

1    **Measurement report: Characteristics of**  
2    **nitrogen-containing organics in PM<sub>2.5</sub> in Urumqi,**  
3    **northwest China: differential impacts of**  
4    **combustion of fresh and old-age biomass**  
5    **materials**

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25 **Abstract:** Nitrogen-containing organic compounds (NOCs) are abundant and  
26 important aerosol components deeply involved in the global nitrogen cycle. However,  
27 the sources and formation processes of NOCs remain largely unknown, particularly in  
28 the city (Urumqi, China) farthest from the ocean worldwide. Here, NOCs in PM<sub>2.5</sub>  
29 collected in Urumqi over a one-year period were characterized by ultrahigh-resolution  
30 mass spectrometry. The abundance of CHON compounds (mainly oxygen-poor  
31 unsaturated aliphatic-like species) in the positive ion mode was higher in the warm  
32 period than in the cold period, which was largely attributed to the contribution of fresh  
33 biomass material combustion (e.g., forest fires) associated with amidation of  
34 unsaturated fatty acids in the warm period, rather than the oxidation processes. However,  
35 CHON compounds (mainly nitro-aromatic species) in the negative ion mode increased  
36 significantly in the cold period, which was tightly related to old-age biomass  
37 combustion (e.g., dry straws) in wintertime Urumqi. For CHN compounds, alkyl nitriles  
38 and aromatic species showed higher abundance in the warm and cold periods,  
39 respectively. Alkyl nitriles can be derived from fresh biomass material combustion  
40 associated with the dehydration of amides (the main CHON compounds in the warm  
41 period). In contrast, aromatic species were tightly related to old-age biomass burning.  
42 These findings further suggested different impacts of the combustion of fresh- and old-  
43 age biomass materials on NOC compositions in different seasons. The overall results  
44 shed light on the mechanisms by which fresh and old-age biomass materials release  
45 different NOCs during combustion.

46 **Keywords:** Aerosols, Organic nitrogen, Molecular composition, Fresh biomass, Old-

47 age biomass

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49 **1. Introduction**

50 Fine particulate matter (PM<sub>2.5</sub>) is a typical atmospheric pollutant that can affect the  
51 global climate system, as well as urban air quality and human health (Seinfeld et al.,  
52 2016; Wang et al., 2021a). Organic aerosol (OA) contributes significantly (20–90%) to  
53 PM<sub>2.5</sub> mass concentration in most polluted areas worldwide (Zhang et al., 2007; Han et  
54 al., 2023). Up to 77% of molecules in OA include nitrogen-containing functional groups  
55 (Ditto et al., 2020; Kenagy et al., 2021), which have been suggested to play important  
56 roles in the formation, transformation, acidity, and hygroscopicity of OA (Xu et al.,  
57 2020; Wang et al., 2017b; Laskin et al., 2009). Moreover, the further oxidation or  
58 nitrification of some nitrogen-containing organic compounds (NOCs) and volatile  
59 organic compounds (VOCs) by ozone (O<sub>3</sub>), hydroxyl radical (•OH), and nitrogen oxides  
60 (NO<sub>x</sub>) can lead to an increase in the health hazards of OA (Franze et al., 2005; Bandowe  
61 and Meusel, 2017). Nitrated amino acids and nitrated PAHs are two representative  
62 hazard NOCs (Franze et al., 2005; Bandowe and Meusel, 2017). Thus, the identification  
63 of aerosol NOCs at the molecular level is important for improving our understanding  
64 of the precursors, sources, and formation processes of nitrogen-containing OA.

65 Previous observations in urban, rural, marine, and forest areas have suggested that  
66 the molecular composition and relative abundance of aerosol NOCs were spatially  
67 different (Samy and Hays, 2013; Jiang et al., 2022; Lin et al., 2012; Xu et al., 2023;  
68 Zeng et al., 2021; Zhang et al., 2022; Zeng et al., 2020). These differences can be mainly

69 attributed to the diverse sources and formation mechanisms of aerosol NOCs.

70 Commonly reported primary sources include combustion process releases and natural

71 emissions (e.g., soils, plant debris, pollen, and ocean) (Song et al., 2022; Wang et al.,

72 2017b; Cape et al., 2011; Lin et al., 2023). In addition, aerosol NOCs can also be tightly

73 associated with secondary formation processes involving the reactions of reactive

74 nitrogen with VOCs or particle-phase CHO compounds (Bandowe and Meusel, 2017;

75 Zarzana et al., 2012; Laskin et al., 2014). For example, laboratory experiments have

76 found that the oxidation of isoprene and  $\alpha$ -/ $\beta$ -pinene in the presence of  $\text{NO}_x$  can result

77 in the formation of organic nitrates (e.g., methacryloyl peroxy nitrate, dihydroxynitrates,

78 and monohydroxynitrates) (Surratt et al., 2010; Rollins et al., 2012; Nguyen et al., 2015).

79 The reduced nitrogen species (e.g.,  $\text{NH}_3$ ,  $\text{NH}_4^+$ , and organic amines) have been

80 demonstrated to contribute to the formation of NOCs through "carbonyl-to-imine"

81 transformations in the laboratory experiments (Zarzana et al., 2012; Laskin et al., 2014).

82 In the field observation studies, NOCs in particulate matter were analyzed at the

83 molecular level to indicate their sources and formation mechanisms (Jiang et al., 2022;

84 Lin et al., 2012; Zhong et al., 2023). Xu et al. (2023) characterized the variations of

85 molecular compositions in urban road  $\text{PM}_{2.5}$ , suggesting that organic nitrates increased

86 largely through the interactions of atmospheric oxidants, reactive gas-phase organics,

87 and aerosol liquid water. Several field studies conducted in Beijing (China) and

88 Guangzhou (China) also suggested that the molecular compositions and formation of

89 NOCs were tightly associated with environmental conditions (Jiang et al., 2022; Lin et

90 al., 2012; Xie et al., 2020). Generally, most studies on aerosol NOCs were performed

91 in economically developed regions, as well as in forest and marine areas (Jiang et al.,  
92 2022; Wang et al., 2017a; Ditto et al., 2022b; Altieri et al., 2016; Xu et al., 2020; Liu et  
93 al., 2023; Zhang et al., 2022; Zeng et al., 2020). In contrast, few studies have  
94 investigated the sources and atmospheric transformation of NOCs in the northwest  
95 border urban of China (e.g., Urumqi) with fragile ecology and harsh environmental  
96 conditions (e.g., cold winter and dry summer), which may hinder our comprehensive  
97 and in-depth understanding of the formation process of NOCs in ambient aerosols.

98 Biomass burning emissions were widely reported in the source identification of  
99 aerosol NOCs in northern and southwestern China because of heating and cooking  
100 needs (Zhong et al., 2023; Wang et al., 2021c; Chen et al., 2017). A recent observation  
101 study in urban Tianjin suggested that most CHON compounds in wintertime PM<sub>2.5</sub>  
102 originated from biomass burning (Zhong et al., 2023). The CHN<sub>2</sub> compounds have been  
103 identified in biomass burning OA (BBOA) (Laskin et al., 2009; Wang et al., 2017b).  
104 Moreover, the high temperature generated by biomass burning can facilitate the release  
105 of ammonia, a process that caused the reaction of carboxylic acids (e.g., oleic acid) with  
106 ammonia to form amides and alkyl nitriles (Radzi Bin Abas et al., 2004; Simoneit et al.,  
107 2003). Interestingly, we found that biomass burning in rural China typically includes  
108 fresh biomass materials (e.g., forest fires) and old-age biomass materials (e.g., straw  
109 after autumn harvest, fallen leaf, and deadwood). Fresh biomass is rich in oils and  
110 proteins, whereas old-age biomass materials are usually oligotrophic due to the transfer  
111 of nutrients to tender tissues or fruits (Jian et al., 2016; Xu and Xiao, 2017). Thus, NOCs  
112 released from different types of biomass combustion may vary in molecular

113 compositions. However, there are large gaps in our current knowledge about the  
114 impacts of fresh and old-age biomass burning on NOCs in ambient aerosols.

115 Urumqi (northwest China) is the largest inland city farthest from the ocean in the  
116 world, which is becoming increasingly prominent due to the national strategy of the  
117 "One Belt, One Road." The city and neighboring countries have a dry summer that can  
118 easily trigger forest fires (Bátori et al., 2018; Xu et al., 2021), while the winter is  
119 freezing with intensive old-age biomass and fuel combustion for heating (Ren et al.,  
120 2017). In this study, we presented one-year ambient measurements of the chemical  
121 compositions in PM<sub>2.5</sub> collected from Urumqi. The specific aims of this study are (1) to  
122 investigate the molecular-level speciation of functionalized organic nitrogen  
123 compounds via high-resolution mass spectrometry with positive (ESI+) and negative  
124 (ESI-) ionizations and (2) to investigate the potential sources and formation processes  
125 for NOCs with a special focus on the relative influences of fresh and old-age biomass  
126 burning in different seasons.

127

## 128 **2. Materials and methods**

### 129 **2.1. Study site description and sample collection**

130 The study was conducted in Urumqi city, which has an average altitude of 800 m.  
131 The region has an arid temperate continental climate with an annual mean temperature  
132 of  $7.4 \pm 13.9$  °C and an annual mean rainfall of 27.8 mm. The sampling site is located  
133 in the suburban area (Boda campus of Xinjiang University) of the city (87.75°E,  
134 43.86°N) (**Figure S1**), which is characterized by low population and traffic density.

135 This is because Urumqi is relatively vast and sparsely populated compared to developed  
136 coastal cities in China (Qizhi et al., 2016). Additionally, the area is surrounded by  
137 mountains on three sides, resulting in the difficulty in diffusing air pollutants. The  
138 dominant forest trees in this area are *Picea schrenkiana*, *Betula tianschanica* Rupr.,  
139 *Populus talassica* Kom., and *Ulmus pumila* L.. The dry climate and strong sunlight in  
140 the warm period ( $18.81 \pm 6.4^{\circ}\text{C}$ , **Table S1**) would be the main culprits of forest fires in  
141 the local and nearby areas. In the cold period ( $-1.96 \pm 11.26^{\circ}\text{C}$ ) (**Table S1**), the  
142 centralized heating and old-age biomass burning may be the main contributors to local  
143 air pollution. Thus, it provides an unexpected opportunity to investigate the potentially  
144 differential impacts of fresh and old-age biomass burning on aerosol NOCs.

145 A high-volume air sampler (Series 2031, Laoying, China) was set up on the  
146 rooftop of a building (School of Geology and Mining Engineering, Xinjiang University).  
147 PM<sub>2.5</sub> samples ( $n = 73$ ) were collected every five days with a duration of  $\sim 24$  h onto  
148 prebaked ( $450^{\circ}\text{C}$  for  $\sim 10$  h) quartz fiber filters (Pallflex, Pall Corporation, USA) from  
149 1 March 2018 to 26 February 2019. One blank filter was collected every month ( $n =$   
150 12). All filter samples were stored at  $-30^{\circ}\text{C}$  until further analysis. During the sampling  
151 campaigns, the meteorological data (e.g., temperature and relative humidity) and the  
152 concentrations of O<sub>3</sub> and NO<sub>x</sub> were recorded hourly from the adjacent environmental  
153 monitoring station. These hourly data were then averaged to obtain daily values to  
154 match the sampling time of PM<sub>2.5</sub>. In addition, the trajectories (72 h) of air masses  
155 arriving at the sampling site at each sampling event were calculated to investigate the  
156 potential influence of pollutant transport on aerosol NOCs.

157

158 **2.2. Chemical analysis**

159 A portion of each filter sample was extracted twice using 3 mL methanol (LC-MS  
160 grade, CNW Technologies Ltd.) under sonication in a chilled ice slurry (~4 °C). The  
161 extracted solutions were filtered through a polytetrafluoroethylene syringe filter (0.22  
162 µm, CNW Technologies GmbH). Subsequently, the extracts were concentrated to 300  
163 µL with a gentle stream of gaseous nitrogen (Shanghai Likang Gas Co., Ltd). The final  
164 extracts were analyzed using an ultra-performance liquid chromatography quadrupole  
165 time-of-flight mass spectrometry equipped with an electrospray ionization (ESI) source  
166 (UPLC-ESI-QToFMS, Waters Acquity Xevo G2-XS) in both ESI+ and ESI- modes  
167 (Wang et al., 2021b). It should be pointed out that UPLC-ESI-MS (i.e., TOF-only) was  
168 used to identify molecular formulas of organic matter, while the functional groups of  
169 the target molecule formulas were deciphered by UPLC-ESI-MS/MS (i.e., tandem mass  
170 spectrometry). Ions obtained from *m/z* 50–700 were assigned molecule formulas by  
171 assuming hydrogen or sodium adducts in ESI+ mode and deprotonation in ESI- mode.  
172 Detailed chromatographic conditions, parameter selection, and quality control were  
173 displayed in the Supplement (Sect. S1). Notably, there may be differences in ionization  
174 efficiencies between compound types. However, the exact impacts of ionization  
175 efficiency on multifunctional compounds in a complex mixture are uncertain and  
176 difficult to evaluate (Ditto et al., 2022b; Yang et al., 2023). Thus, the intercomparison  
177 across compound relative abundance without considering potentially differentiated  
178 ionization efficiency was conducted in this study, which was similar to many previous

179 studies (Xu et al., 2023; Jiang et al., 2022).

180 For the measurement of inorganic ions, a portion of each filter sample was  
181 ultrasonically extracted with Milli-Q water (18 MΩ cm) (3 mL) in an ice-water bath  
182 (~4 °C). The extract solutions were then filtered via a polytetrafluoroethylene syringe  
183 filter (0.22 μm, Millipore, Billerica, MA). The concentrations of water-soluble  
184 inorganic ions, including NO<sub>3</sub><sup>−</sup>, SO<sub>4</sub><sup>2−</sup>, Cl<sup>−</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, and NH<sub>4</sub><sup>+</sup> in the samples  
185 were determined using an ion chromatograph system (Dionex Aquion, Thermo  
186 Scientific, USA) (Xu et al., 2022a; Lin et al., 2023).

187

### 188 **2.3. Compound categorization and predictions of ALW, pH, and hydroxyl radical**

189 The molecular formulas identified by UPLC-ESI-QToFMS were classified into  
190 several major compound classes based on their elemental compositions (i.e., C, H, O,  
191 and N), primarily including CHO, CHON, and CHN groups in the ESI+ mode and CHO,  
192 CHON, CHOS and CHONS groups in the ESI− mode (Wang et al., 2017b). CHOS and  
193 CHONS compounds were also detected in the ESI− mode, with numbers of 398 and  
194 112, respectively (**Table S2**). As this study focused mainly on NOCs, sulfur-containing  
195 species were not discussed. Unless stated otherwise, all of the detected molecules were  
196 reported as neutral molecules. The double-bond equivalent (DBE) and carbon oxidation  
197 state (OS<sub>C</sub>) were calculated to reflect the unsaturation degree of the organics and the  
198 composition evolution of organics that underwent oxidation processes, respectively  
199 (details in **Sect. S2**) (Kroll et al., 2011; Xu et al., 2023). The identified compounds can  
200 be further classified into four subgroups based on the number of carbon atoms and OS<sub>C</sub>

201 value (Kroll et al., 2011; Xu et al., 2023). Briefly, semi-volatile oxidized organic aerosol  
202 (SV-OOA) and low-volatility oxidized organic aerosol (LV-OOA) were associated with  
203 multi-step oxidation reactions, with  $OS_C$  values between  $-1$  and  $+1$  and molecular  
204 formulas less than 13 carbon atoms. BBOA has  $OS_C$  values ranging from  $-0.5$  to  $-1.5$   
205 and more than seven carbon atoms. Compounds with  $OS_C$  values less than  $-1$  and  
206 carbon atoms above 20 may be related to hydrocarbon-like organic aerosol (HOA).  
207 Additionally, the modified aromaticity index ( $AI_{mod}$ ) was also calculated to indicate the  
208 aromaticity of organic compounds (details in **Sect. S2**) (Koch and Dittmar, 2006). The  
209 van Krevelen diagrams and  $AI_{mod}$  values have been proposed to further classify organic  
210 matter categories (Xu et al., 2023; Su et al., 2021), according to which the identified  
211 five subgroups included saturated-like molecules (Sa,  $H/C \geq 2.0$ ), unsaturated aliphatic-  
212 like molecules (UA,  $1.5 \leq H/C < 2.0$ ), highly unsaturated-like molecules (HU,  $AI_{mod} \leq$   
213 0.5 and  $H/C < 1.5$ ), highly aromatic-like molecules (HA,  $0.5 < AI_{mod} \leq 0.66$ ), and  
214 polycyclic aromatic-like molecules (PA,  $AI_{mod} > 0.66$ ). Furthermore, it has been  
215 suggested that the above subgroups can be subdivided into O-poor and O-rich  
216 compounds depending on their O/C ratio (**Table S8**) (Merder et al., 2020; Zhong et al.,  
217 2023).

218 A thermodynamic model (ISORROPIA-II) was applied to predict the mass  
219 concentration of aerosol liquid water (ALW) and the value of pH with particle-phase  
220 ion concentrations as well as ambient temperature and relative humidity as the inputs,  
221 as detailed in our previous publications (Xu et al., 2020; Xu et al., 2023; Xu et al.,  
222 2022b). The model output results based on our data set showed that 94% and 90% of

223  $\text{NO}_3^-$  were in the aerosol phase in the cold and warm periods, respectively. Hence, the  
224 predictions of pH and ALW were conducted without considering gaseous nitric acid  
225 (Guo et al., 2015; Wang et al., 2021c). 78% and 21% of  $\text{NH}_4^+$  were in the aerosol phase  
226 in the cold and warm periods, respectively. Moreover, it is important to note that  
227 gaseous  $\text{NH}_3$  measurements were not conducted and ammonia partitioning was not  
228 considered in this study. Thus, a bias correction of 1 pH unit was applied to calculate  
229 the aerosol pH values (Guo et al., 2015; Wang et al., 2021c). The concentrations of  
230 ambient  $\cdot\text{OH}$  were predicted using empirical formula (Ehhalt and Rohrer, 2000; Wang  
231 et al., 2020).

232

### 233 **3. Results and discussion**

#### 234 **3.1. Overall molecular characterization of organic aerosols**

235 **Figures 1a and 1c** show the mass spectra of organic compounds detected in ESI+  
236 and ESI-, respectively. More compounds were identified in ESI+ (1885 molecular  
237 formulas) than in ESI- (1091 molecular formulas) (**Table S2**), which was similar to  
238 previous reports about the molecular characteristic of biomass burning aerosols and  
239 urban aerosols (Jiang et al., 2022; Wang et al., 2017b). The molecular weights of the  
240 compounds with relatively high signal intensity mainly ranged from 100 Da to 500 Da  
241 in ESI+, which was larger than those (100–300 Da) observed in the urban (Changchun,  
242 Guangzhou, and Shanghai) (Wang et al., 2021a) and agriculture (Suixi) (Wang et al.,  
243 2017b) regions of China. In contrast, the species with the strong signal intensity fell  
244 between 100 Da and 300 Da in ESI-. This mass range detected in Urumqi organic

aerosols was comparable to previous observations in urban (Xi'an) aerosols (Han et al., 2023) but significantly lower than that in firework-related urban (Beijing) aerosols (300–400 Da) (Xie et al., 2020). On average, the molecular number and relative abundance of CHON compounds (150–500 Da) were dominant in ESI+, accounting for 45.57% of the total molecular number and  $62.70 \pm 6.83\%$  of the total signal intensity (**Figures 1a** and **Table S2**). CHO compounds were the second most abundant categories ( $28.76 \pm 4.75\%$  of the total signal intensity), followed by CHN compounds. However, previous observations conducted in Shanghai, Guangzhou, and Changchun suggested that the compounds in ESI+ were dominated by CHN and CHON species (Wang et al., 2021a). In ESI–, although the number of CHON compounds was less than CHO, the relative abundance of CHON compounds (150–250 Da) was higher (**Figures 1d** and **Table S2**). The finding was consistent with the results obtained in Shanghai and Changchun but different from the case in Guangzhou (Wang et al., 2021a). The average H/C ratios of CHO (1.62–1.66) and CHON (1.79–1.83) compounds in ESI+ mode (**Table S3**) were higher than those (0.94–1.13 for CHO and 1.27–1.47 for CHON) in Changchun, Shanghai, and Guangzhou (Wang et al., 2021a). However, the average O/C ratios of CHO (0.25–0.3) and CHON (0.22–0.3) compounds in ESI+ mode (**Table S3**) were less than those (0.42–0.43 for CHO and 0.27–0.45 for CHON) in the urban areas (Shanghai and Guangzhou) (Wang et al., 2021a). Overall, these dissimilarities in molecular characteristics of organic aerosols between Urumqi and other areas may be attributed to their different sources and formation mechanisms.

267 **Figures 1b and 1d** show the time series of the fractional distributions of various  
268 organic matter categories in different ion modes. The abundance of CHO compounds  
269 in ESI+ exhibited a temporal variation similar to that of CHON compounds ( $r = 0.51$ ,  
270  $P < 0.01$ ), with increased levels in the warm period. This indicated that CHO  
271 compounds may be important precursors for the formation of NOCs (via reactions in  
272 the gas- and/or particle-phases) or that they have similar origins. Previous simulation  
273 experiments have demonstrated that higher temperatures increase the concentration of  
274 oxygenated organic molecules, while lower temperatures can allow less oxidized  
275 species to condense (Stolzenburg et al., 2018; Frege et al., 2018). In addition, solar  
276 radiation and atmospheric oxidation capacity are also important factors promoting the  
277 formation of more oxygenated organic molecules (Li et al., 2022; Liu et al., 2022). Air  
278 temperature, radiation, and atmospheric oxidation capacity were much higher in the  
279 warm period than in the cold period in Urumqi (**Table S1**) (Wan et al., 2021), which  
280 may be partly responsible for increased abundances of CHO and CHON compounds in  
281 the warm period. However, the abundance of CHN compounds tended to increase from  
282 the warm period to the cold period. Since the ESI+ mode is highly sensitive to  
283 protonatable species, organic amines were expected to predominate the CHN  
284 compounds (Han et al., 2023; Wang et al., 2021a). It is well documented that the  
285 formation of amine salt in the particle phase is tightly associated with aerosol acidity  
286 and water (Liu et al., 2023). Thus, the reduced pH value and increased ALW level in  
287 the cold period (**Table S1**) provided greater potential for converting gaseous amines  
288 into particles.

289 In ESI- mode, the abundances of CHON and CHO compounds exhibited a  
290 significantly increased level in the cold period (**Figure 1d**), a variation pattern which  
291 was completely opposite to the case in ESI+ mode. The ESI- mode is more sensitive  
292 to deprotonatable compounds like nitrophenols, organic nitrates, organosulfates, and  
293 organic acids (Jiang et al., 2022; Lin et al., 2012). The formations of these compounds  
294 were highly impacted by ALW and aerosol acidity (Ma et al., 2021; Smith et al., 2014;  
295 Zhou et al., 2023; Xu et al., 2023). However, Urumqi has dry and dusty weather,  
296 particularly in warm period, resulting in a quite low ALW concentration ( $1.86 \pm 1.90$   
297  $\mu\text{g m}^{-3}$ ) in the warm period (**Table S1**). Moreover, the calculated mean pH value was  
298  $6.86 \pm 1.71$  (**Table S1**) during the warm period, which implies that the fine aerosol  
299 particles in the warm period in Urumqi was neutral or slightly alkaline. Obviously, the  
300 aerosol characteristics of the warm period in Urumqi may hinder the formation of these  
301 organic compounds measured in ESI- mode. In contrast, the increased ALW  
302 concentration and decreased pH value during the cold period can facilitate the  
303 formation of CHO and CHON compounds through the partitioning of gas-phase species  
304 to the particles and subsequent aqueous phase reactions (Xu et al., 2020; Xu et al., 2023).  
305 Furthermore, the total signal intensity of CHO compounds was significantly correlated  
306 with that of CHON ( $r = 0.62, P < 0.01$ ), indicating that they may have similar origins  
307 or that CHO compounds may serve as important precursors for CHON compound  
308 formation. In general, the differentiated seasonal variation patterns for the different  
309 types of NOCs measured here can be attributed to the unique meteorological conditions  
310 in Urumqi and different ionization mechanisms in ESI+ and ESI- modes. The sources

311 and formation mechanisms of NOCs will be further discussed in the following sections.

312

313 **3.2. Seasonally differential sources and formation mechanisms of CHON**  
314 **compounds**

315 CHON compounds can be derived from the reactions between CHO species and  
316 reactive nitrogen species ( $\text{NO}_x$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$ ) (Lee et al., 2016; De Haan et al., 2017),  
317 as also partly implied by significant positive correlations ( $r = 0.51\text{--}0.62$ ,  $P < 0.01$ )  
318 between total signal intensity of CHO and CHON compounds in both ESI+ and ESI-  
319 modes. Thus, CHO compounds were further classified based on their  $\text{OS}_C$  values to  
320 preliminarily explore their origins and linkages with CHON compound formation  
321 (**Figures 2a and 2b**). In ESI+ mode, the  $\text{OS}_C$  values of the detected CHO compounds  
322 ( $-1.75$  to  $0.5$ ) were higher than those of primary vehicle exhausts ( $-2.0$  to  $-1.9$ ) (Aiken  
323 et al., 2008), likely indicating a weak (or indirect) contribution of primary vehicle  
324 exhausts to CHO molecules in Urumqi. The signal intensity of BBOA dominated the  
325 total OA signal intensity and was higher in the warm period than in the cold period  
326 (**Figure 2e**). However, previous studies conducted in China (e.g., Beijing, Xi'an,  
327 Shanghai, and Liaocheng) suggested that biomass burning was more significant in the  
328 cold seasons (Li et al., 2023; Wang et al., 2017a; Chen et al., 2017; Wang et al., 2009;  
329 Wang et al., 2018; Zhang et al., 2022). Furthermore, we found that the oxygen-poor  
330 unsaturated aliphatic compounds showed a high signal intensity in the warm period and  
331 that the signal intensities of all categories of compounds in the warm period were  
332 weakly correlated with atmospheric oxidants (i.e.,  $\text{O}_3$  and  $\cdot\text{OH}$ ) ( $r < 0.1$ ,  $P > 0.05$ ).

333 Thus, the formation or source of CHO compounds in the warm period may not be  
334 mainly controlled by high atmospheric oxidation but rather by biomass burning, which  
335 was distinguished from previous reports (Duan et al., 2020; Kondo et al., 2007; Zhang  
336 et al., 2023). This consideration was also supported by the fact that there were  
337 significantly more fire spots in the warm period than in the cold period (**Figure 3**). It  
338 should be noted that the materials used for biomass burning in the cold period in rural  
339 China are typically old-age plant tissues, such as dead branches of pine trees, dead  
340 branches of shrubs, corn straw, and rice straw (**Figure S3**), while biomass burning in  
341 the warm season is mainly attributed to forest fires or wildfires (relatively fresh  
342 biomass). Accordingly, a large number of fresh biomass material burning occurred from  
343 April to October each year in the neighboring countries (e.g., Kazakhstan) (Xu et al.,  
344 2021) or regions of Urumqi (due to drought) (**Figure 3**) may be largely responsible for  
345 high CHO compound abundance in the warm period.

346 The CHO species in ESI- had higher OS<sub>C</sub> (-1.85 to 1.1) than those in ESI+ (-1.85  
347 to 0.25) (**Figures 2c and 2d**), which was consistent with a recent study conducted in  
348 Guangzhou, China (Zou et al., 2023). The predominant subgroups of CHO in ESI-  
349 were BBOA (66.4% of total signal intensity) and SV-OOA (23.1% of total signal  
350 intensity), which was different from the observation in Shanghai (dominated by SV-  
351 OOA and LV-OOA) (Wang et al., 2017a). Additionally, some specific saturated and  
352 unsaturated aliphatic CHO substances (i.e., C<sub>12-18</sub>H<sub>n</sub>O<sub>2</sub>) in ESI- showed higher  
353 abundance in the warm season than in the cold season, which was contrary to the  
354 variation pattern of other CHO compounds. These C<sub>12-18</sub>H<sub>n</sub>O<sub>2</sub> compounds were found

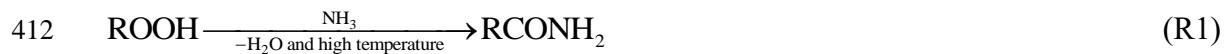
355 to be mainly fatty acids, such as stearic acid ( $C_{18}H_{36}O_2$ ), oleic acid ( $C_{18}H_{34}O_2$ ),  
356 linolelaidic acid ( $C_{18}H_{32}O_2$ ), palmitic acid ( $C_{16}H_{32}O_2$ ), and palmitoleic acid ( $C_{16}H_{30}O_2$ )  
357 (**Figure S4**), all of which usually accumulate in plants, particularly *Suaeda*  
358 *aralocaspica* (W. Hogg and T. Gillan, 1984; Wang et al., 2011). Interestingly, this plant  
359 was widely distributed in Central Asia as well as on the southern edge of the Junggar  
360 Basin in Xinjiang, China (Wang et al., 2011). Although fatty acids can also originate  
361 from food cooking (Zhao et al., 2007), there seem to be no seasonal differences in  
362 cooking behavior locally. Thus, these results further confirmed our consideration that  
363 the abundance of CHO compounds in the warm period was highly impacted by fresh  
364 biomass material burning (e.g., forest fires or wildfires).

365 CHON molecules in ESI+ were mainly identified as unsaturated aliphatic-like  
366 compounds that are oxygen poor (**Figures 4a and 4b**), accounting for more than 70%  
367 of the total signal intensities of CHON species (**Figure S5**). The signal intensity of  
368 CHON species in ESI+ was greater in the warm period than in the cold period (**Figure**  
369 **4e**). Moreover, BBOA contributed to 56.9 % of the total CHON signal intensity in the  
370 warm period (**Figure S6**). These characteristics of CHON compounds were similar to  
371 those of CHO. Considering a significant positive correlation ( $r = 0.62$ ,  $P < 0.01$ )  
372 between the total signal intensity of CHO and CHON compounds in ESI+, we thus  
373 concluded that primary sources (i.e., fresh biomass material burning) were also one of  
374 the main sources of CHON compounds. In this study, CHON compounds with O/N <  
375 3 contributed  $76.48 \pm 1.11\%$  of total CHON species in ESI+ (**Figure S7**), which was  
376 much larger than the results observed in urban Tianjin in winter (less than 20%) (Zhong

377 et al., 2023). In particular, C<sub>16</sub>H<sub>33</sub>ON, C<sub>18</sub>H<sub>37</sub>ON, C<sub>18</sub>H<sub>35</sub>ON, C<sub>18</sub>H<sub>33</sub>ON, C<sub>18</sub>H<sub>31</sub>ON, 378 and C<sub>20</sub>H<sub>33</sub>ON showed a high abundance, together accounting for 55.04 ± 7.09 % of 379 the total CHON abundance (**Table S4**). The carbon number of these compounds was 380 consistent with that of fatty acids mentioned above; moreover, their abundances showed 381 a positive correlation ( $r = 0.43\text{--}0.81$ ,  $P < 0.01$ ) with the abundances of corresponding 382 fatty acids in the warm period. In contrast, these CHON compounds only showed a 383 weak correlation ( $r = -0.24 \sim 0.33$ ) with atmospheric oxidants (e.g., •OH, O<sub>3</sub>, and NO<sub>x</sub>). 384 Thus, the formation mechanism of biomass burning-related NOCs in Urumqi during 385 the warm period may be the interaction between fatty acids and reduced nitrogen 386 species (e.g., NH<sub>3</sub>) rather than the oxidation pathway involving CHO compounds and 387 NO<sub>x</sub>.

388 A recent laboratory study has suggested that NH<sub>3</sub> produced during the thermal 389 degradation of amino acids can react with oleic acid from the pyrolysis of triglycerides 390 to form amides (R1) (Ditto et al., 2022a). As discussed above, the combustion of fresh 391 biomass materials (e.g., forest fires or wildfires) can release abundant fatty acids. In 392 addition, wildfires can also emit large amounts of NH<sub>3</sub>, with an average emission factor 393 more than twice the NH<sub>3</sub> emission factor of agricultural fires (Tomsche et al., 2023). 394 According to MS/MS analysis (**Table S5**), potential fatty acid-derived NOCs were 395 indeed identified as amides. Thus, we proposed that the high temperature generated 396 during wildfires or forest fires provides suitable conditions for the reaction of 397 carboxylic acids and NH<sub>3</sub> to form amides. The specific process was presented in **Figure** 398 **5** (Pathway 1). It has been suggested that atmospheric oxidants can oxidize olefins (R2

399 and R3) to form hydroxyl nitrates and carbonyl nitrates (Perring et al., 2013). Therefore,  
 400 fatty acids (oleic acid as a representative) released from fresh biomass material burning  
 401 may also rely on oxidation pathways to form NOCs (**Figure 5**, Pathway 2). It is worth  
 402 noting that some products with double bonds after the amidation of unsaturated fatty  
 403 acids can continue to undergo the reactions of R2 and R3 in the atmosphere, resulting  
 404 in the formation of nitrooxy amides (**Figure 5**, Pathway 3). However, we found that the  
 405 abundance of oleic acid-derived amides via Pathway 1 in the warm period was more  
 406 than 100 times higher than that of NOCs with  $-ONH_2$  (thus, the impact of ionization  
 407 efficiency is expected to be less than 100 times) from Pathways 3. In the cold period,  
 408 the abundance of fatty acids-derived amides decreased dramatically (**Figure 5** and  
 409 **Figure S8**). Thus, the overall results demonstrated that the combustion of fresh biomass  
 410 materials indeed contributed significantly to aerosol NOCs (e.g., amides) in the warm  
 411 period in Urumqi.



415 The CHON species detected in ESI $^-$  were mainly aromatic-like compounds,  
 416 whose signal intensities were significantly greater in the cold period than in the warm  
 417 period (**Figures 4c,4e** and **Figure S5**). Moreover, we found that several nitro-aromatic  
 418 compounds, including  $C_6H_5O_3N$ ,  $C_6H_5O_4N$ ,  $C_7H_7O_3N$ ,  $C_7H_7O_4N$ ,  $C_7H_5O_5N$ , and  
 419  $C_8H_9O_3N$  (confirmed by their authentic standards in the LC/MS analysis), contributed

420 up to 50% of the total CHON (ESI- mode) intensity (**Table S6**). Other NOCs with  
421 relatively high signal intensity were mainly O<sub>4.6</sub>N<sub>2</sub> species (contributed up to 25%),  
422 such as C<sub>6</sub>H<sub>4</sub>O<sub>5</sub>N<sub>2</sub>, C<sub>7</sub>H<sub>4</sub>O<sub>7</sub>N<sub>2</sub>, C<sub>7</sub>H<sub>6</sub>O<sub>5</sub>N<sub>2</sub>, and C<sub>7</sub>H<sub>6</sub>O<sub>6</sub>N<sub>2</sub>, which have been suggested  
423 to be associated with secondary photochemical or multiphase chemical processes  
424 (Harrison et al., 2005; Cecinato et al., 2005; Salvador et al., 2021). However, the  
425 abovementioned nitro-aromatic compounds including C<sub>6</sub>H<sub>5</sub>O<sub>3</sub>N (nitrophenol),  
426 C<sub>6</sub>H<sub>5</sub>O<sub>4</sub>N (nitrocatechol), C<sub>7</sub>H<sub>7</sub>O<sub>3</sub>N (methyl-nitrophenol), and C<sub>7</sub>H<sub>7</sub>O<sub>4</sub>N (methyl-  
427 nitrocatechol) were primarily identified as tracers of straw and wood burning (old-age  
428 biomass materials commonly used in suburban and rural China) (Iinuma et al., 2010;  
429 Kourtchev et al., 2016). A study about molecular characterization (ESI- mode) of  
430 water-soluble aerosols emitted from the combustion of old-age biomass materials (i.e.,  
431 dry corn straw, rice straw, and pine branches) and coal showed that OA from old-age  
432 biomass burning typically contained much more nitro compounds and/or organonitrates  
433 than that from coal, while OA from coal-smoke contained more sulfur-containing  
434 compounds (Song et al., 2018). Thus, the old-age biomass burning associated with  
435 winter heating rather than coal combustion may contribute a significant amount of  
436 aerosol NOCs (e.g., nitrophenols) in wintertime Urumqi. However, it does not  
437 necessarily suggest that the importance of multiphase chemistry in the formation of  
438 NOCs was ignorable, as indicated by relatively high signal intensity of O<sub>4.6</sub>N<sub>2</sub> species.  
439 In general, the differential molecular characteristics of CHON species in different  
440 seasons in Urumqi can largely attributed to different impacts of the combustion of fresh-  
441 and old-age biomass materials.

442

443 **3.3. CHN molecule evidence of fresh and old-age biomass burning in different**  
444 **periods.**

445 **Figures 6a and 6b** present the van Krevelen diagram of CHN compounds in the  
446 cold and warm periods. The CHN<sub>1</sub> compounds with relatively high signal intensity  
447 mainly contained 7–20 carbon atoms, among which C<sub>5</sub>H<sub>5</sub>N(CH<sub>2</sub>)<sub>n</sub>, C<sub>9</sub>H<sub>7</sub>N(CH<sub>2</sub>)<sub>n</sub>, and  
448 C<sub>13</sub>H<sub>9</sub>N(CH<sub>2</sub>)<sub>n</sub> were dominant ( $78.68 \pm 7.59\%$  of the total signal intensity of CHN<sub>1</sub>  
449 compounds in the cold period, **Table S7**). C<sub>5</sub>H<sub>5</sub>N(CH<sub>2</sub>)<sub>n</sub> could be identified as pyridine  
450 and its homologues, which have been detected in freshly discharged BBOA (Dou et al.,  
451 2015). Additionally, the abundance of C<sub>5</sub>H<sub>5</sub>N(CH<sub>2</sub>)<sub>n</sub> was positively correlated with that  
452 of C<sub>9</sub>H<sub>7</sub>N(CH<sub>2</sub>)<sub>n</sub>, C<sub>13</sub>H<sub>9</sub>N(CH<sub>2</sub>)<sub>n</sub>, and nitro-aromatic compounds mentioned above ( $r =$   
453 0.46–0.81,  $P < 0.01$ ), particularly in the cold period with old-age biomass burning for  
454 heating. We further found that both the total signal intensity and aromaticity of CHN<sub>1</sub>  
455 species were much higher in the cold period (AI<sub>mod</sub> of 0.52) than in the warm period  
456 (AI<sub>mod</sub> of 0.35) (**Figure 6** and **Figure S9**). It has been suggested that old-age leaves  
457 contain more aromatic compounds compared to fresh leaves (Jian et al., 2016). Thus,  
458 the overall results implied that old-age biomass burning had an important contribution  
459 to the variation of CHN<sub>1</sub> compounds. In particular, the intensity of CHN<sub>1</sub> compounds  
460 was significantly negatively correlated with the concentration of O<sub>3</sub> and ·OH ( $r = -0.44$   
461  $\sim -0.53$ ,  $P < 0.01$ ), suggesting that atmospheric oxidation processes were the potential  
462 pathway for amine removal rather than the sources of particle amine salts (Zahardis et  
463 al., 2008; Qiu and Zhang, 2013). This result differed from the previous case, which

464 showed that the formation processes of CHN<sub>1</sub> and its homologs in Guangzhou (South  
465 China) were tightly related to photo-oxidation processes (Jiang et al., 2022). The CHN<sub>2</sub>  
466 species showed a similar temporal variation pattern to the CHN<sub>1</sub> species. Moreover, the  
467 abundances of total CHN<sub>2</sub> and major components (C<sub>8-11</sub>H<sub>8</sub>N<sub>2</sub>(CH<sub>2</sub>)<sub>n</sub>, C<sub>10</sub>H<sub>14</sub>N<sub>2</sub>(CH<sub>2</sub>)<sub>n</sub>,  
468 C<sub>10</sub>H<sub>16</sub>N<sub>2</sub>(CH<sub>2</sub>)<sub>n</sub> and C<sub>5</sub>H<sub>8</sub>N<sub>2</sub>(CH<sub>2</sub>)<sub>n</sub>) were positively correlated with that of total CHN<sub>1</sub>  
469 ( $r = 0.55\text{--}0.90$ ,  $P < 0.01$ ), but negatively correlated with the concentration of O<sub>3</sub>  
470 and ·OH ( $r = -0.43 \sim -0.60$ ,  $P < 0.01$ ). Clearly, old-age biomass burning, particularly  
471 in the cold period, also exerted significant impacts on the abundance of CHN<sub>2</sub>  
472 compounds, which was also supported by several previous studies (Laskin et al., 2009;  
473 Wang et al., 2017b; Song et al., 2022). A study about molecular characterization (ESI+  
474 mode) of humic-like substances emitted from the combustion of old-age biomass  
475 materials (i.e., dry corn straw, rice straw, and pine branches) and coals showed that OA  
476 from old-age biomass burning typically contained much more CHN<sub>2</sub> compounds (55–  
477 64%) than that from coal (20–37%), while OA from coal-smoke showed more CHN<sub>1</sub>  
478 compounds (78–84%) compared to that from old-age biomass materials (15–22%)  
479 (Song et al., 2022). In this study, the signal intensity of CHN<sub>1</sub> compounds in the cold  
480 period was about 40% higher than that in the warm period, while that of CHN<sub>2</sub>  
481 compounds showed a 160% increase from the warm period to the cold period. Thus,  
482 although the contribution of fossil fuel (e.g., coal) combustion to NOCs in the cold  
483 period cannot be ignored, our results at least suggested that the biomass burning-derived  
484 CHN compounds showed a more significant increase compared to coal combustion-  
485 derived compounds from the warm period to the cold period in Urumqi.

Interestingly, we found some CHN species with 16–20 carbon atoms showed higher abundance in the warm period than in the cold period, a pattern opposite to that of all other CNH compounds (**Figure 6c**). These  $C_{16-20}N_1H_x$  compounds were further identified as alkyl nitriles (**Table S5**) (Simoneit et al., 2003). In addition, the carbon number of the identified alkyl nitriles was consistent with those of amides previously proposed to be produced by fresh biomass burning. Thus, we proposed that fresh biomass material burning in the warm period may provide a continuous high-temperature environment to promote the dehydration of amides (**Figure 5**, Pathway 4). These alkyl nitriles with double bonds can continue to undergo the reactions of R2 and R3 (**Figure 5**, Pathway 5). However, the signal intensity of the nitrooxy products in the warm period was insignificantly correlated with the concentration of  $O_3$ ,  $\cdot OH$ , and  $NO_x$  ( $P > 0.05$ ), likely indicating a weak influence of atmospheric oxidation on alkyl nitrile removal in this site. The high-temperature dehydration of amides (e.g., erucamide) to form alkyl nitriles (e.g., erucyl nitrile) has been demonstrated by Simoneit et al. (2003) in a laboratory simulation experiment. A study on BBOA also showed that alkyl nitriles can be serve as indicators of biomass burning in the ambient atmosphere (Radzi Bin Abas et al., 2004). Furthermore, the abundance of identified alkyl nitriles initially increased from March and peaked in September and October (**Figure S10**), a pattern which was consistent with the interannual variation in wildfire areas (more in the warm period) in Central Asian countries (Xu et al., 2021). Although cooking is also a potential source of alkyl nitriles (Schauer et al., 1999), this activity does not have seasonal differences. In contrast, the dramatically increased abundance of aromatic CNH

508 compounds in the cold period (**Figure S9**) can be attributed to the aqueous reactions of  
509 amines emitted from old-age biomass material and coal combustion with acidic  
510 substances, as indicated by significant correlations ( $r = 0.61\text{--}0.95$ ,  $P < 0.01$ ) between  
511 total CHN abundance and  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  concentrations. These findings further  
512 confirmed that the NOCs from the combustion of fresh biomass materials in the warm  
513 period in suburban Urumqi were compositionally different from those from old-age  
514 biomass burning in the cold period.

515

#### 516 **4. Conclusions**

517 The complexity of NOCs restricts our understanding of its sources and formation  
518 processes. In this study, the molecular compositions of organic aerosols in  $\text{PM}_{2.5}$   
519 collected in Urumqi over a one-year period were systematically characterized in both  
520 ESI- and ESI+ modes, with a major focus on NOCs. A large amount of NOCs were  
521 identified, showing that NOCs in relatively highly oxidative and reduced forms can be  
522 roughly distinguished via these two ionization modes. Based on the identification of  
523 molecular markers of amides and alkyl nitriles (much higher in the warm period) and  
524 the analysis of their formation mechanisms (less contribution of atmospheric oxidation),  
525 we highlighted the important contribution of combustion of fresh biomass materials  
526 such as forest fires and wildfires to NOCs in the warm season in Urumqi. In contrast,  
527 the dramatically increased abundances of aromatic CNH compounds and nitro-aromatic  
528 CHON compounds (mainly nitrophenols) in the cold period were tightly associated  
529 with the impacts of old-age biomass material burning. These results were illustrated in

530 a diagram (**Figure 7**).

531 Biomass materials in rural China were typically old-age plant tissues, as  
532 mentioned above. Fresh biomass materials (e.g., green vegetation) with the enrichment  
533 of oils and proteins can exist in forest fires or wildfires. Indeed, previous studies have  
534 suggested that biomass burning can lead to the formation of aerosol amines and nitriles.  
535 However, field observation studies have yet to pay attention to the differences in aerosol  
536 NOCs emitted from the combustion of fresh and old-age biomass materials. For the  
537 first time, our results reveal that fresh biomass material combustion can contribute more  
538 amines and nitriles than old-age biomass material combustion. Generally, this study  
539 provides field evidence on the differential impacts of the combustion of fresh and old-  
540 age biomass materials on aerosol NOCs, improving our current understanding of the  
541 molecular compositions of organic nitrogen aerosols in a vast territory with a sparse  
542 population in Northwest China. Moreover, according to the fact that the studied site is  
543 highly affected by combustion emissions of different types of biomass materials, future  
544 work is needed to deeply understand the quantitative contributions of different types of  
545 biomass burning to OA in China.

546

547 **Data availability.** The data in this study are available at  
548 <https://doi.org/10.5281/zenodo.10453929>

549

550 **Competing interests.** The authors declare no conflicts of interest relevant to this study.

551

552 **Supplement.** Details of chemical analysis and data processing, eight tables (Tables  
553 S1–S8), and ten extensive figures (Figures S1–S10).

554

555 **Author contributions.** YX designed the study. YJM, TY, and HWX performed field  
556 measurements and sample collection; YJM and TY performed chemical analysis; YX  
557 and YJM performed data analysis; YX and YJM wrote the original manuscript; and YX,  
558 YJM, HWX, and HYX reviewed and edited the manuscript.

559

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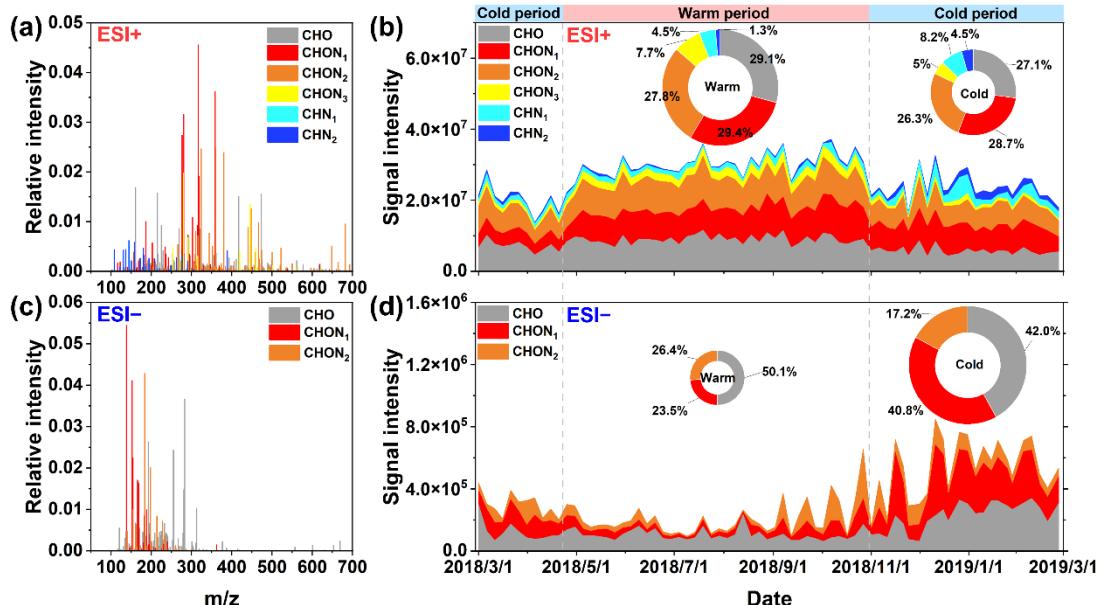
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957

958 **Figure 1.**

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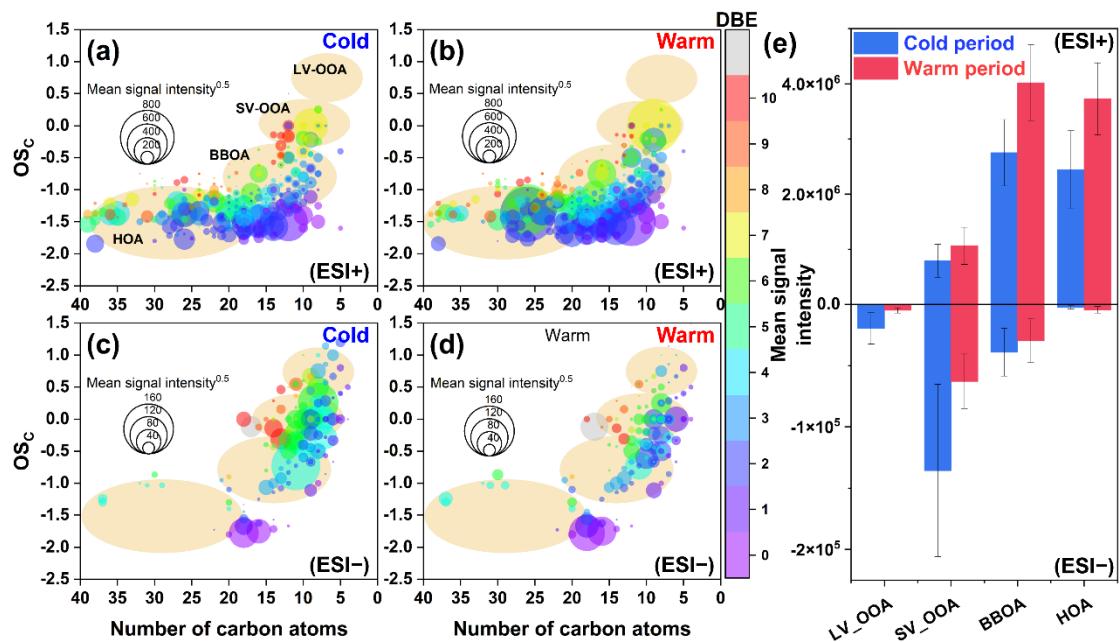
960 **Figure 1.** The reconstructed mass spectrum distribution of the detected species in  $\text{PM}_{2.5}$ 

961 in (a) ESI+ and (c) ESI- modes during the whole campaign. Temporal variations in the  
 962 fractional distribution of classified compounds in (b) ESI+ and (d) ESI- modes. The  
 963 ring diagrams inside the panel show the signal intensity fractions of classified  
 964 compounds, the size of which is proportional to the total signal intensity of all species  
 965 detected in  $\text{PM}_{2.5}$  in different periods.

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969 **Figure 2.**

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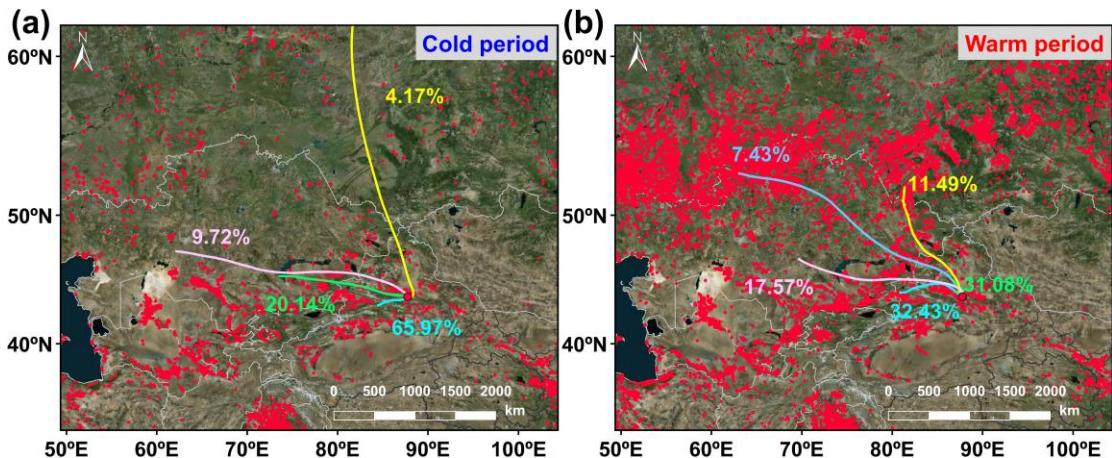
971 **Figure 2.** OSc values of CHO molecules detected in (a and b) ESI+ and (c and d) ESI-  
 972 modes in PM<sub>2.5</sub> collected from different periods (cold vs. warm). The size and color of  
 973 the circle indicate the mean signal intensity and DBE value of compounds, respectively.  
 974 The light-orange background indicates the areas of LV-OOA, SV-OOA, BBOA, and  
 975 HOA (Kroll et al., 2011), according to which (e) the mean signal intensity of classified  
 976 compounds was calculated for samples from different periods.

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981 **Figure 3**

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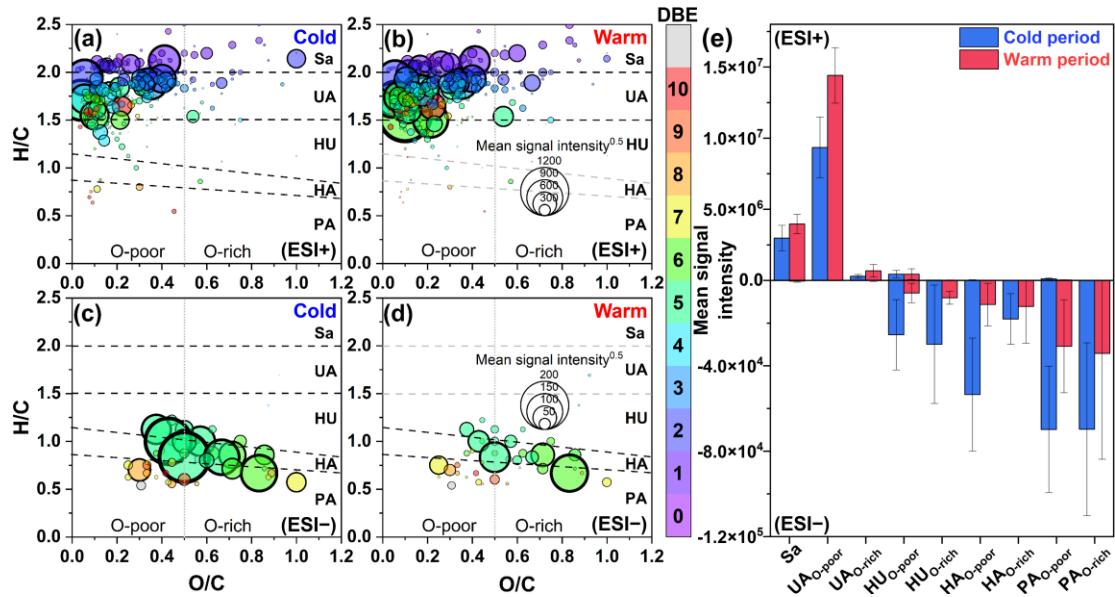
983 **Figure 3.** The 3-day (72 h) back trajectories illustrating the typical air mass flows to  
 984 the study site during the (a) warm and (b) cool periods. Fire spots were shown in red,  
 985 which was created based on NASA active fire data (VIIRS 375 m,  
 986 [https://firms.modaps.eosdis.nasa.gov/active\\_fire/](https://firms.modaps.eosdis.nasa.gov/active_fire/)). The map was derived from  
 987 ©MeteoInfoMap (version 3.6.2) (Chinese Academy of Meteorological Sciences, China).

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992 **Figure 4.**

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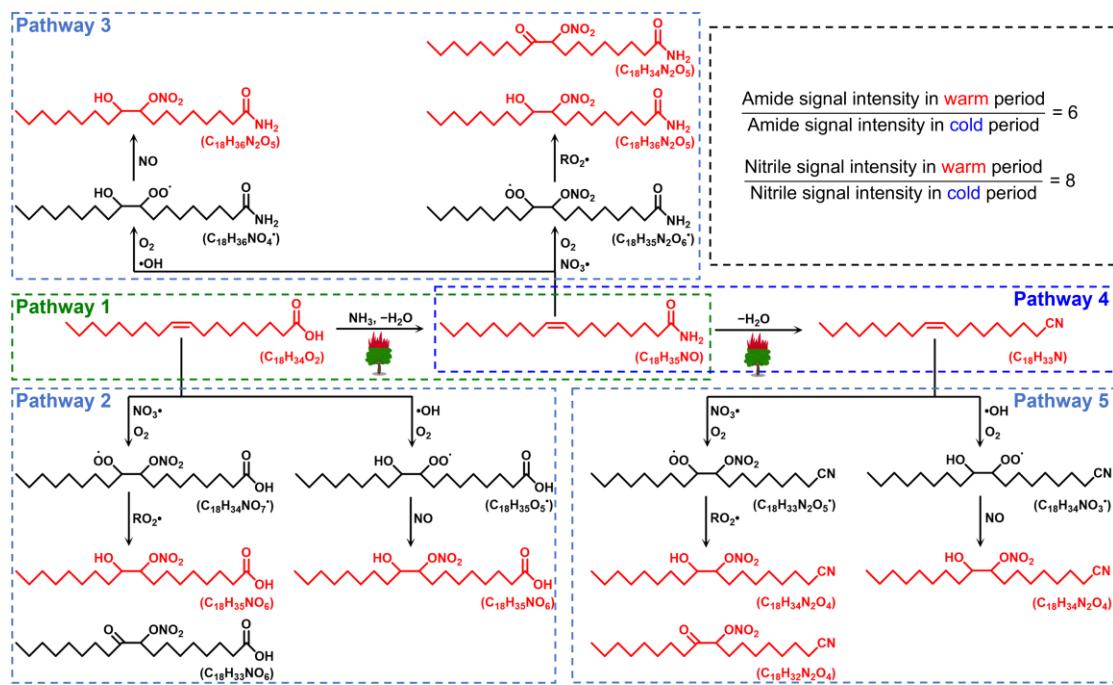
994 **Figure 4.** Van Krevelen diagrams of CHON molecules detected in (a and b) ESI+ and  
 995 (c and d) ESI- modes in PM<sub>2.5</sub> collected from different periods (cold vs. warm). The  
 996 subgroups in the panel include saturated-like (Sa), unsaturated aliphatic-like (UA),  
 997 highly unsaturated-like (HU), highly aromatic-like (HA), and polycyclic aromatic-like  
 998 (PA) compounds, further distinguishing between oxygen-poor and oxygen-rich  
 999 compounds with an oxygen to carbon ratio of 0.5. The size and color of the circle  
 1000 indicate the mean signal intensity and DBE value of compounds, respectively. The (e)  
 1001 mean signal intensity of classified compounds was calculated for samples from  
 1002 different periods.

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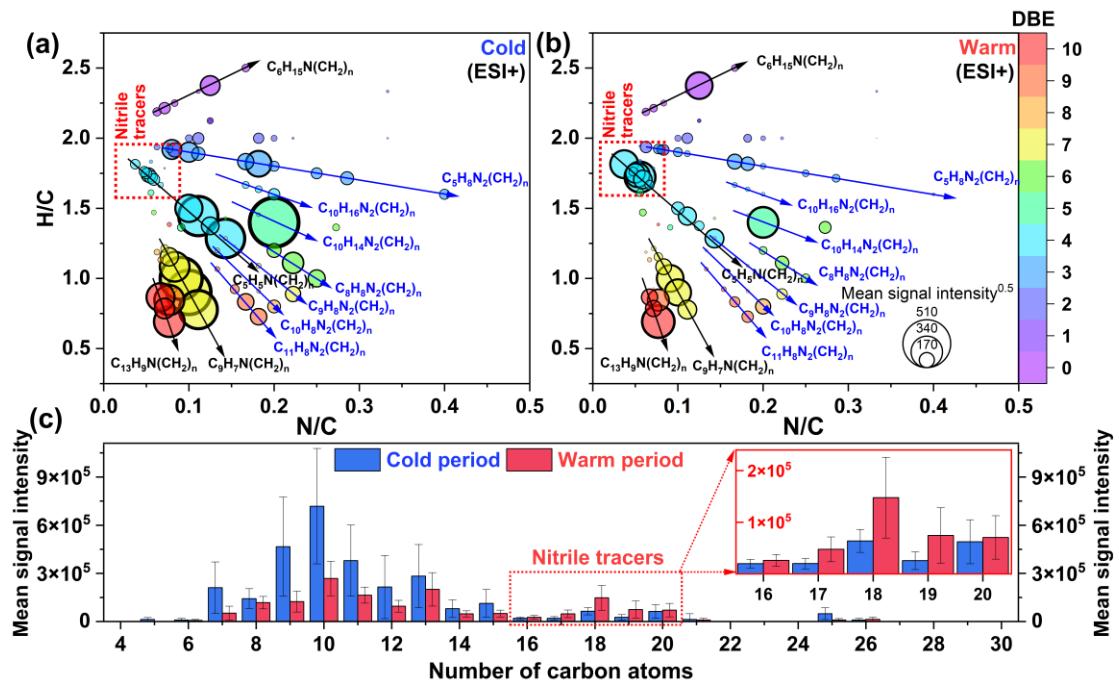
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**Figure 5.**

1009 **Figure 5.** Proposed pathways for the reactions of carboxylic acids (oleic acid as a  
 1010 representative) with reactive nitrogen and atmospheric oxides to form the observed  
 1011 NOCs in PM<sub>2.5</sub> under the influence of the high temperature generated during wildfires  
 1012 or forest fires. Compounds observed in PM<sub>2.5</sub> were shown in red.

1014 **Figure 6.**

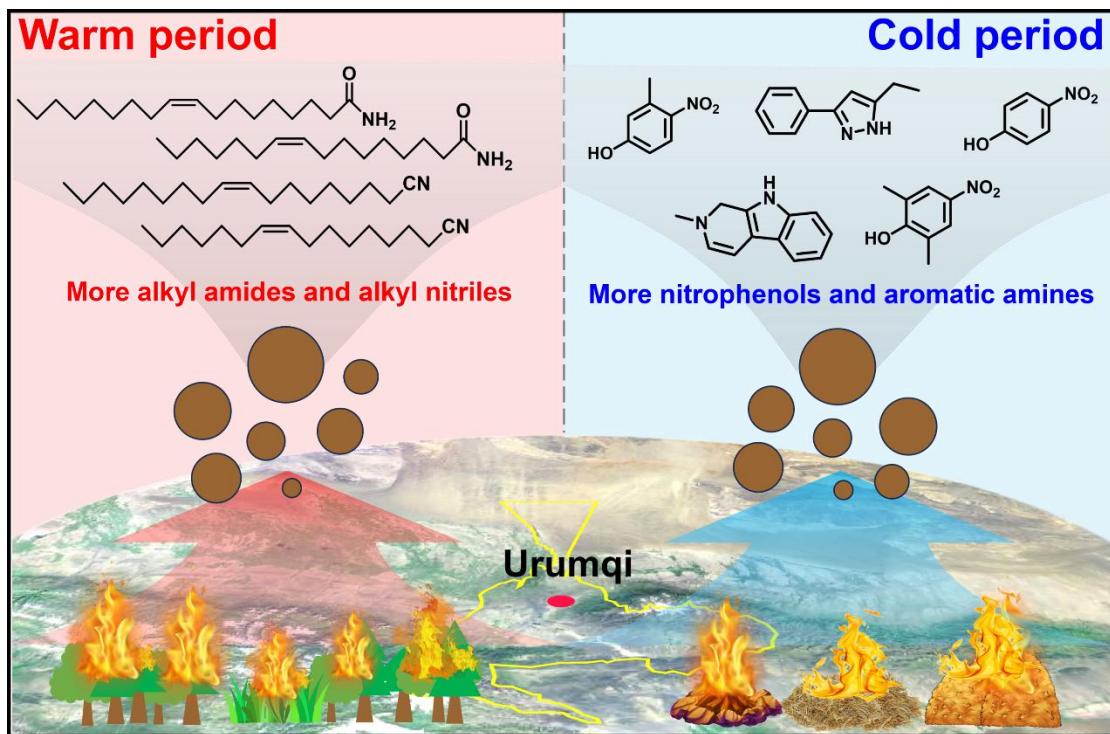


1015

1016 **Figure 6.** Van Krevelen diagrams of CHN molecules detected in PM<sub>2.5</sub> collected from  
1017 the (a) cold and (b) warm periods. The size and color of the circle indicate the mean  
1018 signal intensity and DBE value of compounds, respectively. The mean signal intensity  
1019 distributions of (c) carbon atoms in CHN molecules detected in PM<sub>2.5</sub> collected from  
1020 the cold and warm periods

1021

1022 **Figure 7.**



1023

1024 **Figure 7.** Conceptual picture showing the differential impacts of combustion of fresh  
1025 and old-age biomass materials on aerosol NOCs in suburban Urumqi. The map was  
1026 derived from <sup>©</sup>Baidu Maps (BIDU, China).