

1 Global source apportionment of aerosols into major emission regions
2 and sectors over 1850–2017

3
4
5 Yang Yang^{1*}, Shaoxuan Mou¹, Hailong Wang², Pinya Wang¹, Baojie Li¹, Hong Liao¹

6
7
8
9 ¹Joint International Research Laboratory of Climate and Environment Change (ILCEC),
10 Jiangsu Key Laboratory of Atmospheric Environment Monitoring and Pollution Control,
11 Jiangsu Collaborative Innovation Center of Atmospheric Environment and Equipment
12 Technology, School of Environmental Science and Engineering, Nanjing University of
13 Information Science and Technology, Nanjing, Jiangsu, China

14 ²Atmospheric, Climate, and Earth Sciences Division, Pacific Northwest National Laboratory,
15 Richland, Washington, USA

16
17
18
19
20
21 *Correspondence to yang.yang@nuist.edu.cn

22 **Abstract**

23 Anthropogenic emissions of aerosols and precursor gases have been changing significantly in
24 the past few decades across the world. In this study, an explicit aerosol source tagging system
25 (EAST) is merged into the Energy Exascale Earth System Model version 1 (E3SMv1) to
26 quantify the variations in anthropogenic aerosol concentrations, source contributions, and
27 their subsequent radiative impact in four major emission regions on the globe during 1850–
28 1980, 1980–2010 and 2010–2017. In North America and Europe, changes in anthropogenic
29 PM_{2.5} were mainly caused by changes in emissions from local energy and industrial sectors.
30 The local industrial sector caused the most increase in PM_{2.5} in East Asia during 1980–2010
31 and decrease during 2010–2017. In South Asia, the increase in energy-related emissions
32 dominated the rise of PM_{2.5} levels during 1980–2017. During 1850–1980, the increases in
33 emissions from North America contributed to the increase in European PM_{2.5} burden by 1.7
34 mg m⁻² and the sources from the Europe were also responsible for the PM_{2.5} burden increase
35 in East Asia and South Asia by about 1.0 mg m⁻². During 1980–2010, East Asia contributed
36 to an increase of 0.4–0.6 mg m⁻² in PM_{2.5} burden in North America and Europe, while South
37 Asian contributed about 0.3 mg m⁻². During 2010–2017, the contributions from East Asia to
38 the PM_{2.5} burdens in the North America, Europe and South Asia declined by 0.3–0.6 mg m⁻²
39 due to Clean Air actions in China, while the contributions from South Asia still increased due
40 to the continuous increase in emissions in South Asia. The historical changes in aerosols had
41 an impact on effective radiative forcing through aerosol-radiation interactions (ERF_{ari}).
42 During 1980–2010, a decline in North American aerosols resulted in a positive ERF_{ari} change
43 (warming effect) in Europe and a decline of aerosols in Europe caused a warming effect in

44 Russia and northern China. The changes in ERF_{ari} from the increase and decrease of aerosols
45 in China during 1980–2010 and 2010–2017, respectively, are comparable in magnitude. The
46 continuous aerosol increases in South Asia from 1980 to 2017 resulted in negative ERF_{ari}
47 (cooling) changes in South Asia, Southeast Asia, and southern China.

48 **1. Introduction**

49 Atmospheric aerosols, also known as particulate matter (PM), have significant influences
50 on air quality and human health (Anenberg et al., 2010; Finlayson-Pitts and Pitts, 1997; Li et
51 al., 2017; Pöschl, 2005). Aerosols also affect the energy budget of the earth system by
52 scattering and/or absorbing solar radiation, thus directly affecting the climate (Gao et al., 2022;
53 Yang et al., 2020a, 2023; Wang et al., 2023). Meanwhile, they may act as cloud condensation
54 nuclei and/or ice nuclei, changing cloud characteristics and atmospheric water cycle, which
55 indirectly affect the climate (Liao et al., 2015; Lohmann and Feichter, 2005; Rosenfeld et al.,
56 2008; Yang et al., 2022a). Due to the absorption of solar radiation, aerosol-induced heating can
57 strengthen temperature inversion and increase the atmospheric stability, which inhibits the
58 vertical mixing and transport of aerosols and leads to a further increase in near-surface aerosol
59 concentrations (Chen et al., 2021; Lou et al., 2019). Therefore, knowing the sources of aerosols
60 and their variations have become a vital direction in the field of environmental and atmospheric
61 sciences.

62 Human activities have a great influence on global aerosol distributions and compositions.
63 For example, many countries have taken various air quality control measures at different stages
64 of their economic development, causing distinct historical temporal changes of aerosol
65 emissions across the world. Since the start of industrialization, anthropogenic emissions of
66 aerosols and precursor gases have substantially increased, which significantly affected the
67 atmospheric environment and the Earth's energy balance (Carslaw et al., 2013). European and
68 North American countries became major contributors of pollutant emissions. Since the 1980s,
69 coal emissions have declined steadily in Europe and North America owing to the legislation

70 and effective environmental policies to reduce local anthropogenic emissions of aerosol and
71 precursor gases (Smith et al., 2011). In contrast to North America and Europe, the coal
72 consumption in China and India has experienced a substantial increase and anthropogenic
73 emissions from these regions continued to rise (Hoesly et al., 2018; Lim et al., 2020). Zheng et
74 al. (2018) also reported that due to active clean air policies and the emission control of power
75 plants and industry, anthropogenic emissions of PM_{2.5} (particulate matter less than 2.5 µm in
76 diameter) from China have significantly decreased by 33% during 2013–2017. However,
77 countries in South Asia still rely on coal and petroleum and thus aerosol emissions from South
78 Asia have kept increasing in recent years (Li et al., 2017).

79 Regional aerosol pollution can be induced by both local emissions and long-range
80 transport of pollutants across regions, countries or even continents, which impose a far-
81 reaching impact on air quality and human health (Akimoto, 2003; Anenberg et al., 2014; Jaffe
82 et al., 1999; Lin et al., 2014; Liu et al., 2009; Zhang et al., 2017). Studies reported that the air
83 quality in Europe is largely impacted by the long-range aerosol transport from North America
84 (Stohl and Trickl, 1999; Yang et al., 2018a, 2020b). Asian anthropogenic emissions in spring
85 also have a significant effect on the aerosol concentrations in North American (Jaffe et al.,
86 1999). Moreover, studies found that air pollution from Africa and Europe moved eastward and
87 merged with Asian pollution, affecting air quality in western North America (Liu et al., 2005;
88 Chin et al., 2007). Yang et al. (2017) also found that remote sources contributed the most to
89 the regions with low emissions through long-range transport, which further impacted the local
90 climate. Therefore, relying on the domestic emission control alone may be insufficient to
91 prevent air pollution due to the long-distance transport of air pollutants (Liu et al., 2009). A

92 study revealed that approximately 12% of global premature deaths caused by PM_{2.5} were
93 related to non-local air pollutants (Zhang et al., 2017). About 16% of premature deaths in India
94 caused by PM_{2.5} were attributed to aerosol transport from external source regions (David et al.,
95 2019). Within each emission source region, aerosols also come from different emission sectors.
96 Many scientific control measures and policies are implemented based on the source attribution
97 of air pollutants from individual sectors. Hence, it is of great significance to quantify source
98 contributions of long-range transport of aerosols from major emission regions of the world as
99 well as aerosols from major emission sectors.

100 Anthropogenic emissions of aerosols and precursor gases have changed significantly in
101 different source regions over the past century. Few studies focus on the source attributions of
102 aerosols across the globe over such a long period of time. In this study, we focus on the changes
103 in aerosols and emission source region and sector contributions in major source regions (i.e.,
104 North America, Europe, East Asia, South Asia) during the three important periods of emission
105 changes since industrialization (1850–1980, 1980–2010 and 2010–2017) based on the
106 Energy Exascale Earth System Model version 1 (E3SMv1), equipped with an explicit aerosol
107 source tagging system (E3SMv1-EAST).

108 **2. Methods**

109 **2.1. Model description and experimental design**

110 To study variations in historical anthropogenic aerosols in the major source regions, the
111 E3SMv1 developed by US Department of Energy (DOE) (Golaz et al., 2019) is used in this
112 study. The E3SMv1 model is updated on the basis of Community Atmosphere Model version
113 5 (CAM5) in order to explore several key emerging issues in the field of environment and

114 climate, and is a branch of the widely-used Community Earth System Model (CESM) (Rasch
115 et al., 2019). E3SMv1 consists of atmosphere, land surface, ocean, sea ice, and river model
116 components. It features numerous upgrades to aerosol, turbulence, chemical, and cloud-related
117 processes, offering multiple spatial resolution options. The model can run simulations for
118 decades or more at higher resolution to help understand past, present, and future changes in
119 Earth's behavior, and to explore how the atmosphere interacts with other components of the
120 Earth system. Aerosol microphysics and interactions with stratiform clouds are treated with the
121 four-mode Modal Aerosols Module (MAM4) (Liu et al., 2016), which predicts the mass and
122 number concentrations of sulfate, black carbon (BC), primary organic matter (POM),
123 secondary organic aerosol (SOA), marine organic aerosol, mineral dust and sea salt (Wang et
124 al., 2020). EAMv1 applies the “Morrison and Gettelman version 2” (MG2) two-moment bulk
125 microphysics parameterization for stratiform clouds (Gettelman and Morrison, 2015). It allows
126 aerosol-cloud interactions in all stratiform and shallow convective clouds but neglects in deep
127 convective clouds (Rasch et al., 2019). Liquid cloud drop activation is based on Abdul-Razzak
128 and Ghan (2000) and a classical nucleation theory-based ice nucleation parameterization for
129 the heterogeneous ice formation in mixed-phase clouds follows Y. Wang et al. (2014).
130 Hygroscopicity are specified for soluble aerosols to calculate the particle size based on relative
131 humidity. The aerosols are assumed to mix internally in the same aerosol mode and externally
132 between modes when calculating the aerosol optical properties (Ghan and Zaveri, 2007). The
133 model has been applied to investigate the variations in anthropogenic and natural aerosols
134 related to the air-sea interactions (Yang et al., 2022b; Zeng et al., 2021). Compared to the
135 regional model, the E3SMv1 with an aerosol tagging tool introduced in this study is more

136 suitable for the simulation of transboundary and intercontinental transport of aerosols across
137 the globe. In this study, the model is configured at its standard horizontal spatial resolution of
138 approximately 1° with 72 vertical layers.

139 Global emissions of aerosols and precursor gases used in the simulations are obtained
140 from the CMIP6 (the Coupled Model Intercomparison Project Phase 6) datasets (Hoesly et al.,
141 2018; van Marle et al., 2017). However, the anthropogenic emissions in China are replaced
142 with MEIC (Multi-resolution Emission Inventory for China) inventory, which fully considers
143 the implementation of clean air actions over China since the 2010s (Gao et al., 2022, 2023; Li
144 et al., 2021; Zheng et al., 2018). Following previous studies (Ren et al., 2021; Yang et al.,
145 2018a), the near-surface concentrations of $PM_{2.5}$ here are estimated as the sum of sulfate, BC,
146 POM and SOA concentrations. Effective radiative forcing (ERF) refers to the change of the
147 net radiative flux at the top of the atmosphere (TOA) after the external forcing is applied. In
148 this study, ERF due to aerosol-radiation interactions (ERF_{ari}) for the individual tagged source
149 regions is derived as the difference in TOA net radiative fluxes from a pair of diagnostic
150 radiation calculations with and without the particular tagged aerosols from the source regions
151 for the all-sky condition following Ghan (2013).

152 This study focuses on the variations in source region and sector contributions in four major
153 emission regions of the world (North America, Europe, East Asia and South Asia) during the
154 three key historical periods of emission changes (1850–1980, 1980–2010 and 2010–2017).
155 Four simulations with monthly anthropogenic emissions of aerosols and precursors fixed at the
156 1850, 1980, 2010 and 2017 levels, respectively, are conducted. All simulations are performed
157 for one year following 6-month model spin-up. Greenhouse gas concentrations, solar insolation,

158 sea surface temperature and sea ice extent are prescribed at the 2000 levels. The meteorological
159 fields including 3-dimensional temperature, specific humidity, and winds are nudged toward
160 the MERRA-2 (Modern-Era Retrospective Analysis for Research and Applications, version 2)
161 reanalysis (Gelaro et al., 2017) in year 2017 at a 6-hourly relaxation timescale.

162 **2.2. Explicit aerosol source tagging system**

163 Source apportionment aims to quantify the contributions to aerosols from specific sources.
164 To examine the source-receptor relationships of aerosols, we implemented the Explicit Aerosol
165 Source Tagging (EAST) in E3SMv1, which play a critical role in attributing aerosol
166 concentrations to their respective emission sources. The EAST follows the BC source-tagging
167 technique introduced in H. Wang et al. (2014), sulfate source-tagging method used in Yang et
168 al. (2017) and other carbonaceous aerosol-tagging applied in Yang et al. (2018a), which was
169 previously implemented in the Community Atmosphere Model version 5 (CAM5-EAST). This
170 tagging system is different from the traditional emission sensitivity method that zero out or
171 perturb emissions from a given source region or sector in sensitivity simulations along with a
172 baseline simulation, which has to assume a linear response to emission perturbation and
173 requires many additional simulations for estimating the contributions from multiple sources
174 (Wang et al., 2014). EAST independently considers all aerosol physical, chemical, and
175 dynamical processes for each tagged sources through introducing additional aerosol-related
176 variables within one simulation and it does not rely on a linear response to emission
177 perturbations. These capabilities make it physically more accurate and time saving than the
178 sensitivity experiments. This tagging method has previously been adopted in regional models

179 and has now implemented in the global E3SMv1 model to better understand the
180 intercontinental transport from sources outside the regional domain.

181 In this study, totally 18 tags are set for each anthropogenic species of aerosols and
182 precursors. Each of the 4 major source regions, including North America (NAM), Europe
183 (EUR), East Asia (EAS) and South Asia (SAS), has 4 tags for the energy transformation and
184 extraction (ENE), industrial combustion and processes (IND), residential, commercial and
185 other (RCO) and the rest of anthropogenic emission sectors (RST). One tag is assigned to the
186 anthropogenic emissions from rest of the world (ROW) and the last tag is allocated for all the
187 natural/biogenic sources including open biomass burning, volcanic emissions and oceanic
188 emissions.

189 **2.3. Model evaluation**

190 In order to evaluate the performance of E3SMv1 in reproducing the aerosol concentrations,
191 Fig. 2 compares the simulated near-surface PM_{2.5} concentrations with the observations from
192 the Interagency Monitoring of Protected Visual Environments (IMPROVE) over the U.S., the
193 European Monitoring and Evaluation Programme (EMEP) over Europe, the U.S. embassies
194 and consulates in India and the China National Environmental Monitoring Center (CNEMC)
195 over China in year 2017. The model successfully reproduces the spatial distribution of PM_{2.5}
196 concentrations, with relatively high concentrations in eastern China, India and low
197 concentrations in the U.S. and Europe. The spatial correlation coefficient (R) between the
198 E3SMv1 simulated PM_{2.5} concentrations and observations is +0.80. The model well reproduces
199 the PM_{2.5} concentrations in the U.S. with the normalized mean biases (NMB) of -11%.
200 However, it largely underestimates the PM_{2.5} concentrations in China and Europe, which has

201 also been revealed in several studies (e.g., Gao et al., 2018; Gao et al., 2023; Navinya et al.,
202 2020; Zeng et al., 2021). This discrepancy is partly because E3SMv1 only considers limited
203 aerosol species (BC, POM, SOA and sulfate) without including nitrate and ammonium aerosols
204 in the aerosol module. On the other hand, the overestimated wet scavenging at the mid- and
205 high latitudes and underestimated gas-to-particle conversion can also lead to the low bias (Zeng
206 et al., 2021). The evaluation in 2010 also shows similar high correlation and biases (Fig. S3).

207 In order to evaluate the model performance in reproducing the historical changes in
208 aerosol concentrations during the important periods of emission changes, the variations in near-
209 surface PM_{2.5} concentrations are compared with observations (Fig. S4) and MERRA-2
210 reanalysis (Fig. S5). The model well reproduces the decreases in PM_{2.5} concentrations in the
211 eastern U.S. and Europe and the increases in East Asia and South Asia during 1980–2010, with
212 the spatial R of 0.78 between model results and MERRA-2 data. The model also well simulates
213 the aerosol decline in North America, Europe, and East Asia and a continuous increase in South
214 Asia during 2010–2017, with the R of 0.81 between model results and observational data.

215 However, the model simulation does capture the increases in PM_{2.5} in southwestern Canada
216 and eastern Russia. It is because the wildfire emissions were kept unchanged during the
217 simulation, while wildfires occurred more frequently in these regions during the analyzed time
218 period (Jolly et al., 2015; Goss et al., 2020), leading to the increases in PM_{2.5} in observations.

219 **3. Results**

220 **3.1. Historical changes in aerosols over major source regions**

221 Figure 3 shows the variations in anthropogenic emissions of sulfur dioxide (SO₂) during
222 the three key periods of historical emission changes from the tagged source regions. Since

223 industrialization, anthropogenic SO₂ emissions had rising trends during 1850–1980, especially
224 in Europe and North America. Due to the implementation of clean air actions in western
225 countries, SO₂ emissions in North America and Europe declined considerably during 1980–
226 2010, while the emissions in East Asia and South Asia continued to increase. After 2010, China
227 issued several clean air policies, which led to significant decreases in anthropogenic SO₂
228 emissions in East Asia, while the SO₂ emissions in South Asia kept increasing during 2010–
229 2017. The changes in anthropogenic BC and organic carbon (OC) emissions are similar to those
230 of SO₂ (shown in Figure S1 and S2).

231 The changes in near-surface mass concentrations (Fig. 4) and column burdens (Fig. 5) of
232 anthropogenic PM_{2.5} contributed by the tagged source regions during the focused three
233 historical time periods follow the corresponding changes in anthropogenic emissions. Column
234 burden refers to the concentration of aerosols contained in the air column above a unit area
235 with a top at 60 km, which can better reflect the aerosol transport within the air column and is
236 more related to the aerosol radiative effect. The near-surface concentration of aerosols
237 represents the concentration of aerosols in the air near the surface (from ~~1000~~the surface to
238 997 hPa for model layer), which is more related to air quality and human health. Local
239 anthropogenic emission changes drove the PM_{2.5} to reach its peak in 1980 in North America
240 and Europe and then to decrease. The maximum PM_{2.5} appeared in 2010 in East Asia and the
241 anthropogenic PM_{2.5} level continued growing in South Asia during 1850–2017, consistent with
242 previous studies (Dey et al.,2020; Guttikunda et al., 2022; Singh et al.,2023).

243 To explore which aerosol species contributed the most to the changes in PM_{2.5} during the
244 focused three historical time periods, Figs. 6 and 7 illustrate the relative contributions of

245 individual aerosols to the simulated changes in near-surface $PM_{2.5}$ mass concentrations and
246 column burdens, respectively, in four major emission regions. In general, the historical changes
247 in anthropogenic $PM_{2.5}$ were primarily driven by the changes in sulfate. In North America, the
248 contribution of sulfate to near-surface $PM_{2.5}$ concentration (column burden) rose from 7%
249 (11%) in 1850 to 67% (81%) in 1980, then dropped to 52% (67%) in 2017. In Europe, sulfate
250 contribution changed from 24% (34%) in 1850 to 71% (85%) in 1980, then decreased to 50%
251 (68%) in 2017. In East Asia, sulfate contribution changes from 2% (6%) in 1850 to 51% (71%)
252 in 1980, then decreased to 33% (56%) in 2017. It is interesting that the $PM_{2.5}$ levels in East
253 Asia increased during 1980–2010, but the sulfate contribution decreased in this time period. It
254 is because the carbonaceous aerosols increased remarkably, which reduced the fractional
255 contribution of sulfate. The sulfate contribution to $PM_{2.5}$ concentration (column burden)
256 increased throughout the period of 1850–2017, from 2% (5%) to 42% (62%) in South Asia.
257 Note that the carbonaceous aerosols, especially POM, dominated the $PM_{2.5}$ in all four targeted
258 regions in 1850, resulting from the high heating demand from the residential sector.

259 **3.2 Changes in contributions from major source regions and sectors**

260 Figure 8 shows the relative contributions from local and remote anthropogenic sources to
261 the near-surface concentrations and column burdens of $PM_{2.5}$ in the four targeted regions in
262 2010. Local sources dominated the near-surface anthropogenic $PM_{2.5}$ concentrations over the
263 high emission regions including eastern China, eastern U.S. and Indo-Gangetic Plain, with
264 local contributions being higher than 90%. In the regions with low emissions, such as the
265 Tibetan Plateau, the anthropogenic $PM_{2.5}$ concentrations were largely contributed by the long-
266 range transport of aerosols. The spatial distributions of burden contribution are similar to those

267 of corresponding contribution near the surface, but the long-range transport contributed more
268 to the column burden than to the near-surface contribution due to the more efficient pollutant
269 transport in the free troposphere than within the boundary layer. The long-range transport
270 contributes 30%–35% of PM_{2.5} burden in East Asia and South Asia and 50–55% in North
271 America and Europe, much higher than the 10%–25% for the near-surface concentrations over
272 the four targeted regions.

273 Since both local and remote emissions can contribute to the anthropogenic PM_{2.5}, it is
274 valuable to know the historical changes in these contributions, especially by the local sources
275 from individual emission sectors and by remote sources from major emission regions. Figure
276 9 illustrates changes in the local source contributions from major emission sectors to the near-
277 surface concentrations and column burdens of PM_{2.5} during 1850–2017. In North America and
278 Europe, the historical changes in anthropogenic PM_{2.5} were largely induced by changes in
279 emissions from the local energy (ENE) sector, followed by the industry (IND) sector, which
280 increased before 1980 and decreased afterward. In East Asia, ENE, IND and residential (RCO)
281 sector emissions all had significant contributions to the increases in PM_{2.5} concentration and
282 burden from 1850 to 1980. Then the contribution from local IND sector showed the largest
283 increases from 1980 to 2010 and decreases from 2010 to 2017. In South Asia, the
284 anthropogenic PM_{2.5} increases from 1850 to 1980 were mainly attributed to the RCO emission
285 increases. After that, increases in ENE emissions dominated the rising PM_{2.5} levels in South
286 Asia during 1980–2017.

287 Figure 10 presents changes in remote emission contributions from the tagged source
288 regions to the column burdens of PM_{2.5} during 1850–2017. The contributions from long-range

289 transport to the near-surface concentrations and their historical variations over the four major
290 emissions regions are relatively small, which were also reported in previous studies (e.g., Yang
291 et al., 2018b) and are not discussed here. During 1850–1980, the long-range transport
292 contributions to the PM_{2.5} burdens show increases and the contributions from ROW increased
293 the most among the tagged source regions over all four targeted receptor regions. Note that
294 aerosol emissions from North America contributed to the increase in European PM_{2.5} burden
295 by 1.7 mg m⁻² and sources from Europe were also responsible for the PM_{2.5} burden increase by
296 1.0 mg m⁻² in East Asia and 1.1 mg m⁻² in South Asia. During 1980–2010, the long-range
297 transport from North America and Europe decreased, but that from East Asia and South Asia
298 increased. East Asia contributed 0.4–0.6 mg m⁻² to the PM_{2.5} burden increases in North
299 America, while Europe and South Asia contributed about 0.3 mg m⁻². In East Asia, 1.6 mg m⁻²
300 of the PM_{2.5} burden increase was attributed to South Asian sources and 0.8 mg m⁻² of the
301 PM_{2.5} burden increase in South Asia during this time period was due to increases in East Asian
302 emissions. From 2010 to 2017, owing to the clean air actions in China, contributions from East
303 Asia to the PM_{2.5} burdens in the other three targeted regions decreased by 0.3–0.6 mg m⁻².
304 However, due to the continuous increases in South Asian emissions, South Asia still
305 contributed to the PM_{2.5} burden increase in East Asia by 0.4 mg m⁻² during this short time
306 period.

307 **3.3 Changes in effective radiative forcing due to aerosol-radiation interactions**

308 The variation in aerosols can have a significant impact on ERF through aerosol-radiation
309 and aerosol-cloud interactions. Figure 11 shows changes in ERF due to aerosol-radiation
310 interactions (ERF_{ari}) at the top of the atmosphere (TOA) that can be attributed to changes in

311 anthropogenic emissions from the tagged regions in the three key periods during 1850–2017.
312 Due to the increases in aerosols from 1850 to 1980, a large negative ERF_{ari} was located over
313 the major source regions and their downwind areas, with maximum ERF_{ari} changes being larger
314 than 2 W m^{-2} over eastern U.S., Europe and eastern China. In 2010, there were positive ERF_{ari}
315 changes (warming effect) by a maximum of 2 W m^{-2} in North America and Europe compared
316 to 1980, which were due to the decreases in anthropogenic aerosols in these two regions. The
317 positive ERF_{ari} changes due to the decrease in North American aerosols extended across the
318 North Atlantic and caused an increase in incoming radiation by up to 0.5 W m^{-2} in Europe.
319 Similarly, the decrease in aerosols from Europe also led to a positive ERF_{ari} change by up to
320 0.5 W m^{-2} in the downwind regions including Russia and northern China during 1980–2010.
321 This is also revealed by previous studies that aerosol reduction has caused fast warming in
322 downwind regions (Urdiales-Flores et al., 2023). Increases in aerosols in China during 1980–
323 2010 and decreases during 2010–2017 produced negative (cooling) and positive (warming)
324 changes in ERF_{ari} , respectively, over eastern China and North Pacific, which largely
325 contradicted each other. The continuously growing aerosols in South Asia induced negative
326 ERF_{ari} changes (cooling) in South Asia, Southeast Asia and southern China during both 1980–
327 2010 and 2010–2017. Note that in this study we only quantify the ERF_{ari} for the major emission
328 regions based on the source tagging technique. The quantification of ERF due to aerosol-cloud
329 interactions (ERF_{aci}) requires additional simulations, which could be further examined in future
330 studies.

331 **4. Conclusions and Discussions**

332 Since the start of industrialization, aerosols have changed significantly in different regions
333 of the world driven by global economic development and air pollution control measures. It is
334 of great significance to quantify the contributions of aerosols from major emission source
335 regions and sectors during the key periods of substantial emission changes. In this study, the
336 Explicit Aerosol Source Tagging (EAST) technique is implemented in E3SMv1 to quantify the
337 variations in the concentrations, source contributions, and the subsequent effective radiative
338 forcing of anthropogenic aerosols in four major source regions (NAM, EUR, EAS and SAS)
339 during three key historical periods of emission changes (1850–1980, 1980–2010 and 2010–
340 2017).

341 Following the corresponding anthropogenic emission changes, $PM_{2.5}$ concentrations
342 reached its peak in 1980 in North America and Europe, while the peak of $PM_{2.5}$ in East Asia
343 occurred in 2010. The $PM_{2.5}$ from anthropogenic sources in South Asia continued growing
344 during 1850–2017. These changes in anthropogenic $PM_{2.5}$ were primarily dominated by
345 changes in sulfate aerosol. In North America and Europe, historical changes in anthropogenic
346 $PM_{2.5}$ were mainly caused by changes in emissions from local energy sector, followed by the
347 industrial sector, which increased from 1850 to 1980 and decreased afterward. In East Asia,
348 energy, industrial, and residential emissions contributed significantly to the increase in $PM_{2.5}$
349 from 1850 to 1980, then the local industrial sector caused the most increase from 1980 to 2010,
350 and declined from 2010 to 2017. For South Asia, the increase in $PM_{2.5}$ was mainly due to
351 emission changes in the residential sector from 1850 to 1980, then the increase in energy-
352 related emissions became dominant to the rise of $PM_{2.5}$ levels during 1980–2017.

353 Regional aerosol pollution comes from both local emissions and long-range transport of
354 remote emissions. Local emissions contribute the most in regions with high emissions, while
355 in regions with low emissions the long-distance transport plays an important role. Due to the
356 more efficient transport of air pollutants in the free troposphere, contributions of long-range
357 transport to column burden are greater than to the near-surface concentration over all four
358 targeted receptor regions. From 1850 to 1980, increases in emissions from North America
359 contributed to the increase in European $\text{PM}_{2.5}$ burden by 1.7 mg m^{-2} and emissions from Europe
360 were also responsible for the $\text{PM}_{2.5}$ burden increase by 1.0 mg m^{-2} in East Asia and 1.1 mg m^{-2}
361 in South Asia. From 1980 to 2010, the long-range transport from North America and Europe
362 decreased, while those from East Asia and South Asia increased. East Asia contributed 0.4 –
363 0.6 mg m^{-2} to the $\text{PM}_{2.5}$ burden increases in North America, while Europe and South Asia
364 contributed about 0.3 mg m^{-2} . In East Asia, 1.6 mg m^{-2} of the $\text{PM}_{2.5}$ burden increase was
365 attributed to South Asian sources and 0.8 mg m^{-2} of the $\text{PM}_{2.5}$ burden increase in South Asia
366 during this time period was due to the increases in East Asian emissions. From 2010 to 2017,
367 the contributions from East Asia to the $\text{PM}_{2.5}$ burdens in the other three targeted regions
368 declined by 0.3 – 0.6 mg m^{-2} due to Clean Air actions in China. However, due to the continuous
369 increase of emissions in South Asia, the $\text{PM}_{2.5}$ burden in East Asia still increased by 0.4 mg m^{-2} .
370

371 Changes in aerosols can have a significant impact on ERF, which further imposes an
372 impact on climate. Large negative ERF_{ari} appeared over the major source regions and their
373 downwind areas during 1850–1980 due to the increases in aerosol emissions, with maximum
374 ERF_{ari} changes being larger than 2 W m^{-2} over eastern North America, Europe and eastern

375 China. From 1980 to 2010, a positive ERF_{ari} change caused by a decline in North American
376 aerosols extended over the North Atlantic, resulting in a warming of up to 0.5 W m^{-2} in Europe.
377 Meanwhile, a decline of aerosols in Europe also caused a warming of up to 0.5 W m^{-2} in Russia
378 and northern China. The changes in ERF_{ari} from the increase (from 1980 to 2010) and decrease
379 (from 2010 to 2017) of aerosols in China had an opposite sign. The continuous aerosol
380 increases in South Asia from 1980 to 2017 resulted in negative ERF_{ari} changes in South Asia,
381 Southeast Asia, and southern China.

382 This study provides an in-depth analysis of historical changes in anthropogenic aerosol
383 concentrations, compositions, source contributions and radiative impacts in the four major
384 emission source regions of the globe, which has important implications for the pollution
385 prevention/control measures and decision making for global collaboration. The spatial
386 distribution and changes in anthropogenic aerosols are similar to those reported in previous
387 studies (Hoesly, et al., 2018; Lim, et al., 2020). However, we also note that the E3SMv1 model
388 underestimates the near-surface $PM_{2.5}$ concentrations in Europe and East Asia, which could
389 lead to an underestimate of the corresponding radiative and climate impact. Our analysis
390 focuses on aerosols from anthropogenic emissions; however, with the increasing attention to
391 air quality in many countries around the world, anthropogenic aerosol concentrations are
392 declining and contributions from biomass burning aerosols are becoming more and more
393 important. The source contributions and impacts of biomass burning aerosols will be
394 investigated in our future work. Also, this study only quantifies the ERF_{ari} for individual major
395 emission regions based on the source tagging technique and radiation diagnostic calculations.

396 The quantification of ERF_{aci} requires additional simulations, which could be further examined
397 in future studies.

398 **Author contributions.** YY designed the research, added the tagging code and performed
399 simulations; SM analyzed the data. All authors including HW, PW, and HL discussed the
400 results and wrote the paper.

401

402 **Code and data availability.** The E3SMv1 model is available at <https://github.com/E3SM->
403 [Project/E3SM](https://github.com/E3SM-Project/E3SM)(last access:1 October 2022) (<https://doi.org/10.11578/E3SM/dc.20180418.36>;
404 E3SM Project, 2018). Our results can be made available upon request.

405

406 **Acknowledgments.** HW acknowledges the support by the U.S. Department of Energy (DOE),
407 Office of Science, Office of Biological and Environmental Research (BER), as part of the
408 Earth and Environmental System Modeling program. The Pacific Northwest National
409 Laboratory (PNNL) is operated for DOE by the Battelle Memorial Institute under contract
410 DE-AC05-76RLO1830.

411

412 **Financial support.** This study was supported by the National Key Research and
413 Development Program of China (grant 2020YFA0607803 ~~and 2019YFA0606800~~), ~~the~~
414 ~~National Natural Science Foundation of China (grant 41975159)~~, Jiangsu Science Fund for
415 Distinguished Young Scholars (grant BK20211541), and the Jiangsu Science Fund for
416 Carbon Neutrality (grant ~~no.~~ BK20220031).

417

418 **Conflict of interest.**

419 At least one of the (co-)authors is a member of the editorial board of ACP.

420 **References**

- 421
- 422 Abdul-Razzak, H. and Ghan, S. J.: A parameterization of aerosol activation, 2. Multiple
423 aerosol types, *J. Geophys. Res.*, 105, 6837–6844,
424 <https://doi.org/10.1029/1999JD901161>, 2000.
- 425
- 426 Akimoto, H.: Global Air Quality and Pollution, *Science*, 302, 1716–1719,
427 <https://doi.org/10.1126/science.1092666>, 2003.
- 428
- 429 Anenberg, S. C., Horowitz, L. W., Tong, D. Q., and West, J. J.: An Estimate of the Global
430 Burden of Anthropogenic Ozone and Fine Particulate Matter on Premature Human
431 Mortality Using Atmospheric Modeling, *Environ. Health Perspect.*, 118, 1189–1195,
432 <https://doi.org/10.1289/ehp.0901220>, 2010.
- 433
- 434 Anenberg, S. C., West, J. J., Yu, H., Chin, M., Schulz, M., Bergmann, D., Bey, I., Bian, H.,
435 Diehl, T., Fiore, A., Hess, P., Marnmer, E., Montanaro, V., Park, R., Shindell, D.,
436 Takemura, T., and Dentener, F.: Impacts of intercontinental transport of anthropogenic
437 fine particulate matter on human mortality, *Air. Qual. Atmos. Health.*, 7, 369–379,
438 <https://doi.org/10.1007/s11869-014-0248-9>, 2014.
- 439
- 440 Carslaw, K. S., Gordon, H., Hamilton, D. S., Johnson, J. S., Regayre, L. A., Yoshioka, M.,
441 and Pringle, K. J.: Aerosols in the Pre-industrial Atmosphere, *Curr. Clim. Change Rep.*,
442 3, 1–15, <https://doi.org/10.1007/s40641-017-0061-2>, 2017.
- 443
- 444 Chen, D., Liao, H., Yang, Y., Chen, L., and Wang, H.: Simulated aging processes of black
445 carbon and its impact during a severe winter haze event in the Beijing-Tianjin-Hebei
446 region, *Sci. Total Environ.*, 755, 142712,
447 <https://doi.org/10.1016/j.scitotenv.2020.142712>, 2021.
- 448
- 449 Chin, M., Diehl, T., Ginoux, P., and Malm, W.: Intercontinental transport of pollution and
450 dust aerosols: implications for regional air quality, *Atmos. Chem. Phys.*, 7, 5501–5517,
451 <https://doi.org/10.5194/acp-7-5501-2007>, 2007.
- 452
- 453 David, L. M., Ravishankara, A. R., Kodros, J. K., Pierce, J. R., Venkataraman, C., and
454 Sadavarte, P.: Premature mortality due to PM_{2.5} over India: Effect of atmospheric
455 transport and anthropogenic emissions, *Geohealth*, 3, 2–10,
456 <https://doi.org/10.1029/2018GH000169>, 2019.
- 457
- 458 Dey, S., Purohit, B., Balyan, P., Dixit, K., Bali, K., Kumar, A., Imam, F., Chowdhury, S.,
459 Ganguly, D., Gargava, P., and Shukla, V. K.: A satellite-based high-resolution (1-km)
460 ambient PM_{2.5} database for India over two decades (2000–2019): applications for air
461 quality management, *Remote Sens.*, 12, 3872, <https://doi.org/10.3390/rs12233872>, 2022.
- 462

463 Finlayson-Pitts, B. J., and Pitts, J. N.: Tropospheric Air Pollution: Ozone, Airborne Toxics,
464 Polycyclic Aromatic Hydrocarbons, and Particles, *Science*, 276, 1045–1051,
465 <https://doi.org/10.1126/science.276.5315.1045>, 1997.
466

467 Gao, M., Beig, G., Song, S., Zhang, H., Hu, J., Ying, Q., Liang, F., Liu, Y., Wang, H., Lu, X.,
468 Zhu, T., Carmichael, G. R., Nielsen, C. P., McElroy, M. B.: The impact of power
469 generation emissions on ambient PM_{2.5} pollution and human health in China and India,
470 *Environ. Int.*, 121, 250-259, <https://doi.org/10.1016/j.envint.2018.09.015>, 2018.
471

472 Gao, J., Yang, Y., Wang, H., Wang, P., Li, H., Li, M., Ren, L., Yue, X., and Liao, H.: Fast
473 climate responses to emission reductions in aerosol and ozone precursors in China
474 during 2013–2017, *Atmos. Chem. Phys.*, 22, 7131–7142, <https://doi.org/10.5194/acp-22-7131-2022>, 2022.
475

476

477 Gao, J., Yang, Y., Wang, H., Wang, P., Li, B., Li, J., Wei, J., Gao, M., and Liao, H.: Climate
478 responses in China to domestic and foreign aerosol changes due to clean air actions
479 during 2013–2019, *npj Clim. Atmos. Sci.*, 6, 160, <https://doi.org/10.1038/s41612-023-00488-y>, 2023.
480

481

482 Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., Randles, C. A.,
483 Darmenov, A., Bosilovich, M. G., Reichle, R., Wargan, K., Coy, L., Cullather, R.,
484 Draper, C., Akella, S., Buchard, V., Conaty, A., da Silva, A. M., Gu, W., Kim, G.,
485 Koster, R., Lucchesi, R., Merkova, D., Nielsen, J. E., Partyka, G., Pawson, S., Putman,
486 W., Rienecker, M., Schubert, S. D., Sienkiewicz, M., and Zhao, B.: The Modern-Era
487 Retrospective Analysis for Research and Applications, Version 2 (MERRA-2), *J.*
488 *Climate*, 30, 5419–5454, <https://doi.org/10.1175/JCLI-D-16-0758.1>, 2017.
489

490 Ghan, S. J.: Technical Note: Estimating aerosol effects on cloud radiative forcing, *Atmos.*
491 *Chem. Phys.*, 13, 9971–9974, <https://doi.org/10.5194/acp-13-9971-2013>, 2013.
492

493 Ghan, S. J. and Zaveri, R. A.: Parameterization of optical properties for hydrated internally
494 mixed aerosol, *J. Geophys. Res.-Atmos.*, 112, D10201,
495 <https://doi.org/10.1029/2006JD007927>, 2007.
496

497 Golaz, J. C., Caldwell, P. M., Van Roekel, L. P., Petersen, M. R., Tang, Q., Wolfe, J. D.,
498 Abeshu, G., Anantharaj, V., Asay-Davis, X. S., Bader, D. C., Baldwin, S. A., Bisht, G.,
499 Bogenschutz, P. A., Branstetter, M., Brunke, M. A., Brus, S. R., Burrows, S. M.,
500 Cameron-Smith, P. J., Donahue, A. S., Deakin, M., Easter, R. C., Evans, K. J., Feng, Y.,
501 Flanner, M., Foucar, J. G., Fyke, J. G., Griffin, B. M., Hannay, C., Harrop, B. E.,
502 Hoffman, M. J., Hunke, E. C., Jacob, R. L., Jacobsen, D. W., Jeffery, N., Jones, P. W.,
503 Keen, N. D., Klein, S. A., Larson, V. E., Leung, L. R., Li, H. Y., Lin, W., Lipscomb, W.
504 H., Ma, P., Mahajan, S., Maltrud, M. E., Mامتjanov, A., McClean, J. L., McCoy, R.
505 B., Neale, R. B., Price, S. F., Qian, Y., Rasch, P. J., Eyre, J. E. J. R., Riley, W. J.,
506 Ringler, T. D., Roberts, A. F., Roesler, E. L., Salinger, A. G., Shaheen, Z., Shi, X.,

507 Singh, B., Tang, J., Taylor, M. A., Thornton, P. E., Turner, A. K., Veneziani, M., Wan,
508 H., Wang, H., Wang, S., Williams, D. N., Wolfram, P. J., Worley, P. H., Xie, S., Yang
509 Y., Yoon, J., Zelinka, M. D., Zender, C. S., Zeng, X., Zhang, C., Zhang, K., Zhang, Y.,
510 Zheng, X., Zhou, T., and Zhu, Q.: The DOE E3SM Coupled Model Version 1:
511 Overview and Evaluation at Standard Resolution, *J. Adv. Model. Earth Syst.*, 11, 2089–
512 2129. <https://doi.org/10.1029/2018MS001603>, 2019.

513

514 Goss, M., Swain, D. L., Abatzoglou, J. T., Sarhadi, A., Kolden, C. A., Williams A. P., and
515 Diffenbaugh, N. S.: Climate change is increasing the likelihood of extreme autumn
516 wildfire conditions across California, *Environ. Res. Lett.*, 15, 094016,
517 <https://doi.org/10.1088/1748-9326/ab83a7>, 2020.

518

519 Guttikunda, S., and Nishadh, K. A.: Evolution of India's PM_{2.5} pollution between 1998 and
520 2020 using global reanalysis fields coupled with satellite observations and fuel
521 consumption patterns, *Environ. Sci. Atmos.*, 2, 1502-1515,
522 <https://doi.org/10.1039/D2EA00027J>, 2022.

523

524 Hoesly, R. M., Smith, S. J., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen, T.,
525 Seibert, J. J., Vu, L., Andres, R. J., Bolt, R. M., Bond, T. C., Dawidowski, L., Kholod,
526 N., Kurokawa, J., Li, M., Liu, L., Lu, Z., Moura, M. C. P., O'Rourke, P. R., and Zhang,
527 Q.: Historical (1750–2014) anthropogenic emissions of reactive gases and aerosols from
528 the Community Emissions Data System (CEDS), *Geosci. Model Dev.*, 11, 369–408,
529 <https://doi.org/10.5194/gmd-11-369-2018>, 2018.

530

531 Jaffe, D., Anderson, T., Covert, D., Kotchenruther, R., Trost, B., Danielson, J., Simpson, W.,
532 Berntsen, T., Karlsdottir, S., Blake, D., Harris, J., Carmichael, G., and Uno, I.: Transport
533 of Asian air pollution to North America, *Geophys. Res. Lett.*, 26, 711–714,
534 <https://doi.org/10.1029/1999GL900100>, 1999.

535

536 Jolly, W. M., Cochrane, M. A., Freeborn, P. H., Holden, Z. A., Brown, T. J., Williamson, G.
537 J., and Bowman, D. M. J. S.: Climate-induced variations in global wildfire danger from
538 1979 to 2013, *Nat. Commun.*, 6, 7537, <https://doi.org/10.1038/ncomms8537>, 2015.

539

540 Li, C., McLinden, C., Fioletov, V., Krotkov, N., Carn, S., Joiner, J., Streets, D., He, H., Ren,
541 X., Li, Z., and Dickerson, R.: India Is Overtaking China as the World's Largest Emitter
542 of Anthropogenic Sulfur Dioxide, *Sci. Rep.*, 7, 14304, <https://doi.org/10.1038/s41598-017-14639-8>, 2017.

544

545 Li, H., Yang, Y., Wang, H., Li, B., Wang, P., Li, J., and Liao, H.: Constructing a
546 spatiotemporally coherent long-term PM_{2.5} concentration dataset over China using a
547 machine learning approach, *Sci. Total Environ.*, 765, 144263,
548 <https://doi.org/10.1016/j.scitotenv.2020.144263>, 2021.

549

550 Liao, H., Chang, W., and Yang, Y.: Climatic effects of air pollutants over China: A review,
551 *Adv. Atmos. Sci.*, 32, 115–139, <https://doi.org/10.1007/s00376-014-0013-x>, 2015.

552

553 Lim, C.-H., Ryu, J., Choi, Y., Jeon, S. W., and Lee, W.-K.: Understanding global PM_{2.5}
554 concentrations and their drivers in recent decades (1998–2016), *Environ. Int.*, 144,
555 106011, <https://doi.org/10.1016/j.envint.2020.106011>, 2020.

556

557 Lin, J., Pan, D., Davis, S. J., Zhang, Q., He, K., Wang, C., Streets, D. G., Wuebbles, D. J.,
558 and Guan, D.: China’s international trade and air pollution in the United States, *Proc.*
559 *Natl. Acad. Sci.*, 111, 1736–1741, <https://doi.org/10.1073/pnas.1312860111>, 2014.

560

561 Liu, J., and Mauzerall, D.: Estimating the average time for inter-continental transport of air
562 pollutants, *Geophys. Res. Lett.*, 32, L11814, <https://doi.org/10.1029/2005GL022619>,
563 2005.

564

565 Liu, J., Mauzerall, D. L., Horowitz, L. W., Ginoux, P., and Fiore, A. M.: Evaluating inter-
566 continental transport of fine aerosols: (1) Methodology, global aerosol distribution and
567 optical depth, *Atmos. Environ.*, 43, 4327–4338,
568 <https://doi.org/10.1016/j.atmosenv.2009.03.054>, 2009.

569

570 Lohmann, U., and Feichter, J.: Global indirect aerosol effects: a review, *Atmos. Chem. Phys.*,
571 5, 715–737, <https://doi.org/10.5194/acp-5-715-2005>, 2005.

572

573 Lou, S., Yang, Y., Wang, H., Smith, S. J., Qian, Y., and Rasch, P. J.: Black carbon amplifies
574 haze over the North China Plain by weakening the East Asian winter monsoon,
575 *Geophys. Res. Lett.*, 46, 452–460, <https://doi.org/10.1029/2018GL080941>, 2019.

576

577 Navinya, C. D., Vinoj, V. and Pandey, S. K.: Evaluation of PM_{2.5} Surface Concentrations
578 Simulated by NASA’s MERRA Version 2 Aerosol Reanalysis over India and its relation
579 to the Air Quality Index, *Aerosol Air Qual. Res.*, 20, 1329–1339,
580 <https://doi.org/10.4209/aaqr.2019.12.0615>, 2020.

581

582 Pöschl, U.: Atmospheric Aerosols: Composition, Transformation, Climate and Health
583 Effects, *Angew. Chem. Int. Ed.*, 44, 7520–7540,
584 <https://doi.org/10.1002/anie.200501122>, 2005.

585

586 Rasch, P. J., Xie, S., Ma, P.-L., Lin, W., Wang, H., Tang, Q., Burrows, S. M., Caldwell, P.,
587 Zhang, K., Easter, R. C., Cameron-Smith, P., Singh, B., Wan, H., Golaz, J.-C., Harrop,
588 B. E., Roesler, E., Bacmeister, J., Larson, V. E., Evans, K. J., Qian, Y., Taylor, M.,
589 Leung, L. R., Zhang, Y., Brent, L., Branstetter, M., Hannay, C., Mahajan, S.,
590 Mامتjanov, A., Neale, R., Richter, J. H., Yoon, J.-H., Zender, C. S., Bader, D.,
591 Flanner, M., Foucar, J. G., Jacob, R., Keen, N., Klein, S. A., Liu, X., Salinger, A. G.,
592 Shrivastava, M., and Yang, Y.: An overview of the atmospheric component of the

593 Energy Exascale Earth System Model, *J. Adv. Model. Earth Sy.*, 11, 2377–2411,
594 <https://doi.org/10.1029/2019MS001629>, 2019.

595

596 Ren, L., Yang, Y., Wang, H., Wang, P., Chen, L., Zhu, J., and Liao, H.: Aerosol transport
597 pathways and source attribution in China during the COVID-19 outbreak, *Atmos. Chem.*
598 *Phys.*, 21, 15431–15445, <https://doi.org/10.5194/acp-21-15431-2021>, 2021.

599

600 Rosenfeld, D., Lohmann, U., Raga, G. B., Kulmala, M., Fuzzi, S., Reissell, A., and Andreae,
601 M. O.: Flood or Drought: How Do Aerosols Affect Precipitation?, *Science*, 321, 1309–
602 1313, <https://doi.org/10.1126/science.1160606>, 2008.

603

604 Singh, T., Matsumi, Y., Nakayama, T., Hayashida, S., Patra, P. K., Yasutomi, N., Kajino, M.,
605 Yamaji, K., Khatri, P., Takigawa, M., Araki, H., Kurogi, Y., Kuji, M., Muramatsu, K.,
606 Imasu, R., Ananda, A., Arbain, A. A., Ravindra, K., Bhardwaj, S., Kumar, S., Mor, S.,
607 Dhaka, S. K., Dimri, A. P., Sharma, A., Singh, N., Bhatti, M. S., Yadav, R., Vatta, K.,
608 and Mor, S.: Very high particulate pollution over northwest India captured by a high-
609 density in situ sensor network, *Sci. Rep.*, 13, 13201, [https://doi.org/10.1038/s41598-023-](https://doi.org/10.1038/s41598-023-39471-1)
610 [39471-1](https://doi.org/10.1038/s41598-023-39471-1), 2023.

611

612 Smith, S. J., van Aardenne, J., Klimont, Z., Andres, R. J., Volke, A., and Arias, S. D.:
613 Anthropogenic sulfur dioxide emissions: 1850–2005, *Atmos. Chem. Phys.*, 11, 1101–
614 1116, <https://doi.org/10.5194/acp-11-1101-2011>, 2011.

615

616 Stohl, A., and Trickl, T.: A textbook example of long-range transport: Simultaneous
617 observation of ozone maxima of stratospheric and North American origin in the free
618 troposphere over Europe, *J. Geophys. Res.*, 104, 30445–30462,
619 <https://doi.org/10.1029/1999JD900803>, 1999.

620

621 [Urdiales-Flores, D., Zittis, G., Hadjinicolaou, P., Osipov, S., Klingmüller, K., Mihalopoulos,](https://doi.org/10.1038/s41612-023-00423-1)
622 [N., Kanakidou, M., Economou, T., and Lelieveld, J.: Drivers of accelerated warming in](https://doi.org/10.1038/s41612-023-00423-1)
623 [Mediterranean climate-type regions, *npj. Clim. Atmos. Sci.*, 6, 97,](https://doi.org/10.1038/s41612-023-00423-1)
624 <https://doi.org/10.1038/s41612-023-00423-1>, 2023.

625

626 van Marle, M. J. E., Kloster, S., Magi, B. I., Marlon, J. R., Daniau, A.-L., Field, R. D.,
627 Arneeth, A., Forrest, M., Hantson, S., Kehrwald, N. M., Knorr, W., Lasslop, G., Li, F.,
628 Mangeon, S., Yue, C., Kaiser, J. W., and van der Werf, G. R.: Historic global biomass
629 burning emissions for CMIP6 (BB4CMIP) based on merging satellite observations with
630 proxies and fire models (1750–2015), *Geosci. Model Dev.*, 10, 3329–3357,
631 <https://doi.org/10.5194/gmd-10-3329-2017>, 2017.

632

633 Wang, H., Rasch, P. J., Easter, R. C., Singh, B., Zhang, R., Ma, P.-L., Qian, Y., Ghan, S. J.,
634 and Beagley, N.: Using an explicit emission tagging method in global modeling of
635 source-receptor relationships for black carbon in the Arctic: Variations, sources, and

636 transport pathways, *J. Geophys. Res. Atmos.*, 119, 12888–12909,
637 <https://doi.org/10.1002/2014JD022297>, 2014.

638

639 Wang, H., Easter, R. C., Zhang, R., Ma, P.-L., Singh, B., Zhang, K., Ganguly, D., Rasch, P.
640 J., Burrows, S. M., Ghan, S. J., Lou, S., Qian, Y., Yang, Y., Feng, Y., Flanner, M.,
641 Leung, L. R., Liu, X., Shrivastava, M., Sun, J., Tang, Q., Xie, S., and Yoon, J.-H.:
642 Aerosols in the E3SM Version 1: New developments and their impacts on radiative
643 forcing, *J. Adv. Model. Earth Syst.*, 12, e2019MS001851,
644 <https://doi.org/10.1029/2019MS001851>, 2020.

645

646 Wang, P., Yang, Y., Xue, D., Ren, L., Tang, J., Leung, L. R., and Liao, H.: Aerosols overtake
647 greenhouse gases causing a warmer climate and more weather extremes toward carbon
648 neutrality, *Nat. Commun.*, 14, 7257, <https://doi.org/10.1038/s41467-023-42891-2>, 2023.

649

650 Wang, Y., Liu, X., Hoose, C., and Wang, B.: Different contact angle distributions for
651 heterogeneous ice nucleation in the Community Atmospheric Model version 5, *Atmos.*
652 *Chem. Phys.*, 14, 10411–10430, <https://doi.org/10.5194/acp-14-10411-2014>, 2014.

653

654 Yang, Y., Wang, H., Smith, S. J., Easter, R., Ma, P.-L., Qian, Y., Yu, H., Li, C., and Rasch,
655 P. J.: Global source attribution of sulfate concentration and direct and indirect radiative
656 forcing, *Atmos. Chem. Phys.*, 17, 8903–8922, [https://doi.org/10.5194/acp-17-8903-](https://doi.org/10.5194/acp-17-8903-2017)
657 [2017](https://doi.org/10.5194/acp-17-8903-2017), 2017.

658

659 Yang, Y., Wang, H., Smith, S. J., Zhang, R., Lou, S., Qian, Y., Ma, P., and Rasch P. J.:
660 Recent intensification of winter haze in China linked to foreign emissions and
661 meteorology, *Sci. Rep.*, 8, 2107, <https://doi.org/10.1038/s41598-018-20437-7>, 2018a.

662

663 Yang, Y., Wang, H., Smith, S. J., Zhang, R., Lou, S., Yu, H., Li, C., and Rasch, P. J.: Source
664 Apportionments of Aerosols and Their Direct Radiative Forcing and Long-Term Trends
665 Over Continental United States, *Earth's Future*, 6, 793–808,
666 <https://doi.org/10.1029/2018EF000859>, 2018b.

667

668 Yang, Y., Ren, L., Li, H., Wang, H., Wang, P., Chen, L., Yue, X., and Hong, L.: Fast climate
669 responses to aerosol emission reductions during the COVID-19 pandemic, *Geophys.*
670 *Res. Lett.*, 47, e2020GL089788, <https://doi.org/10.1029/2020GL089788>, 2020a.

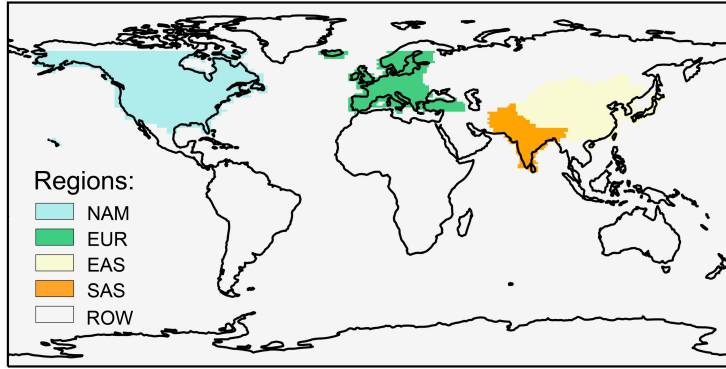
671

672 Yang, Y., Lou, S., Wang, H., Wang, P., and Liao, H.: Trends and source apportionment of
673 aerosols in Europe during 1980–2018, *Atmos. Chem. Phys.*, 20, 2579–2590,
674 <https://doi.org/10.5194/acp-20-2579-2020>, 2020b.

675

676 Yang, Y., Ren, L., Wu, M., Wang, H., Song, F., Leung, L. R., Hao, X., Li, J., Chen, L., Li,
677 H., Zeng, L., Zhou, Y., Wang, P., Liao, H., Wang, J., and Zhou, Z.-Q.: Abrupt emissions
678 reductions during COVID-19 contributed to record summer rainfall in China, *Nat.*
679 *Commun.*, 13, 959, <https://doi.org/10.1038/s41467-022-28537-9>, 2022a

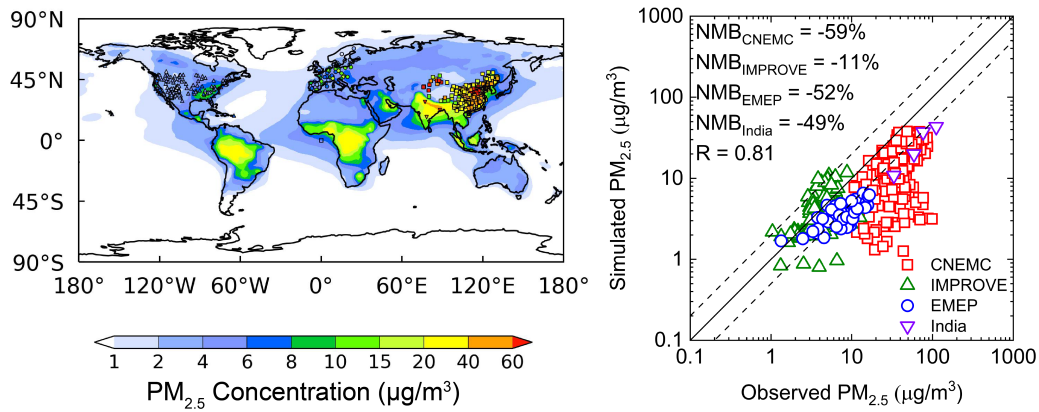
680
681 Yang, Y., Zeng, L., Wang, H., Wang, P., and Liao, H.: Dust pollution in China affected by
682 different spatial and temporal types of El Niño, *Atmos. Chem. Phys.*, 22, 14489–14502,
683 <https://doi.org/10.5194/acp-22-14489-2022>, 2022b.
684
685 Yang, Y., Zeng, L., Wang, H., Wang, P., and Liao, H.: Climate effects of future aerosol
686 reductions for achieving carbon neutrality in China, *Sci. Bull.*, 68, 902–905,
687 <https://doi.org/10.1016/j.scib.2023.03.048>, 2023..
688
689 Zeng, L., Yang, Y., Wang, H., Wang, J., Li, J., Ren, L., Li, H., Zhou, Y., Wang, P., and Liao,
690 H.: Intensified modulation of winter aerosol pollution in China by El Niño with short
691 duration, *Atmos. Chem. Phys.*, 21, 10745–10761, [https://doi.org/10.5194/acp-21-10745-](https://doi.org/10.5194/acp-21-10745-2021)
692 [2021](https://doi.org/10.5194/acp-21-10745-2021), 2021.
693
694 Zhang, Q., Jiang, X., Tong, D., Davis, S. J., Zhao, H., Geng, G., Feng, T., Zheng, B., Lu, Z.,
695 Streets, D. G., Ni, R., Brauer, M., van Donkelaar, A., Martin, R. V., Huo, H., Liu, Z.,
696 Pan, D., Kan, H., Yan, Y., Lin, J., He, K., and Guan, D.: Transboundary health impacts
697 of transported global air pollution and international trade, *Nature*, 543, 705–709,
698 <https://doi.org/10.1038/nature21712>, 2017.
699
700 Zheng, B., Tong, D., Li, M., Liu, F., Hong, C., Geng, G., Li, H., Li, X., Peng, L., Qi, J., Yan,
701 L., Zhang, Y., Zhao, H., Zheng, Y., He, K., and Zhang, Q.: Trends in China’s
702 anthropogenic emissions since 2010 as the consequence of clean air actions, *Atmos.*
703 *Chem. Phys.*, 18, 14095–14111, [https://doi.org/10.5194/acp-18-14095-](https://doi.org/10.5194/acp-18-14095-2018)
704 [2018](https://doi.org/10.5194/acp-18-14095-2018), 2018.



705

706 **Figure 1.** Tagged source regions (NAM: North America; EUR: Europe; EAS: East Asia; SAS:

707 South Asia; ROW: rest of the world).



708

709

710

711

712

713

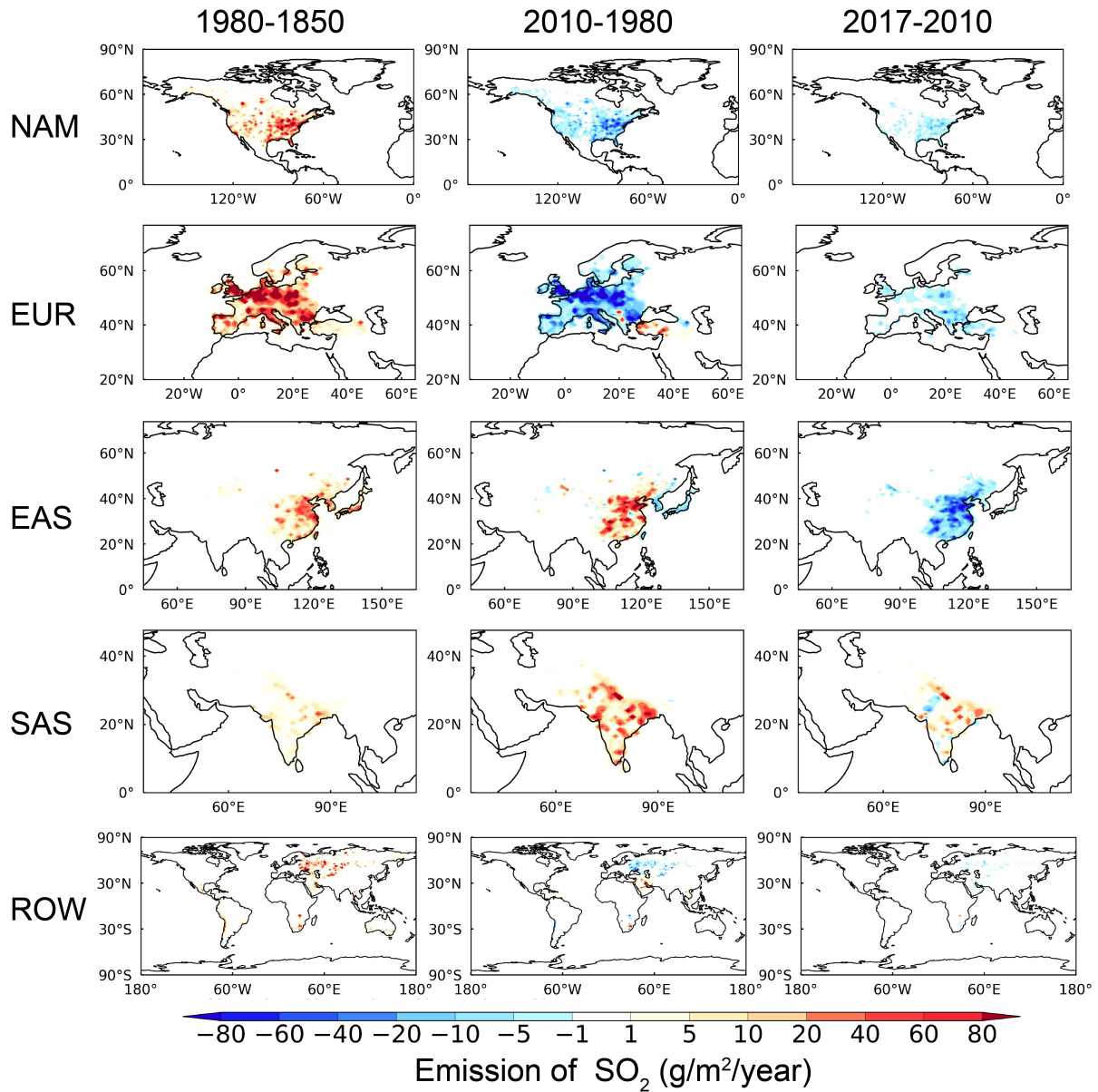
714

715

716

717

Figure 2. Spatial distribution (left panel) and scatter plot (right) between the simulated and observed annual mean near-surface $PM_{2.5}$ concentrations ($\mu g/m^3$) in 2017. Observational data are from IMPROVE (triangle), EMEP (circle), India (inverted triangle) and CNEMC (square). The solid line marks the 1:1 ratio and dashed lines mark the 1:2 and 2:1 ratios. Normalized mean bias (NMB) and correlation coefficient (R) between observation and simulation are shown on the right panel. $NMB = 100\% \times \sum(M_i - O_i) / \sum O_i$, where M_i and O_i are the modeled and observed values at site i , respectively.

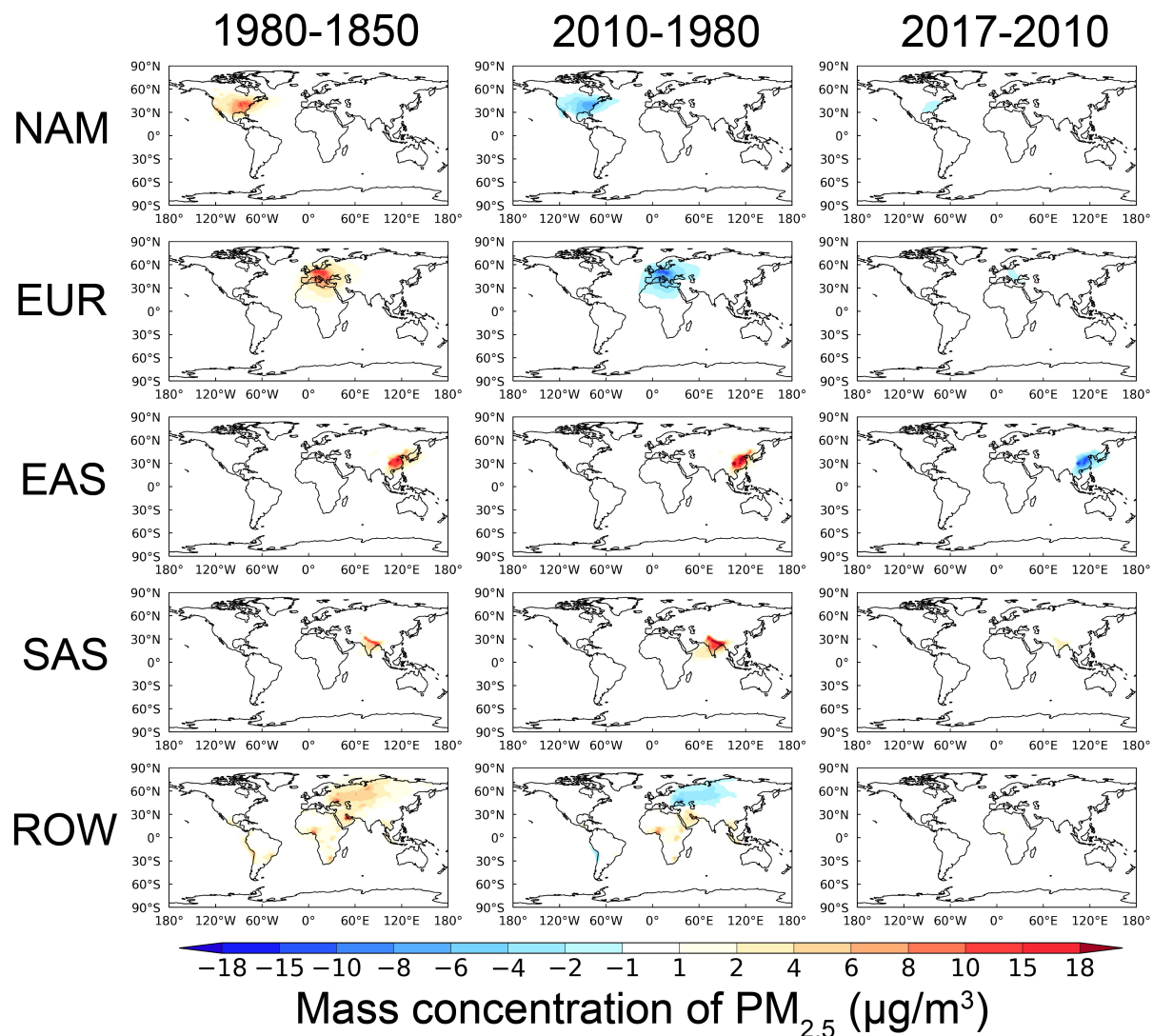


718

719 **Figure 3.** Changes in anthropogenic sulfur dioxide (SO₂) emissions (g/m²/year) between 1850
 720 and 1980 (left), between 1980 and 2010 (middle), and between 2010 and 2017 (right) in the 5
 721 tagged source regions (NAM, EUR, EAS, SAS and ROW from top to bottom).

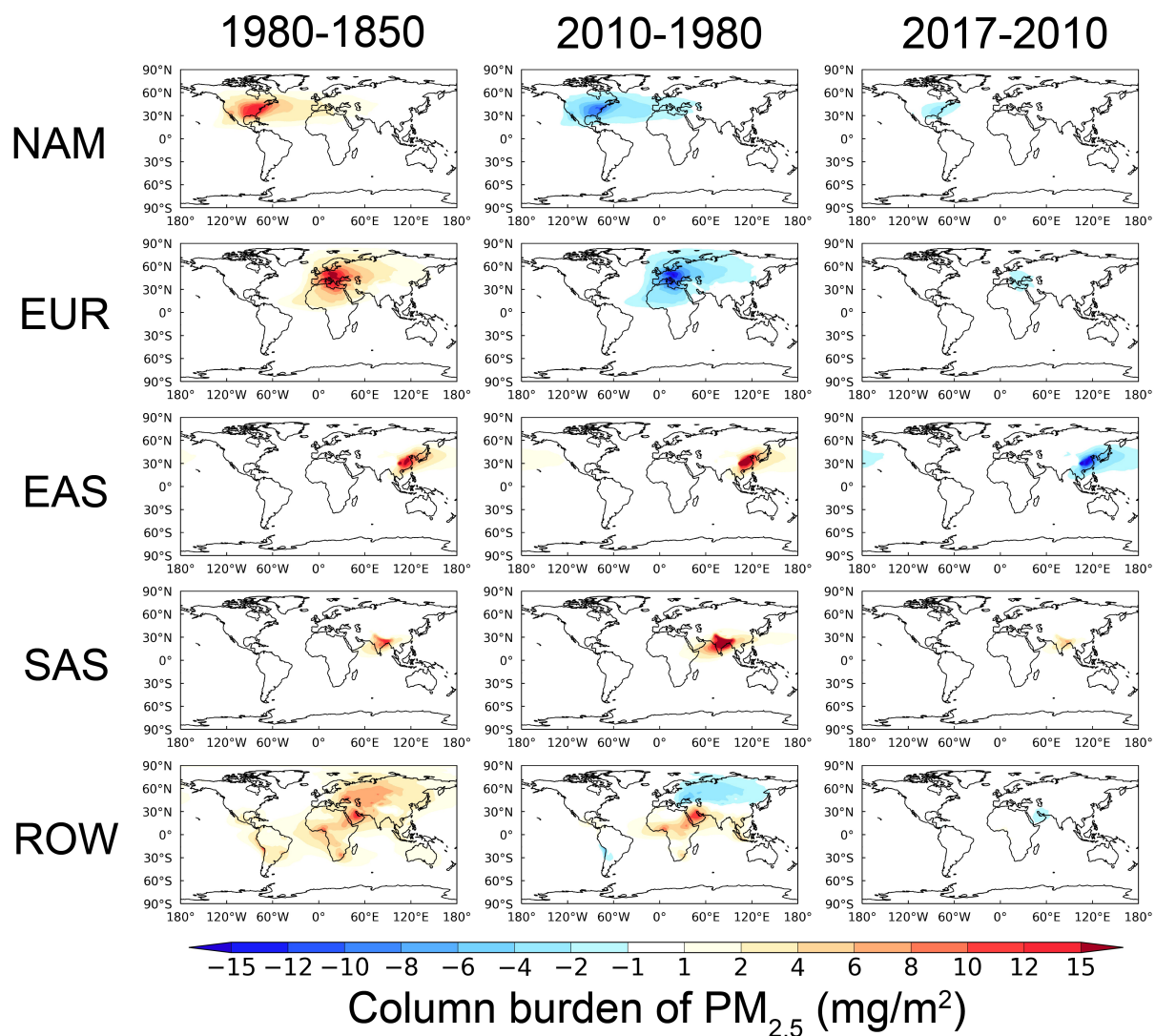
722

723



724

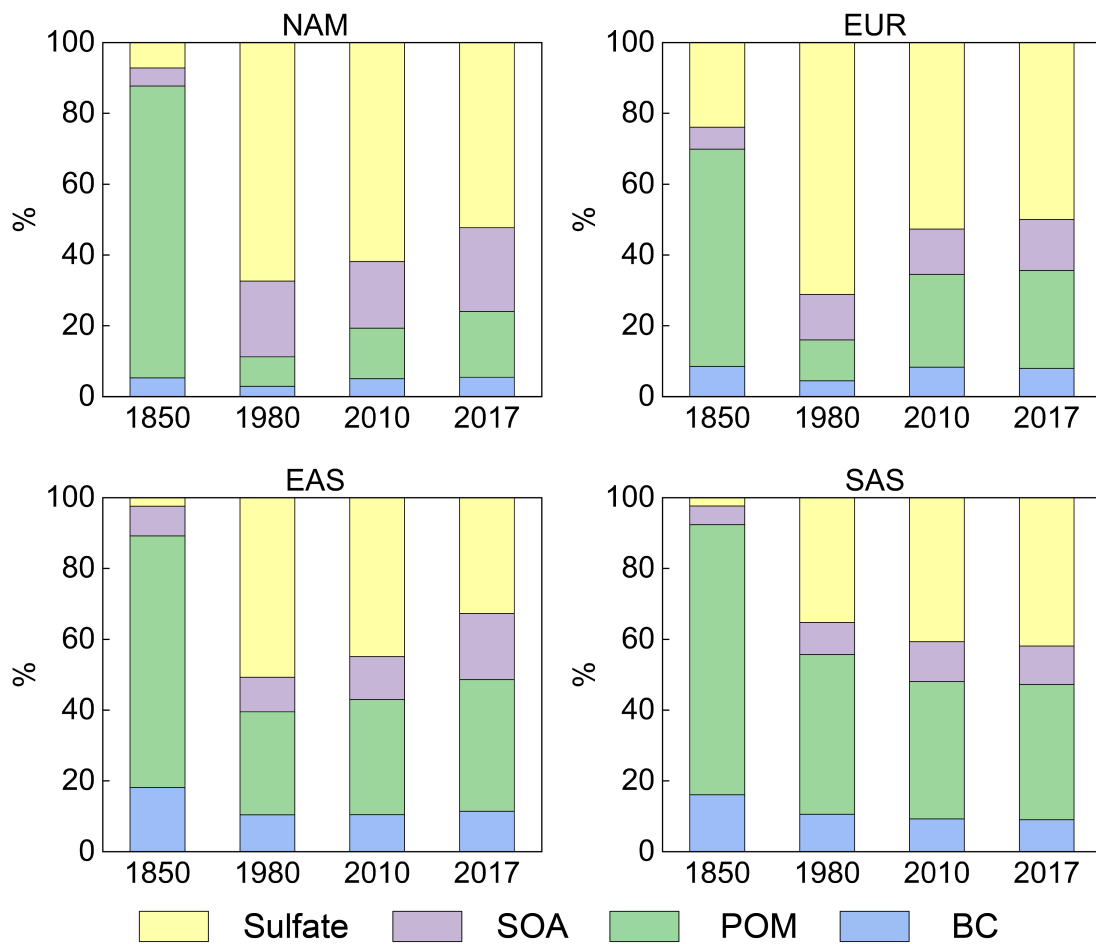
725 **Figure 4.** Changes in near-surface mass concentration ($\mu\text{g}/\text{m}^3$) of anthropogenic $\text{PM}_{2.5}$
 726 contributed by the 5 tagged source regions (NAM, EUR, EAS, SAS and ROW from top to
 727 bottom) between 1850 and 1980 (left), between 1980 and 2010 (middle), and between 1980
 728 and 2017 (right).
 729



730

731 **Figure 5.** Changes in column burden (mg/m²) of anthropogenic PM_{2.5} contributed by the 5
 732 tagged source regions (NAM, EUR, EAS, SAS and ROW from top to bottom) between 1850
 733 and 1980 (left), between 1980 and 2010 (middle), and between 2010 and 2017 (right).

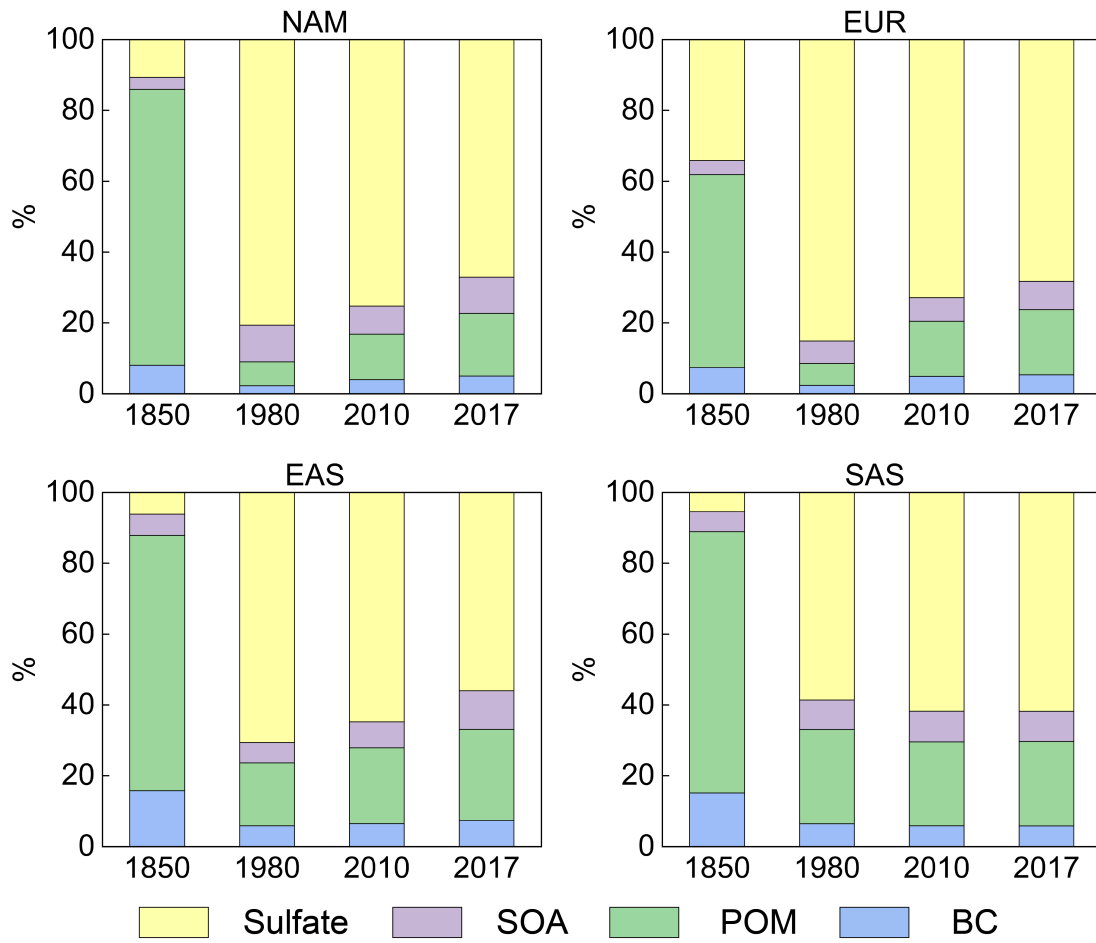
734



735

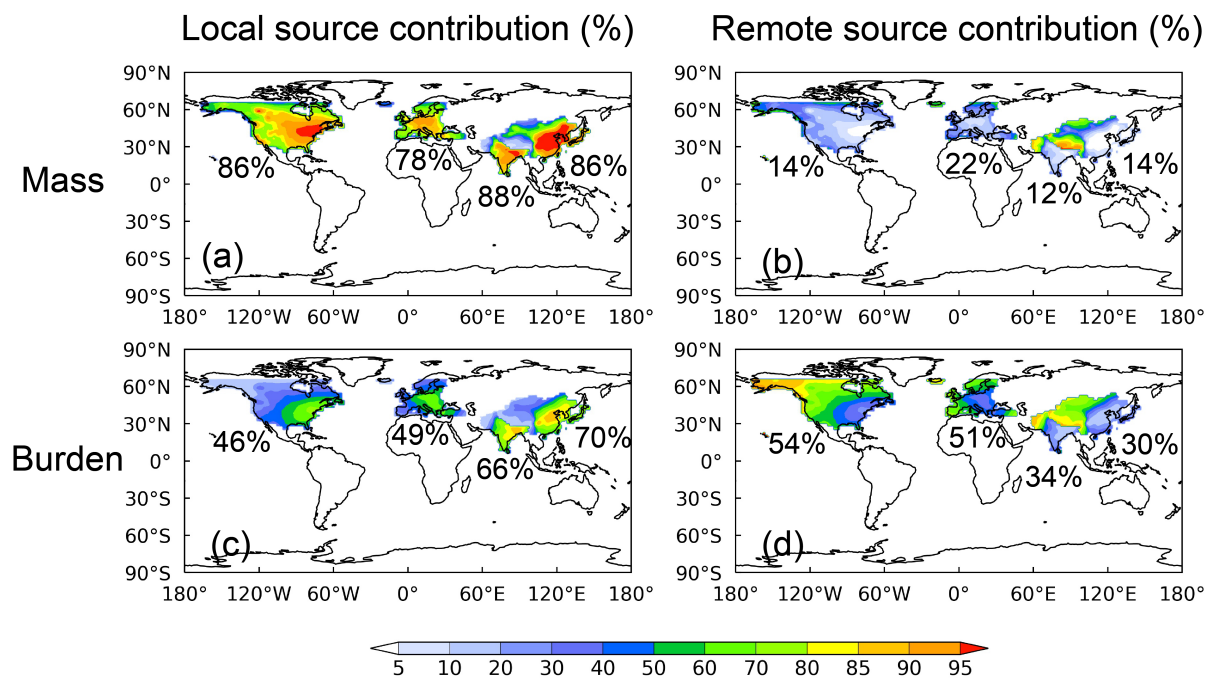
736 **Figure 6.** Percentage contributions (%) of aerosol species including BC, POM, SOA and
 737 sulfate to the near-surface mass concentrations of PM_{2.5} averaged over the four major emission
 738 source regions (NAM, EUR, EAS and SAS) in the focused four years (1850, 1980, 2010 and
 739 2017).

740



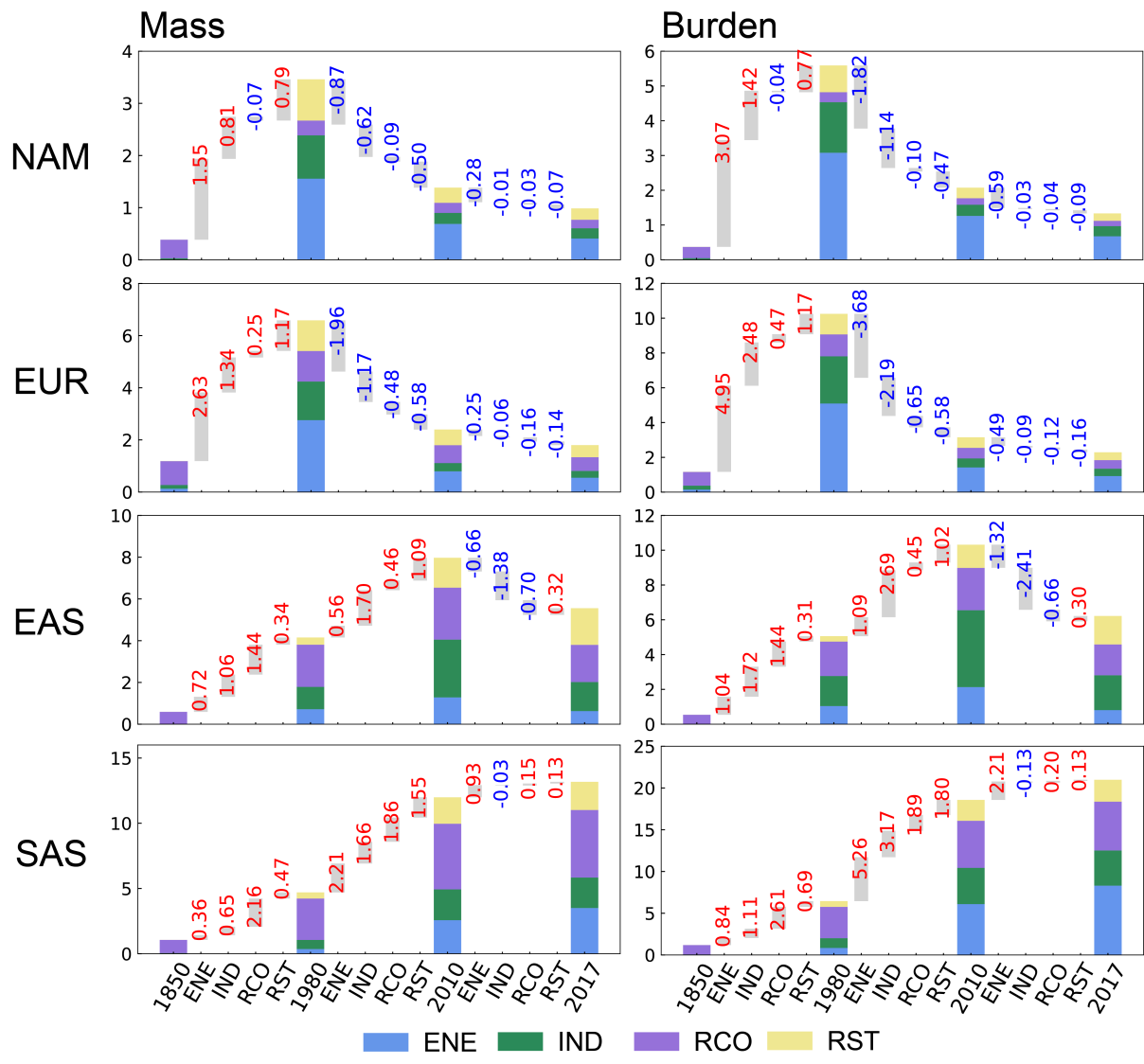
741

742 **Figure 7.** Percentage contributions (%) of aerosol species including BC, POM, SOA and
 743 sulfate to the column burden of PM_{2.5} averaged over the four major emission source regions
 744 (NAM, EUR, EAS and SAS) in the focused four years (1850, 1980, 2010 and 2017).
 745



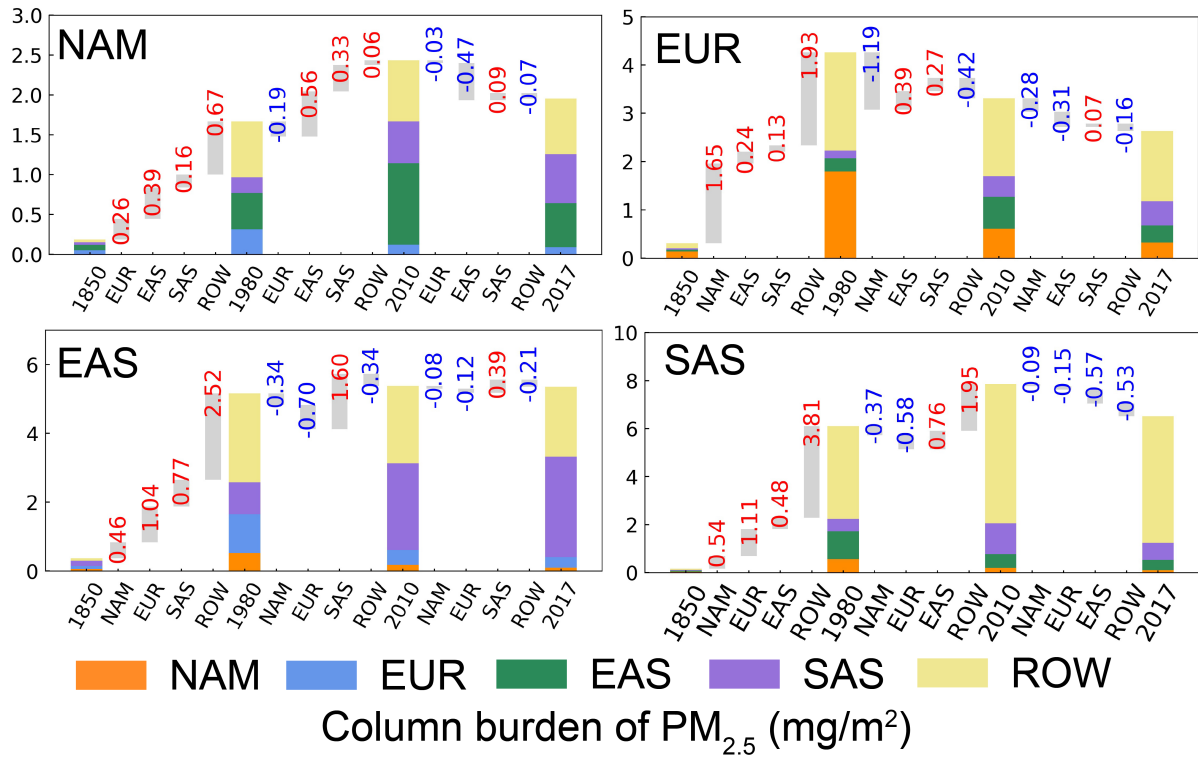
746

747 **Figure 8.** Relative contributions (%) from (a, c) local and (b, d) remote anthropogenic
 748 emissions to the near-surface mass concentrations and column burdens of PM_{2.5} in the four
 749 targeted regions (NAM, EUR, EAS and SAS) in 2010. Numbers marked on the figure are the
 750 regional average over the four individual targeted regions.
 751



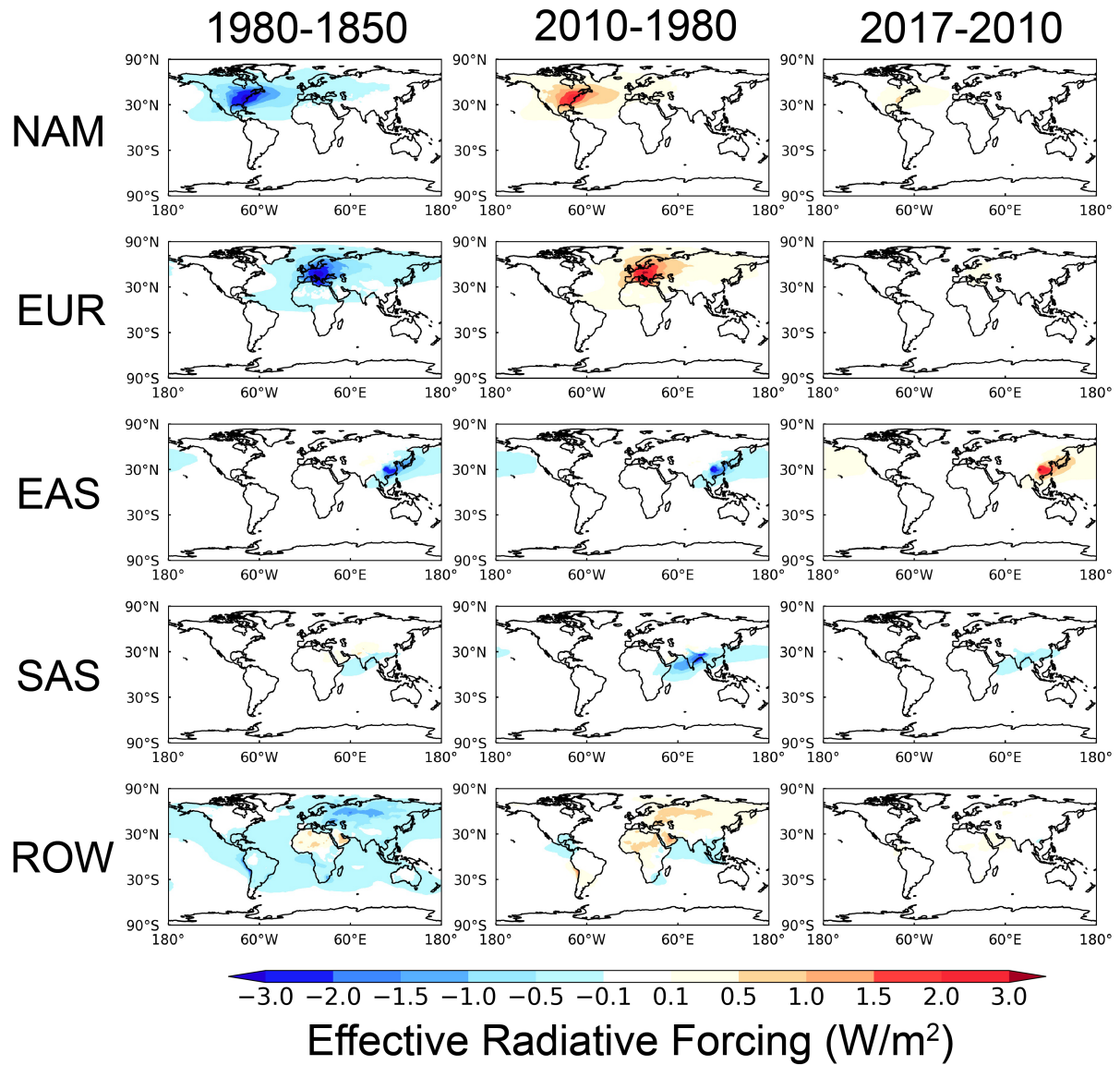
752
 753
 754
 755
 756
 757
 758
 759

Figure 9. Local source contributions from four individual emission sectors (ENE, IND, RCO and RST) to the near-surface mass concentrations ($\mu\text{g}/\text{m}^3$, left) and column burdens (mg/m^2 , right) of anthropogenic $\text{PM}_{2.5}$ in the four targeted regions (NAM, EUR, EAS and SAS from top to bottom) for years 1850, 1980, 2010 and 2017 (in color bars). Grey bar and numbers between two years show the change in sector contributions. Positive values are shown in red and negative values are shown in blue.



760
761
762
763
764

Figure 10. Same as Figure 9, but for contributions from the remote tagged source regions to the column burdens (mg/m²) of anthropogenic PM_{2.5} in the four targeted regions (NAM, EUR, EAS and SAS).



765
 766 **Figure 11.** Changes in effective radiative forcing (W m^{-2}) at the top of the atmosphere due to
 767 aerosol-radiation interactions between 1850 and 1980 (left), between 1980 and 2010 (middle),
 768 and between 2010 and 2017 (right) attributed to the changes in anthropogenic emissions from
 769 NAM, EUR, EAS, SAS and ROW (from top to bottom).