### No.: egusphere-2023-2913

Title: Measurement Report: Elevated excess-NH<sub>3</sub> can promote the redox reaction to produce HONO: Insights from the COVID-19 pandemic

# **Reviewer #1:**

### **General Comments:**

In this study, the authors analyzed the chemical composition changes during the pandemic in ten urban and rural sites, and compared the HONO concentration level before and during the emission control period. The authors found that the HONO decline was relatively insignificant compared to its precursors and a detailed calculation shows that the enhanced production rate of aqueous phase reaction partially offset the effect of lower precursors. By comparing the atmospheric acids and bases concentrations, the authors suggested that the enhanced level of NH<sub>3</sub> and elevated aerosol pH due to less acidic components in the atmosphere was the reason for the higher HONO production rate. It can be one of the possible reasons, while there are several important issues that the authors did not have enough discussion or provide clear explanation. Some analysis and explanations are too simplified to give the assessment of the quality of this study.

Thank you for your careful reading of our paper and valuable comments and

suggestions. We believe that we have adequately addressed your comments. To facilitate your review, we used yellow highlights for your comments, green highlights for Reviewer #2, and red color indicating our own corrections in the manuscript.

#### Major issues:

1. The direct emission HONO was estimated based on the vehicle emission factors and NO<sub>x</sub> concentration level, which should reflect a general situation of normal human activities. However, during the pandemic, the emission factors could change very significantly if only necessary activities were allowed to be carried out. The authors did not mention emission profile change before and during the pandemic, which could lead to the overestimation of the effect of other pathways.

**Response:** Thanks for your comments. We determined whether it is necessary to calculate vehicle emissions, and a supplementary HONO emission factor table from vehicle emissions was added to support the selection of factors in the revised version:

"HONO can be released directly into the atmosphere through vehicle exhaust (Burling et al., 2010; Veres et al., 2010). The lifetime of HONO in the atmosphere is relatively short, so vehicle emissions significantly contribute to urban atmospheric HONO (Chen et al., 2023; Liu et al., 2021a). Considering that there has been a significant reduction in vehicle emissions in urban areas during DC. Additionally, the R-PY site is far from roads. Thus, vehicle emissions may not be the primary source of HONO for the U-ZK site during DC and R-PY sites during entire periods. To further validate the above conclusions, the conditional bivariate probability function diagrams of NO<sub>2</sub> at U-ZK and R-PY sites during PC and DC are depicted in Figure S2. NO<sub>2</sub> predominantly originated from long-distance transport at the U-ZK site during DC and the R-PY site during both PC and DC. Consequently, vehicle emissions are only calculated for the U-ZK site during the PC.

Here we use the HONO/NO<sub>x</sub> ratio to estimate HONO concentration, which is generally considered to be the vehicle emission factor (Kramer et al., 2020; Hao et al., 2020; Yu et al., 2022) for HONO. The calculation formula is as follows:

$$[\text{HONO}_{\text{emi}}] = 0.8\% \times [\text{NO}_{x}]$$
(1)

where [HONO<sub>emi</sub>] and [NO<sub>x</sub>] represent the HONO concentration emitted by vehicles and the observed NO<sub>x</sub> concentration, respectively. Regarding previous studies (Table S3), 0.8% was selected as the vehicle emission factor, considering differences in vehicle type, fuel composition, and other factors (Kramer et al., 2020; Hao et al., 2020; Huang et al., 2017)."



Figure S2. Result of conditional bivariate probability function plots: NO2 at U-ZK and

R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The color scale bar

<u>1</u>	Table S3. Sur	nmary of vehicle emission	factors.
Observation site	Period	Emission factor (%)	Reference
Beijing	<mark>2020</mark>	<mark>0.79</mark>	(Meng et al., 2020)
Hong Kong	<mark>2015</mark>	<mark>0.4–1.8</mark>	<mark>(Yun et al., 2017)</mark>
Hong Kong	<mark>2011</mark>	<mark>0.5–1.6</mark>	<mark>(Xu et al., 2015)</mark>
Kiesberg Tunnel	<mark>2001</mark>	<mark>0.8</mark>	<mark>(Kleffmann et al., 2003)</mark>
Kiesberg Tunnel	<mark>1997</mark>	<mark>0.3–0.8</mark>	(Kurtenbach and Wiesen, 2001)
<b>Guangzhou</b>	<mark>2019</mark>	<mark>1.31</mark>	(Li et al., 2021b)

represents NO<sub>2</sub> concentration.

2. Supplement Line 107: it is very challenging to pick a representative OH concentration to represent the general situation. The authors also suggested in the

introduction that OH radical concentration could change during emission control as part of atmospheric oxidizing capacity changes. While the authors did not mention such an approach in their HONO production calculation. In addition to other reaction pathways, another possibility is the change of reaction rates, like OH concentration levels and higher temperature (the authors only mentioned H and K temperature dependence but did not mention  $k_1$  temperature dependence, which could be important). The authors should fully discuss the possibilities of the changes in reaction rate and possible sinks.

**Response:** Thank you for your valuable comments.

Firstly, we have modified the method for determining •OH concentration in the revised manuscript:

"•OH concentration was simulated according to the empirical model (Hu et al., 2022; Wang et al., 2025):

$$[OH] = 4.1 \times 10^9 \times \frac{J(O^1D) \times J(NO_2) \times (140 \times [NO_2] + 1) + [HONO] \times J(HONO)}{0.41 \times [NO_2]^2 + 1.7 \times [NO_2] + 1 + [NO] \times k_{NO+OH} + [HONO] \times k_{NO+OH}} (12)$$

where, J (O<sup>1</sup> D), J (NO<sub>2</sub>), and J (HONO) are the photolysis rates calculated using the TUV model (v5.2; available at http://cprm.acom.ucar.edu/Models/TUV/). The calculated •OH concentration varied from  $0.1 \times 10^6$  to  $4 \times 10^6$  molecule/cm<sup>3</sup> at U-ZK and  $0.1 \times 10^6$  to  $5 \times 10^6$  molecule/cm<sup>3</sup> t R-PY, which was comparable to the levels in other cities of North China (Li et al., 2018; Fuchs et al., 2017; Yang et al., 2017). Since there is no photolysis at night, the •OH concentration was assumed to be  $0.8 \times 10^6$  molecule/cm<sup>3</sup> (Wang et al., 2022)."

Secondly, the sources of HONO were recalculated to better investigate the changes

in HONO between PC and DC periods:

#### **Text S4 Sources of HONO**

#### 4.1 Direct emission

HONO can be released directly into the atmosphere through vehicle exhaust (Burling et al., 2010; Veres et al., 2010). The lifetime of HONO in the atmosphere is relatively short, so vehicle emissions significantly contribute to urban atmospheric HONO (Chen et al., 2023; Liu et al., 2021a). Considering that there has been a significant reduction in vehicle emissions in urban areas during DC. Additionally, the R-PY site is far from roads. Thus, vehicle emissions may not be the primary source of HONO for the U-ZK site during DC and R-PY sites during entire periods. To further validate the above conclusions, the conditional bivariate probability function diagrams of NO<sub>2</sub> at U-ZK and R-PY sites during PC and DC are depicted in Figure S2. NO<sub>2</sub> predominantly originated from long-distance transport at the U-ZK site during DC and the R-PY site during both PC and DC. Consequently, vehicle emissions are only calculated for the U-ZK site during the PC.

Here we use the HONO/NO<sub>x</sub> ratio to estimate HONO concentration, which is generally considered to be the vehicle emission factor (Kramer et al., 2020; Hao et al., 2020; Yu et al., 2022) for HONO. The calculation formula is as follows:

$$[\text{HONO}_{\text{emi}}] = 0.8\% \times [\text{NO}_{\text{x}}]$$
(1)

where [HONO<sub>emi</sub>] and [NO<sub>x</sub>] represent the HONO concentration emitted by vehicles

and the observed  $NO_x$  concentration, respectively. Regarding previous studies (Table S3), 0.8% was selected as the vehicle emission factor, considering differences in vehicle type, fuel composition, and other factors (Kramer et al., 2020; Hao et al., 2020; Huang et al., 2017).

4.2 Homogeneous reaction of NO and •OH

The reaction between NO and •OH is the primary gas-phase reaction source of HONO at high NO concentrations, and the production rate contribution (P<sub>OH+NO</sub>) for this reaction can be calculated as:

$$P_{OH+NO} = k_{OH+NO}[OH][NO]$$
(2)

where  $k_{OH+NO}$  (7.2 × 10<sup>-12</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup>) is the rate constant for the reactions at 298K (Li et al., 2012). •OH concentration was simulated according to the empirical model (Hu et al., 2022; Wang et al., 2025):

$$[OH] = 4.1 \times 10^9 \times \frac{J(O^1D) \times J(NO_2) \times (140 \times [NO_2] + 1) + [HONO] \times J(HONO)}{0.41 \times [NO_2]^2 + 1.7 \times [NO_2] + 1 + [NO] \times k_{NO+OH} + [HONO] \times k_{NO+OH}} (3)$$

where, J (O<sup>1</sup> D), J (NO<sub>2</sub>), and J (HONO) are the photolysis rates calculated using the TUV model (v5.2; available at http://cprm.acom.ucar.edu/Models/TUV/). The cloud optical depth value for the TUV model was adjusted so that the predicted UVB radiation intensity matched the observations (Lyu et al., 2019; Wang et al., 2022b). The calculated •OH concentration varied from  $0.1 \times 10^6$  to  $4 \times 10^6$  molecule/cm<sup>3</sup> at U-ZK and  $0.1 \times 10^6$  to  $5 \times 10^6$  molecule/cm<sup>3</sup> t R-PY, which was comparable to the levels in other cities of North China (Li et al., 2018; Fuchs et al., 2017; Yang et al., 2017). Since there is no

photolysis at night, the •OH concentration was assumed to be  $0.8 \times 10^6$  molecule/cm<sup>3</sup> (Wang et al., 2022).

4.3 Heterogeneous conversion of NO<sub>2</sub> to HONO

#### 4.3.1 Heterogeneous dark reactions

The heterogeneous conversion of  $NO_2$  to HONO on the ground ( $P_{ground}$ ) and on the aerosol surface ( $P_{aerosol}$ ) was calculated based on parameters obtained from experiments or observations.

$$P_{\text{ground}} = \frac{1}{8} \gamma_1 \times [\text{NO}_2] \times C_{\text{NO}_2} \times \frac{S_g}{V}$$
(4)

$$\mathbf{P}_{\text{aerosol}} = \frac{1}{4} \gamma_2 \times [\text{NO}_2] \times \mathbf{C}_{\text{NO}_2} \times \frac{\mathbf{S}_{\text{a}}}{\mathbf{V}}$$
(5)

$$\frac{S_g}{V} = \frac{1}{MLH}$$
(6)

$$C_{NO_2} = \sqrt{\frac{8RT}{\pi M}}$$
(7)

where  $C_{NO2}$  is the average molecular velocity of NO<sub>2</sub> molecule (m s<sup>-1</sup>); R is the ideal gas constant; T is the temperature (K); M is the molecular weight of NO<sub>2</sub> (kg mol<sup>-1</sup>); MLH is the height of the mixed layer, which is determined to be 50 m due to HONO formation on the ground level and its short lifetime (Liu et al., 2020b); S<sub>a</sub>/V is the surface area to volume ratio of aerosol, estimated by Su et al. (Su et al., 2008).

## 4.3.2 Heterogeneous photo-enhanced reactions

The heterogeneous photo-enhanced reactions of NO2 on the surface of the ground

 $(P_{ground + hv})$  and the surface of the aerosol  $(P_{aerosol + hv})$  were calculated following (Zhang

et al., 2020a):

$$P_{\text{ground+hv}} = \frac{1}{8} \times C_{\text{NO}_2} \times \frac{1}{\text{MLH}} \times \gamma_1 \times \frac{J_{\text{NO}_2}}{J_{\text{NO}_2 \text{more}}} \times [\text{NO}_2]$$
(8)

$$\mathbf{P}_{\text{aerosol+hv}} = \frac{1}{4} \times \mathbf{C}_{\text{NO}_2} \times \frac{\mathbf{S}_{a}}{\mathbf{V}} \times \gamma_2 \times \frac{\mathbf{J}_{\text{NO}_2}}{\mathbf{J}_{\text{NO}_2 \text{ noon}}} \times [\text{NO}_2]$$
(9)

where JNO<sub>2</sub> and JNO<sub>2, noon</sub> are the photolysis rate of NO<sub>2</sub> and the photolysis rate of NO<sub>2</sub> at noon during the day, respectively.

 $y_1$  and  $y_2$  are the absorption coefficient of NO<sub>2</sub> on the ground and aerosol surface, respectively, which is assumed to be  $4 \times 10^{-6}$  (Yu et al., 2022; Zhang et al., 2021; Zhang et al., 2020a). Moreover, we discuss the uncertainties based on the recommended values of  $2 \times 10^{-6}$ – $1 \times 10^{-5}$  as upper and lower bounds(Chen et al., 2023; VandenBoer et al., 2013; Wong et al., 2011). Results show (Figure S3) that the uncertainties for P<sub>ground</sub>, P<sub>aerosol</sub>, P<sub>groung+hv</sub>, and P<sub>aerosol+hv</sub> are -50% to 150%, -50% to 151%, -20% to 120%, and -50% to 121% at the U-ZK, respectively. At the R-PY, the uncertainties for P<sub>ground</sub>, P<sub>aerosol</sub>, P<sub>groung+hv</sub>, and P<sub>aerosol+hv</sub> are -50% to 150%, -50% to 151%, -20% to 120%, and

#### 4.4 Nitrate photolysis

The nitrate photolysis ( $P_{nitrate}$ ) was calculated based on the measured nitrate concentration ( $NO_3^-$ ) from  $PM_{2.5}$  and nitrate photolysis rate ( $J_{nitrate \rightarrow HONO}$ ):

$$\mathbf{P}_{\text{nitrate}} = \mathbf{J}_{\text{nitrate} \to \text{HONO}} \times [\text{NO}_3^-] \tag{10}$$

where the  $J_{nitrate \rightarrow HONO}$  was simulated by normalizing UV values, when the Zenit Angle is 0°,  $J_{nitrate \rightarrow HONO}$  varied within the range of  $1.22 \times 10^{-5}$  to  $4.84 \times 10^{-4}$  s<sup>-1</sup>, with an average value of  $8.24 \times 10^{-5}$  s<sup>-1</sup> (Bao et al., 2018)."

Unfortunately, for MLH, S<sub>a</sub>/V, and the relationship between k1 and temperature, as there were no observational data or scientifically established estimation methods, this study did not consider their variations. This omission may lead to differences in conclusions and warrants further investigation in future research.



Figure S2. Result of conditional bivariate probability function plots: NO<sub>2</sub> at U-ZK and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The color scale bar represents NO<sub>2</sub> concentration.



Figure S3. HONO production rate using different uptake rates of NO<sub>2</sub> at the U-ZK and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. (a)P<sub>ground</sub>, (b) P<sub>aerosol</sub>, (c) P<sub>ground+hv</sub>, and (d) P<sub>aerosol+hv</sub>

Observation site	Period	Emission factor (%)	Reference
Beijing	<mark>2020</mark>	<mark>0.79</mark>	<mark>(Meng et al., 2020)</mark>
Hong Kong	<mark>2015</mark>	<mark>0.4–1.8</mark>	<mark>(Yun et al., 2017)</mark>
Hong Kong	<mark>2011</mark>	<mark>0.5–1.6</mark>	<mark>(Xu et al., 2015)</mark>
Kiesberg Tunnel	<mark>2001</mark>	<mark>0.8</mark>	(Kleffmann et al., 2003)
Kiesberg Tunnel	<mark>1997</mark>	<mark>0.3–0.8</mark>	(Kurtenbach and Wiesen, 2001)
<mark>Guangzhou</mark>	<mark>2019</mark>	<mark>1.31</mark>	(Li et al., 2021b)

Table S3. Summary of vehicle emission factors.

3. It is also questionable about the contribution of NH<sub>3</sub> concentration changes to the total pH changes. Temperature, relative humidity, and other salts could also contribute to pH changes. It was not mentioned how the sensitivity tests of Line 264-275 were done and the interpretation of the results was also unclear. The authors did not give a complete pH comparison like NH<sub>x</sub> levels, only provided two sites in Figure 4. The authors mentioned the increase of pH 0.4 and 0.1 for U-ZK and R-PY sites respectively. However, based on the NH<sub>3</sub> levels shown in Table 1 and the relationship mentioned in Song et al. (2019):  $\partial pHi/\partial [NH_3(g)] \approx 0.4/[NH_3(g)]$ , the NH<sub>3</sub> concentration changes was only responsible for 0.13-unit pH change in U-ZK (less than half). The pH changes of most sites, if only considering NH<sub>3</sub> levels changes in Table 1, can be calculated to be around 0.1 with the exception of R-SQ where NH<sub>3</sub> concentration nearly doubled.

**Response:** Sorry for the misunderstanding. The formula in Song's study only considers the effect of NH<sub>3</sub> on the pH value of particulate matter and does not take into account other substances such as TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, T, etc., which have a greater impact on pH value. Therefore, when the NH<sub>3</sub> value in this study is brought into the formula, there is a different conclusion obtained. To explore the dominant factors that determine the high pH during the DC, we have added a detailed description of the sensitivity tests of pH to input data:

"To explore the dominant factors that determine the local particle pH level and result in the high pH during the DC, sensitivity tests of pH to chemical species (i.e.,  $TNH_x$ , TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, TCl, TNa,  $K^+$ , Ca<sup>2+</sup>, and Mg<sup>2+</sup>) and meteorological parameters (i.e., To explore the dominant factors that determine the local particle pH level and result in the high pH during the DC, sensitivity tests of pH to chemical species (i.e.,  $TNH_x$ ,  $TH_2SO_4$ , TNO<sub>3</sub>, TCl, TNa,  $K^+$ , Ca<sup>2+</sup>, and Mg<sup>2+</sup>) and meteorological parameters (i.e., T and RH) were performed. A given range for a variable (i.e., TNH<sub>x</sub>) with corresponding average values of other parameters (i.e., TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, TCl, TNa, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, T, and RH) was input into the model and simulated to compare its effects on pH. As shown in Fig. S7, pH increases with the cation concentrations (i.e.,  $TNH_x$ ,  $Na^+$ ,  $K^+$ ,  $Ca^{2+}$ , and  $Mg^{2+}$ ) increasing as well as the anion concentrations (i.e., TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, and Cl<sup>-</sup>), T and RH decreasing. According to the average values of input data during PC (Blue line in Fig. S7) and DC (Red line in Fig. S7) at U-ZK and R-PY sites respectively, the changes in pH ( $\Delta$ pH in Fig. 5) indicate that the decrease in TNH<sub>x</sub> concentration and the increase in T in DC led to a decrease in pH values ( $\Delta pH$ : 0.09 at U-ZK and 0.08 at R-PY sites) compared to PC. However, this effect was outweighed by the decrease in  $TH_2SO_4$  ( $\Delta pH$ : 0.07 and 0.8 at U-ZK and R-PY sites, respectively) and TNO<sub>3</sub> (ΔpH: 0.05 and 0.4 at U-ZK and R-PY sites, respectively) concentrations as well as the increase in  $K^+$  ( $\Delta pH$ : 0.03 at U-ZK and 0.2 at R-PY site) and  $Mg^{2+}$  ( $\Delta pH$ : 0.01 at U-ZK and 0.04 at R-PY site) concentrations in the DC, and resulting in an overall increase in pH values in the DC. Furthermore, the relationship between particle pH with the concentrations of Required-NH<sub>x</sub>, and Excess-NH<sub>x</sub>, which considers all chemical components, is investigated to examine the dominant factor on the increasing pH in DC. As shown in Fig. 6, the higher Excess-NH<sub>x</sub> concentrations in the DC led to higher increases in pH values ( $\Delta pH$ : 1 at U-ZK and 0.5 at R-PY site) than those in PC ( $\Delta pH$ : 0.3 at U-ZK and 0.2 at R-PY site), thus Excess-NHx concentrations may be the key factor in promoting





Figure 5. Changes of pH ( $\Delta$ pH) through the sensitivity tests (Figure S5 and S6) by

changing parameters between PC and DC at the a. U-ZK and b. R-PY sites.



Figure 6. Particle pH corresponds to increasing  $\text{TNH}_x$  at U-ZK and R-PY sites to examine the effects of major indicators of  $\text{NH}_3$  (i.e.,  $\text{TNH}_x$ , Required- $\text{NH}_x$ , and Excess- $\text{NH}_x$ ) on aerosol acidity. Particle pH was calculated by using a wide range of  $\text{TNH}_x$  (25–130 µg/m<sup>3</sup>) and average values of other parameters in PC and DC of U-ZK and R-PY sites. The concentrations of  $\text{TNH}_x$ , Required- $\text{NH}_x$ , and Excess- $\text{NH}_x$  with corresponding pH values are marked by a hollow box, hollow circle, and arrow respectively. The yellow and blue background colors correspond to the  $\text{NH}_x$ -poor and  $\text{NH}_x$ -rich, respectively.



Figure S7. Sensitivity tests of pH to each factor. The vertical bar represents the mean values before (PC) and during (DC) the COVID-19 outbreak. A given range for a variable (i.e.,  $TNH_x$ ) with corresponding average values of other parameters (i.e.,  $TH_2SO_4$ ,  $TNO_3$ , TCl, TNa,  $K^+$ ,  $Ca^{2+}$ ,  $Mg^{2+}$ , T, and RH) was simulated to compare its effects on pH.

Additionally, we have added a complete pH comparison of ten sites:

•

"Diurnal patterns of particle pH in PC and DC at ten sites are summarized in Fig. 4 with their average values listed in Table S9. PM<sub>2.5</sub> shows consistent moderate acidity, with mean values in the range of 4.2–5.1, which were close to the values in previous studies (Table S9). Compared to the PC, the particle pH at ten sites increased obviously in the DC, with the highest increase of 0.5 (U-ZK) and 0.3 (R-PY) at urban and rural sites, respectively, which were the subject of in-depth discussion in the following text."



- 3 lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of
- 4 statistical data, respectively

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mean – sta	ndard).			
	Sites .	Periods	<mark>рН</mark>	References
<mark>Urban</mark>	Sanmenxia	<mark>Jan–Feb 2020</mark>	$4.6 \pm 0.5 / 4.8 \pm 0.9$	This study
	<mark>Zhoukou</mark>	Jan–Feb 2020	${\color{red}{4.6\pm0.6/5.1\pm0.4}}$	
	Zhumadian	<mark>Jan–Feb 2020</mark>	$4.6 \pm 0.3/4.8 \pm 1.2$	
	<mark>Xinyang</mark>	Jan–Feb 2020	$4.2 \pm 0.3/4.6 \pm 1.3$	
Rural	<mark>Anyang</mark>	Jan–Feb 2020	$4.5 \pm 0.4 / 4.6 \pm 0.8$	-
	<mark>Xinxiang</mark>	<mark>Jan–Feb 2020</mark>	$4.8 \pm 0.5 / 4.9 \pm 0.9$	
	<mark>Puyang</mark>	Jan–Feb 2020	$4.8 \pm 0.3/5.1 \pm 0.9$	
	<mark>Jiaozuo</mark>	Jan–Feb 2020	$4.9 \pm 0.5/5.1 \pm 0.8$	
	<mark>Shangqiu</mark>	Jan–Feb 2020	$4.5 \pm 0.3/4.7 \pm 0.8$	
	<mark>Nanyang</mark>	<mark>Jan–Feb 2020</mark>	$4.2 \pm 0.5/4.4 \pm 0.7$	
<mark>Urban</mark>	Beijing	<mark>Jan–Feb 2015</mark>	<mark>4.5</mark>	<mark>(Guo et al., 2017)</mark>
	Beijing	Dec 2016	$4.3 \pm 0.4$	<mark>(Liu et al., 2017)</mark>
	<b>Beijing</b>	Feb 2017	$4.5 \pm 0.7$	(Ding et al., 2019)
	<mark>Tianjin</mark>	Dec–Jun 2015	<mark>4.9 ± 1.4</mark>	<mark>(Shi et al., 2017)</mark>
	<mark>Tianjin</mark>	Aug 2015	$3.4 \pm 0.5$	<mark>(Shi et al., 2019)</mark>
	<mark>Hohhot</mark>	Winter <b>Winter</b>	<mark>5.7</mark>	(Wang et al., 2019)
	Mt. Tai	Summer	$2.9 \pm 0.5$	<mark>(Liu et al., 2021b)</mark>
	<mark>Taoyuan</mark>	<mark>Nov 2017–Jan 2018</mark>	$5.1 \pm 1.0$	<mark>(Duan et al., 2021)</mark>
	<mark>Zhengzhou</mark>	<mark>Jan 2018</mark>	<mark>4.5</mark>	<mark>(Wang et al., 2020)</mark>
	<mark>Anyang</mark>	<mark>Jan 2018</mark>	<mark>4.8</mark>	<mark>(Wang et al., 2020)</mark>
Mountain	Mt. Tai	Summer	$3.6 \pm 0.7$	<mark>(Liu et al., 2021b)</mark>
Rural	Shanglan	Nov 2017–Jan 2018	$5.5 \pm 1.1$	(Duan et al., 2021)

Table S9. Comparison of the particle pH values in this study (PC/DC) and other sites (mean or mean  $\pm$  standard).

4. Figure 4, the maximum and minimum values provided little information of the whole pH variations. A box and whisker plot are more useful to identify the general trends and variations. And there were frequent situations of maximum pH higher than 7, which could not be explained by higher NH<sub>3</sub> concentrations. Instead, it could be from the strong influence of dust components. If that situation happened frequently enough (hard to judge now based on the information given), it could be the dust components that are actually responsible for the high pH.

Response: Thank you for your suggestions. We redrew Figure 4 as a box diagram and replaced

it in the revised version. After examining the raw data, we found that the pH data higher than 7 mainly concentrated in clean air with low pollutant concentrations. Additionally, some data had RH levels below 30%, which could lead to significant errors in the model. Thus, ISORROPIA-II was rerun only using data with  $RH \ge 30\%$  in the revised version.



4 statistical data, respectively.

5. It should also be mentioned that the approach of the authors used to estimate  $AWC_{org}$  is sensitive to the parameters chosen, such as OM/OC ratio, density, and kappa parameter. Normally, the term  $AWC_{org}$  is small enough so that its influence is limited, while it is possible the uncertainty associated with the parameters chosen became big enough when inorganic salts become depleted and the relative contribution of OM got enhanced.

**Response:** Thank you for your comment. We supplemented the selection criteria for calculating parameters in the revised manuscript: "AWC<sub>org</sub> is the particle water associated with the organic matters predicted using the following method:

$$AWC_{org} = \frac{m_s}{\rho_s} \frac{k_{org}}{\left(\frac{1}{RH} - 1\right)}$$
(2.2)

where  $m_s$  is the mass concentration of organic matter (OM = OC × *f*). The *f* is the conversion factor of OC, which is dependent on the extent of OM oxidation and secondary organic aerosol formation (Chow et al., 2015). Studies on the ratio of OM/OC in fourteen cities in China suggested that the mean value of *f* was  $1.59 \pm 0.18$  during the winter season in Northern China (Xing et al., 2013), and thus we adopted 1.6 as the *f* in this study. *k*<sub>org</sub> is the organic hygroscopicity parameter and depends on organic functionality, water solubility, molecular weight, and oxidation level. Han et al. (2022) have reported that the *k*<sub>org</sub> generally increased with O: C ratios, with a range of 0–0.3 for 23 organics, including carboxylic acids, amino acids, sugars, and alcohols. Gunthe

et al, (2011) estimated a $k_{org} = 0.06 \pm 0.01$ for the effective average hygroscopicity of
the non-refractory organic particulate matter in the aerosols in Beijing. Our previous
study has estimated that the uncertainties of $k_{org}$ value (0.06) for pH in the range of 0–
0.3 only lead to $-1-3\%$ errors, which can be neglected (Wang et al., 2023a). Therefore,
the value of 0.06 was selected in this paper. $\rho_s$ is the organic density, which was chosen
to be 1.35 g/cm <sup>3</sup> following previous studies (Table S2)."

Observation site	Period	$\rho_s(g/cm^3)$	Reference
Beijing	Dec 2016	<mark>1.4</mark>	<mark>(Liu et al., 2017)</mark>
<b>Tianjin</b>	Dec-Jun 2015	<mark>1.3</mark>	(Shi et al., 2017)
Xi'an	Nov-Dec 2012	<mark>1.4</mark>	<mark>(Guo et al., 2017)</mark>
<mark>Hohhot</mark>	Winter 2015	<mark>1.35</mark>	<mark>(Wang et al., 2019)</mark>
Northeastern USA	Feb-Mar 2015	<mark>1.4</mark>	<mark>(Guo et al., 2016)</mark>
Crete, Greece	Aug-Nov 2012	<mark>1.35</mark>	(Bougiatioti et al., 2016)
<mark>Alabama, USA</mark>	<mark>Jun-Jul 2013</mark>	<mark>1.4</mark>	<mark>(Guo et al., 2015)</mark>
Georgia, USA	Aug-Oct 2016	<mark>1.4</mark>	<mark>(Nah et al., 2018)</mark>

**Table S2**. The value of  $\rho_s$  in other studies.

### Minor issues:

6. The definition of TNHx is different in Line113 and Line 228.

Response: Thank you for your careful reading of our paper. The formula is used

uniformly in the new version:

$$\text{TNH}_{\text{x}} = 17 \times (\frac{[\text{NH}_{4}^{+}]}{18} + \frac{[\text{NH}_{3}]}{17})$$

7. Line 42, the study cited is the result based on a field campaign.

**Response:** Thank you for your comment. We added more references: "Nitrous acid (HONO) is a critical precursor of hydroxyl radical (OH), contributing to more than 60%

of OH production (Alicke, 2003; Platt et al., 1980; Kleffmann et al., 2005)."

8. Figure 2, the max and min as error bars provide little information about the

general trends, and there are negative values.

Response: Thank you for your comments. We redrew Figure 2 as a box diagram and

the negative values were removed after quality control.





location of ten sites and their size represents the concentration of NH<sub>3</sub>; In each box, the top, middle, and bottom lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.

9. Line 215, it is hard to judge if agricultural activity got weakened or not. The NH<sub>3</sub> concentration change could be due to less farm activity like less frequent animal feces cleaning, relatively higher temperature or a different regional transportation pattern.

Response: Thank you for your valuable suggestions. We have removed the speculation:

# 1 Manuscript

# 2 Measurement Report: Elevated excess-NH<sub>3</sub> can

**3 promote the redox reaction to produce HONO:** 

- 4 Insights from the COVID-19 pandemic
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# 20 Abstract

21	HONO plays a crucial role as a precursor to OH radicals in the tropospheric atmosphere.
22	
22	The incongruity between HONO concentration and NO <sub>x</sub> emissions during the COVID-
23	19 pandemic remains puzzling. Here, we show evidence from field observations of ten
24	sites in China that there was a noticeable increase in NH3 concentrations during the
25	COVID-19 pandemic. In addition to the meteorological conditions, the significant
26	decrease in sulfate and nitrate concentrations enhanced the portioning of $NH_4^+$ to $NH_3$ .
27	Such conditions enable enhanced particle pH values, which in turn accelerate the redox
28	reactions between NO <sub>2</sub> and SO <sub>2</sub> to form HONO. This mechanism partly explains the
29	lower reduction of HONO concentration than that of NO <sub>2</sub> concentration during the
30	pandemic and highlights the importance of coordinating the control of SO <sub>2</sub> , NO <sub>x</sub> , and
31	NH <sub>3</sub> emissions.
32	Keywords: Ammonia, HONO, Gas-particle portioning, Acidity, COVID-19 pandemic

# **1. Introduction**

35	Nitrous acid (HONO) is a critical precursor of hydroxyl radical (OH), contributing
36	to more than 60% of OH production (Alicke, 2003; Platt et al., 1980; Kleffmann et al.,
37	2005). The OH can react with carbon monoxide, nitrogen oxides (NO <sub>x</sub> ), sulfur dioxide
38	(SO <sub>2</sub> ), and volatile organic compounds to produce secondary pollutants such as ozone
39	(O <sub>3</sub> ) and PM <sub>2.5</sub> (particulate matter with an aerodynamic diameter less than or equal to
40	2.5 $\mu$ m), thereby affecting air quality, human health, and global climate change (Li et
41	al., 2021a; Wang et al., 2023b; Lu et al., 2018)
42	High concentrations of HONO are present in urban daytime atmospheres, and
43	exploring its sources has become a hot and challenging topic in the field of atmospheric
44	chemistry (Jiang et al., 2022; Xu et al., 2019). Various sources of atmospheric HONO
45	have been identified, including combustion processes (e.g., vehicle emissions) (Kramer
46	et al., 2020; Liao et al., 2021; Li et al., 2021b), direct emissions from soil (Su and Zhang,
47	2011; Oswald et al., 2013; Meusel et al., 2018), homogeneous reactions between NO
48	and OH radicals (Pagsberg et al., 1997; Atkinson et al., 2004), heterogeneous reactions
49	of NO <sub>2</sub> on aerosols and ground surfaces (Zhang et al., 2020a; McFall et al., 2018; Liu
50	et al., 2014, 2020a), and photolysis of nitrate (Spataro and Ianniello, 2014; Scharko et
51	al., 2014; Romer et al., 2018; Ye et al., 2017; Shi et al., 2021). During the pandemic
52	control periods, there was a substantial reduction in vehicle traffic flow and industrial
53	emissions, leading to a decrease of more than 60% in NO <sub>x</sub> emissions in eastern China
54	(Huang et al., 2021a). It was initially expected that the concentration of HONO would
55	also decrease proportionally. However, Liu et al. (2020b) observed that the decrease in

HONO concentration during the pandemic period was only 31%, which was significantly lower than the reductions in NO (62%) and NO<sub>2</sub> (36%). Furthermore, the observed concentrations of HONO during the COVID-19 pandemic in 2020 were higher than those during the corresponding period in 2021 in Beijing (Luo et al., 2023). These findings suggest the existence of a considerable unknown source of HONO during the COVID-19 pandemic period.

62 Ammonia (NH<sub>3</sub>) is a primary alkaline gas in the atmosphere, capable of influencing the pH level of particulate matter and plays a crucial role in the atmospheric nitrogen 63 64 cycle (Gu et al., 2022; Xu et al., 2020; Gong et al., 2011). Several studies have indicated 65 that NH<sub>3</sub> can promote the formation of HONO by promoting the hydrolysis of NO<sub>2</sub> (Xu 66 et al., 2019) or the redox reaction of NO<sub>2</sub> with SO<sub>2</sub> (Liu et al., 2023). Moreover, 67 previous studies have reported that NH<sub>3</sub> concentrations in the atmosphere, particularly 68 in rural areas, significantly increased during the pandemic (Xu et al., 2022; Cui, 2023; 69 Zhang et al., 2020b). Consequently, the rise in NH<sub>3</sub> may contribute to the enhanced formation of HONO (Huang et al., 2021a). Unfortunately, there is currently a lack of 70 71 research on the relationship between enhanced NH<sub>3</sub> and HONO during the COVID-19 pandemic period. 72

To address this, online observational data on the chemical composition of PM<sub>2.5</sub>, gaseous pollutants, and meteorological conditions at ten sites in China before and during the COVID-19 pandemic period were analyzed to investigate the variation in NH<sub>3</sub> concentrations and particle pH, and explore the promoting effect of increased pH values on HONO formation.

# 78 2. Materials and methods

## 79 **2.1 Observation sites**

Online measurements were conducted at four urban and six rural sites in Henan Province, China from January 1 to February 29, 2020, including Sanmenxia (U-SMX), Zhoukou (U-ZK), Zhumadian (U-ZMD), and Xinyang (U-XY), as well as rural locations including Anyang (R-AY), Xinxiang (R-XX), Jiaozuo (R-JZ), Shangqiu (R-SQ), Nanyang (R-NY), and Puyang (R-PY). Descriptions and the spatial distribution of these ten sites can be found in Table S1 and Fig. S1.

### 86 **2.2 Measurements**

87 The aerosol and gas monitor (MARGA, Metrohm, Switzerland) was used to analyze the hourly water-soluble ions (Na<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Mg<sup>2+</sup>, Ca<sup>2+</sup>, Cl<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, and SO<sub>4</sub><sup>2-</sup>) in PM<sub>2.5</sub>, 88 89 as well as gaseous species (NH<sub>3</sub>, HNO<sub>3</sub>, HCl, and HONO) at ten sampling sites. The 90 MARGA instrument is widely used (Chen et al., 2017; Stieger et al., 2019; Twigg et al., 91 2022). A detailed description of the instrument and QA/QC can be found in Text S1. In 92 brief, the atmospheric sample passes through a PM<sub>2.5</sub> cut-off head, and both particles 93 and gases enter a wet rotating dissolution device for diffusion. Subsequently, the 94 particles in the sample undergo hygroscopic growth and condensation in an aerosol 95 supersaturated vapor generator, followed by collection and ion chromatographic 96 analysis. The gases in the sample are oxidized by H<sub>2</sub>O<sub>2</sub> in the dissolution device, 97 absorbed into a liquid solvent, and then entered the gas sample collection chamber for

98	ion chromatographic quantification. The range of minimum detection limits for water-
99	soluble ions was between 0.002 $\mu$ g/m <sup>3</sup> (Cl <sup>-</sup> ) to 0.081 $\mu$ g/m <sup>3</sup> (NH <sub>4</sub> <sup>+</sup> ). Uncertainties of 20%
100	are assumed for the detection of $NH_3$ and $NH_4^+$ , while uncertainties of 10% are assumed
101	for other components (Wang et al., 2020, 2022a). In addition, a detailed description of
102	HONO measurement using this system can be found in Text S2. Overall, the limit of
103	detection for HONO was 4 pptv and the uncertainty was estimated to be $\pm 20\%$ .
104	The data for NO <sub>2</sub> and SO <sub>2</sub> were obtained from a series of instruments provided by
105	Thermo Fisher Scientific (USA). The hourly concentrations of organic carbon (OC) in
106	PM <sub>2.5</sub> were analyzed using a carbon analyzer (Model 4, Sunset Laboratory., USA).
107	Detailed descriptions of the NO <sub>2</sub> , SO <sub>2</sub> , and carbon analyzers can be found in Text S3.
108	The smart weather stations (LUFFTWS500, Sutron, Germany) were utilized for
109	synchronized observation of meteorological parameters including pressure,
110	temperature (T), and relative humidity (RH).

# 111 **2.3 Data analysis.**

# 112 **2.3.1 pH prediction.**

The thermodynamic model ISORROPIA-II was used to estimate the pH value of the particles (Fountoukis, 2007) by inputting RH, T, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, total ammonia ( $TNH_x = 17 \times (\frac{[NH_4^+]}{18} + \frac{[NH_3]}{17})$ ), total sulfuric acid (TH<sub>2</sub>SO<sub>4</sub>, SO<sub>4</sub><sup>2-</sup>), total sodium (TNa, Na<sup>+</sup>), total chlorine (TCl, Cl<sup>-</sup>), and total nitrate (TNO<sub>3</sub> = NO<sub>3</sub><sup>-</sup> + HNO<sub>3</sub>). The model has two calculation modes: the forward mode and reverse mode, and the aerosol 118 dissolution systems can be set to simulate a metastable state (aqueous phase) or stable 119 state (aqueous and solid phase). Studies have shown that the forward mode is less 120 affected by instrument measurement errors than the reverse mode (Ding et al., 2019; 121 Song et al., 2018). Additionally, the minimum average RH of about 55% was recorded during the sampling period at the ten sites. Thus, ISORROPIA-II was run in the forward 122 123 model for the aerosol system in the metastable condition and only used data with RH  $\geq$ 30% for simulation accuracy (Ding et al., 2019; Song et al., 2018). The ISORROPIA 124 model calculated the particle hydrate ion concentration per volume of air  $(H_{air}^{+})$  and 125 aerosol water associated with inorganic matter (AWC<sub>inorg</sub>). The pH value was calculated 126 127 using the following equation (Bougiatioti et al., 2016):

128 
$$pH = -\log_{10}H_{aq}^{+} = -\log_{10}\frac{1000H_{air}^{+}}{AWC_{inorg} + AWC_{org}}$$
(2.1)

129 where the modeled concentrations for AWC<sub>inorg</sub> and  $H_{air}^+$  are  $\mu g/m^3$ , and AWC<sub>org</sub> is the

130 particle water associated with the organic matters predicted using the following method:

131 
$$AWC_{org} = \frac{m_s}{\rho_s} \frac{k_{org}}{\left(\frac{1}{RH} - 1\right)}$$
(2.2)

where  $m_s$  is the mass concentration of organic matter (OM = OC × *f*). The *f* is the conversion factor of OC, which is dependent on the extent of OM oxidation and secondary organic aerosol formation (Chow et al., 2015). Studies on the ratio of OM/OC in fourteen cities in China suggested that the mean value of *f* was  $1.59 \pm 0.18$ during the winter season in Northern China (Xing et al., 2013), and thus we adopted 1.6 as the *f* in this study. *k*<sub>org</sub> is the organic hygroscopicity parameter and depends on organic functionality, water solubility, molecular weight, and oxidation level. Han et al. (2022)

have reported that the  $k_{org}$  generally increased with O: C ratios, with a range of 0–0.3 139 140 for 23 organics, including carboxylic acids, amino acids, sugars, and alcohols. Gunthe et al. (2011) estimated a  $k_{org} = 0.06 \pm 0.01$  for the effective average hygroscopicity of 141 the non-refractory organic particulate matter in the aerosols in Beijing. Our previous 142 study has estimated that the uncertainties of  $k_{org}$  value (0.06) for pH in the range of 0– 143 0.3 only lead to -1-3% errors, which can be neglected (Wang et al., 2023a). Therefore, 144 the value of 0.06 was selected in this paper.  $\rho_s$  is the organic density, which was chosen 145 to be  $1.35 \text{ g/cm}^3$  following previous studies (Table S2). 146

### 147 **2.3.2 The sources of HONO**

The sources of HONO include direct emission ( $P_{emi}$ ), the homogeneous reaction of NO and •OH ( $P_{OH+NO}$ ), the heterogeneous reaction of NO<sub>2</sub> on the ground ( $P_{ground}$ ) and aerosol ( $P_{aerosol}$ ), the photo-enhanced heterogeneous reaction of NO<sub>2</sub> on the ground ( $P_{ground+hv}$ ) and aerosol ( $P_{aerosol+hv}$ ), and nitrate photolysis ( $P_{nitrate}$ ). The detailed calculation method is described in the Supplementary Material (Text S4, Table S3, Figs. S2 and S3)

## 154 **2.3.3 Redox reaction of NO<sub>2</sub> with SO<sub>2</sub>.**

The redox reaction of NO<sub>2</sub> with SO<sub>2</sub> (R<sub>1</sub>) is considered a crucial potential source of high concentrations of HONO in Northern China (Cheng et al., 2019; Wang et al., 2016b):

158 
$$S(IV) + 2NO_2 + H_2O \rightarrow S(VI) + 2H^+ + 2NO_2^-$$
(R<sub>1</sub>)

159 the rate expression for reaction  $(R_1)$  was estimated to:

160 
$$d[S(VI)] / dt = k_1[NO_2][S(VI)],$$
 (2.3)

161 the rate constant  $k_1$  value is pH dependent, e.g., for pH, 5,  $k_1 = (1.4 \times 10^5 + 1.24 \times 10^7)/2$ 

162  $M^{-1} s^{-1}$ . For  $k_1$  values under other pH conditions and other related information, please

163 refer to Text S5, Table S4, and Table S5.

# 164 **3. Results and discussion**

# 165 **3.1 Variations of NH<sub>3</sub>, NH<sub>4</sub><sup>+</sup> and TNH<sub>x</sub>.**

166 The temporal variations of  $NH_3$ ,  $NH_4^+$ , and  $TNH_x$  at 10 sampling sites in the pre-COVID-19 pandemic period (PC, January 1 to 23, 2020) and during the COVID-19 167 pandemic period (DC, January 24 to February 29, 2020) are presented in Fig. 1, with 168 169 their average concentration listed in Table 1. In general, rural sites exhibited higher concentrations of NH<sub>3</sub>,  $NH_4^+$ , and  $TNH_x$  compared to urban sites, except for the R-NY 170 171 site. This finding is consistent with previous studies conducted in Zhengzhou (Wang et 172 al., 2020), Shanghai (Chang et al., 2019), and Quzhou (Feng et al., 2022a), owing to 173 the intense agricultural ammonia emissions. The highest concentrations of NH<sub>3</sub> and 174 TNH<sub>x</sub> were recorded at site R-JZ, with average values of  $25.3 \pm 11.5$  and  $40.8 \pm 20.1$  $\mu$ g/m<sup>3</sup>, respectively. Site R-AY had the highest NH<sub>4</sub><sup>+</sup> concentration, measuring 19.3 ± 175 176 12.9  $\mu$ g/m<sup>3</sup>. Note that the current study area exhibited higher NH<sub>3</sub> levels compared to 177 other regions (Table S6), which probably was attributed to the highest NH<sub>3</sub> emissions of Henan Province in China, primarily from nitrogen fertilizer application and livestock 178 179 farming (Wang et al., 2018; Ma, 2020).

180	Compared to the PC, NH <sub>3</sub> concentrations increased in the DC at all sites. Notably,
181	rural sites experienced more significant increases in NH3 concentrations than urban
182	sites, which was similar to the trend in Shanghai (Xu et al., 2022). The largest increases
183	in NH <sub>3</sub> concentrations were observed at R-SQ (71%, 7.3 $\mu$ g/m <sup>3</sup> ) and U-ZK (37%, 4.8
184	$\mu g/m^3)$ for rural and urban sites, respectively. In contrast, the concentrations of $NH_4^{\scriptscriptstyle +}$
185	and $\text{TNH}_x$ decreased in the DC with the largest reduction at rural site R-PY (51%, 12.9
186	$\mu$ g/m <sup>3</sup> ) and urban site U-ZMD (48%, 9.3 $\mu$ g/m <sup>3</sup> ). Regarding TNH <sub>x</sub> , rural sites exhibited
187	larger reductions, with site R-SQ experiencing the largest decrease of 37% (4.7 $\mu$ g/m <sup>3</sup> ).
188	Figure 2 illustrates the spatial distribution and the diurnal variation of NH <sub>3</sub>
189	concentrations at the ten sites before and during the pandemic. NH <sub>3</sub> concentrations in
190	most sites exhibited an unimodal trend in PC that NH3 concentrations gradually
191	increased after sunrise, reaching a peak around noon (11:00–12:00), and then decreased
192	to a valley around 16:00–17:00. This diurnal pattern is similar to NH <sub>3</sub> variations
193	observed in urban areas of Houston, USA, as a result of the natural emissions from
194	vegetation and soil during photosynthesis (Gong et al., 2011). However, other studies
195	have recorded a significant NH <sub>3</sub> peak during the early morning of 8:00–10:00 (Ellis et
196	al., 2011; Meng et al., 2018; Gu et al., 2022), suggesting the influence of vehicle
197	emissions (Gong et al., 2011; Gu et al., 2022), residual NH3 mixing, soil or plant
198	emissions (Ellis et al., 2011), and dew volatilization (Wentworth et al., 2016; Huang et
199	al., 2021b). Therefore, the NH <sub>3</sub> in urban and rural areas of this study was probably less
200	affected by NH <sub>3</sub> emissions from vehicles, different from the recent studies in megacities
201	of China (e.g., Beijing and Shanghai) (Gu et al., 2022; Wu et al., 2023; Zhang et al.,
202 2020b). In addition to the transport from agricultural emissions, urban NH<sub>3</sub> in this 203 region might also originate from other non-agricultural sources, such as wastewater 204 treatment, coal combustion, household waste, urban green spaces, and human 205 excrement (Chang et al., 2019).

206 During the COVID-19 pandemic, the diurnal variation of NH<sub>3</sub> in both urban and rural 207 sites still maintained an unimodal distribution. The peak values in urban sites remained 208 consistent with PC levels, further demonstrating that the influence of vehicles on NH<sub>3</sub> in urban areas was limited. Notably, the peak time of NH<sub>3</sub> in rural sites shifted 1–2 hours 209 210 earlier compared to the trend in PC. Ammonia in rural areas primarily originates from 211 nitrogen fertilizer application, livestock, and poultry breeding (Feng et al., 2022b; 212 Meng et al., 2018), which are significantly influenced by T and RH (Liu et al., 2023). 213 Table S7 and Fig. S4 reveal that there was an increased T and a decreased RH at rural 214 sites in the DC than the PC, which could accelerate the evaporation of NH<sub>3</sub> and thus 215 potentially lead to earlier peak NH<sub>3</sub> concentrations.

#### 216 **3.2 Gas-to-particle conversion of NH**<sub>3</sub>

# The increased NH<sub>3</sub> accompanying decreased NH<sub>4</sub><sup>+</sup> in the DC suggests that the gasparticle partition of NH<sub>3</sub>/ NH<sub>4</sub><sup>+</sup> may determine the elevated NH<sub>3</sub> concentrations. Meteorological parameters, including RH and T, play a crucial role in the gas-particle partitioning of NH<sub>3</sub> (Liu et al., 2023; Xu et al., 2020). Therefore, the higher T and lower RH in the DC (Table S7 and Fig. S4) favored the conversion of NH<sub>4</sub><sup>+</sup> to NH<sub>3</sub>, resulting in a decrease in $\varepsilon$ (NH<sub>4</sub><sup>+</sup>) ([NH<sub>4</sub><sup>+</sup>]/([NH<sub>3</sub>] + [NH<sub>4</sub><sup>+</sup>]) compared to those in the PC (Table

223 S7).

NH<sub>3</sub> primarily enters particles to neutralize acidic ions (Wang et al., 2020; Xu et al., 2020; Liu et al., 2017; Ye et al., 2011; Wells, 1998). Accordingly, the concentrations of required ammonia (Required-NH<sub>x</sub>) and excess ammonia (Excess-NH<sub>x</sub>) were calculated based on the acidic substances as follows (Wang et al., 2020):

228  
Required-NH<sub>x</sub> = 17 × 
$$(\frac{[SO_4^{2^-}]}{48} + \frac{[NO_3^{-}]}{63} + \frac{[CI^-]}{35.5} + \frac{[HNO_3]}{64} + \frac{[HCI]}{36.5})$$
 (3.1)  
 $-17 \times (\frac{[Na^+]}{23} + \frac{[K^+]}{39} + \frac{[Ca^{2^+}]}{20} + \frac{[Mg^{2^+}]}{12})$ 

229  $Excess-NH_x = TNH_x - Required-NH_x$  (3.2)

where [W] represents the concentration of the substance ( $\mu g/m^3$ ). The significant linear fitting (R<sup>2</sup> is greater than 0.96, and the slope is close to 1) in Fig. S5 demonstrates that the anions and cations at each site were close to the equilibrium state. Therefore, the organic acids in PM<sub>2.5</sub> may have less effect on NH<sub>3</sub> and NH<sub>4</sub><sup>+</sup> and were not considered in Formula 3.1.

235 As shown in Fig. 3 and Table S8, compared to those in the PC, the concentration of 236 Required-NH<sub>x</sub> in the DC significantly decreased (ranging from 37% at site R-JZ to 58% 237 at site R-PY), while the concentration of Excess-NH<sub>x</sub> increased (ranging from 9% at 238 site R-AY to 78% at site R-SQ). The reduction in the concentrations of sulfate and 239 nitrate (Fig. S6) was responsible for the decrease in the concentration of Required- $NH_x$ . 240 To sum up, in addition to meteorological conditions, the substantial reduction in anthropogenic emissions of SO<sub>2</sub>, NO<sub>x</sub>, and other pollutants in the DC had led to a 241 242 decrease in acidic substances (e.g., sulfate and nitrate) in particles, in turn, resulting in 243 more gas-phase NH<sub>3</sub> concentration remaining in the atmosphere.

**3.3 Particle pH before and during COVID-19** 

245	Diurnal patterns of particle pH in PC and DC at ten sites are summarized in Fig. 4
246	with their average values listed in Table S9. PM <sub>2.5</sub> shows consistent moderate acidity,
247	with mean values in the range of 4.2–5.1, which were close to the values in previous
248	studies (Table S9). Compared to the PC, the particle pH at ten sites increased obviously
249	in the DC, with the highest increase of 0.5 (U-ZK) and 0.3 (R-PY) at urban and rural
250	sites, respectively, which were the subject of in-depth discussion in the following text.
251	To explore the dominant factors that determine the local particle pH level and result
252	in the high pH during the DC, sensitivity tests of pH to chemical species (i.e., $TNH_x$ ,
253	TH <sub>2</sub> SO <sub>4</sub> , TNO <sub>3</sub> , TCl, TNa, $K^+$ , Ca <sup>2+</sup> , and Mg <sup>2+</sup> ) and meteorological parameters (i.e., T
254	and RH) were performed. A given range for a variable (i.e., TNH <sub>x</sub> ) with corresponding
255	average values of other parameters (i.e., $TH_2SO_4$ , $TNO_3$ , $TCl$ , $TNa$ , $K^+$ , $Ca^{2+}$ , $Mg^{2+}$ , T,
256	and RH) was input into the model and simulated to compare its effects on pH. As shown
257	in Fig. S7, pH increases with the cation concentrations (i.e., $TNH_x$ , $Na^+$ , $K^+$ , $Ca^{2+}$ , and
258	$Mg^{2+}$ ) increasing as well as the anion concentrations (i.e., $TH_2SO_4$ , $TNO_3$ , and $CI^-$ ), T
259	and RH decreasing. According to the average values of input data during PC (Blue line
260	in Fig. S7) and DC (Red line in Fig. S7) at U-ZK and R-PY sites respectively, the
261	changes in pH ( $\Delta$ pH in Fig. 5) indicate that the decrease in TNH <sub>x</sub> concentration and the
262	increase in T in DC led to a decrease in pH values ( $\Delta$ pH: 0.09 at U-ZK and 0.08 at R-
263	PY sites) compared to PC. However, this effect was outweighed by the decrease in
264	TH <sub>2</sub> SO <sub>4</sub> (ApH: 0.07 and 0.8 at U-ZK and R-PY sites, respectively) and TNO <sub>2</sub> (ApH:
265	0.05 and 0.4 at U-ZK and R-PY sites respectively) concentrations as well as the



274 factor in promoting the pH values.

### **3.4 The influence of pH on HONO.**

The observed HONO concentrations decreased by 18% and 54% at U-ZK (0.8 ppb) 276 277 and R-PY (0.9 ppb) sites in the DC, respectively, compared to those (1.0 and 2.2 ppb) 278 in the PC. Moreover, all the known HONO production sources rates including Pemi, POH + NO, Pground, Pground+hy, Paerosol, Paerosol+hy, and Pnitrate (Fig. 7) show a decreasing trend from 279 280 PC to DC, with the total reductions of 42% and 80% for U-ZK and R-PY, respectively. 281 At the U-ZK, Pground+hv decreased the most (84%), while at the R-PY, Pnitrate had the 282 largest decrease about 85%, which was speculated to be related to the decrease of NO<sub>x</sub> 283 and NO<sub>3</sub><sup>-</sup> concentration in DC. Note that the reduction rates in the overall known source 284 and almost individual sources were greater than the reduction rates in HONO 285 concentrations (Figs. 7 and 8), thus we hypothesized that there should be other sources 286 capable of promoting HONO production. Soil emission has been demonstrated to be a major source of HONO, which is affected by temperature to some extent (Liu et al., 2020b, 2020c). However, there was no significant positive correlation with temperature in Fig. S8, and temperatures did not exceed 10°C during the study periods, suggesting that soil emission may not be a major contributor to HONO. Note that there were positive correlations between HONO with SO<sub>2</sub>, Excess-NH<sub>x</sub>, SO<sub>4</sub><sup>2-</sup>, and pH (Fig. S8) indicating that the R<sub>1</sub> reaction might form an amount of HONO and contribute to less reduction in the observed HONO concentrations.

294 Considering that  $R_1$  mainly reacts in the liquid phase, the calculated reaction rates of 295  $R_1$  under the conditions of RH > 60% in the PC and DC periods are illustrated in Figs. 296 8 and S9. Despite the decrease in NO<sub>2</sub> and SO<sub>2</sub> concentrations in the DC, the increase 297 in particle pH, increasing HSO<sub>3</sub> concentration in the aqueous phase, promoted the  $R_1$ 298 reaction rates by 58% and 59% at U-ZK and R-PY (Figure 8), respectively. Consequently, the enhanced R<sub>1</sub> reaction might prevent a large decrease in HONO (18% 299 300 at U-ZK and 53% at R-PY) under the conditions of a significant reduction in vehicle 301 emissions and a decline of 66% and 69% in NO<sub>2</sub> concentrations at U-ZK and R-PY, respectively. 302

#### 303 **3.5 Uncertainty**

According to sensitivity tests of pH (Fig. S7) and  $R_1$  (Fig. S10), pH increases with the concentrations of cations (TNH<sub>x</sub>, TNa, K<sup>+</sup>, Ca<sup>2+</sup>, and Mg<sup>2+</sup>) and OC increasing as well as anions (TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, and Cl<sup>-</sup>) concentrations, T, and RH decreasing.  $R_1$ reaction rate increases with the concentrations of AWC, NO<sub>2</sub>, SO<sub>2</sub>, pH, and pressure, 308 while increasing as well as T decreasing. Therefore, two extreme scenarios (i.e., the 309 maximum and minimum rate scenarios) were evaluated to estimate the uncertainty of 310 pH, and  $R_1$  based on the measurement uncertainties at the U-ZK and R-PY sites. Figure 311 S11 suggests that the two extreme scenarios can be led to -10-7% and -71-125%312 uncertainties at the U-ZK site and -10-7% and -78-123% uncertainties at the R-PY 313 site for pH and  $R_1$ , respectively.

## 314 **4. Conclusions**

315 Elevated NH<sub>3</sub> concentration was observed during the COVID-19 pandemic at both 316 urban and rural sites in China. In addition to the rise in T and decrease in RH during the 317 COVID-19 pandemic, which favored the conversion of  $NH_4^+$  to  $NH_3$ , the significant decrease in sulfate and nitrate concentrations led to the decline in Required-NH<sub>x</sub> and 318 319 was beneficial to the particle-phase NH<sub>4</sub><sup>+</sup> portioning to gas-phase NH<sub>3</sub>. Furthermore, 320 under the environmental conditions of increased NH<sub>3</sub> concentration and decreased acidic substance concentration, the pH values increased by 0.5 and 0.3 at U-ZK and R-321 322 PY increased during the pandemic, respectively. Consequently, the high pH values 323 accelerated the formation rate of HONO through the oxidation-reduction reaction of NO<sub>2</sub> with SO<sub>2</sub> (an increase of 58% at U-ZK and 59% at R-PY, respectively), partially 324 325 compensating for the decrease in HONO concentration caused by the decline in vehicle 326 emissions, NO<sub>2</sub> and NO<sub>3</sub><sup>-</sup> concentrations during the COVID-19 pandemic.

## 327 **5. Implications**

328 HONO plays a crucial role as a precursor to OH radicals in the tropospheric atmosphere (Xue, 2022). There have been significant observations of high HONO 329 330 concentrations in urban areas during the daytime, leading to a growing interest in 331 understanding its sources in atmospheric chemistry (Jiang et al., 2022; Xu et al., 2019). 332 The heterogeneous reaction mechanism of NO<sub>2</sub> on aerosol surfaces is currently the 333 focus of research on HONO sources, particularly in regions with elevated levels of 334 atmospheric particulate matter, where it could potentially become a major contributor to HONO production (Zhang et al., 2022; Liao et al., 2021). One of the pathways for 335 heterogeneous reactions on aerosol surfaces is the redox reaction of NO<sub>2</sub> with SO<sub>2</sub>. 336 337 However, the significance of this reaction in HONO production in the real atmosphere 338 is often overlooked, as it relies on the high pH of aerosols (Ge et al., 2019). In recent 339 years, there has been increasing attention on the enhancing effect of NH<sub>3</sub> on the redox 340 reaction, with laboratory experiments demonstrating its ability to generate substantial amounts of HONO (Ge et al., 2019). This study highlights the importance of this 341 342 reaction based on actual atmospheric observations. Furthermore, numerous studies 343 have indicated that if control over NH<sub>3</sub> emissions continues to relax while SO<sub>2</sub> and NO<sub>2</sub> 344 emissions decrease, the particle pH in future China is expected to rise steadily (Xie et al., 2020; Song et al., 2019; Wang et al., 2020). Consequently, the redox reaction of 345 NO<sub>2</sub> with SO<sub>2</sub> could become a significant source of HONO in China. Therefore, it is 346 347 crucial to coordinate the control of SO<sub>2</sub>, NO<sub>x</sub>, and NH<sub>3</sub> emissions to avoid a rapid 348 increase in the particle pH.

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350 Data availability: All the data presented in this article can be accessed through
351 https://zenodo.org/records/10273539. (Zhang, 2023).

352

353 A	uthor c	ontributions.	XZ Data	Curation,	Writing -	Origin	al Draft,	Visualization.
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354 LW, NW, SM, and DZ Investigation, Visualization, Data Curation. DZ, HZ, and MW

355 Investigation. SW Conceptualization, Data Curation, Supervision. RZ Data Curation,

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359

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## 656 Figures:



#### 657

Figure 1. Temporal variations of a.  $NH_3$ , b.  $NH_4^+$ , and c.  $TNH_x$  at the urban and rural

659 sites before (PC) and during (DC) the COVID-19 outbreak, respectively. The shaded

areas of the curve represent the maximum and minimum values.



664 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.



666 Figure 3. Box diagram of changes in Required-NH<sub>x</sub> at ten sites before (PC) and during

667 (DC) the COVID-19 outbreak. In each box, the top, middle, and bottom lines represent

the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower

669 whiskers represent the 90 and 10 percentiles of statistical data, respectively.





#### 675 Figure 5. Changes of pH ( $\Delta$ pH) through the sensitivity tests (Figure S5 and S6) by

#### 676 changing parameters between PC and DC at the a. U-ZK and b. R-PY sites.



678 Figure 6. Particle pH corresponds to increasing TNH<sub>x</sub> at U-ZK and R-PY sites to examine the effects of major indicators of  $NH_3$  (i.e.,  $TNH_x$ , Required- $NH_x$ , and Excess-679 680  $NH_x$ ) on aerosol acidity. Particle pH was calculated by using a wide range of  $TNH_x$ 681  $(25-130 \,\mu\text{g/m}^3)$  and average values of other parameters in PC and DC of U-ZK and R-PY sites. The concentrations of  $TNH_x$ , Required-NH<sub>x</sub>, and Excess-NH<sub>x</sub> with 682 corresponding pH values are marked by a hollow box, hollow circle, and arrow 683 respectively. The yellow and blue background colors correspond to the NH<sub>x</sub>-poor and 684 NH<sub>x</sub>-rich, respectively. 685



Figure 7. Comparison of HONO sources at a. U-ZK and b. R-PY sites before (PC) and

during (DC) the COVID-19 outbreak. The calculation method can be found in Text S4.



Figure 8. Decline ratios of a. NO<sub>2</sub>, b. HONO concentration, and c. HONO production rate at U-ZK and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The center point represents the mean value, and the upper and lower whiskers represent the 95% confidence interval of the mean.

**Table:** 

 $\label{eq:constraint} 696 \qquad \text{Table 1. Changes in concentrations (mean $\pm$ standard deviation) of $NH_3$, $NH_4^+$, and $NH_4^$ 

TNH<sub>x</sub> at ten sites during entire periods (Average), before (PC), and during (DC) the

698 COVID-19 outbreak.

Sites	Substances	Average (µg/m <sup>3</sup> )	PC ( $\mu g/m^3$ )	DC ( $\mu g/m^3$ )
U-SMX	NH <sub>3</sub>	$13.8\pm10.8$	$12.6 \pm 10.1$	$14.5 \pm 11.1$
	$\mathrm{NH}_4^+$	$10.9\pm7.2$	$14.2 \pm 7.2$	$8.8\pm 6.5$
	TNH <sub>x</sub>	$22.9\pm14.1$	$24.9 \pm 14.5$	$21.7\pm13.8$
U-ZK	NH <sub>3</sub>	$15.6\pm8.3$	$12.7 \pm 6.5$	$17.4\pm8.8$
	$\mathrm{NH}_4^+$	$13.6\pm9.3$	$19.1\pm8.4$	$10.3\pm8.1$
	TNH <sub>x</sub>	$28.6\pm13.7$	$30.9\pm12.8$	$27.1\pm14.0$
U-ZMD	NH <sub>3</sub>	$13.1\pm8.4$	$11.6 \pm 8.2$	$14.0\pm8.4$
	$\mathrm{NH}_4^+$	$13.9\pm9.8$	$19.6\pm10.3$	$10.3\pm7.5$
	TNH <sub>x</sub>	$25.7\pm14.6$	$30.3 \pm 15.1$	$22.8\pm13.5$
U-XY	NH <sub>3</sub>	$7.0\pm4.3$	$5.7 \pm 4.0$	$7.9\pm4.3$
	$\mathrm{NH}_4^+$	$11.0\pm7.7$	$15.4 \pm 7.6$	$8.3\pm6.5$
	TNH <sub>x</sub>	$17.6\pm9.8$	$20.6\pm10.1$	$15.7\pm9.2$
R-AY	NH <sub>3</sub>	$19.0\pm8.4$	$17.9\pm8.3$	$19.7\pm8.4$
	$\mathrm{NH}_4^+$	$19.3\pm12.9$	$26.4 \pm 13.7$	$15.0\pm10.3$
	TNH <sub>x</sub>	$36.6\pm18.2$	$41.7\pm20.4$	$33.4\pm16.0$
R-XX	NH <sub>3</sub>	$21.7\pm10.2$	$18.1\pm9.3$	$23.8\pm10.1$
	$\mathrm{NH}_4^+$	$15.9\pm10.4$	$20.6\pm11.0$	$13.0\pm8.8$
	TNH <sub>x</sub>	$34.9\pm17.0$	$35.1\pm18.8$	$34.8\pm15.8$
R-PY	NH <sub>3</sub>	$19.8\pm9.4$	$16.8 \pm 8.1$	$21.7\pm9.6$
	$\mathrm{NH}_4^+$	$17.4 \pm 11.8$	$25.3 \pm 12.6$	$12.4\pm8.0$
	TNH <sub>x</sub>	$35.2\pm17.8$	$39.4 \pm 19.8$	$32.6 \pm 15.7$
R-JZ	NH <sub>3</sub>	$25.3\pm11.5$	$24.1 \pm 11.5$	$25.9 \pm 11.4$
	$\mathrm{NH}_4^+$	$17.3 \pm 11.3$	$22.7 \pm 11.6$	$14.2\pm9.9$
	TNH <sub>x</sub>	$40.8\pm20.1$	$42.9\pm22.8$	$33.5\pm18.2$
R-SQ	NH <sub>3</sub>	$15.0\pm7.9$	$10.3 \pm 5.2$	$17.7\pm7.9$
	$\mathrm{NH}_4^+$	$13.4\pm8.5$	$18.9\pm8.6$	$10.3\pm6.7$
	TNH <sub>x</sub>	$26.3\pm13.2$	$25.5 \pm 14.0$	$26.8\pm12.7$
R-NY	NH <sub>3</sub>	$5.5 \pm 3.1$	$4.3 \pm 2.7$	$6.2 \pm 3.2$
	$\mathrm{NH}_4^+$	$10.2\pm6.9$	$13.3 \pm 7.2$	$8.4\pm6.1$
	TNH <sub>x</sub>	$14.8\pm8.5$	$16.0\pm9.5$	$14.1\pm7.8$

## 1 Supplement materials:

## 2 Text S1 Detailed description of the aerosol and gas monitor.

3	The aerosol and gas monitor (MARGA, Metrohm, Switzerland) was used to analyze
4	the hourly water-soluble ions (Na <sup>+</sup> , NH4 <sup>+</sup> , K <sup>+</sup> , Mg <sup>2+</sup> , Ca <sup>2+</sup> , Cl <sup>-</sup> , NO <sub>3</sub> <sup>-</sup> , and SO <sub>4</sub> <sup>2-</sup> ) in
5	PM <sub>2.5</sub> , as well as gaseous species (NH <sub>3</sub> , HNO <sub>3</sub> , HCl, and HONO) at ten sampling sites.
6	The atmospheric sample passes through a PM <sub>2.5</sub> cut-off head, and both particles and
7	gases enter a wet rotating dissolution device for diffusion. Subsequently, the particles
8	in the sample undergo hygroscopic growth and condensation in an aerosol
9	supersaturated vapor generator, followed by collection and ion chromatographic
10	analysis. The gases in the sample are oxidized by H2O2 in the dissolution device,
11	absorbed into a liquid solvent, and then entered the gas sample collection chamber for
12	ion chromatographic quantification. During this process, the sample is extracted
13	through a liquid diffusion filter, where interfering acidic and alkaline gases are removed
14	To achieve high collection efficiency, the airflow containing loaded ions then enters an
15	aerosol supersaturated collector until the particles can be injected into the ion
16	chromatograph. The ion chromatography system utilizes either suppressor or non-
17	suppressor conductivity detection methods for ion analysis. Before running the samples
18	the ion chromatograph system needs to be calibrated using standard solutions. By
19	comparing the data obtained from the sample with the data obtained from known
20	standard solutions, the identification and quantification of sample ions can be

21 performed. The data acquisition system generates chromatograms, and the 22 chromatography software further converts each peak in the chromatogram into sample 23 concentrations and outputs the results.

24 **QA/QC** 

25 The instrument underwent daily checks and maintenance, which typically involved ensuring the stability of the internal standard response and maintaining a relative error 26 27 within  $\pm 10\%$  between the measured and theoretical concentrations of the internal 28 standard. The system's data acquisition and transmission were carefully examined, 29 along with monitoring the instrument's status information and collected data. This 30 included checking parameters such as sampling flow rate, chromatographic column 31 pressure, column temperature, conductivity, target compound peak retention time, and 32 peak width to ensure their normal functioning. Regular replacement of consumables 33 used by the instrument was carried out at predetermined intervals and frequencies. 34 Additionally, standard curve measurements and calibration were performed in each 35 season to guarantee the accuracy of the instrument's data. Calibration curve verification 36 was performed at least once per quarter. A standard series containing at least 6 calibration points, including zero concentration, was prepared using standard solutions. 37 The concentration range of the calibration curve was set according to the actual 38 39 environmental concentration levels and determined by manual injection. The obtained calibration curve had a linear correlation coefficient (r) of  $\geq 0.995$ . If this requirement 40 was not met, the rationality of the internal standard solution concentration settings 41

42	would be checked. When key components such as the quantitative loop,
43	chromatographic column, or suppressor are replaced, a new calibration curve will be
44	promptly established. After establishing the new calibration curve, the sample sequence
45	in the analysis software would be updated. The minimum detection limit was
46	determined as follows: 0.002 $\mu$ g/m <sup>3</sup> (Cl <sup>-</sup> ), 0.081 $\mu$ g/m <sup>3</sup> (NH <sub>4</sub> <sup>+</sup> ), 0.02 $\mu$ g/m <sup>3</sup> (NO <sub>3</sub> <sup>-</sup> ), 0.06
47	$\mu$ g/m <sup>3</sup> (SO <sub>4</sub> <sup>2-</sup> ), 0.002 $\mu$ g/m <sup>3</sup> (Na <sup>+</sup> ), 0.08 $\mu$ g/m <sup>3</sup> (K <sup>+</sup> ), 0.06 $\mu$ g/m <sup>3</sup> (Ca <sup>2+</sup> ) and 0.007 $\mu$ g/m <sup>3</sup>
48	( $Mg^{2+}$ ). According to the research about the uncertainties in MARGA measurements,
49	such as a 20% uncertainty for MARGA measurements (Song et al., 2018), an error of
50	10% for detecting $SO_4^{2-}$ , $NO_3^{-}$ , and $HNO_3$ , and 15% for $NH_3$ and $NH_4^+$ (Rumsey et al.,
51	2014), we set the uncertainties of 20% for $NH_3$ and $NH_4^+$ , 10% for other components.
52	Due to the complex conditions encountered during the actual sampling process,
53	including variations in calibration slopes, sampling volumes, solvent concentrations,
54	temperature, atmospheric pressure, and sampling flow rates at different sampling points,
55	the assumed values mentioned above may not accurately reflect the actual situation.

#### 56 Text S2 HONO measurement

57 The HONO monitoring method adopted in MARGA is the wet-flow diffusion tube method (WEDD) in the diffusion tube method (Zellweger, 1999; Takeuchi et al, 2013), 58 59 which is a common method for measuring HONO in wet chemistry and has high 60 absorption efficiency. The device adopts a vertical setting, through the diffusion tube, 61 the air in the atmosphere is pulled upward from the bottom, and the absorbent liquid is 62 transported to the top of the diffusion tube through the air pump. When flowing under the action of gravity, a thin absorbent liquid film will be generated on the inner surface 63 64 of the tube by the tension. The absorbent liquid film will absorb HONO, and the solution at the bottom of the diffusion tube will be sucked out through the air pump. 65 Then it is sent to the ion chromatography for analysis. The integration time of the 66 67 sample mainly depends on the running time of the ion chromatography, which is about 5-30 min (Zellweger, 1999; Takeuchi et al, 2013). Based on the original, some scholars 68 69 developed the flow injection-chemiluminescence method and used it together with 70 WEDD for the measurement of HONO. The detection limit is about 0.03  $\mu$ g/m<sup>3</sup> 71 (Mikuska et al., 2008; Zhao et al., 2010). In addition, HONO observations measured 72 with this AIM-IC system agree well with HONO observations measured with the other 73 systems (VandenBoer et al., 2014). Therefore, it is feasible to measure HONO using 74 this instrument.

#### Text S3 Detailed description of the NO<sub>2</sub>, SO<sub>2</sub> and carbon analyzer.

The NO<sub>2</sub> analyzer utilized the chemiluminescence technique to measure the 76 concentration of NO<sub>2</sub> in the air. This involved converting NO<sub>2</sub> to NO using a 77 molybdenum converter, and then quantifying the NO concentration. The principle 78 79 behind the SO<sub>2</sub> analyzer involved measuring the amount of ultraviolet light emitted 80 during the decay of high-energy state SO<sub>2</sub>. This emitted light was used to calculate the 81 concentration of SO<sub>2</sub>. 82 The carbon analyzer principle is primarily based on the NIOSH-5040 method, which 83 involves analyzing the thermal optical transmittance of quartz filter samples. It employs 84 a calibrated non-dispersive infrared sensor to detect the evolving carbon. Under controlled conditions with inert helium gas, carbon formed during a gradually 85 increasing temperature gradient is referred to as OC, while carbon evolved under a 86

87 mixture of 90% helium.

#### 88 **Text S4 Sources of HONO**

#### 89 4.1 Direct emission

90	HONO can be released directly into the atmosphere through vehicle exhaust
91	(Burling et al., 2010: Veres et al., 2010). The lifetime of HONO in the atmosphere is
02	relatively short so vehicle emissions significantly contribute to urban atmospheric
)2	relatively short, so vehicle emissions significantly contribute to urban autospheric
93	HONO (Chen et al., 2023; Liu et al., 2021a). Considering that there has been a

- 94 significant reduction in vehicle emissions in urban areas during DC. Additionally, the R-PY site is far from roads. Thus, vehicle emissions may not be the primary source of 95 96 HONO for the U-ZK site during DC and R-PY sites during entire periods. To further 97 validate the above conclusions, the conditional bivariate probability function diagrams 98 of NO<sub>2</sub> at U-ZK and R-PY sites during PC and DC are depicted in Figure S2. NO<sub>2</sub> 99 predominantly originated from long-distance transport at the U-ZK site during DC and 100 the R-PY site during both PC and DC. Consequently, vehicle emissions are only calculated for the U-ZK site during the PC. 101 102 Here we use the HONO/NO<sub>x</sub> ratio to estimate HONO concentration, which is 103 generally considered to be the vehicle emission factor (Kramer et al., 2020; Hao et al., 104 2020; Yu et al., 2022) for HONO. The calculation formula is as follows: 105  $[\text{HONO}_{\text{emi}}] = 0.8\% \times [\text{NO}_{\text{x}}]$ (1)106 where [HONO<sub>emi</sub>] and [NO<sub>x</sub>] represent the HONO concentration emitted by vehicles 107 and the observed NO<sub>x</sub> concentration, respectively. Regarding previous studies (Table
- 108 S3), 0.8% was selected as the vehicle emission factor, considering differences in vehicle
- 109 type, fuel composition, and other factors (Kramer et al., 2020; Hao et al., 2020; Huang
- 110 et al., 2017).
- 111 4.2 Homogeneous reaction of NO and •OH
- 112 The reaction between NO and •OH is the primary gas-phase reaction source of
- 113 HONO at high NO concentrations, and the production rate contribution (POH+NO) for this

115 
$$P_{OH+NO} = k_{OH+NO}[OH][NO]$$
(2)

- 116 where  $k_{OH+NO}$  (7.2 × 10<sup>-12</sup> cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup>) is the rate constant for the reactions at
- 117 298K (Li et al., 2012). •OH concentration was simulated according to the empirical
- 118 model (Hu et al., 2022; Wang et al., 2025):

119 
$$[OH] = 4.1 \times 10^9 \times \frac{J(O^1D) \times J(NO_2) \times (140 \times [NO_2] + 1) + [HONO] \times J(HONO)}{0.41 \times [NO_2]^2 + 1.7 \times [NO_2] + 1 + [NO] \times k_{NO+OH} + [HONO] \times k_{NO+OH}} (3)$$

- 120 where, J (O<sup>1</sup> D), J (NO<sub>2</sub>), and J (HONO) are the photolysis rates calculated using the
- 121 TUV model (v5.2; available at http://cprm.acom.ucar.edu/Models/TUV/). The cloud
- 122 optical depth value for the TUV model was adjusted so that the predicted UVB radiation
- 123 intensity matched the observations (Lyu et al., 2019; Wang et al., 2022b). The calculated
- <sup>124</sup>•OH concentration varied from  $0.1 \times 10^6$  to  $4 \times 10^6$  molecule/cm<sup>3</sup> at U-ZK and  $0.1 \times 10^6$
- 125  $10^6$  to  $5 \times 10^6$  molecule/cm<sup>3</sup> t R-PY, which was comparable to the levels in other cities
- 126 of North China (Li et al., 2018; Fuchs et al., 2017; Yang et al., 2017). Since there is no
- 127 photolysis at night, the •OH concentration was assumed to be  $0.8 \times 10^6$  molecule/cm<sup>3</sup>
- 128 (Wang et al., 2022).
- 129 4.3 Heterogeneous conversion of NO<sub>2</sub> to HONO
- 130 4.3.1 Heterogeneous dark reactions
- 131 The heterogeneous conversion of NO<sub>2</sub> to HONO on the ground (P<sub>ground</sub>) and on the
- 132 aerosol surface (P<sub>aerosol</sub>) was calculated based on parameters obtained from experiments
- 133 or observations.
134 
$$\mathbf{P}_{\text{ground}} = \frac{1}{8} \gamma_1 \times [\text{NO}_2] \times \mathbf{C}_{\text{NO}_2} \times \frac{\mathbf{S}_{\text{g}}}{\mathbf{V}}$$
(4)

135 
$$\mathbf{P}_{\text{aerosol}} = \frac{1}{4} \gamma_2 \times [\text{NO}_2] \times \mathbf{C}_{\text{NO}_2} \times \frac{\mathbf{S}_a}{\mathbf{V}}$$
(5)

136 
$$\frac{S_g}{V} = \frac{1}{MLH}$$
 (6)

137 
$$C_{NO_2} = \sqrt{\frac{8RT}{\pi M}}$$
(7)

138 where  $C_{NO2}$  is the average molecular velocity of NO<sub>2</sub> molecule (m s<sup>-1</sup>); R is the ideal

139 gas constant; T is the temperature (K); M is the molecular weight of NO<sub>2</sub> (kg mol<sup>-1</sup>);

140 MLH is the height of the mixed layer, which is determined to be 50 m due to HONO

141 formation on the ground level and its short lifetime (Liu et al., 2020b);  $S_a/V$  is the

142 surface area to volume ratio of aerosol, estimated by Su et al. (Su et al., 2008).

- 143 4.3.2 Heterogeneous photo-enhanced reactions
- 144 The heterogeneous photo-enhanced reactions of NO<sub>2</sub> on the surface of the ground
- 145  $(P_{ground + hv})$  and the surface of the aerosol  $(P_{aerosol + hv})$  were calculated following (Zhang
- 146 et al., 2020a):

147 
$$\mathbf{P}_{\text{ground+hv}} = \frac{1}{8} \times \mathbf{C}_{\text{NO}_2} \times \frac{1}{\text{MLH}} \times \gamma_1 \times \frac{\mathbf{J}_{\text{NO}_2}}{\mathbf{J}_{\text{NO}_{2,\text{noon}}}} \times [\text{NO}_2]$$
(8)

148 
$$\mathbf{P}_{\text{aerosol+hv}} = \frac{1}{4} \times \mathbf{C}_{\text{NO}_2} \times \frac{\mathbf{S}_a}{\mathbf{V}} \times \gamma_2 \times \frac{\mathbf{J}_{\text{NO}_2}}{\mathbf{J}_{\text{NO}_{2,\text{noon}}}} \times [\text{NO}_2]$$
(9)

149 where JNO<sub>2</sub> and JNO<sub>2, noon</sub> are the photolysis rate of NO<sub>2</sub> and the photolysis rate of NO<sub>2</sub>

- 150 at noon during the day, respectively.
- 151  $y_1$  and  $y_2$  are the absorption coefficient of NO<sub>2</sub> on the ground and aerosol surface,

- 152 respectively, which is assumed to be  $4 \times 10^{-6}$  (Yu et al., 2022; Zhang et al., 2021; Zhang
- 153 et al., 2020a). Moreover, we discuss the uncertainties based on the recommended values
- 154 of  $2 \times 10^{-6}$ -1  $\times 10^{-5}$  as upper and lower bounds(Chen et al., 2023; VandenBoer et al.,
- 155 2013; Wong et al., 2011). Results show (Figure S3) that the uncertainties for P<sub>ground</sub>,
- 156  $P_{aerosol}$ ,  $P_{groung+hv}$ , and  $P_{aerosol+hv}$  are -50% to 150%, -50% to 151%, -20% to 120%, and
- 157 -50% to 121% at the U-ZK, respectively. At the R-PY, the uncertainties for P<sub>ground</sub>,
- 158 P<sub>aerosol</sub>, P<sub>groung+hv</sub>, and P<sub>aerosol+hv</sub> are -50% to 150%, -50% to 151%, -20% to 120%, and
- 159 –50% to 121%, respectively.
- 160 4.4 Nitrate photolysis
- 161 The nitrate photolysis (P<sub>nitrate</sub>) was calculated based on the measured nitrate

162 concentration (NO<sub>3</sub>) from PM<sub>2.5</sub> and nitrate photolysis rate ( $J_{nitrate \rightarrow HONO}$ ):

163  $\mathbf{P}_{\text{nitrate}} = \mathbf{J}_{\text{nitrate} \to \text{HONO}} \times [\text{NO}_3^-]$ (10)

164 where the  $J_{\text{nitrate} \rightarrow \text{HONO}}$  was simulated by normalizing UV values, when the Zenit Angle

- 165 is 0°,  $J_{\text{nitrate}\rightarrow\text{HONO}}$  varied within the range of  $1.22 \times 10^{-5}$  to  $4.84 \times 10^{-4}$  s<sup>-1</sup>, with an average
- 166 value of  $8.24 \times 10^{-5}$  s<sup>-1</sup> (Bao et al., 2018).
- 167

## 168 Text S5 Estimation of HONO formation rate

169 The redox reaction of  $NO_2$  with  $SO_2$  (R<sub>1</sub>) is considered a crucial potential source of

170 high concentrations of HONO in Northern China (Wang et al., 2016b; Cheng, 2016):

171 
$$S(IV) + 2NO_2 + H_2O \rightarrow S(IV) + 2H^+ + 2NO_2^-$$
(R<sub>1</sub>)

172 The rate expression for the reaction was estimated to:

173
$$d[S(VI)]/dt = k_1[NO_2][S(VI)],$$
 (11)174where the  $k_1 = (1.4 \times 10^5 + 1.24 \times 10^7)/2 M^{-1}s^{-1}$  for the pH range < 5;175 $k_1 = (23.25 \times (pH-5) + 1.4 + 124)/2 \times 10^5 M^{-1}s^{-1}$  for the pH range 5 < pH < 5.3;176 $k_1 = (23.25 \times (pH-5) + 1.4 + 12.6 \times (pH-5.3) + 124)/2 \times 10^5 M^{-1}s^{-1}$  for the pH range 5.3 <177pH < 5.8;178 $k_1 = (12.6 \times (pH-5.3) + 124+20)/2 \times 10^5 M^{-1}s^{-1}$  for the pH range 5.8 < pH < 8.7;179and  $k_1 = (2 \times 10^6 + 1.67 \times 10^7)/2 M^{-1}s^{-1}$  for the pH range pH > 8.7. (Seinfeld et al., 1998)180In the above calculation formulas, the concentration of gas in the liquid is determined181by Henry's constant (H\*). The calculation formula is in Table S4. SO<sub>2</sub> has a dissociation182equilibrium in the solution, producing HSO<sub>3</sub><sup>-</sup> and SO<sub>3</sub><sup>2-</sup>. The ionization constants (K)183are shown in the following Table S5. The H\* and K are temperature-dependent. The184values are given in Tables S4 and S5 under the condition of 298K, converted to the185value under the actual temperature using the following calculation formula:

186 
$$H(T) \text{ or } K(T) = H(T_{298K}) \text{ or } K(T_{298K}) \exp\left[-\frac{\Delta H_{298K}}{R}(\frac{1}{T} - \frac{1}{298K})\right]$$
(12)

187 where H(T), K(T),  $H(T_{298K})$ , and  $K(T_{298K})$  represent the H\* and K at actual 188 temperature and 298 K, respectively.

Influences of ionic strength on R<sub>1</sub> were not considered because of the high values
predicted by the ISORROPIA-II model during the sampling periods (Cheng et al.,
2016). To evaluate the effects of mass transport, the formulation of a standard resistance

192 model was adopted:

193 
$$\frac{1}{R_{\rm H,aq}} = \frac{1}{R_{\rm aq}} + \frac{1}{J_{\rm aq,lim}}$$
(13)

where  $R_{H, aq}$  is the sulfate production rate,  $R_{aq}$  is the aqueous-phase reaction rate and J<sub>aq,lim</sub> is the limiting mass transfer rate. which could be calculated by the formulas as follows:

197 
$$J_{aq,lim} = \min\{J_{aq}(SO_2), J_{aq}(X)\}$$
 (14)

198 
$$J_{aq}(X) = k_{MT}(X) \cdot [X]$$
 (15)

where [X] refers to the aqueous-phase concentrations of SO<sub>2</sub> or the oxidants O<sub>xi</sub> calculated by the equation in Table S4. The mass transfer rate coefficient  $k_{MT}(X)$  (s<sup>-1</sup>) can be calculated by:

202 
$$k_{_{\rm MT}} = \left[\frac{R_{\rm p}^2}{3D_{\rm g}} + \frac{4R_{\rm p}}{3\alpha\nu}\right]^{-1}$$
(16)

where  $R_p$  is the aerosol radius,  $D_g$  is the gas-phase molecular diffusion coefficient (0.2 cm<sup>2</sup> s<sup>-1</sup> at 293K), *v* is the mean molecular speed of X (3×10<sup>4</sup> cm s<sup>-1</sup>), and *a* is the mass accommodation of X on the droplet surface, and we adopted values of 0.11 and 2E<sup>-4</sup> for SO<sub>2</sub> and NO<sub>2</sub>, respectively referring to Cheng et al. (Cheng, 2016).





Figure S1. Sampling point map in Henan Province, China. © 2019 National Geomatics
Center of China. i.e., urban sites at Sanmenxia (U-SMX), Zhoukou (U-ZK), Zhuamdian
(U-ZMD) and Xinyang (U-XY), rural sites at Anyang (R-AY), Xinxiang (R-XX),
Puyang (R-PY), Jiaozuo(R-JZ), Shangqiu (R-SQ) and Nanyang (R-NY). All rights
reserved.



216 Figure S2. Result of conditional bivariate probability function plots: NO<sub>2</sub> at U-ZK and

- 217 R-PY sites before (PC) and during (DC) the COVID-19 outbreak. The color scale bar
- 218 represents NO<sub>2</sub> concentration.





221 Figure S3. HONO production rate using different uptake rates of NO<sub>2</sub> at the U-ZK

- and R-PY sites before (PC) and during (DC) the COVID-19 outbreak. (a)P<sub>ground</sub>, (b)
- 223 Paerosol, (c) Pground+hv, and (d) Paerosol+hv
- 224





Figure S4. Daily changes in temperature and relative humidity (RH) in rural sites before (PC) and during (DC) the COVID-19 outbreak, the error bar represents the standard deviation. The upper and lower whiskers represent the standard deviation.



232 Figure S5. The equilibrium state of anions and cations at ten sites before (PC) and

233 during (DC) the COVID-19 outbreak.



Figure S6. Concentrations of the water-soluble ions at the ten sites before (PC) and

236 during (DC) the COVID-19 outbreak.



242 TH<sub>2</sub>SO<sub>4</sub>, TNO<sub>3</sub>, TCl, TNa, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, T, and RH) was simulated to compare its

243 effects on pH.









(b)





Figure S8. Relationship between HONO and main influencing factors during (DC) the COVID-19 outbreak at U-ZK and R-PY sites. In each box, the top, middle, and bottom lines represent the 75, 50, and 25 percentiles of statistical data, respectively; the upper and lower whiskers represent the 90 and 10 percentiles of statistical data, respectively.





Figure S10. Sensitivity of HONO product rate to each factor. The vertical bar represents the mean values before (PC) and during (DC) the COVID-19 outbreak. The real-time measured values of a variable and the average values of other parameters were input into the production rate of the  $R_1$  reaction.



264

265 Figure S11. pH and  $R_1$  uncertainties at the U-ZK and R-PY sites are based on two

266 extreme scenarios of sensitivity to measurement uncertainty.

268 Tables

Table S1. Descriptions of the ten sampling sites in Henan Province, China.

Observation	Classifications	Abbreviations	Coordinates	Locations	Surrounding
sites					environment
Sanmenxia	Urban site	U-SMX	34.79 °N, 111.16 °E	Sanmenxia Environmental Protection Bureau	Roads, residential
					areas
Zhoukou	Urban site	U-ZK	33.65° N, 114.65° E	Chuanhui District People's Government	Roads, residential
					areas
Zhumadian	Urban site	U-ZMD	33.01° N, 114.01° E	Huanghuai College	Roads, residential
					areas, shopping malls
Xinyang	Urban site	U-XY	32.14° N, 114.09° E	Xinyang Museum	Roads, residential
					areas, shopping malls
Anyang	Rural site	R-AY	36.22°N, 114.39° E	Baizhuang Town Xindian North Street China	Highways, villages,
				Resources Gas (Andan Station)	farmland
Xinxiang	Rural site	R-XX	35.38° N, 114.30° E	Banzao Township Central School in Yanjin	Villages, farmland
				County	
Puyang	Rural site	R-PY	36.15° N, 115.10° E	Nanle County Longwang Temple Station	Villages, farmland
Jiaozuo	Rural site	R-JZ	35.02° N, 113.35° E	The Second River Bureau of Jiefeng Village,	Villages, farmland
				Beiguo Township, Wuxi County	
Shangqiu	Rural site	R-SQ	34.56° N, 115.61° E	Liangyuan Huanghe Gudao National Forest	Highways, villages,
				Park	farmland
Nanyang	Rural site	R-NY	32.68° N, 111.70° E	Nanyang Tangshan Park	Villages, farmland

274		Table S2. T	The value of $\rho_s$ in other	r studies.
	Observation site	<b>Period</b>	$\rho_s(g/cm^3)$	Reference
-	Beijing	Dec 2010	<mark>5 1.4</mark>	(Liu et al., 2017)
	<b>Tianjin</b>	Dec-Jun 20	015 1.3	<mark>(Shi et al., 2017)</mark>
	<mark>Xi'an</mark>	Nov-Dec 20	012 1.4	<mark>(Guo et al., 2017)</mark>
	Hohhot	Winter 20	15 1.35	(Wang et al., 2019)
	Northeastern USA	Feb-Mar 20	)15 <u>1.4</u>	<mark>(Guo et al., 2016)</mark>
	Crete, Greece	Aug-Nov 2	012 1.35	(Bougiatioti et al., 2016)
	<mark>Alabama, USA</mark>	<mark>Jun-Jul 20</mark>	<mark>13</mark> 1.4	<mark>(Guo et al., 2015)</mark>
-	Georgia, USA	Aug-Oct 20	<mark>)16</mark> 1.4	(Nah et al., 2018)
275				
276	<mark>Ta</mark>	ible S3. Summa	ary of vehicle emissio	n factors.
	Observation site	Period	Emission factor (%)	Reference
	Beijing	<mark>2020</mark>	<mark>0.79</mark>	(Meng et al., 2020)
	Hong Kong	<mark>2015</mark>	<mark>0.4–1.8</mark>	<mark>(Yun et al., 2017)</mark>
	Hong Kong	<mark>2011</mark>	<mark>0.5–1.6</mark>	(Xu et al., 2015)
	Kiesberg Tunnel	<mark>2001</mark>	<mark>0.8</mark>	(Kleffmann et al., 2003)
	Kiesberg Tunnel	<mark>1997</mark>	<mark>0.3–0.8</mark>	(Kurtenbach and Wiesen, 2001)
	<b>Guangzhou</b>	<mark>2019</mark>	<mark>1.31</mark>	<mark>(Li et al., 2021b)</mark>
277				
278	Table S4. C	onstants for cal	culating the apparent	Henry's constant (H*).
	Equilibriu	ım	H (M atm <sup>-1</sup> ) at 298K	- \(\Delta H_{298K} / R (K) \)
	$SO_2(g) \leftrightarrow$	· SO <sub>2</sub> (aq)	1.23	3145.3
	$NO_2(g) \leftarrow$	NO <sub>2</sub> (aq)	1.00E-02	2516.2
279				
280	Table St	5. Constants for	calculating the ionization	ation constants (K).
	Equilibriu	Im	K (M) at 298K	- \(\Delta H_{298K} / R (K) \)
	SO <sub>2</sub> ·H <sub>2</sub> O	$\leftrightarrow \mathrm{H^{+} + HSO_{3}^{-}}$	1.30E-02	1960
	$\mathrm{HSO}_{3}^{-} \leftrightarrow$	$H^{+} + SO_{3}^{2-}$	6.60E-08	1500

Sampling sites	Seasons	Years	$NH_3$ (ug/m <sup>3</sup> )	Sites	References
Delhi, India	Winter	2013– 2015	$(\mu g) = 19.2 \pm 3.5$	Urban	(Saraswati et al., 2019)
Osaka, Japan	Winter	2015	$1.5\pm0.7$	Urban	(Huy et al., 2017)
Toronto, Canada	Winter	2007	$0.8\pm0.5$	Urban	(Hu et al., 2014)
Kanpur, India	Winter	2007	$21.7\pm5.8$	Urban	(Behera and Sharma, 2010)
Nanjing	Winter	2014	6.7	Urban	(Wang et al., 2016b)
Yangtze River Delta	Winter	2019	$9.3\pm4.0$	Urban	(Wang et al., 2021)
Shanghai	Winter	2014	$2.8 \pm 1.0$	Urban	(Wang et al., 2018b)
Tianjin Xi'an	Winter Winter	2015 2012	$\begin{array}{c} 12.0\\ 17.5\pm9.1 \end{array}$	Urban Urban	(Shi et al., 2019) (Wang et al., 2016a)
Fujian	Winter	2016	$12.8\pm4.8$	Urban	(Wu et al., 2017)
Beijing	Winter	2015	$15.1\pm2.9$	Urban	(Wang et al., 2016a)
Beijing	Winter	2017	$13.1\pm1.6$	Urban	(Zhang et al., 2020b)
Beijing	Winter	2020	$19.9\pm3.8$	Urban	(Zhang et al., 2020b)
Taoyuan	Winter	2017-	$1.7\pm1.9$	Urban	(Duan et al., 2021)
Zhengzhou	Winter	2018	$19.0\pm4.0$	Rural	(Wang et al., 2020)
Quzhou	Winter	2019	$29.5\pm2.2$	Rural	(Feng et al., 2022)
Gucheng	Winter	2016	9.3	Rural	(Xu et al., 2019)
Chongming	Winter	2019–	$9.3\pm4.0$	Rural	(Lv et al., 2022)
Shanglan	Winter	2020 2017– 2018	$2.5\pm2.6$	Rural	(Duan et al., 2021)

282Table S6. Comparisons of NH3 concentrations (mean  $\pm$  standard deviation) ( $\mu$ g/m³)283from studies in other cities.

Table S7. The concentration (mean  $\pm$  standard deviation) of relative humidity (RH),

287	temperature (T), $\epsilon$ (NH <sub>4</sub> <sup>+</sup> ) at the ten sites before (PC) and during (DC) the COVID-19
288	outbreak.

Sites	Substances	Total	PC	DC
U-SMX	RH (%)	$54.8 \pm 18.0$	$60.6\pm16.5$	$51.2\pm18.0$
	T (°C)	$5.6\pm4.2$	$3.1 \pm 2.1$	$7.0 \pm 4.4$
	$\epsilon(\mathrm{NH_4^+})$	$0.43\pm0.20$	$0.54\pm0.18$	$0.36\pm0.18$
U-ZK	RH (%)	$70.1\pm21.9$	$73.6\pm14.5$	$69.4\pm22.4$
	T (°C)	$6.4\pm4.3$	$3.8\pm2.3$	$7.0 \pm 4.5$
	$\epsilon(\mathrm{NH_4}^+)$	$0.43\pm0.20$	$0.59\pm0.14$	$0.32\pm0.17$
U-ZMD	RH (%)	$74.9\pm23.3$	$84.4 \pm 17.8$	$68.9\pm24.4$
	T (°C)	$5.6\pm4.6$	$2.9\pm2.7$	$7.4\pm4.8$
	$\epsilon(\mathrm{NH_4^+})$	$0.48\pm0.21$	$0.62\pm0.17$	$0.38\pm0.18$
U-XY	RH (%)	$77.0\pm22.1$	$86.7\pm13.3$	$74.3\pm23.3$
	T (°C)	$7.7\pm4.5$	$4.7\pm2.2$	$8.5\pm4.6$
	$\epsilon(\mathrm{NH_4^+})$	$0.55\pm0.21$	$0.71\pm0.14$	$0.45\pm0.18$
R-AY	RH (%)	$62.2\pm17.9$	$70.1\pm14.9$	$57.2\pm17.8$
	T (°C)	$2.6\pm0.9$	$-0.2 \pm 2.5$	$4.4\pm4.7$
	$\epsilon(\mathrm{NH_4^+})$	$0.46\pm0.17$	$0.57\pm0.15$	$0.39\pm0.14$
R-XX	RH (%)	$63.0\pm17.0$	$68.7 \pm 14.6$	$59.5\pm17.5$
	T (°C)	$2.9\pm4.6$	$0.3\pm2.8$	$4.4\pm4.8$
	$\epsilon(\mathrm{NH_4^+})$	$0.40\pm0.17$	$0.52\pm0.16$	$0.35\pm0.14$
R-PY	RH (%)	$63.6\pm18.0$	$71.5\pm14.6$	$57.6\pm18.0$
	T (°C)	$1.7\pm4.8$	$-0.8 \pm 3.2$	$3.6\pm4.9$
	$\epsilon(\mathrm{NH_4^+})$	$0.43\pm0.17$	$0.58\pm0.13$	$0.34\pm0.13$
R-JZ	RH (%)	$56.3 \pm 18.5$	$62.0\pm16.7$	$52.8 \pm 18.7$
	T (°C)	$4.1\pm4.4$	$1.7 \pm 2.6$	$5.6 \pm 4.7$
	$\epsilon(\mathrm{NH_4^+})$	$0.37\pm0.14$	$0.46\pm0.13$	$0.32\pm0.13$
R-SQ	RH (%)	$63.2\pm15.6$	$67.5\pm12.6$	$60.5\pm17.0$
	T (°C)	$4.2\pm4.5$	$2.0\pm2.9$	$5.6 \pm 4.7$
	$\epsilon(\mathrm{NH_4^+})$	$0.45\pm0.19$	$0.63\pm0.12$	$0.35\pm0.14$
R-NY	RH (%)	$75.9 \pm 19.1$	$79.3\pm17.7$	$73.9\pm19.6$
	T (°C)	$5.7\pm3.8$	$3.6\pm2.6$	$6.9\pm3.9$
	$\epsilon(\mathrm{NH_4^+})$	$0.59\pm0.19$	$0.73\pm0.12$	$0.52\pm0.18$

290 Table S8. The concentration (mean  $\pm$  standard deviation) of required ammonia

291	Required-NH <sub>x</sub>	) and excess ammonia	(Excess-NH <sub>x</sub> )	) at the ten	sites before	(PC)	) and
			\	/		<b>`</b>	/

Sites	Substances	Total (µg/m <sup>3</sup> )	PC (µg/m <sup>3</sup> )	DC ( $\mu g/m^3$ )
U-SMX	Required-NH4 <sup>+</sup>	$9.1 \pm 7.1$	$12.7 \pm 7.1$	$7.0 \pm 6.2$
	Excess-NH <sub>4</sub> <sup>+</sup>	$14.7\pm11.2$	$13.6\pm10.4$	$15.3 \pm 11.6$
U-ZK	Required-NH4 <sup>+</sup>	$15.2\pm9.6$	$21.4\pm8.6$	$11.6\pm8.4$
	Excess-NH <sub>4</sub> <sup>+</sup>	$14.6\pm8.3$	$11.9\pm6.0$	$16.1\pm9.0$
U-ZMD	Required-NH4 <sup>+</sup>	$13.9\pm9.8$	$19.4\pm9.8$	$10.4\pm8.0$
	$Excess-NH_4^+$	$12.8\pm8.7$	$11.6\pm8.2$	$13.6\pm8.8$
U-XY	Required-NH4 <sup>+</sup>	$10.2\pm7.5$	$14.6\pm7.3$	$7.4 \pm 6.2$
	Excess-NH4 <sup>+</sup>	$7.8\pm4.6$	$6.5\pm4.4$	$8.7\pm4.5$
R-AY	Required-NH4 <sup>+</sup>	$17.1 \pm 12.4$	$23.9\pm13.4$	$12.8\pm9.5$
	Excess-NH <sub>4</sub> <sup>+</sup>	$21.2\pm9.4$	$20.2\pm9.2$	$21.9\pm9.4$
R-XX	Required-NH4 <sup>+</sup>	$13.5\pm9.6$	$18.0\pm9.8$	$10.7\pm8.2$
	Excess-NH <sub>4</sub> <sup>+</sup>	$23.3 \pm 11.4$	$19.6\pm10.8$	$25.6 \pm 11.2$
R-PY	Required-NH4 <sup>+</sup>	$13.8 \pm 11.0$	$22.1\pm12.5$	$9.3\pm6.6$
	Excess-NH <sub>4</sub> <sup>+</sup>	$22.3\pm10.8$	$17.5\pm8.6$	$25.0\pm11.0$
R-JZ	Required-NH4 <sup>+</sup>	$15.4\pm10.4$	$20.3\pm10.6$	$12.5 \pm 9.1$
	$Excess-NH_4^+$	$27.5\pm12.9$	$26.0 \pm 13.1$	$28.4 \pm 12.7$
R-SQ	Required-NH4 <sup>+</sup>	$13.2\pm9.1$	$19.1\pm8.9$	$9.9\pm7.3$
	$Excess-NH_4^+$	$15.1 \pm 8.6$	$10.1\pm5.4$	$17.9\pm8.7$
R-NY	Required-NH4 <sup>+</sup>	$9.9\pm6.6$	$13.0\pm6.9$	$8.1\pm5.8$
	Excess-NH <sub>4</sub> <sup>+</sup>	$6.0\pm3.6$	$4.4\pm3.3$	$6.9 \pm 3.4$

292 during (DC) the COVID-19 outbreak.

294 Table S9. Comparison of the particle pH values in this study (PC/DC) and other sites

		<del>u).</del>		
	<mark>Sites</mark>	Periods	pH	References
<mark>Urban</mark>	Sanmenxia	Jan–Feb 2020	$\frac{4.6 \pm 0.5 / 4.8 \pm 0.9}{100}$	This study
	<mark>Zhoukou</mark>	Jan–Feb 2020	$4.6 \pm 0.6 / 5.1 \pm 0.4$	
	Zhumadian	Jan–Feb 2020	$4.6 \pm 0.3/4.8 \pm 1.2$	
	<mark>Xinyang</mark>	Jan–Feb 2020	$4.2 \pm 0.3/4.6 \pm 1.3$	
Rural	Anyang	Jan–Feb 2020	$4.5 \pm 0.4 / 4.6 \pm 0.8$	
	<mark>Xinxiang</mark>	Jan–Feb 2020	$4.8 \pm 0.5 / 4.9 \pm 0.9$	
	<mark>Puyang</mark>	Jan–Feb 2020	$4.8 \pm 0.3 / 5.1 \pm 0.9$	
	<mark>Jiaozuo</mark>	Jan–Feb 2020	$4.9 \pm 0.5/5.1 \pm 0.8$	
	<mark>Shangqiu</mark>	<mark>Jan–Feb 2020</mark>	$4.5 \pm 0.3/4.7 \pm 0.8$	
	<b>Nanyang</b>	Jan–Feb 2020	$4.2 \pm 0.5/4.4 \pm 0.7$	
<mark>Urban</mark>	Beijing	Jan–Feb 2015	<mark>4.5</mark>	<mark>(Guo et al., 2017)</mark>
	<b>Beijing</b>	Dec 2016	$4.3 \pm 0.4$	(Liu et al., 2017)
	<b>Beijing</b>	Feb 2017	$4.5 \pm 0.7$	(Ding et al., 2019)
	<mark>Tianjin</mark>	Dec–Jun 2015	<mark>4.9 ± 1.4</mark>	<mark>(Shi et al., 2017)</mark>
	Tianjin	Aug 2015	$3.4 \pm 0.5$	<mark>(Shi et al., 2019)</mark>
	Hohhot	Winter	<mark>5.7</mark>	(Wang et al., 2019)
	Mt. Tai	Summer	$2.9 \pm 0.5$	<mark>(Liu et al., 2021b)</mark>
	Taoyuan	<mark>Nov 2017–Jan 2018</mark>	$\frac{5.1 \pm 1.0}{1}$	<mark>(Duan et al., 2021)</mark>
	Zhengzhou	<mark>Jan 2018</mark>	<mark>4.5</mark>	(Wang et al., 2020)
	Anyang	Jan 2018	4.8	(Wang et al., 2020)
Mountain	Mt. Tai	Summer	$3.6 \pm 0.7$	(Liu et al., 2021b)
Rural	<b>Shanglan</b>	Nov 2017–Jan 2018	$5.5 \pm 1.1$	(Duan et al., 20 <mark>21)</mark>

295 (mean or mean  $\pm$  standard).

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