

Cyclones enhance the transport of sea spray aerosols to the high atmosphere in the Southern Ocean

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Abstract: Cyclones are expected to increase the vertical transport of sea spray aerosols (SSAs), which may significantly impact the climate by increasing cloud condensation nuclei (CCN) and cloud droplets (N_d) population.

In this study, a high-time resolution (1h) aerosol monitoring was carried out in the mid- and high Southern Hemisphere from 23th February 2018 to 4th March 2018. The characteristics of SSAs during three cyclones were observed in the cruise. The results showed that SSAs level in the low atmosphere didn't increase with the wind speed during cyclone processes, which was different from the anticipated scenario that SSAs concentration increased with wind speed. However, the size of SSA particles during cyclones was larger than that in the no-cyclone periods. It seems that the generation of SSAs was enhanced during cyclones, but SSAs concentration near the sea surface increased scarcely. The upward transport proportion was calculated according to the wind stress and sea-salt flux between cyclone and non-cyclone periods. It indicated that more than 23.4% of the SSAs were transported upward by cyclone processes during Event 1, and 36.2% and 38.9% in Event 2 and Event 3, respectively. The upward transport of SSAs was the main reason why SSAs concentration didn't increase in the low atmosphere. The transport of SSAs to the high atmosphere during cyclones may increase the CCN burden additionally in the marine boundary layer, which may affect the regional climate. This study highlights the importance of SSAs transport to the high atmosphere by cyclone and extends the knowledge of SSAs generation and impact factor during the cyclone period in marine atmosphere.

Keywords: Sea spray aerosols (SSAs); Cyclone; Southern Ocean (SO); Transport

29 **1. Introduction**

30 Sea spray aerosols (SSAs) were one of the largest sources of primary aerosols in the marine
31 atmosphere, making a significant contribution to aerosols in the marine atmosphere (McInnes et al.,
32 1996). It is reported that the annual global SSAs flux was estimated to be $1.01 \times 10^4 \text{Tg yr}^{-1}$ (Gong
33 et al., 2002). Pure sea salt mostly consisted of NaCl and a mixture of one or more other salts, such
34 as Mg, K, Ca, sulfates and traces of organic materials (Thomas et al., 2022). SSAs were considered
35 to be the most important contributor to aerosol light scattering in the marine boundary layer (MBL)
36 (Quinn and Coffman, 1999; Takemura et al., 2002). In addition, as the source of cloud condensation
37 nuclei (CCN), SSAs can alter solar radiation reflection, extent the lifetime of clouds and further
38 impact the global climate (Pierce and Adams, 2006). The influence of SSAs on cloud properties was
39 thought to be particularly intense over remote ocean region devoid of continental particles. Studies
40 have found that The largest SSA CCN number fractions, up to 65%, were observed in the high
41 southern latitudes (40° S to 70° S) at low supersaturation (0.1%) (Quinn et al., 2017).

42 SSAs are generated predominantly by the action of wind on the ocean (Stokes et al., 2013) as
43 the major mechanism of SSAs production is bubble bursting at sea surface as a result of wind stress
44 (Monahan and Muirheartaigh, 1980). Wind stress on the sea surface forms waves. Then bubbles
45 are generated and return to the sea surface, creating whitecaps. Subsequently, the bubble bursts and
46 jets droplets into the atmosphere. Hence, the sea salt particle size and concentration of SSAs was
47 significantly depended on the wind speed (McDonald et al., 1982). However, some studies have
48 showed that wind speed was not the sole impact factor of SSAs production, as humidity, temperature
49 and sea-air temperature would also impact SSAs generation (Cole et al., 2003; Shi et al., 2022; Liu
50 et al., 2020). The generation of SSAs in the marine atmosphere have been investigated in previous

51 studies, but the production of SSAs during extreme weather (such as cyclones) in the mid- and high
52 Southern Hemisphere is still lack of knowledge.

53 Westerlies in the Southern Hemisphere fundamentally control regional patterns of air
54 temperature and also regulate ocean circulation, heat transport and carbon uptake (Goyal et al.,
55 2021). Moreover, the zone of westerlies is prone to cyclones which dominate the precipitation
56 pattern of the mid- and high Southern Hemisphere (Mycoy et al., 2020). Southern Ocean (SO) plays
57 an important role in global carbon cycles and climate changes (Gruber et al., 2019). Furthermore,
58 the SO is scarcely affected by human activities, in which the influence of SSAs on CCN is
59 particularly strong in this region.

60 Cyclones may carry large water volumes and impose strong winds, which have a significant
61 impact on marine aerosols, especially on SSAs. (Fang et al., 2009). Air convergence due to the
62 reduction of pressure caused by cyclones may also affect SSAs concentration. Typically, higher
63 frequency of cyclone was developed during the summer season than in other seasons. It is reported
64 that 959 cyclones occurred in the SO during summer time from 2004 to 2008 (Liu et al., 2012).

65 In summary, the impact of cyclones on the emission of SSAs can't be ignored. SSAs can direct
66 absorbing and scattering of solar radiation. Additionally, sea spray aerosol is an important source of
67 CCN, which plays a significant role in regulating global warming, but it remains unclear that how
68 does the cyclone impact SSAs emission. Cyclones developing in the westerlies and SO result in the
69 decrease of pressure, air convergence, strong winds and heavy precipitation, which alter the
70 emission of SSAs and thus affects regional climate in the mid- and high latitudes of the Southern
71 Hemisphere. However, the lack of direct observations makes it challenge to deep insight into this
72 question. Generally, the observation of cyclones is commonly performed at fixed points on land

73 (Badarinath et al., 2008), but such observations cannot be used to investigate the effect of cyclones
74 on SSAs over remote ocean regions. However, high-time resolution observation technology is now
75 available on research vessels to carried out the SSAs monitoring and understand the SSAs behavior
76 during cyclone processes.

77 In this study, SSAs characteristics were observed with high-time resolution during three
78 cyclones in the SO to determine the transport of SSAs to the high atmosphere by cyclones. The
79 concentration and particle size of SSAs were measured simultaneously for the first time with high-
80 time resolution (1h) in the mid- and high Southern Hemisphere from 23th February to 4th March
81 2018. The results provide a new insight into the effect of cyclones on the generation and vertical
82 transport of SSAs in the mid- and high latitudes of the Southern Hemisphere.

83 **2. Methodology**

84 **2.1 Observational sites**

85 Observations were carried out on board the R/V “Xuelong” during the 34th Chinese Antarctica
86 Expedition Research Cruise from 23th February and 4th March 2018. The observation covered with
87 a large portion of the SO (40°S to 73°S, 170°E to 124°W, seen in Fig. S1).

88 **2.2 SSAs measurement**

89 Aerosol composition was monitored with a temporal resolution of 1 h using an in-situ gas and
90 aerosol composition monitoring system (IGAC, Model S-611 <http://www.machine-shop.com.tw/>).
91 To minimize the impact of ship emissions, the sampling inlet connected to the monitoring
92 instruments was fixed on a mast (20 m above the sea surface) located at the bow of the research
93 vessel. Note that the major pollutants were from the chimney, which is located at the stern of the

94 R/V and about 25 meters above the sea level. Hence, the pollution emissions from the vessel mainly
95 located at the downwind of the sampling inlet, especially when the vessel is running forward. As
96 high-time-resolution observations were used in this study, the self-contaminations from the vessel
97 have been eliminated from the measurement results. The wind speed and wind directions were also
98 monitoring during the observation period, which were used to determine if the observations were
99 affected by the self-contaminations or not. A total suspended particulate sample inlet was also
100 positioned at the top of the mast. All aerosol observational instruments were connected by
101 conductive silicone tubing with an inner diameter of 1.0 cm.

102 The IGAC monitoring system consisted of three main units, including a Wet Annular Denuder
103 (WAD), a Scrub and Impact Aerosol Collector (SIAC) and an ion chromatograph with a sampling
104 flow of 16.7 LPM. The collection of acidic and basic gases relies on the diffusion and absorption of
105 gases into a downward flowing aqueous solution. The SIAC was positioned at an angle to facilitate
106 the collection of enlarged particles. Ultrapure water was fed continuously into the nozzle at 1.2
107 mL.min⁻¹ and heated to 140 °C to generate stream, which was sprayed directly towards the particle-
108 laden air to improve the humidity of flue gases. Fine particles were enlarged and subsequently
109 accelerated through a conical-shaped impaction nozzle and collected on the impaction plate. The
110 gas and aerosol liquid samples from the WAD and SIAC were drawn separately by a pair of syringe
111 pumps. The samples were then analyzed for anions and cations by an online ion chromatography
112 (IC) system (Dionex ICS-3000). The injection loop size was 500µL for both anions and cations
113 (Young et al., 2016). Six to eight concentrations of standard solutions were used for calibration
114 purposes, depending on the target concentration (R^2 values above 0.997). The detection limits for
115 Na⁺ concentration was 0.03 µg. L⁻¹ (aqueous solution).

116 2.3 SSA particle size measurement

117 A single particle mass spectrometer (SPAMS) was used to measure the SSAs particle size
118 distribution. A nafion tube dryer was placed at the inlet of SPAMS to remove the moisture of
119 sampling gas. Details of the methods used for aerosol detection and the operational procedure for
120 the on-board SPAMS have been described carefully in the previous study (Li et al., 2014). The
121 performance of particle size distribution determination using SPAMS has been confirm (Yan et al.,
122 2016; Li et al., 2014; Ma et al; 2016). A PM_{2.5} collector was deployed to remove particles larger
123 than 2.5 μm . Fine particles were drawn into the vacuum system through a critical orifice and then
124 accelerated and focused to form a particle beam. Particles with specific velocities then passed
125 through two Nd: YAG lasers (532 nm). The aerodynamic diameter of single particle was calculated
126 by the particle velocity. The particle size detected by the SPAMS was calibrated using polystyrene
127 latex spheres (PSL, Duke Scientific Corp., Palo Alto) with diameter of 0.2, 0.3, 0.5, 0.75, 1.0, 2.0,
128 and 2.5 μm (Li et al., 2011).

129 2.4 Meteorological parameters

130 Meteorological parameters such as wind speed (WS), wind direction and temperature etc. were
131 measured continuously using an automated meteorological station mounted on the R/V "Xuelong".
132 Weather map data, including sea surface pressure and total precipitation, was obtained from the fifth
133 generation ECMWF reanalysis for the global climate and weather (ERA5.
134 <https://cds.climate.copernicus.eu/>). Satellite cloud maps were obtained from the Level-1 and
135 Atmosphere Archive and Distribution System Distributed Active Archive Center (LAADS DAAC
136 data product MOD021KM. <https://ladsweb.modaps.eosdis.nasa.gov/>).

137 **2.5 Undisturbed SSA concentration estimates during the cyclone period**

138 Undisturbed SSAs (U-SSAs) is defined as the SSAs generation supposed without cyclone
139 impact in the marine boundary layer during the cyclone process, which was determined by the wind
140 stress and sea-salt flux. The upward transport proportion of SSA was estimated by comparing U-
141 SSA concentration with the concentration of SSAs during cyclone period. U-SSA concentration
142 during the cyclone period were estimated in two ways as follow.

143 The momentum flux at the air-sea interface, also called wind stress, is an important part of the
144 interaction between ocean and atmosphere, which reflects the friction and drag effect between the
145 two fluids. Wind stress is the energy source of SSAs generation. The momentum flux at the air-sea
146 interface can be calculated using the following equation (Toffoli et al., 2012):

147
$$\tau = \rho_a C_d U_{10}^2 \quad (1)$$

148 Where, ρ_a is the air density, U_{10} is the wind speed measured at 10 m above the sea surface, and
149 C_d is a drag coefficient, which can be expressed as follow:

150
$$C_d = (a + bU_{10}) \times 10^{-3} \quad (2)$$

151 Where, a is 0.96 and b is 0.06.

152 According to the difference of wind stress between cyclonic and non-cyclonic periods,
153 combining with the concentration of SSAs during non-cyclonic periods, U-SSA (wind stress)
154 concentration can be obtained.

155
$$U\text{-SSA}_{(\text{wind stress})} = \frac{\tau_{\text{cy}}}{\tau_{\text{non-cy}}} * \text{SSA}_{(\text{non-cy})} \quad (3)$$

156 For the indirect production of SSAs through the formation and bursting of bubbles, the SSAs
157 flux function dF_0/d_r (particles $\text{m}^{-2} \text{s}^{-1} \text{mm}^{-1}$) which expresses the rate of sea water droplet generation
158 per unit area of sea surface per increment of particle radius, is given by Monahan et al. (1986) as

159 Eq. (4):

$$160 \quad \text{SSA flux} = \frac{dF_0}{dr} = 1.373 U_{10}^{3.41} r^{-3} (1 + 0.057 r^{1.05}) \times 10^{1.19e^{-B^2}} \quad (4)$$

161 Where, $B = (0.38 - \log r) / 0.65$, r is the particle radius.

162 Then, U-SSA_(Sea-salt flux) concentration can be obtained as follow:

$$163 \quad \text{U-SSA}_{(\text{Sea-salt flux})} = \frac{\text{SSA flux}_{\text{cy}}}{\text{SSA flux}_{\text{non-cy}}} * \text{SSA}_{(\text{non-cy})} \quad (5)$$

164 3. Results and discussion

165 3.1 Meteorology and cyclone events

166 The observation region in the SO was defined by the outermost closed isobar surrounding the
167 cyclone area center (Wernli and Schwierz, 2006). Rainfall has numerous aspects impact of SSAs
168 production. Generally, raindrops falling onto the sea surface can produce SSA particles directly or
169 indirectly, either from bubbles entrained by the drop or SSA particles produced by the splashed
170 drops (Blanchard and Woodcock, 1957). However, raindrops can also function as efficient scavenger
171 of particles in the atmosphere (Lewis and Schwartz, 2004). Hence, the precipitation period was
172 extracted, when the transport of SSAs by cyclones was discussed in this study. As known that
173 relative humidity also has an impact on SSAs, high relative humidity was presented in this study,
174 which basically reached the deliquescence point (about RH:75%) of NaCl (Cole et al, 2003). In this
175 case, the change of relative humidity has little effect on the particle size. Three cyclone events were
176 observed during the cruise (Fig. S2). Na⁺ derived from SSAs is an important component of marine
177 atmospheric aerosols (Teinila et al., 2014) and is generally considered to be a marker of SSAs in the
178 marine atmosphere (Yeatman et al., 2001). Hence, the relationship between Na⁺ concentrations and
179 meteorological factors were discussed in this study, seen in Fig. 1.

180 The first cyclone was generated in the mid- Southern Hemisphere (45°S, 150°E), and gradually

181 moved to eastwards (Fig. 2). As the cyclone approached, the R/V “Xuelong” sampled a northwest
182 warm and humid air mass followed by precipitation. As the research vessel entered the cyclone area
183 (Event 1. shadow area in Fig. 1) at about 15:00 24/2/18 (UTC time), air pressure suddenly dropped
184 from 1003 hpa to 961 hpa and wind speed significantly enhanced, comparing with the non-cyclone
185 area (average wind speed increased from 11.7 m s⁻¹ to 14.8 m s⁻¹). However, the average Na⁺
186 concentration during this cyclone event remained relatively constant as the WS increased, changing
187 from 1529 ng m⁻³ to 1706 ng m⁻³. At about 23:00 25/2/18, the research vessel left the cyclone area.
188 Note that wind speed dropped sharply between 13:00 and 23:00 25/2/18 (average wind speed
189 decreased from 14.8 ms⁻¹ to 9.3 ms⁻¹) and this was matched by a rapid decrease in SSAs
190 concentration (from 1706 ng m⁻³ to 343 ng m⁻³, seen in Fig. 1b).

191 The vessel encountered another cyclone area at 10:00 26/2/18 and immediately turned to the
192 southeast, leaving the cyclone area at 22:00 26/2/18 (Event 2, seen in Fig. 1). During Event 2, the
193 research vessel did not pass through the center of the cyclone. However, it was also affected by the
194 cyclone, as the atmospheric pressure dropped from 983 hpa to 973 hpa and the average wind speed
195 increased from 13.5 m s⁻¹ to 15.5 m s⁻¹. Similar to Event 1, the average Na⁺ concentration during
196 the cyclone period remained relatively constant, or even decreased from 2810 ng m⁻³ to 2354 ng m⁻³,
197 as the WS increased. During the event 2, the dominant air flow was cold and westerly thus there
198 was only a little precipitation (Fig. 2).

199 While the research vessel moved southeast and arrived at sea ice edge of the high SO, Na⁺
200 concentration was much lower than the value during the first two cyclone events, which suggested
201 that low air temperature and sea ice coverage reduced SSAs generation (Fig. S3) (Yan et al., 2020).
202 Between 18:00 1/3/18 and 04:00. 4/3/18, the research vessel encountered the third cyclone (Event

203 3). Wind speed increased from 7.5 m s^{-1} to 21.5 m s^{-1} and the air pressure dropped from 986 hpa to
204 960 hpa (lowest). Similarly, the average Na^+ concentration during this cyclone period showed little
205 increase (changing from 255 ng m^{-3} to 335 ng m^{-3}). The third cyclone was relatively stable and
206 moved slowly, but the cyclone only brought a small amount of precipitation in the wind shear region.

207 **3.2 SSA properties during cyclone processes**

208 Correlation coefficients between different compositions of sea spray aerosols in the atmospheric
209 were shown in Table S1. Na^+ correlated well with Mg^{2+} , K^+ , Ca^{2+} and SO_4^{2-} , implying Na^+ has a
210 good representation of SSAs. The variation of Na^+ concentrations in different latitude regions is
211 presented in Fig. S4. Positive correlations between Na^+ concentrations and wind speeds were found
212 in the low-middle latitudes ($20^\circ\text{S} - 40^\circ\text{S}$) ($R=0.59$, Fig.S4), where the atmospheric pressure
213 remained stable (Fig. S5). This suggests that that SSAs generation was greatly influenced by the
214 wind speed. However, the correlation between Na^+ concentrations and wind speed was relatively
215 low in middle-high latitudes ($40^\circ\text{S}-60^\circ\text{S}$) and in the polar region ($60^\circ\text{S}-74^\circ\text{S}$) ($R=0.45$ and 0.05 ,
216 respectively), where unstable atmosphere state or cyclone occurred frequency in these areas (Fig.
217 S5), suggesting that cyclone may affect the relationship between wind speed and SSAs
218 concentration in the marine atmosphere.

219 The relationship between WS and Na^+ concentration in different meteorological conditions are
220 illustrated in Fig. 3. To further investigate the influence of the cyclone on SSA concentrations in the
221 mid- and high Southern Hemisphere, a non-cyclone period (April 5th and 6th) with stable pressure
222 and relative humidity but without precipitation, was selected as a control period (defined as normal
223 period).

224 It is readily apparent that Na^+ concentrations and SSAs increased with the WS during the
225 control period (Fig. 3a, $R = 0.74$). Positive correlations between Na^+ concentrations and WS were
226 also presented during non-cyclone effects in event 1, event 2 and event 3 ($R = 0.65, 0.64$ and 0.50 ,
227 respectively, seen in Fig. 3b, 3c and 3d), which was in good agreement with previous study (O'Dowd
228 and de Leeuw, 2007). It is worth noting that the correlation between Na^+ concentration and WS
229 during Event 3 were lower than the value during the other two cyclone events. This was caused by
230 the low temperature and sea ice coverage in the high SO, which weakened the influence of WS on
231 SSA generation (Yan et al., 2020).

232 In contrast, poor correlations between Na^+ concentration and wind speed were found during all
233 the three cyclone periods ($R = -0.32, 0.15$ and 0.44) and precipitation periods ($R = 0.08$ and -0.02 .
234 Fig. 3b, 3c and 3d). During the cyclone periods, Na^+ concentration changed irregularly as the WS
235 increased, suggesting that rainfall altered the effect of wind stress on SSA generation. The effect of
236 precipitation on the formation of SSA was complicated and WS may not be the critical factor that
237 affected SSAs emission during precipitation process. Further studies of how does precipitation affect
238 SSAs are required.

239 It is interesting that an obvious correlation between WS and Na^+ concentration was not presented
240 during cyclone process with high wind speed. Na^+ concentration during cyclone periods was even
241 lower than those during non-cyclone periods. That means the generation of SSAs did not enhanced
242 during the cyclone process or the SSAs was transported by the cyclone. The generation and transport
243 of SSAs during cyclone process was further discussed in the following section.

244 3.3 SSA particle size distribution

245 Generally, SSA generation increased with wind speed, however in this study it was found that
246 higher wind speed did not result in higher levels of SSAs during cyclone process. It seems that the
247 generation of SSAs was suppressed during cyclone. It is necessary to determine whether the
248 emission of SSAs in the cyclonic periods was higher than that in the non-cyclone periods. Feng et
249 al. (2017) and Liu et al. (2020) reported that both SSAs particle size and the concentration increased
250 with increasing wind speed. As the WS increased from 3.4 to 10 m s⁻¹, a 7–10 fold increase in
251 atmospheric sea salt concentration was observed. Log-normal distributions predicted a 30-fold
252 increase in the concentration (µg/m³) of particles larger than 1 x 10⁻⁹ g (10 µm radius) and a 50-fold
253 increase in the concentration of particles larger than 1 x 10⁻⁸ g (20 µm radius) (McDonald et al.,
254 1982). If the particle size of SSAs increased with increasing wind speed, it indirectly confirmed that
255 the concentration of SSAs also increase.

256 The size distributions of SSAs observed during the three cyclone events are presented in Table
257 S2 and Fig. 4. During Event 1, the difference between the number of SSA particles larger than 1.2
258 µm observed in cyclone and non-cyclone periods was about 11%. The change of SSA size
259 distribution during Event 2 and Event 3 were consistent with that during Event 1 (about 6% and 5%,
260 respectively). The mean size of SSA particles was larger during cyclone period than that during no-
261 cyclone period. These results revealed that cyclones in mid- and high Southern Hemisphere
262 enhanced SSAs generation. However, the increase of SSAs concentration was not presented as
263 expected when high wind speed occurred during cyclone period, suggesting that SSAs may be
264 transported or diluted in the lower atmosphere.

265 3.4 Estimation of the upward transport proportion of SSAs by cyclone

266 The mid- and high latitude of Southern Hemisphere, especially in the Antarctic region is one
267 of the most pristine in the world and serves as an important proxy for the pre-industrial atmosphere,
268 which was less affected by human activity. Hence, anthropogenic aerosols account for a small
269 proportion of the total aerosol population. In the SO, aerosols are typically derived from natural
270 sources, including primary particles (sea spray and bursting bubbles), which make up the vast
271 majority of the aerosol mass. In this region cyclones tend to occur in summer, generating more SSAs
272 due to high WS. The observation results suggested that air convergence caused by the cyclone may
273 result in considerable quantities of SSAs being transported vertically to high atmosphere, which can
274 partly explain why the mean number concentration of CCN/cloud droplets (N_d) in the SO in summer
275 is much higher than that in winter (Mycoy et al., 2020).

276 As mentioned above, the size of SSAs was larger during cyclone events than that in no-cyclone
277 period. However, the level of SSAs in the low atmosphere hardly increased with wind speed during
278 the cyclone process. It's likely that considerable SSAs were transported upward by air convergence
279 due to cyclone. When large number of SSAs were transported to the upper air, the SSAs in high
280 atmosphere enhanced solar radiation reflected back to space by modulating the N_d , which in turn
281 changed cloud reflectivity even without any changes of cloud macrostructure (Twomey, 1977).
282 Furthermore, cloud microphysical processes were also altered with changing CCN/ N_d (Albrecht,
283 1989). These two effects, summarized in Fig. 5, can ultimately affect the radiation balance of the
284 earth system in the mid- and high latitudes of Southern Hemisphere (Quinn and Bates, 2011). Thus
285 the effect of cyclone on SSAs generation, especially in the polar region, can't be neglected.

286 It is difficult to precisely estimate the proportion of SSAs transported vertically directly.

287 However, the differences of wind stress and sea-salt flux between cyclone and non-cyclone periods
288 can be used to calculate the undisturbed concentrations of Na⁺ (U-SSA concentration) during the
289 cyclone period. This can be used to quantify the upward transport proportion of SSAs.

290 Fig. S6 and S7 show the differences of wind stress and sea-salt flux between cyclone and non-
291 cyclone periods. The estimated proportion of vertically transport of SSAs, using the wind stress
292 method and the sea-salt flux method, are presented in Table 1. According to the calculation results,
293 more than 23.4% of the SSAs were transported upward by cyclone process during Event 1, and 36.2%
294 and 38.9% in Event 2 and Event 3, respectively. The upward transport proportion of SSAs estimated
295 using the bubble method were higher than those estimated using the wind stress method for all the
296 three cyclone events. As the research vessel was located at the high SO and close to the Antarctica
297 during Event 3, the upward transport proportion estimated using the bubble method was the highest,
298 reaching 56.6%, which was much higher than the results estimated for Event 1 (39.9%) and Event
299 2 (42.8%).

300 The high transportation ratio in Event 3 was agree well with the results of previous study which
301 reported that the largest contribution of SSAs to CCN (up to 65%) was observed in the high southern
302 latitudes (Quinn et al., 2017). Another factor affecting the estimated result in Event 3 was that R/V
303 “Xuelong” located at the high SO. As the sea state was typically not fully developed in such a
304 situation, the energy flux from the air to the ocean may differ from that under steady state conditions,
305 which may affect wave breaking and SSAs production (Lewis and Schwartz, 2004). These
306 circumstances can lead to the overestimation of vertical transport proportion of SSAs. In summary,
307 the results suggested that in the mid- and high the Southern Hemisphere, a significant proportion of
308 SSAs was transported upward and subsequently potentially affected regional climate change.

309 The influence of cyclone on SSAs in the tropics, characterized by stronger and more intricate
310 cyclonic system, was not covered by this paper. Further studies of how does SSAs concentration
311 change in the tropical cyclone area are required. However, the observational results presented in this
312 study extend the current knowledge of the impact of cyclone on marine aerosol emission in the mid-
313 and high Southern Hemisphere and their potential climate effect.

314 **Conclusions**

315 An underway aerosol monitoring system was used to determine the aerosol composition and
316 size distribution during different cyclone events in the mid- and high Southern Hemisphere in order
317 to access the potential effects of cyclone on SSAs emission. Three cyclone events were observed
318 during the 34th Chinese Antarctica Expedition Research Cruise from 23th February 2018 to 4th March
319 2018.

320 It was expected that the high wind speeds produced during the cyclone events would increase
321 the generation of SSAs. However, the SSAs levels increase in the low atmosphere were not observed
322 during these cyclone events. It indicated that considerable SSAs were transported upward to the
323 high atmosphere due to the cyclone. According to the wind stress and sea-salt flux between cyclone
324 and non-cyclone periods, the calculation indicated that more than 23% of SSAs were transported
325 upwards to the high atmosphere, with the highest proportion observed in the Southern Ocean
326 (ranging from 39% to 55%). Vertical transport of SSAs can be regarded as an important source of
327 CCN in the marine boundary layer.

328 The effect of cyclone on SSAs emission was indirect and complicated. Therefore, future work
329 is required to investigate the effect of varying intensity of the cyclone on SSAs emission and SSAs
330 generation mechanism during precipitation, as well as their potential climate effect in different

331 regions.

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341 **Data availability**

342 The data discussed in this manuscript are available from the following websites:

343 <https://doi.org/10.5281/zenodo.7912911>.

344 **Author contributions.**

345 SJ analyzed the results and wrote the paper. JY conducted the observations, proposed the
346 research ideas and wrote the paper. SW and SZ contributed considerably to the interpretation of the
347 results. MZ and QL conducted the on-board observations and data analyses. SX applied the
348 calculations of sea ice distribution and Metrological data. HY and SD contributed to the observation
349 data analyses.

350 **Competing interests**

351 The authors declare that they have no conflict of interest.

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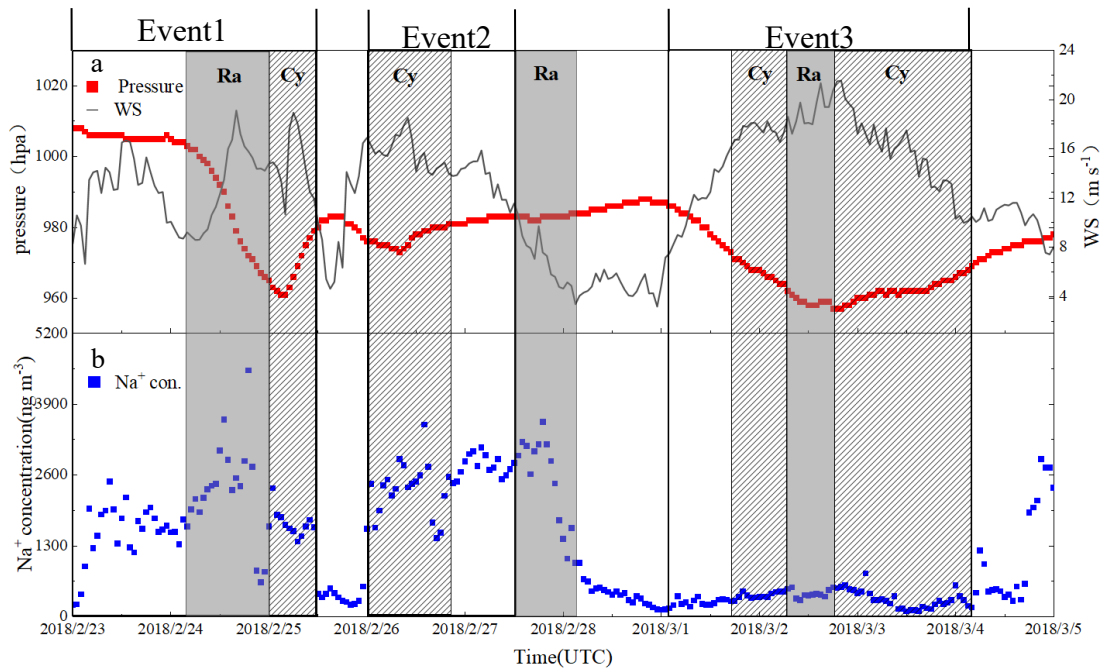
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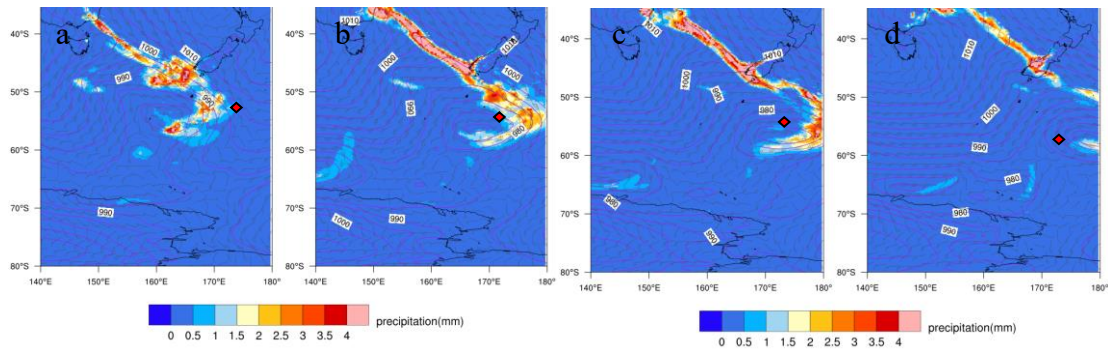
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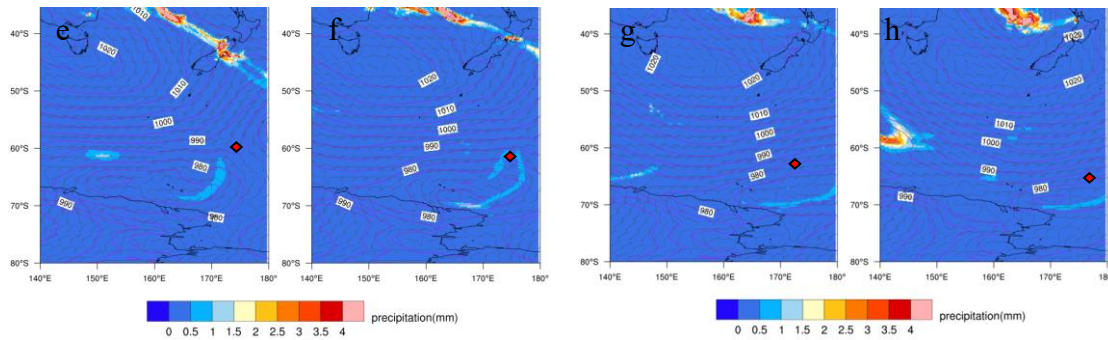
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Figure 1. Temporal distributions of Na⁺ and relevant meteorological parameters obtained during the period 23 February to 4 March 2018 in the cyclone area of the Southern Ocean. (a) Time series of atmospheric pressure (hpa) and wind speed (WS, m s⁻¹). (b) Time series of Na⁺ concentrations (ng m⁻³). Shading indicates: Ra - precipitation periods; Cy - cyclone periods. No shading corresponds to non-cyclone periods.

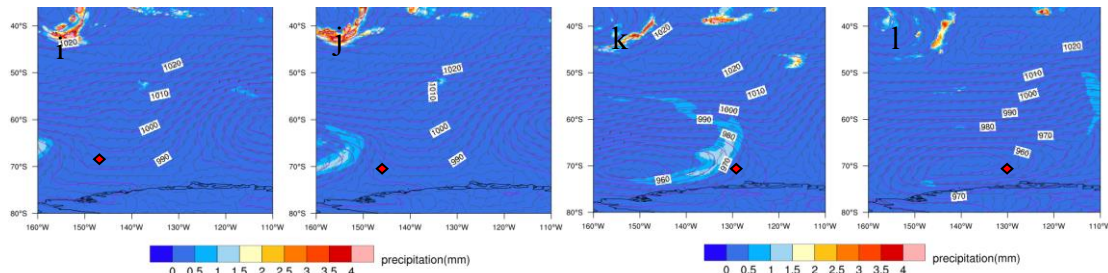
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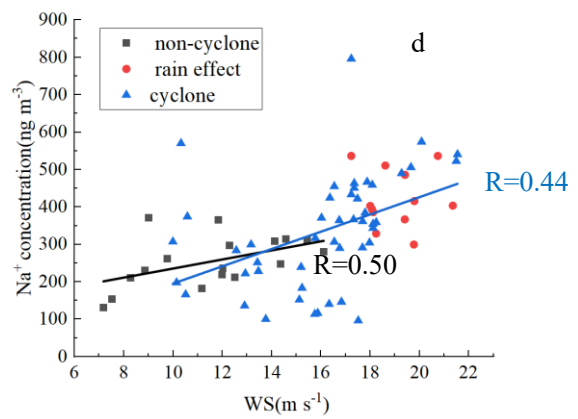
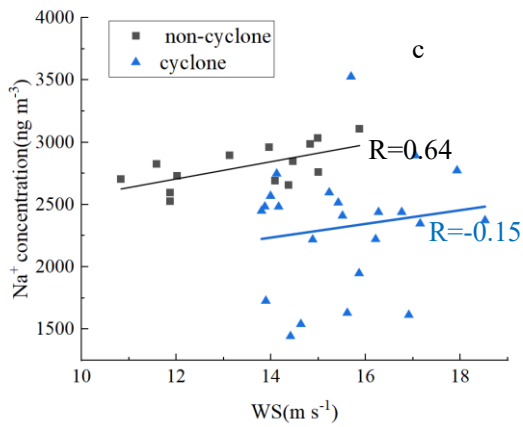
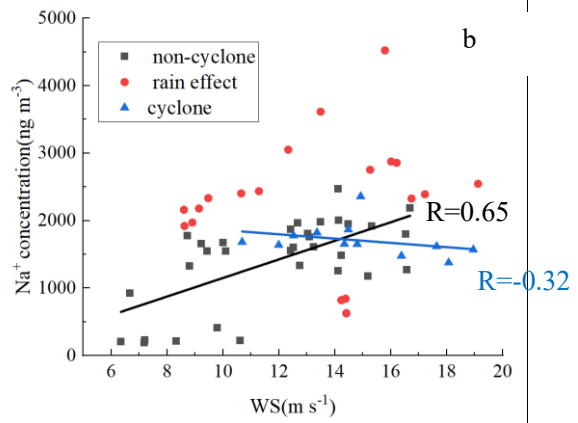
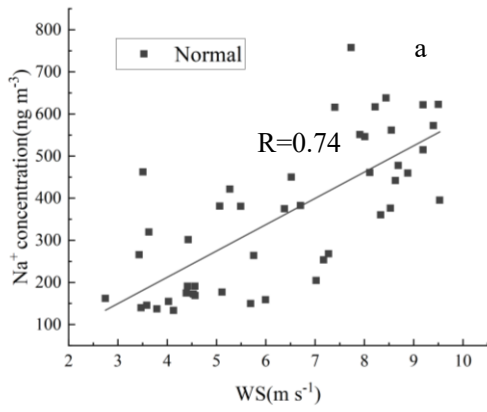
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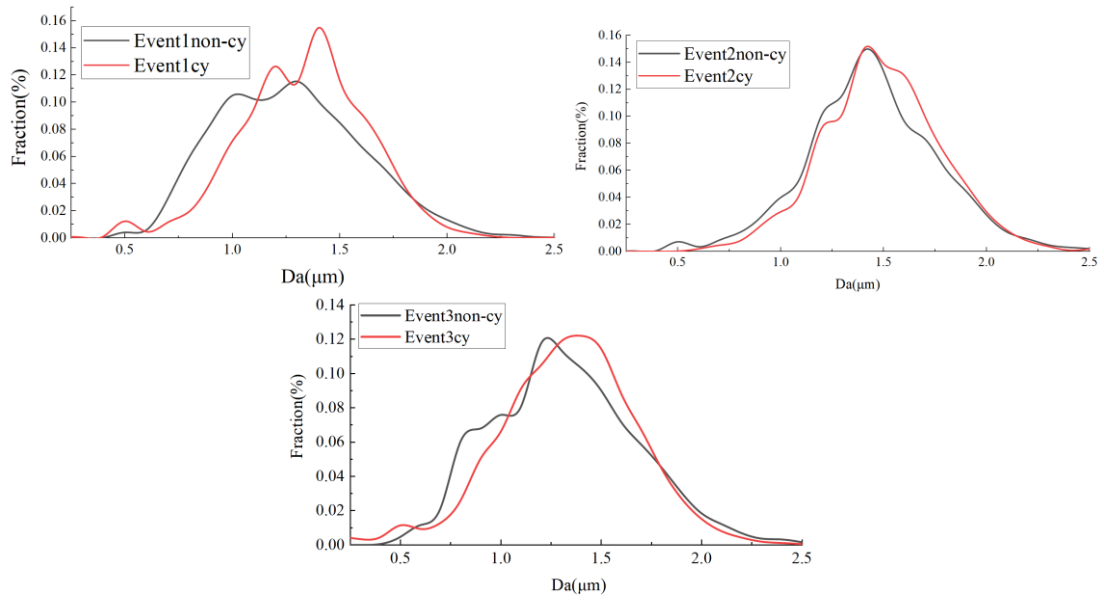
Figure 2. Sea surface pressure (hpa) and total precipitation (mm) maps for the three observed cyclone events. Event 1: (a) 14:00 24/2/18; (b) 20:00 24/2/18; (c) 01:00 25/2/18; (d) 08:00 25/2/18. Event 2: (e) 16:00 25/2/18; (f) 00:00 26/2/18; (g) 20:00 26/2/18; (h) 18:00 26/2/18. Event 3: (i) 01:00 1/3/18 (j) 18:00 1/3/18 (k) 13:00 2/3/18; (l) 22:00 2/3/18. The red diamond represents the position of the research ship. All times are UTC. The coastline of Antarctica is seen at the bottom of each figure.



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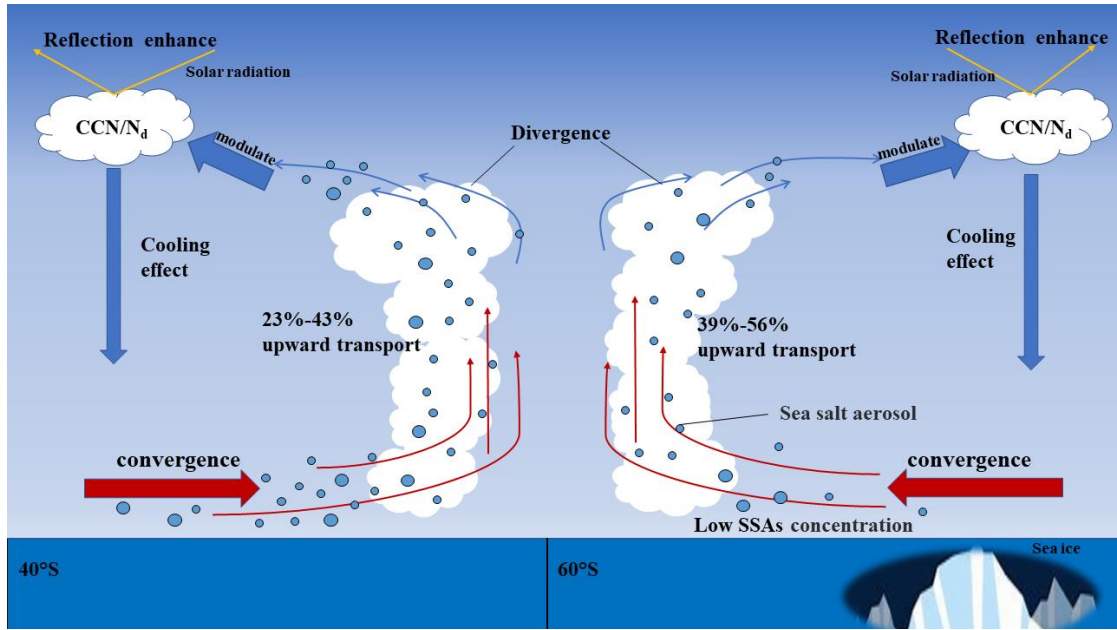
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Figure 3. Correlation between Na^+ concentration and wind speed under different meteorological conditions. (a) A non-cyclone “normal” period (i.e., stable air pressure and relative humidity, constant air mass, no precipitation). (b) Event 1 (cyclone, non-cyclone and raining periods). (c) Event 2 (cyclone and non-cyclone periods). (d) Event 3 (cyclone, non-cyclone and raining periods).



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Figure 4. SSA size distributions (in terms of fractional percent) for cyclone and non-cyclone periods during the three observed Events.



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Figure 5. Schematic diagram illustrating the impact of cyclone on SSA generation and transport and the resulting climate effects.

Table 1. Estimation of SSAs vertical transport proportion by assessing the difference of wind stress and Sea-salt flux.

	Event 1	Event 2	Event 3
Quotient of wind stress	1.689	1.310	2.153
Quotient of sea-salt flux	2.156	1.463	3.031
Average Na ⁺ con. (non-cy) ng/m ³	1273.19	2816.90	254.76
Average Na ⁺ con. (cy) ng/m ³	1647.31	2353.74	334.94
Estimated U-SSA _(wind stress) con. ng/m ³	2151.05	3689.30	548.52
Estimated U-SSA _(Sea-salt flux) con. ng/m ³	2745.16	4113.74	772.14
Estimated upward transport _(wind stress)	23.4%	36.2%	38.9%
Estimated upward transport _(Sea-salt flux)	39.9%	42.8%	56.6%

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