Modeling 2D gravity-driven flow in unsaturated porous media for different infiltration rates

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Abstract. The gravity-driven flow in an unsaturated porous medium remains one of the most important unsolved problems in multiphase flow. Sometimes a diffusion-like flow, known as stable flow, with a uniform wetting front is observed, but at other times, it is unstable with distinct preferential pathways. The formation of an unstable wetting front in a porous medium depends on many factors, including the type of the porous medium, the initial saturation, and the applied infiltration rate. As the infiltration rate increases, the wetting front first transitions from stable to unstable at low infiltration rates and then from unstable to stable at high infiltration rates. We propose a governing equation and its discretized form, the semi-continuum model, to describe this significant non-monotonic transition. We show that the semi-continuum model is able to capture the influx dependence together with the correct finger width and spacing. Moreover, we demonstrate that the instability of the wetting front is closely related to the saturation overshoot in one dimension. Finally, we show that the flow can still be preferential even when the porous medium is completely wetted.

1 Introduction

In the field of hydrology, the gravity-driven multiphase flow in porous media, typically involving the flow of water into soil, remains a long-standing and unsolved problem. In the early 1980s, a series of seminal papers provided important insights into this phenomenon (Diment et al., 1982; Diment and Watson, 1983, 1985). A decade later, the influence of soil matrix hydrophobicity on water movement was described (Dekker and Ritsema, 1994). A detailed overview of the research on porous media flow up to the year 2000 is presented in de Rooij (2000).

The infiltration of water into soil is an extremely complicated physical phenomenon that exhibits two kinds of flow behavior: diffusion-like and finger-like flow (de Rooij, 2000; DiCarlo, 2013; Xiong, 2014). In the case of diffusion-like flow, a stable wetting front is observed and is referred to as stable flow. Conversely, finger-like flow, characterized by an unstable wetting front, is called unstable flow. The lack of a well physically-based and experimentally verified model of soil water movement is a major obstacle to the development of rainfall-runoff models at the hydrological scale. This issue has received consistent attention for many years, as evidenced by the recommendation that rainfall-runoff models should include a robust model of soil water movement (Kutílek and Nielsen, 1994). Traditionally, the standard concept of diffusion-like flow based on ~~the~~

Richards' equation (Richards, 1931) has been used for modeling water movement in soil. However, this concept is inadequate for modeling finger-like flow (DiCarlo, 2013). Consequently, there is an ongoing search for a concept that can describe both types of flow, given its substantial application potential in soil science (Lake, 1989; Bundt et al., 2000; DiCarlo, 2013; Xiong, 2014) and other fields (Sutherland and Chase, 2008; Vafai, 2011). This paper is devoted to the description of such a model.

The instability of the wetting front is accompanied by preferential flow, where most of the water moves through the preferential pathways, leaving much of the porous medium dry even after hours of uniform infiltration. This type of flow is characterized by the formation of fingers (DiCarlo, 2013). A finger consists of two parts: an undersaturated tail and an oversaturated tip, and this non-monotonicity of saturation is known as saturation overshoot. Therefore, wetting front instability and associated saturation overshoot have been at the center of attention for several decades (Saffman and Taylor, 1958; Chuoke et al., 1959; Smith, 1967; Hill and Parlange, 1972). Since then, a huge number of laboratory and field experimental works have become available. Some of the works concern 3D experiments (Glass et al., 1990; Yao and Hendrickx, 1996), but most are performed in 1D (long vertical tubes) (DiCarlo, 2004, 2007, 2010; Aminzadeh and DiCarlo, 2010) and in 2D (Hele-Shaw cells) (Smith, 1967; Glass et al., 1988, 1989a, b, c; Liu et al., 1994; DiCarlo et al., 1999; Glass et al., 2000; Bauters et al., 2000; Sililo and Tellam, 2005; Rezanezhad et al., 2006; Wei et al., 2014; Cremer et al., 2017; Pales et al., 2018; Chen et al., 2022; Liu et al., 2023) due to simpler realization.

It turns out that flow in an unsaturated porous medium has many unexpected features. For example, a non-monotonic dependence of the wetting front velocity and finger width on the initial saturation is observed (Bauters et al., 2000). At lower initial saturation, the wetting front is unstable with slow and wide fingers. With increasing initial saturation, the fingers first narrow and speed up and then slow and widen again until a stable wetting front is observed. Another non-intuitive behavior is the dependence on applied influx (Glass et al., 1989c; Yao and Hendrickx, 1996; DiCarlo, 2013), which is crucial for understanding the precipitation-runoff relationship in hydrology. Glass et al. (1989c) ~~infiltrated into~~ built a two-dimensional chamber with a thickness of $1\,\mathrm{cm}$ filled with a homogeneous porous medium. Water was uniformly applied at a constant flux q_{top} at the top boundary. They observed a stable wetting front when the flux was close to the saturated conductivity, but with decreasing applied influx, the flow tended to become significantly preferential. Yao and Hendrickx (1996) performed similar experiments, but they infiltrated water into large three-dimensional columns with diameters of $30\,\mathrm{cm}$ and $100\,\mathrm{cm}$, and the applied flux was much lower. They demonstrated that as the flux decreases, the finger-like flow disappears and a stable wetting front reappears. This behavior was maintained regardless of the type of homogeneous sand used. Therefore, preferential flow is observed only within a certain range of infiltration flux. Note that while homogeneous soil does not actually exist, as there is always some level of heterogeneity, this term commonly refers to soils, such as sands, where the characteristics of soil hydraulic properties appear uniform from a macroscopic perspective.

Approaches to modeling unsaturated porous media flow can be divided into two categories: (1) microscale models, which are developed for small-scale where the Darcy-Buckingham law (Buckingham, 1907) can be applied, and (2) macroscale models, which focus directly on large-scale where the Darcy-Buckingham law is not applicable. Macroscale models, such as the ARM model (Liu et al., 2005; Liu, 2022), cannot capture individual fingers because the computational grid is too coarse, so multiple fingers are included within each computational element. In contrast, microscale models do not have this limitation but are

very computationally intensive. A typical example of a microscale model is ~~the~~ Richards' equation, which combines the mass balance law and the Darcy-Buckingham law. Richards' equation is diffusive in nature, as it is unable to model a non-monotonic saturation profile in the case of uniform infiltration rate with a smooth and non-decreasing retention curve (Fürst et al., 2009). Therefore, many extensions of ~~the~~ Richards' equation, known as continuum models, have been proposed (Hassanizadeh et al., 2002; Eliassi and Glass, 2002; Brindt and Wallach, 2020; Cueto-Felgueroso et al., 2020; Beljadid et al., 2020; Roche et al., 2021; Ommi et al., 2022a, b). Other approaches include discrete (pore-scale) models (Lenormand et al., 1988; Primkulov et al., 2018; Wei et al., 2022) and combinations of discrete and continuum approaches (Glass and Yarrington, 1989, 2003; Vodák et al., 2022).

Any microscale model designed primarily for small-scale processes should be first thoroughly validated using small-scale experiments. The opposite approach, i.e., validating with in-field experiments without proper small-scale validation, can lead to fundamentally incorrect conclusions. For instance, Tesař et al. (2004) developed a model based on upscaling Richards' equation. The model was validated using outflow data from the Liz catchment, covering an area of $1\,\mathrm{km}^2$ during the vegetation season of 1999. However, subsequent simulations showed that water always flowed through the entire porous medium regardless of initial and boundary conditions, revealing the model's inaccuracies. Therefore, the validation of microscale models should be focused on small-scale experiments, such as those performed in the laboratory, which provide a detailed analysis of flow behavior. Well-defined boundary and initial conditions in laboratory experiments enable a thorough examination of simulation details, allowing to determine whether the model accurately captures preferential and diffusion flow. This makes the validation of the model more thorough and reliable. Following rigorous validation, one can then transition to more complex in-field experiments.

A validation procedure on a small-scale through detailed laboratory experiments was proposed by DiCarlo (2013). Specifically, the author suggested criteria for evaluating which model is the "most appropriate" (see section 6.6 in DiCarlo (2013)). According to the author, the model should have a minimum of adjustable parameters, be capable of producing the diffusive character of the flow as seen in ~~the~~ Richards' equation, match observed 1D profiles well, and finally, be able to predict 2D and 3D preferential flow in terms of finger widths and finger spacings. Our aim is to demonstrate that the semi-continuum model proposed by Vodák et al. (2022) succeeds in this type of evaluation. The authors developed the semi-continuum model and its formal limit in the form of a partial differential equation with a Prandtl-type hysteresis operator (Visintin, 1993) under the derivative. It was shown that the semi-continuum model was able to correctly reproduce experiments of flow into a long vertical tube (Kmec et al., 2019). In Kmec et al. (2021), the model was used to replicate the transition between unstable and stable wetting fronts for increasing initial saturation. Along with this, the model was shown to correctly capture the finger persistence and the flow across a heterogeneous porous medium. Finally, the strong non-monotonic dependence of the wetting front on the initial saturation for a point source infiltration was captured well (Kmec et al., 2023). According to the suggested model evaluation, the semi-continuum model has successfully addressed all aspects except for predictions of 2D and 3D preferential flow. The model has already reproduced the dependency on initial saturation in 2D (Kmec et al., 2021, 2023). However, the final aspect required for full validation is to accurately capture the dependence on the infiltration rate in 2D/3D.

There are a few models that are able to capture the transition between diffusion and finger-like flow for various infiltration rates in 1D, but analyzing the 2D case is much more complex. Although the 2D experiments are over 30 years old (Glass et al., 1989c; Yao and Hendrickx, 1996), they have not yet been successfully reproduced by any model. Finding a model capable of simulating this complex transition thus remains a major challenge (DiCarlo, 2013). One of the most promising and recent attempts to simulate the dependence on applied influx in 2D was proposed by Beljadid et al. (2020), who introduced a nonlocal model endowed with an entropy function. This model extends ~~the~~ Richards' equation by incorporating a fourth-order spatial derivative of saturation. The authors demonstrated the transition from finger-like to diffusion-like flow under high applied fluxes. Moreover, the model showed good agreement with experimentally measured finger widths for large fluxes (Glass et al., 1989c). However, inconsistencies arise at very low fluxes, where the wetting front does not stabilize as expected from experiments (Yao and Hendrickx, 1996). In contrast, decreasing the infiltration rate results in thinner fingers, maintaining the preferential character of the flow. Therefore, matching experiments in 1D does not guarantee a match in 2D/3D (Cueto-Felgueroso and Juanes, 2009; Beljadid et al., 2020).

In this paper, we aim to demonstrate the ability of the semi-continuum model to accurately capture the 2D transition from stable to unstable flow for low infiltration fluxes (Yao and Hendrickx, 1996) and the transition from unstable to stable flow for high infiltration fluxes (Glass et al., 1989c), along with making correct predictions of preferential flow in terms of experimentally measured finger widths and spacings. In addition, we will investigate the relationship between the saturation overshoot in 1D and the wetting front instability in 2D to demonstrate further agreement with experimental observations.

2 Retention curve and its sample size dependence

In soil physics, the relationship between saturation S and pressure P exhibits strong hysteresis and is known as the retention curve. The retention curve consists of two main branches: the wetting and draining branches. Various approaches have been proposed to model the hysteresis, including (Mualem, 1976; Lenhard and Parker, 1987; Parker and Lenhard, 1987; Beliaev and Hassanizadeh, 2001; McNamara, 2014; Schweizer, 2017; Abreu et al., 2019).

The shape of the retention curve strongly depends on the size of the sample on which the measurement is performed (Larson and Morrow, 1981; Mishra and Sharma, 1988; Zhou and Stenby, 1993; Perfect et al., 2004; Hunt et al., 2013; Ghanbarian et al., 2015; Silva et al., 2018). As the sample size decreases, the pore size variability within the sample also decreases, and the retention curve becomes flatter (Silva et al., 2018). Therefore, it is useful to examine the influence of pore size variability on the shape of the retention curve. This influence has been studied by Pražák et al. (1999). They demonstrated that the main draining branch becomes flatter as the pore size variability decreases, which is consistent with flattening the retention curve as sample size decreases.

Although it is well known that the retention curve is dependent on the sample size of the porous medium (Ghanbarian et al., 2015), the implementation of this dependence is not common in flow modeling. Moreover, other characteristics of the porous medium, such as permeability and porosity, are also dependent on the sample size (Mishra and Sharma, 1988; Ewing et al., 2010; Ghanbarian et al., 2017, 2021; Esmaeilpour et al., 2021). This sample size dependence can be integrated into

models by transferring this dependency to the individual elements or volumes of the discretization mesh. This means that these elements or volumes represent a real sample of the porous medium and carry information about its physical characteristics. This approach fundamentally differs from standard numerical schemes for partial differential equations, where the mesh serves only a mathematical role and ignores the fact that individual elements/volumes represent the real domain. However, some authors have already considered this aspect in modeling porous media. For instance, White et al. (2006) estimated a lower limit of finite elements and then used this size in their model. They argue that the use of smaller elements would not be appropriate because it would lead to a violation of the continuum assumptions. Note that in the semi-continuum model, the sample size dependence of the retention curve is not neglected, as it is a crucial feature of the model. This is further discussed in Sect. 3.4.

3 Methods

In this section, we first introduce the Prandtl-type hysteresis operator and the resulting governing partial differential equation, which was derived as a formal limit of the semi-continuum model (Vodák et al., 2022). We then present a proper discretization of the governing equation along with the discretization of the Prandtl-type hysteresis operator to provide a detailed description of the semi-continuum model. This model describes the movement of the wetting liquid in a porous medium; specifically, it is a multiphase flow model used for modeling unsaturated porous media flow.

3.1 Prandtl-type hysteresis operator

The Prandtl-type hysteresis operator P_H [Pa] is the pressure operator defined by the following differential inequality (Visintin, 1993):

$$(K_{PS}\partial_t S - \partial_t P_H)(P_H - v) \geq 0, \quad \forall v \in [C_2, C_1], \; P_H \in [C_2, C_1], \tag{1}$$

where S [−] denotes the saturation of the wetting phase and K_{PS} [Pa], C_1 [Pa], and C_2 [Pa] are constants. A detailed description of this operator's characteristics can be found in Visintin (1993), specifically on page 16.

Figure 1 illustrates the form of the Prandtl-type hysteresis operator for $K_{PS} = 10^5 \, \text{Pa}$. Hysteresis implies that the pressure value P depends not only on the current saturation S but also on its history. The Prandtl operator, described by Eq. (1), defines how the pressure value is assigned as a function of saturation. Assume the initial pressure value is C_1. When the saturation is non-decreasing over time, the pressure value remains at C_1. Otherwise, the pressure value "jumps" from C_1 to C_2 over time, with the rate of this transition determined by K_{PS}. The same principle applies when the pressure value is C_2. The value remains at C_2 when saturation is non-increasing; otherwise, it jumps from C_2 to C_1. These transitions, represented by non-vertical scanning curves, are marked in black in Fig. 1. The value K_{PS} indicates a large gradient of these scanning curves, causing the pressure to change significantly between values C_1 and C_2 with a negligible change in saturation.

Although the behavior of the Prandtl operator is relatively straightforward, its mathematical description is not trivial, as the community is usually not familiar with differential inequalities. For more details, we refer to Vodák et al. (2022), specifically its supplement, where we have demonstrated that the mathematical description follows the explanation of the Prandtl operator

provided above. We have also shown that the Prandtl operator can be generalized using non-decreasing functions instead of constants C_1 and C_2.
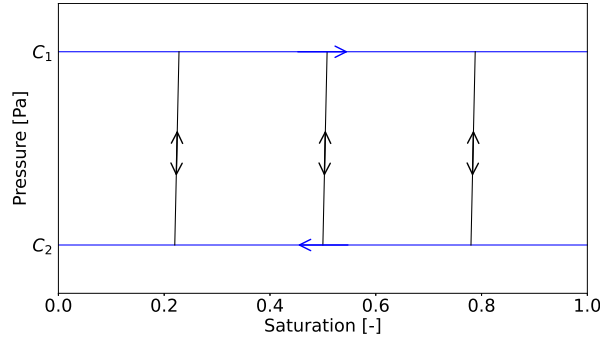
Figure 1. Prandtl-type hysteresis operator P_H. Blue lines denote constants C_1 and C_2, and black lines are non-vertical scanning curves with large gradient $K_{PS} = 10^5 \, \text{Pa}$.

3.2 Governing equation

The governing equation is given by Eq. (2). It is a partial differential equation containing the Prandtl-type hysteresis operator P_H (Fig. 1) under the spatial derivative.

$$\theta \partial_t S + \text{div} \left[\frac{\kappa}{\mu} \sqrt{k(S^-)} \sqrt{k(S^+)} \left(\rho \boldsymbol{g} - \nabla P_H \right) \right] = 0, \quad S^\pm(x_0, t) = \lim_{x \to x_0^\pm} S(x, t). \tag{2}$$

In this equation, the porous medium is characterized by its porosity $\theta \, [-]$, intrinsic permeability $\kappa \, [\text{m}^2]$, and relative permeability $k(S) \, [-]$. The wetting phase (liquid) is characterized by its saturation $S \, [-]$, density $\rho \, [\text{kg m}^{-3}]$, dynamic viscosity $\mu \, [\text{Pa s}]$, and pressure $P \, [\text{Pa}]$ defined by operator P_H. In a porous material that is not completely filled with liquid, the pressure P represents the capillary pressure, which is the tensile stress by which the liquid is held in pores. This pressure P in the liquid phase is ~~less~~ lower than the pressure in the non-wetting phase (gas), which ~~is assumed to be zero, and therefore~~ serves as the reference for expressing the liquid pressure. As a result, the pressure P becomes negative. The vector $\boldsymbol{g} = (0, 0, g)$, where $g \, [\text{m s}^{-2}]$ denotes the acceleration due to gravity, thus the gravity acts only in the third dimension. Note that $k(S^-)$ and $k(S^+)$ denote the left and right limits of relative permeability in the spatial variable. The discontinuity in saturation arises from the use of a geometric mean conductivity (Vodák et al., 2022). In the case of continuous saturation, $S^- = S^+$, which results in $\sqrt{k(S^-)} \sqrt{k(S^+)} = k(S)$.

The Prandtl-type hysteresis operator causes Eq. (2) to switch between parabolic and hyperbolic types in the case of an unsaturated porous medium. When the pressure value is defined by values C_1 or C_2 (blue lines in Fig. 1), the pressure-saturation relation is constant, resulting in $\nabla P_H = 0$. Thus, the equation becomes a hyperbolic differential equation. Otherwise, when the pressure is given by the scanning curves (black lines in Fig. 1), the equation is a parabolic differential equation.

It is well known that different types of flow, saturated and unsaturated, can occur in parallel in the porous medium (Brandhorst et al., 2021). In a fully saturated medium, the pressure-saturation relation is no longer defined by the Prandtl-type hysteresis operator P_H. In this case, the pressure becomes hydrostatic pressure and takes on positive values. Using the hydrostatic pressure in Eq. (2) instead of P_H, we obtain Laplace's equation as $k(S) = 1$ and $\partial_t S = 0$. Since we focus on unsaturated flow, hydrostatic pressure is not implemented and the case of a fully saturated medium is not studied further.

3.3 Discretization of the porous medium

We want to simulate experiments in a two dimensional Hele-Shaw cell of a porous medium; hence, a 2D discretization is used. The porous medium is a rectangle of size $A \times B$, where A and B denote the horizontal and vertical widths of the porous medium, respectively. The porous medium is represented by a square mesh consisting of $N \times M$ blocks (finite volumes) of size $\Delta x \times \Delta x$. These blocks retain the character of the porous medium.

3.4 Discretization of the Prandtl-type hysteresis operator

~~In soil physics, the relationship between saturation S and pressure P exhibits strong hysteresis and is known as the retention curve. The retention curve consists of two main branches: the wetting and draining branches. The shape of the retention curve strongly depends on the size of the sample on which the measurement is performed (Larson and Morrow, 1981; Larson and Morrow, 1988; As the sample size decreases, the pore size variability within the sample also decreases, and the retention curve becomes flatter, which was experimentally confirmed in Silva et al. (2018). This sample size dependence of the retention curve is taken into account in the discretization of the Prandtl-type hysteresis operator given by Eq. (1).~~

The discretization of the Prandtl-type hysteresis operator given by Eq. (1) has already been described in Kmec et al. (2023). Its discretized version is the capillary pressure operator $P(S)$, which satisfies $P(S) \to P_H(S)$ as $\Delta x \to 0$. The well-known fact that the shape of the retention curve depends on the sample size (Ghanbarian et al., 2015; Silva et al., 2018) is incorporated into our model so that the retention curve depends on the block size Δx. We refer to this discretization as the scaling of the retention curve. The proposed scaling is explained below; however, for a detailed mathematical and physical justification, we refer to Vodák et al. (2022). For the reference block size Δx_0, the reference retention curve is given by the van Genuchten equation (Genuchten, 1980):

$$P_0^w(S) = -\frac{1}{\alpha_w}\Big(S^{\frac{n_w}{1-n_w}} - 1\Big)^{\frac{1}{n_w}}, \qquad P_0^d(S) = -\frac{1}{\alpha_d}\Big(S^{\frac{n_d}{1-n_d}} - 1\Big)^{\frac{1}{n_d}}, \tag{3}$$

where P_0^w is the main wetting branch, P_0^d is the main draining branch, α_w, n_w are parameters of the main wetting branch, and α_d, n_d are parameters of the main draining branch. For a block size $\Delta x < \Delta x_0$, the main wetting and draining branches are scaled as follows:

$$P^w(S, \Delta x) = \frac{\Delta x}{\Delta x_0} P_0^w(S) + P_0^w(0.5)\Big(1 - \frac{\Delta x}{\Delta x_0}\Big), \qquad P^d(S, \Delta x) = \frac{\Delta x}{\Delta x_0} P_0^d(S) + P_0^d(0.5)\Big(1 - \frac{\Delta x}{\Delta x_0}\Big). \tag{4}$$

Obviously, for $\Delta x = \Delta x_0$, the retention curve is given by Eq. (3). For $\Delta x \to 0$, the retention curve converges to the Prandtl-type hysteresis operator P_H so that $C_1 = P_0^w(0.5)$ and $C_2 = P_0^d(0.5)$. Constants C_1 and C_2 in the semi-continuum model represent the water entry and air entry values, respectively (Vodák et al., 2022). Instead of the midpoint $S = 0.5$ in Eq. (4), it

210   is possible to choose, for example, an inflection point of the main branches. However, the effect on the results is negligible because the flux is calculated relative to the pressure gradient. Note that the reference block size Δx_0 is a parameter of the semi-continuum model and its determination can be obtained, for example, through calibration experiments.

Figure 2 shows the capillary pressure operator $P(S)$ for different block sizes. It can be clearly seen that for $\Delta x \to 0$, the operator $P(S)$ converges to the Prandtl-type hysteresis operator P_H shown in Fig. 1. For non-zero but very small Δx (see

215   black lines in Fig. 2), both main branches take the form of step-like functions that are almost constant for $S \in (0, 1)$. Since decreasing the block size reduces pore size variability, the proposed linear scaling defined by Eq. (4) perfectly aligns with the findings of Pražák et al. (1999), where the authors demonstrated a step-like form of the main draining branch for a hypothetical porous medium without pore size variability.
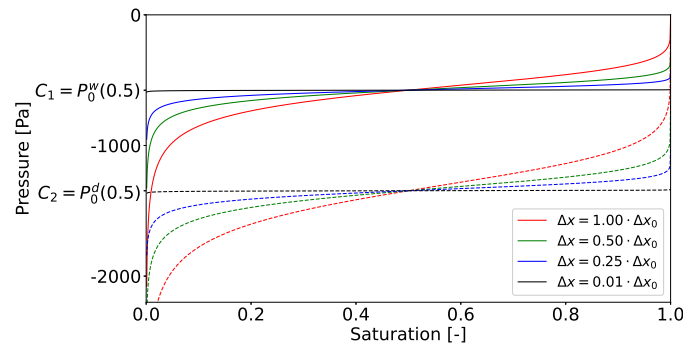
Figure 2. The scaling of the retention curve for different block sizes Δx. The solid lines denote the main wetting branches and the dashed lines denote the main draining branches. The parameters $\alpha_w, n_w, \alpha_d, n_d$ are given in Table 1. For $\Delta x \to 0$, the retention curve converges to the Prandtl-type hysteresis operator P_H, as both main branches take the form of horizontal lines C_1 and C_2.

220   ~~Decreasing the block size results in a decrease of the pore size variability. Therefore, it is useful to examine the influence of pore size variability on the shape of the retention curve. Such influence has been studied by Pražák et al. (1999). They demonstrated that the main draining branch becomes flatter as the pore size variability decreases. Moreover, they showed that for a hypothetical porous medium with identical pores (i.e., without pore size variability), the main draining branch has a step-like form such that it is constant for $S \in (0, 1)$. This perfectly aligns with the proposed scaling of the retention curve defined by Eq. (4). For non-zero, but very small Δx (see black lines in Fig. 2), both main branches take the form of step-like~~

225   ~~functions that are almost constant for $S \in (0, 1)$.~~

Note that all scanning curves are non-vertical straight lines, as illustrated in Fig. 1. This approach is similar to the play-type hysteresis (Schweizer, 2017). ~~There are~~ Various approaches to modeling the hysteresis between saturation and pressure ~~(Mualem, 1976; Mualem, 1987; Mualem, 1987; Mualem, 2001; Mualem, 2014; Mualem, 2019), and these methods~~ could po-

tentially be implemented in our model. However, using a more complex hysteresis model is not beneficial, as we have achieved
good agreement with experiments using the simpler model.

~~Although it is well known that the retention curve is dependent on the sample size of the porous medium (Ghanbarian et al., 2015), the implementation of this dependence is not common in flow modeling. Moreover, other characteristics of the porous medium, such as permeability and porosity, are also dependent on the sample size (Mishra and Sharma, 1988; Mishra and Sharma, 2010; Mishra and In the semi-continuum model, the sample size dependence of the retention curve is implemented by a linear scaling of the retention curve. This means that individual blocks of the discretization mesh represent a real sample of the porous medium, and these blocks carry information about the physical characteristics. This approach fundamentally differs from standard numerical schemes for partial differential equations, where the mesh serves only a mathematical role and ignores the fact that individual elements/volumes represent the real domain. However, some authors have already considered this aspect in modeling porous media. For instance, White et al. (2006) estimated a lower limit of finite elements and then used this size in their model. They argue that the use of smaller elements would not be appropriate because it would lead to a violation of the continuum assumptions.~~

3.5 Discretization of the governing equation – the semi-continuum model

Each block of the discretized porous medium is denoted by indices (i,j) representing the corresponding row and column. $S_t(i,j)\,[-]$ and $P_t(i,j)\,[\mathrm{Pa}]$ represent saturation and pressure of the wetting phase (liquid) within block (i,j) at time t, respectively. These values are assumed to be constant within each block. Moreover, $q_{t\,(i_2,j_2)}^{(i_1,j_1)}\,[\mathrm{m\,s^{-1}}]$ denotes the flux of the wetting phase from block (i_1,j_1) to block (i_2,j_2) at time t.

The semi-continuum model, which is the discretization of the governing Eq. (2), consists of three consecutive steps: saturation update, pressure update, and flux update. First, the saturation in each block is updated according to the discretized mass balance law for a given time step Δt and the block size Δx:

$$S_{t+\Delta t}(i,j) = S_t(i,j) + \frac{\Delta t}{\theta}\frac{1}{\Delta x}\left(q_{t\,(i,j)}^{(i-1,j)} - q_{t\,(i+1,j)}^{(i,j)} + q_{t\,(i,j)}^{(i,j-1)} - q_{t\,(i,j+1)}^{(i,j)}\right). \tag{5}$$

The second step is to update the pressure in each block to obtain the pressure at time $t+\Delta t$, i.e., $P_{t+\Delta t}$. The pressure is updated according to the hysteretic capillary pressure operator $P(S)$, whose main wetting and draining branches are given by Eq. (4) and scanning curves are illustrated in Fig. 1.

The third and final step is the flux update. We adopt the following form of the relative permeability function $k(S)$ (Genuchten, 1980):

$$k(S) = S^\lambda \left[1 - \left(1 - S^{\frac{n}{n-1}}\right)^{\frac{n-1}{n}}\right]^2, \tag{6}$$

where $\lambda\,[-]$ is a free parameter and $n\,[-]$ is a parameter of the retention curve given by Eq. (3). To simplify the notation, we introduce the so-called effective permeability as $\gamma(S) = \kappa k(S)$. The flux between blocks is updated using the discretized

version of Darcy-Buckingham law (Bear, 1972):

$$260 \quad q_{t+\Delta t\,(i_2,j_2)}^{(i_1,j_1)} = \begin{cases} \frac{1}{\mu}\sqrt{\gamma(S_{t+\Delta t}(i_1,j_1))\gamma(S_{t+\Delta t}(i_2,j_2))}\left(\rho g - \frac{P_{t+\Delta t}(i_2,j_2)-P_{t+\Delta t}(i_1,j_1)}{\Delta x}\right), & \text{for } j_1 = j_2,\ i_2 = i_1+1 \\ \frac{1}{\mu}\sqrt{\gamma(S_{t+\Delta t}(i_1,j_1))\gamma(S_{t+\Delta t}(i_2,j_2))}\left(0 - \frac{P_{t+\Delta t}(i_2,j_2)-P_{t+\Delta t}(i_1,j_1)}{\Delta x}\right), & \text{for } i_1 = i_2,\ j_2 = j_1+1 \\ 0, & \text{otherwise} \end{cases} \quad (7)$$

The acceleration due to gravity is included only for the vertical fluxes. Unsurprisingly, the fluxes between non-neighboring blocks are set to zero. The geometric mean is used to average the effective permeability between blocks. This aligns with the approach utilized in governing Eq. (2), where the geometric mean is also used. Moreover, this type of averaging is also consistent with the findings of Jang et al. (2011). After updating the fluxes between neighboring blocks, we update the time
265 $t = t + \Delta t$ and return to the saturation update given by Eq. (5).

If the fluxes between blocks are too large, especially if the flux is close to the saturated conductivity $K_S = \frac{\kappa}{\mu}\rho g$, the saturation may exceed one. This occurs due to the inherent nature of the semi-continuum model and is often the case for other models as well. For example, in Cueto-Felgueroso and Juanes (2009), a "compressibility term" is used for the capillary energy-saturation dependence. This term becomes dominant near saturation close to one, so it prevents the saturation from increasing any further.
270 We use a different approach; the magnitude of the flux to the block can be at most so large that the saturation does not exceed one. This straightforward approach is only possible because of the simple numerical scheme used.

According to Eq. (5) and Eq. (7), if a standard retention curve without scaling is used, i.e., it does not converge to a Prandtl-type hysteresis operator P_H, the semi-continuum model degenerates into a numerical scheme for solving the classical Richards' equation. The crucial difference in our model is that the shape of the retention curve depends on the block size Δx. From a
275 mathematical point of view, in the case of an unsaturated porous medium, ~~the~~ Richards' equation is diffusive in nature, as it is a parabolic differential equation (DiCarlo, 2013). This is not the case for the governing Eq. (2), which is a hyperbolic-parabolic differential equation.

4 Results

4.1 Numerical setup

280 The parameters used for simulations of the infiltration dependence experiments are given in Table 1. The porous medium used for simulations is 20/30 sand. The parameter $\lambda = 0.8$, which is consistent with measurements (Schaap and Leij, 2000). It is important to emphasize that we use the same parameters as in Kmec et al. (2023), where the semi-continuum model accurately reproduced the experiments reported in Bauters et al. (2000). This includes a parameter of the semi-continuum model, the reference block size $\Delta x_0 = \frac{10}{12}$, which was calibrated for 20/30 sand in Kmec et al. (2023) using the experiments of Bauters
285 et al. (2000). This decision was made to demonstrate that the semi-continuum model can simulate different flow phenomena without additional parameter adjustments. Hence, our aim is not to optimize the parameters to achieve the best agreement with experiments. Let us note that the slope of the scanning curves K_{PS} does not affect the results if it is chosen to be large enough.

The differences between solutions are negligible for $K_{PS} \geq 10^5\,\mathrm{Pa}$, hence we set K_{PS} to this value. ~~Here, the lower limit for K_{PS} is used.~~

Table 1. Parameters used for reproducing the flow dependence on different infiltration rates. Parameters for 20/30 sand were adopted from Schroth et al. (1996) and DiCarlo (2004).

Parameter	Symbol	Value
Horizontal width of the chamber	A	50 cm
Vertical width of the chamber	B	50 cm
Reference block size	Δx_0	$\frac{10}{12}$ cm
Block size	Δx	0.25 cm
Porosity	θ	0.35
Density of water	ρ	1000 kg m^{-3}
Dynamic viscosity of water	μ	9×10^{-4} Pa s
Intrinsic permeability	κ	2.294×10^{-10} m^2
Relative permeability exponent	λ	0.8
Acceleration due to gravity	g	9.81 m s^{-2}
Wetting curve parameter	α_w	0.177 cm^{-1}
Wetting curve parameter	n_w	6.23
Draining curve parameter	α_d	0.0744 cm^{-1}
Draining curve parameter	n_d	8.47
Slope of scanning curves	K_{PS}	10^5 Pa
Initial saturation	S_{in}	0.01
Residual saturation	S_{rs}	0.05

290   The scheme of the numerical setup is shown in the left panel of Fig. 3. Initial and boundary conditions are set to be consistent with the experiments we aim to reproduce (Yao and Hendrickx, 1996; Glass et al., 1989c). The porous medium is initially dry with an initial saturation $S_{in} = 0.01$, and all the blocks begin on the main wetting branch. A constant infiltration rate q_{top} is applied to the entire top boundary. A total of 18 different infiltration rates q_{top} are used, with the lowest influx equal to $0.001\,\mathrm{cm\,min^{-1}}$, and the highest influx equal to the saturated conductivity $K_S = 15\,\mathrm{cm\,min^{-1}}$. The lateral boundaries of the

295   porous medium are impenetrable, so the lateral fluxes are set to zero. For the bottom boundary flux q_{bot}, the outflow of water into the air is prescribed. The following implementation is ~~used~~ applied so that it does not affect the flow above the bottom boundary. This allows us to isolate the behavior of the model from the influence of the bottom boundary condition.

$$q_{\text{bot}} := q_{t\,(\text{out})}^{(\text{N},j)} = \begin{cases} 0 & \text{for } S_t \leq S_{rs} \\ \frac{1}{\mu}\gamma(S_t(\text{N},j))\left(\rho g + \frac{P_t(\text{N},j)}{\Delta x}\right), & j = 1,\ldots,\text{M}, \quad \text{for } S_t > S_{rs} \end{cases}, \tag{8}$$

where N denotes the bottom row index. Thus, the flux from the bottom boundary is set to zero if the saturation of the corresponding block does not exceed the residual saturation S_{rs}; otherwise, it is non-zero. Residual saturation refers to the maximum amount of water that the porous medium can retain against the force of the gravity. The value of $S_{rs} = 0.05$ corresponds to the experimentally measured residual saturation for 20/30 sand in Bauters et al. (2000). This implementation of bottom boundary flux is similar to a free discharge (Šimůnek and Suarez, 1994) and has already been used for the semi-continuum model in Kmec et al. (2021).
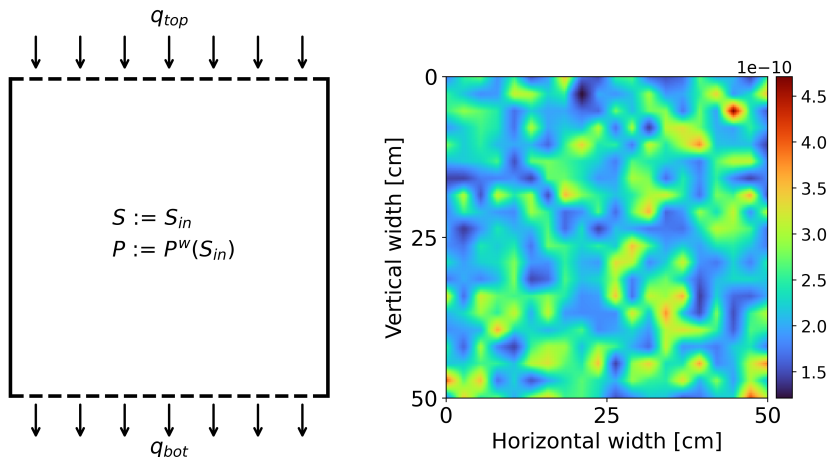
Figure 3. Left panel: The scheme of the numerical setup. Right panel: The distribution of spatially correlated intrinsic permeability. The average value of κ equals approximately $2.294 \times 10^{-10}\,\mathrm{m^2}$ and the distribution satisfies $\kappa_{\max}/\kappa_{\min} \approx 4$. The values are colored according to the color bar on the right.

The measurement of macroscopic properties of a homogeneous porous medium, such as permeability or porosity, is typically performed by averaging microscopic quantities over the domain (White et al., 2006; Ghanbarian et al., 2021). Therefore, to ~~To~~ obtain a more realistic description of a porous medium, such as 20/30 sand, it is convenient to slightly perturb its characteristics. ~~Therefore, a small distribution of~~ In this study, we employ a slightly distributed and spatially correlated intrinsic permeability ~~is included~~ (see the right panel of Fig. 3). To achieve this, we first generate a normally distributed array for larger blocks (2.5 cm). This array is then zoomed on smaller blocks of size Δx using linear interpolation, which ensures a spatially correlated distribution. The resulting values represent a multiplication factor for the intrinsic permeability, adjusted so that the probability of being n times smaller is the same as being n times larger. The specific implementation can be found in Kmec (2023a). A similar distribution has been used in Cueto-Felgueroso and Juanes (2009); Gomez et al. (2013); Kmec et al. (2023). The effect of intrinsic permeability distribution is further discussed in Sect. 5.4.

4.2 Evolution of the saturation profile

Figure 4 shows the evolution of the saturation profile at six different times for $q_{\mathrm{top}} = 0.05\,\mathrm{cm\,min^{-1}}$. Times are displayed in the upper left corner for each frame. Initially, a stable wetting front with small frontal perturbations develops at $t = 300\,\mathrm{s}$.

These perturbations then grow into long persistent fingers. Finally, when the fingers reach the bottom of the chamber, the water flows out of the chamber through preferential pathways so that the most of the porous medium remains dry. This evolution of the wetting front instability is consistent with the experimental observation (DiCarlo, 2013).
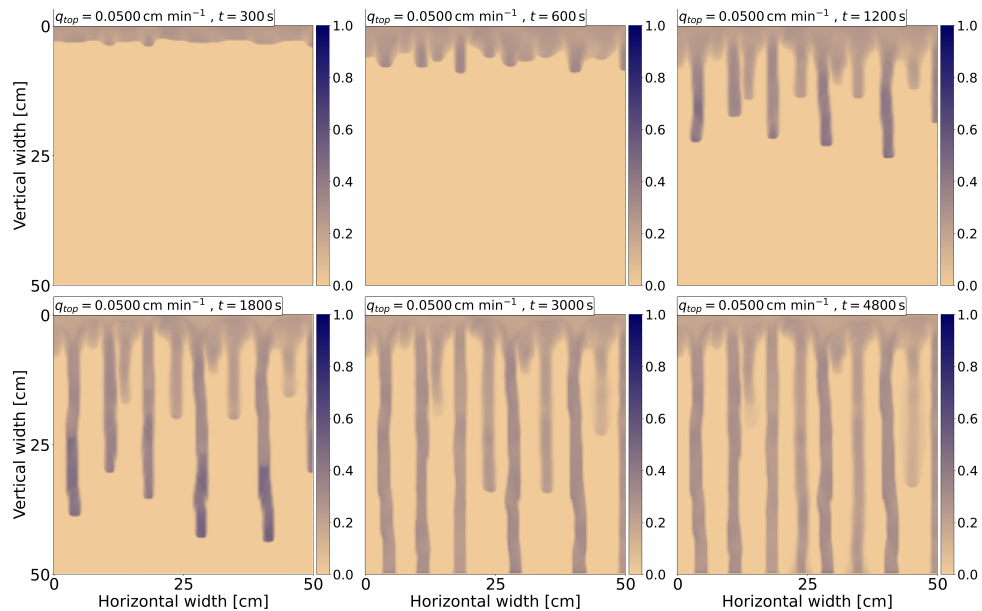
Figure 4. Evolution of saturation profile for $q_{\text{top}} = 0.05\,\text{cm}\,\text{min}^{-1}$ at six different times. Times are displayed in the upper left corner for each frame. Saturation is colored according to the color bar on the right.

4.3   Dependence of flow on infiltration rate

A total of 18 simulations for infiltration rates in range $q_{\text{top}} = 0.001 - 15\,\text{cm}\,\text{min}^{-1}$ are performed. Saturation profiles for nine infiltration rates are shown in Fig. 5. For clarity, saturation profiles for all infiltration rates are provided in Appendix A1 (Figs. A1 and A2). The time for each flux is selected so that the saturation reaches $40\,\text{cm}$ from the upper boundary. The flux and corresponding time are displayed in the upper left corner for each frame. At low fluxes, the transition from a stable wetting front to finger-like flow is clearly observed. As the flux approaches the hydraulic conductivity, the fingers widen, and a stable wetting front develops. To the best of our knowledge, this is the first model capable of simulating this non-trivial transition. More detailed views are available in the videos of transient simulations corresponding to each applied influx (Kmec, 2023b).

The saturation overshoot is evident for fluxes between ~~0.05~~0.075 $- 5\,\text{cm}\,\text{min}^{-1}$. Even at lower fluxes, for which unstable behavior persists, the saturation overshoot is observed. However, its magnitude (i.e., the saturation difference between finger tip and tail) is very small, and thus not visible in Fig. 5. This phenomenon is further discussed in Sect. 4.6. Additionally, a saturated wetting front emerges for infiltration rates lower than the saturated conductivity K_S, as shown in Fig. A2. This is a direct consequence of the developed saturation overshoot.
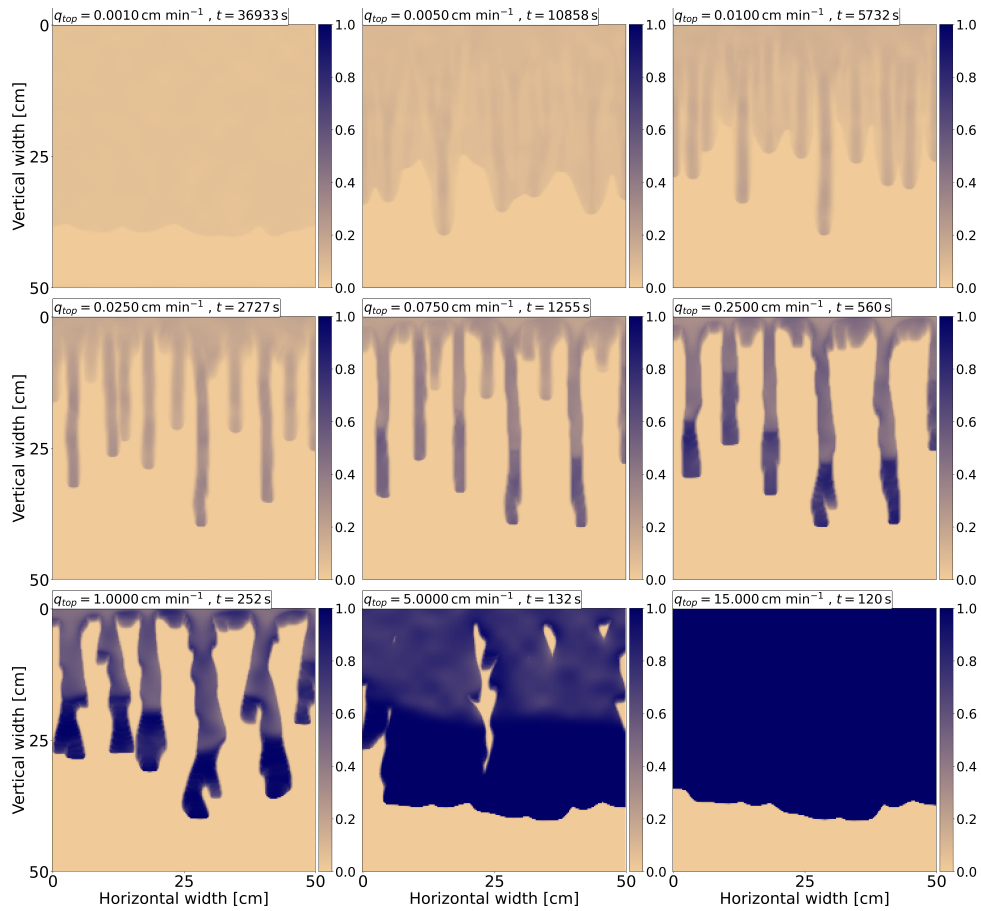
Figure 5. Saturation profiles for nine infiltration rates. For each frame, the influx is displayed together with the simulation time in the upper left corner. The transition from stable to unstable and back to stable flow is clearly observed. Saturation is colored according to the color bar on the right.

In cases of unstable flow, two fingers can merge, or one finger can split into two (Glass et al., 1989b, c; Rezanezhad et al., 2006). Both scenarios are reproduced here, but for detailed observation, we recommend viewing the videos of transient simulations (Kmec, 2023b). Merging can be observed at $q_{\text{top}} = 5\,\text{cm}\,\text{min}^{-1}$, where two wide fingers merge. However, finger merging is also noticeable at lower fluxes. Splitting can be seen at $q_{\text{top}} = 2.5\,\text{cm}\,\text{min}^{-1}$. Although these are minor details, such experimental consistency is beneficial.

4.4 Finger width as function of influx

DiCarlo (2013) plotted experimentally measured finger widths as a function of influx in a single graph, incorporating data for both low (Yao and Hendrickx, 1996) and high (Glass et al., 1989c) infiltration rates. The measured finger widths are shown in

Figure 2 in DiCarlo (2013) along with the predicted finger widths using standard theory (Chuoke et al., 1959; Parlange and Hill, 1976). The observed results can be summarized as follows:

- A stable wetting front is observed at very low fluxes, where the finger width is equal to the chamber width. This is not predicted by standard theory (Chuoke et al., 1959; Parlange and Hill, 1976).

- As the influx increases, a rapid decrease in finger width is observed, followed by a long flat valley of almost constant finger widths for different fluxes – specifically, the finger width first slightly decreases and then increases.

- As the influx approaches the saturated conductivity, the finger widths increase again, followed by stable flow.

- Fingers in contact with the edge of the chamber are narrower and thinner and are not included in the analysis of DiCarlo (2013). Simulations performed by the semi-continuum model are consistent with this observation; thus, these fingers are not included in the following analysis as well.

To calculate the finger width for a specific applied influx, we use the saturation profile at the time of the simulation when the water reaches the bottom of the chamber. This ensures that the bottom boundary condition does not influence the obtained results. For each saturation profile, all fingers are segmented as shown in Fig. A3 in Appendix A2. For preferential flow, this segmentation is straightforward because all the fingers are fully developed. It is worth noting that the fully developed fingers are persistent over time, so segmented fingers are not influenced by longer simulation time. In the case of stable flow, the approach used by DiCarlo (2013) is followed, where the entire saturation profile is assumed to be one "finger". However, the so-called intermediate flow – the transition between stable and unstable flow – is more complicated because some fingers are narrow while others are diffusively expanding. Therefore, completely objective segmentation is not possible in this case. After segmenting the fingers, the average finger width is calculated for each applied influx.

Figure 6 shows calculated finger widths for 18 infiltration rates. Red, green and blue dots indicate fluxes for which we observe stable, intermediate, and unstable flow, respectively. The results are consistent with experimentally observed behavior – a nearly constant finger width is observed for fluxes between $0.01 - 2.5\,\mathrm{cm\,min^{-1}}$, followed by stable flow for very low and very high applied fluxes. Even a slight increase in the finger width for fluxes above $0.01\,\mathrm{cm\,min^{-1}}$ is in perfect agreement with experiments (see Figure 2 in Glass et al. (1989c)). The specific case is at $q_{\mathrm{top}} = 5\,\mathrm{cm\,min^{-1}}$ (see Fig. 5), where two fingers are first developed and then both fingers merge at a depth of approximately $30\,\mathrm{cm}$. In this case, the average width is calculated from one merged finger.

According to the evaluation suggested by DiCarlo (2013), the model should be able to predict preferential flow in terms of finger widths and finger spacings. It has already been demonstrated that finger widths are captured well. To demonstrate the model's capability in terms of finger spacings, the number of fingers can be investigated. Glass et al. (1989c) calculated the number of fingers, including those in contact with the edge of the chamber. They reported that, in the case of preferential flow, the number of fingers does not change significantly with different fluxes, varying between four and six. Given the size of the chamber used in their experiments ($30\,\mathrm{cm}$), the expected number of fingers for the $50\,\mathrm{cm}$ used in our simulations is
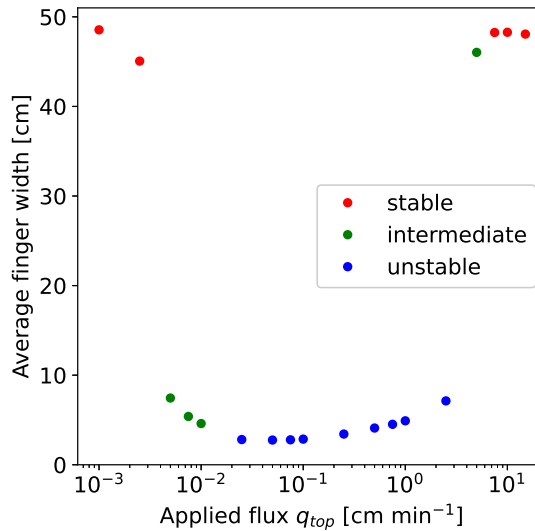
Figure 6. Finger width as a function of the applied influx. A stable wetting front is observed at either very low or very high applied fluxes – in this case, the "finger width" is approximately equal to the horizontal width of the chamber. The color of dots indicate the type of flow: stable, intermediate and unstable flow.

approximately between seven and ten. This corresponds well to the number of fingers developed in the simulations, which
375   ranges from six to ten for preferential flow. Hence, the spacing between the fingers is also well captured.

4.5   Preferential flow as function of influx

To calculate the degree of preferential water flow, the by-pass ratio approach is used, which is defined as the ratio of the preferential flow rate to the total flow rate (Kneale and White, 1984). First, the inflow is calculated for each block corresponding to the horizontal section at a depth of $30\,\mathrm{cm}$. These inflow values are then divided by the top boundary flux q_{top} to normalize
380   them to one. The normalized values represent the by-pass ratio. The simulation time is always chosen to be sufficiently long to ensure that extending the simulation time period further does not affect the calculated values. Note that if the flow is uniform throughout the porous medium, the by-pass ratio is equal to one everywhere.

    The left panel of Fig. 7 shows the by-pass ratio in the horizontal section at a depth of $30\,\mathrm{cm}$ for three different influxes. The color indicates the type of flow: red, green and blue denote examples of stable, intermediate and unstable flow, respectively.
385   The by-pass ratio for stable flow equals one almost everywhere with no significant flow preferences. For unstable flow, the by-pass ratio corresponds to the developed fingers. Water flows only through the fingers, and outside the flow is zero. The intermediate case is quite surprising: water flows through the entire porous medium, but the flow is still highly preferential. This is counterintuitive since the porous medium is fully wetted in this case, yet preferential pathways are formed through which the most of the water flows. The origin of these pathways can be seen in Fig. 5 for $q_{\mathrm{top}} = 0.005\,\mathrm{cm\,min^{-1}}$, where they
390   appear as a slight increase in saturation. For better illustration, the right panel of Fig. 7 shows the saturation profile at a longer

simulation time ($t = 24000\,\mathrm{s}$) with the maximum value of the color bar adjusted to make the pathways more visible. Developed pathways are well observed and do not disappear even when the porous medium is completely wetted. Comparing the by-pass ratio for $q_{\mathrm{top}} = 0.005\,\mathrm{cm\,min^{-1}}$ (the left panel of Fig. 7) and the corresponding saturation profile (the right panel of Fig. 7), we can clearly see that the highest by-pass ratio corresponds to the most saturated parts of the porous medium.
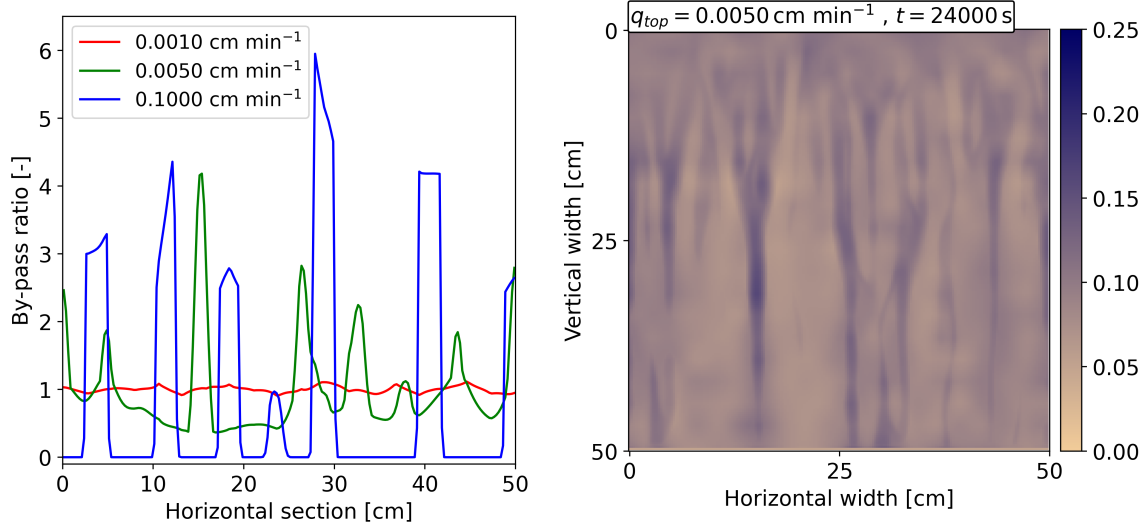
Figure 7. Left panel: The by-pass ratio, defined as the ratio of the preferential flow rate to the total flow rate, is shown for stable (red), intermediate (green) and unstable flow (blue) at a depth of $30\,\mathrm{cm}$ in the porous medium. There are no significant preferences for stable flow, while pronounced preferential pathways develop for unstable flow. Such behavior is expected for both cases. A surprise, however, is the intermediate case, where water still flows preferentially even when the porous medium is fully wetted. Right panel: Saturation profile for $q_{\mathrm{top}} = 0.005\,\mathrm{cm\,min^{-1}}$ at time $t = 24000\,\mathrm{s}$. The simulation corresponds to the intermediate case in the left panel of this figure. Saturation is colored according to the color bar on the right. The pathways of saturation are clearly visible and correspond to the highest by-pass ratio values.

395   We conjecture that this is a very important observation about the problem of preferential flow in unsaturated porous media because experimental measurements are highly limited for this case. Using the semi-continuum model, more details can be observed, making it easier to understand the origin of the preferential pathways for different boundary conditions. For completeness, by-pass ratios for all performed simulations are depicted in Fig. A4. It is evident that a similar scenario, as for $q_{\mathrm{top}} = 0.005\,\mathrm{cm\,min^{-1}}$, holds for another two intermediate cases, i.e. for $q_{top} = 0.0075$ and $0.01\,\mathrm{cm\,min^{-1}}$.

400   To compare the degree of preferential water flow for different boundary fluxes, it is convenient to represent it by a single value. Again, the horizontal section at a depth of $30\,\mathrm{cm}$ is used. We calculate the smallest number of blocks through which at least $50\,\%$ of the total amount of water flows at this horizontal section. For each influx, the length is then calculated as $\Delta x \times n_B$, where n_B denotes the calculated number of blocks. In the case of uniform flow, the length is equal to half the horizontal width of the porous medium, i.e., $25\,\mathrm{cm}$. The smaller the length is, the more the preferential flow dominates.

405   Figure 8 shows the length of the porous medium through which $50\,\%$ of the water flows, depending on influx. For stable flow, half of the total water flows through almost $25\,\mathrm{cm}$, while for unstable flow, it ranges between $5.50 - 8.75\,\mathrm{cm}$. An exception is at $q_{\mathrm{top}} = 2.5\,\mathrm{cm\,min}^{-1}$, where the length is $12.50\,\mathrm{cm}$ due to significantly wider fingers compared to lower fluxes. For $q_{\mathrm{top}} = 0.005 - 0.01\,\mathrm{cm\,min}^{-1}$ (intermediate case), the length is similar to the values of unstable flow, indicating significant preferential flow. This aligns with the by-pass ratio analysis.
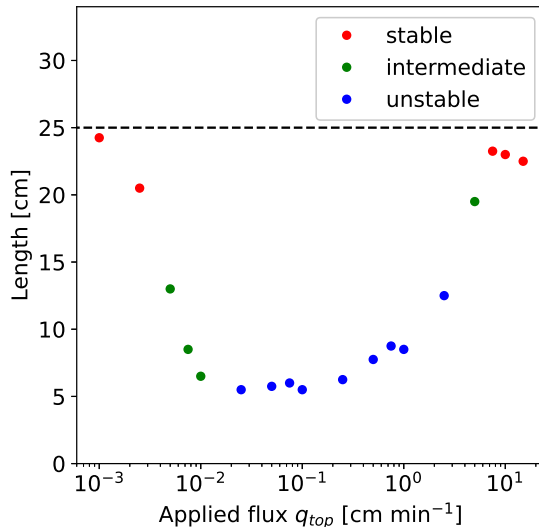
Figure 8. The length of the porous medium in the horizontal section at a depth of $30\,\mathrm{cm}$ through which $50\,\%$ of the water flows is plotted against the influx. For each influx, the simulation times are chosen such that the flow through the porous medium does not change any further. The dashed line denotes the case of uniform flow, i.e., $25\,\mathrm{cm}$. The color of the dots indicates the type of flow: stable, intermediate and unstable flow.

4.6   Wetting front instability

410

     The cause of preferential flow is the saturation and pressure overshoot (Eliassi and Glass, 2001; Egorov et al., 2003; DiCarlo, 2013). The flux range for the saturation overshoot in 1D (DiCarlo, 2004) experimentally corresponds to the flux range for preferential flow in higher dimensions (Yao and Hendrickx, 1996; Glass et al., 1989c). The same applies to different initial saturation levels (DiCarlo, 2004; Bauters et al., 2000). This simplifies the analysis of wetting front instability, as we can

415   switch to 1D. To analyze whether this experimentally confirmed dependence is replicated by the semi-continuum model, 1D simulations are performed using the same parameters as for 2D simulations. However, the distribution of the intrinsic permeability is not included to avoid influencing the results.

     The left panel of Fig. 9 shows saturation profiles for six different applied fluxes q_{top}. The range for which the saturation overshoot occurs is the same for 1D and 2D simulations. For the lowest influx, $q_{\mathrm{top}} = 0.001\,\mathrm{cm\,min}^{-1}$, the profile is stable

420   without saturation overshoot. For $q_{\mathrm{top}} = 0.010\,\mathrm{cm\,min}^{-1}$, the saturation overshoot is formed and becomes more pronounced

with increasing influx up to $q_{\text{top}} = 1.000\,\text{cm}\,\text{min}^{-1}$. The saturation overshoot then becomes less pronounced until it disappears completely, and a stable profile is observed again for the highest influx $q_{\text{top}} = 15.00\,\text{cm}\,\text{min}^{-1}$. This is in good agreement with 1D experiments (DiCarlo, 2010).
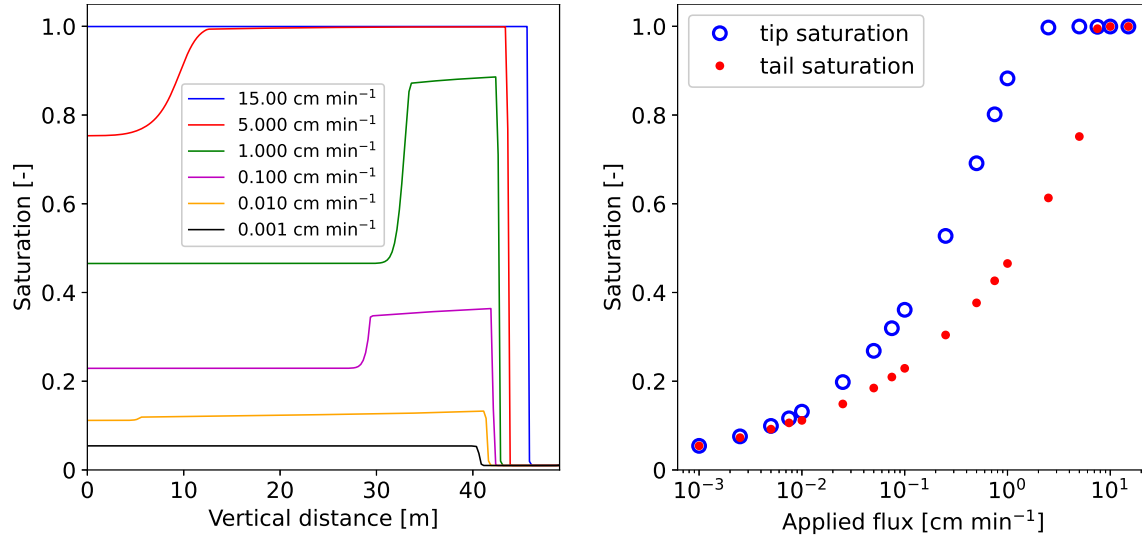
Figure 9. Left panel: Saturation profiles for various applied fluxes. The saturation overshoot is not developed for the lowest flux ($q_{\text{top}} = 0.001\,\text{cm}\,\text{min}^{-1}$) and the highest flux ($q_{\text{top}} = 15.00\,\text{cm}\,\text{min}^{-1}$). Right panel: Finger tip and tail saturations for 18 different infiltration rates. The occurrence of the saturation overshoot is consistent with 2D preferential flow.

The right panel of Fig. 9 shows the saturation of finger tip and tail for 18 different infiltration rates. The results are again in agreement with 1D experiments (DiCarlo, 2010). Moreover, the flux range connection between 1D overshoot and 2D preferential flow is well captured by the semi-continuum model. When preferential flow is observed in 2D, significant saturation overshoot is developed in 1D. Conversely, for diffusion-like flow in 2D, no overshoot occurs in 1D. ~~However,~~ The most interesting ~~is the~~ are intermediate case~~s~~ for flux range between $q_{\text{top}} = 0.005 - 0.010\,\text{cm}\,\text{min}^{-1}$, where ~~a~~ very small saturation overshoot~~s~~ ~~is~~ are formed, with magnitudes between 0.007 and 0.02. Even a small saturation overshoot with a magnitude below 0.02 indicates that the flow is preferential in 2D. For ~~$q_{\text{top}} = 0.0005\,\text{cm}\,\text{min}^{-1}$~~ $q_{\text{top}} = 0.005\,\text{cm}\,\text{min}^{-1}$, the magnitude of the saturation overshoot equals only 0.007, yet the flow remains preferential. ~~even in this case. Such a close connection is quite surprising.~~ This close relationship is quite surprising.

5 Discussion

5.1 Preferential flow in the case of fully wetted porous medium

The semi-continuum model is the first to correctly predict both diffusion-like and finger-like flow for various applied fluxes, as well as accurately predict experimentally measured finger widths and spacings. The well-captured connection between

saturation overshoot in 1D and preferential flow in 2D has been also demonstrated. In addition to these results, we have observed that flow can be highly preferential even when the porous medium is fully wetted and water flows through the entire medium. This surprising behavior is illustrated for ~~$q_{\text{top}} = 0.0005\,\text{cm}\,\text{min}^{-1}$~~ $q_{\text{top}} = 0.005\,\text{cm}\,\text{min}^{-1}$ in Fig. 7. The flow

440 remains preferential because pathways with slightly increased saturation develop during infiltration into a dry porous medium and persist over time. This slight increase in saturation is enough to make the flow preferential, causing water to flow faster through these pathways. This is due to the power-law nature of the relative permeability function. The left panel of Fig. 10 shows the by-pass ratio along with the saturation and relative permeability at a depth of $30\,\text{cm}$. Clearly, the highest by-pass ratio values correspond to the highest relative permeability values. A slight difference is caused by the distribution of the intrinsic

445 permeability, as demonstrated in the right panel of Fig. 10, where the effective permeability closely follows the by-pass ratio.
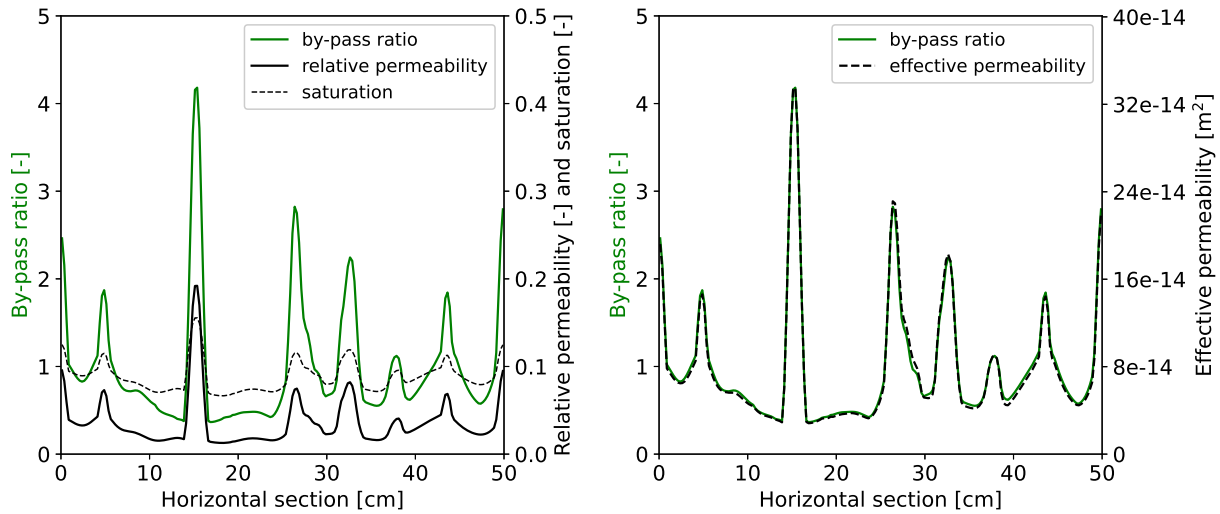
Figure 10. Left panel: The by-pass ratio at a depth of $30\,\text{cm}$ of the porous medium is plotted for $q_{\text{top}} = 0.005\,\text{cm}\,\text{min}^{-1}$ at time $t = 24000\,\text{s}$, together with the saturation and the corresponding relative permeability. By-pass ratio values are plotted on the left y-axis, and values of saturation and relative permeability are plotted on the right y-axis. Relative permeability values are multiplied by 100, since these values are much smaller compared to the saturation. It can be clearly seen that the by-pass ratio correlates with relative permeability. Right panel: The by-pass ratio at a depth of $30\,\text{cm}$ of the porous medium is plotted for $q_{\text{top}} = 0.005\,\text{cm}\,\text{min}^{-1}$ at time $t = 24000\,\text{s}$, together with the effective permeability. By-pass ratio values are plotted on the left y-axis, and values of the effective permeability are plotted on the right y-axis. Effective permeability closely follows the by-pass ratio in this case.

5.2 Formation of saturation overshoot and stabilization of wetting front

Borrowing the term "hold-back pile-up effect" from Eliassi and Glass (2001), saturation overshoot occurs when water cannot enter the dry porous medium, causing water to be held back. This hold back effect subsequently leads to water piling up above the interface between the wet and dry porous medium. As the amount of water increases, the pressure gradient between the wet and dry parts of the porous medium also increases, allowing water to advance. The same principle applies to the semi-

450

continuum model, where the hold-back effect is caused by appropriate averaging of conductivity. However, it is not the only factor governing the formation of the saturation overshoot in the model. The saturation overshoot is formed by two factors: (1) the geometric mean for averaging the permeability and (2) the scaling of the retention curve. The geometric mean plays a crucial role in creating the hold-back effect, which consequently forms the saturation overshoot. This effect occurs due to the very low relative permeability between blocks, which is a direct consequence of the applied geometric mean. However, without the scaling of the retention curve, the overshoot would disappear as the block size $\Delta x \rightarrow 0$ (Vodák et al., 2022). Note that the geometric mean can be replaced by any type of averaging that produces a small value if one of the averaged numbers is small, such as the harmonic mean. The geometric mean is used in the model because it is more appropriate in the case of a random stratified medium, while the harmonic mean is more suitable for a medium which is stratified perpendicular to the flow direction (Jang et al., 2011).

A crucial question arises: Why does the wetting front stabilize at low or high infiltration rates in the semi-continuum model? It is known that the stability of the 2D wetting front correlates with the saturation overshoot in 1D (DiCarlo, 2013). Since the semi-continuum model implies the same conclusions, as demonstrated in Sect. 4.6, it is sufficient to study the origin of the saturation overshoot in 1D. This significantly simplifies the analysis as we do not have to consider the spatial variability of flow in 2D or 3D. When a small amount of water flows into a dry porous medium, the conductivity of this part increases rapidly due to the power-law nature of the relative permeability function. Using the parameters in Table 1, when the saturation increases from 0.01 to 0.05, the relative permeability increases approximately 170 times. At very low infiltration rates, water has enough time to proceed downwards, so that the conductivity of the dry part of the porous medium increases significantly, which subsequently allows more water to flow downwards. In principle, this means that water does not have sufficient time to pile up at low infiltration rates. On the other hand, at very high infiltration rates, the porous medium is fully saturated, and hence the saturation overshoot cannot occur. This is well observed in Fig. 9 for $q_{\text{top}} = 15 \, \text{cm} \, \text{min}^{-1}$, where the saturation is equal to one everywhere. As the applied flux decreases, the porous medium is not fully saturated and there is room for the saturation overshoot to arise. This can be seen for $q_{\text{top}} = 5 \, \text{cm} \, \text{min}^{-1}$; the saturation at the finger tip is still equal to one, but the influx is not high enough to fully saturate the entire porous medium, and the saturation overshoot develops. This observation is in agreement with the experiments of DiCarlo (2010). The author showed that the saturation of the finger tip approaches unity at very high infiltration fluxes for which a stable flow is observed (see figure 3 in DiCarlo (2010) for details).

According to comprehensive experimental works of DiCarlo (2004, 2010), saturation and pressure are constant at the finger tail under uniform top boundary condition. However, some studies, such as Cho et al. (2005), have shown the opposite. Assuming constant saturation and pressure in the finger tail, the flux between the blocks is stabilized and is given by the value of top boundary flux q_{top}. Using the Darcy-Buckingham law given by Eq. (7), we can calculate the saturation in the finger tail:

$$k(S_{\text{tail}}) = \frac{q_{\text{top}}}{K_S}, \tag{9}$$

where S_{tail} denotes the saturation in the finger tail within the relative permeability function. Therefore, S_{tail} is independent of the initial saturation, which is in agreement with experimental measurements (Fritz, 2012). A necessary and sufficient condition for saturation overshoot to form is an increase in saturation of the block above the value S_{tail}. The saturation overshoot will

485 then develop as the saturation of this block drops to S_{tail} over time. Moreover, for $q_{\text{top}} = K_S$, it follows from Eq. (9) that S_{tail} equals one. Consequently, the flow will always be stable in the semi-continuum model as long as the influx $q_{\text{top}} \geq K_S$ because the porous medium is fully saturated in this case, and thus saturation overshoot cannot arise. This has been demonstrated in both 1D (Fig. 9) and 2D simulations (Fig. 5). Note that this is consistent with the stability condition (Saffman and Taylor, 1958; Parlange and Hill, 1976; DiCarlo, 2013), which predicts stable flow for $q_{\text{top}} \geq K_S$.

5.3 Importance of the geometric mean in terms of water entry value

The porous medium comprises many pores, each characterized by a specific water entry value determined by the Young-Laplace equation based on its principal radii. The shape of the main wetting branch of the retention curve results from the combination of various pore water entry values. The pores with the smallest radii determine the lowest pressure P_{low} and the corresponding lowest saturation S_{low}, marking the beginning of the main wetting branch. Although this is a necessary
495 physical characteristic of the porous medium, it has not been implemented into the retention curve used for simulations. This implies that water can enter dry porous medium at a pressure equal to $-\infty$. In this case, the Young-Laplace equation gives $r_1^{-1} + r_2^{-1} = \infty$, where r_1 and r_2 are the principal radii of curvature of the pore. This relationship is valid only if at least one of these radii equals zero. Such a hypothetical pore is obviously unable to conduct water. The same reasoning applies to the air entry value and the corresponding main draining branch.
500 The initial saturation used in simulations is higher than S_{low}, thus excluding the beginning of the main wetting branch does not affect the obtained results. One might argue that even if the chosen initial saturation is above S_{low}, it would still be reasonable to define the value S_{low} in the retention curve, ensuring the model's validity for lower initial saturation. This is clearly not implemented in our model, and the retention curve satisfies $P \to -\infty$ for $S \to 0$. However, the calculated flux in the semi-continuum model equals zero in this case, which aligns with a hypothetical pore with zero radii. This applies due
505 to the application of the geometric mean of the relative permeability and is not valid for the more common arithmetic mean. To illustrate, consider two blocks denoted by indices $1, 2$: one fully saturated ($S_1 = 1$) and the second block with saturation decreasing towards zero, i.e., $S_2 \to 0$. Assume κ, μ and Δx are equal to one, as these values are independent of saturation and thus does not affect the limiting process. Note that a fully saturated block satisfies $k(S_1) = 1$ and $P^w(S_1) = 0$. The horizontal flux q between blocks, given by Eq. (7), then simplifies to:

$$510 \quad \lim_{S_2 \to 0} q = \lim_{S_2 \to 0} \left[-\sqrt{k(S_2)} P^w(S_2) \right] = 0. \tag{10}$$

If the arithmetic mean is used instead of the geometric mean, the limit satisfies:

$$\lim_{S_2 \to 0} q = \lim_{S_2 \to 0} -\frac{1 + k(S_2)}{2} P^w(S_2) = +\infty. \tag{11}$$

For completeness, the numerical limiting process for both means is depicted in Fig. 11 using the parameters specified in Table 1. For the geometric mean, the limit approaches zero, confirming that the flux indeed equals zero for $S_2 \to 0$, while for
515 the arithmetic mean, the flux approaches infinity. In the case of the geometric mean, the block with zero saturation thus cannot

conduct water and represents a hypothetical pore with zero radii. This is consistent with the Young-Laplace equation that yields a pore with zero radii for $P = -\infty$. On the other hand, with the application of the arithmetic mean, unrealistic behavior would occur: the flux would rapidly increase as saturation decreases, which is clearly not physically correct. Using the arithmetic mean, it is necessary to cut off the retention curve (e.g., by using a water entry value) for low saturation values to avoid an

520   increase in flux when saturation decreases. However, the increase in flux in Fig. 11 occurs already for saturation $S_2 < 0.57$. We conjecture that the geometric mean or the harmonic mean is thus a more reasonable type of averaging (Jang et al., 2011).
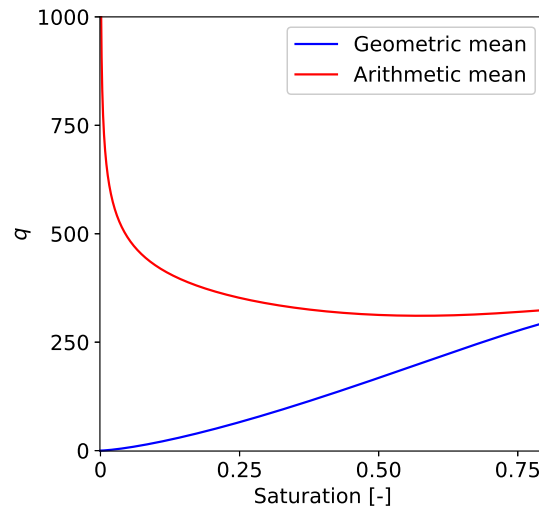
Figure 11. Calculated horizontal flux between a fully saturated block ($S_1 = 1$) and a block with saturation decreasing towards zero ($S_2 \to 0$). The flux equals zero for $S_2 \to 0$ using the geometric mean, while it approaches infinity using the arithmetic mean.

Incorporating the point $[S_{low}, P_{low}]$ into the main wetting branch results in a non-smooth retention curve. However, the Richards' equation remains unconditionally stable for saturation values above S_{low}, where the retention curve is smooth. Therefore, to develop saturation overshoot with the Richards' equation, it is necessary to violate one of the assumptions of the

525   stability proof derived in Fürst et al. (2009). It is possible to define a bottleneck (zero flux) using a water or air entry pressure (Tesař et al., 2004), which causes non-smoothness of the retention curve. Alternatively, using a non-monotonic influx at the upper boundary can also lead to the formation of saturation overshoot (Steinle and Hilfer, 2017). We conjecture that the model should ideally generate the overshoot without the need for a threshold incorporated into the model. Forming the saturation overshoot should be a resulting property of the model.

530   ## 5.4   Effect of intrinsic permeability distribution

~~The measurement of macroscopic properties of a homogeneous porous medium, such as permeability or porosity, is usually performed by averaging the microscopic quantities over the domain (White et al., 2006; White et al., 2021). For more realistic simulations, it is appropriate to use the spatial variation of the continuum quantity. Therefore, a small distribution of the intrinsic permeability was used in simulations, which is typical for models for unsaturated porous media flow (Cueto-Felgueroso and Juanes, 2009).~~

535 One may wonder whether the formation of the saturation overshoot depends on the intrinsic permeability distribution used. As evidenced by 1D simulations in Fig. 9, the saturation overshoot is not caused by the distribution of intrinsic permeability. This corresponds with experimental findings (DiCarlo, 2013), where the formation of saturation overshoot is not determined by heterogeneity. Moreover, Kmec et al. (2023) demonstrated that the distribution has no significant effect on the flow, because the nature of the flow remains the same even for eight different intrinsic permeability distributions.

540   Hence, the question arises: Why use the distribution of intrinsic permeability, and how does this distribution affect the development of preferential pathways? Even a slight distribution of intrinsic permeability can cause water not to flow uniformly through the entire porous medium in 2D/3D. This is expected because no heterogeneity is introduced in the governing equation; hence, water will always flow uniformly unless some additional heterogeneity is introduced. This is demonstrated for 2D simulations in Fig. 12, where saturation profiles for four different infiltration rates without a distribution of intrinsic permeability are

545 shown. The overshoot is observed at $q_{\text{top}} = 0.05$ and $0.25\,\text{cm}\,\text{min}^{-1}$, but not at very low and very high infiltration rates, which is consistent with the previous simulations. Therefore, it is evident that the intrinsic permeability distribution is not the cause of the formation of the saturation overshoot in 2D, but causes water to flow preferentially. To summarize: (1) When using only the distribution of intrinsic permeability without incorporating the geometric mean and the scaling of the retention curve, the overshoot will not be formed. Consequently, the flow is diffusive in this scenario, and water always flows throughout the entire

550 porous medium. (2) When using only the geometric mean and the scaling of the retention curve, the overshoot can be formed depending on initial and boundary conditions. However, the flow remains uniform throughout the entire porous medium. By combining (1) and (2), the saturation overshoot can be formed, and water then does not flow uniformly.
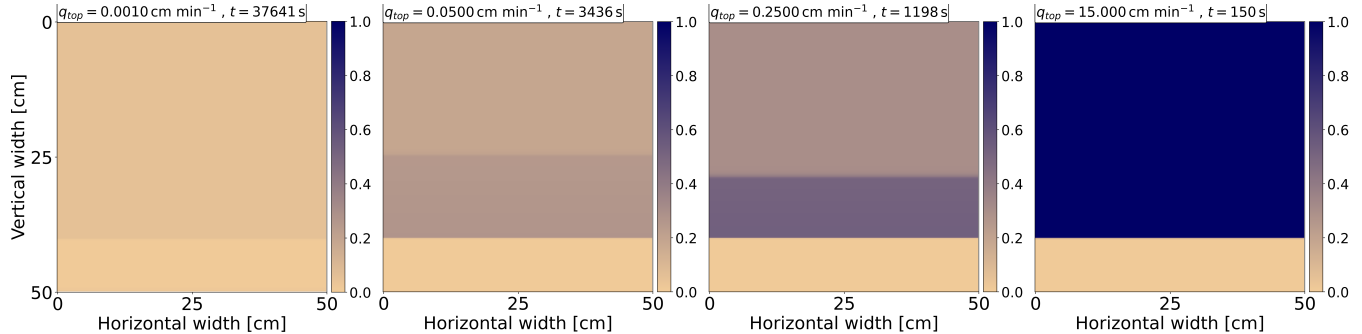
Figure 12. Saturation profiles for four different infiltration rates $q_{\text{top}} = 0.001, 0.05, 0.25, 15\,\text{cm}\,\text{min}^{-1}$. The distribution of intrinsic permeability is not used. Each frame displays the influx and simulation time. Saturation is colored according to the color bar on the right.

Moreover, when the distribution is employed in the semi-continuum model but the saturation overshoot does not occur (e.g., in the case of low influx), water tends to flow diffusively. This aligns well with experimental observations. Other similar

555 models fail in this case, as water flows preferentially even when the saturation overshoot is not formed. In our case, if the overshoot is formed, preferential flow is observed. If there is no overshoot, preferential flow disappears. The distribution of intrinsic permeability is included to make water flow non-uniformly (preferentially) throughout the entire porous medium, but this non-uniformity occurs only when the physics of the semi-continuum model allows it, i.e., when the overshoot is formed.

5.5 Model validation and future plans

560 The semi-continuum model has been shown to be consistent with well-known experiments in 1D (Kmec et al., 2019) and 2D (Kmec et al., 2021, 2023). Based on the results presented in this paper, we conjecture that the model has been validated by core experiments performed in unsaturated homogeneous porous media (Glass et al., 1988, 1989b, c, a; Selker et al., 1992; Liu et al., 1994; Yao and Hendrickx, 1996; Bauters et al., 2000; DiCarlo, 2004; Sililo and Tellam, 2005; Rezanezhad et al., 2006; DiCarlo, 2007, 2010). For heterogeneous porous medium, the model has been applied (Kmec et al., 2021) to simulate

565 water infiltration experiments into a layered porous medium (Rezanezhad et al., 2006). Therefore, the semi-continuum model is successfully validated according to DiCarlo's approach (DiCarlo, 2013).

A crucial aspect of the model is to account for the retention curve sensitivity to the dimension of the laboratory sample (Ghanbarian et al., 2015). This sensitivity is addressed by incorporating a single parameter, the reference block size Δx_0, into the semi-continuum model. This makes the simulations sensitive to this parameter, but once Δx_0 is appropriately set

570 through calibration, the simulations become independent of the block size Δx. Therefore, the results are consistent regardless of the computational mesh size used. The convergence of moisture profiles in 1D and 2D for Δx varying over two orders of magnitude is demonstrated in Vodák et al. (2022).

The semi-continuum model can be adapted for more complex natural conditions by incorporating different initial and boundary conditions. This adaptability makes the model applicable to large-scale in-field experiments, which is highly desirable for

575 hydrological applications. A complete validation of the semi-continuum model ensures the accuracy and reliability of its results on a large scale. However, the semi-continuum model is computationally intensive, thus the development of a method to speed up its numerical scheme is necessary. Our future plan is to develop a new numerical scheme based on the Lattice Boltzmann method, as it has been successfully applied to ~~the~~ Richards' equation (Ginzburg et al., 2004). This will allow to simulate large-scale experiments effectively.

580 ## 6   Conclusions

It has long been known that the behavior of the wetting front in an initially dry and homogeneous porous medium strongly depends on the applied infiltration flux. The wetting front remains stable at both low and high infiltration fluxes but becomes unstable within a certain intermediate flux range. The instability is characterized by the formation of fingers that vary in width, velocity, and spacing. Despite decades of experimental observations, no model has yet been developed to reliably capture this

585 dependence. In this paper, we introduced a partial differential equation that includes a Prandtl-type hysteresis operator under the spatial derivative. This equation represents a formal limit of the semi-continuum model of liquid transport in a porous medium. The semi-continuum model accurately captures the complex behavior of infiltration flux dependence. This includes the transition from diffusion-like to finger-like flow and back to diffusion-like flow as the infiltration flux increases. It also correctly predicts finger widths and spacings. In addition, the model helps explain preferential flow and provides insight into

590 the formation of an unstable wetting front.

Appendix A

A1 Dependence of flow on infiltration rate

A total of 18 simulations for infiltration rates ranging from $q_{\mathrm{top}} = 0.001\,\mathrm{cm\,min}^{-1}$ to $q_{\mathrm{top}} = 15\,\mathrm{cm\,min}^{-1}$ are presented here. Figures A1 and A2 show a snapshot of the saturation profiles for all applied fluxes. The time for each influx is chosen so that the saturation reaches $40\,\mathrm{cm}$ from the upper boundary. The applied influx and simulation time are displayed at the upper left corner of each frame. The transition between stable and unstable flow is clearly observed.
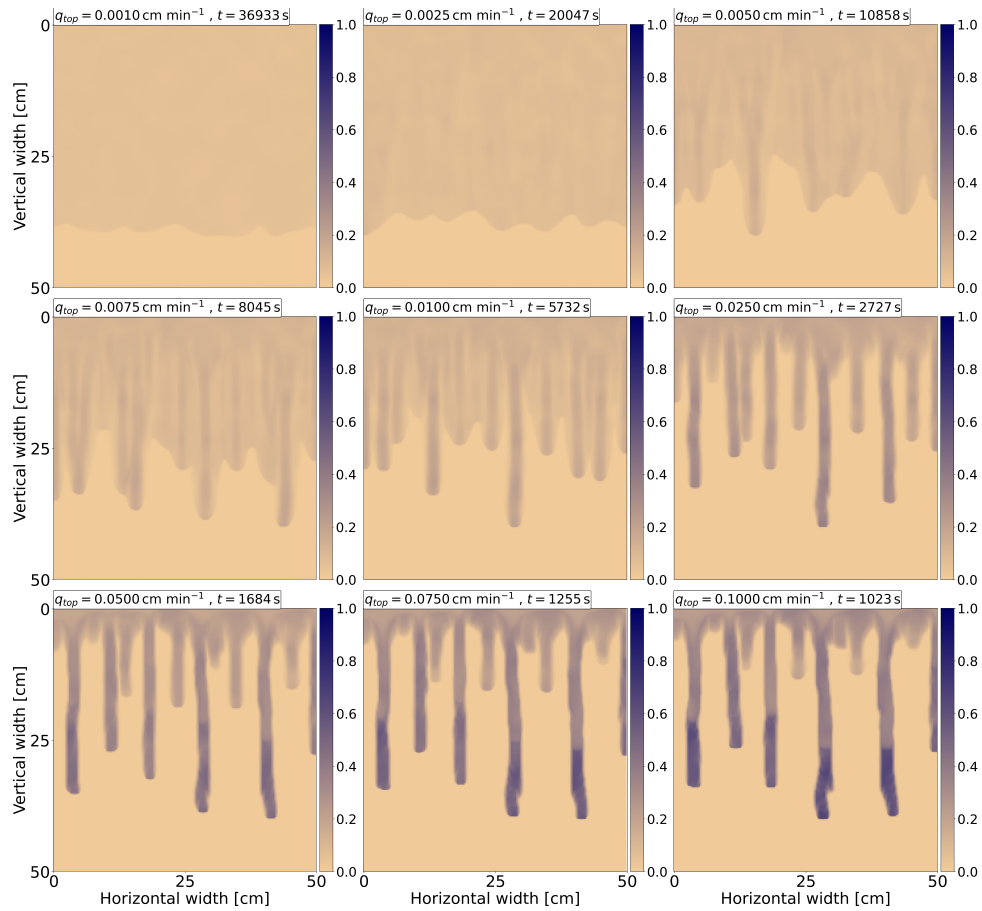
Figure A1. Saturation profiles for nine different infiltration rates ranging from $q_{\mathrm{top}} = 0.001\,\mathrm{cm\,min}^{-1}$ to $0.1\,\mathrm{cm\,min}^{-1}$. Each frame displays the influx and simulation time. Saturation is colored according to the color bar on the right.
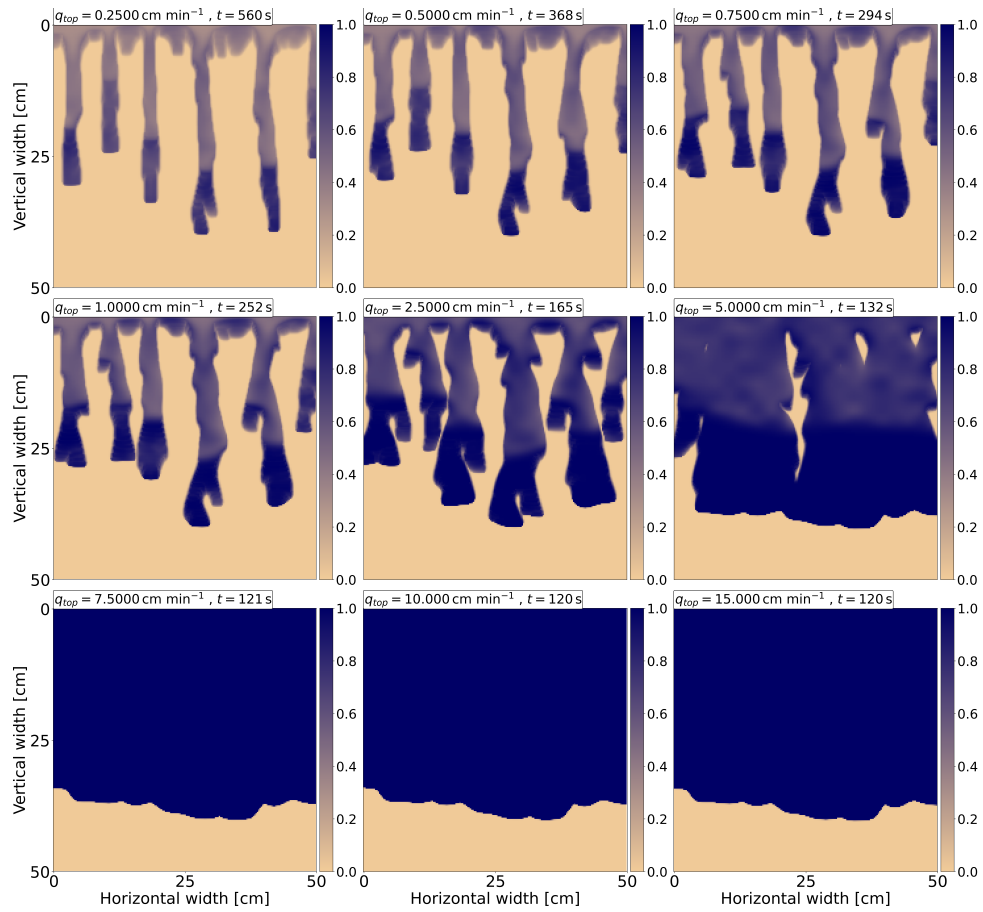
Figure A2. Saturation profiles for nine different infiltration rates ranging from $q_{\mathrm{top}} = 0.25\,\mathrm{cm}\,\mathrm{min}^{-1}$ to $15\,\mathrm{cm}\,\mathrm{min}^{-1}$. Each frame displays the influx and simulation time. Saturation is colored according to the color bar on the right.

A2 Finger segmentations
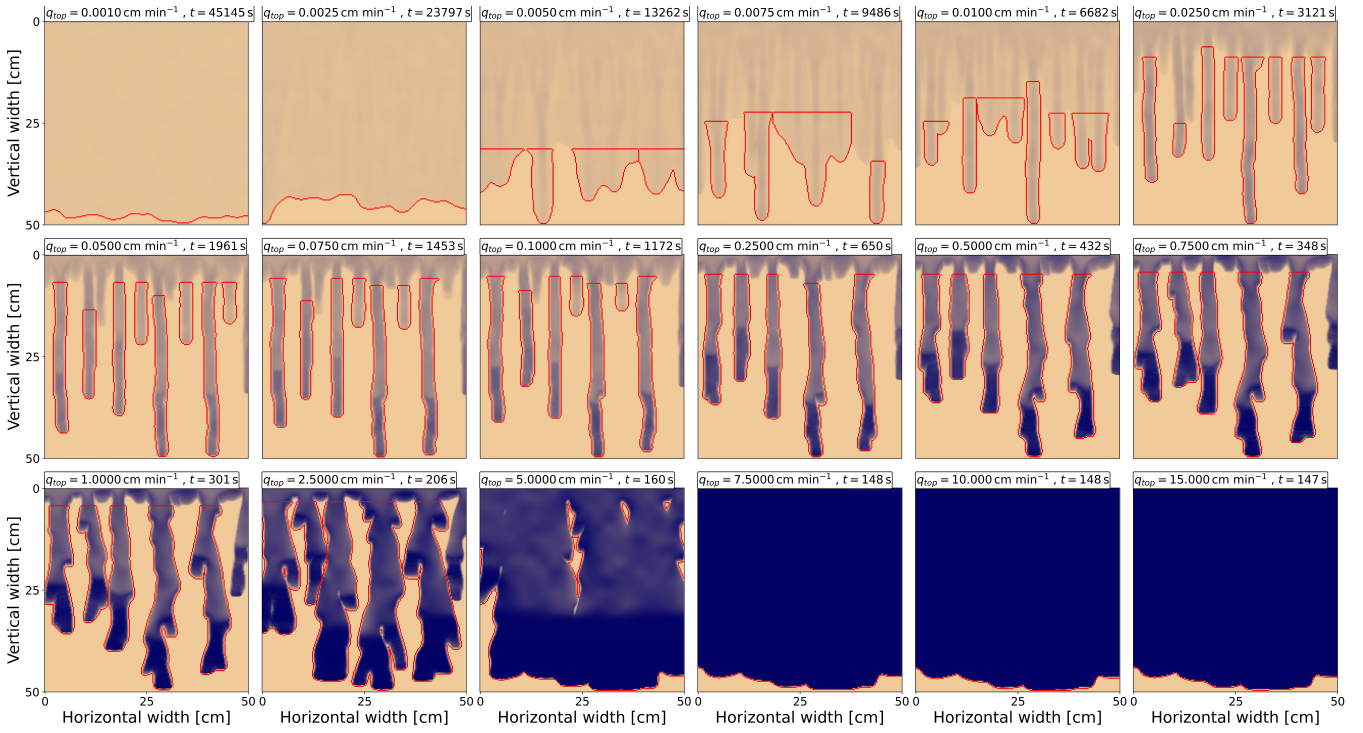
Figure A3. Finger segmentations for 18 different infiltration rates ranging from $q_{\text{top}} = 0.001\,\text{cm}\,\text{min}^{-1}$ to $q_{\text{top}} = 15\,\text{cm}\,\text{min}^{-1}$. The segmentations are marked in red. The time for each influx is chosen so that the saturation reaches the bottom of the chamber. The applied influx and simulation time are displayed at the upper left corner of each frame.

A3 By-pass ratio

Figure A4 shows the by-pass ratio at a depth of $30\,\text{cm}$ for 18 different infiltration rates ranging from $q_{\text{top}} = 0.001\,\text{cm}\,\text{min}^{-1}$
to $q_{\text{top}} = 15\,\text{cm}\,\text{min}^{-1}$. The by-pass ratio is marked in red and is plotted along with the saturation (in blue) across the same horizontal section. The highest values of the by-pass ratio correlate with the locations of the highest saturation values. However, this correlation decreases if the saturation is high enough, as seen for $q_{\text{top}} \geq 7.5\,\text{cm}\,\text{min}^{-1}$. In such cases, the slight change in saturation does not significantly affect the relative permeability values, and therefore, the by-pass ratio remains also unaffected. Instead, the dominant factor influencing the flow is the intrinsic permeability of the porous medium.

Note that the porous medium is first fully saturated for $q_{\text{top}} = 7.5\,\text{cm}\,\text{min}^{-1}$ and $q_{\text{top}} = 10\,\text{cm}\,\text{min}^{-1}$ (see Fig. A2). However, the saturation then decreases because the bottom boundary flux q_{bot} approximately equals $K_S = 15\,\text{cm}\,\text{min}^{-1}$, which is larger than q_{top}. In this scenario, the saturation should correspond to the calculated value S_{tail} given by Eq. (9). Considering the distribution of the intrinsic permeability, the average value of the calculated S_{tail} at a depth of $30\,\text{cm}$ is 0.8540 and 0.9218

for $q_{\text{top}} = 7.5\,\text{cm}\,\text{min}^{-1}$ and $q_{\text{top}} = 10\,\text{cm}\,\text{min}^{-1}$, respectively. The average values from simulations at a depth of $30\,\text{cm}$ are

610   0.8552 and 0.9225, showing that the saturation indeed corresponds to the calculated values S_{tail}. Therefore, the saturation at the finger tail always tends to decrease to the value obtained by Eq. (9). Moreover, for $q_{\text{top}} = 15\,\text{cm}\,\text{min}^{-1}$, the porous medium remains fully saturated even when water flows from the bottom boundary, as $S_{\text{tail}} = 1$ in this case.
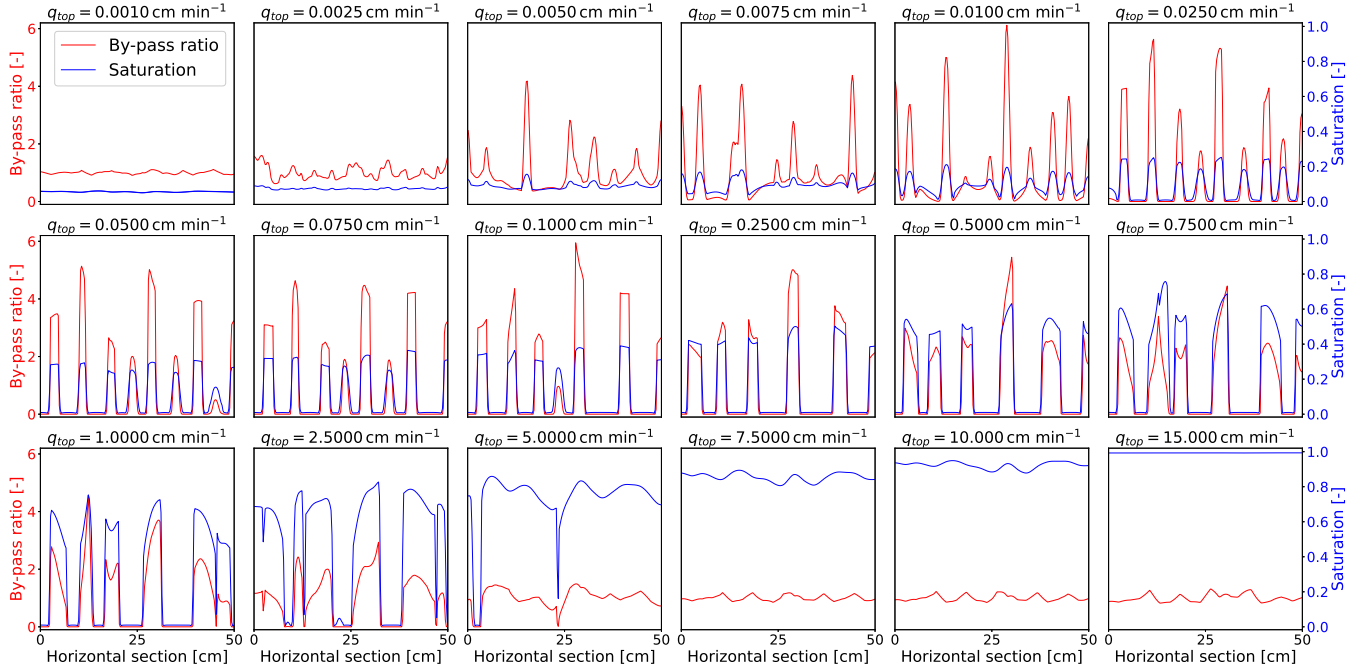
Figure A4. The by-pass ratio at a depth of $30\,\text{cm}$ of the porous medium for 18 different infiltration rates ranging from $q_{\text{top}} = 0.001\,\text{cm}\,\text{min}^{-1}$ to $q_{\text{top}} = 15\,\text{cm}\,\text{min}^{-1}$. The by-pass ratio is marked in red and is plotted along with the saturation (in blue) across the same horizontal section. By-pass ratio values are plotted on the left y-axis and values of saturation are plotted on the right y-axis. The corresponding influx is displayed at the top of each frame.

Code and data availability. The software code used to produce the simulations is written in Python and can be downloaded from Kmec (2023a). The code supports 1D, 2D and 3D simulations. The simulation data required to create the plots included in the manuscript can

615   be downloaded from Kmec and Šír (2024a, b). Please do not hesitate to contact us if you encounter any problems while downloading the software code and simulation data.

Video supplement. Videos of the transient simulations for all 2D cases can be downloaded from Kmec (2023b).

Author contributions. JK and MS wrote the manuscript, JK implemented the computer code, ran and analysed the simulations.

Competing interests. The authors declare that they have no conflict of interest.

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References

Abreu, E., Bustos, A., Ferraz, P., and Lambert, W.: A Relaxation Projection Analytical–Numerical Approach in Hysteretic Two-Phase Flows in Porous Media, J. Scient. Comp., 79, 1936–1980, 2019.

Aminzadeh, B. and DiCarlo, D. A.: The Transition between Sharp and Diffusive Wetting Fronts as a Function of Imbibing Fluid Properties, Vadose Zone J., 9(3), 588–596, https://doi.org/10.2136/vzj2009.0072, 2010.

Bauters, T. W. J., DiCarlo, D. A., Steenhuis, T., and Parlange, J.-Y.: Soil water content dependent wetting front characteristics in sands, J. Hydrol., 231-232, 244–254, https://doi.org/10.1016/S0022-1694(00)00198-0, 2000.

Bear, J.: Dynamics of Fluids in Porous Media, American Elsevier Publishing Company, 1972.

Beliaev, A. Y. and Hassanizadeh, S. M.: A Theoretical Model of Hysteresis and Dynamic Effects in the Capillary Relation for Two-phase Flow in Porous Media, Transp. Porous Media, 43, 487–510, 2001.

Beljadid, A., Cueto-Felgueroso, L., and Juanes, R.: A continuum model of unstable infiltration in porous media endowed with an entropy function, Adv. Water Resour., 144, 103 684, https://doi.org/10.1016/j.advwatres.2020.103684, 2020.

Brandhorst, N., Erdal, D., and Neuweiler, I.: Coupling saturated and unsaturated flow: comparing the iterative and the non-iterative approach, Hydrol. E. Sys. Sci., 25, 4041–4059, https://doi.org/10.5194/hess-25-4041-2021, 2021.

Brindt, N. and Wallach, R.: The moving-boundary approach for modeling 2D gravity-driven stable and unstable flow in partially wettable soils, Water Resour. Res., 56(5), e2019WR025 772, https://doi.org/10.1029/2019WR025772, 2020.

Buckingham, E.: Studies on the Movement of Soil Moisture. Bulletin 38., USDA Bureau of Soils, Washington: Government Printing Office, 1907.

Bundt, M., Albrecht, A., Froidevaux, P., Blaser, P., and Flühler, H.: Impact of preferential flow on radionuclide distribution in soil., Environ. Sci. Technol., 44(18), 3895–3899, 2000.

Chen, L., Qiu, Q., Wang, P., Zhang, X., and Zhang, Z.: Visualization study on preferential flow in highly saturated and super hydrophilic porous media by combining dye tracking and infrared imaging, J. Hydrol., 612, 128 077, https://doi.org/10.1016/j.jhydrol.2022.128077, 2022.

Cho, H., G. H. de Rooij, G. H., and Inoue, M.: The Pressure Head Regime in the Induction Zone During Unstable Nonponding Infiltration: Theory and Experiments, Vadose Zone Journal, 4, 908–914, https://doi.org/10.2136/vzj2004.0158, 2005.

Chuoke, R. L., van Meurs, P., and van der Poel, C.: The instability of slow, immiscible, viscous liquid-liquid displacements in permeable media., Trans. AIME., 216, 88–194, 1959.

Cremer, C. J. M., Schuetz, C., Neuweiler, I., Lehmann, P., and Lehmann, E. H.: Unstable Infiltration Experiments in Dry Porous Media, Vadose Zone J., 16 (7), https://doi.org/10.2136/vzj2016.10.0092, 2017.

Cueto-Felgueroso, L. and Juanes, R.: A phase field model of unsaturated flow., Water Resour. Res., 45(10), W10 409, https://doi.org/10.1029/2009WR007945, 2009.

Cueto-Felgueroso, L., Suarez-Navarro, M. J., Fu, X., and Juanes, R.: Numerical Simulation of Unstable Preferential Flow during Water Infiltration into Heterogeneous Dry Soil, Water, 12(3), 909, https://doi.org/10.3390/w12030909, 2020.

de Rooij, G. H.: Modeling fingered flow of water in soils owing to wetting front instability: a review., J. Hydrol., 231-232, 277–294, https://doi.org/10.1016/S0022-1694(00)00201-8, 2000.

Dekker, L. W. and Ritsema, C. J.: How water moves in a water repellent sandy soil: 1. Potential and actual water repellency, Water Resources Research, 30, 2507–2517, https://doi.org/10.1029/94WR00749, 1994.

DiCarlo, D. A.: Experimental measurements of saturation overshoot on infiltration, Water Resour. Res., 40(4), W04 215, https://doi.org/10.1029/2003WR002670, 2004.

DiCarlo, D. A.: Capillary pressure overshoot as a function of imbibition flux and initial water content., Water Resour. Res., 43(8), W08 402, https://doi.org/10.1029/2006WR005550, 2007.

DiCarlo, D. A.: Can continuum extensions to multiphase flow models describe preferential flow?, Vadose Zone J., 9(2), 268–277, https://doi.org/10.2136/vzj2009.0099, 2010.

DiCarlo, D. A.: Stability of gravity-driven multiphase flow in porous media: 40 Years of advancements, Water Resour. Res., 49, 4531–4544, https://doi.org/10.1002/wrcr.20359, 2013.

DiCarlo, D. A., Bauters, T. W. J., Darnault, C. J. G., Steenhuis, T., and Parlange, J.-Y.: Lateral expansion of preferential flow paths in sands., Water Resour. Res., 35(2), 427–434, 1999.

Diment, G. A. and Watson, K. K.: Stability analysis of water movement in unsaturated porous materials: 2. Numerical studies, Water Resources Research, 19, 1002–1010, https://doi.org/10.1029/WR019i004p01002, 1983.

Diment, G. A. and Watson, K. K.: Stability Analysis of Water Movement in Unsaturated Porous Materials: 3. Experimental Studies, Water Resources Research, 21, 979–984, https://doi.org/10.1029/WR021i007p00979, 1985.

Diment, G. A., Watson, K. K., and Blennerhassett, P. J.: Stability analysis of water movement in unsaturated porous materials: 1. Theoretical considerations, Water Resources Research, 18, 1248–1254, https://doi.org/10.1029/WR018i004p01248, 1982.

Egorov, A. G., Dautov, R. Z., Nieber, J. L., and Sheshukov, A. Y.: Stability analysis of gravity-driven infiltrating flow., Water Resour. Res., 39(9), 1266, 2003.

Eliassi, M. and Glass, R. J.: On the continuum-scale modeling of gravity-driven fingers in unsaturated porous media: The inadequacy of the Richards equation with standard monotonic constitutive relations and hysteretic equations of state., Water Resour. Res., 37, 2019–2035, 2001.

Eliassi, M. and Glass, R. J.: On the porous-continuum modeling of gravity-driven fingers in unsaturated materials: Extension of standard theory with a hold-back-pile-up effect, Water Resour. Res., 38(11), 16–1–16–11, https://doi.org/10.1029/2001WR001131, 2002.

Esmaeilpour, M., Ghanbarian, B., Liang, F., and Liu, H.-H.: Scale-dependent permeability and formation factor in porous media: Applications of percolation theory, Fuel, 301, 121 090, https://doi.org/10.1016/j.fuel.2021.121090, 2021.

Ewing, R. P., Hu, Q., and Liu, C.: Scale dependence of intragranular porosity, tortuosity, and diffusivity, Water Resour. Res., 46, https://doi.org/10.1029/2009WR008183, 2010.

Fritz, S.: Experimental investigations of water infiltration into unsaturated soil – analysis of dynamic capillary effects, Diploma Thesis, University of Stuttgart, Germanys, 2012.

Fürst, T., Vodák, R., Šír, M., and Bíl, M.: On the incompatibility of Richards' equation and finger-like infiltration in unsaturated homogeneous porous media, Water Resour. Res., 45(3), W03 408, https://doi.org/10.1029/2008WR007062, 2009.

Genuchten, M. T. V.: A closed-form equation for predicting the hydraulic conductivity of unsaturated soils, Soil Sci. Soc. Am. J., 44, 892–898, https://doi.org/10.2136/sssaj1980.03615995004400050002x, 1980.

Ghanbarian, B., Taslimitehrani, V., Dong, G., and Pachepsky, Y. A.: Sample dimensions effect on prediction of soil water retention curve and saturated hydraulic conductivity, J. Hydrol., 528, 127–137, https://doi.org/10.1016/j.jhydrol.2015.06.024, 2015.

Ghanbarian, B., Taslimitehrani, V., and Pachepsky, Y. A.: Accuracy of sample dimension-dependent pedotransfer functions in estimation of soil saturated hydraulic conductivity, CATENA, 149, 374–380, https://doi.org/10.1016/j.catena.2016.10.015, 2017.

Ghanbarian, B., Esmaeilpour, M., Ziff, R. M., and Sahimi, M.: Effect of Pore-Scale Heterogeneity on Scale-Dependent Permeability: Pore-Network Simulation and Finite-Size Scaling Analysis, Water Resour. Res., 57, e2021WR030 664, https://doi.org/10.1029/2021WR030664, 2021.

Ginzburg, I., Carlier, J.-P., and Kao, C.: Lattice Boltzmann approach to Richards' equation, in: Computational Methods in Water Resources: Volume 1, vol. 55 of *Developments in Water Science*, pp. 583–595, Elsevier, https://doi.org/10.1016/S0167-5648(04)80083-2, 2004.

Glass, R. J. and Yarrington, L.: Analysis of wetting front instability using modified invasion percolation theory, Eos Trans. AGU, 70, 1117, 1989.

Glass, R. J. and Yarrington, L.: Mechanistic modeling of fingering, nonmonotonicity, fragmentation, and pulsation within gravity/buoyant destabilized two-phase/unsaturated flow, Water Resour. Res., 39(3), 1058, https://doi.org/10.1029/2002WR001542, 2003.

Glass, R. J., Parlange, J.-Y., and Steenhuis, T. S.: Wetting front instability as a rapid and farreaching hydrologic process in the vadose zone P.F. Germann (Ed.), Rapid and farreaching hydrologic processes in the vadose zone, J. Contam. Hydrol., 3(2-4), 207–226, https://doi.org/10.1016/0169-7722(88)90032-0, 1988.

Glass, R. J., Oosting, G. H., and Steenhuis, T. S.: Preferential solute transport in layered homogeneous sands as a consequence of wetting front instability., J. Hydrol., 110(1-2), 87–105, https://doi.org/10.1016/0022-1694(89)90238-2, 1989a.

Glass, R. J., Parlange, J.-Y., and Steenhuis, T. S.: Mechanism for finger persistence in homogenous unsaturated, porous media: Theory and verification, Soil Sci., 148(1), 60–70, https://doi.org/10.1097/00010694-198907000-00007, 1989b.

Glass, R. J., Parlange, J.-Y., and Steenhuis, T. S.: Wetting front instability. 2. Experimental determination of relationships between system parameters and two-dimensional unstable flow field behavior in initially dry porous media, Water Resour. Res., 25(6), 1195–1207, https://doi.org/10.1029/WR025i006p01195, 1989c.

Glass, R. J., Cann, S., King, J., Baily, N., Parlange, J.-Y., and Steenhuis, T. S.: Wetting front instability in unsaturated porous media: a three-dimensional study in initially dry sand., Transp. Por. Med., 5, 247–268, 1990.

Glass, R. J., Conrad, S. H., and Peplinksi, W.: Gravity-destabilized nonwetting phase invasion in macroheterogeneousporous media: Experimental observations of invasion dynamics and scale analysis., Water Resour. Res., 36(11), 3121–3137, 2000.

Gomez, H., Cueto-Felgueroso, L., and Juanes, R.: Three-dimensional simulation of unstable gravity-driven infiltration of water into a porous medium., J. Comput. Phys, 238, 217–239, https://doi.org/10.1016/j.jcp.2012.12.018, 2013.

Hassanizadeh, S. M., Celia, M. A., and Dahle, H. K.: Dynamic effects in the capillary pressure-saturation relationship and its impact on unsaturated flow, Vadose Zone J., 1, 38–57, https://doi.org/10.2136/vzj2002.3800, 2002.

Hill, D. E. and Parlange, J. Y.: Wetting front instability in layered soils., Soil Sci. Soc. Am. Proc., 36, 697–702, 1972.

Hunt, A. G., Ewing, R. P., and Horton, R.: What's wrong with soil physics, Soil Sci. Soc. Am. J., 77, 1877– 1887, https://doi.org/10.2136/sssaj2013.01.0020, 2013.

Jang, J., Narsilio, G. A., and Santamarina, J. C.: Hydraulic conductivity in spatially varying media - a pore-scale investigation., Geophysical J. International, 184, 1167–1179, 2011.

Kmec, J.: Semi-continuum model, https://doi.org/10.5281/zenodo.10117915, 2023a.

Kmec, J.: Videos for: Modeling 2D gravity-driven flow in unsaturated porous media for different infiltration rates, https://doi.org/10.5281/zenodo.10090841, 2023b.

Kmec, J., Fürst, T., Vodák, R., and Šír, M.: A semi-continuum model of saturation overshoot in one dimensional unsaturated porous media flow, Scient. Rep., 9, 8390, https://doi.org/10.1038/s41598-019-44831-x, 2019.

Kmec, J., Fürst, T., Vodák, R., and Šír, M.: A two dimensional semi-continuum model to explain wetting front instability in porous media, Scient. Rep., 11, 3223, https://doi.org/10.1038/s41598-021-82317-x, 2021.

Kmec, J., Šír, M., Fürst, T., and Vodák, R.: Semi-continuum modeling of unsaturated porous media flow to explain Bauters' paradox, Hydrol. E. Sys. Sci., 27, 1279–1300, https://doi.org/10.5194/hess-27-1279-2023, 2023.

740   Kmec, J. and Šír, M.: Simulation data for (part 1): Modeling 2D gravity-driven flow in unsaturated porous media for different infiltration rates, https://doi.org/10.5281/zenodo.13768956, 2024a.

Kmec, J. and Šír, M.: Simulation data for (part 2): Modeling 2D gravity-driven flow in unsaturated porous media for different infiltration rates, https://doi.org/10.5281/zenodo.13769113, 2024b.

Kneale, W. R. and White, R. E.: The movement of water through cores of a dry (cracked) clay-loam grassland topsoil, J. Hydrol., 67,
745   361–365, https://doi.org/10.1016/0022-1694(84)90251-8, 1984.

Kutílek, M. and Nielsen, D.: Soil Hydrology, Catena Verlag, Germany, 1994.

Lake, L.: Enhanced Oil Recovery, Prentice Hall: Englewood Cliffs, 1989.

Larson, R. G. and Morrow, N. R.: Effects of sample size on capillary pressures in porous media, Powder Technology, 30(2), 123–138, https://doi.org/10.1016/0032-5910(81)80005-8, 1981.

750   Lenhard, R. J. and Parker, J. C.: A model for hysteretic constitutive relations governing multiphase flow: 2. Permeability-saturation relations., Water Resour. Res., 23(12), 2197–2206, 1987.

Lenormand, R., Touboul, E., and Zarcone, C.: Numerical models and experiments on immiscible displacement in porous media, J. Fluid Mech., 189, 165–187, https://doi.org/10.1017/S0022112088000953, 1988.

Liu, H.-H.: The Large-Scale Hydraulic Conductivity for Gravitational Fingering Flow in Unsaturated Homogenous Porous Media: A Review
755   and Further Discussion, Water, 14, https://doi.org/10.3390/w14223660, 2022.

Liu, H.-H., Zhang, R., and Bodvarsson, G. S.: An active region model for capturing fractal flow patterns in unsaturated soils: Model development, J. Contaminant Hydrol., 80, 18–30, https://doi.org/10.1016/j.jconhyd.2005.07.002, 2005.

Liu, Y., Steenhuis, T. S., and Parlange, J. Y.: Formation and persistence of fingered flow fields in coarse grained soils under different moisture contents., J. Hydrol., 159, 187–195, https://doi.org/10.1016/0022-1694(94)90255-0, 1994.

760   Liu, Y., Zhang, S., and Liu, H.-H.: The relationship between fingering flow fraction and water flux in unsaturated soil at the laboratory scale, J. Hydrol., 622, 129 695, https://doi.org/10.1016/j.jhydrol.2023.129695, 2023.

McNamara, H.: An estimate of energy dissipation due to soil-moisture hysteresis, Water Resources Research, 50, 725–735, https://doi.org/10.1002/2012WR012634, 2014.

Mishra, B. K. and Sharma, M. M.: Measurement of pore size distributions from capillary pressure curves, American Institute of Chemical
765   Engineers Journals, 34(4), 684–687, https://doi.org/10.1002/aic.690340420, 1988.

Mualem, Y.: A new model for predicting the hydraulic conductivity of unsaturated porous media, Water Resour. Res., 12(3), 513–522, https://doi.org/10.1029/WR012i003p00513, 1976.

Ommi, S. H., Sciarra, G., and Kotronis, P.: A phase field model for partially saturated geomaterials describing fluid–fluid displacements. Part I: The model and one-dimensional analysis, Adv. Water Resour., 164, 104 170, https://doi.org/10.1016/j.advwatres.2022.104170, 2022a.

770   Ommi, S. H., Sciarra, G., and Kotronis, P.: A phase field model for partially saturated geomaterials describing fluid–fluid displacements, Part II: Stability analysis and two-dimensional simulations, Adv. Water Resour., 164, 104 201, https://doi.org/10.1016/j.advwatres.2022.104201, 2022b.

Pales, A. R., Li, B., Clifford, H. M., Kupis, S., Edayilam, N., Montgomery, D., W.-Z. Liang, M. Dogan, N. T., Martinez, N., Moysey, S., Powell, B., and Darnault, C. J. G.: Preferential flow systems amended with biogeochemical components: imaging of a two-dimensional study, Hydrol. E. Sys. Sci., 22(4), 2487–2509, https://doi.org/10.5194/hess-22-2487-2018, 2018.

Parker, J. C. and Lenhard, R. J.: A model for hysteretic constitutive relations governing multiphase flow: 1. Saturation-pressure relations, Water Resour. Res., 23(12), 2187–2196, https://doi.org/10.1029/WR023i012p02187, 1987.

Parlange, J. and Hill, D. E.: Theoretical analysis of wetting front instability in soils., Soil Sci., 122(4), 236–239, 1976.

Perfect, E., McKay, L. D., Cropper, S. C., Driese, S. G., Kammerer, G., and Dane, J. H.: Capillary Pressure–Saturation Relations for Saprolite: Scaling With and Without Correction for Column Height, Vadose Zone J., 3(2), 493–501, https://doi.org/10.2136/vzj2004.0493, 2004.

Pražák, J., Šir, M., and Tesař, M.: Retention cruve of simple capillary networks, J. Hydrol. Hydromech., 47, 117–131, 1999.

Primkulov, B. K., Talman, S., Khaleghi, K., Shokri, A. R., Chalaturnyk, R., Zhao, B., MacMinn, C. W., and Juanes, R.: Quasistatic fluid-fluid displacement in porous media: Invasion-percolation through a wetting transition, Phys. Rev. Fluids, 3, 104 001, https://doi.org/10.1103/PhysRevFluids.3.104001, 2018.

Rezanezhad, F., Vogel, H.-J., and Roth, K.: Experimental study of fingered flow through initially dry sand, Hydrol. E. Sys. Sci. D., 3(4), 2595–2620, hal-00298761, 2006.

Richards, L. A.: Capillary conduction of liquid through porous media, Physics, 1, 318–333, https://doi.org/10.1063/1.1745010, 1931.

Roche, W. J., Murphy, K., and Flynn, D. P.: Modelling preferential flow through unsaturated porous media with the Preisach model and an extended Richards Equation to capture hysteresis and relaxation behaviour, Journal of Physics: Conference Series, 1730, 012 002, https://doi.org/10.1088/1742-6596/1730/1/012002, 2021.

Saffman, P. G. and Taylor, G. I.: The penetration of a fluid into a porous medium or Hele-Shaw cell containing a more viscous liquid, Proceedings of the Royal Society of London. Series A. Mathematical and Physical Sciences, 245, 312–329, https://doi.org/10.1098/rspa.1958.0085, 1958.

Schaap, M. and Leij, F.: Improved Prediction of Unsaturated Hydraulic Conductivity with the Mualem-van Genuchten Model, Soil Science Society of America Journal, 64, 843–851, https://doi.org/10.2136/sssaj2000.643843x, 2000.

Schroth, M., Ahearn, S., Selker, J., and Istok, J.: Characterization of Miller-similar silica sands for laboratory hydrologic studies, Soil Sci. Soc. Am. J., 60, 1331–1339, https://doi.org/10.2136/sssaj1996.03615995006000050007x, 1996.

Schweizer, B.: Hysteresis in porous media: Modelling and analysis, Interfaces and Free Boundaries, 19 (3), 417–447, https://doi.org/10.4171/IFB/388, 2017.

Selker, J., Parlange, J.-Y., and Steenhuis, T.: Fingered flow in two dimensions: 2. Predicting finger moisture profile, Water Resour. Res., 28, 2523–2528, https://doi.org/10.1029/92WR00962, 1992.

Sililo, O. T. N. and Tellam, J. H.: Fingering in Unsaturated Zone Flow: A Qualitative Review with Laboratory Experiments on Heterogeneous Systems, Ground Water, 38(6), 864–871, https://doi.org/10.1111/j.1745-6584.2000.tb00685.x, 2005.

Silva, M. L. N., Libardi, P. L., and Gimenes, F. H. S.: Soil water retention curve as affected by sample height, Rev. Bras. Cienc. Solo, 42, e0180 058, https://doi.org/10.1590/18069657rbcs20180058, 2018.

Smith, W. O.: Infiltration in sands and its relation to groundwater recharge., Water Resour. Res., 3(2), 539–555, 1967.

Steinle, R. and Hilfer, R.: Hysteresis in relative permeabilities suffices for propagation of saturation overshoot: A quantitative comparison with experiment., Phys. Rev. E, 95, 043 112, 2017.

Sutherland, K. and Chase, G.: Filters and Filtration Handbook, 5th Ed., Elsevier, Oxford, 2008.

810 Tesař, M., Šír, M., Pražák, J., and Lichner, L.: Instability driven flow and runoff formation in a small catchment, Geologica Acta, 2(1), 147–156, https://doi.org/10.1344/105.000001435, 2004.

Vafai, K.: Porous media: Applications in Biological Systems and Biotechnology., Taylor & Francis., 2011.

Visintin, A.: Differential models of hysteresis, New York: Springer, 1993.

Vodák, R., Fürst, T., Šír, M., and Kmec, J.: The difference between semi-continuum model and Richards' equation for unsaturated porous
815 media flow, Scient. Rep., 12, 7650, https://doi.org/10.1038/s41598-022-11437-9, 2022.

Wei, H., Zhu, X., Liu, X., Yang, H., Tao, W.-Q., and Chen, L.: Pore-scale study of drainage processes in porous media with various structural heterogeneity, International Communications in Heat and Mass Transfer, 132, 105 914, https://doi.org/10.1016/j.icheatmasstransfer.2022.105914, 2022.

Wei, Y., Cejas, C. M., Barrois, R., Dreyfus, R., and Durian, D. J.: Morphology of Rain Water Channeling in Systematically Varied Model
820 Sandy Soils, Physical Review Applied, 2, 044 004, https://doi.org/10.1103/PhysRevApplied.2.044004, 2014.

White, J. A., Borja, R. I., and Fredrich, J. T.: Calculating the effective permeability of sandstone with multiscale lattice Boltzmann/finite element simulations, Acta Geotechnica, 1, 195–209, https://doi.org/10.1007/s11440-006-0018-4, 2006.

Xiong, Y.: Flow of water in porous media with saturation overshoot: A review, J. Hydrol., 510, 353–362, https://doi.org/10.1016/j.jhydrol.2013.12.043, 2014.

825 Yao, T. and Hendrickx, J. M. H.: Stability of wetting fronts in dry homogeneous soils under low infiltration rates, Soil Sci. Soc. Am. J., 60, 20–28, https://doi.org/10.2136/sssaj1996.03615995006000010006x, 1996.

Zhou, D. and Stenby, E. H.: Interpretation of capillary-pressure curves using invasion percolation theory, Transport Porous Med., 11, 17–31, https://doi.org/10.1007/BF00614632, 1993.

Šimůnek, J. and Suarez, D. L.: Two-dimensional transport model for variably saturated porous media with major ion chemistry., Water
830 Resour. Res., 30, 1115–1133, 1994.