



1 Response of Marine Primary Producers to Olivine 2 Additions

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8 **Abstract.** Carbon dioxide removal (CDR) technologies are gaining increasing attention as a
9 potential solution to reduce atmospheric CO₂ concentrations and combat climate change. Ocean
10 alkalinity enhancement (OAE) seeks to enhance the ocean's CO₂ absorption capacity by
11 introducing powdered minerals or dissolved alkaline substances into the surface ocean.
12 Nevertheless, the impact of OAE on marine ecosystems remains largely uncharted. In this study,
13 we explored the impact of olivine on a diversity of cosmopolitan primary producers, including
14 Coccolithophores, Diatoms, Dinophyceae, *Micromonas* sp., *Prochlorococcus* sp., and
15 *Synechococcus* sp. Here, we show that most primary producers were not impacted negatively by
16 the concentrations of olivine additions that were applied despite olivine additions increasing nickel
17 concentrations in our cultures. Additions of olivine did not lead to growth inhibition but resulted
18 in a slight increase in growth in most cultures, with picoplankton benefiting the most. However,
19 phytoplankton responses were species-specific and subject to the media used in a combination
20 with the olivine addition. Additionally, it is essential to mention the pitfalls and concerns
21 associated with our experimental setup, particularly regarding the impact of the medias and
22 considerations of carbonate chemistry. Our findings raise confidence in applying olivine for
23 carbonation to generate CO₂ removal without harming primary producers; however, future studies
24 should include tests on an ecosystem level to investigate potential effects on different trophic levels
25 and natural settings.

26 1 Introduction

27 Ocean alkalinity enhancement (OAE), a carbon dioxide removal (CDR) technology, is gaining
28 substantial attention as a strategy to reduce atmospheric CO₂ concentrations and thus mitigate
29 climate change (Fawzy et al., 2020; IPCC, 2023; Pörtner et al., 2023). OAE aims to increase the
30 alkalinity of surface ocean waters and sediments of entire regions, enhancing the ocean's capacity



31 to absorb more CO₂ from the atmosphere (Campbell et al., 2022; Bach et al., 2019), while at the
32 same time mitigating ocean acidification (Kheshgi, 1995; Doney et al., 2009; Egleston et al., 2010;
33 Heinze et al., 2015; Renforth and Henderson, 2017; Bach et al., 2019).

34 Mineral-based OAE utilizes the addition of ground alkaline minerals to ocean waters to increase
35 seawater alkalinity (Renforth and Henderson, 2017). OAE approaches are based on the natural
36 process of rock weathering as a key process of stabilizing Earth's temperature (Urey, 1957; Walker
37 et al., 1981; Berner et al., 1983; Volk, 1987) and are closely linked to another CDR technology
38 known as enhanced weathering (EW) (Köhler et al., 2010; Goll et al., 2021; Calabrese et al., 2022).
39 EW processes enhance weathering rates by, for instance, spreading ground silicate rocks or basaltic
40 rocks on terrestrial land, leveraging natural weathering mechanisms (Renforth et al., 2015;
41 Meysman and Montserrat, 2017; Beerling et al., 2020). The permeance of CO₂ removal achieved
42 through OAE is estimated to last more than ten thousand years, depending on the duration of
43 HCO₃⁻ in solution and the absence of reverse reactions due to a changed equilibrium (Archer, 2005;
44 Rau, 2011; Middelburg et al., 2020) Notably, reverse reactions of the carbonate-silicate weathering
45 are reduced in ocean regions with low salinity, faster mixing, and dilution (Archer, 2005; Kirchner
46 et al., 2020; Bertagni and Porporato, 2022).

47 Rising concentrations of CO₂ in the atmosphere have profound impacts on the marine environment
48 (IPCC, 2023; Smith et al., 2023), leading to, for example, heat waves and ocean acidification
49 (Doney et al., 2009; Oliver et al., 2018). These changes cause significant alterations to marine
50 biodiversity, including changes in primary producer community composition and activity (Doney
51 et al., 2012). A change in the primary producer community in response to alkaline substances, such
52 as olivine and quicklime, has been suggested (Bach et al., 2019), with certain taxonomic groups
53 being favored over others. The addition of olivine was, for instance, suggested to promote a
54 diatom- and cyanobacteria-rich environment, producing a "Greener" ocean, as these groups have
55 high requirements for one or more elements present in olivine, including silicon (Si), iron (Fe),
56 and nickel (Ni). On the other hand, the addition of quicklime was suggested to lead to a
57 coccolithophore-rich environment, generating a "Whiter" ocean (Bach et al., 2019).

58 Primary producers comprise the first trophic level of the marine food web and play a crucial role
59 as the foundation of global ecosystems (Michaels and Silver, 1988; Longhurst et al., 1995;
60 Buesseler, 1998). Marine phytoplankton is responsible for approximately half of Earth's carbon



61 fixation (Longhurst et al., 1995; Field et al., 1998) and a driver of the biological carbon pump
62 exporting CO₂ to waters below the mixed layer, thus removing CO₂ on longer timescales (Volk
63 and Hoffert, 1985). The transfer of organic matter by the biological carbon pump creates a vertical
64 dissolved inorganic carbon (DIC) gradient, enhancing the ocean's ability to absorb CO₂ from the
65 atmosphere through air-sea exchange to replace the DIC that was converted into organic matter
66 (Kwon et al., 2009). Approximately 20% of carbon fixed during primary production in the euphotic
67 zone is transported into the deep ocean through various mechanisms (Dunne et al., 2007; Devries
68 and Weber, 2017; Nowicki et al., 2022).

69 The primary producers we have chosen to work with are cosmopolitan and quantitatively important
70 in the marine environment (Waterbury et al., 1979; Chisholm et al., 1988; Quere et al., 2005;
71 Worden et al., 2015). Prochlorococcus and Synechococcus are cyanobacteria, a diverse phylum of
72 photoautotrophic prokaryotes (Chisholm et al., 1988; Chisholm et al., 1992). The eukaryotic
73 primary producers in this study included coccolithophores (*Emiliania huxleyi*), diatoms
74 (*Thalassiosira weissflogii*, *Skeletonema marinoi*), dinophyceae (*Scrippsiella trochoidea*) and a
75 green alga (*Micromonas commoda*). Coccolithophores are calcifying organisms with exoskeletons
76 made of calcium carbonate called coccoliths, making them particularly vulnerable to ocean
77 acidification (Iglesias-Rodriguez et al., 2008; Meyer and Riebesell, 2015). They are ecologically
78 and biogeochemically essential and play a vital role in the marine biological carbon pump (Poulton
79 et al., 2013; De Vries et al., 2021). Diatoms are responsible for about one-fifth of Earth's
80 photosynthesis (Treguer et al., 1995). They are present in all marine environments and are
81 responsible for generating as much organic carbon as all the global rainforests combined, which
82 illustrates their importance for the global carbon cycle (Treguer et al., 1995; Field et al., 1998). In
83 the ocean, diatoms account for approximately 40 % of the total primary production (Treguer et
84 al., 1995; Tréguer and Pondaven, 2000; Sarthou et al., 2005). Diatoms require silicate to construct
85 their siliceous cell walls. Therefore, elevated silicate concentrations in the ocean could promote
86 the proliferation of diatoms in comparison to non-siliceous types of phytoplankton, an aspect
87 potentially relevant in the context of OAE-mineral additions (Tréguer and Pondaven, 2000).
88 Dinophyceae, coccolithophores, and diatoms initiated a major shift in the global carbon cycle,
89 which began an era of increasing atmospheric O₂ concentrations and declining atmospheric CO₂
90 concentrations (Katz et al., 2005). Dinophyceae are motile phytoplankton, allowing them to
91 migrate in the water column (Margalef, 1978). It is a diverse group of organisms in which some



92 representatives can produce toxins, bioluminesce, or even be parasitic or symbiotic (Hackett et al.,
93 2004). Lastly, *Micromonas* sp. is a unicellular picoplanktonic ($\leq 2 \mu\text{m}$ cell diameter) marine green
94 alga which thrives in most ocean environments, from coastal to open waters (Not et al., 2004;
95 Worden et al., 2009; Cuvelier et al., 2017). Several studies over the past decades have observed
96 that there has been an increased abundance of picophytoplankton species, including *Micromonas*
97 spp. In comparison, larger phytoplankton have decreased in abundance. These changes are strongly
98 linked to climate change, which could lead to reduced biological production at higher trophic levels
99 and points towards the importance of *Micromonas* spp. in a future ocean (Li et al., 2009; Mckie-
100 Krisberg and Sanders, 2014; Van Baren et al., 2016).

101 In this study, we explored the impact of olivine additions (grain sizes $< 63 \mu\text{m}$), on ten strains of
102 primary producers, including coccolithophores, diatoms, dinophyceae, the green alga *Micromonas*
103 *commoda*, as well as on cyanobacteria (*Prochlorococcus*, and *Synechococcus*). We monitored the
104 response of primary producers using their overall growth and snapshots of carbon fixation rates in
105 response to olivine additions. Those biological parameters were complemented with quantification
106 of trace metal and ion concentrations, including Si, Ni, magnesium (Mg), manganese (Mn),
107 calcium (Ca), and Fe, to explore potential ecotoxicological or growth-stimulating effects on
108 primary producer cultures. In addition, we assessed how olivine additions affected the carbonate
109 system during the experimental periods. Taken together, we hope that our study can inform and
110 contribute to the design of experiments in relevant ecosystems, model explorations of OAE
111 applicability, and decision-making on the applicability of olivine additions for OAE as a tool to
112 mitigate climate change.

113 **2 Experimental methods**

114 **2.1 Mineral analysis**

115 Olivine used for our addition experiments were obtained from Norway. To increase the surface
116 area of the raw minerals, olivine minerals was initially crushed in a Laboratory Jaw Crusher (Type
117 BB51, Retsch, Haan, Germany) and then further crushed using a Planetary Ball Mill PM 100
118 (Retsch, Haan, Germany) for 5-10 minutes at 500 Relative Centrifugal Force (RCF). The crushed
119 materials were sieved using a $63 \mu\text{m}$ pore size sieve to select particles with diameters at or below



120 <63 μm (Horowitz, 1985; Kersten and Smedes, 2002). The olivine was then ground with ball mills,
121 requiring less energy than other ultra-fine grinding methods (Summers et al., 2005).

122 The specific surface area of the particles was determined using a Brunauer–Emmett–Teller (BET)
123 approach. Samples were analyzed in duplicates (Tab. 1); a Micromeritics Gemini 2390
124 (Micromeritics, Norcross, USA) was used to determine further the particle sizes (Particle
125 Analytical ApS, Hørsholm, Denmark). The results for olivine showed the following average
126 surface areas (BET analysis) $1.81\text{ m}^2\text{ g}^{-1}$ (Tab. 1).

127 Major and trace elements of the applied olivine were analyzed with inductively coupled plasma
128 optical emission spectroscopy (ICP-OES, Agilent 7700 ICP-MS, Agilent Technologies, Santa
129 Clara, USA). Olivine (forsterite) is an ultramafic rock primarily composed of Mg_2SiO_4 (Bragg and
130 Brown, 1926). The major element composition analysis displayed in Tab. 1 confirms the
131 composition of olivine. Olivine is primarily composed of Mg (29.65 %) and Si (18.97 %).

132 Table 1. Specific surface area (BET, $\text{m}^2\text{ g}^{-1}$) and major element composition of olivine (%).

Minerals	BET, $\text{m}^2\text{ g}^{-1}$	Major elements in %						
		Na	Mg	Al	Si	K	Ca	Fe
Olivine	1.81	0.006	29.654	0.079	18.966	0.005	0.085	5.191

133 Olivine’s trace metal content, a concentration of Mn ($728.286\text{ mg kg}^{-1}$), and a concentration of
134 zinc (Zn) (29.1 mg kg^{-1}) was detected in olivine. Both trace metals are crucial for primary
135 producers' growth (Blaby-Haas and Merchant, 2017) (Tab. 2). Ni, arsenic (As), and chromium (Cr)
136 can negatively affect the growth of primary producers (Cervantes et al., 2001; Tripathi and Poluri,
137 2021; Guo et al., 2022). Olivine contained a concentration of As (0.066 mg kg^{-1}), and of Cr (453.3
138 mg kg^{-1}), and Ni (2709.2 mg kg^{-1}), all of which could impair ecosystem health. It is essential to
139 consider that Mn, Zn, As, Cr, and Ni toxicity can vary depending on the specific aquatic setting.
140 Guidelines on acceptable trace metal concentrations might, therefore, vary.

141 Table 2. Trace element concentrations of olivine in mg kg^{-1} applied in the experiments.

Minerals	Trace elements in mg kg^{-1}							
	Cr	Mn	Co	Ni	Cu	Zn	As	Pb
Olivine	453.294	728.286	106.306	2709.208	3.308	29.133	0.066	0.646

142

143 **2.2 Strains and cultivation conditions**



144 The primary producers used in this study included cyanobacteria, dinophyceae, chlorophyta,
145 diatoms, and coccolithophores. We explored how they responded to abrupt exposure to olivine
146 additions, thus imitating a water column mineral addition. The following strains were all related
147 to the open oceans, specifically the Atlantic and Pacific oceans. The primary producers were pure
148 cultures purchased from Roscoff Culture Collection (CNRS-Sorbonne Université, Station
149 Biologique, Place G. Tessier 29680 Roscoff, France) or Bigelow, National Center for Marine
150 Algae and Microbiota (Bigelow, East Boothbay, ME, USA).

151 Table 3. Overview of the phytoplankton strains, cultivation media, geographic origin, experimental period [d], and if the
152 phytoplankton cultures were grown on a shaker during the experimental period.

Species	Strain Name	Cultivation Media	Ocean Region	Runtime of experiment [d]	Shaking
<i>Synechococcus sp.</i>	A15-62 Clonal	PCR S11 Red Sea Medium (Roscoff Culture Collection, 2019) adopted from (Rippka et al., 2000)	North Atlantic, Tropical Atlantic	25	Shaking
<i>Synechococcus sp.</i>	EUM Syn10	PCR S11 Red Sea Medium (Roscoff Culture Collection, 2019) adopted from (Rippka et al., 2000)	North Atlantic, Tropical Atlantic	10	No Shaking
<i>Synechococcus sp.</i>	WH8102	PCR S11 Red Sea Medium (Roscoff Culture Collection, 2019) adopted from (Rippka et al., 2000)	North Atlantic, Caribbean Sea	9	Shaking
<i>Prochlorococcus sp.</i>	NATL2-M98	PCR S11 Red Sea Medium (Roscoff Culture Collection, 2019) adopted from (Rippka et al., 2000)	North Atlantic	34 (41) *	Shaking
<i>Prochlorococcus sp.</i>	EQPAC1-C	PCR S11 Red Sea Medium (Roscoff Culture Collection, 2019) adopted from (Rippka et al., 2000)	Pacific Ocean, Equatorial Pacific	9	Shaking
<i>Scrippsiella trochoidea</i>	CCMP1331	L1 medium (Biglow, 2020, DOC-051.000) adapted from (Guillard and Hargraves, 1993), silicate excluded	Unknown	44	No Shaking
<i>Micromonas commoda</i>	CCMP489	K medium (Biglow, 2020, DOC-050.000) adapted from (Keller, 1985; Keller et al., 1987)	North Atlantic, Sargasso Sea	9	No Shaking
<i>Thalassiosira weissflogii**</i>	CCMP1051	L1 medium (Biglow, 2020, DOC-051.000) adapted from (Guillard and Hargraves, 1993), silicate included	North Pacific	42	Shaking
<i>Skeletonema marinoi**</i>	CCMP2092	L1 medium (Biglow, 2020, DOC-051.000) adapted from (Guillard and Hargraves, 1993), silicate included	North Atlantic, Adriatic Sea	42	Shaking
<i>Emiliania huxleyi***</i>	AC448	Adapted K medium, (Biglow, 2020, DOC-050.000) adapted from (Keller, 1985; Keller et al., 1987)****	North Atlantic	32	Shaking

153 *Dolomite addition of NATL2-M98 was incubated for an extended 41 days.

154 **Diatom

155 ***Coccolithophore



156 **** Adaptions of K medium involving division of the additions of components by two.

157 The seawater for L1, K, and their adaptations have been taken from the Danish coastal North Sea
158 (55°17'49.4" N 8°39'05.2" E) at the beginning of high tide on June 23, 2020, and July 6, 2021.
159 Seawater salinity was between 30.9 – 31.0. The seawater was prefiltered with a bottle top 0.2 µm
160 PES membrane filter (Avantor VWR® Radnor, Pa, USA). The primary producers were cultivated
161 in biological triplicates. Salinity varied between the media. PCR S11 had a salinity of 36.0, while
162 the others were mixed with North Sea water and had a salinity between 31.4 and 31.5.

163

164

165 **2.3 Culture experiments**

166 To ensure consistent and comparable results in cultivating the 10 primary producers affected by
167 the olivine additions, we kept the primary producers in identical conditions in a cultivation room
168 with a day/night cycle of 8 h of light and 16 h of darkness. The laboratory tables were consistently
169 exposed to an average light intensity of 42.47 lux (lumen/m²) measured with LI-250A Light Meter
170 (LICOR®, Lincoln, USA) in the light period and maintained a controlled room temperature of 20.0
171 C°- 20.3° C throughout the experiments. All primary producers were cultivated in 1 L or 2 L
172 conical flasks, with headspaces in exchange with the atmosphere. 400 mL of the algal-specific
173 medium was mixed with 100 mL of the desired, pre-grown strain prior to the experimental start.
174 Samples were continually collected at the same time, between 9:30 a.m. and 11:00 a.m. The
175 primary producers, namely CCMP489, EUM Syn10, and CCMP1331, were not subjected to
176 shaking during the cultivation experiments. CCMP489 and EUM Syn10 were cultivated in a
177 growth chamber Model KBW 400 (Binder, Tuttlingen, Germany), where shakers were not used
178 but were still cultivated under conditions similar to those mentioned. The remaining primary
179 producers were subjected to shaking at 80 relative centrifugal force (RPG) using a Laboshaker (C.
180 Gerhardt, Königswinter, Germany).

181 Carbonate chemistry was investigated to assess the influence of olivine additions on primary
182 producers' growth. Ground additions of 0.5 g L⁻¹ were added for olivine in the preliminary phase
183 of the experiment, giving a solid-liquid ratio of 0.0005 kg L⁻¹. This amount has been chosen based
184 on our earlier experiments (Rønning, in prep, 2024), in which we could obtain and maintain



185 significant increases in TA and pH over a 35-day runtime. Olivine was not detected to fully
186 dissolve in our experiments. Control batches for each strain were cultivated under the same
187 conditions as those with mineral additions but without mineral supplements.

188 Parameters that were investigated during the cultivation study of the 10 primary producers of
189 olivine additions were relative fluorescence units (RFU) or optical density (OD 680) (S2-
190 supplemental material) and pigment concentration. RFU and pigment analysis was conducted on
191 a Triology® Fluorometer (Model #7200-00) using a chlorophyll-a (chl-a) in vivo blue module
192 (Model #7200-043, both Turner Designs, San Jose, CA, USA), set to use a method for chl-a. Chl-
193 a samples were collected alongside every sample session to monitor growth over time by filtering
194 the samples onto GF/F filters (GE Healthcare Life Sciences, Whatman, USA). Filtered volumes
195 for chl-a were adjusted to the density of the primary producer cultures. Filters were stored in
196 darkened 15 mL LightSafe centrifuge tubes (Merck, Rahway, NJ, USA) at -20 °C until further
197 analysis. A calibration was generated for pigment analysis using a standard stock solution based
198 on chl-a from spinach (Sigma Aldrich, Burlington, USA). The samples were submerged in 8 mL
199 of 90% acetone overnight at a temperature of 5°C; the next day, 1 mL of the overnight liquid was
200 added into a 1.5 mL glass vial (Mikrolab Aarhus A/S, Aarhus, Denmark) and then analyzed on the
201 Triology® Fluorometer. RFU measurements were taken at each sampling point to monitor the
202 real-time growth of primary producers. 1 mL samples were placed into glass vials with a pipette
203 and analyzed in the Triology® Fluorometer. However, due to the unavailability of the Triology®
204 fluorometer, two primary producers, namely the CCMP1331 and A15-62 clonal strains, had their
205 growth analyzed by measuring the optical density at OD 680 with a Multiscan GO
206 spectrophotometer (Thermo Fisher Scientific, Waltham, MA, USA).

207 **2.4 Carbon fixation in the cultures**

208 We determined carbon fixation rates in the cultures to assess the potential impacts of olivine
209 additions on primary production. Samples were incubated with 20 µg mL⁻¹ 98 % sodium
210 bicarbonate ⁻¹³C before starting the experiments (Lot #MBB9999 and #MBBD2322; Sigma-
211 Aldrich, Burlington, USA). An initial set of biological triplicate samples was collected at the
212 beginning of the experiments before the additions of olivine as T0. After the total incubation time
213 for each primary producer (Tab. 3), samples were filtered onto pre-combusted GF/F filters (GE
214 Healthcare Life Sciences, Whatman, USA). Filters were stored at -20 °C until further analysis.



215 Filters were then acidified and dried before analysis using an Elemental Analyzer Flash 2000
216 (Thermo Fisher Scientific, Waltham, MA, USA) coupled to an isotope ratio mass spectrometer
217 Finnigan Delta V Advantage (Thermo Fisher Scientific, Waltham, MA, USA), as described earlier
218 (Reeder et al., 2022).

219 **2.5 Carbonate system parameters (DIC / pH / TA)**

220 Parameters measured within the carbonate system, including DIC, pH, and total TA, were collected
221 alongside every sample session to assess the effects of olivine additions on the primary producers'
222 effect on the carbonate system. The samples were taken up with a syringe that was equipped with
223 an 8-10 cm Iso-Versinic tube 4 mm (Saint Gobain, Courbevoie, France) and filtered with 0.22 μm
224 syringe filters (Avantor VWR[®] Radnor, Pa, USA). The first 1-3 mL that went through the syringe
225 filter were discarded due to the production residues of CO_2 . The filtered samples were stored in
226 amber glass vials with closed caps bonded with a PTFE-faced silicone liner to ensure a secure
227 sealing (Thermo Fisher Scientific, Waltham, MA, USA) at 5°C until further analysis. TA and pH
228 samples were measured using a Titrand 809 (Metrohm, Herisau, Switzerland) and an attached pH
229 microelectrode, 6.0234.100 (Metrohm, Herisau, Switzerland), which was calibrated against four
230 standard buffers daily before use (DuraCal pH buffer solutions, 2.00, 4.00, 7.00 Reagecon, Clare,
231 Ireland, and a 10.00 from Hamiltoncompany, Reno, USA). TA was measured based on Dickson
232 SOP3b (Dickson et al., 2007) using an open cell titration with the Titrand 809. To obtain
233 alkalinity measurements, 5 mL of each sample were analyzed on the Titrand 809 with mixing
234 and 0.1 mol L⁻¹ (0.1 N) HCl (Merck, Rahway, NJ, USA) while 0.6 mol L⁻¹ NaCl was added under
235 stirring. The analysis used the Tiamo 2.5 software (Metrohm, Herisau, Switzerland). To verify the
236 accuracy of the TA laboratory protocol a certified seawater reference was applied (Batch 187,
237 Scripps Institution of Oceanography, University of California, San Diego, USA). DIC samples
238 were analyzed using an AS-C5 DIC analyzer (ApolloSciTech, Newark, Delaware, USA) with a
239 laser-based CO_2 detector. Before analyzing the samples, 2 mM NaHCO_3 was used as a standard,
240 following the manufacturer's protocol, with three standard volumes at 0.9 mL, 1.2 mL, and 1.5 mL
241 in addition to a certified seawater reference for CO_2 measurements (Batch 187, Scripps Institution
242 of Oceanography, University of California, San Diego, USA). For analysis, a sample volume of
243 0.9 mL was extracted from each sample; the best of 3 out of a maximum of 5 measurements within
244 the precision range of 0.02 on the AS-C6L DIC analyzer determined a single DIC point.



245 **2.6 Trace metal concentrations**

246 For the elemental analysis of the water, samples were analyzed for concentrations of dissolved
247 trace elements, including aluminum (Al), As, copper (Cu), Fe, Mg, Mn, and Ni, using an Agilent
248 7900 Quadrupole Inductively Coupled Plasma (ICP) Mass Spectrometers instrument (Agilent
249 Technologies, Ca, USA). Calibration standards (20, 40, 60, 80, and 100 mg L⁻¹) were prepared in
250 5% nitric acid from 1000 mg L⁻¹ ICP standards. Samples were prepared in 5% nitric acid and stored
251 at room temperature until further analysis (Olesik, 2013). Samples were collected both at the start
252 and at the end of the experiments, following the procedure described above. However, elemental
253 data for strain CCMP489 and EUM Syn103 are not available.

254 **2.7 Nutrient concentrations**

255 During the experimental runs, we monitored the development of nitrate (NO₃⁻), nitrite (NO₂⁻), and
256 phosphate (PO₄³⁻). Samples were collected alongside every session for analysis. The nutrient
257 samples were taken up with a syringe and filtered with (Avantor VWR® Radnor, Pa, USA) 0.22
258 µm syringe filters and stored in 1.5 mL Eppendorf microtubes (Eppendorf, Hamburg, Germany)
259 at -20°C until further analysis. The following standards were used: a stock solution of KNO₃ for
260 NO₃⁻, NaNO₂ for NO₂⁻, and KH₂PO₄ for PO₄³⁻. Standard dilution series were at 0, 5, 10, 20, and
261 30 µmol. These solutions served as reliable references for the analysis of the respective nutrients.
262 The samples were analyzed with a Multiscan GO spectrophotometer (Thermo Fisher Scientific,
263 Waltham, MA, USA) on OD at 680 nanometers, according to (Grasshoff et al., 1999) protocol.
264 Samples are to be found in raw data on request.

265 **3 Results**

266 This study aimed to examine the impact of small amounts of ground olivine additions on isolated
267 primary producers. To achieve this, we conducted controlled experiments on primary producers to
268 evaluate the ability of OAE in both coastal and open ocean environments. We aimed to discern the
269 impact of olivine additions on the growth of these primary producers.

270 **3.1 Responses of carbonate chemistry to olivine additions**

271 **3.1.1 Total alkalinity over the course of the experiments**



TA concentrations displayed variability in cultures of different species, as exemplified by the three *Synechococcus* strains (Fig. 1. a-c), while differences between primary producers were generally detectable. A similar pattern was observed in *Synechococcus* sp. strain A15-62 Clonal (Fig. 1. a) and *Prochlorococcus* sp. strain NATL2-M98 (Fig. 1. d). The influence of olivine additions on TA and the control groups varied between strains. Olivine additions had a bigger effect on TA concentrations during the experimental period for *Synechococcus* sp. strain WH8102 cultures compared to the control (Fig. 1. c). Among the *Synechococcus* sp. strains examined the most pronounced variations in carbonate chemistry were observed in the case of strain A15-62 Clonal, where olivine additions elevated concentrations to $3159 \mu\text{mol kg}^{-1}$ on day 6 (Fig. 1. a). It is worth highlighting that the olivine treatments, including the control, experienced a decline in TA concentrations throughout the experiment (Fig. 1. a). The EUM Syn10 strain showed similar TA concentration fluctuations between the control and olivine-treated groups. The olivine addition exhibited a low decline, decreasing from 3158 to $3121 \mu\text{mol kg}^{-1}$ at the end of the experimental period (Fig. 1. b). The control group showed a decrease in TA concentrations from the beginning to the end of the experimental period, with fluctuations occurring intermittently between days 2 and 9 (Fig. 1. b). In the *Synechococcus* strain WH8102, olivine additions and the control strain had very similar patterns showing declines in TA concentrations during the experimental phase (Fig. 1. c). The trend in TA concentrations was considerably similar among the olivine addition and the control for the *Prochlorococcus* sp. strains, both experiencing a decline from the beginning to the end of the experiment (Fig. 1. d-e). However, olivine additions showed some fluctuations, with TA concentrations increasing in the first few days of the experiments, peaking at day 4 with $3236 \mu\text{mol kg}^{-1}$ before declining again (Fig. 1. d). Similar observations appeared for olivine additions for the *Prochlorococcus* sp. strain EQPAC1-C, where TA concentrations increased from $3015 \mu\text{mol kg}^{-1}$ at day 0 to $3373 \mu\text{mol kg}^{-1}$ at day 2 before subsequently declining (Fig. 1. e). For *Scrippsiella trochoidea*, CCMP1331, TA concentrations increased for the olivine treatment and the control group during the experimental period, with olivine additions reaching the maximum peak at $3087 \mu\text{mol kg}^{-1}$ (Fig. 1. f). For *Micromonas commoda*, the fluctuations in the olivine treatments closely resembled those in the control group from day 3 until the end of the cultivation period (Fig. 1. g). For diatoms, the control displayed the highest values on TA concentrations, both in terms of general concentrations and maximum values, as exemplified by CCMP1051's $3880 \mu\text{mol kg}^{-1}$ (Fig. 1. h) and CCMP2092's $3827 \mu\text{mol kg}^{-1}$ (Fig. 1. i). For the *Emiliania huxleyi* strain,



ACC448, the influence of olivine additions on TA concentrations varied, similarly did the control.
For ACC448, olivine reached its maximum peak of 2742 $\mu\text{mol kg}^{-1}$ on day 3 (Fig. 1. j).

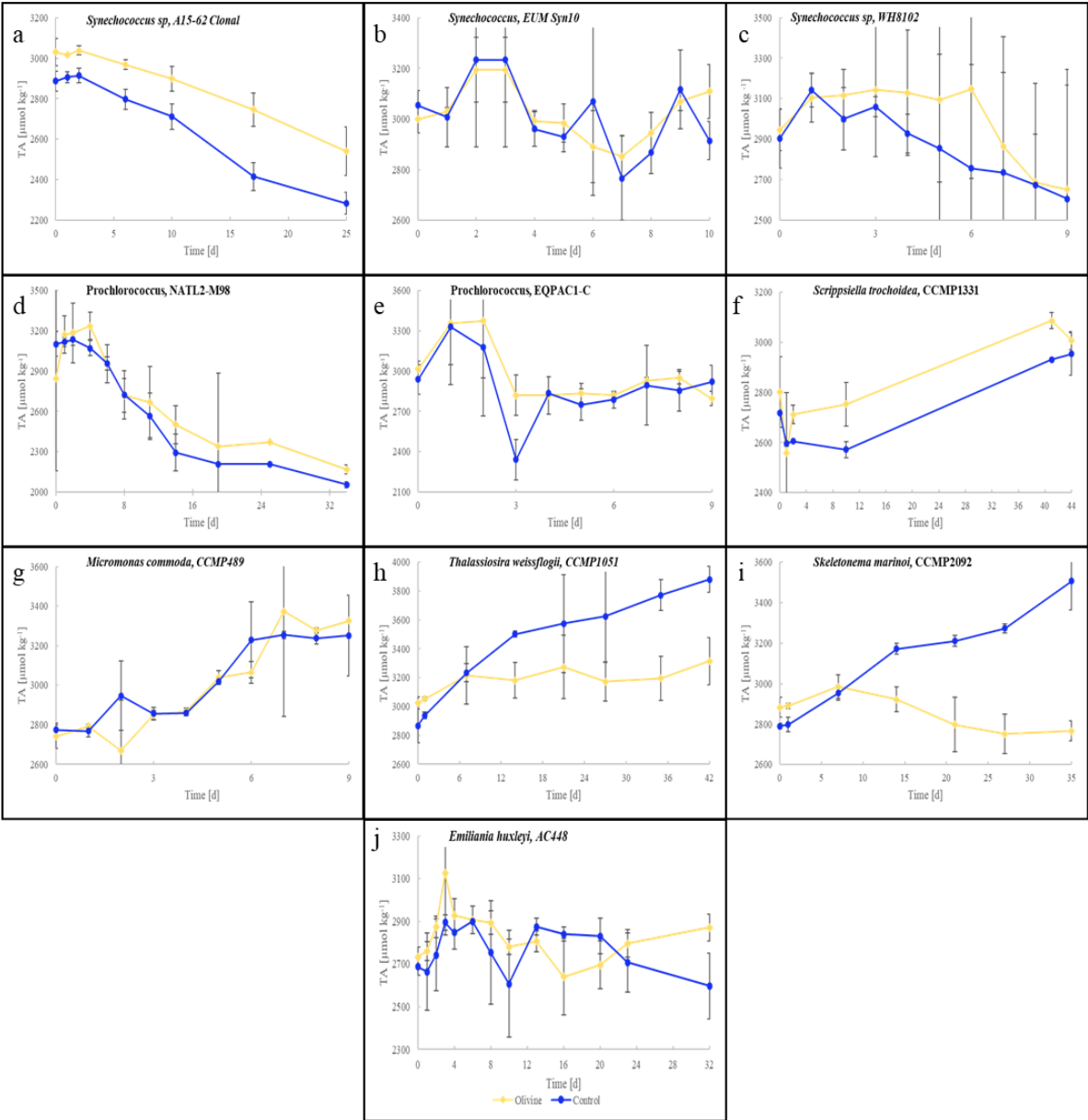


Figure 1. Total alkalinity concentrations ($\mu\text{mol kg}^{-1}$) in cultures exposed to olivine (yellow, diamond) additions and the control without any olivine supplement (blue, circle). The standard deviation is represented with grey bars.

3.1.2 pH changes over the course of the experiments



309 Within our study, we observed pH across the primary producers examined. Despite the addition of
310 olivine being predicted to be the primary driver of pH changes, some cultures also exhibited
311 changes parallel with these in their control groups. The strains showing such analogous pH patterns
312 include *Synechococcus* sp. EUM Syn10 (Fig. 2. b), *Scrippsiella trochoidea* CCMP1331 (Fig. 2. f),
313 *Micromonas commoda* CCMP489 (Fig. 2. g), and the diatom strains CCMP1051 (Fig. 2. h), and
314 CCMP2092 (Fig. 2. i). In multiple but not all cultivations, we noted an initial elevation in pH
315 resulting from olivine additions compared to the control group (Fig. 2. a, d, e, f, and j). Certain
316 incubations displayed an initial pH increase for specific olivine additions surpassing 0.2 relative
317 to the control (Fig. 2. a and j). The *Synechococcus* sp. strain A15-62 Clonal displayed a pH increase
318 from the introduction of olivine until the end of the cultivation period, ranging from 7.83 to 8.05
319 pH (Fig. 2. a). Similarly, in the *Prochlorococcus* sp. strain, EQPAC1-C pH increased from 7.8 to
320 7.92 (Fig. 2. e) after olivine additions. In contrast, pH levels declined from the initial incubation
321 to the end of the cultivation period in response to olivine additions for the two diatom strains. For
322 the diatom strain CCMP1051, pH decreased from 8.36 to 8.09 (Fig. 2. h), and strain CCMP2092
323 decreased from 8.60 to 7.76 pH (Fig. 2. i) during an incubation period of 35 days. *Emiliania*
324 *Huxleyi*, AC448 showed a pH decline from 7.99 to 7.86 in response to the olivine additions during
325 the experimental period (Fig. 2. j), the control strain maintained a pH range between 7.77 and 7.81
326 throughout the experimental period, with fluctuations occurring intermittently during the
327 cultivation (Fig. 2. j). In the case of the pH changes for both the olivine additions and the control
328 strain of *Synechococcus* sp. WH8102 varied during the 9-day cultivation period, fluctuating
329 between 7.68 and 8.04 pH for the olivine additions and between 7.68 and 8.08 for the control strain
330 (Fig. 2. c).

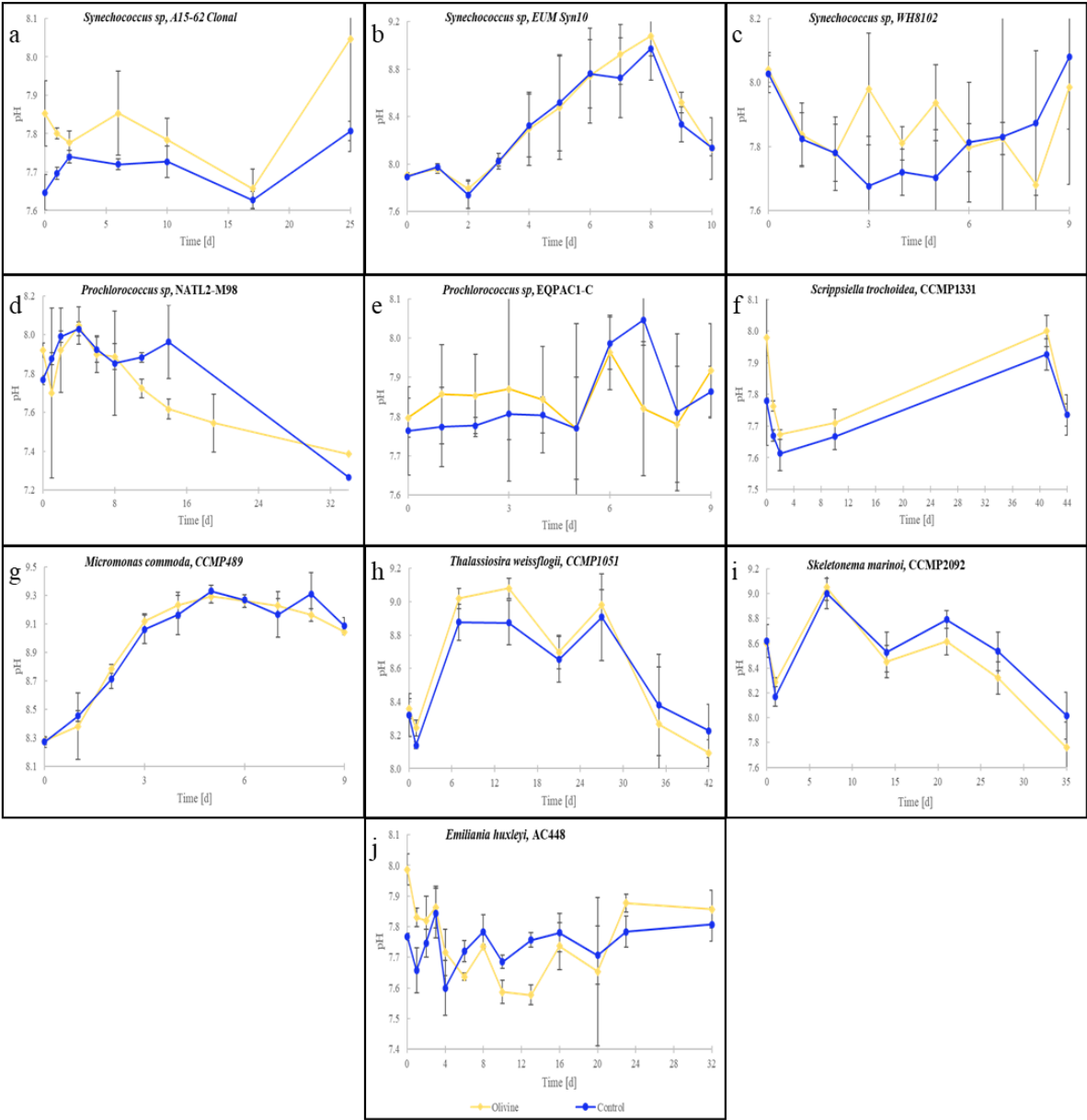


Figure 2. pH in cultures exposed to olivine (yellow, diamond) additions and the control without any olivine supplement (blue, circle). The standard deviation is represented with grey bars.

3.1.3 Dissolved inorganic carbon changes over the course of the experiments

The influence of olivine supplements on DIC concentrations within enclosed systems demonstrated substantial variability (Fig. 3). Throughout most of the experimental duration, the



337 highest DIC concentrations were consistently observed in the control of the diatom strains
338 CCMP1051 and CCMP2092 (Fig. 3. h and i). Among the *Synechococcus sp.* strains investigated,
339 variations in carbonate chemistry were observed, with the strain A15-62 Clonal displaying the
340 most visible changes from the beginning to the end of the experiment (Fig. 3. a). The addition of
341 olivine influenced the DIC concentration by peaking at $2038.39 \mu\text{mol kg}^{-1}$, whereas the DIC
342 concentration of the control strain remained consistently lower throughout the experiment (Fig. 3.
343 a). For the EUM Syn10 strain, the control alongside olivine additions had comparable
344 developments in DIC concentrations (Fig. 3. b). In the case of *Synechococcus sp.* strain WH8102,
345 olivine additions yielded a higher effect than the control group over the experimental period. The
346 olivine addition reached a maximum peak of $3096 \mu\text{mol kg}^{-1}$ on day 5 (Fig. 3. c). Across the three
347 *Synechococcus sp.* strains, the addition of olivine induced more pronounced changes in DIC
348 concentrations in strains A15-62 Clonal and WH8102 compared to the control group, while strain
349 EUM Syn10 exhibited similar changes to the control (Fig. 3. a-c). For the *Prochlorococcus sp.*
350 strains, DIC concentrations have similarities between the olivine addition and the control.
351 Nevertheless, the olivine additions resulted in the highest maximum DIC concentrations for both
352 *Prochlorococcus sp.* strains NATL2-M98 at $2707.75 \mu\text{mol kg}^{-1}$ on day 1 (Fig. 3. d), and EQPAC1-
353 C at $2782.60 \mu\text{mol kg}^{-1}$ on day 2 (Fig. 3. e), as well as for *Scrippsiella trochoidea* CCMP1331 at
354 $3420.12 \mu\text{mol kg}^{-1}$ on day 44 (Fig. 3. f) compared to the control strain. Olivine additions seemed
355 to have a negative impact on the DIC concentrations on the *Micromonas commoda*, CCMP489
356 strain throughout the experimental duration, compared to the control strain (Fig. 3. g.). We
357 observed that olivine additions had a negative impact on DIC concentrations in the diatom strains
358 compared to the control strain. Explicitly, both the overall increase and the maximum
359 concentrations increased more in the control strain than in the olivine-treated strains (CCMP1051,
360 $3501.42 \mu\text{mol kg}^{-1}$; Fig. 3. h; and CCMP2092, $3734.43 \mu\text{mol kg}^{-1}$; Fig. 3. i). Consistent with the
361 patterns observed in TA concentrations (Fig. 1). The DIC concentrations for the *Emiliania huxleyi*
362 ACC448 strain, olivine additions, had a considerable influence, with olivine having a pronounced
363 effect until day 12 compared to the control group (Fig. 3. j).

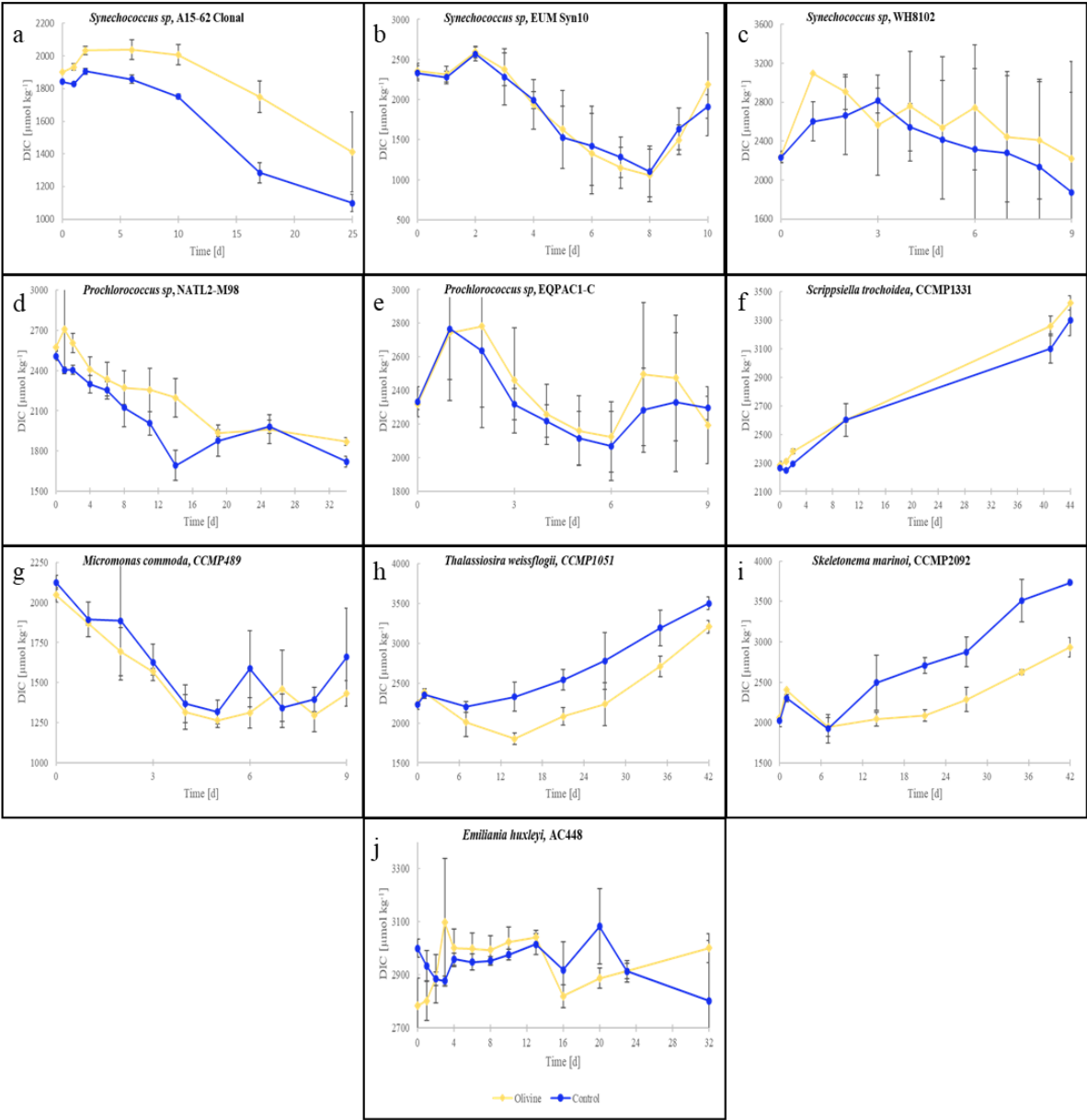


Figure 3. Dissolved inorganic carbon concentrations ($\mu\text{mol kg}^{-1}$) in cultures exposed to olivine (yellow, diamond) additions and the control without any olivine supplement (blue, circle). The standard deviation is represented with grey bars.

3.2 Impact of olivine additions on primary producer growth

Chl-a concentrations were utilized as the primary indicator for assessing the growth patterns of the primary producers. The chl-a concentrations reflected the primary production activity during the



370 experiments. Fluorescence measurements were taken to obtain data supporting chl-a
371 concentrations, including RFU and OD (S1-supplemental material). Overall, we did not find a
372 solid response to olivine additions compared to untreated controls concerning chl-a development
373 over the course of the experiment, except *Emiliana huxleyi* AC448, which did not show
374 indications of an increase after adding olivine (Fig. 4, j). Strains generally and expectedly showed
375 species-specific growth patterns with chl-a maxima ranging from $0.55 \mu\text{g L}^{-1}$ to $3222.99 \mu\text{g L}^{-1}$
376 between the cultivations (Fig. 4). The maximum chl-a concentration for *Synechococcus sp.* strains
377 under olivine additions ranged from $84.10 \mu\text{g L}^{-1}$ on day 25 (Fig. 4. a), $14.71 \mu\text{g L}^{-1}$ on day 8 (Fig.
378 4. b), and $18.25 \mu\text{g L}^{-1}$ max (Fig. 4. c) on day 6. The control group demonstrated a higher chl-a
379 concentration compared to the olivine additions for the *Synechococcus sp.* strain WH8102
380 throughout the experiment, except on day 6, the control group reached a maximum of $21.70 \mu\text{g L}^{-1}$
381 (Fig. 4. c). For *Prochlorococcus sp.* strains, the chl-a concentrations for olivine additions were
382 between 614.91 on day 8 (Fig. 4. d), whereas the control group reached a maximum on 1430.00
383 $\mu\text{g L}^{-1}$ and were generally at higher chl-a concentration than the olivine additions (Fig. 4. d). for
384 the strain EQPAC1-C it reached a max on day 9 at $216.15 \mu\text{g L}^{-1}$ (Fig. 4. e). *Scrippsiella*
385 *trochoidea*, CCMP1331, reached a peak chl-a concentration of $69.27 \mu\text{g L}^{-1}$ on day 41 following
386 the olivine addition (Fig. 4. f). *Micromonas commoda*, CCMP489, exhibited a maximum of the
387 olivine addition at $2947.73 \mu\text{g L}^{-1}$ on day 6, whereas the untreated control reached a peak at
388 2860.40 on day 6 (Fig. 4. g). The diatom strains' chl-a concentration peaked with maxima in the
389 untreated controls between 3222.99 on day 21 for strain CCMP1051 (Fig. 4. h) and $3031.09 \mu\text{g L}^{-1}$
390 on day 21 for strain CCMP2092 (Fig. 4. i). The maximum chl-a concentrations of the *Emiliana*
391 *huxleyi* strain were at $0.55 \mu\text{g L}^{-1}$ on day 23, reached in the control group (Fig. 4. j). Olivine
392 addition was associated with low chl-a concentrations for the *Emiliana huxleyi* AC448 at $0.04 \mu\text{g}$
393 L^{-1} (Fig. 4. j).

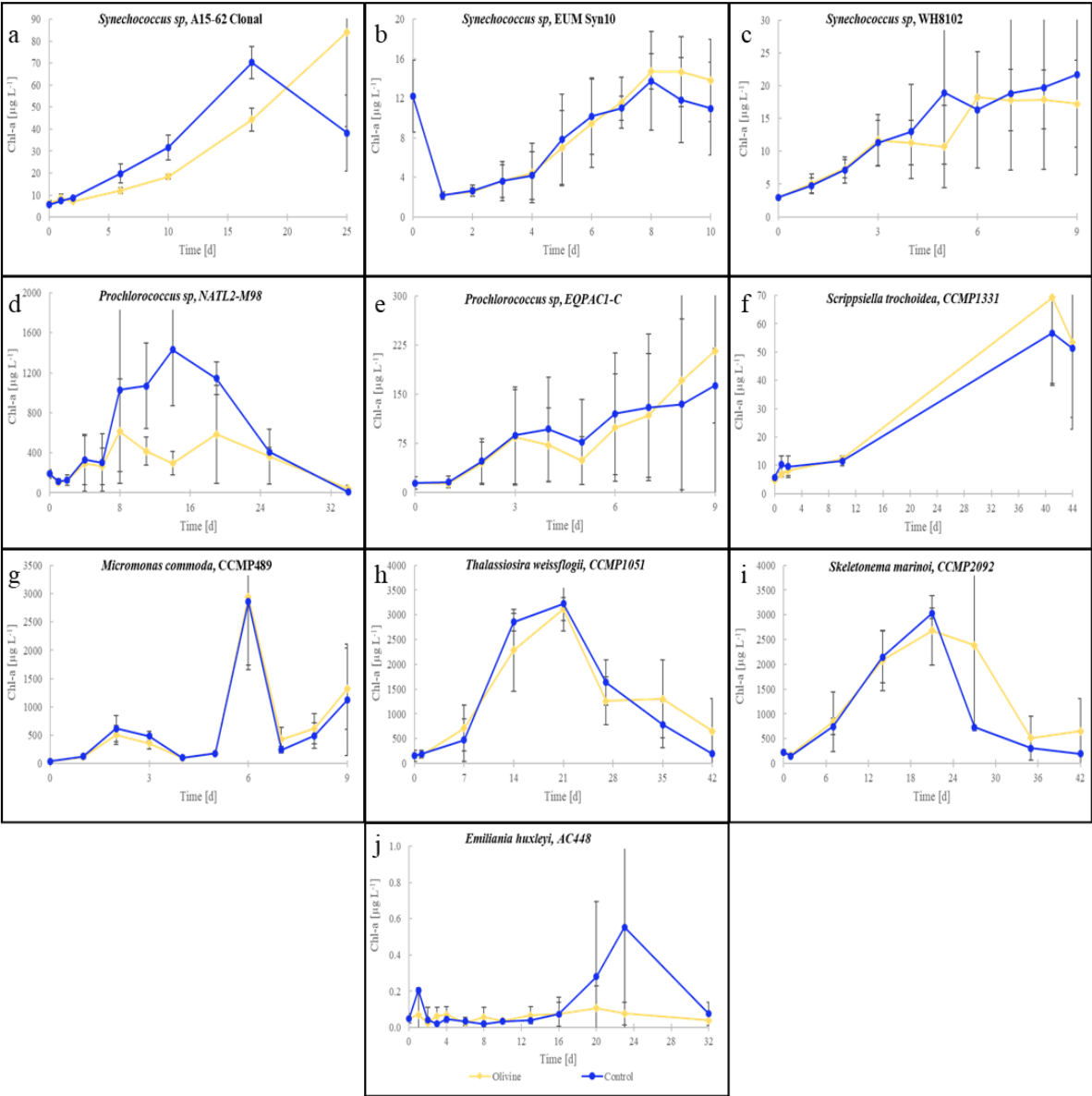


Figure 4. Chl-a concentration ($\mu\text{g L}^{-1}$) in cultures exposed to olivine (yellow, diamond) additions and the control experiments without any olivine supplement (blue, circle). The standard deviation is represented with grey bars.

3.3 Impact of olivine additions on carbon fixation rates

In contrast to a relatively low impact of olivine additions on chl-a concentrations, as described above, we do not see a uniform response of carbon fixation rates to alkaline mineral additions. We identified species-specific responses in carbon fixation rates in response to the addition of olivine



(Fig. 5). The *Synechococcus sp.* showed carbon fixation rates in our experiment, ranging between 61.86213 C L⁻¹ d⁻¹ reached by the olivine addition (Fig. 5. a), 188.47021 C L⁻¹ d⁻¹ reached by the olivine addition (Fig. 5. b), and 159.55055 μmol C L⁻¹ d⁻¹ reached by the untreated control (Fig. 5. c). The carbon fixation rates at the end of the experiment were substantially higher than those at T0 for all three *Synechococcus sp.* strains (Fig. 5. a-c). For the *prochlorococcus sp.* NATL2-M98 strain, the carbon fixation rate reached 27.23686 μmol C L⁻¹ d⁻¹ by the untreated control group and 20.84817 μmol C L⁻¹ d⁻¹ in the olivine addition treatment (Fig. 5. d). The EQPAC1-C strain reached the highest carbon fixation rate with olivine addition, reaching 195.51166 μmol C L⁻¹ d⁻¹ (Fig. 5. e). The carbon fixation rates at the end of the experiment were substantially higher than those at T0 for the two *prochlorococcus sp.* strains (Fig. 5. d-e). *Scrippsiella trochoidea* exhibited a carbon fixation rate of 0.69189 μmol C L⁻¹ d⁻¹ following olivine addition (Fig. 5. f), whereas the control group demonstrated a carbon fixation rate at 0.35906 μmol C L⁻¹ d⁻¹ after 44 days, marginally surpassing the carbon fixation rate measured at T0 0.31303 μmol C L⁻¹ d⁻¹ (Fig. 5. f). *Micromonas commoda* showed a carbon fixation rate of 453.97966 μmol C L⁻¹ d⁻¹ following olivine addition (Fig. 5. g), the carbon fixation rates of both the control and the olivine treatment considerably exceeded those at T0 (Fig. 5. g). The diatoms displayed carbon fixation rates following olivine additions, reaching 69.08472 μmol C L⁻¹ d⁻¹ for the CCMP1051 strain (Fig. 5. h) and 44.18524 μmol C L⁻¹ d⁻¹ for the CCMP2092 strain (Fig. 5. i). In comparison, the untreated control showed carbon fixation rates of 56.29830 μmol C L⁻¹ d⁻¹ for the strain CCMP1051 (Fig. 5. h) and the strain CCMP2092 at 41.36381 μmol C L⁻¹ d⁻¹ (Fig. 5. i). Both diatoms demonstrated substantially higher carbon fixation rates at the end of the experiment compared to those at T0 (Fig. 5. h-i). The *Emiliania huxleyi* strain AC448 exhibited a carbon fixation rate of 1.40959 μmol C L⁻¹ d⁻¹ following olivine addition (Fig. 5. j), whereas the untreated control had a carbon fixation rate of 0.15304 μmol C L⁻¹ d⁻¹, a rate much closer to T0 of 0.02907 μmol C L⁻¹ d⁻¹ in contrast to the olivine addition (Fig. 5. j).

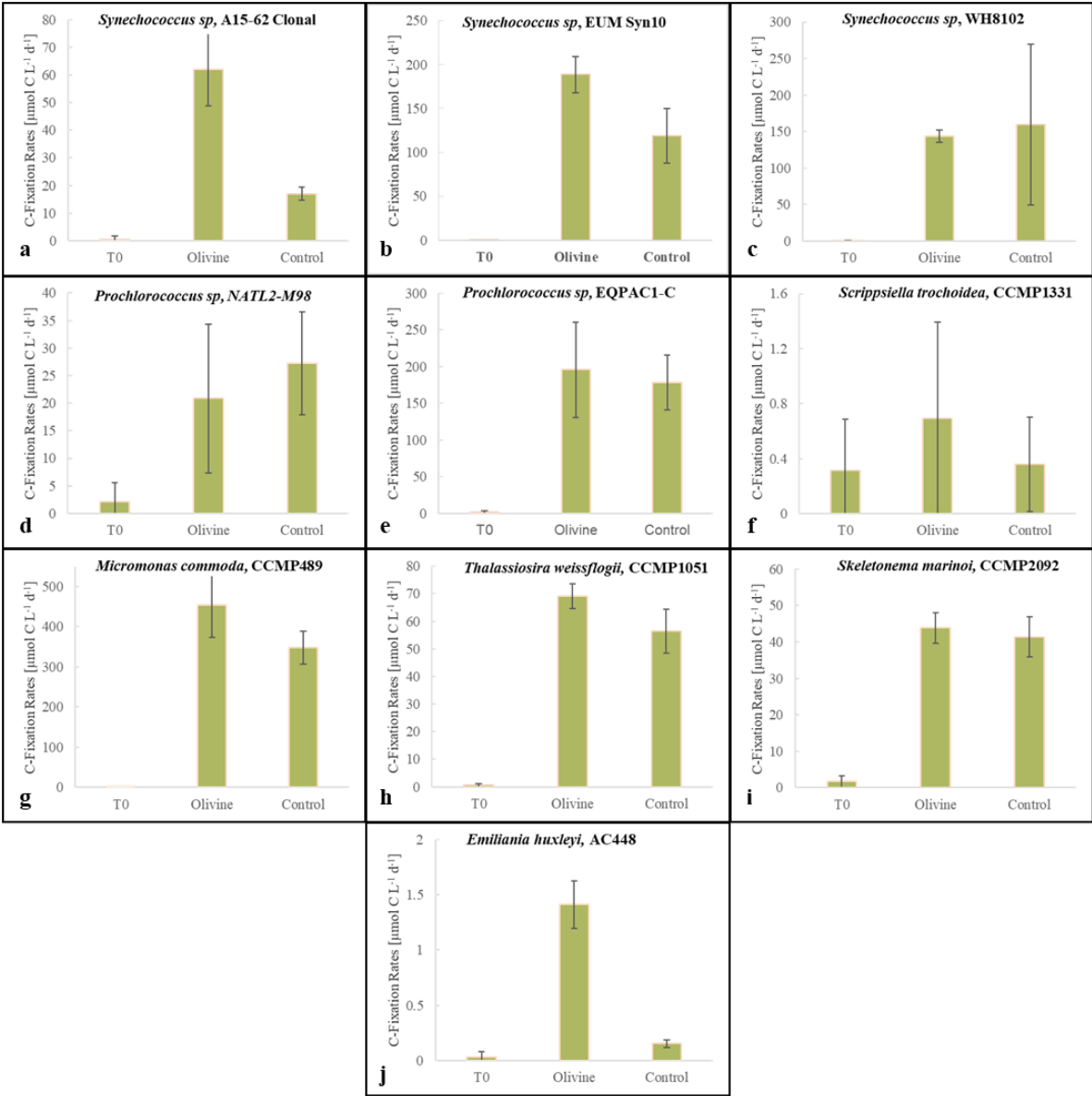


Figure 5. Carbon fixation rates ($\mu\text{mol C L}^{-1} \text{d}^{-1}$). T0 displays the initial biological triplicates at the beginning of the experiments. Olivine display the effect of ground alkaline minerals treatment at the end of the experimental period, while the control strain was cultivated without any alkaline mineral additions. The standard deviation is represented with grey bars.

3.4 Development of nickel content during the experimental period

Overall, the nickel concentrations observed in the various cultivation experiments were considerably higher in mineral addition experiments compared to the controls (Fig. 6). For



instance, in both *Synechococcus sp.* strains exhibited nickel concentrations at the end of the experiment were observed. The strain A15-62 Clonal reached a nickel concentration of 1258.59 $\mu\text{g L}^{-1}$ (Fig. 6. a), while strain WH8102 showed a concentration 536.63 $\mu\text{g L}^{-1}$ (Fig. 6. b). However, in *Scrippsiella trochoidea* we found a decrease in nickel concentration since the start of the experiment, with nickel concentrations declining by 262.57 $\mu\text{g L}^{-1}$ following the olivine addition reaching a nickel concentration of 36.96 $\mu\text{g L}^{-1}$ at the end of the experiment (Fig. 6. c). The nickel concentration in the *Prochlorococcus sp.* NATL2-M98 experiment reached substantial concentrations, measuring 2347.08 $\mu\text{g L}^{-1}$, and experienced an increase of 1730.25 $\mu\text{g L}^{-1}$ after the addition of olivine (Fig. 6. d). Conversely, the *Prochlorococcus sp.* EQPAC1-C, experienced a decrease in nickel concentration following the olivine addition at the beginning of the experiment with a decrease of 400.61 $\mu\text{g L}^{-1}$, ending at a nickel concentration of 93.45 $\mu\text{g L}^{-1}$ (Fig. 6. e). At the end point of the experiment, *Emiliania huxleyi* strain AC448 exhibited a nickel concentration of 583.38 $\mu\text{g L}^{-1}$ following the addition of olivine. This concentration reflected a decrease of 325.78 $\mu\text{g L}^{-1}$ from the initial olivine treatment (Fig. 6. f). The diatoms exhibited varying responses in nickel concentrations following olivine addition. For the CCMP1051 strain, the nickel concentration decreased by 208.33 $\mu\text{g L}^{-1}$ to 972.07 $\mu\text{g L}^{-1}$ towards the end of the experiment (Fig. 6. g), while the CCMP2092 strain experienced an increase in the nickel concentration by 115.60 $\mu\text{g L}^{-1}$ increasing to 840.43 $\mu\text{g L}^{-1}$ since the olivine treatment (Fig. 6. h).

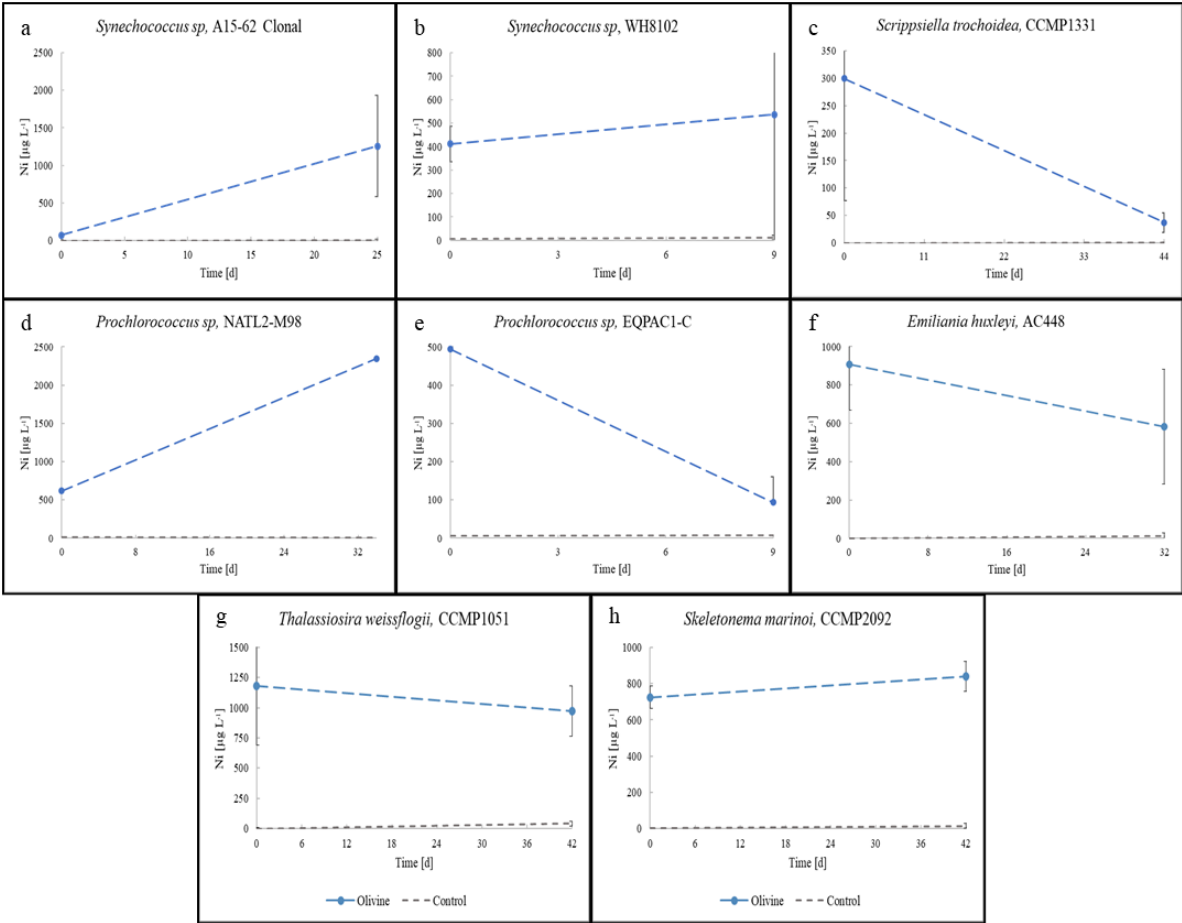


Figure 6. Nickel concentration ($\mu\text{g L}^{-1}$) in cultures exposed to olivine (blue, circle, long dash) additions and the control experiments without any olivine supplement (blue, small dash). The standard deviation is represented with grey bars.

4 Discussion

4.1 Experimental considerations and concerns

Our methodology in this experiment warrants consideration for its potential impact on the overall results. The use of different media for the various algae species may have influenced TA development (Fig. 1), particularly in the case of cyanobacteria cultured in PCR S11 Red Sea Medium (Tab. 3). The inclusion of HEPES buffer in PCR S11 Red Sea Medium could have mitigated the alkalization effect on pH. HEPES acts as a pH buffer, thereby stabilizing the aquatic cultivation media and promoting consistency in experimental conditions (McFadden and Melkonian, 1986). Additionally, the use of silica in the L1 medium, which was incorporated on



463 for the diatom to enhance their growth (Tab. 3), might have resulted in co-precipitation within the
464 cultivation medium following the addition of the alkaline mineral.

465 The addition of sodium bicarbonate, known to act as a buffer to maintain alkalinity in the solution,
466 introduces an additional layer of concern. Sodium bicarbonate is recognized as a compound that
467 effectively increases alkalinity (Furtado et al., 2011). Its inclusion prior to the olivine additions
468 likely influenced the experimental conditions across all cultivations for the alkalinity. Therefore,
469 the presence of sodium bicarbonate needs to be carefully considered when interpreting the results.
470 Another apparent consideration is the potential impact of nutrient concentrations on alkalinity
471 within the different media used in our experiments. The medias contain varying concentrations of
472 nitrate (NO_3^-) and nitrite (NO_2^-) and ammonia (NH_4^+) (Tab. 3). Furthermore, the k and L1 media
473 consist of seawater from the North Sea, potentially containing higher nutrient concentrations from
474 land sources, while the PCR S11 Red Sea Medium is based on deionized water. This variability in
475 nutrient composition raises concerns about the extent to which these nutrients have affected
476 alkalinity concentrations, as a release of NO_3^- and NO_2^- decreases alkalinity and a release of NH_4^+
477 increases alkalinity (Wolf-Gladrow et al., 2007). Therefore, the differing nutrient compositions of
478 the medias may have contributed to variations in alkalinity concentrations detected during our
479 experiments.

480 The shaking or no shaking could potentially impact the growth of the primary producers in relation
481 to the nickel concentrations (Tab. 3). However, it appears that this variable did not have a
482 noticeable effect on either the growth of the organisms or the concentration of nickel (Fig. 4 and
483 6), as for example *Scrippsiella trochoidea* was not mixed and it does show good growth compared
484 to most of the other primary producers.

485 This study is one of the first investigations directly examining the response of naturally occurring
486 olivine additions on both primary producers and general carbonate chemistry (S2-supplemental
487 material). In contrast to other, earlier studies (Delacroix et al., 2023; Hutchins et al., 2023), we
488 worked directly with olivine, rather than other forms of alkalinity. This decision was made to
489 mimic the process of OAE through the addition of mineral-based substances in aquatic conditions.
490 By doing so, we aimed to replicate the complex interactions and potential impurities present in
491 naturally occurring olivine, which may be challenging to explore when using artificial alkaline
492 products. The presence of sand and other non-alkaline components in the olivine might further



493 impact organisms and aquatic ecosystems. Moreover, the smaller grain size in this study,
494 approximately $\sim 63\mu\text{m}$, distinguishes it from and earlier study Fuhr et al. (2022) in which grain
495 sizes ranging from $100\text{--}125\mu\text{m}$ and additions ranging between 25g and 100g per L were applied.
496 While our additions were significantly lower compared to other studies, we used smaller grain
497 sizes, thus increasing the reactive surface of the particles.

498 **4.2 Potential effects of Ni on phytoplankton cultures**

499 Our study could not identify fatal effects or growth inhibition on the cultivated primary producers
500 exposed to olivine minerals despite increases in Ni concentrations for several species observed in
501 the media following the olivine additions (Fig. 6). Neither chl-a concentrations nor carbon fixation
502 rates were impacted compared to the other olivine addition and the control group. Interestingly,
503 recent studies investigating the toxic impact of Ni on phytoplankton species yielded similar
504 findings (Guo et al., 2022; Hutchins et al., 2023). Still, we observed an immediate and continuous
505 increase in Ni concentrations from ground olivine. A recent study by Xin et al. (2023) directly
506 investigated the impact of various concentrations of Ni on algae after instantaneous exposure (Xin
507 et al., 2023). The study showed varying responses of phytoplankton strains to changes in Ni
508 concentrations, with some strains being more impacted than others. Specifically, the
509 coccolithophore *Emiliana huxleyi* and the dinoflagellate *Amphidinium carterae* were only mildly
510 impacted. The investigated strain of *Emiliana huxleyi* did not show many signs of growth for
511 either the control or the olivine addition (Fig. 4 and 5. j), but it did show a decline of nickel
512 concentration in the aquatic media (Fig. 6. f). In our study on a dinoflagellate, carbon fixation rates
513 in *Scrippsiella trochoidea* were observed to be the lowest compared to T0 (Fig. 5f). Interestingly,
514 while nickel concentrations declined during the experiment (Fig. 6. c), chl-a concentrations
515 increased (Fig. 4. f). This suggests that *Scrippsiella trochoidea* potentially were not impacted by
516 the nickel concentrations in the media. Instead, it is possible that the *Scrippsiella trochoidea* had
517 an alternative metabolic pathway to support its growth and biomass production.

518 In contrast, the diatom *Thalassiosira weissflogii* displayed a high inhibition in growth rate with
519 increasing Ni concentrations, specifically at the three highest concentrations at $20\mu\text{mol L}^{-1}$, 50
520 $\mu\text{mol L}^{-1}$, and $100\mu\text{mol L}^{-1}$. These findings imply that the diatom *Thalassiosira weissflogii* has a
521 lower tolerance to Ni than the other cultures tested (Xin et al., 2023). Endpoint concentrations
522 during our experiments were towards the lower end of the range tested in that study, and ranged



523 between 0.6 – 40.0 $\mu\text{mol L}^{-1}$ (Figure 6), depending on cultivation media and primary producers.
524 These concentrations are still several orders of magnitude higher than ocean average Ni
525 concentration in surface waters at 2.1 nmol kg^{-1} , with a maximum of 11 nmol kg^{-1} in the deep sea
526 (Bruland, 1980). Still, our study can confirm the results by Xin et al. (2023), with even high
527 concentrations of Ni not significantly impacting the growth of the primary producers resulting
528 from the olivine additions.

529 In contrast, Hutchins et al. (2023) observed similar patterns regarding the response of different
530 phytoplankton taxa towards Ni exposure. Their study describes that most of the investigated
531 primary producers were irresponsive to the exposure of the applied Ni concentrations.
532 Interestingly, the synthetic olivine dissolution applied in Hutchins et al. (2023) had lower Ni
533 concentrations compared to that in Xin et al. (2023); the experimental designs hinder a precise
534 comparison between the studies discussed. This is in accordance with other recent observations
535 studying OAE scenarios regarding trace metals (Guo et al., 2022). Guo et al. (2022) specifically
536 focused on elevated Ni exposure in a broad range of phytoplankton groups; similar to our study,
537 they did not observe any strong effects across the investigated taxa. It is important to note that we
538 are unaware of the specific chemical species of dissolved Ni affecting the physiology of primary
539 producers. However, evidence suggests that phytoplankton are primarily sensitive to free Ni^{2+} ions
540 rather than the total concentration of dissolved Ni (Guo et al., 2022).

541 As Flipkens et al. (2021) proposed, establishing guidelines for trace metals, including Ni and Cr,
542 would be highly advantageous to ensure the safe implementation of coastal enhanced weathering.
543 Extending these guidelines to cover OAE would be equally beneficial, helping to mitigate
544 uncertainties in marine-based CDR while considering various trophic levels (Flipkens et al., 2021).
545 Additionally, to avoid the ecological concerns regarding potentially toxic trace metals from
546 alkaline minerals in OAE processes, trace metals like Ni could be extracted, for instance, from
547 olivine. (Santos et al., 2015).

548 **4.3 Effects of olivine additions**

549 Our experiments suggest that adding olivine is either slightly beneficial or did not negatively
550 impact the tested primary producers, with no harm to growth parameters such as the carbon fixation
551 rates, and chl-a concentrations (Fig. 4 and 5). Our primary producers were cultivated in specific



552 growth media that provided all necessary nutrients for growth (Tab. 3), ensuring they had abundant
553 essential components for optimal growth. This approach minimized bias and sensitivity, which can
554 be more pronounced in nitrogen and phosphate-limiting marine environments, but it also added a
555 layer of concerns and considerations to our experiments as mentioned above in the discussion
556 section 4.1. Consequently, the addition of olivine is expected to have a lesser impact on growth in
557 these conditions. Additionally, our primary producers were cultivated in closed systems that
558 experience no natural mixing and blending with sea- or freshwater. Residence times of the olivine
559 additions in the water column are relatively brief (He and Tyka, 2023), implying that dissolution
560 in natural settings is expected to occur more rapidly than in our experiments. Our study deliberately
561 introduced substantial quantities of pulverized olivine to conservatively explore their impact on
562 primary producers.

563 Consequently, we anticipate that the olivine additions in our experiments likely encountered
564 considerably higher concentrations of dissolved components from the olivine additions than under
565 typical field or natural conditions. Our experiments primarily emphasized the responses of primary
566 producers regarding a select set of physiological indicators and carbonate chemistry. Future
567 investigations should broaden the scope to assess the effects on various trophic levels and
568 encompass phytoplankton belonging to different central taxonomic groups.

569 **4.4 Where could ocean alkalinity enhancement be applied?**

570 Our study has revealed the potential use of olivine. To determine suitable regions for implementing
571 OAE practices, it is crucial to comprehend the regional and local variations in primary producers
572 and the aquatic chemistry.

573 In our experiments, we have demonstrated that regions characterized by abundant phytoplankton
574 populations, particularly picophytoplankton (cyanobacteria and *Micromonas*), tended to exhibit
575 either positive or neutral responses to olivine additions factors like growth rates, carbon fixation
576 rates, and TA concentrations. The diatoms control group exhibited slightly higher increases in
577 chlorophyll-a concentrations compared to the olivine additions (Fig. 4). However, olivine
578 additions had a greater impact on carbon fixation rates for the diatoms (Fig. 5). Interestingly,
579 physical processes such as eddies, coastal upwelling, wind, or convective mixing in oceanic
580 mechanisms can promote the occurrence of diatom blooms or larger cells of primary producers as



581 nutrient-rich waters enter the surface from the deeper ocean (Walsh et al., 1978; Olaizola et al.,
582 1993; Falkowski et al., 1998) and most likely also.

583 However, satellite measurements indicate a global decline in diatoms between 1998 and 2012
584 within the surface mixed layers (Rousseaux and Gregg, 2015). This decline is critical since diatoms
585 contribute approximately 40% to marine primary productivity and serve as a primary source of
586 carbon export to the deep ocean (Tréguer et al., 2018). Consequently, an increase in nutrients from
587 for example olivine in targeted regions, driven by the aforementioned physical processes, can
588 enhance CO₂ export to the deeper ocean by promoting diatom growth. In contrast, nutrient-poor
589 waters, mainly populated with small primary producers like picocyanobacteria, particularly in
590 central ocean gyres (Falkowski et al., 1998), result in minimal carbon escape from remineralization
591 (Azam, 1998).

592 With this, satellite data analysis can aid in identifying regions of interest (Righetti et al., 2019).
593 Specifically, satellite imagery has proven instrumental in identifying four major groups of
594 phytoplankton (Kramer and Siegel, 2019). Furthermore, these satellite-derived measurements can
595 contribute to understanding net primary production in oceanic regions.

596 (Longhurst et al., 1995; Behrenfeld and Falkowski, 1997). Combined with our dataset, this is
597 essential knowledge when considering additions olivine in coastal or open oceans. Based on our
598 findings, we suggest considering locations such as the Labrador Sea, the North Atlantic, and the
599 Eastern China Sea to explore OAE practices further. The following steps of OAE research should
600 explore the underlying mechanisms of growth responses in phytoplankton, extend to ecosystem
601 levels, and identify potential synergies and trade-offs between different alkaline minerals and
602 primary producers to inform the progress of sustainable OAE strategies for marine ecosystems
603 (Eisaman et al., 2023).

604 **5. Conclusion**

605 We aimed to assess the influence of olivine additions on various cosmopolitan primary producers,
606 considering a moderate range of solid-liquid ratios. Our data suggests that olivine additions did
607 not lead to growth inhibition across cultivated cultures, with the picoplankton benefiting the most.

608 Our findings further indicate that Ni exposure from olivine additions did not inhibit phytoplankton
609 growth in our cultures. While our study provides valuable insights into the responses of certain



610 cosmopolitan primary producers to olivine additions, it is important to recognize its limitations.
611 Our study focused on phytoplankton responses in controlled laboratory conditions, which may not
612 fully capture the complexities of natural ecosystems. Future research should aim to incorporate a
613 community or ecosystem-level perspective to better understand the broader impacts of olivine
614 additions in natural settings.

615 Additionally, it is essential for future studies to address potential pitfalls and concerns associated
616 with our experimental setup, particularly regarding considerations of carbonate chemistry. By
617 addressing these limitations, future research can provide more robust and comprehensive insights
618 into the effects of olivine additions on marine ecosystems. **Data availability:** Data are presented in the
619 supplement.

620 **Author contribution:** JR and CL designed the study. JR, ZK and MJ carried out laboratory
621 experiments and analyses. JR analyzed the data and wrote the manuscript together with CRL, and
622 input of all co-authors.

623 **Competing interests:** At least one of the (co-)authors is a member of the editorial board of
624 Biogeosciences.

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633



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