- **Measurement Report: Elevated atmospheric ammonia**
- 2 may promote the particle pH and HONO formation:
- 3 Insights from the COVID-19 pandemic
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Abstract

HONO plays a crucial role as a precursor to OH radicals in the tropospheric atmosphere.
The incongruity between HONO concentration and NO _x emissions during the COVID-
19 pandemic remains puzzling. Here, we show evidence from field observations of ten
sites in China that there was a noticeable increase in NH3 concentrations during the
COVID-19 pandemic. In addition to the meteorological conditions, the significant
decrease in sulfate and nitrate concentrations enhanced the portioning of NH_4^+ to NH_3 .
Sensitivity analysis indicated that the decrease in anion concentrations (especially
sulfate and nitrate) and the increase in cation concentrations during the COVID-19
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1. Introduction

36	Nitrous acid (HONO) is a critical precursor of hydroxyl radical (OH), contributing
37	to more than 60% of OH production (Alicke, 2003;Platt et al., 1980;Kleffmann et al.,
38	2005). The OH can react with carbon monoxide, nitrogen oxides (NO _x), sulfur dioxide
39	(SO ₂), and volatile organic compounds to produce secondary pollutants such as ozone
40	(O ₃) and PM _{2.5} (particulate matter with an aerodynamic diameter less than or equal to
41	$2.5~\mu m$), thereby affecting air quality, human health, and global climate change (Li et
42	al., 2021a; Wang et al., 2023b; Lu et al., 2018)
43	High concentrations of HONO are present in urban daytime atmospheres, and
44	exploring its sources has become a hot and challenging topic in the field of atmospheric
45	chemistry (Jiang et al., 2022;Xu et al., 2019). Various sources of atmospheric HONO
46	have been identified, including combustion processes (e.g., vehicle emissions) (Kramer
47	et al., 2020;Liao et al., 2021a;Li et al., 2021b), direct emissions from soil (Su and Zhang
48	2011;Oswald et al., 2013;Meusel et al., 2018), homogeneous reactions between NO and
49	OH radicals (Pagsberg, 1997;Atkinson and Rossi, 2004), heterogeneous reactions of
50	NO ₂ on aerosols and ground surfaces (Zhang et al., 2020a;McFall et al., 2018;Liu et al.,
51	2014;Liu et al., 2020a), and photolysis of nitrate (Spataro and Ianniello, 2014;Scharko
52	et al., 2014;Romer et al., 2018;Ye et al., 2017;Shi et al., 2021). During the pandemic
53	control periods, there was a substantial reduction in vehicle traffic flow and industrial
54	emissions, leading to a decrease of more than 60% in NO _x emissions in eastern China
55	(Huang et al., 2021a). It was initially expected that the concentration of HONO would
56	also decrease proportionally. However, Liu et al. (2020b) observed that the decrease in

HONO concentration during the pandemic period was only 31% (from 1.5 ppb to 0.9) ppb), which was significantly lower than the reductions in NO (62%, from 26.3 to 4.2 ppb) and NO₂ (36%, from 15.5 to 6.2 ppb). Furthermore, the observed concentrations of HONO during the COVID-19 pandemic in 2020 were higher than those during the corresponding period in 2021 in Beijing (Luo et al., 2023). These findings suggest the existence of a considerable unknown source of HONO during the COVID-19 pandemic period. Ammonia (NH₃) is a primary alkaline gas in the atmosphere, capable of influencing the pH level of particulate matter and plays a crucial role in the atmospheric nitrogen cycle (Gu et al., 2022; Xu et al., 2020; Gong et al., 2011). Several studies have indicated that NH₃ can promote the formation of HONO by promoting the hydrolysis of NO₂ (Xu et al., 2019) or the redox reaction of NO₂ with SO₂ (Liu et al., 2023). Moreover, previous studies have reported that NH₃ concentrations in the atmosphere, particularly in rural areas, significantly increased during the pandemic (Xu et al., 2022). Consequently, the rise in NH₃ may contribute to the enhanced formation of HONO (Huang et al., 2021a). Unfortunately, there is currently a lack of research on the relationship between enhanced NH₃ and HONO during the COVID-19 pandemic period. To address this, online observational data on the chemical composition of PM_{2.5}, gaseous pollutants, and meteorological conditions at ten sites in China before and during the COVID-19 pandemic period were analyzed to investigate the variation in NH₃ concentrations and particle pH, and explore the promoting effect of increased pH values on HONO formation.

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2. Materials and methods

2.1 Observation sites

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- Online measurements were conducted at four urban and six rural sites in Henan Province, China from January 1 to February 29, 2020, including Sanmenxia (U-SMX), Zhoukou (U-ZK), Zhumadian (U-ZMD), and Xinyang (U-XY), as well as rural locations including Anyang (R-AY), Xinxiang (R-XX), Jiaozuo (R-JZ), Shangqiu (R-
- 85 SQ), Nanyang (R-NY), and Puyang (R-PY). Descriptions and the spatial distribution
- of these ten sites can be found in Table S1 and Fig. S1.

2.2 Measurements

The aerosol and gas monitor (MARGA, Metrohm, Switzerland) was used to analyze the hourly water-soluble ions (Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Cl⁻, NO₃, and SO₄²) in PM_{2.5}, as well as gaseous species (NH₃, HNO₃, HCl, and HONO) at ten sampling sites. The MARGA instrument is widely used (Chen et al., 2017;Stieger et al., 2019;Twigg et al., 2022). A detailed description of the instrument and QA/QC can be found in Text S1. In brief, the atmospheric sample passes through a PM_{2.5} cut-off head, and both particles and gases enter a wet rotating dissolution device for diffusion. Subsequently, the particles in the sample undergo hygroscopic growth and condensation in an aerosol supersaturated vapor generator, followed by collection and ion chromatographic analysis. The gases in the sample are oxidized by H₂O₂ in the dissolution device, absorbed into a liquid solvent, and then entered the gas sample collection chamber for

99 ion chromatographic quantification. The range of minimum detection limits for watersoluble ions was between 0.002 μg/m³ (Cl⁻) to 0.081 μg/m³ (NH₄). Uncertainties of 20% 100 are assumed for the detection of NH₃ and NH₄, while uncertainties of 10% are assumed 102 for other components (Wang et al., 2020; Wang et al., 2022). In addition, a detailed 103 description of HONO measurement using this system can be found in Text S2. Overall, 104 the limit of detection for HONO was 4 pptv and the uncertainty was estimated to be \pm 105 20%. The data for NO₂ and SO₂ were obtained from a series of instruments provided by 106 107 Thermo Fisher Scientific (USA). The hourly concentrations of organic carbon (OC) in PM_{2.5} were analyzed using a carbon analyzer (Model 4, Sunset Laboratory., USA). 108 Detailed descriptions of the NO₂, SO₂, and carbon analyzers can be found in Text S3. 109 110 The smart weather stations (LUFFTWS500, Sutron, Germany) were utilized for synchronized observation of meteorological parameters including pressure,

2.3 Data analysis.

temperature (T), and relative humidity (RH).

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2.3.1 pH prediction.

115 The thermodynamic model ISORROPIA-II was used to estimate the pH value of the particles (Fountoukis, 2007) by inputting RH, T, K⁺, Ca²⁺, Mg²⁺, total ammonia 116 $(TNH_x = 17 \times (\frac{[NH_4^+]}{18} + \frac{[NH_3]}{17}))$, total sulfuric acid (TH_2SO_4, SO_4^{2-}) , total sodium 117 (TNa, Na⁺), total chlorine (TCl, Cl⁻), and total nitrate (TNO₃ = NO₃ + HNO₃). The 118

model has two calculation modes: the forward mode and reverse mode, and the aerosol dissolution systems can be set to simulate a metastable state (aqueous phase) or stable state (aqueous and solid phase). Studies have shown that the forward mode is less affected by instrument measurement errors than the reverse mode (Ding et al., 2019;Song et al., 2018). Additionally, the minimum average RH of about 55% was recorded during the sampling period at the ten sites. Thus, ISORROPIA-II was run in the forward model for the aerosol system in the metastable condition and only used data with RH \geq 30% for simulation accuracy (Ding et al., 2019;Song et al., 2018). The ISORROPIA model calculated the particle hydrate ion concentration per volume of air (Hair⁺) and aerosol water associated with inorganic matter (AWC_{inorg}). The pH value was calculated using the following equation (Bougiatioti et al., 2016):

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$$pH = -\log_{10}H_{aq}^{+} = -\log_{10}\frac{1000H_{air}^{+}}{AWC_{inorg} + AWC_{org}}$$
(2.1)

where the modeled concentrations for AWC_{inorg} and H_{air}^{+} are $\mu g/m^3$, and AWC_{org} is the particle water associated with the organic matters predicted using the following method:

$$AWC_{org} = \frac{m_s}{\rho_s} \frac{k_{org}}{\left(\frac{1}{RH} - 1\right)}$$
 (2.2)

where m_s is the mass concentration of organic matter (OM = OC × f). The f is the conversion factor of OC, which is dependent on the extent of OM oxidation and secondary organic aerosol formation (Chow et al., 2015). Studies on the ratio of OM/OC in fourteen cities in China suggested that the mean value of f was 1.59 ± 0.18 during the winter season in Northern China (Xing et al., 2013), and thus we adopted 1.6 as the f in this study. k_{org} is the organic hygroscopicity parameter and depends on organic

functionality, water solubility, molecular weight, and oxidation level. Han et al. (2022) have reported that the k_{org} generally increased with O: C ratios, with a range of 0–0.3 for 23 organics, including carboxylic acids, amino acids, sugars, and alcohols. Gunthe et al. (2011) estimated a $k_{org} = 0.06 \pm 0.01$ for the effective average hygroscopicity of the non-refractory organic particulate matter in the aerosols in Beijing. Our previous study has estimated that the uncertainties of k_{org} value (0.06) for pH in the range of 0–0.3 only lead to –1–3% errors, which can be neglected (Wang et al., 2023a). Therefore, the value of 0.06 was selected in this paper. ρ_s is the organic density, which was chosen to be 1.35 g/cm³ following previous studies (Table S2).

2.3.2 The sources of HONO

The sources of HONO include direct emission (P_{emi}), the homogeneous reaction of NO and •OH (P_{OH+NO}), the heterogeneous reaction of NO₂ on the ground (P_{ground}) and aerosol (P_{aerosol}), the photo-enhanced heterogeneous reaction of NO₂ on the ground (P_{ground+hv}) and aerosol (P_{aerosol+hv}), and nitrate photolysis (P_{nitrate}). The detailed calculation method is described in the Supplementary Material (Text S4, Table S3, Figs. S2 and S3).

Soil emission has been demonstrated to be a major source of HONO, which is affected by temperature to some extent (Liu et al., 2020c; Liu et al., 2020b). However, during the sampling periods, there was no significant positive correlation between HONO concentration and temperature (Fig. S4). In addition, temperatures did not exceed 10°C, under which the soil HONO emission rate is generally considered to be zero (Zhang et

al., 2023). Furthermore, the equilibrium gas-phase concentration over an aqueous solution of nitrous acid, [HONO]*, a key parameter controlling the exchange of HONO 162 between the gas and aqueous phase in soil, is calculated according to Su et al.(2011). 163 The results indicate that the temperature difference between PC and DC periods only 164 led to approximately a 0.01% concentration change. On the other hand, studies on the 165 166 sources of HONO in the North China Plain of China during winter consistently showed that soil HONO emissions contribute around 1% (Liu et al., 2020c;Liu et al., 167 2020b; Zhang et al., 2023). Therefore, this study does not consider soil HONO 168 169 emissions.

2.3.3 Redox reaction of NO₂ with SO₂.

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- 171 The redox reaction of NO₂ with SO₂ (R₁) is considered a crucial potential source of
- high concentrations of HONO in Northern China (Cheng et al., 2019; Wang et al., 2016): 172

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$$S(IV) + 2NO_2 + H_2O \rightarrow S(VI) + 2H^+ + 2NO_2^-$$
 (R₁)

the rate expression for reaction (R_1) was estimated to: 174

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$$d[S(VI)] / dt = k_1[NO_2][S(VI)], \qquad (2.3)$$

- the rate constant k_1 value is pH dependent, e.g., for pH, 5, $k_1 = (1.4 \times 10^5 + 1.24 \times 10^7)/2$ 176
- M⁻¹ s⁻¹. For k₁ values under other pH conditions and other related information, please 177
- 178 refer to Text S5, Table S4, and Table S5.

3. Results and discussion

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3.1 Variations of NH₃, NH₄ and TNH_x.

181 The temporal variations of NH₃, NH₄, and TNH_x at 10 sampling sites in the pre-COVID-19 pandemic period (PC, January 1 to 23, 2020) and during the COVID-19 182 183 pandemic period (DC, January 24 to February 29, 2020) are presented in Fig. 1, with 184 their average concentration listed in Table 1. In general, rural sites exhibited higher 185 concentrations of NH₃, NH₄, and TNH_x compared to urban sites, except for the R-NY 186 site. This finding is consistent with previous studies conducted in Zhengzhou (Wang et 187 al., 2020), Shanghai (Chang et al., 2019), and Quzhou (Feng et al., 2022), owing to the intense agricultural ammonia emissions. The highest concentrations of NH₃ and TNH_x 188 189 were recorded at site R-JZ, with average values of 25.3 ± 11.5 and 40.8 ± 20.1 µg/m³, respectively. Site R-AY had the highest NH₄ concentration, measuring 19.3 ± 12.9 190 191 μg/m³. Note that the current study area exhibited higher NH₃ levels compared to other 192 regions (Table S6), which probably was attributed to the highest NH₃ emissions of 193 Henan Province in China, primarily from nitrogen fertilizer application and livestock 194 farming (Wang et al., 2018; Ma, 2020). 195 Compared to the PC, NH₃ concentrations increased in the DC at all sites. Notably, 196 rural sites experienced more significant increases in NH₃ concentrations than urban 197 sites, which was similar to the trend in Shanghai (Xu et al., 2022). The largest increases in NH₃ concentrations were observed at R-SQ (71%, 7.3 µg/m³) and U-ZK (37%, 4.8 198 μg/m³) for rural and urban sites, respectively. In contrast, the concentrations of NH₄ 199

and TNH_x decreased in the DC with the largest reduction at rural site R-PY (51%, 12.9 $\mu g/m^3$) and urban site U-ZMD (48%, 9.3 $\mu g/m^3$). Regarding TNH_x, rural sites exhibited larger reductions, with site R-SQ experiencing the largest decrease of 37% (4.7 µg/m³). Figure 2 illustrates the spatial distribution and the diurnal variation of NH₃ concentrations at the ten sites before and during the pandemic. NH₃ concentrations in most sites exhibited an unimodal trend in PC that NH₃ concentrations gradually increased after sunrise, reaching a peak around noon (11:00–12:00), and then decreased to a valley around 16:00–17:00. This diurnal pattern is similar to NH₃ variations observed in urban areas of Houston, USA, as a result of the natural emissions from vegetation and soil during photosynthesis (Gong et al., 2011). However, other studies have recorded a significant NH₃ peak during the early morning of 8:00–10:00 (Ellis et al., 2011; Meng et al., 2018; Gu et al., 2022), suggesting the influence of vehicle emissions (Gong et al., 2011; Gu et al., 2022), residual NH₃ mixing, soil or plant emissions (Ellis et al., 2011), and dew volatilization (Wentworth et al., 2016; Huang et al., 2021b). Therefore, the NH₃ in urban and rural areas of this study was probably less affected by NH₃ emissions from vehicles, different from the recent studies in megacities of China (e.g., Beijing and Shanghai) (Gu et al., 2022; Wu et al., 2023; Zhang et al., 2020b). In addition to the transport from agricultural emissions, urban NH₃ in this region might also originate from other non-agricultural sources, such as wastewater treatment, coal combustion, household waste, urban green spaces, and human excrement (Chang et al., 2019).

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During the COVID-19 pandemic, the diurnal variation of NH₃ in both urban and rural

sites still maintained an unimodal distribution. The peak values in urban sites remained consistent with PC levels, further demonstrating that the influence of vehicles on NH₃ in urban areas was limited. Notably, the peak time of NH₃ in rural sites shifted 1–2 hours earlier compared to the trend in PC. Ammonia in rural areas primarily originates from nitrogen fertilizer application, livestock, and poultry breeding (Feng et al., 2022;Meng et al., 2018), which are significantly influenced by T and RH (Liu et al., 2023). Table S7 and Fig. S5 reveal that there was an increased T and a decreased RH at rural sites in the DC than the PC, which could accelerate the evaporation of NH₃ and thus potentially lead to earlier peak NH₃ concentrations.

3.2 Gas-to-particle conversion of NH₃

The increased NH₃ accompanying decreased NH₄ in the DC suggests that the gas-particle partition of NH₃/ NH₄ may determine the elevated NH₃ concentrations. Meteorological parameters, including RH and T, play a crucial role in the gas-particle partitioning of NH₃ (Liu et al., 2023; Xu et al., 2020). Therefore, the higher T and lower RH in the DC (Table S7 and Fig. S5) favored the conversion of NH₄ to NH₃, resulting in a decrease in $\varepsilon(NH_4^+)$ ($[NH_4^+]/([NH_3] + [NH_4^+])$ compared to those in the PC (Table S7). NH₃ primarily enters particles to neutralize acidic ions (Wang et al., 2020;Xu et al., 2020; Liu et al., 2017; Ye et al., 2011; Wells, 1998). Accordingly, the concentrations of required ammonia (Required-NH_x) and excess ammonia (Excess-NH_x) were calculated based on the acidic substances as follows (Wang et al., 2020):

Required-NH_x = 17 ×
$$(\frac{[SO_4^{2-}]}{48} + \frac{[NO_3^-]}{63} + \frac{[Cl^-]}{35.5} + \frac{[HNO_3]}{64} + \frac{[HCl]}{36.5})$$

$$-17 × (\frac{[Na^+]}{23} + \frac{[K^+]}{39} + \frac{[Ca^{2+}]}{20} + \frac{[Mg^{2+}]}{12})$$
(3.1)

$$Excess-NH_{x} = TNH_{x} - Required-NH_{x}$$
 (3.2)

where [W] represents the concentration of the substance ($\mu g/m^3$). The significant linear fitting (R^2 is greater than 0.96, and the slope is close to 1) in Fig. S6 demonstrates that the anions and cations at each site were close to the equilibrium state. Therefore, the organic acids in $PM_{2.5}$ may have less effect on NH_3 and NH_4^+ and were not considered in Formula 3.1.

As shown in Fig. 3 and Table S8, compared to those in the PC, the concentration of Required-NH_x in the DC significantly decreased (ranging from 37% at site R-JZ to 58% at site R-PY), while the concentration of Excess-NH_x increased (ranging from 9% at site R-AY to 78% at site R-SQ). The reduction in the concentrations of sulfate and nitrate (Fig. S7) was responsible for the decrease in the concentration of Required-NH_x. To sum up, in addition to meteorological conditions, the substantial reduction in anthropogenic emissions of SO₂, NO_x, and other pollutants in the DC had led to a decrease in acidic substances (e.g., sulfate and nitrate) in particles, in turn, resulting in more gas-phase NH₃ concentration remaining in the atmosphere.

3.3 Particle pH before and during COVID-19

Diurnal patterns of particle pH in PC and DC at ten sites are summarized in Fig. 4 with their average values listed in Table S9. PM_{2.5} shows consistent moderate acidity, with mean values in the range of 4.2–5.1, which were close to the values in previous

263 studies (Table S9). Compared to the PC, the particle pH at ten sites increased obviously in the DC, with the highest increase of 0.5 (U-ZK) and 0.3 (R-PY) at urban and rural 264 265 sites, respectively, which were the subject of an in-depth discussion in the following 266 text. 267 To explore the dominant factors that determine the local particle pH level and result 268 in the high pH during the DC, sensitivity tests of pH to chemical species (i.e., TNH_x, TH₂SO₄, TNO₃, TCl, TNa, K⁺, Ca²⁺, and Mg²⁺) and meteorological parameters (i.e., T 269 and RH) were performed. A given range for a variable (i.e., TNH_x) with corresponding 270 average values of other parameters (i.e., TH₂SO₄, TNO₃, TCl, TNa, K⁺, Ca²⁺, Mg²⁺, T, 271 272 and RH) was input into the model and simulated to compare its effects on pH. As shown in Fig. S8, pH increases with the cation concentrations (i.e., TNH_x, Na⁺, K⁺, Ca²⁺, and 273 Mg²⁺) increasing as well as the anion concentrations (i.e., TH₂SO₄, TNO₃, and Cl⁻), T 274 275 and RH decreasing. According to the average values of input data during PC (Blue line in Fig. S8) and DC (Red line in Fig. S8) at U-ZK and R-PY sites respectively, the 276 277 changes in pH (Δ pH in Fig. 5) indicate that the decrease in TNH_x concentration and the 278 increase in T in DC led to a decrease in pH values (\Delta pH: 0.09 at U-ZK and 0.08 at R-PY sites) compared to PC. However, this effect was outweighed by the decrease in 279 280 TH₂SO₄ (ΔpH: 0.07 and 0.8 at U-ZK and R-PY sites, respectively) and TNO₃ (ΔpH: 0.05 and 0.4 at U-ZK and R-PY sites, respectively) concentrations as well as the 281 increase in K⁺ (Δ pH: 0.03 at U-ZK and 0.2 at R-PY site) and Mg²⁺ (Δ pH: 0.01 at U-ZK 282 and 0.04 at R-PY site) concentrations in the DC, and resulting in an overall increase in 283 pH values in the DC. Furthermore, the relationship between particle pH with the 284

concentrations of Required-NH_x, and Excess-NH_x, which considers all chemical components, is investigated to examine the dominant factor on the increasing pH in DC. As shown in Fig. 6, the higher Excess-NH_x concentrations in the DC led to higher increases in pH values (ΔpH: 1 at U-ZK and 0.5 at R-PY site) than those in PC (ΔpH: 0.3 at U-ZK and 0.2 at R-PY site), thus Excess-NH_x concentrations may be the key factor in promoting the pH values.

3.4 The influence of pH on HONO.

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The observed HONO concentrations decreased by 18% and 54% at U-ZK (0.8 ppb) and R-PY (0.9 ppb) sites in the DC, respectively, compared to those (1.0 and 2.2 ppb) in the PC. Moreover, all the known HONO production sources rates including P_{emi}, P_{OH} + NO, Pground, Pground+hy, Paerosol, Paerosol+hy, and Pnitrate (Fig. 7, Fig S9 and S10) show a decreasing trend from PC to DC, with the total reductions of 38% (from 30% to 45%) in the scenario with the minimum and maximum uncertainty, respectively) and 79% (from 77% to 82% in the scenario with the minimum and maximum uncertainty, respectively) for U-ZK and R-PY, respectively. At the U-ZK, $P_{ground+hv}$ decreased the most (84%), while at the R-PY, P_{nitrate} had the largest decrease about 85%, which was speculated to be related to the decrease of NO_x and NO₃ concentration in DC. Note that the reduction rates in the overall known source and almost individual sources were greater than the reduction rates in HONO concentrations (Figs. 7 and 8), thus we hypothesized that there should be other sources capable of promoting HONO production. There were positive correlations between HONO with SO₂, Excess-NH_x, SO₄²⁻, and pH (Fig. S12) indicating that the R₁ reaction might form an amount of HONO and contribute to less reduction in the observed HONO concentrations. Considering that R₁ mainly reacts in the liquid phase, the calculated reaction rates of R₁ under the conditions of RH > 60% in the PC and DC periods are illustrated in Figs. 8 and S12. Despite the decrease in NO₂ and SO₂ concentrations in the DC, the increase in particle pH, increasing HSO₃ concentration in the aqueous phase, promoted the R₁ reaction rates by 58% and 59% at U-ZK and R-PY (Figure 8), respectively. Consequently, the enhanced R₁ reaction might prevent a large decrease in HONO (18% at U-ZK and 53% at R-PY) under the conditions of a significant reduction in vehicle emissions and a decline of 66% and 69% in NO₂ concentrations at U-ZK and R-PY, respectively.

3.5 Uncertainty

According to sensitivity tests of pH (Fig. S8) and R_1 (Fig. S12), pH increases with the concentrations of cations (TNH_x, TNa, K⁺, Ca²⁺, and Mg²⁺) and OC increasing as well as anions (TH₂SO₄, TNO₃, and Cl⁻) concentrations, T, and RH decreasing. R_1 reaction rate increases with the concentrations of AWC, NO₂, SO₂, pH, and pressure, while increasing as well as T decreasing. Therefore, two extreme scenarios (i.e., the maximum and minimum rate scenarios) were evaluated to estimate the uncertainty of pH, and R_1 based on the measurement uncertainties at the U-ZK and R-PY sites. Figure S13 suggests that the two extreme scenarios can lead to -10-7% and -71-125% uncertainties at the U-ZK site and -10-7% and -78-123% uncertainties at the R-PY site for pH and R_1 , respectively. Even considering the above uncertainty in Fig. 8, it can still be observed that during the DC period, the decrease in HONO was less than that of NO₂, and the rate of the R_1 reaction increased.

Considering the conclusions of this study are based solely on observational data,

there are certain limitations. For example, only the changes in the R₁ reaction of PM_{2.5} were calculated, without considering variations in components, pH values, and R₁ reaction rates of coarse particles. Additionally, although this study selected scenarios with RH > 60% to calculate the R₁ reaction to ensure the presence of a liquid phase, it is evident that this approach overlooks some R₁ reactions. Furthermore, due to thermodynamic model calculations of pH values, changes in the mixed state of particle components, and the omission of organic acids, alongside the absence of gaseous HNO₃ and HCl in this study, these factors may lead to inaccuracies in pH value simulations and uncertainty in R₁ calculations(Pye et al., 2020;Haskins et al., 2018;Nah et al., 2018). Therefore, there is a certain degree of uncertainty in the conclusions regarding the growth of R₁ reactions in this paper. Nevertheless, by calculating the changes in R₁ reactions, this study provides a possible explanation for the relatively small decrease in HONO during the epidemic period.

4. Conclusions

Elevated NH₃ concentration was observed during the COVID-19 pandemic at both urban and rural sites in China. In addition to the rise in T and decrease in RH during the COVID-19 pandemic, which favored the conversion of NH₄⁺ to NH₃, the significant decrease in sulfate and nitrate concentrations led to the decline in Required-NH_x and was beneficial to the particle-phase NH₄⁺ portioning to gas-phase NH₃. Furthermore, under the environmental conditions of increased anion concentrations (especially sulfate and nitrate) and increased cation concentrations, the pH values increased by 0.5

and 0.3 at U-ZK and R-PY sites increased during the pandemic, respectively. Consequently, the high pH values accelerated the formation rate of HONO through the oxidation-reduction reaction of NO₂ with SO₂ (an increase of 58% at U-ZK and 59% at R-PY, respectively), partially compensating for the decrease in HONO concentration caused by the decline in vehicle emissions, NO₂ and NO₃ concentrations during the COVID-19 pandemic.

5. Implications

HONO plays a crucial role as a precursor to OH radicals in the tropospheric atmosphere (Xue, 2022). There have been significant observations of high HONO concentrations in urban areas during the daytime, leading to a growing interest in understanding its sources in atmospheric chemistry (Jiang et al., 2022;Xu et al., 2019). The heterogeneous reaction mechanism of NO₂ on aerosol surfaces is currently the focus of research on HONO sources, particularly in regions with elevated levels of atmospheric particulate matter, where it could potentially become a major contributor to HONO production (Zhang et al., 2022;Liao et al., 2021b). One of the pathways for heterogeneous reactions on aerosol surfaces is the redox reaction of NO₂ with SO₂. However, the significance of this reaction in HONO production in the real atmosphere is often overlooked, as it relies on the high pH of aerosols (Ge et al., 2019). In recent years, there has been increasing attention on the enhancing effect of NH₃ on the redox reaction, with laboratory experiments demonstrating its ability to generate substantial amounts of HONO (Ge et al., 2019). This study highlights the importance of this

reaction based on actual atmospheric observations. Furthermore, numerous studies have indicated that if control over NH₃ emissions continues to relax while SO₂ and NO₂ emissions decrease, the particle pH in future China is expected to rise steadily (Xie et al., 2020;Song et al., 2019;Wang et al., 2020). Consequently, the redox reaction of NO₂ with SO₂ could become a significant source of HONO in China. Therefore, it is crucial to coordinate the control of SO₂, NO_x, and NH₃ emissions to avoid a rapid increase in the particle pH.

380 Data availability: All the data presented in this article can be accessed through 381 https://zenodo.org/records/10273539. (Zhang, 2023). 382 383 Author contributions. XZ Data Curation, Writing - Original Draft, Visualization. 384 385 LW, NW, SM, and DZ Investigation, Visualization, Data Curation. DZ, HZ, and MW Investigation. SW Conceptualization, Data Curation, Supervision. RZ Data Curation, 386 387 Funding acquisition. All people are involved in the discussion of the results. 388 389 **Competing interest.** The authors declare no competing financial interest. 390 **Acknowledgments.** This work was supported by the China Postdoctoral Science 391 392 Foundation (2023M733220), the Zhengzhou PM_{2.5} and O₃ Collaborative Control and Monitoring Project (20220347A), and the National Key Research and Development 393 394 Program of China (No. 2017YFC0212403).

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Figures:

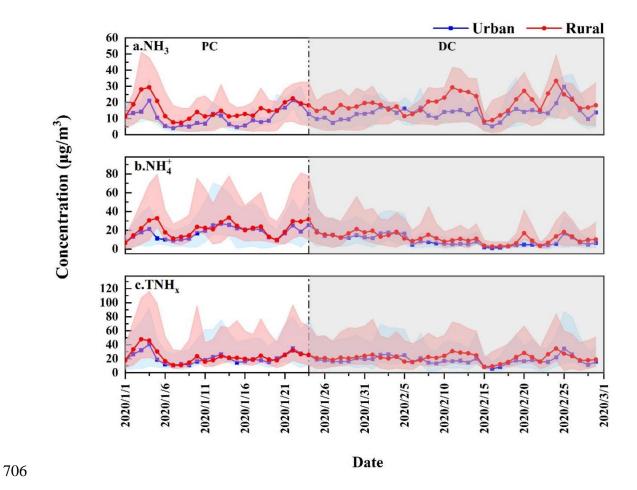


Figure 1. Temporal variations of a. NH_3 , b. NH_4^+ , and c. TNH_x at the urban and rural sites before (PC) and during (DC) the COVID-19 outbreak, respectively. The shaded areas of the curve represent the maximum and minimum values.

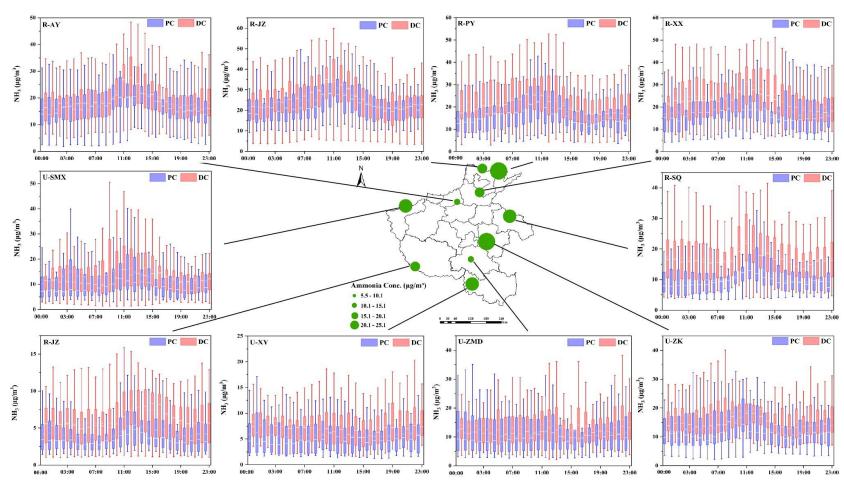


Figure 2. Daily variation of NH₃ concentrations at ten sites before (PC) and during (DC) the COVID-19 outbreak. The green dots represent the location of ten sites and their size represents the concentration of NH₃; In each box, the top, middle, and bottom lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.

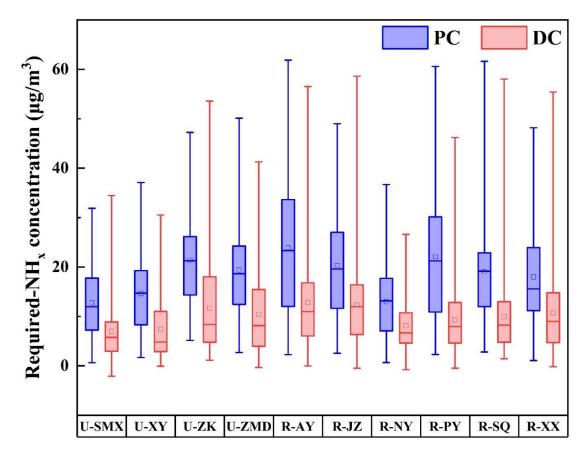


Figure 3. Box diagram of changes in Required-NH $_x$ at ten sites before (PC) and during (DC) the COVID-19 outbreak. In each box, the top, middle, and bottom lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.

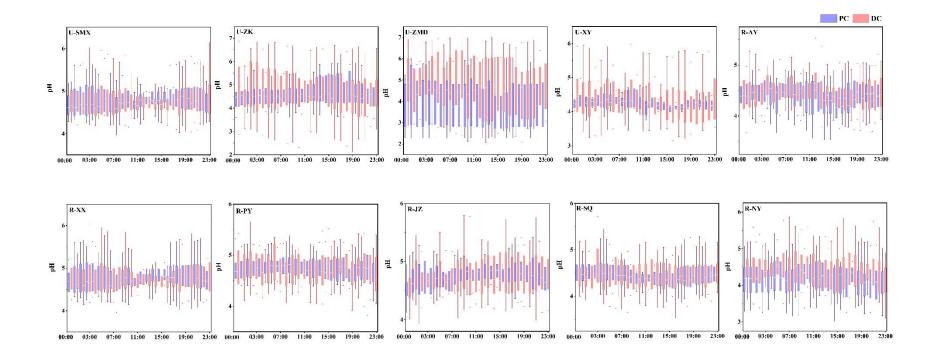


Figure 4. Diurnal patterns of pH at ten sites before (PC) and during (DC) the COVID-19 outbreak. In each box, the top, middle, and bottom lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.

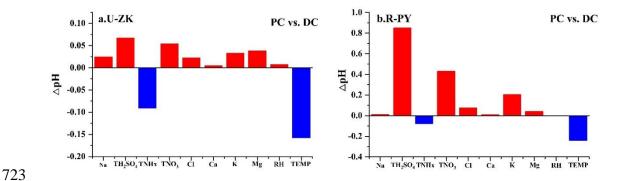


Figure 5. Changes of pH (Δ pH) through the sensitivity tests (Figure S5 and S6) by changing parameters between PC and DC at the a. U-ZK and b. R-PY sites.

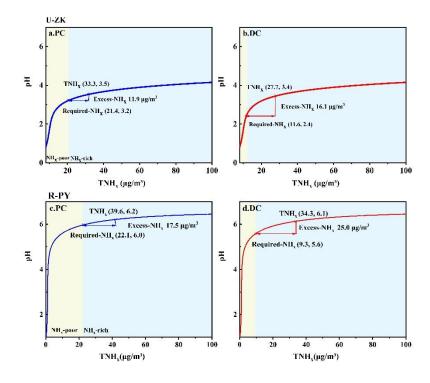


Figure 6. Particle pH corresponds to increasing TNH_x at U-ZK and R-PY sites to examine the effects of major indicators of NH₃ (i.e., TNH_x, Required-NH_x, and Excess-NH_x) on aerosol acidity. Particle pH was calculated by using a wide range of TNH_x (25–130 μg/m³) and average values of other parameters in PC and DC of U-ZK and R-PY sites. The concentrations of TNH_x, Required-NH_x, and Excess-NH_x with corresponding pH values are marked by a hollow box, hollow circle, and arrow respectively. The yellow and blue background colors correspond to the NH_x-poor and NH_x-rich, respectively.

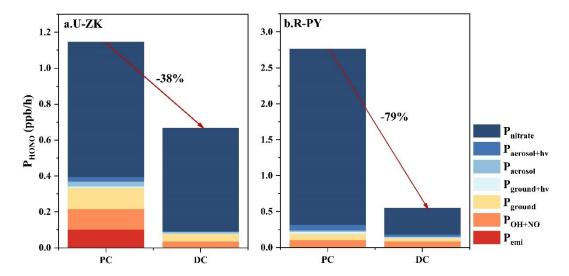


Figure 7. Comparison of HONO sources at a. U-ZK and b. R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The calculation method can be found in Text S4.

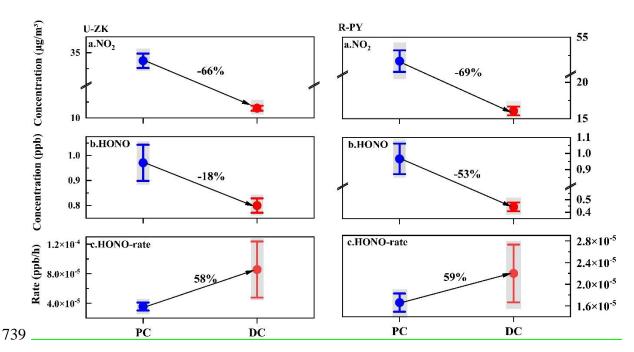


Figure 8. Decline ratios of a. NO_2 , b. HONO concentration, and c. HONO production rate at U-ZK and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The center point represents the mean value, and the upper and lower whiskers represent the 95% confidence interval of the mean. The shadows in the figure represent the uncertainties of NO_2 measurement ($\pm 10\%$), HONO measurement ($\pm 20\%$), and the

745 HONO formation rate of R_1 reaction (-78–123%), respectively.

Table:

Table 1. Changes in concentrations (mean \pm standard deviation) of NH₃, NH₄⁺, and TNH_x at ten sites during entire periods (Average), before (PC), and during (DC) the COVID-19 outbreak.

Sites	Substances	Average ($\mu g/m^3$)	PC (μ g/m³)	DC (μ g/m³)
U-SMX	NH_3	13.8 ± 10.8	12.6 ± 10.1	14.5 ± 11.1
	NH_4^+	10.9 ± 7.2	14.2 ± 7.2	8.8 ± 6.5
	TNH_x	22.9 ± 14.1	24.9 ± 14.5	21.7 ± 13.8
U-ZK	NH_3	15.6 ± 8.3	12.7 ± 6.5	17.4 ± 8.8
	NH_4^+	13.6 ± 9.3	19.1 ± 8.4	10.3 ± 8.1
	TNH_x	28.6 ± 13.7	30.9 ± 12.8	27.1 ± 14.0
U-ZMD	NH_3	13.1 ± 8.4	11.6 ± 8.2	14.0 ± 8.4
	NH_4^+	13.9 ± 9.8	19.6 ± 10.3	10.3 ± 7.5
	TNH_x	25.7 ± 14.6	30.3 ± 15.1	22.8 ± 13.5
U-XY	NH_3	7.0 ± 4.3	5.7 ± 4.0	7.9 ± 4.3
	NH_4^+	11.0 ± 7.7	15.4 ± 7.6	8.3 ± 6.5
	TNH_x	17.6 ± 9.8	20.6 ± 10.1	15.7 ± 9.2
R-AY	NH_3	19.0 ± 8.4	17.9 ± 8.3	19.7 ± 8.4
	NH_4^+	19.3 ± 12.9	26.4 ± 13.7	15.0 ± 10.3
	TNH_x	36.6 ± 18.2	41.7 ± 20.4	33.4 ± 16.0
R-XX	NH_3	21.7 ± 10.2	18.1 ± 9.3	23.8 ± 10.1
	NH_4^+	15.9 ± 10.4	20.6 ± 11.0	13.0 ± 8.8
	TNH_x	34.9 ± 17.0	35.1 ± 18.8	34.8 ± 15.8
R-PY	NH_3	19.8 ± 9.4	16.8 ± 8.1	21.7 ± 9.6
	NH_4^+	17.4 ± 11.8	25.3 ± 12.6	12.4 ± 8.0
	TNH_x	35.2 ± 17.8	39.4 ± 19.8	32.6 ± 15.7
R-JZ	NH_3	25.3 ± 11.5	24.1 ± 11.5	25.9 ± 11.4
	NH_4^+	17.3 ± 11.3	22.7 ± 11.6	14.2 ± 9.9
	TNH_x	40.8 ± 20.1	42.9 ± 22.8	33.5 ± 18.2
R-SQ	NH_3	15.0 ± 7.9	10.3 ± 5.2	17.7 ± 7.9
	NH_4^+	13.4 ± 8.5	18.9 ± 8.6	10.3 ± 6.7
	TNH_x	26.3 ± 13.2	25.5 ± 14.0	26.8 ± 12.7
R-NY	NH_3	5.5 ± 3.1	4.3 ± 2.7	6.2 ± 3.2
	NH_4^+	10.2 ± 6.9	13.3 ± 7.2	8.4 ± 6.1
	TNH_x	14.8 ± 8.5	16.0 ± 9.5	14.1 ± 7.8