

# A new 3D full-Stokes calving algorithm within Elmer/Ice (v9.0)

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**Abstract.** A new calving algorithm is developed in the glacier model Elmer/Ice that allows unrestricted calving and terminus advance in 3D. The algorithm uses the meshing software Mmg to implement anisotropic remeshing and allow mesh adaptation at each timestep. The development of the algorithm along with the implementation of the crevasse depth law produces a new full-Stokes calving model capable of simulating calving and terminus advance across an array of complex geometries. Using a synthetic tidewater glacier geometry the model is tested to highlight the numerical model parameters that can alter calving when using the crevasse depth law. For a system with no clear attractor at a pinning point, model timestep and mesh resolution are shown to alter the simulated calving. In particular vertical mesh resolution has a large impact, increasing calving, as the frontal bending stresses are better resolved. However, when the system has a strong attractor, provided by basal pinning points, numerical model parameters have a limited effect on the terminus evolution. The new algorithm is capable of implementing unlimited terminus advance and retreat as well as unrestricted calving geometries, applying any vertically varying melt distribution to the front, use in conjunction with any calving law or potentially advecting variables downstream. In overcoming previous technical hurdles, the algorithm opens up the opportunity to improve both our understanding of the physical processes and our ability to predict calving.

## 1 Introduction

One of the largest sources of uncertainty in predictions of future sea level rise is the magnitude of losses from the Greenland and Antarctic Ice Sheets via ice discharge and iceberg calving (IPCC, 2023). In recent years, significant advances have been made in understanding calving processes and their relationship with ice dynamics (Benn and Åström, 2018), but this has yet to translate into the adoption of a reliable, universal ‘calving law’ in continuum ice sheet models. This largely reflects the contrast between the complexity of calving processes, which are influenced by stresses in three dimensions, and the simplified, vertically-integrated stress fields required in model simulations of long-term and large-scale ice sheet evolution. There is a need, therefore, to develop robust, physically based calving models in three dimensions, which can then be used to develop simpler calving parameterisations required for ice sheet models that apply approximations to the Stokes equations.

Currently there are two main types of 3D calving modelling methods with the capability to simulate the evolution of glacier calving fronts through time. The first is Discrete Element Modelling (DEM), which is commonly referred to as particle modelling. An example model is the Helsinki Discrete Element Model (HiDEM) which predicts calving from first principles by treating the ice as elastically-bonded individual particles where the bonds fracture if a failure threshold is exceeded (Åström et al., 2014). HiDEM can accurately simulate a wide range of individual calving styles but is unsuitable for modelling longer-term glacier evolution because it does not include viscous deformation (Åström et al., 2013; van Dongen et al., 2020; Benn et al., 2023). Conversely, glaciers can be modelled as continua in 3D using a model such as Elmer/Ice. Elmer/Ice is a Finite Element Method (FEM) model that treats the ice sheet as a continuum and solves the flow and stress fields using the full Stokes equations. Using the position-based crevasse-depth calving law, Elmer/Ice has been shown to accurately predict calving front evolution at Store Glacier (Sermeq Kujalleq) in response to changing fjord conditions (Todd et al., 2018; Benn et al., 2023). Although the computational requirements of Elmer/Ice remain high, they are significantly lower than the requirements of HiDEM and multiple years of calving front evolution can be simulated with reasonable computational resources (Todd et al., 2018, 2019).

However, several issues remained with using the Elmer/Ice calving model as implemented by Todd et al. (2019) to run multi-year simulations. Firstly, the lateral corners of the ice front (i.e., where the calving front meets the fjord wall) were fixed in place. This is a reasonable assumption for stable glaciers such as Store Glacier (Sermeq Kujalleq) where only seasonal movement of the calving front needs to be captured (Todd et al., 2018). However, in the most extreme warming experiments conducted for Store Glacier the simulated retreat caused model breakdown due to the restriction that the lateral boundaries of the front were fixed in space (Todd et al., 2019). The assumption of fixed lateral boundaries does not hold for fast-retreating glaciers such as Jakobshavn Isbræ (also Sermeq Kujalleq in Greenlandic) where the terminus position may change year on year by several kilometers (Joughin et al., 2020). Furthermore, the Todd et al. (2019) model relied on a 2D vertically extruded mesh, which is problematic where the ice-front is non-vertical, such as where submarine melt produces undercutting (Todd et al., 2019). As a result, ice-front retreat leads to degeneracy of the mesh elements, which causes irrecoverable breaking of the calving model. The importance of submarine melt is often associated with its known ability to increase calving (O’Leary and Christoffersen, 2013). However, investigating the importance of submarine plume melt undercutting on calving is only possible in a 3D calving model because of the importance of terminus geometry and as such has remained understudied. Finally, Todd et al. (2018) assumed projectability of the calving front and calved icebergs, namely the post-calving ice front does not contain any complex re-entrants and can therefore be projected onto a transverse plane. This is clearly not always the case at real glaciers, particularly those with complex front geometries such as Bowdoin Glacier (Kangerluarsuup Sermia), Greenland.

This paper presents a new calving algorithm that rectifies these limitations. It utilises 3D remeshing which is implemented using the open-source remeshing algorithm Mmg (Dapogny et al., 2014). The level set method is used to define areas of calved ice (Osher and Fedkiw, 2001; Sethian, 1999) and the new calving front is physically implemented using Mmg. In this paper, the key steps and processes involved in the new algorithm are outlined, leading into an explanation of how the algorithm resides

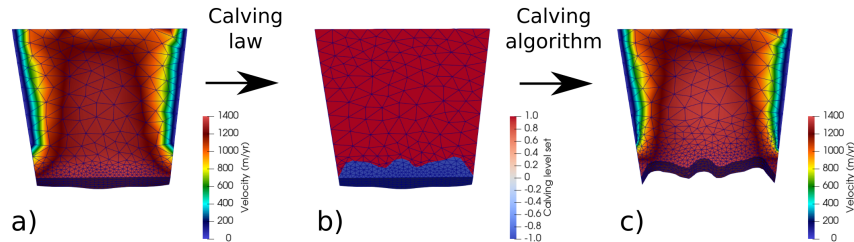
within the framework of a typical glacier simulation. The capabilities of the new algorithm are described and illustrated using a set of synthetic glacier geometries with a particular focus on the use of the crevasse depth calving law. The new algorithm allowed a more detailed study using the crevasse depth law to assess the influence of model setup on predicted calving along with demonstrating the theoretical implications of the calving law. Furthermore, the often-neglected numerical model parameters with the potential to alter modelled calving are investigated, providing a sound basis for rational parameter choices in future studies using real world domains. The impact of each numerical model parameter is investigated and discussed within the context of the crevasse depth calving law. Finally, the capabilities of the new model are concluded.

## 2 Modelling calving in a continuum

When modelling calving it is important to distinguish between the calving law, calving algorithm and calving model (Fig. 1). The *calving law* is the function by which calving is predicted, in this case the crevasse depth (CD) law (Nick et al., 2010; Benn et al., 2007). The *calving algorithm* takes the prediction from the calving law and implements this within the model, ultimately leading to the removal of calved ice and the resulting alteration of the domain. Consequently, a calving algorithm is not tied to a particular calving law. Instead, it is the technical implementation, within a given ice-flow model, of a theoretical calving event provided by the independent calving law. Putting the calving law and calving algorithm together gives us the *calving model*. This is an important distinction to make within glacier modelling, as both calving laws and calving algorithms are limited in their functionality at this time. Calving laws in 3D can be thought of as resulting from a combination of our current understanding of the physical processes behind calving and our ability to condense them into a mathematical function compatible with the ice-flow model. Calving algorithms can be viewed as the capability of our models to implement the calving law's prediction on a given domain by removing the predicted calved ice. This is surprisingly non-trivial especially for more advanced calving laws that can produce convoluted calving geometries. This paper focuses on the development of a new calving algorithm that substantially increases our ability to realise calving in a 3D glacier model. This is primarily motivated by our previous limited ability to simulate substantial retreat and advance of fast-flowing glaciers in 3D. The objective is that once the previous technical hurdles have been overcome, our understanding and capability to test calving laws and physical processes can correspondingly be improved.

In this paper, we implement the CD calving law in connection to our newly developed algorithm, defining our calving model. The CD law was chosen because it is the only currently available calving law that is explicitly based on physical processes, with proven capability in predicting calving on Greenland tidewater glaciers (Todd et al., 2019; Amaral et al., 2020; Benn et al., 2023). The CD law is based on the idea that calving occurs when crevasses penetrate the full thickness of the glacier, or some prescribed part thereof (Benn et al., 2007; Nick et al., 2010). For simplicity and ease of implementation, crevasse depth is predicted using the zero-stress approach introduced by Nye (1957) and modified by Todd et al. (2018), which assumes

# Calving model



**Figure 1.** Illustrating the difference between the calving model, algorithm and law. a) The initial flow solution. The calving law predicts calving based off the solved variables on the initial geometry. b) The predicted calving based on the calving law. This prediction is in the form of a level set variable which is then fed to the calving algorithm. Nodes with a negative value (in blue) show areas to be calved and those with positive values (in red) show the remaining glacier domain. c) The new geometry with all the variables interpolated across (from a)) showing the interpolated velocity field. The calving model represents the combination of both the calving law and algorithm allowing you to go from a) to c).

that fracturing is possible wherever the largest principal stress  $\sigma_1$  (the largest eigenvalue of the Cauchy stress) is extensional (positive). This approach assumes negligible stress concentrations at crack tips, which is reasonable for fields of closely-spaced crevasses. Calving is thus a function of the large-scale stress field of the glacier and particularly regions of high extensional stress. It is important to emphasise that the CD law does not predict individual crevasses or impose discontinuities on the model glacier. Rather, the zero-stress function defines bounding surfaces of crevasse fields (i.e. the base in the case of surface crevasses fields, the top in the case of basal crevasses fields), which are then used to predict the position of the glacier front at a particular timestep.

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### 3 Implementation of calving in a continuum model

Implementing any calving law in a continuum model requires the new position of the ice front to be prescribed at each timestep. In vertically-integrated models this is relatively straightforward because the ice front is by definition vertical before and after calving and the model physics can be solved on a 2D mesh. The problem becomes substantially more difficult in three dimensions when the front can have complex vertical profiles due to melt-undercutting or deformation.

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Representing the change in the glacier domain via calving can either be done by an explicit change in a domain (i.e., modification of a mesh) or by using a ‘calving variable’ that represents the new terminus position. Computationally it is much cheaper to use a variable since only this variable needs to be manipulated prior to calculation of the new flow solution. This approach

110 is becoming increasingly popular, especially among lower-level models where computational efficiency is the priority (e.g.,  
Bondzio et al., 2016). Usually, this is done using a level set method (Osher and Fedkiw, 2001; Sethian, 1999; Bondzio et al.,  
2016) where a surface is defined from a signed distance and moves based on an advection equation. The main drawback of this  
method is the need to solve intra-element dynamics if a level set surface is to be followed exactly since the hyperplane will  
cut through any form of element. No current model has the capability to solve this problem, which would require boundary  
115 conditions to be applied across the hyperplane (e.g., Bondzio et al., 2016). Instead, intersected elements are marked, those  
beyond are marked as ice free and those in the glacier domain marked as ice. Boundary conditions are instead applied across  
the ice-free boundary of the intersected elements, so the actual modelled calving front generally deviates from the level set  
surface (Bondzio et al., 2016).

120 Alternatively, explicit domain change is more computationally expensive and significantly more complicated to implement.  
Simplistically, this can be achieved by deforming the mesh (i.e., moving nodes but keeping the mesh topology intact), but  
this can only be done for non-complex geometries and becomes less feasible with more dimensions and more significant de-  
formations. Complex problems require complete remeshing where element edges are realigned along the new calving front.  
Todd and Christoffersen (2014) produced a flowline 2D model with complete remeshing that allowed a non-vertical terminus,  
125 but concluded that all three dimensions needed to be considered for accurate calving representation. A similar approach was  
taken by Berg and Bassis (2022) where the vertical dimension was seen as imperative to the study of calving dynamics and ice  
history. The most advanced example is the 3D extruded remeshing implemented by Todd et al. (2018). Despite the increased  
complexity, several advantages emerge from this approach. For example, complete remeshing allows finer resolution to be  
maintained near the calving front if the terminus moves greatly over the course of the simulation. Secondly, the position of  
130 the calving front resembles that defined by the calving law much more closely, allowing implementation of more complicated  
position-based calving laws.

The CD law is designed to predict the position of the ice front at each timestep based on the ice geometry and stress field.  
However, the predicted front may also be influenced by numerical model parameters such as timestep and mesh resolution.  
135 Like any computational method, FEM requires a model timestep that represents the time gap between the calculation of the  
model physics. Usually, the timestep must be small enough to meet the Courant-Friedrichs-Lewy (CFL) condition (Courant  
et al., 1928) as it applies to the evolution of the free surfaces. Furthermore, changes in the timestep can alter the numerical  
solution with coarser timesteps often reducing accuracy. Similarly, changes in the mesh resolution can alter the numerical  
solution of the problem with smaller elements allowing greater gradients across the domain. Limited research has reported on  
140 how numerical model parameters influence model results, the experiments conducted in this study provided an opportunity to  
analyse the affect of these parameters on both the new calving algorithm and the CD law.

## 4 The new calving algorithm: methods and capabilities

145 In this paper, we illustrate the features and capability of the new algorithm using a simple synthetic geometry broadly representative of a Greenlandic glacier but with reduced computational costs. The synthetic glacier is dominated by strong shear margins with super buoyant sections of the terminus. All of the experiments in this paper were completed on a local desktop with Intel Xeon Processor E3-1245 v5 using eight threads. The control simulation took 45 minutes with simulation time scaling uniformly with an increase in timesteps and exponentially with an increase in 3D mesh resolution. The full geometry and input  
150 files can be found in the official Elmer/Ice repository (<https://github.com/ElmerCSC/elmerfem>) or in Wheel (2023a) in the test case `elmerice/Tests/Calving3D_lset`. Full details of the associated geometry and boundary conditions are detailed in Appendix A. Rather than giving the unrestricted details and coding design associated with the algorithm, the key methodological choices and model capabilities are presented here highlighting their benefits and drawbacks. A comprehensive breakdown of the algorithm and minor coding choices is provided in the supplement or in Wheel (2023b). Detailed user documentation are provided  
155 in the Elmer/Ice repository or in Wheel (2023a).

### 4.1 Explicit domain modification

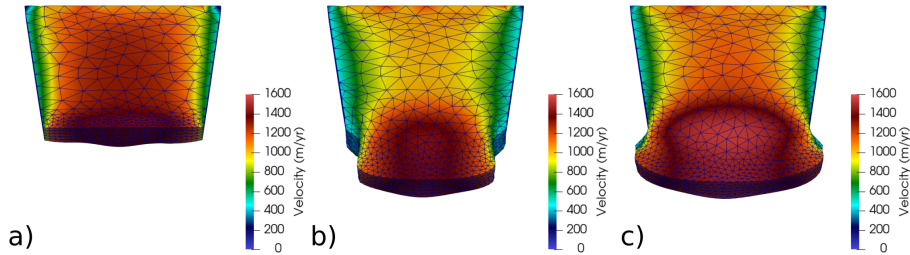
The domain of a glacier actively evolves through three mechanisms: 1) the movement of the glacier terminus via ice flow countered by frontal melting; 2) retreat of the terminus via iceberg calving and 3) the elevation of the surface and base of the glacier can evolve through ice thickness advection, melting or accumulation. The implementation of the movement of  
160 the terminus, base and surface boundaries is achieved through deforming the mesh, while calving is implemented through remeshing.

### 4.2 Mesh deformation: terminus advance and free surfaces

The Lagrangian advection of the terminus for a given timestep is computed using the velocity solution. For constricted domains such as glaciers flowing down a fjord, the advance of the lateral margins take account of local geometry. Across the majority  
165 of the terminus, the displacement vector ( $\mathbf{d}$ ) is the velocity multiplied by the timestep. However, where the calving front and the lateral boundaries meet, the displacement vector cannot be based purely on the velocity since the terminus advance must be constrained by the fjord walls and by imposed kinematic constraints. This could be solved as a contact problem analogously to the grounding line (Durand et al., 2009), but this would increase computational requirements as the non-linearity of the problem increases. Instead, fjord walls are user defined before the simulation. For each timestep after the calculation of the  
170 flow solution, the velocity is reprojected along the predefined fjord walls (see Fig. 5 for solver order). As such the displacement in the x and y plane at the lateral margins ( $\mathbf{d}_l$ ) is defined by

$$\mathbf{d}_l = |\mathbf{u}| \times \mathbf{f}, \quad (1)$$

where  $\mathbf{u}$  is the velocity and  $\mathbf{f}$  is the direction of the fjord wall. Although this is very computationally efficient, it provides some sources of error. Firstly, the normal vector is poorly defined on the lateral corners where the terminus meets the lateral



**Figure 2.** Two examples showing the capabilities of the front advance routines over one year simulations. Calving is suppressed for both simulations so the resultant geometry is purely from glacier advance. a) The initial geometry prior to simulation. Final geometry and velocity field for b) a fjord with a narrowing and c) a widening fjord. The full simulations can be seen in Supplementary Videos 1 and 2.

175 boundary. Additionally, even though the velocity follows the local lateral boundary, this does not guarantee that it will follow the boundary further downstream as the glacier advances. Therefore, it may lead to artificial mass change as the imposed kinematics of the lateral boundary corners are not guaranteed to obey the incompressibility condition. This artificial mass change is negligible, especially when compared to mass changes associated with remeshing. This provides a much better solution than having the lateral corners fixed in place and time (Todd et al., 2018). The mesh is deformed horizontally using the predicted

180 front advance as a boundary condition. Similarly, the kinematic free surface is calculated for the surface and base before the mesh is deformed vertically using these solutions. This accounts for the glacier surface mass balance, while grounding-line evolution is solved as a contact problem (Durand et al., 2009). In order to simulate terminus advance through complex fjord geometries, the model has the capability to transfer boundary elements from the terminus to the lateral boundary. For example, when the glacier advances down a narrowing fjord, the Lagrangian movement of the terminus means it will make contact with

185 the prescribed fjord wall. Here, if all nodes in a boundary element reach the fjord wall, the model will change the conditions applied to the element from those of the terminus to those of the lateral boundary. The model can replicate realistic advance at widening and narrowing fjords whilst maintaining the appropriate boundary assignment (Fig. 2).

Similarly to the terminus advance, the retreat due to submarine melting can be applied to the calving face. It is applied as

190 a scalar variable where the direction of melt is always assumed to be normal to the terminus nodes based on adjacent elements. Following the Lagrangian implementation of the glacier advance, melt is prescribed such that

$$\mathbf{d} = \mathbf{u} - \hat{\mathbf{m}}, \quad (2)$$

$$\mathbf{d}_t = |\mathbf{u}| \times \mathbf{f} - \hat{\mathbf{m}}, \quad (3)$$

195 where  $\hat{\mathbf{m}}$  is the melt normal to the front. As the terminus normal will not follow the lateral boundary some artificial mass change is introduced at the lateral boundary corners. Again, this is negligible when compared to mass changes from remeshing. Melting

with any vertical or horizontal profile can be implemented and the calculation of which is independent of the calving algorithm. Given the simplicity of this method there are very few issues that can arise when applying melt. Degenerate elements will only be produced if the melt per timestep is larger than the element length in the normal direction. In other words, Eq. 2 must comply with the CFL condition for the front.

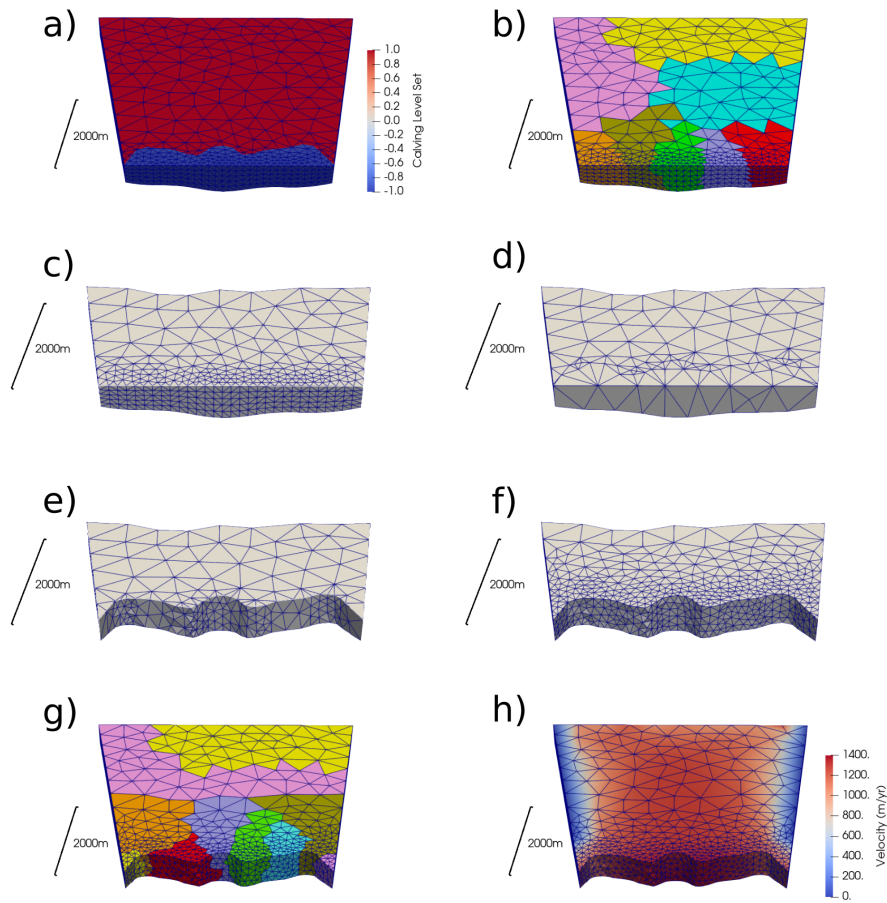
### 4.3 Calving through remeshing

The calving algorithm is defined as the implementation of calving, or the removal of ice from the glacier front. It takes a level set or signed distance variable where the zero contour is the new calving front to produce a new mesh onto which all the model variables are interpolated. Importantly, the new calving algorithm is not limited by iceberg or frontal geometries and consequently is not tied to a particular calving law. Given its physical basis and use as a position-based law, the CD law is a particularly complex calving law to apply in a 3D model. It therefore provides a high benchmark and simpler rate based laws could easily be applied. The CD law is implemented following Todd et al. (2018) but improvements allowing non-projectible calving have been made through use of a level set function (Osher and Fedkiw, 2001; Sethian, 1999). For further details see the user documentation detailed in the Code and Data Availability. To overcome these issues, the remeshing software Mmg (version 5.5.4 or later) is used in the new calving algorithm to produce a fully 3D domain without the need for vertical extrusion (Dapogny et al., 2014). To use the calving algorithm, Mmg must be compiled and linked with Elmer.

Currently, remeshing is completed in two separate steps. The first step realigns element edges along the zero level set contour. This will be referred to as implementing the level set variable. The second stage is complete anisotropic remeshing to improve the mesh quality where a user defined aspect ratio produces elongated elements in the horizontal plane. The full remeshing algorithm is visualized in Fig. 3 and outlined in Fig. 4. To reduce the computational requirements, only the area within a user defined distance from the terminus is remeshed. In a parallel run the Elmer mesh partitions must first be gathered onto one process (Fig. 3b-c) as Mmg must be run in serial. Essentially, for both remeshing stages, the nodes and elements on the upstream partition boundary of the gathered mesh are fixed and cannot be altered. This means that when converted back into the Elmer mesh format, the new mesh still has the same partition boundaries with the upstream parts of the mesh which have not been altered. In a parallel run, the gathered mesh must be redistributed using the library Zoltan (Devine et al., 2009) or using ParMetis which provides much better balancing for larger jobs. A rebalancing algorithm aims to rebalance the mesh evenly in terms of computational requirements among all active processes whilst trying to reduce parallel communication (Fig. 3f). After the rebalancing, variables from the old mesh are interpolated across the new mesh in parallel (Figs. 3, 4). Surface and bottom boundary variables are projected from the old mesh. Both surfaces must maintain projectability to allow the free surface to be solved. However, with the possibility of complete remeshing, there are occasionally sections of the surface and bottom boundaries that are not covered by the old mesh. Nodes here are individually extrapolated.

Iceberg calving and fracture happens far below the typical glacier model timesteps and consequently, secondary calving is often omitted from calving simulations. Secondary calving occurs when iceberg removal leads to a geometry change that is





**Figure 3.** A visualisation of the steps involved during remeshing in the calving algorithm for a simulation run on eight cores. a) The 3D glacier mesh showing the calving level set variable defining the predicted calving. The reduced range of the calving level set variable is just to clearly show the predicted calving. b) The same distributed mesh showing the eight partitions. c) The gathered mesh on one core. The upstream size is defined by the user defined remeshing distance from the new calving front. d) The first remeshing step where element edges are aligned along the new calving front as predicted by the calving level set variable shown in a). e) After the first remeshing step the nodes with negative calving level set values are cut (e.g. glacier ice that would calve as icebergs is removed). f) The second remeshing step where the mesh quality is improved. g) After remeshing the mesh is rebalanced so each partition has a computationally equivalent partition. h) The variables are interpolated to the new mesh in parallel and the old mesh is deallocated from memory.

inherently mechanically unstable so produces further calving. To overcome this issue, the new calving algorithm can insert additional small timesteps, to facility secondary calving where the new glacier geometry is assessed to see if it produces further calving. The size of the additional timestep and calving volume threshold to invoke extra timesteps are user defined. These additional calving timesteps allow the calving cycle to be complete prior to the model moving onto the next true timestep.

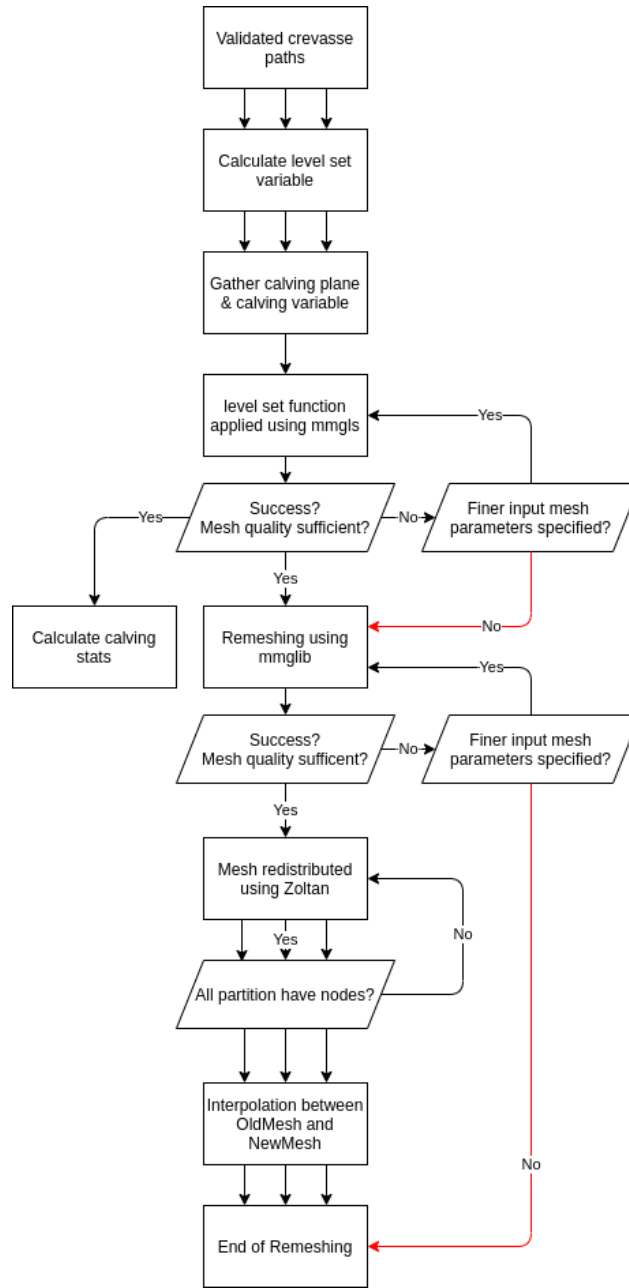
#### 235 4.4 Robustness of algorithm

Given the complexity of potential geometries arising from calving at a tidewater glacier, it can be expected that instances of remeshing failure will occur during simulations. Remeshing failure can be defined as the inability to produce a mesh of sufficient quality to allow the continuation of the simulation. It is not possible to prevent this for every scenario that may arise, so additional focus must be placed on the robustness of the calving algorithm to cope with remeshing failure. A major advantage  
240 of performing the level set variable implementation and anisotropic remeshing in two steps is the ability to isolate a source of potential failure. If there is an issue with either step, the process is retried with finer mesh input parameters. This often results in success (Fig. 4). This feature allows the user to specify multiple input parameters, such as the minimum element edge length, which can be iterated through until remeshing success occurs.

245 Remeshing failure can occur for several reasons. The first reason is that Mmg is unable to return a mesh. Secondly, remeshing failure can occur if the element quality does not meet the user defined minimum quality. However, there can be times when level set implementation fails for all user specified input parameters (Fig. 4). Calving cannot occur in this case since the new terminus boundary is not defined. Remeshing still occurs to try to improve the mesh quality. Similarly, if remeshing fails on all input parameters, calving does not occur even if the level set variable implementation has been successful (Fig. 4). Both situa-  
250 tions can lead to the glacier falsely advancing. This is not a major issue as calving will likely occur on the subsequent time steps.

A final potential issue with the remeshing can result if a mesh passes through the various quality checks but has some physical imperfection or poor element quality that leads to problems in the Stokes solver. Element quality checks attempt to prevent such instances but are not fool proof. Such imperfections often cause unrealistic velocity solutions leading to an exaggerated  
255 stress distributions that in turn can cause unrealistic calving events. An additional step in the calving algorithm has been added to check the new flow solution. The velocity convergence of the velocity solution, maximal velocity and divergence from the previous timestep are checked against user defined limits, which ensures that a mesh imperfection does not slip through. Finally, model check pointing, by regularly saving the model state, allows for easy recovery in case of simulation breakdown.

260 If the flow solution is determined to be inadequate (non-converged), the mesh deforming and calving solvers are suppressed except the anisotropic remeshing routines. The mesh input variables are refined and the flow solution residuals are removed, so that the velocity is calculated from scratch. These refinements result in a finer mesh and an extension of the remeshing distance. The latter is extended in case the fixed elements are causing the flow-convergence issue. Additionally, the model time is set back to the time at the start prior to the current, problematic, time step plus one second (Fig. 5). The additional second allows  
265 the time dependent solvers to be rerun while changing the model time an insignificant amount in a glacier context. An extra time step is added to the required time steps for the simulation. Following the subsequent remeshing step the mesh quality will usually have improved enough to provide a flow solution. At this point, the solver will unpause the mesh deforming and calving solvers and reset the mesh input variables to their original values (Fig. 5).



**Figure 4.** Diagram outlining the steps involved in the calving algorithm. A visual representation of an example simulation can be seen in Fig. 3. Red arrows indicate paths where calving is suppressed. Single arrows indicate stages which occur in serial and three arrows indicate stages that can occur in parallel.

## 5 Typical simulation

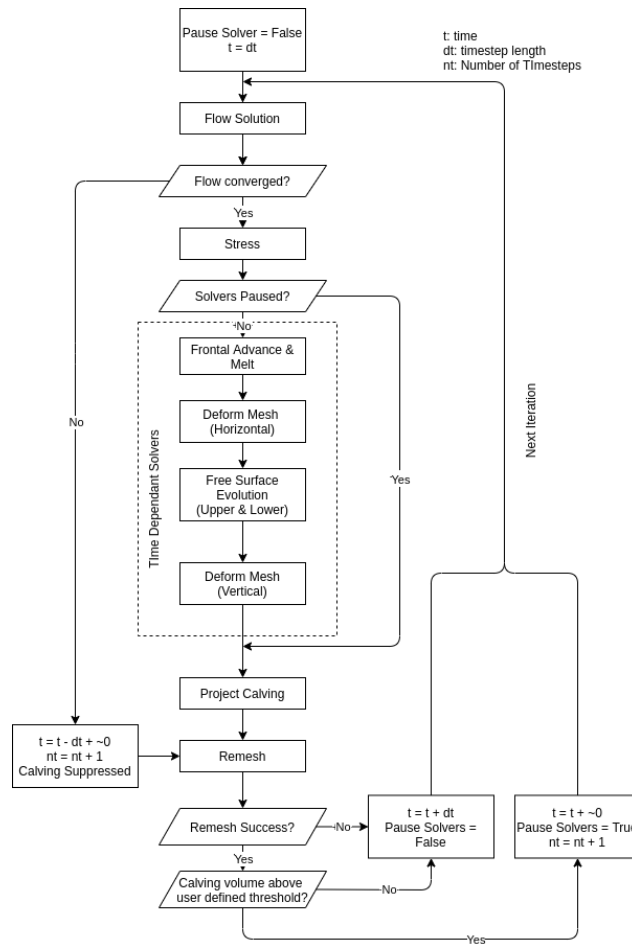
270 A typical simulation is now outlined, putting together the new advance, calving and remeshing mechanisms (Fig. 5). After model initialization and set up, the solvers follow the new algorithm. First, the velocity field is solved. It is then checked to assess it for any abnormalities. If abnormalities exist, the recovery mechanisms outlined above are followed.

Assuming the flow solution converges, it is used to solve the stress fields from which calving is predicted by the CD law. 275 After the calving prediction, the top and bottom free surfaces are solved using the built-in Elmer free-surface solver. As outlined previously the terminus advance is calculated from the flow solution. Using these variables the glacier mesh is deformed both vertically and horizontally. Ideally, the flow solution would be recalculated after the front adjustment since the calculated velocities are based on a different geometry. This means that, currently, the calving prediction is based on the stress field that was calculated for a slightly different geometry. This purely explicit in time approach is used currently to save computational 280 requirements. Unless timesteps are extremely large for the size of domain this should not impact the results.

After the mesh deformation stage, the calving algorithm is called. If remeshing is unsuccessful at any stage, calving is suppressed and the model moves onto the next time step. Following recovery through successful remeshing, any solvers that were previously paused are turned back on and the timestep is checked to make sure it is the original input. If remeshing is initially 285 successful but only insignificant calving occurs the timestep is not altered. However, if an iceberg calves above the user defined threshold, mesh deforming solvers are paused, the timestep is reduced and an additional timestep is added to the model to allow the simulation to run for the required time (Fig. 5). To summarise, the model timestep can be altered for two reasons: 1) as a safety check following an inadequate flow solution allowing the model timestep to be rerun or 2) the reduction of the timestep following a large calving event to check if subsequent calving can occur. Henceforth, the latter is referred to as adaptive time 290 stepping.

## 6 Model parameter experiments

Predictions of calving front position depend not only upon the physical setup and tunable parameters such as crevasse penetration threshold, but also on numerical model parameters such as timestep and mesh density. Numerical model parameters can potentially affect the results through the calving algorithm or the characteristics of the chosen calving law. Experiments were 295 run to assess the impact on the calving algorithm by varying timestep magnitude and mesh density in runs using a simple rate-based law. These showed insignificant variations in predicted retreat rate ( $\pm 0.01$  m/d; Appendix B). Consequently, the focus of the experimentation is on the variability provided by the CD law, which provides a better representation of calving processes. Conceptually, the CD law predicts the stable ice-front position for any given position, and tends to converge on attractors at pinning points (Benn et al., 2023). In order to best show how non-physical factors within the numerical system can alter the 300 predicted calving, two synthetic domains were created. The first has a sloping bed with no bed irregularities producing a system lacking a strong attractor. The second has two basal rises (i.e. potential pinning points) near the initial terminus position which



**Figure 5.** Diagram showing the stages involved in a typical simulation using the calving algorithm and front advance routines in tandem.

act as an attractor within the system. The model setup is described in Appendix A. Note the location of the basal rises produces an asymmetric geometry. The control value for each parameter is described at the start of the relevant section. Mean retreat rates were calculated as the mean displacement of the terminus per timestep along the y-axis.

## 305 6.1 Timestep magnitude

The use of a position-based calving law, which calculates the ice-front location from an instantaneous state of the evolving glacier geometry and stress field, means its solution could be affected by the assigned timestep. Although rate-based laws will also be timestep dependent often through the evolution of the glacier free surfaces, they do not inherently aim to predict the attractor point in the retreat. As noted above, the duplicate experiments using a rate-based law were insensitive to timestep  
 310 changes. In contrast, the CD law aims to predict the position at which the glacier stabilises so will potentially be sensitive to the timestep when the glacier is in a transient state.

Four additional simulations with timesteps of 2d, 0.55d, 0.33d and 0.25d were run to complement the control simulation with the timestep of 1d (Fig. 6a, b). All the simulations ran for a total of 100d. The initial calving event on the first timestep remains similar irrespective of assigned timestep. For all runs, the glacier underwent retreat, but the rate of retreat increased as the timestep decreased when no pinning points were present. There is no clear convergence with each reduction in the timestep leading. In contrast, when pinning points are present, the magnitude of retreat no longer increased with a decrease in timestep beyond 0.33d. Further reduction of the timestep to 0.25d does not yield further retreat. Notably when pinning points are present, all the experiments showed very similar terminus positions and shapes (Fig. 6b).

320

The increased calving modelled when smaller timesteps were applied can be explained by the increased number of times that calving was predicted. As the timestep reduces, the frequency of computing the stress field on a unique geometry increases. This increased frequency of calving prediction increases the modelled calving because it heightens the probability of successful calving. For a glacier system with no strong attractor, calving is highly dependent on the timestep. When an attractor is added, in the form of pinning points, the predicted calving is much less sensitive to the model timestep. Diminished returns are seen and the retreat did not increase when reducing the timestep beyond 0.33d. Care should therefore be taken when using the CD law to model transient behaviour to choose the appropriate model timestep. The timestep is of less importance when an attractor is present as the general form of the retreat is consistent regardless of the timestep.

325

## 6.2 Adaptive timestep

The calving algorithm can add additional, shorter timesteps if large calving events occur to determine whether the calving-induced change in geometry will lead to further, immediate calving (Fig. 5). The control simulation could add up to three additional small timesteps of  $1 \times 10^{-10}$  yr if the calving threshold, set as  $1 \times 10^7$  cubic meters, was reached at the prior timestep. Four further simulations were done for each domain, the first with the adaptive time stepping deactivated and the other three experiments were set to enable up to one, five or ten additional timesteps (Fig. 6k, l). The calving threshold at which the adaptive time stepping was invoked and the additional small timestep size were not changed from the control.

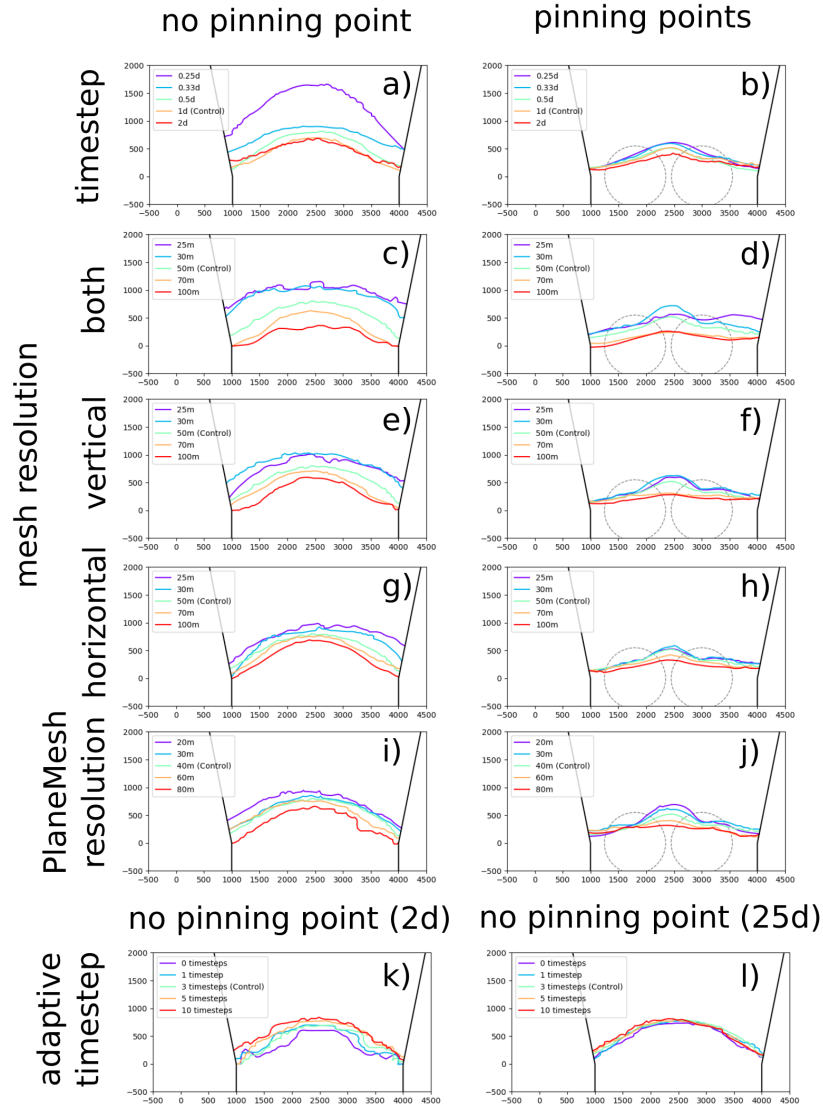
335

Without the adaptive timestep the terminus initially retreated at a slower rate but ultimately reached the same stable point as the control regardless of the domain (Fig. 6k, l). The positions for the more sensitive domain without pinning points are shown (Fig. 6k, l) but when pinning points are present the ultimate terminus also matches the control. Conversely, with more adaptive timesteps the terminus retreated at a quicker rate from the unstable starting geometry. Over time, the terminus positions slowly converged towards the same position where at 25d they are almost identical (Fig. 6l). From 25d until the end of the simulation at 100d the terminus positions do not diverge.

340

Adaptive time stepping intuitively makes sense based on our knowledge of secondary calving and so is important if a timeseries of terminus positions is wanted. If, instead, only the ultimate terminus position for a longer-term simulation is required, it is

345



**Figure 6.** Terminus positions at the end of the experiments altering numerical model parameters. a, c, e, g, i, k, l) were run on the domain with no pinning points and b, d, f, h, j) with basal pinning points. The numerical model parameter altered for each simulation shown were: a, b) timestep magnitude, c, d) horizontal and vertical mesh resolution, e, f) horizontal resolution, g, h) vertical resolution, i, j) PlaneMesh resolution and k, l) adaptive timestepping. The mean retreat rates are presents in Table 1.

unlikely that adaptive time stepping will change the result. Therefore, when using the adaptive time stepping functionality, it is important to consider the topic of investigation and timescales of interest for the simulation.

### 6.3 Glacier mesh density

As previously stated, the choice of mesh density is usually determined by the CFL criterion corresponding to the assigned  
350 timestep in glacier simulations. Increased horizontal and vertical resolution can refine the flow and stress solution, especially  
at areas of high shear or tensile stress. In the simulations shown in this paper, the smallest elements were closest to the glacier  
terminus where bending and extensional stresses are prominent and essential for determining calving locations. Increased mesh  
density comes at a computational cost, increasing exponentially with a component of between two to four as the resolution  
increases. This is because doubling the resolution quadruples the number of elements in the mesh. Usually mesh density is dic-  
355 tated by computational resources, but the small synthetic geometry used in this testing allows us to further investigate changes  
in the mesh resolution.

Experiments with varying mesh densities were conducted for both domains (Fig. 6e-h). The control had a vertical and hor-  
zontal mesh resolution of 50m at the terminus, and additional simulations were run with mesh resolutions of 25m, 30m, 70m  
360 and 100m. Vertical and horizontal mesh resolution were investigated independently and collectively producing 12 experiments  
for each domain (Fig. 6e-h).

Increased glacier mesh density lead to increased retreat across the terminus in the absence of a pinning point (Fig. 6e). When  
pinning points were present, although increased mesh density lead to increased retreat, the retreat is limited to where the glacier  
365 is pinned. A reduction in the glacier mesh density limits retreat on both domains. Independently changing the horizontal reso-  
lution had a limited impact on calving in either domain. Vertical mesh resolution has a far larger impact accounting for most  
the retreat seen when mesh resolution is increased across both planes (Fig. 6g, h).

The vertical resolution had a much larger impact on calving because this parameter alters the penetration of basal crevasses as  
370 – depending on the resolution – more or fewer elements are present near the waterline. Greater resolution allows the ice-cliff  
imbalances and bending stresses to be resolved in better detail, as the ice close to the waterline is depicted in more detail.  
In the simulations with finer resolution, multiple elements are present above the waterline but in the coarser mesh only one  
node is above this point. If the ice-cliff imbalance and bending stresses above the waterline are resolved in more detail, more  
extensional stress is captured above the waterline. Consequently, principal stress values are higher allowing surface crevasses  
375 to penetrate more of the ice column. Similar retreat to the experiments shown here could be replicated by locally increasing  
the vertical resolution 100m either side of the waterline. The importance of resolving the ice-cliff imbalances and bending  
stresses highlights two key details that need to be considered when applying the CD calving law. The first is that resolving  
the velocity using the full-Stokes flow is essential to account for the bending stresses. The second is that local vertical mesh  
refinement could potentially be valuable for reducing computational costs while successfully simulating calving dynamics.



380 This would follow in a similar vein to the understanding of the horizontal anisotropic remeshing being important in ice-sheet scale modelling to resolve the flow dynamics at ice streams (e.g., Gillet-Chaulet et al., 2012).

#### 6.4 Plane mesh density

When calving is predicted the crevasse field is mapped onto a 2D mesh, known as the PlaneMesh, the resolution of which is independent of the 3D glacier mesh. Since there are no partial differential equations being solved on it, the resolution of the  
385 PlaneMesh does not greatly change the computational cost of the algorithm. Its use, however, is not fully parallel so does not scale as well as the 3D glacier mesh.

For each domain, four simulations with PlaneMesh resolutions of 20m, 30m, 60m and 80m were run for comparison against the control that had a resolution of 40m (Fig. 6i, j). Increased PlaneMesh density led to more calving, particularly when no pinning  
390 points were present. Here, increased PlaneMesh resolution showed a linear increase in calving volume with no clear reducing pattern. When pinning points were present the retreat was limited to the location the basal rises even as the resolution increased.

Increasing the PlaneMesh resolution increases the number of points at which the vertical ice column is assessed for the computation of crevasse penetration (Fig. 6 i, j). By reducing the space between these points, the location of the crevasse field  
395 inducing calving can be more accurately determined. This will often shift the crevasses up glacier slightly when the modelled crevasse location is between PlaneMesh elements. The maximum upstream refinement in crevasse location is determined by the element size.

The low computational cost of the PlaneMesh routine means that a finer mesh resolution than the glacier mesh should always be aimed for. It is difficult to determine how fine a resolution to apply, especially when there is a relationship with the  
400 3D glacier mesh resolution as well. The lower variation in retreat seen when the horizontal glacier mesh resolution is altered suggests a interdependence between the two meshes that warrants further investigation. When using the CD law, the resolution of crevasse mapping does affect calving and so the sensitivity to a given set-up should be considered in future glaciological applications of the algorithm. Glaciers in transient states will be more heavily influenced by mesh resolution. The PlaneMesh  
405 resolution should remain consistent between experiments as well.

#### 6.5 Summary of the influence of numerical model parameters on calving

In summary, this set of experiments shows that changes in timestep length, mesh density, and plane mesh can affect predicted calving using the CD law. Differences in predicted calving are independent of the new algorithm that shows very little variability when using a rate based law (Appendix B). Importantly, however, changes in these parameters have a much smaller  
410 impact on the predicted ice-front position when a pinning point is present. That is, the calving predicted by the CD law centres on attractors for all chosen values of parameters when the system contains pinning points. In contrast, when such an attractor is absent, the model runs simulate different rates of retreat depending on parameter choices. This suggests that the choice of

**Table 1.** Summary of the mean retreat rates for 100d experiments testing the numerical model parameters that can alter calving.

		Mean retreat rate (m/d)	
Numerical model parameter	Experiment	No pinning points	Pinning points
Control		4.36	3.13
Timestep	0.25d	12.16	3.67
	0.33d	7.27	3.57
	0.5d	5.50	2.99
	2d	4.52	2.51
Mesh resolution (both)	25m	9.56	4.46
	30m	9.12	4.39
	70m	2.03	1.57
	100m	0.73	1.32
(vertical)	25m	6.83	3.48
	30m	7.72	3.86
	70m	2.89	2.48
	100m	1.28	2.10
(horizontal)	25m	6.70	3.40
	30m	5.66	4.35
	70m	3.94	2.63
	100m	2.51	2.21
PlaneMesh	20m	6.29	3.89
	30m	5.01	3.90
	60m	4.00	2.75
	80m	2.33	2.40
Adaptive time	0 timesteps	4.18	3.11
	1 timesteps	4.37	2.92
	5 timesteps	4.45	2.88
	10 timesteps	4.55	3.32

temporal and spatial model resolution is of greatest importance where transient glacier behaviour is of interest when using the CD law.

## 415 7 Summary of a model capabilities and potential

Putting together the new calving algorithm along with the upgrades in the calving projection gives us a new model with unparalleled capabilities of simulating calving over a 3D continuum. New features of the model include:

1. Unlimited advance or retreat can now be simulated in 3D.
2. Unrestricted 3D calving geometries can be utilised by the model.
- 420 3. Any calving law can be implemented
4. Features or variables can be advected as part of the mesh.
5. Any vertically and horizontally varying melt field can be applied to the glacier front.

### 7.1 Unlimited retreat and advance

The ability to model unrestricted retreat and advance in 3D is a major step forward for simulating the dynamics of calving  
425 glaciers. This allows simulations of longer term glacier change where the terminus retreats many kilometers, which was previously impossible (Todd et al., 2018). Before these developments, 3D glacier simulations had been limited to stable glaciers that do not undergo large seasonal variability (Todd et al., 2019; Cook et al., 2023). The ability to simulate unlimited retreat and advance presents an opportunity to model any glacier in the world.

### 7.2 Unrestricted calving

430 The ability to calve unrestricted geometries of icebergs both in the horizontal and vertical planes should be treated as a distinct feature of this calving algorithm. This means any potential configuration of calving or front geometry is possible. Again, this ability opens up the possibility to model any complex scenario or situation seen in the real world. For example, glaciers with complex front geometries such as Bowdoin Glacier (Van Dongen et al., 2019) or large fan-shaped ice tongues which are non-projectable can now be modelled. The new calving algorithm also offers the possibility to create unrestricted synthetic calving  
435 geometries to explore how the glacier dynamics respond to forced calving events.

### 7.3 Flexible implementation of calving laws

The calving algorithm is not restricted to the CD law implementation outlined above. Any calving law could be implemented through the production of a level set variable or signed distance variable that is given in the calving algorithm. Despite the relative ease by which a new calving law could be implemented, only the CD calving law has been used up to now. This is  
440 because it is currently the only only calving law based on physical processes (Benn et al., 2017). Other popular calving laws are based on calving rates as opposed to calving position and they could be used in conjunction with this calving algorithm. The calving algorithm provides an easily accessible framework for which various calving laws could be compared in 3D. More likely, alterations or improvements will need to be added to the CD law as coupled modelling of tidewater glaciers

advances (Cook et al., 2023). Possible ways in which the CD law could be developed include the incorporation of ice history  
445 via advection of damage, and implementing more sophisticated methods of calculating crevasse depths. However, this new  
algorithm provides the best framework to approach these problems, as 3D modelling is no longer restricted by technical  
hurdles. As such we can now focus on improving the calving laws along with assessing which missing processes are important  
in calving prediction.

#### 7.4 Feature advection

450 The advances in remeshing techniques allow complete anisotropic remeshing of a particular glacier part or complete glacier.  
This allows mesh quality to be maintained even if nodes are moved in a Lagrangian manner. Some issues remain related to  
element degeneracy at the lateral margins, but this does not apply to ice shelves such as those extending from Thwaites Glacier  
which are laterally unconstricted (Scambos et al., 2017). There is very exciting potential for use of the remeshing techniques  
to advect variables such as damage downstream in order to better predict calving, especially on ungrounded ice sheets (Cook  
455 et al., 2023).

#### 7.5 Use for melt simulations

Beyond the calving component of the new algorithm, the ability to model the effects of submarine melt on glacier dynamics  
in 3D is very novel. The limited availability of full-Stokes glacier models implies that the effects of 3D melt fields on glacier  
dynamics are rarely researched. The ability of the algorithm to have a non-projectable front means any melt field could be  
460 applied for any length of time without model breakdown. On its own, the calving algorithm is limited in applying a set melt  
field to the terminus. Future work should focus on coupling the glacier model, with the calving algorithm, with ocean/plume  
models (Cook et al., 2023).

The most advanced method of calculating frontal melt in Elmer is the coupled hydrology model developed by Cook et al.  
465 (2020). This model couples ice flow with the Glacier Drainage System (GlaDS) module in Elmer/Ice, uses predicted subglacial  
meltwater discharge to drive a 1D plume model and determine patterns of frontal melting, and the CD law to predict consequent  
calving. This work employed the Todd et al. (2018) calving algorithm, and significant further development may be required to  
reproduce this effort with the new calving algorithm. Future coupling work in Elmer should focus not just on hydrology but  
also fjord circulation (Cook et al., 2023). The lack of coupling of fjord models with glacier models means there are often large  
470 uncertainties when applying melt fields to glacier models. Melt profiles are often derived from buoyant plume theory (Slater  
et al., 2017), but the lack of 3D fjord modelling neglects horizontal flow across the front of the glacier. Coupling should aim  
to be with a high-resolution fjord model such as MITgcm (Cook et al., 2023). Although computationally expensive, it seems  
futile to solve the glacier dynamics in detail but neglect the same detail with the fjord model. However, many issues such as  
congruent time-step sizing would need to be resolved.

## 475 7.6 Future parallelisation

A parallel calving algorithm is currently in the development stage. Conceptually it follows the serial calving routine but undertakes all computationally expensive routines in parallel. In some ways this vastly simplifies the calving algorithm as calving is always implemented in parallel rather than switching between serial and parallel routines. This cuts out the need to gather and redistribute the mesh along with reducing the complexity of the additional new functionality. This would allow increased  
480 scalability of the algorithm allowing it to be used in large scale simulations as core models on Elmer/Ice have been shown to scale well for high performance computing (Gagliardini et al., 2013). However, large-scale testing has shown the need for remeshing to occur in serial is not an insurmountable problem at present, as solving the Stokes equations is still the major computational requirement for any simulation.

## 8 Conclusions

485 The new calving algorithm has been shown to be capable of simulating unrestricted calving and terminus advance. This marks a major step forward in our ability to model and therefore understand calving dynamics. Importantly, the new algorithm and its use as part of Elmer/Ice remains computationally light compared to DEMs such as HiDEM and fills the gap between models based on first principles and the widely used SSA-style ice-sheet models.

490 An assessment of the numerical model parameters and their potential to alter calving predicted by the CD law revealed numerical decisions can have a large impact on calving for systems lacking a strong attractor showing the sensitivity of the CD law when the glacier is in a transient state. For systems with strong attractors, such as those with pinning points, the influence of model parameter choice is limited. As such, modellers should be aware of the sensitivity of the system of interest when choosing numerical model parameters. Importantly, calving predicted by the CD law is very coherent when an attractor is  
495 present regardless of modelling decisions.

*Code and data availability.* The data associated with this study is made available as supplementary information. The new algorithm including current and future releases is available on the official Elmer/Ice GitHub repository <https://github.com/ElmerCSC/elmerfem>. The exact code used in this study is available on Zenodo at <https://zenodo.org/records/10182705> (Wheel 2023a). The user guides can be accessed from GitHub at [https://github.com/ElmerCSC/elmerfem/tree/calving\\_meshadapt/elmerice/Solvers/Documentation](https://github.com/ElmerCSC/elmerfem/tree/calving_meshadapt/elmerice/Solvers/Documentation). The three documents associated with the new calving algorithm are `Calving3D_lset.md`, `CalvingRemeshMMG.md` and `CalvingGlacierAdvance3D.md`.  
500

*Video supplement.* Videos of the control experiment for each domain are provided as Supplementary Video 1 and 2.

## Appendix A: Synthetic glacier setup

The synthetic glacier domain extends 5km upstream from the calving front and has a terminus width of 3km (Fig. A1). It flows through a narrowing fjord that has an upstream boundary width of 5km. The fjord geometry projected beyond the initial domain has parallel sidewalls. An initial 3D tetrahedral mesh was created using the meshing software Gmsh. The mesh consisted of 5 layers of 100m resolution that were extruded between surface and bed maps. The horizontal resolution at the terminus was 100m before increasing to 500m at the inflow boundary (Fig. A1 a-c). Two bed geometries were created to produce two domains, one with a constant downglacier slope and the other with two additional pinning points near the terminus (Fig. A1 d, e). The first domain had the formulation

$$B_h(x, y) = InteriorDepth + WidthDepth + TerminusRise, \quad (A1)$$

while the second had the addition to pinning points to give

$$B_h(x, y) = InteriorDepth + WidthDepth + Bump1 + Bump2 + TerminusRise, \quad (A2)$$

where

$$Depth(y) = B_0 + ym_b, \quad B_0 = -550, \quad (A3)$$

$$WidthDepth(x, y) = \frac{|x - 2500|}{(3000 + \frac{y}{5})} \times 400, \quad (A4)$$

$$Bump1(x, y) = \exp(H_b \times \frac{(x - 1800)^2 + y^2}{r_b})^2, \quad H_b = 100, \quad r_b = 550, \quad (A5)$$

$$Bump2(x, y) = \exp(H_b \times \frac{(x - 3000)^2 + y^2}{r_b})^2, \quad H_b = 100, \quad r_b = 550, \quad (A6)$$

$$TerminusRise(y) = \exp(H_t \times \frac{-y}{l_t})^2, \quad H_t = 100, \quad l_t = 1000, \quad (A7)$$

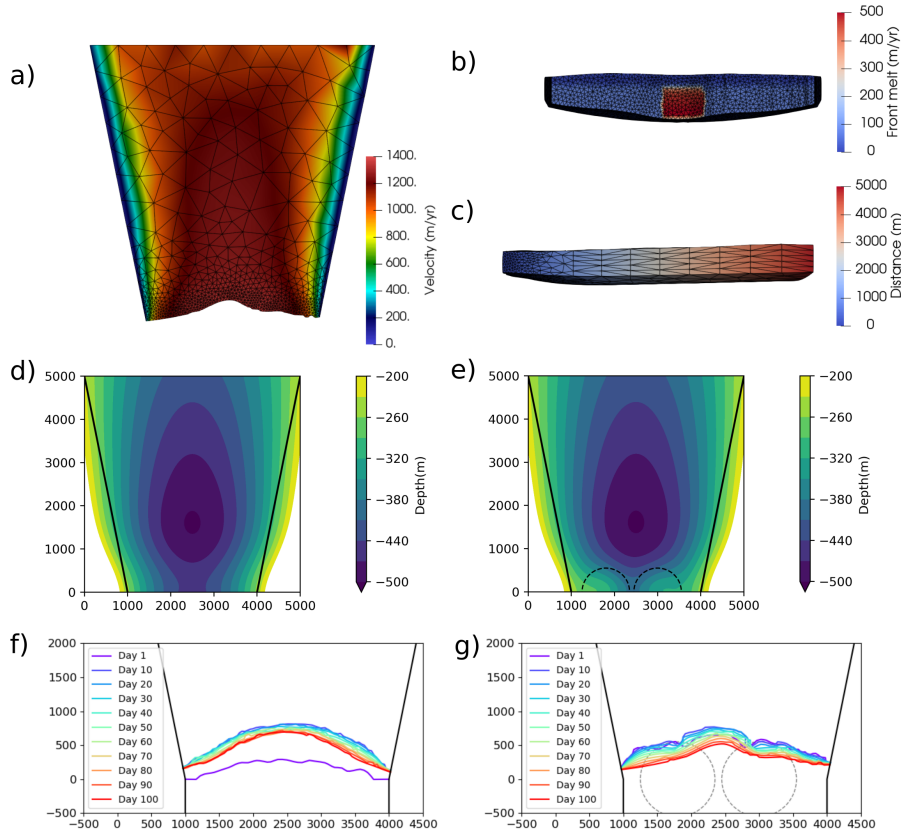
and where  $B_h$  is the bedrock height in meters,  $m_b$  is 1/40 the gradient of the bed,  $x$  is the x coordinate,  $y$  is the y coordinate and  $B_0$  is the bed height at  $y = 0$ . For the bumps,  $H_b$  is the height of the bump and  $r_b$  is the radius of bump. The surface geometry was identical for both domains and is given by

$$S_h(y) = S_0 + ym_s, \quad S_0 = 50, \quad (A8)$$

where  $S_h$  is the surface height and  $m_s$  is 1/40 the gradient of the surface.

### A1 Boundary conditions

The synthetic glacier domain had six boundary conditions that consist of the calving front of the glacier ( $\Gamma_{front}$ ), both lateral margins ( $\Gamma_{left}$ ,  $\Gamma_{right}$ ), the interior inflow ( $\Gamma_{inflow}$ ), the base ( $\Gamma_{base}$ ) and the top surface ( $\Gamma_{surf}$ ). The boundary conditions applied to each is discussed below.



**Figure A1.** The synthetic geometry setup. a) Topdown view of the mesh at the end of the 100d simulation with pinning points. Mesh element size is 50m at the terminus increasing to 500m further upstream. The velocity field is displayed. b) Front view of the from the same simulation showing the horizontal and vertical element size of 50m at the terminus. The melt plume can be seen in the centre of the terminus. c) Side view of the same simulation showing the 50m mesh density at the terminus and 500m at the inflow boundary. d) The bed elevation with no pinning points. e) The bed elevation with pinning points near the terminus. The terminus evolution of during 100d simulation with f) no pinning points and g) pinning points. The full control simulations for each domain can be seen in Supplementary Videos 3 and 4.

Since the lateral boundary represents the fjord walls, a no-penetration condition is implemented at the sidewall margins of the model domain. A simple linear friction law, which has a constant slip coefficient ( $\beta$ ) of  $1e-2 \text{ MPa m}^{-1}$  was applied as a Neumann boundary condition. Therefore, the lateral boundary conditions are

$$u_{\perp} = 0, \quad \text{on } \Gamma_{left}, \Gamma_{right}, \quad (\text{A9})$$

$$\sigma_{||} = -u_{||}\beta, \quad \text{on } \Gamma_{left}, \Gamma_{right}, \quad (\text{A10})$$

535 where  $u$  is the velocity component,  $\sigma$  the stress component and the perpendicular and tangential components are shown by  $\perp$  and  $\parallel$  respectively. A non-linear Weertman friction law is applied to the base with a slip coefficient of  $1e-4 \text{ MPa m}^{-1/3} \text{ yr}^{1/3}$  and an exponent of 3. The base boundary condition is complicated by the grounding line dynamics of the glacier. The grounding line of the glacier is solved as a contact problem following Favier et al. (2012). If the glacier is grounded the boundary condition is similar to the lateral boundary with a non-penetration condition applied:

$$540 \quad u_{\perp} = 0, \quad \text{on } \Gamma_{base}, \quad (\text{A11})$$

$$\sigma_{\parallel} = C u_b^{1/m}, \quad \text{on } \Gamma_{base}, \quad (\text{A12})$$

where  $C$  is the Weertman slip coefficient,  $u_b$  is the basal velocity and  $m$  is the Weertman exponent. If the glacier is ungrounded, no friction is applied and the glacier is free to move vertically:

$$\sigma_{\perp} = \min(-\rho_w g h, 0), \quad \text{on } \Gamma_{base}, \quad (\text{A13})$$

$$545 \quad \sigma_{\parallel} = 0, \quad \text{on } \Gamma_{base}, \quad (\text{A14})$$

where  $\rho_w$  is the density of the water,  $g$  is the gravitational acceleration and  $h$  is the depth below the water level. Similar to ungrounded ice, no friction is applied to the glacier calving face. Since it is in contact with the fjord water body a normal stress is applied below the water level.

$$\sigma_{\perp} = \min(-\rho_w g h, 0), \quad \text{on } \Gamma_{term}, \quad (\text{A15})$$

$$550 \quad \sigma_{\parallel} = 0, \quad \text{on } \Gamma_{term}. \quad (\text{A16})$$

The inflow boundary has a fixed velocity of 1000m/yr and this is assumed to be constant throughout the vertical ice column. Therefore, the inflow boundary condition is simply

$$|\mathbf{u}| = u_{in}, \quad \text{on } \Gamma_{inflow}, \quad (\text{A17})$$

where  $\mathbf{u}$  is the velocity vector and  $u_{in}$  is 1000 m/yr. The surface boundary condition is stress-free and surface mass balance is not considered. A submarine melt condition is added to the glacier front boundary or calving front ( $\Gamma_{term}$ ) where a central plume is present along with background melt (Fig. A1 b). The plume is taken from the summer plumes modelled by Kajanto et al. (2023) at Illulissat Fjord. The plume profile was then normalised given the much lower maximum flow velocity at the synthetic geometry of 1500m/yr compared to the much larger speed seen at Jakobshavn Isbrae ( $> 10,000$ m/yr). Maximal plume melt was set to 500m/yr and the profile adjusted accordingly. No melt was applied to the floating ice on the base.

## 560 **A2 Ice properties**

A constant temperature of  $-20^{\circ}\text{C}$  is set throughout the domain and the associated ice properties are based on Glen's flow, which is then used to solve the full-Stokes equations (Cuffey and Paterson, 2010). The feedback of the temperature dependency of Glen's flow law is calculated using the Arrhenius equation and the rate factors are detailed in (Cuffey and Paterson, 2010).



### A3 Model solvers and parameters

565 For the transient simulation, the model is run forward for 100 days at timesteps of 1d, giving a total of 100 timesteps. The full-Stokes flow is solved and from this solution the Cauchy stress tensor across the domain is computed. Although there are no surface mass balance conditions, the surface and base free surfaces are solved as a kinematic boundary condition so that the glacier can evolve in response to the flow solution. Similarly, the new front advance routines outlined are used to predict the advance of the glacier down the fjord. As a consequence, the mesh is deformed twice, first vertically and then longitudinally.  
570 The longitudinal mesh deformation is limited to 1500m from the calving front of the glacier.

The new calving algorithm as outlined in the main section is applied and the front 1500m of the glacier is remeshed at each timestep. The anisotropic remeshing metric had a minimum horizontal resolution of 50m, which increases to 500m further inland. It has a constant vertical resolution of 50m. The adaptive time stepping present in the calving algorithm was active,  
575 with a maximum number of added timesteps set to three. Timesteps were only added if large calving events occurred to capture any consecutive calving from changes in the domain. The calved iceberg threshold was set to  $1e7$  cubic meters for the adaptive time stepping to be activated. The 2D PlaneMesh that crevasses were mapped onto had a grid size of 40m.

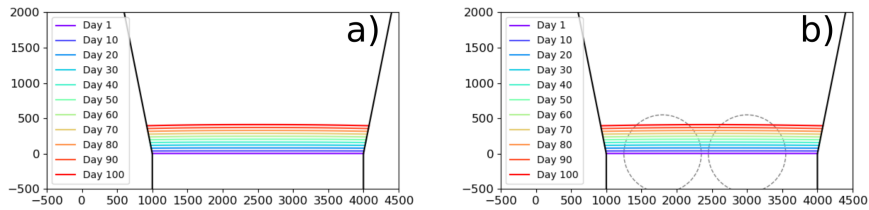
The crevasse depth (CD) law is modified to make the setup more sensitive to calving and highlight potential numerical in-  
580 fluences on calving. This is achieved by reducing the *crevasse penetration threshold* to 92.5%. Here, full thickness calving occurs when either: 1) a surface crevasse penetrates 92.5% of the ice column between the surface and water line or 2) surface and basal crevasses extend 92.5% of the entire water column. Although it is known that CD law can underestimate calving (e.g., Choi et al., 2018; Todd et al., 2019; Cook et al., 2023) this should not be concluded in this instance. Instead, the alteration of the calving law should be considered as an increase to the sensitivity of a synthetic glacier so that the effects of varying  
585 parameters on calving dynamics can be clearly identified in the following simulation tests. No conclusion on the accuracy of the CD law can be made for a synthetic scenario.

### Appendix B: Rate-based calving law experiments

In order to distinguish between the any variability arising from the CD law and that of the new calving algorithm further experiments were run using a simple rate-based law. The new calving law took the form

$$590 \quad C = V_s - \hat{R}, \tag{B1}$$

where  $C$  is the predicted calving magnitude,  $V_s$  the surface velocity at the terminus and  $\hat{R}$  the prescribed retreat rate. The prescribed retreat rate is applied normal to the terminus and was set to 1500m/yr for all experiments. The setup excluding the calving law was identical to that described in Appendix A and the control variables kept consistent. Complimentary experiments to those using the CD law in the main text were run. This produced the mesh density experiments with minimum  
595 element sizes of 25m, 30m, 70m and 100m. Experiments changing the timestep were the same as the main text using timesteps



**Figure B1.** Terminus positions through 100d control experiments using the rate-based law on a) domain without pinning points and b) the domain with pinning points. The outline of the pinning points is shown by the dashed circles. Only the control experiments are shown as the results were too consistent to overlay on each other.

of 0.25d, 0.33d, 0.5d and 2d. Finally, the use of the adaptive time stepping in the algorithm was set to 0d, 1d, 5d and 10d.

Predicted retreat and calving varied little between the rate-based experiments with a mean of 4.06m/d. Maximum and minimum values were only 0.01m/d greater or less than the mean (Table B1). The rate of retreat throughout the experiments was consistent and did not vary between the two different setup domains where the presence of pinning points had no influence on predicted calving (Fig. B1). Simulated retreat was slightly lower than the prescribed  $R$  value (4.11m/d) but was consistent between experiments.

Very little variability in the simulated calving can be attributed to the new calving algorithm and instead variability in the main test can be attributed to the CD law. The trivial variation in the rate-based law experiments is due to the Hausdorff distance prescribed in the remeshing where slight errors are introduced into the surface remapping. It is likely the remeshing is the cause of insignificant difference between the mean retreat rate and the prescribed retreat rate (0.05m/d). However, given the consistency of the implemented retreat, this issue is easily overcome by adjustment of the input  $R$  value.

*Author contributions.* IW wrote the algorithm with technical support from JT and TZ. IW designed and analysed the experiments with guidance from DB and AC. IW wrote the manuscript with contributions from DB and AC. All authors contributed and approved the final manuscript.

*Competing interests.* The authors declare that they have no conflict of interest

**Table B1.** Summary of the mean retreat rates for 100d experiments testing the numerical model parameters using a rate-based calving law. The numerical model parameters tested match those in the main text.

		Mean retreat rate (m/d)		
Numerical model parameter	Experiment	No pinning points	Pinning points	
Control		4.07	4.06	
Timestep	0.25d	4.06	4.06	
	0.33d	4.06	4.06	
	0.5d	4.07	4.06	
	2d	4.07	4.06	
Mesh resolution (both)	25m	4.07	4.06	
	30m	4.06	4.06	
	70m	4.07	4.06	
	100m	4.06	4.06	
	(vertical)	25m	4.05	4.05
		30m	4.06	4.05
		70m	4.06	4.06
		100m	4.05	4.06
(horizontal)	25m	4.07	4.06	
	30m	4.06	4.06	
	70m	4.06	4.05	
	100m	4.06	4.05	
Adaptive time	0 timesteps	4.07	4.07	
	1 timesteps	4.06	4.06	
	5 timesteps	4.06	4.06	
	10 timesteps	4.06	4.06	

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