

1 **Insights into Soil NO Emissions and the Contribution to Surface** 2 **Ozone Formation in China**

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13 **Keywords:** Soil NO emissions; Ground-level ozone; BDSNP; OSAT

14
15 **Abstract.** Elevated ground-level ozone concentrations have emerged as a major
16 environmental issue in China. Nitrogen oxide (NO_x) is a key precursor to ozone formation.
17 Although control strategies aimed at reducing NO_x emissions from conventional combustion
18 sources are widely recognized, soil NO_x emissions (mainly as NO) due to microbial processes
19 have received little attention. The impact of soil NO emissions on ground-level ozone
20 concentration is yet to be evaluated. This study estimated soil NO emissions in China using
21 the Berkeley-Dalhousie soil NO_x parameterization (BDSNP) algorithm. A typical modeling
22 approach was used to quantify the contribution of soil NO emissions to surface ozone
23 concentration. The Brute-force method (BFM) and the Ozone Source Apportionment
24 Technology (OSAT) implemented in the Comprehensive Air Quality Model with extensions
25 (CAMx) were used. The total soil NO emissions in China for 2018 were estimated to be
26 1157.9 Gg N, with an uncertainty range of 715.7~1902.6 Gg N. Spatially, soil NO emissions
27 are mainly concentrated in Central China, North China, Northeast China, northern Yangtze
28 River Delta (YRD) and eastern Sichuan Basin, with distinct diurnal and monthly variations
29 that are mainly affected by temperature and the timing of fertilizer application. Both the BFM
30 and OSAT results indicate a substantial contribution of soil NO emissions to the maximum
31 daily 8-hour (MDA8) ozone concentrations by 8~12.5 μg/m³ on average for June 2018, with
32 the OSAT results consistently higher than BFM. The results also showed that soil NO
33 emissions led to a relative increase in ozone exceedance days by 10%~43.5% for selected
34 regions. Reducing soil NO emissions resulted in a general decrease in monthly MDA8 ozone
35 concentrations, and the magnitude of ozone reduction became more pronounced with
36 increasing reductions. **However, even with complete reductions in soil NO emissions,**
37 **approximately 450.3 million people are still exposed to unhealthy ozone levels, necessitating**

38 multiple control policies at the same time. This study highlights the importance of soil NO
39 emissions for ground-level ozone concentrations and the potential of reducing NO emissions
40 as a future control strategy for ozone mitigation in China.

41 1. Introduction

42 A substantial decrease in the atmospheric fine particulate matter (PM_{2.5}) concentrations has
43 been witnessed during the past decade in China (Zhai et al., 2019; Xiao et al., 2020; Maji,
44 2020) while the ground-level ozone (O₃) concentrations do not exhibit a steady downward
45 trend (Lu et al., 2020; Lu et al., 2021; Wang et al., 2022a; Sun et al., 2021). Because high
46 ozone concentration increases respiratory and circulatory risks (Malley et al., 2017; Cakaj et
47 al., 2023; Wang et al., 2020) and reduces crop yields (Feng et al., 2019; Lin et al., 2018;
48 Mukherjee et al., 2021; Montes et al., 2022), the coordinate control of PM_{2.5} and O₃ was
49 proposed as part of the 14th Five-year plan (State Council, 2021). A continuous increase in
50 summertime surface ozone was observed across China's nationwide monitoring network from
51 2013 to 2019, followed by an unprecedented decline in 2020 (except for Sichuan Basin) (Sun
52 et al., 2021), which is equally attributed to meteorology and anthropogenic emissions
53 reductions (Yin et al., 2021). As a secondary air pollutant, ozone is generated by the
54 photochemical oxidation of volatile organic compounds (VOC) in the presence of nitrogen
55 oxides (NO_x = NO + NO₂), both of which are considered ozone precursors. The control
56 strategies to mitigate ozone pollution in China focused on reducing NO_x emissions at an early
57 stage and started to stress the control of VOCs emissions in recent years (e.g., the 2020 action
58 plan on VOCs mitigation), including control of fugitive emissions, stringent emissions
59 standards, and substituting raw materials with low VOCs content (Ecology, 2020). Ding et al.
60 (2021) concluded that for North China Plain (NCP), a region that experienced the most severe
61 PM_{2.5} and ozone pollution in China, reductions in NO_x emissions are essential regardless of
62 VOC reduction.

63 Existing control strategies for NO_x emissions are almost exclusively targeted at combustion
64 sources, for example, power plants, industrial boilers, cement production, and vehicle
65 exhausts (Sun et al., 2018; Ding et al., 2017; Diao et al., 2018). However, NO_x emissions
66 from soils (mainly as NO), as a result of microbial processes (e.g., nitrification and
67 denitrification), could make up a substantial fraction of the total NO_x emissions (Lu et al.,
68 2021; Drury et al., 2021), yet is often overlooked. In California, soil NO_x emissions in July
69 accounted for 40% of the state's total NO_x emissions (when using an updated estimation
70 algorithm) and resulted in 23% of enhanced surface ozone concentration (Sha et al., 2021).
71 However, a wide range of annual soil NO_x emissions from 8,685 tons (as NO₂, (Guo et al.,
72 2020) to 161,100 metric tons of NO_x-N (Almaraz et al., 2018) were reported depending on
73 different methods. Romer et al. (2018) estimated that nearly half of the increase in hot-day

74 ozone concentration in a forested area of the rural southeastern United States is attributable to
75 the temperature-induced increases in NO_x emissions, mostly likely due to soil microbes.

76 Soil NO emissions are affected by many factors, including nitrogen fertilizer application, soil
77 organic carbon content, soil temperature, humidity, and pH (Vinken et al., 2014; Yan et al.,
78 2005; Wang et al., 2021; Skiba et al., 2021). The amount of nitrogen fertilizer application in
79 China was estimated to account for one-third of the global nitrogen fertilizer application
80 (Heffer and Prud'homme, 2016), with most of the land under high nitrogen deposition (Liu et
81 al., 2013; Lü and Tian, 2007). Therefore, soil NO emissions in China are expected to be
82 significant, and their impacts on ozone pollution need to be systematically evaluated. So far,
83 only a limited number of studies have addressed this issue in China (Lu et al., 2021; Shen et
84 al., 2023; Wang et al., 2008; Wang et al., 2022b). Lu et al. (2021) concluded that soil NO
85 significantly reduced the ozone sensitivity to anthropogenic emissions in NCP, therefore,
86 causing a so-called “emissions control penalty”. Wang et al. (2022b) reported NO_x emissions
87 from cropland contributed 5.0% of the maximum daily 8h average ozone (MDA8 O₃) and
88 27.7% of NO₂ concentration in NCP. These studies focused solely on NCP, a region with
89 persistent O₃ pollution in warm seasons (Liu et al., 2020; Lu et al., 2020). The impact of soil
90 NO emissions on ozone concentrations over other regions, for example, the northern Yangtze
91 River Delta (YRD) and Sichuan Basin, where soil emissions are high (see Section 3.1) and
92 ozone pollution is also severe (Shen et al., 2022; Yang et al., 2021), has not been much
93 evaluated in details (Shen et al., 2023). In addition, the method employed in existing studies
94 to evaluate soil NO emissions on ozone concentration is the conventional “brute-force” zero-
95 out approach, which might be inappropriate given the strong nonlinearity of the ozone
96 chemistry (Clappier et al., 2017; Thunis et al., 2019).

97 With the deepening of emissions control measures for power, industrial and on-road sectors,
98 anthropogenic NO_x emissions from combustion sources have decreased at a much faster rate
99 (by 4.9% since 2012) than that from soil (fertilizer application decreases at a rate of 1.5%
100 since 2015, Fig. S1). Therefore, understanding the impacts of soil NO emissions on ground-
101 level ozone concentration, particularly considering the spatial heterogeneities over different
102 regions of China, is of great importance for formulating future ozone mitigation strategies. In
103 this study, soil NO emissions in China for 2018 were estimated based on a most recent soil
104 NO parameterization scheme with updated fertilizer data as input. The spatial and temporal
105 variations of soil NO emissions were described first. Uncertainties associated with estimation
106 of soil NO emissions were discussed. An integrated meteorology and air quality model was
107 applied to quantify the impact of soil NO emissions on surface ozone concentration based on
108 two different methods. Lastly, we evaluated the changes in ozone concentration and exposed
109 population under different emission scenarios to highlight the effectiveness of reducing soil
110 NO emissions as potential control policy. Our results provide insights into developing

111 effective emissions reduction strategies to mitigate the ozone pollution in China.

112 2. Methodology

113 2.1. Estimation of soil NO emissions in China

114 Soil NO emissions were estimated based on the Berkeley-Dalhousie Soil NO_x
115 Parameterization (BDSNP) that is implemented in the Model of Emissions of Gases and
116 Aerosols from Nature (MEGAN) version 3.2 (<https://bai.ess.uci.edu/megan/data-and-code>,
117 accessed on September 1st, 2021). The BDSNP algorithm estimates the soil NO emissions by
118 adjusting a biome-specific NO emissions factor in response to various conditions, including
119 the soil temperature, soil moisture, precipitation-induced pulsing, and a canopy reduction
120 factor (Eq. 1, Rasool et al., 2016):

$$121 \text{NO}_{\text{emission flux}} = A'_{\text{biome}}(N_{\text{avail}}) \times f(T) \times g(\theta) \times P(l_{\text{dry}}) \times \text{CRF}(\text{LAI}, \text{Biome}, \text{Meterology}) \quad \text{Eq. 1}$$

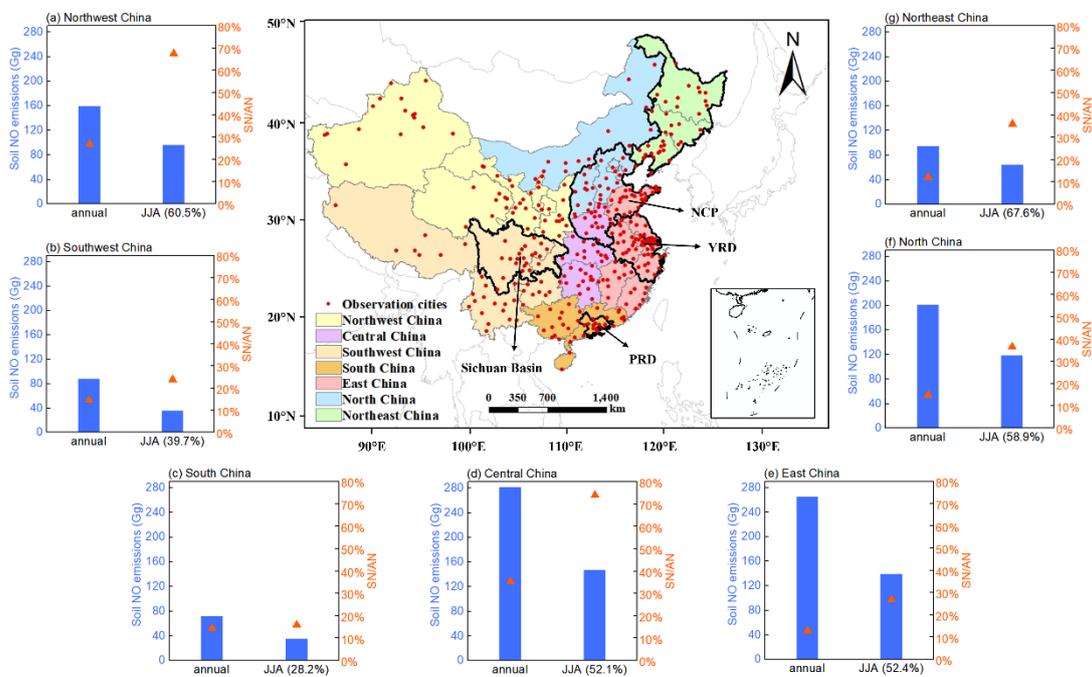
121 where $f(T)$ and $g(\theta)$ is the temperature (T , unit: K) and soil moisture (θ , unit: m^3/m^3)
122 dependence functions, respectively; $P(l_{\text{dry}})$ represents the pulsed soil emissions due to wetting
123 of dry soils; l_{dry} (hours) is the antecedent dry period of a pulse; and CRF describes the canopy
124 reduction factor, which is a function of the leaf area index (LAI, m^2/m^2) and the meteorology.
125 A'_{biome} ($\text{ng N m}^{-2} \text{s}^{-1}$) is the biome-specific emission factor, which is further calculated as Eq.2:

$$A'_{\text{biome}} = A_{w,\text{biome}} + N_{\text{avail}} \times \bar{E} \quad \text{Eq. 2}$$

126 In Eq. 2, $A_{w,\text{biome}}$ ($\text{ng N m}^{-2} \text{s}^{-1}$) is the wet biome-dependent emission factor; N_{avail} is the
127 available nitrogen from fertilizer and deposition; \bar{E} is the emission rate based on an observed
128 global estimates of fertilizer emissions (Rasool et al., 2016). The detailed expressions of these
129 parameters are presented in the Supporting Information. More information on the BDSNP
130 parameterizations can be found in previous studies (Hudman et al., 2012).

131 The default N fertilizer input data provided with the BDSNP algorithm is based on the a
132 (Potter et al., 2010), which gives a number of 19.6 Tg N/a. In this study, we collected fertilizer
133 data from statistical yearbooks at the provincial level. The total amount of pure nitrogen
134 fertilizer (hereafter N fertilizer) applied in the year 2018 is 20.7 Tg N/a, which is similar
135 (5.6% higher) to IFA value. However, besides the N fertilizer, NPK compound fertilizer
136 (containing nitrogen (N), phosphorous (P), and potassium (K)) is being increasingly applied
137 in China. According to the statistical yearbook, the amount of N fertilizer applied decreased
138 from 23.5 Tg in 2010 to 20.7 Tg in 2018 (a relative reduction of 11.9%). In contrast, NPK
139 fertilizer increased from 18.0 in 2010 to 22.7 Tg in 2018 (a relative increase of 26.1%). We
140 assumed one-third of the NPK fertilizer is nitrogen (Liu, 2016); thus, the total amount of
141 nitrogen applied as fertilizer is 28.2 Tg N in 2018, which is 43.9% higher than the value from
142 Potter et al. (2010). We divided China into seven regions for emission analysis at regional

143 scale, namely Northeast China, North China, Central China, East China, South China,
 144 Southwest China, and Northwest China, as indicated by different colors in Fig. 1 (see Table
 145 S1 for the list of provinces in each region). At the regional level, the amount of total fertilizer
 146 differs by as much as 9.1% to 46.4% from the default fertilizer (Table S2).



147
 148 **Figure 1.** Modeling domain and region definitions. Surrounding charts show the annual and
 149 summer (June-July-August, JJA) soil NO emissions and ratio of soil NO to anthropogenic
 150 NO_x emissions for each region.

151 **2.2. Model configurations**

152 A typical modeling approach was applied to evaluate the contribution of soil NO emissions to
 153 surface ozone concentration. The Weather Research and Forecasting (WRF) model (version
 154 3.7, <https://www.mmm.ucar.edu/wrf-model-general>, accessed on December 1st, 2021) and the
 155 Comprehensive Air Quality Model with Extension (CAMx, version 7.0,
 156 <http://www.camx.com/>, accessed on December 1st, 2021) were applied to simulate the
 157 meteorological fields and subsequent ozone concentrations. Table S3 listed the detailed model
 158 configurations for WRF and CAMx. Anthropogenic emissions include the Multi-resolution
 159 Emission Inventory of China for 2017 (MEIC, <http://www.meicmodel.org>, accessed on
 160 December 1st, 2021) and the 2010 European Commission's Emissions Database for Global
 161 Atmospheric Research (EDGAR, <http://edgar.jrc.ec.europa.eu/index.php>, accessed on
 162 December 1st, 2021) for outside China. Biogenic emissions were calculated along with the
 163 soil NO emissions using MEGAN3.2. Open biomass burning emissions are adopted from the
 164 Fire INventory from NCAR version (FINN, version 1.5,
 165 <https://www.acom.ucar.edu/Data/fire/>) with MOZART speciation and converted to CAMx

166 CB05 model species. The gaseous and aerosol modules used in CAMx include the CB05
167 chemical mechanism (Yarwood et al., 2010) and the CF module. The aqueous-phase
168 chemistry is based on the updated mechanism of the Regional Acid Deposition Model
169 (RADM) (Chang et al., 1987). A base case simulation was conducted for June 2018 when soil
170 NO emissions reached maxima (Section 3.1) and ozone pollution was severe over eastern
171 China (Mao et al., 2020; Jiang et al., 2022). Base case model performances have been
172 evaluated in our previous studies (Huang et al., 2021; Huang et al., 2022b). Here we evaluated
173 simulated ozone concentrations using the Pearson correlation coefficient (R), mean bias
174 (MB), root-mean-square error (RMSE), normalized mean bias (NMB), and normalized mean
175 error (NME) against hourly observed ozone concentrations for 365 cities in China. The
176 formula for each of the statistical metrics is given in Table S4. Observed hourly ozone
177 concentrations were obtained from the China National Environmental Monitoring Center.

178 2.3. Brute-force and OSAT

179 In this study, two methods were used to quantify the impact of soil NO emissions on surface
180 ozone concentration during the simulation period. The first is the conventional brute-force
181 method (BFM), which involves comparing the simulated ozone concentration between the
182 base case and a scenario case without soil NO emissions. The difference between these two
183 scenarios was considered to represent the contribution of soil NO emissions to ozone. The
184 second method applies the widely used Ozone Source Apportionment Technology (OSAT)
185 implemented in CAMx (Yarwood et al., 1996), with soil NO emissions being tagged as an
186 individual emission group. OSAT attributes ozone formation to NO_x or VOCs based on their
187 relative availability and apportions NO_x and VOCs emissions by source group/region
188 (Ramboll, 2021). In addition to soil NO emissions, anthropogenic and natural emissions
189 (including biogenic VOC emissions, lightning NO emissions, and open biomass burning)
190 were also tagged as individual emission groups.

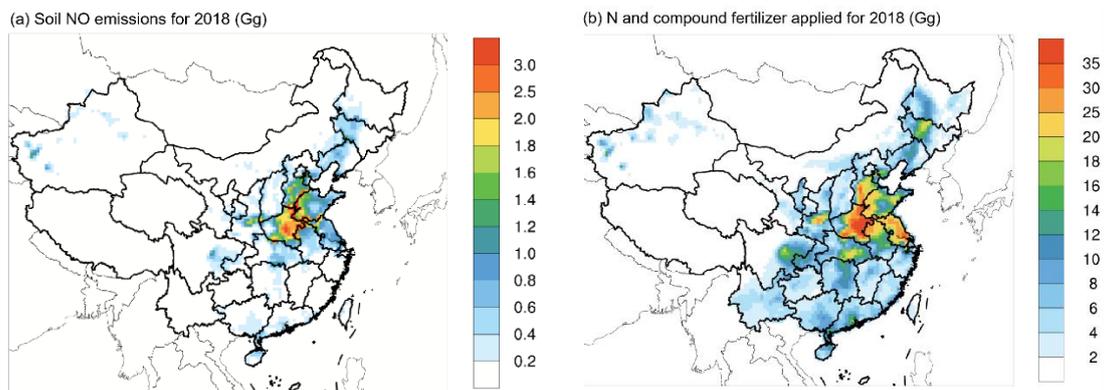
191 3. Results and discussions

192 3.1. Soil NO emissions for 2018 in China

193 3.1.1. Spatial and temporal variations

194 National total soil NO emissions for 2018 is estimated to be 1157.9 Gg N, with an uncertainty
195 range of 715.7~1902.6 Gg N, which will be discussed more in Section 3.1.2. On an annual
196 scale, soil NO emissions accounted for 17.3% of the total anthropogenic NO_x emissions in
197 China for 2017 (based on MEIC inventory). This ratio varies from 12.0% to 35.3% at regional
198 scale. Unlike the anthropogenic NO_x emissions that concentrate over densely populated
199 regions (e.g., NCP, YRD), soil NO emissions are most abundant in Central China, particularly
200 Henan Province and nearby provinces, including Hebei and Shandong in the NCP, Jiangsu

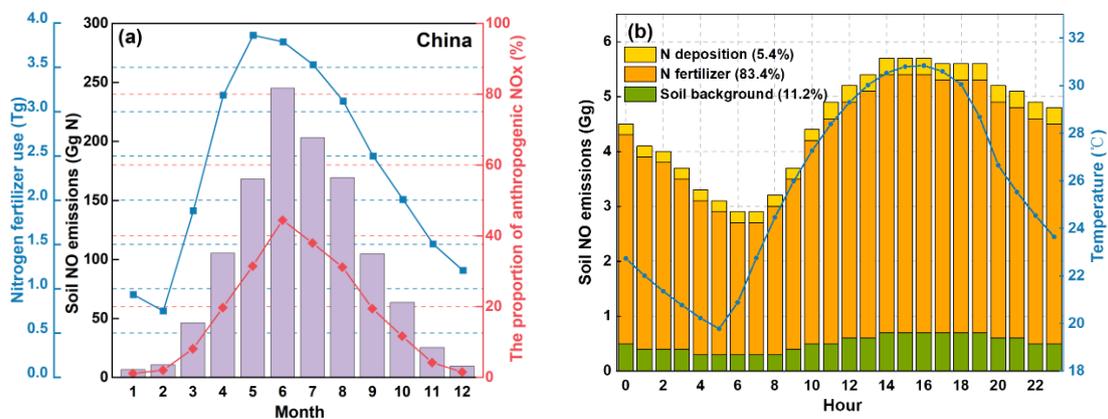
201 and Anhui in northern YRD (Fig. 2a). Other hotspots of soil NO emissions include Northeast
202 China and the eastern part of the Sichuan Basin. As expected, the spatial distribution of soil
203 NO emissions closely mirrors that of the fertilizer application (Fig. 2b). Henan (located in
204 Central China), Shandong (NCP), and Hebei (NCP) are the top three provinces that have the
205 highest fertilizer application (together accounting for 24.1% of national totals in 2018) and
206 thus highest soil NO emissions (together accounting for 35.7%).



207 **Figure 2.** Spatial distribution of (a) soil NO emissions for 2018 and (b) N and compound
208 fertilizer applied for 2018.

209 In terms of the monthly variations, the total soil NO emissions show a unimodal pattern (as
210 shown in Fig. 3a with the highest emissions occurring in the summer months of June, July,
211 and August), except for South China and Southeast China (Fig. S2), where the peak emissions
212 occur in April or May. Soil NO emissions during the summer months account for 28.2%
213 (South China) to 67.6% (Northeast China) of the annual totals (Fig. 1 and Table S5). The
214 shape of monthly soil NO emissions is influenced by temperature and the timing of fertilizer
215 application. The BDSNP algorithm assumes that 75% of the annual fertilizer is applied over
216 the first month of the growing season, with the remaining 25% applied evenly throughout the
217 rest of the growing season. This assumption results in a significant amount of fertilizer being
218 applied from April to August (Fig. 3a). In contrast, anthropogenic NO_x emissions display
219 weaker monthly variations (Zheng et al., 2021). Consequently, the ratio of soil NO emissions
220 to anthropogenic NO_x (SN/AN) is much higher during the summer months. In regions such as
221 Central China and Northwest China, where soil NO emissions are high and anthropogenic
222 NO_x emissions are relatively low, SN/AN reaches 74.0% and 67.5% during the summer
223 months (Fig. 1 and Table S5). In East China and North China, where anthropogenic NO_x
224 emissions are high, SN/AN ranges from 26.8% to 36.5% during the summer months. These
225 findings align with Chen et al. (2022), who reported that soil NO emissions made up 28%
226 of total NO_x (soil NO + anthropogenic NO_x) emissions in summer and could reach 50–90%
227 in isolated areas and suburbs. The substantial contribution of soil NO emissions during the
228 ozone pollution season implies a potentially significant impact on surface ozone

229 concentration. In terms of diurnal variations, soil NO emissions peak in the afternoon due to
 230 diurnal temperature fluctuations. As illustrated by Fig. 3b, the average hourly soil NO
 231 emissions over NCP for June 2018 closely follow the WRF simulated temperature changes.
 232 The BDSNP algorithm identifies three sources of soil nitrogen: background, atmospheric
 233 nitrogen deposition, and fertilizer application, with the latter being the primary contributor. A
 234 decomposition analysis of soil NO emissions for NCP reveals that fertilizer application
 235 accounts for 83.4% of total NO soil emissions (Fig. 3b), while background and atmospheric
 236 nitrogen deposition only contribute for 11.2% and 5.4%, respectively. Thus, although soil NO
 237 emissions are generally considered a “natural” source (Galbally et al., 2008) and are not
 238 currently targeted in NO_x emission mitigation strategies, human fertilizer activities render soil
 239 NO emissions an anthropogenic source.



240 **Figure 3.** (a) Monthly fertilizer (N + compound) applied and soil NO emissions in China and
 241 (b) hourly soil NO emissions for 2018 June in NCP and domain-averaged hourly 2-m
 242 temperature simulated by WRF.

243 3.1.2. Limitations and Uncertainties associated with soil NO emission estimation

244 Although the current BDSNP algorithm is considered more sophisticated than the old YL95
 245 algorithm, it still suffers certain limitations. For example, the current BDSNP
 246 parameterization employs a static classification of “arid” versus “non-arid” soils, upon which
 247 the relationship between soil NO emissions and soil moisture relies (Hudman et al., 2012).
 248 However, recent studies (Sha et al., 2021; Huber et al., 2023) have shown more dynamic
 249 representation of this classification is needed to capture the emission characteristics as
 250 observed by many chamber and atmospheric studies (e.g., Oikawa et al. 2015; Huang et al.
 251 2022a). Huber et al. (2023) also showed that the emission estimated based on the static
 252 classification are very sensitive to the soil moisture and thus could not produce self-consistent
 253 results when using different soil moisture products.

254 In addition to the aforementioned limitation, the estimated soil NO emissions are also
 255 subjected to certain limitations and large uncertainties. The first uncertainty comes from the

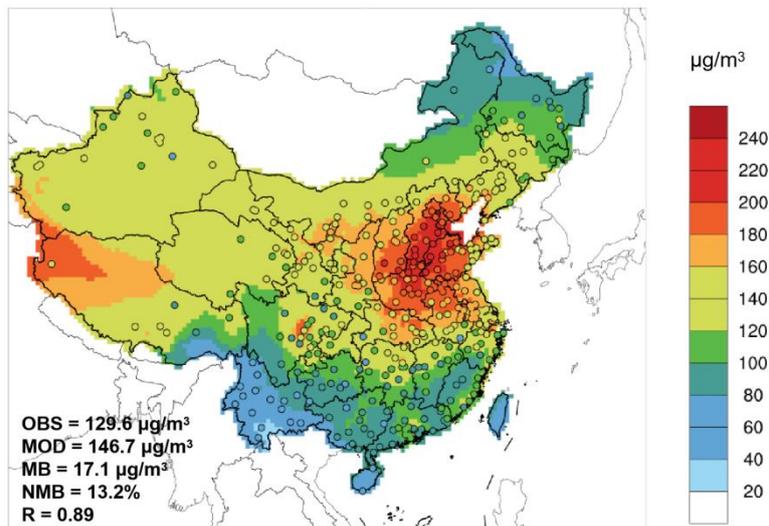
256 amount of fertilizer application, which has been identified as the dominant contributor to soil
257 NO emissions, as mentioned above. According to the global dataset (Potter et al., 2010), the
258 amount of fertilizer applied is 19.6 Tg, which is comparable to the sum of nitrogen fertilizer
259 for 2018 (20.7 Tg) obtained from provincial statistical yearbooks. However, compound
260 fertilizer, usually with a nitrogen, phosphorus, and potassium ratio of 15: 15: 15, has been
261 used more in China. Since 2016, the amount of nitrogen fertilizer has been decreasing
262 annually at an average rate of 4.6%, while the amount of compound fertilizer has been
263 increasing since 2010 at an average rate of 3.3%. The ratio of compound fertilizer to nitrogen
264 fertilizer has increased from 76.4% in 2010 to 109.8% in 2018. Consequently, soil NO
265 emissions may be largely underestimated if the compound fertilizer is not taken into account.
266 Our calculation shows that if only nitrogen fertilizer is considered, the estimated total soil NO
267 emissions are 805.2 Gg N/a for 2018, which is comparable to the value (770 Gg N/a averaged
268 during 2008-2017) reported by Lu et al. (2021), but 30.5% lower than that based on both
269 nitrogen fertilizer and compound fertilizer. Regionally, this underestimation ranges from
270 11.1%~41.5%, with a larger underestimation in Central China and East China (Fig. S3).

271 Another major uncertainty in estimating soil NO emissions is the temperature dependence
272 factor $f(T)$ in Eq.1. According to the BDSNP scheme, soil NO emissions increase
273 exponentially with temperature between 0 and 30°C and reach a maximum when the
274 temperature exceeds 30°C. The default temperature dependence coefficient (i.e., k in Eq. S2)
275 follows the value used in the YL95 scheme, which is 0.103 ± 0.04 . However, as shown by
276 Table 3 in Yienger and Levy (1995), this value is the weighted average of values reported for
277 different land types, which shows a wide range from 0.040 to 0.189. Even for the same crop
278 type (e.g., corn), the value of k could be quite different (0.130 vs. 0.066). We conducted a
279 sensitivity analysis to examine the impact of varying the k value on estimated soil NO
280 emissions. When the k value decreases or increases by 20%, the estimated total soil NO
281 emissions change from 715.7 to 1902.6 Gg N/a, representing a relative difference of -
282 38.2~64.3% deviation from the default value (1157.9 Gg N/a). Using the default k value
283 would result in a large overestimation of simulated NO₂ concentrations over NCP and YRD
284 and underestimation over Northeast China (Fig. S4). According to the total sown areas of
285 farm crops reported in the provincial statistical yearbook, the primary crops grown in these
286 regions are wheat and corn, which have a relatively low k value (0.066~0.073). Therefore, we
287 adjusted k for NCP (reduced by 20%), YRD (reduced by 10%), and Northeast China
288 (increased by 10%). CAMx simulation results show that this adjustment would not
289 significantly affect the simulated MDA8 O₃ concentration but could reduce the NO₂ gap
290 between observation and simulation (Fig. S4-S5). Therefore, we applied this adjustment to
291 soil NO emissions in the following CAMx simulations.

292 3.2. Contribution of soil NO emissions to ground-level ozone

293 3.2.1. Base case model evaluation

294 Fig. 4 shows the monthly averaged MDA8 ozone concentration simulated for June 2018 with
295 observed values presented on top. Overall the model well captured the spatial distribution of
296 MDA8 with a spatial correlation $R = 0.89$. Over the 365 cities in China, the simulated
297 monthly averaged MDA8 ozone concentration is $146.7 \pm 36.1 \mu\text{g}/\text{m}^3$, which is slightly higher
298 than the observed value of $129.6 \pm 37.6 \mu\text{g}/\text{m}^3$ (NMB = 13.2%). Regionally, model shows
299 better performance in Northeast China (MB = $2.4 \mu\text{g}/\text{m}^3$, NMB = 1.9%) and NCP (MB = 13.3
300 $\mu\text{g}/\text{m}^3$, NMB = 7.7%). Over-prediction is observed for Sichuan Basin and YRD (Table S6).
301 Simulated ozone concentration over the northwest Qinghai-Tibet Plateau was also much
302 higher than observed values. Our OSAT results (shown later) show that the high ozone
303 concentration over the Qinghai-Tibet Plateau is mostly contributed by the transport of
304 boundary ozone, which includes both horizontal and vertical (i.e., stratosphere) directions. For
305 regions with high altitude (e.g., the Qinghai-Tibet Plateau), vertical ozone intrusion from the
306 stratosphere is most substantial, which is consistent with the finding by Chen et al. (2023) that
307 the boundary layer height was identified as the most important feature for ozone over the
308 Qinghai-Tibet Plateau.



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310 **Figure 4.** Comparison of simulated (base colors) and observed (scatter points) values of
311 MDA8 ozone in China in June 2018.

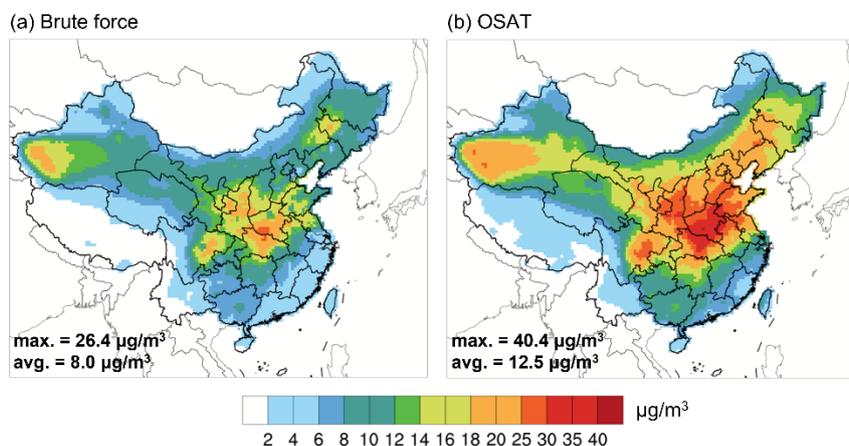
312 3.2.2. Impacts on regional ozone

313 To assess the contribution of soil NO emissions to surface ozone, both the brute-force method
314 (BFM) and the OSAT method were applied, and the results are shown in Fig. 5. Generally, the
315 two methods show consistent ozone contribution from soil NO emissions but with different
316 magnitudes. The BFM method shows widespread ozone enhancement due to soil NO

317 emissions with a spatial pattern that aligns with the distribution of soil NO emissions.
318 Substantial ozone enhancement is found over Central China, Sichuan Basin, northern YRD,
319 and eastern Northeast China. Maximum ozone enhancement (ΔMDA8) due to soil NO
320 emissions is $26.4 \mu\text{g}/\text{m}^3$ with a domain-average value of $8.0 \mu\text{g}/\text{m}^3$. For selected key regions,
321 the ozone contribution ranges from low to high: PRD ($3.8 \pm 1.1 \mu\text{g}/\text{m}^3$), YRD (8.7 ± 4.7
322 $\mu\text{g}/\text{m}^3$), Sichuan Basin ($9.1 \pm 0.9 \mu\text{g}/\text{m}^3$), Northeast ($9.3 \pm 3.0 \mu\text{g}/\text{m}^3$), and NCP (13.9 ± 4.4
323 $\mu\text{g}/\text{m}^3$), respectively. A similar spatial pattern is observed with the OSAT results, but the
324 magnitudes are much higher. Maximum ozone contribution by soil NO emissions reaches
325 $40.4 \mu\text{g}/\text{m}^3$ according to OSAT results, which is 53.0% higher than the brute force method.
326 The corresponding ozone contribution for each selected region is $6.7 \pm 1.2 \mu\text{g}/\text{m}^3$ (PRD), 13.5
327 $\pm 7.4 \mu\text{g}/\text{m}^3$ (Sichuan Basin), $14.5 \pm 4.9 \mu\text{g}/\text{m}^3$ (Northeast China), $16.2 \pm 7.8 \mu\text{g}/\text{m}^3$ (YRD)
328 and $25.7 \pm 5.3 \mu\text{g}/\text{m}^3$ (NCP). The scatter plots between BFM and OSAT results show good
329 correlations (Fig. S6, $R^2 = 0.78\text{--}0.97$), with OSAT results higher by 10%~61%. For YRD,
330 Sichuan Basin, and Northeast, the difference between the OSAT method and BFM increases
331 with the absolute ozone concentration (Fig. S7), while NCP shows the opposite trend. The
332 difference between the two methods reflects the nonlinear ozone response to NO_x emissions.
333 This nonlinearity becomes stronger in regions with larger NO_x concentrations, especially
334 where O₃ production is characterized as NO_x-saturated (or VOC-limited), such as the NCP. In
335 such cases, removing a portion of the NO emissions (e.g., zeroing out soil NO for the BFM
336 simulation) makes O₃ production from the remaining NO emissions more efficient, which
337 lessens the O₃ response. As shown later in Figure 7a, the O₃ response for NCP is more curved
338 (nonlinear) than other regions, consistent with NCP tending to have more NO_x-saturated O₃
339 production. This nonlinear effect also explains smaller O₃ attribution to soil NO by the BFM
340 than OSAT, especially over the NCP. Attributing a secondary pollutant to a primary emission
341 (e.g., O₃ to NO) is inherently tricky with nonlinear chemistry, as Koo et al. (2009) discussed.
342 Therefore, it is useful to present estimates from different methods. The Path Integral Method
343 (PIM) is a source apportionment method that explicitly treats nonlinear responses with
344 mathematical rigor (Dunker et al., 2015). However, applying the PIM is more costly than the
345 BFM or OSAT.

346 In addition to soil NO contribution, OSAT also gives ozone contributions from other source
347 groups, including anthropogenic emissions within China, boundary contribution, natural
348 emissions (e.g., biogenic emissions, open biomass burning, lightning NO_x), and emissions
349 outside China. The spatial distribution for each source category is presented in Fig. S8, and
350 the relative contribution for each selected region is shown in Fig. S9. Overall, boundary
351 transport (56.5%) and anthropogenic emissions (24.0%) contribute most to MDA8 ozone for
352 June 2018. Boundary contribution is high over the western and northern parts of China, while
353 the contribution from anthropogenic emissions is substantial over eastern China, where

354 anthropogenic emissions are extensive. On a national scale, soil NO emissions exhibit a
 355 relative ozone contribution of 9.1%, and regionally this value ranges from 6.1% in PRD to
 356 13.8% in NCP.



357 **Figure 5.** Ozone contribution from soil NO emissions based on (a) brute force method and (b)
 358 OSAT method.
 359

360 We further evaluated the impact of soil NO emissions on the number of ozone exceedances
 361 days (i.e., days with MDA8 O₃ higher than 160 µg/m³) during June 2018 based on the relative
 362 response factor (RRF) method and results from the brute force method. The total number of
 363 ozone exceedances days during June 2018 for the five selected regions ranged from 50 days
 364 in PRD to 985 days in NCP (Table 1). The number of ozone exceedance days per city ranged
 365 from 3.1 days in Sichuan Basin to 18.2 days in NCP, suggesting the severe ozone pollution in
 366 June 2018 over NCP. RRF was first calculated for each city as the ratio of simulated ozone
 367 concentration between the base case and the case with soil NO emissions excluded and
 368 applied to the observed ozone concentrations to obtain adjusted ozone concentrations without
 369 soil NO emissions. Soil NO emissions are estimated to lead to 121 ozone exceedance days in
 370 NCP, followed by 84 days in the Northeast and 70 days in YRD, corresponding to a percent
 371 change of 12.3%, 32.8%, and 10.5%, respectively. In Sichuan Basin, where soil NO emissions
 372 are also substantial, soil NO emissions contribute 30 ozone exceedances days, which accounts
 373 for 43.5% of the total ozone exceedances days. These results suggest the substantial
 374 contribution of soil NO emissions to the number of ozone pollution days over regions with
 375 high soil NO emissions.

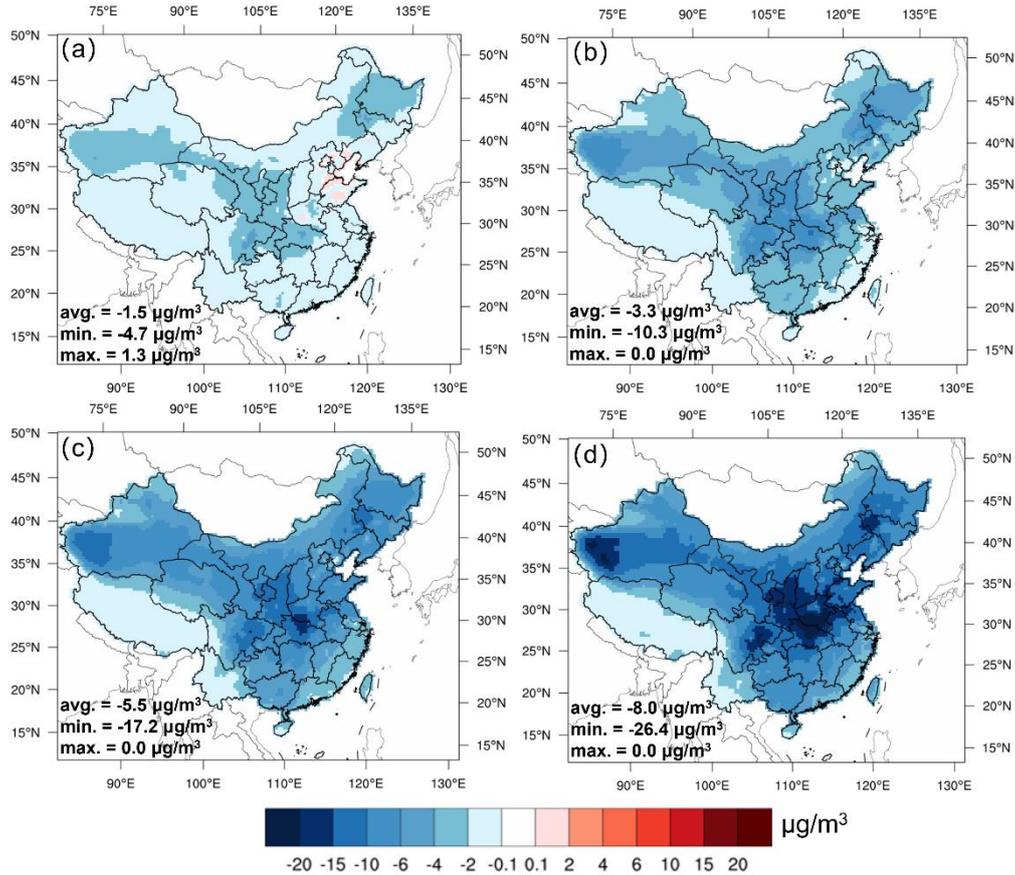
376 **Table 1.** Number of ozone exceedances over selected regions during June 2018.

Region (No. of cities)	Number of ozone exceedance days (% of total days)	Δozone exceedances days when soil NO emissions are removed	% of total ozone exceedances days
NCP (54)	985 (60.8%)	-121	-12.3%
YRD (55)	666 (41.1%)	-70	-10.5%

PRD (9)	50 (18.5%)	-6	-12.0%
Sichuan Basin (22)	69 (10.5%)	-30	-43.5%
Northeast (37)	256 (23.1%)	-84	-32.8%

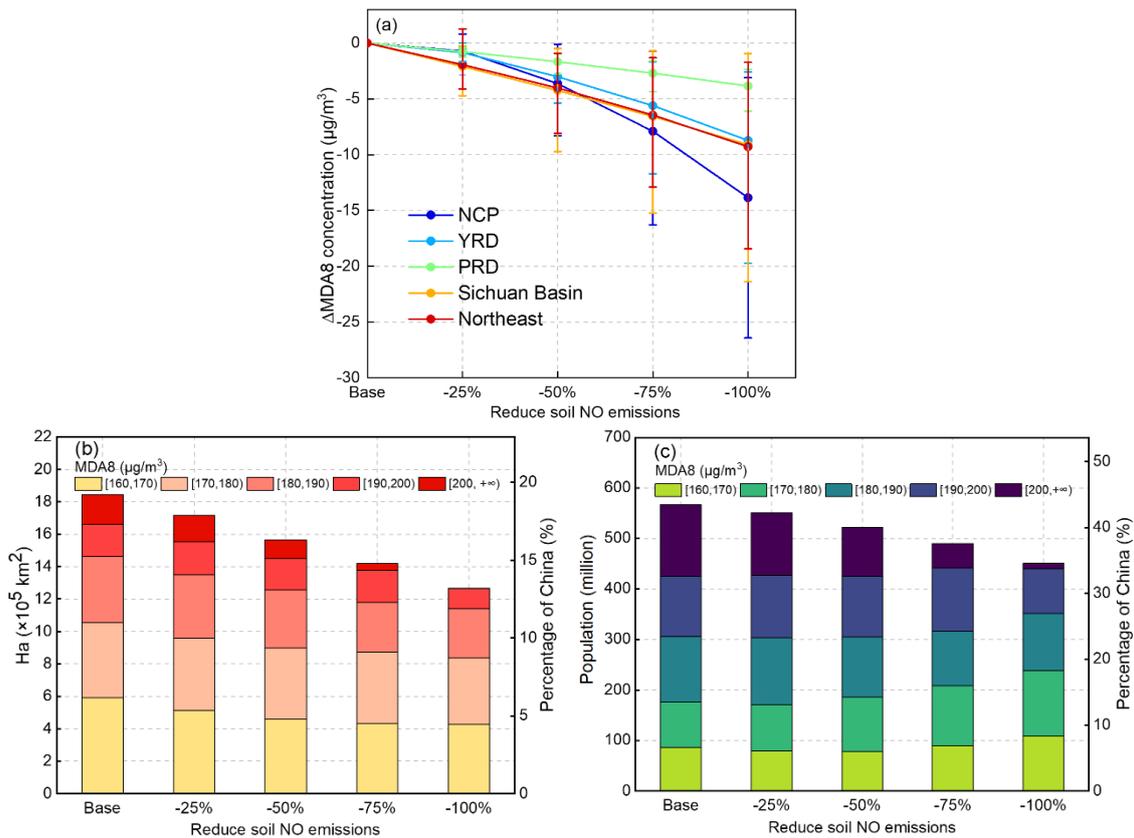
377 3.3. Ozone responses to reductions in soil NO emissions

378 Current NO_x emission control policies primarily target combustion sources, such as power
379 plants (Du et al., 2021) and on-road vehicles (Park et al., 2021). Nitrification inhibitors, such
380 as dicyandiamide (DCD, C₂H₄N₄), have been found to be effective in reducing nitrogen loss,
381 thereby reducing NO emissions from soil (Abalos et al., 2014). Studies have shown that using
382 5% DCD with nitrogen fertilizer can reduce NO emissions by up to 70% (Xue et al., 2022). In
383 light of this, it is important to evaluate the impact of reduced soil NO emissions on ozone
384 concentration. To address this question, four sensitivity simulations were carried out for June
385 2018, with soil NO emissions reduced by 25%, 50%, 75%, and 100% relative to the base
386 case. As shown by Fig. 6, reducing soil NO emissions led to a general decrease in monthly
387 MDA8 ozone concentration (Δ MDA8), with the magnitude of Δ MDA8 becoming more
388 significant with the reduction ratio. With a 25% reduction in soil NO emissions, there was a
389 widespread small decrease in monthly average MDA8 ozone concentration (Δ MDA8: -
390 $1.5 \pm 0.9 \mu\text{g}/\text{m}^3$), except over NCP where ozone showed a slight increase (up to $1.3 \mu\text{g}/\text{m}^3$) in
391 Shandong and Henan province. When soil NO emissions were cut by 50%, Δ MDA8 showed a
392 ubiquitous decrease across entire China with an average Δ MDA8 of $-5.5 \mu\text{g}/\text{m}^3$. When soil
393 NO emissions were removed entirely, the maximum Δ MDA8 could exceed $25 \mu\text{g}/\text{m}^3$ over
394 central China, part of the Sichuan Basin, Northeast China, and Northeast China. Regions with
395 strong ozone responses generally aligned with regions that also had high soil NO emissions.
396 However, it should be noted that the ozone response to soil NO reductions not only depends
397 on the magnitude of soil NO emissions but is also affected by (1) the local ozone formation
398 regime that is further determined by the relative abundance of NO_x and VOCs, and (2)
399 changes in transport of upwind ozone.



400
 401 **Figure 6.** Spatial distribution of Δ MDA8 under (a) 25%, (b) 50%, (c) 75%, and (d) 100%
 402 reductions of soil NO emissions in June 2018.

403 Fig. 7a provides further details on the domain-averaged Δ MDA8 under different reduction
 404 scenarios for the five key regions. As expected, the ozone response in each region increased as the
 405 reduction in the soil NO emissions increased. NCP exhibited the strongest ozone responses to
 406 changes in soil NO emissions, with Δ MDA8 increasing from $-0.7 \pm 0.8 \mu\text{g}/\text{m}^3$ with 25% reductions
 407 to $-13.9 \pm 4.4 \mu\text{g}/\text{m}^3$ when all soil NO emissions were removed. YRD, Sichuan Basin, and
 408 Northeast China exhibit similar ozone responses when soil NO emissions are reduced. Under the
 409 25% scenario, Δ MDA8 ranged from -4.7 to $1.3 \mu\text{g}/\text{m}^3$ for these three regions; with 100% soil NO
 410 reductions, Δ MDA8 ranged from -21.4 to $-0.9 \mu\text{g}/\text{m}^3$. Δ MDA8 in PRD was relatively small. Even
 411 with a 100% reduction, the average Δ MDA8 in PRD was less than $5 \mu\text{g}/\text{m}^3$, which is associated
 412 with the small soil NO emissions in PRD. It is interesting to note that all regions except NCP
 413 exhibited an approximate linear ozone response to changes in soil NO emission reductions. NCP
 414 showed more significant ozone reductions as the reduction ratio increased, suggesting that NCP
 415 would gain more benefits with more aggressive reductions in soil NO emissions compared to other
 416 regions.



417 **Figure 7.** (a) Δ MDA8 concentrations in five key regions under different emission reduction
 418 scenarios (b) Area and (c) population exposed to different ozone levels under different soil
 419 NO emission reduction scenarios.

420 We evaluated the impact of different soil NO emission reduction scenarios on the area and
 421 population exposed to varying ozone levels. The results, presented in Fig. 7b and 7c, revealed
 422 a decrease in coverage and exposed population under high ozone concentrations as soil NO
 423 emissions decrease. The data presented in the plots are for grid cells with monthly MDA8
 424 ozone concentrations exceeding 160 $\mu\text{g}/\text{m}^3$. In the Base scenario, the estimated coverage of
 425 MDA8 ozone exceeding 160 $\mu\text{g}/\text{m}^3$ was $1.84 \times 10^6 \text{ km}^2$, equivalent to 19.2% of the national
 426 land area. The population exposed to ozone concentrations exceeding 160 $\mu\text{g}/\text{m}^3$ amounts to
 427 566.6 million, representing 43.4% of the entire population. The areas with extremely high
 428 ozone concentrations (MDA8 > 200 $\mu\text{g}/\text{m}^3$) account for 1.9% of the national land area, with a
 429 corresponding exposed population of 10.9%, indicating that densely populated areas
 430 experience higher ozone concentrations. When soil NO emissions are halved, there is a 15.2%
 431 reduction in the coverage of non-attainment areas and an 8.0% reduction in the total exposed
 432 population. If soil NO emissions are eliminated, the total area coverage and population
 433 exposed to MDA8 ozone concentrations exceeding 160 $\mu\text{g}/\text{m}^3$ would be $1.27 \times 10^6 \text{ km}^2$ and
 434 450.3 million, respectively, representing 13.2% and 34.5% of the total. Compared to the Base
 435 scenario, a 100% theoretical reduction in soil NO emissions leads to a 31.3% and 20.5%

436 reduction in the exposed area and population under high ozone concentration, respectively,
437 indicating substantial health benefits gained when soil NO emissions are mitigated.
438 Fig. S10-S11 displays similar area and population plots for selected key regions. The overall
439 trends for each sub-region are consistent. With 100% reductions in soil NO emissions, the
440 area with high ozone concentration decreased by 17.8%, 22.3%, 65.4%, and 100% for NCP,
441 YRD, Sichuan Basin, and Northeast. The corresponding values for the exposed population are
442 91.4%, 60.3%, 9.8%, and 0.0%. While the relative change is more significant in Sichuan
443 Basin and Northeast China, NCP and YRD gain more health benefits due to the significantly
444 higher total population for these two regions. However, it is worth noting that even with the
445 complete elimination of soil NO emissions, a total of 450.3 million people are still exposed to
446 ozone levels exceeding the national standard, necessitating multiple control policies at the
447 same time, such as synergistic control of anthropogenic VOC emissions (Chen et al., 2022;
448 Ding et al., 2021).

449 3.4 Comparison with existing studies

450 The soil NO emissions estimated in this study were also compared with values reported by
451 existing studies based on either field measurement or model estimation (Table S7). Previous
452 studies report a wide range of soil NO emissions from 480 to 1375 Gg N and soil NO flux
453 ranging from 10 to 47.5 ng N m⁻² s⁻¹. The soil NO emissions estimated in our study are 1157.9
454 Gg N with the default *k* value and 951.9 Gg N with region-adjusted *k* value, which falls
455 within the upper range of previously reported values. The averaged soil NO flux over NCP in
456 June 2018 estimated in our study is 35.4 ng N m⁻² s⁻¹, which is within the range reported by
457 previous studies (12.9~40.0 ng N m⁻² s⁻¹).

458 The simulated ozone contribution by soil NO emissions is compared with other studies. In
459 California, soil NO was estimated to cause a 23.0% increase in surface O₃ concentrations (Sha
460 et al., 2021). Constrained by satellite measured NO₂ column densities, Wang et al. (2022b)
461 reported MDA8 ozone contribution of 9.0 μg/m³ (relative contribution of 5.4%) from
462 cropland NO_x emissions over NCP during a growing season in 2020. Lu et al. (2021) showed
463 an interactional effect of domestic anthropogenic emissions with soil NO emissions of 9.5 ppb
464 in the NCP during July 2017. In addition, soil NO_x emissions strongly affect the sensitivity of
465 ozone concentrations to anthropogenic sources in the NCP. In a most recent study by Shen et
466 al. (2023), addition of the soil NO_x emissions was shown to result in up to 15 ppb increase of
467 ozone concentration over Xinjiang, Tibet, Inner Mongolia, and Heilongjiang, although a
468 minor reduction was evident over the Yangtze River basin. When soil NO_x emissions were
469 reduced by 30%, ozone concentrations increased by 3-5 ppb over Inner Mongolia,
470 Heilongjiang, Xinjiang, and Tibet, while decreased by 0-2 ppb over the Yangtze River basin.
471 Surprisingly, when soil NO_x emissions were increased by 30%, nearly identical ozone

472 responses were observed.

473 **4. Conclusions**

474 Soil NO emissions are non-negligible NO_x sources, particularly during summer. The
475 importance of soil NO emissions to ground-level ozone concentration in China is much less
476 evaluated than combustion NO_x emissions. In this study, the total national soil NO emissions
477 were estimated to be 1157.9 Gg N in 2018, with a spatial distribution closely following that of
478 fertilizer application. High soil NO emissions were mainly concentrated over Henan,
479 Shandong, and Hebei provinces, which differs from anthropogenic NO_x emissions. Distinct
480 diurnal and seasonal variations in soil NO emissions were simulated, mainly driven by the
481 changes in temperature as well as the timing of fertilizer application. Uncertainty analysis
482 reveals a range of 715.7~1902.6 Gg N of soil NO emissions that warrant further constraints
483 from observations.

484 Using two methods (BFM and OSAT), we evaluated the contribution of soil NO emissions to
485 ground-level ozone concentration for June 2018. Both methods suggest a substantial
486 contribution of soil NO emissions to MDA8 ozone concentrations by 8~12.5 μg/m³ on
487 average for June 2018, with the OSAT results consistently higher than BFM. Soil NO
488 emissions were shown to lead to a relative increase of ozone exceedances days by
489 10.0%~43.5% for selected regions. Reducing soil NO emissions could generally reduce the
490 ground-level ozone concentrations and populations exposed to unhealthy ozone levels
491 (MDA8 > 160 μg/m³), especially over NCP and YRD. With a 50% reduction in soil NO
492 emissions, the coverage of non-attainment areas and the population exposed to unhealthy
493 ozone levels decreased by 15.2% and 8.0%, respectively. However, even with the complete
494 removal of soil NO emissions, approximately 450.3 million populations are still exposed to
495 unhealthy ozone levels, necessitating multiple control policies at the same time, such as
496 synergistic control of anthropogenic VOC emissions.

497 **Data availability.** Data will be made available on request.

498 **Author contributions.** **Ling Huang:** Conceptualization, Formal analysis, Writing – original
499 draft. **Jiong Fang:** Data curation, Formal analysis, Visualization. **Jiaqiang Liao:** Data
500 curation, Formal analysis, Visualization. **Greg Yarwood:** Writing – review & editing. **Hui**
501 **Chen:** Writing – review & editing. **Yangjun Wang:** Writing – review & editing. **Li Li:**
502 Conceptualization, Supervision, Funding acquisition, Writing – review & editing.

503 **Competing interests.** The authors declare that they have no known competing financial
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