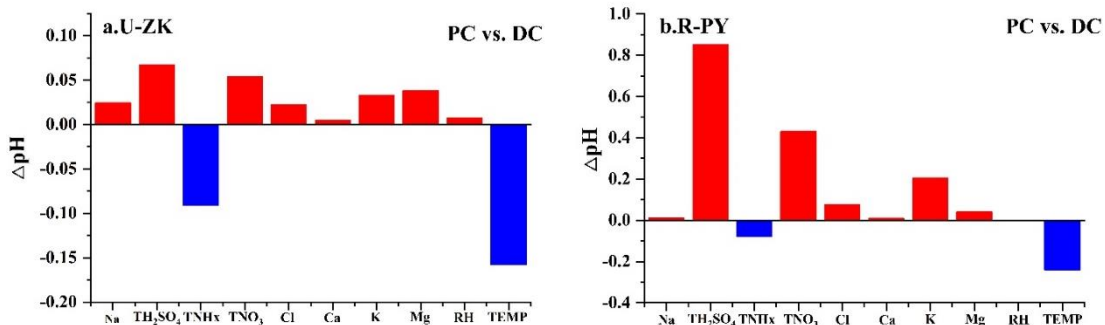
718 the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower

719 whiskers represent the 90 and 10 percentiles of statistical data, respectively.



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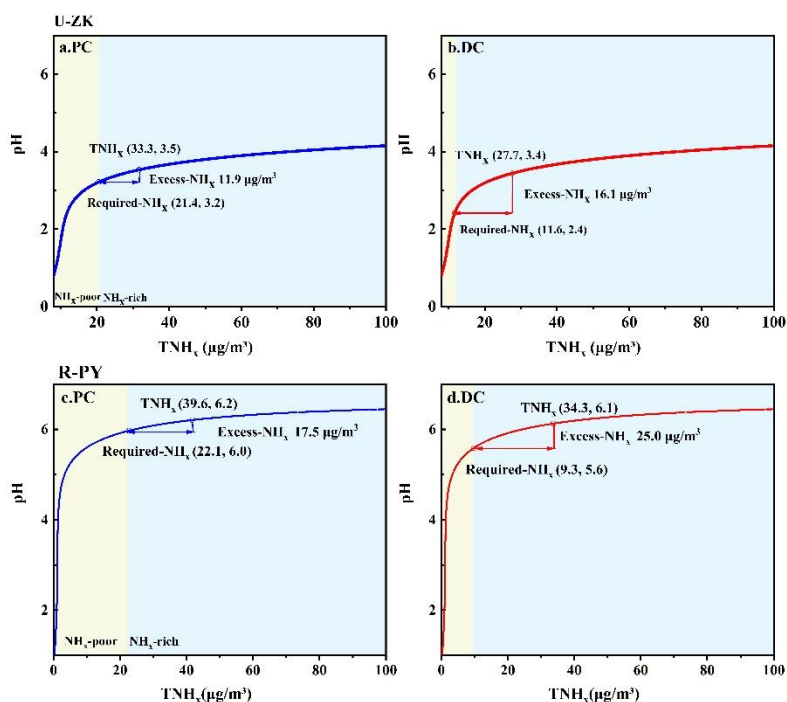
721 Figure 4. Diurnal patterns of pH at ten sites before (PC) and during (DC) the COVID-19 outbreak. In each box, the top, middle, and bottom
 722 lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of
 723 statistical data, respectively.



724

725 Figure 5. Changes of pH (ΔpH) through the sensitivity tests (Figure S5 and S6) by

726 changing parameters between PC and DC at the a. U-ZK and b. R-PY sites.



727

728 Figure 6. Particle pH corresponds to increasing TNH_x at U-ZK and R-PY sites to

729 examine the effects of major indicators of NH₃ (i.e., TNH_x, Required-NH_x, and Excess-

730 NH_x) on aerosol acidity. Particle pH was calculated by using a wide range of TNH_x

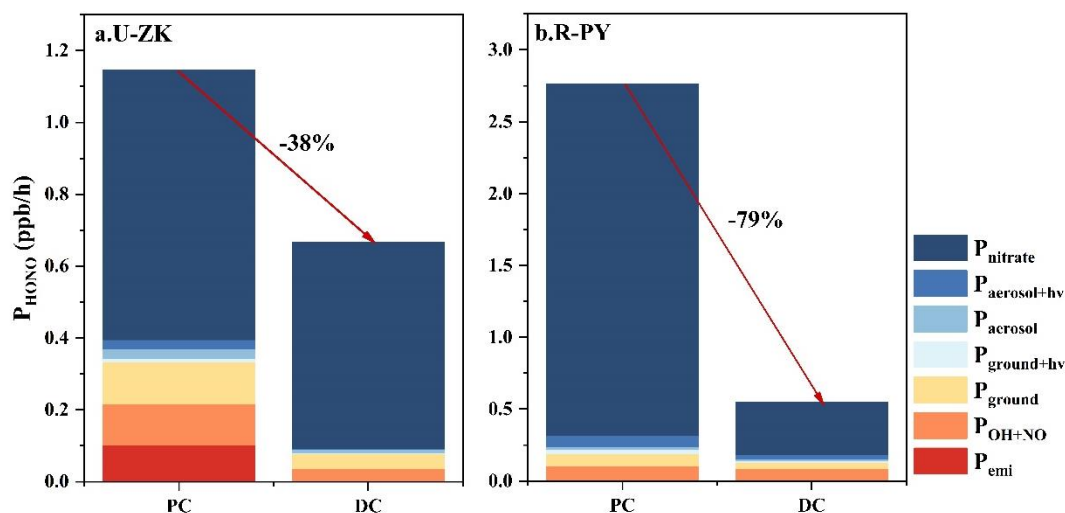
731 (25–130 μg/m³) and average values of other parameters in PC and DC of U-ZK and R-

732 PY sites. The concentrations of TNH_x, Required-NH_x, and Excess-NH_x with

733 corresponding pH values are marked by a hollow box, hollow circle, and arrow

734 respectively. The yellow and blue background colors correspond to the NH_x-poor and

735 NH_x-rich, respectively.

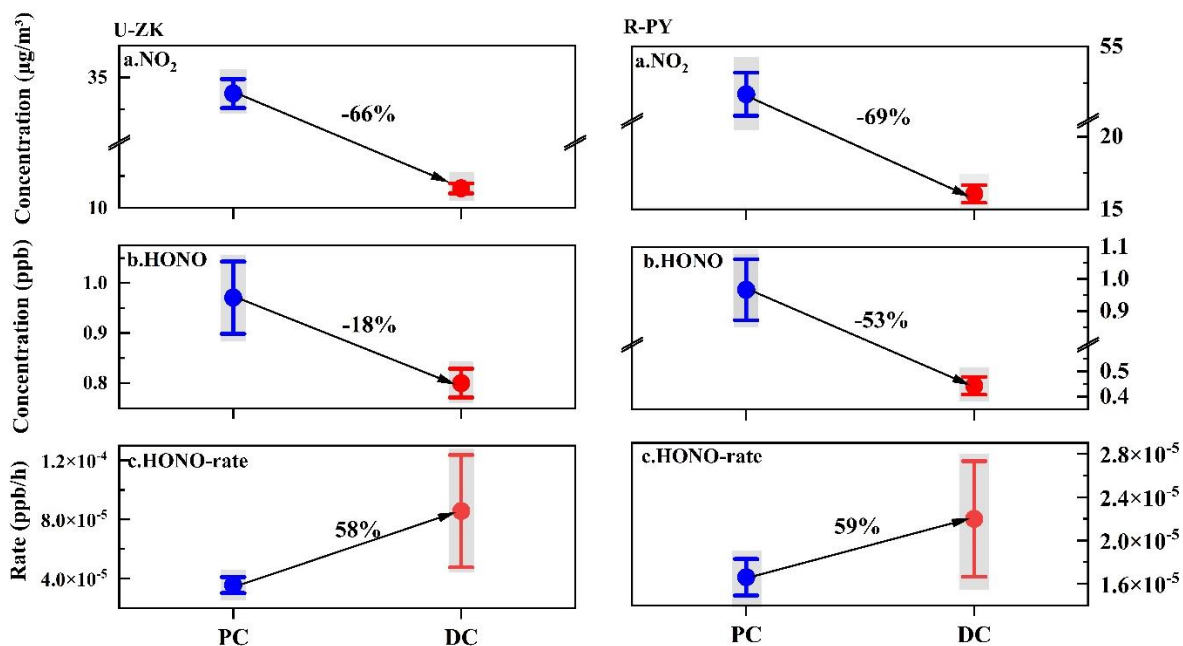


736

737 Figure 7. Comparison of HONO sources at a. U-ZK and b. R-PY sites before (PC) and

738 during (DC) the COVID-19 outbreak. The calculation method can be found in Text S4.

739



740

741 Figure 8. Decline ratios of a. NO_2 , b. HONO concentration, and c. HONO production

742 rate at U-ZK and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The

743 center point represents the mean value, and the upper and lower whiskers represent the

744 95% confidence interval of the mean. The shadows in the figure represent the

745 uncertainties of NO_2 measurement ($\pm 10\%$), HONO measurement ($\pm 20\%$), and the

746 HONO formation rate of R₁ reaction (−78–123%), respectively.

747

748 **Table:**

749 Table 1. Changes in concentrations (mean ± standard deviation) of NH₃, NH₄⁺, and

750 TNH_x at ten sites during entire periods (Average), before (PC), and during (DC) the

751 COVID-19 outbreak.

Sites	Substances	Average (µg/m ³)	PC (µg/m ³)	DC (µg/m ³)
U-SMX	NH ₃	13.8 ± 10.8	12.6 ± 10.1	14.5 ± 11.1
	NH ₄ ⁺	10.9 ± 7.2	14.2 ± 7.2	8.8 ± 6.5
	TNH _x	22.9 ± 14.1	24.9 ± 14.5	21.7 ± 13.8
U-ZK	NH ₃	15.6 ± 8.3	12.7 ± 6.5	17.4 ± 8.8
	NH ₄ ⁺	13.6 ± 9.3	19.1 ± 8.4	10.3 ± 8.1
	TNH _x	28.6 ± 13.7	30.9 ± 12.8	27.1 ± 14.0
U-ZMD	NH ₃	13.1 ± 8.4	11.6 ± 8.2	14.0 ± 8.4
	NH ₄ ⁺	13.9 ± 9.8	19.6 ± 10.3	10.3 ± 7.5
	TNH _x	25.7 ± 14.6	30.3 ± 15.1	22.8 ± 13.5
U-XY	NH ₃	7.0 ± 4.3	5.7 ± 4.0	7.9 ± 4.3
	NH ₄ ⁺	11.0 ± 7.7	15.4 ± 7.6	8.3 ± 6.5
	TNH _x	17.6 ± 9.8	20.6 ± 10.1	15.7 ± 9.2
R-AY	NH ₃	19.0 ± 8.4	17.9 ± 8.3	19.7 ± 8.4
	NH ₄ ⁺	19.3 ± 12.9	26.4 ± 13.7	15.0 ± 10.3
	TNH _x	36.6 ± 18.2	41.7 ± 20.4	33.4 ± 16.0
R-XX	NH ₃	21.7 ± 10.2	18.1 ± 9.3	23.8 ± 10.1
	NH ₄ ⁺	15.9 ± 10.4	20.6 ± 11.0	13.0 ± 8.8
	TNH _x	34.9 ± 17.0	35.1 ± 18.8	34.8 ± 15.8
R-PY	NH ₃	19.8 ± 9.4	16.8 ± 8.1	21.7 ± 9.6
	NH ₄ ⁺	17.4 ± 11.8	25.3 ± 12.6	12.4 ± 8.0
	TNH _x	35.2 ± 17.8	39.4 ± 19.8	32.6 ± 15.7
R-JZ	NH ₃	25.3 ± 11.5	24.1 ± 11.5	25.9 ± 11.4
	NH ₄ ⁺	17.3 ± 11.3	22.7 ± 11.6	14.2 ± 9.9
	TNH _x	40.8 ± 20.1	42.9 ± 22.8	33.5 ± 18.2
R-SQ	NH ₃	15.0 ± 7.9	10.3 ± 5.2	17.7 ± 7.9
	NH ₄ ⁺	13.4 ± 8.5	18.9 ± 8.6	10.3 ± 6.7
	TNH _x	26.3 ± 13.2	25.5 ± 14.0	26.8 ± 12.7
R-NY	NH ₃	5.5 ± 3.1	4.3 ± 2.7	6.2 ± 3.2
	NH ₄ ⁺	10.2 ± 6.9	13.3 ± 7.2	8.4 ± 6.1
	TNH _x	14.8 ± 8.5	16.0 ± 9.5	14.1 ± 7.8

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