



1 **Measurement Report: Optical Characterization, Seasonality, and Sources of**  
2 **Brown Carbon in Fine Aerosols from Tianjin, North China: Year-round**  
3 **Observations**

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9 **Abstract**

10 To investigate the physicochemical characteristics and sources of brown carbon (BrC) in North  
11 China, we collected fine aerosols (PM<sub>2.5</sub>) at an urban site in Tianjin over a 1-year period. We  
12 measured the ultraviolet (UV) light absorption and excitation emission matrix (EEM) fluorescence  
13 of the water-soluble BrC (WSBrC) and the water-insoluble but methanol-soluble BrC (WI-MSBrC)  
14 in the PM<sub>2.5</sub> using a three-dimensional fluorescence spectrometer. Average light absorption  
15 efficiency of both WSBrC (Abs<sub>365, WSBrC</sub>) and WI-MSBrC (Abs<sub>365, WI-MSBrC</sub>) at 365 nm was found  
16 to be highest in winter and distinct from season to season. Averages of biological index (BIX) and  
17 fluorescence index (FI) of WSBrC were lower in summer than in other seasons and opposite to  
18 that of humification index (HIX), which implied that the secondary formation and further chemical  
19 processing of aerosols were intensive during the summer period than in other seasons. Whereas in  
20 winter, the higher HIX together with the higher BIX and FI of WI-MSBrC suggested that the BrC  
21 loading was mainly influenced by primary emissions and was relatively water-soluble. Based on  
22 EEM, the types of chromophores in BrC were divided into humic-like substances (HULIS),  
23 including low-oxygenated and high-oxygenated species, and protein like compounds (PLOM).  
24 The direct radiation absorption caused by WSBrC and WI-MSBrC combinedly in the range of  
25 300–400 nm was accounted for about 40% to the total radiation (range, 300–700 nm), which  
26 emphasizes that the radiation balance of the Earth's climate system is substantially affected by the  
27 BrC and should be considered in the radiative forcing models.

28

29 **1 Introduction**

30 Brown carbon (BrC) is a part of organic aerosol (OA) and has the ability to absorb solar radiation  
31 in the near-ultraviolet (UV) to visible spectrum (Liu et al., 2013). In the range of near-UV/Vis  
32 light (300–500 nm), BrC has a significant effect on radiative forcing in both regional and global  
33 climate (Feng et al., 2013; Jo et al., 2016; Park et al., 2010). However, the warming effect of water-  
34 soluble BrC in the Arctic has been reported to be accounted for about 30% of that of the black  
35 carbon (Yue et al., 2022). BrC not only affects the direct radiative forcing by OA, but also has a  
36 potential impact on indirect radiative forcing due to its hydrophilicity, which influences the  
37 formation of cloud condensation nuclei (CCN) (Andreae and Gelencsér, 2006; Laskin et al., 2015).  
38 In addition, BrC is mostly composed of highly conjugated aromatic ring compounds (such as a  
39 polycyclic aromatic hydrocarbons) and high molecular weight substances with a polar functional



40 group such as oxygen or nitrogen, or a humic-like substances (HULIS), which can cause a risk to  
41 human health. *For example*, carbon-containing aromatic compounds can cause physical weakness,  
42 decreased immunity, arteriosclerosis, etc., which will increase the mortality due to cardiovascular  
43 and cerebrovascular diseases and a variety of cancers such as skin cancer, pharyngeal cancer and  
44 nasal cancer (Diggs et al., 2011; Peters et al., 2008; Hecobian et al., 2010).

45 BrC can be emitted directly from primary sources and produced by chemical reactions of  
46 volatile organic compounds (VOCs) in the atmosphere (Chakrabarty et al., 2010; Jacobson, 1999;  
47 Sareen et al., 2010). Primary sources of BrC are combustion processes such as biomass burning  
48 (Hoffer et al., 2006; Brown et al., 2021), fossil fuel combustion (Jo et al., 2016), automobile  
49 exhaust (Liu et al., 2015) and non-combustion processes such as bioaerosols (plant debris and  
50 fungi) and soil humus (Lin et al., 2014; Rizzo et al., 2013; Rizzo et al., 2011). On the other hand,  
51 secondary BrC can be produced from complex chemical reactions of VOCs emitted from both  
52 anthropogenic and biological origin (Kasthuriarachchi et al., 2020; Li et al., 2020a).

53 After establishing the fact that organic compounds also absorb the light in recent times, the  
54 researchers are paying significant attention to the estimations of physical (optical) and chemical  
55 properties the BrC. Excitation emission matrix (EEM) fluorescence spectroscopy and ultraviolet  
56 spectroscopy are common techniques for studying the optical absorption and fluorescence  
57 chromophore optical and structural characteristics of complex organic materials (Chen et al.,  
58 2016b). Combined spectrophotometric measurement and chemical analysis has been applied to  
59 study the BrC in Xi 'an, Northwest China from July 2008 to June 2009 (Huang et al., 2018).  
60 Recently, the absorption spectroscopy of BrC has been used as a tool for understanding broader  
61 composition and characteristics of the BrC (Satish and Rastogi, 2019).

62 EEM provides some information about the chromophores responsible for the light absorption  
63 of organic matter. The fluorescence technique has been widely applied to measure organics in  
64 terrestrial and oceanic systems (Murphy et al., 2013; Yu et al., 2015), but has not been widely used  
65 in the study of atmospheric aerosols. The composition of humic-like and protein-like components  
66 have been identified from the analysis of chromophores of dissolved organic substances in aquatic  
67 environments (Xie et al., 2020). Fluorescence measurements with higher detection sensitivity are  
68 more useful in classifying BrC compared to absorption measurements, which rely on the shape of  
69 absorption spectra (Laskin et al., 2015). In recent times, the application of fluorescence technology  
70 has been well established to study the chemical composition of atmospheric aerosols (Wu et al.,  
71 2021; Deng et al., 2022; Li et al., 2022; Cao et al., 2022).

72 However, the studies on BrC are very limited because of difficulties in quantitative  
73 measurement of organic components and the optical properties of the BrC (Corbin et al., 2019;  
74 Wang et al., 2022b). In addition, traditional optical instruments do not provide any distinction  
75 between the light absorption by black and brown carbon. Therefore, the indirect approach to  
76 explore the nature and sources of BrC has been developed through its light absorption  
77 characteristics. Of course, such research is just at preliminary stages, and much attention need to  
78 be paid further such as long-term and continuous observations of light absorption characteristics  
79 of water-soluble BrC (WSBrC) and their temporal and spatial variations (Izhar et al., 2020).  
80 Moreover, the investigation of light absorption characteristics of water-insoluble BrC (WIBrC)  
81 that can be extracted into a solvent with higher extraction efficiency is necessary to better  
82 understand the impact of the BrC on climate change (Corbin et al., 2019). In fact, such studies are  
83 very scarce because the selection of solvents and determination of extraction efficiency are  
84 difficult, although different polar chromophores could be extracted by solvent extraction according  
85 to the polarity of solvent and methanol has been used as a common solvent (Chen et al., 2016a).



86 Therefore, the comprehensive study of the optical properties of WSB<sub>r</sub>C and WIB<sub>r</sub>C is highly  
87 necessary to better understand the types of chromophores and optical properties of atmospheric  
88 aerosols, as well as the processes of oxidation and transformation of chromophores at different  
89 locale over the world.

90 China is one of the most polluted areas in the world, and suffering from the absorption and  
91 scattering of solar radiation by atmospheric aerosols that directly affect the energy balance of the  
92 Earth's climate system, especially in North China Plain (Wang et al., 2022a). As an important port  
93 city in the North China Plain, Tianjin, which has a large population, has received a widespread  
94 attention to address the atmospheric environmental issues. Previous studies have shown that BrC  
95 in the atmosphere contributes significantly to the light absorption by aerosols (Deng et al., 2022).  
96 PM<sub>2.5</sub> in the Tianjin area are extremely high and contain a lot of organic matter (OM) (Dong et al.,  
97 2023). In such an environment, BrC is likely to become an important light-absorbing component  
98 of atmospheric aerosols. However, the studies on physicochemical characteristics and sources of  
99 BrC are very limited in the North China Plain, and to the best of our knowledge, the long-term  
100 observations have not been reported yet over the Tianjin region.

101 In this study, we measured the optical properties of WSB<sub>r</sub>C and water-insoluble but methanol-  
102 soluble BrC (WI-MSBrC) in fine aerosols (PM<sub>2.5</sub>) collected from Tianjin, North China over a one  
103 year period during 2018–2019. We discuss the seasonal variations in optical properties of WSB<sub>r</sub>C  
104 and WI-MSBrC and their chromophore composition assessed by three-dimensional fluorescence  
105 spectroscopy. We also assess the relationship between BrC and chemical composition in PM<sub>2.5</sub>,  
106 the possible sources of BrC and the influence of photochemical reaction processes on the BrC.  
107 Thus, this study provides a comprehensive understanding of the optical characteristics, seasonality  
108 and sources of BrC in the Tianjin region, and the need to develop the prevention and control  
109 strategies for the BrC emissions.

## 110 **2 Materials and Methods**

### 111 **2.1 Aerosol sampling**

112 Fine aerosol (PM<sub>2.5</sub>) sampling was conducted in Tianjin, a coastal city located at the lower  
113 reaches of the Haihe River and Bohai Sea and 150 km away from Beijing in the northern part of  
114 China. The sampling took place on the rooftop of a six-storey building at Tianjin University (ND,  
115 39.11°N, 117.18°E) in an urban area of Nankai District, Tianjin. A high-volume air sampler (Tisch  
116 Environmental, TE-6070DX) at a flow rate of 1.0 m<sup>3</sup> min<sup>-1</sup> and pre-combusted (6 hours at 450°C)  
117 quartz fiber filters (Pallflex 2500QAT-UP) were used for continuously collecting the PM<sub>2.5</sub>  
118 samples for 3 days (~72 hours) each during 5 July 2018 to 4 July 2019 ( $n = 121$ ). Filter blanks  
119 were collected twice per season during the sample period, using the same procedure as regular  
120 sampling, but without turning on the sampler pump. The blank filters were left in the filter hood  
121 for 10 minutes. Prior to and after sampling, each filter was dehumidified in a desiccator for 48  
122 hours, and then stored in a pre-combusted glass jar with a Teflon-lined cap in the dark at -20°C  
123 until analysis.

### 124 **2.2 Chemical analysis**

125 Details of the measurements of aerosol OC, EC, total carbon (TC), and WSOC were described  
126 by Wang et al. (Wang et al., 2019) and Dong et al. (Dong et al., 2023). Briefly, concentrations of  
127 the OC and EC were measured with a thermal-optical transmission analyzer (Sunset Laboratory



128 Inc, USA) following the IMPROVE protocol of the protective visual environment. Stable carbon  
129 and nitrogen isotope ratios of total carbon and nitrogen were measured with an elemental analyzer  
130 (EA, Flash 2000HT) coupled with stable isotope ratio mass spectrometer (IrMS, 253 Plus).  
131 Concentrations of  $K^+$  and  $Cl^-$  were determined using ion chromatography (ICS-5000 System,  
132 China, Dai An).

133 BrC was extracted into 30 ml ultrapure water ( $>18.2M\Omega$  cm) using a sample filter disc of 22  
134 mm in diameter under ultrasonication for 30 min. The extracts were filtered through a 0.45  $\mu$ m  
135 polytetrafluoron (PTFE) syringe filter to remove the water-insoluble compounds, and then  
136 transferred into another clean glass bottle. The extracts were used for the light absorption and  
137 fluorescence measurements of WSBrc. While the concentration of WSBrc was considered as the  
138 concentration of water-soluble organic carbon (WSOC).

139 After the extraction of WSBrc, the WI-MSBrC was extracted into 30 ml methanol using the  
140 same filter sample under ultrasonication for 30 min. The extracts were filtered using the same 0.45  
141  $\mu$ m PTFE syringe filter to remove the insoluble particles and filter debris. The methanol extracts  
142 were used for the measurements of optical properties of WI-MSBrC. The concentration of water-  
143 insoluble organic carbon (WIOC) was considered as the concentration of WI-MSBrC, which  
144 calculated as:  $WI-MSBrC = OC - WSOC$ .

## 145 2.3 Optical properties of brown carbon (BrC) analysis

### 146 2.3.1 Light absorption analysis

147 A three-dimensional fluorescence spectrometer (Aqualog, Horiba Scientific) was used to  
148 record the excitation-emission matrices (EEM) spectra and ultraviolet-visible (UV-Vis)  
149 absorption spectra of the solution samples in  $1 \times 1$  cm quartz cuvettes. The instrument parameters  
150 during sample analysis were as follows: The UV-Vis absorption spectra of extracts were recorded  
151 in the wavelength range of 240–700 nm. The UV-visible absorption spectra of the solvents were  
152 also recorded to subtract their contributions from the extract spectra. The EEM was recorded in  
153 the wavelength range of 240–700 nm for excitation and the integration time was 0.1 s with a 1 nm  
154 increment. An increment of 8 pixels (5.04 nm) is used as the emission wavelength interval. Prior  
155 to sample analysis, a fluorescence spectrometer was used to analyze the pure solvents of water and  
156 MeOH to obtain the reference signal.

157 Based on the light absorption spectra, the absorption data are converted to the absorption  
158 coefficient ( $Abs: m^{-1}$ ) following this formula:

$$159 Abs_{\lambda} = (A_{\lambda} - A_{700}) \times V_l / V_a / L \times \ln(10)$$

160 where  $A_{700}$  is the absorption at 700 nm, serving as a reference to account for baseline drift;  
161  $V_l$  is the volume of water or MeOH used for extraction;  $V_a$  is the volume of sampled air;  $L$  is the  
162 optical path length (0.01 m). A factor of  $\ln(10)$  is utilized to convert the log base 10 to a natural  
163 logarithm to obtain a base-e absorption coefficient. To compensate for any baseline shift that may  
164 occur during analysis, absorption at wavelengths below 700 nm is compared to that of 700 nm  
165 where no absorption occurs for ambient aerosol extracts. The average absorption coefficient  
166 between 360 and 370 nm ( $Abs_{365}$ ) is used to represent BrC absorption in order to avoid any  
167 interferences from non-organic compounds (e.g., nitrate) and to be consistent with the literature  
168 values (Huang et al., 2018).

169 Absorption Ångström exponent (AAE, Å) represents the spectral dependence of aerosol light  
170 absorption, indicating that BrC has a great contribution to aerosol Absorption. The spectral



171 dependence of light absorption by chromophores in solution can be described by the following  
172 equation:

$$173 \quad \text{Abs}_\lambda = C \times \lambda^{-\text{AAE}}$$

174 where C is a concentration of extract;  $\lambda$  is the wavelength (nm). The AAE of the filter extracts  
175 is calculated by a formula in the wavelength range of 300–500 nm. The selected range serves two  
176 purposes: (1) to prevent any interferences from non-organic compounds at lower wavelengths; (2)  
177 to ensure a sufficient signal-noise ratio for the investigating samples (Huang et al., 2018).

178 The mass absorption efficiency (MAE:  $\text{m}^2 \text{g}^{-1}$ ) of the filter extract at wavelength of  $\lambda$  can be  
179 characterized as:

$$180 \quad \text{MAE}_\lambda = \text{Abs}_\lambda / M$$

181 where M ( $\mu\text{g m}^{-3}$ ) is the concentration of WSOC for water extracts and that of WIOC for  
182 methanol extracts.

183 The imaginary part ( $k$ ) of the refractive index ( $m = n+ik$ ) is derived with the following  
184 equation:

$$185 \quad k_\lambda = (\text{MAC} \times \rho \times \lambda) / 4\pi$$

186 where MAC is the mass-absorption cross section of WSBrC or WI-MSBrC ( $\text{m}^2 \text{g}^{-1}$ ),  $\rho$  is the  
187 effective density,  $\lambda$  is the wavelength for the computed MAC including WSBrC and WI-MSBrC.  
188 For this study, an effective density of  $1.5 \text{ g m}^{-3}$  is assumed for WSBrC and WI-MSBrC in the  
189 derivation (Liu et al., 2013). MAC values are computed for 365 nm.

### 190 2.3.2 EEM and PARAFAC analysis

191 The raw EEMs were first calibrated for the correction of spectrometer factors, which reflect  
192 the spectrometer deviation and light source, and then for the inner filter correction, following the  
193 procedure described elsewhere (Chen et al., 2019; Gu and Kenny, 2009). Briefly, the inner filter  
194 correction of the EEMs was done based on the UV-Vis light absorbance of the extracts, which was  
195 lower than 0.7 in the calibrated wavelength range and is appropriate (Gu and Kenny, 2009). The  
196 signal intensity of the EEMs was then normalized to the Raman unit (RU) of water (Lawaetz and  
197 Stedmon, 2009). The fluorescence volume (FV,  $\text{RU}\cdot\text{nm}^2/\text{m}^3$ ) of extracts present in the atmosphere  
198 was estimated based on the EEMs at the excitation wavelength ranging from 240 to 700 nm, and  
199 then normalized it (i.e., NFV ( $\text{RU}\cdot\text{nm}^2\cdot[\text{mg/L}]^{-1}$ )) by dividing the FV with the concentration of  
200 WSOC and WIOC in the aerosol [ $\text{mg m}^{-3}$ ]).

201 Various types of chromophores present in the  $\text{PM}_{2.5}$  samples were classified and identified  
202 based on the PARAFAC analysis of the EEMs using the SOLO, the data analysis software.  
203 PARAFAC analysis was performed for each extraction fluid in each season. Ultimately, three  
204 EEM components were determined and assigned to different types of chromophores.

205 Additionally, fluorescence index (FI) was determined by calculating the ratio of emission  
206 intensities at 450 nm and 500 nm after excitation at 370 nm. Contributions from local biological  
207 sources can be characterized by biological index (BIX), which was calculated using the ratio of  
208 emission intensities at 380 and 430 nm following 310 nm excitation (Gao yan and Zhang, 2018).

### 209 2.3.3 Simple forcing efficiency (SFE)

210 It is possible to make a rough estimate of the radiative forcing caused by aerosols using a  
211 simple forcing efficiency (SFE,  $\text{W/g}$ ), which was assessed as described in the literature (Bond and  
212 Bergstrom, 2006; Deng et al., 2022):





$$SFE = \int \frac{dS(\lambda)}{d\lambda} \tau_{atm}^2 (1 - F_c) a_s MAE(\lambda) d\lambda$$

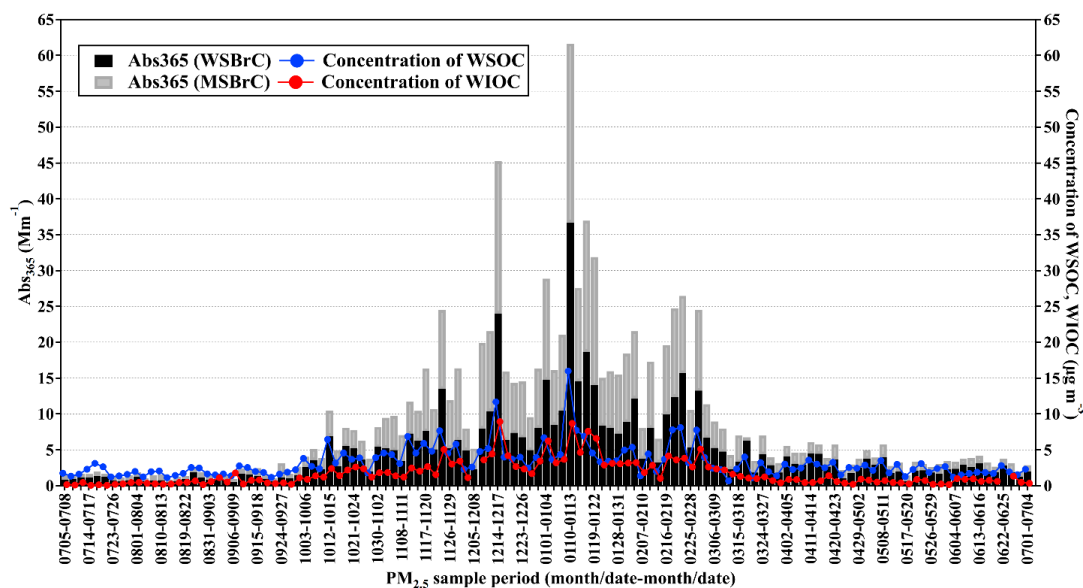
where  $dS/d\lambda$  is the solar irradiance,  $\tau_{atm}$  is the atmospheric transmission (0.79),  $F_c$  is the cloud fraction (approximately 0.6),  $a$  is the surface albedo (average 0.19),  $\beta$  is the backscatter fraction, and MSE and MAE are the mass scattering (can be ignored) and absorption efficiency, respectively.

### 3 Results and discussion

#### 3.1 Characteristics of ultraviolet light absorption of BrC

##### 3.1.1 Absorption coefficient (Abs: $Mm^{-1}$ )

The annual and seasonal concentrations and optical properties of BrC are summarized in Table 1. Temporal variations in absorption coefficient of WSB<sub>BrC</sub> and WI-MSB<sub>BrC</sub> at 365 nm (i.e., Abs<sub>365, WSB<sub>BrC</sub></sub> and Abs<sub>365, WI-MSB<sub>BrC</sub></sub>) together with their concentrations are depicted in Fig. 1. Averages of all measured parameters peaked in winter (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $10.4 \pm 6.76 Mm^{-1}$ , Abs<sub>365, MSB<sub>BrC</sub></sub> =  $10.0 \pm 5.13 Mm^{-1}$ ), followed by a decrease to autumn and spring and the lowest in summer (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $1.47 \pm 0.77 Mm^{-1}$ , Abs<sub>365, WI-MSB<sub>BrC</sub></sub> =  $0.74 \pm 0.25 Mm^{-1}$ ). The lower absorbance in summer might have been caused by extensive oxidation of organics and thus, the enhanced decomposition of some BrC substances, due to high solar light intensity and ambient temperatures. While the increase in the absorption coefficient of BrC in winter might be mainly due to the existence of large amounts of organic aerosols under the unfavorable meteorological conditions. The seasonal variations of Abs<sub>365</sub> in Tianjin were similar to those reported in the southeastern United States, but the Abs<sub>365</sub> was much higher than that (0.3–3.0  $Mm^{-1}$  in 2007) in the southeastern United States (Hecobian et al., 2010) and that in Atlanta and Los Angeles ( $0.88 \pm 0.71$  and  $0.61 \pm 0.38 Mm^{-1}$ , respectively) in summer 2010 (Zhang et al., 2011). However, compared with the Abs<sub>365</sub> of WSB<sub>BrC</sub> ( $14.1 \pm 8.5 Mm^{-1}$ ) in winter 2016 and summer 2017 ( $2.1 \pm 1.0 Mm^{-1}$ ) in Tianjin (Deng et al., 2022), the Abs<sub>365</sub> has slightly decreased in this study. Spatially, the value and contribution of Abs<sub>365</sub> in this study (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $10.4 Mm^{-1}$ , Abs<sub>365, WI-MSB<sub>BrC</sub></sub> =  $10.0 Mm^{-1}$ ) in winter were higher than those reported at different locations in southern China; Nanjing (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $4.84 Mm^{-1}$ , Abs<sub>365, MSB<sub>BrC</sub></sub> =  $7.75 Mm^{-1}$ ) (Xie et al., 2020), Guangzhou in autumn (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $8.8 Mm^{-1}$ ) (Li et al., 2018), and Lhasa in winter (Abs<sub>365, WSB<sub>BrC</sub></sub> =  $1.04 Mm^{-1}$ , Abs<sub>365, MSB<sub>BrC</sub></sub> =  $1.47 Mm^{-1}$ ) (Zhu et al., 2018). Interestingly, the Abs<sub>365</sub> in Tianjin in winter is lower than that in Beijing, Xi'an, which are heavily polluted cities in northern China (Huang et al., 2020; Li et al., 2020b). The higher Abs<sub>365</sub> in winter indicates that the light absorption of BrC in PM<sub>2.5</sub> may have more significant effect on the climate and the photochemical reactions in the atmosphere over Tianjin in winter than in other seasons.



245

246 **Figure 1.** Time series of the light absorption coefficient of water-soluble brown carbon (WSBrC)  
 247 and water-insoluble but methanol-soluble BrC (WI-MSBrC) at 365nm ( $Abs_{365, WSBrC}$  and  $Abs_{365, WI-MSBrC}$ , respectively)  
 248 and their concentrations in Tianjin, North China during 2018 and 2019.  
 249

250 Figure 2a shows the seasonal average absorption spectra of WSBrC and WI-MSBrC at  
 251 wavelengths of 240–700 nm, which shows the common feature of BrC, i.e., the absorption of  
 252 shorter wavelengths increases sharply and significantly. Such feature is different from the  
 253 absorption characteristics of BC, whose AAE is close to 1 and weakly dependent on the  
 254 wavelength. Another evident feature of BrC absorption spectra shown in Figure 2a is that the  
 255 absorption coefficient of WI-MSBrC was always greater than that of WSBrC across the shorter  
 256 wavelengths, which is consistent with the previous studies (Huang et al., 2020; Li et al., 2020b).  
 257 This can be attributed to the difference in types and amounts of chromophores extracted, i.e., more  
 258 chromophores were soluble in methanol (e.g., PAHs from biomass burning and fossil fuel  
 259 combustion), but not in water. It is worthy to note that,  $\pi-\pi^*$  electron transitions in the double  
 260 bonds of aromatic compounds are the primary cause of light absorption in the wavelength range  
 261 of 250–300 nm. Nitroaromatic compounds contribute 60% of absorbance in the 300–400 nm range  
 262 (Hems et al., 2021). According to most of the studies, the absorption of hydroxylation and ring  
 263 cleavage products formed from nitrophenol exceeds 400 nm and nitroaromatics have strong light  
 264 absorption in the visible region (Vidović et al., 2020; Satish and Rastogi, 2019). The electron  
 265 transitions in phenolic arenes, aniline derivatives, polyenes and polycyclic aromatic hydrocarbons  
 266 with two or more rings are responsible for the absorbance in the bands between 270 and 280 nm  
 267 (Baduel et al., 2009). There was a peak of light absorption at 280 nm in WI-MSBrC spectra, but  
 268 not in that of WSBrC, probably because some polycyclic aromatic chromophores were insoluble  
 269 in water but soluble in MeOH.

270 Figs. 3 and 4 show the seasonal variations in the correlations between  $Abs_{365}$  and chemical  
 271 components, including WSBrC and WI-MSBrC,  $K^+$ , and  $Cl^-$ , which are possibly emitted from  
 272 biomass burning. In this study, the light absorption at the wavelength of 365 nm would not be



273 interfered by inorganic substances, so Abs at 365 nm was selected for analysis (Hecobian et al.,  
274 2010). High correlations were found between Abs<sub>365</sub> and WSB<sub>r</sub>C, WI-MSB<sub>r</sub>C, except in summer,  
275 indicating that WSB<sub>r</sub>C and WI-MSB<sub>r</sub>C might have been derived from similar sources, except in  
276 summer, because the light absorption efficiency of organic compounds from different sources were  
277 significantly different. *For example*, precursors and reaction processes affect the light absorption  
278 capacity of SOA in the atmosphere (Zhong and Jang, 2011).

279 In fact, the Abs depends on the amount of BrC availability, but not of total OC content. The  
280 BrC content increases with increasing level of nitrogen containing organics (with aging in presence  
281 of NO<sub>3</sub><sup>-</sup>/NH<sub>3</sub> etc.). So, the higher levels of OC or WSOC might occur due to enhanced primary  
282 emissions and/or secondary formation on those particular days, but the BrC content in that high  
283 amount of OC or WSOC might be less due to either aging and/or less availability of N species to  
284 produce N containing organics in summer.

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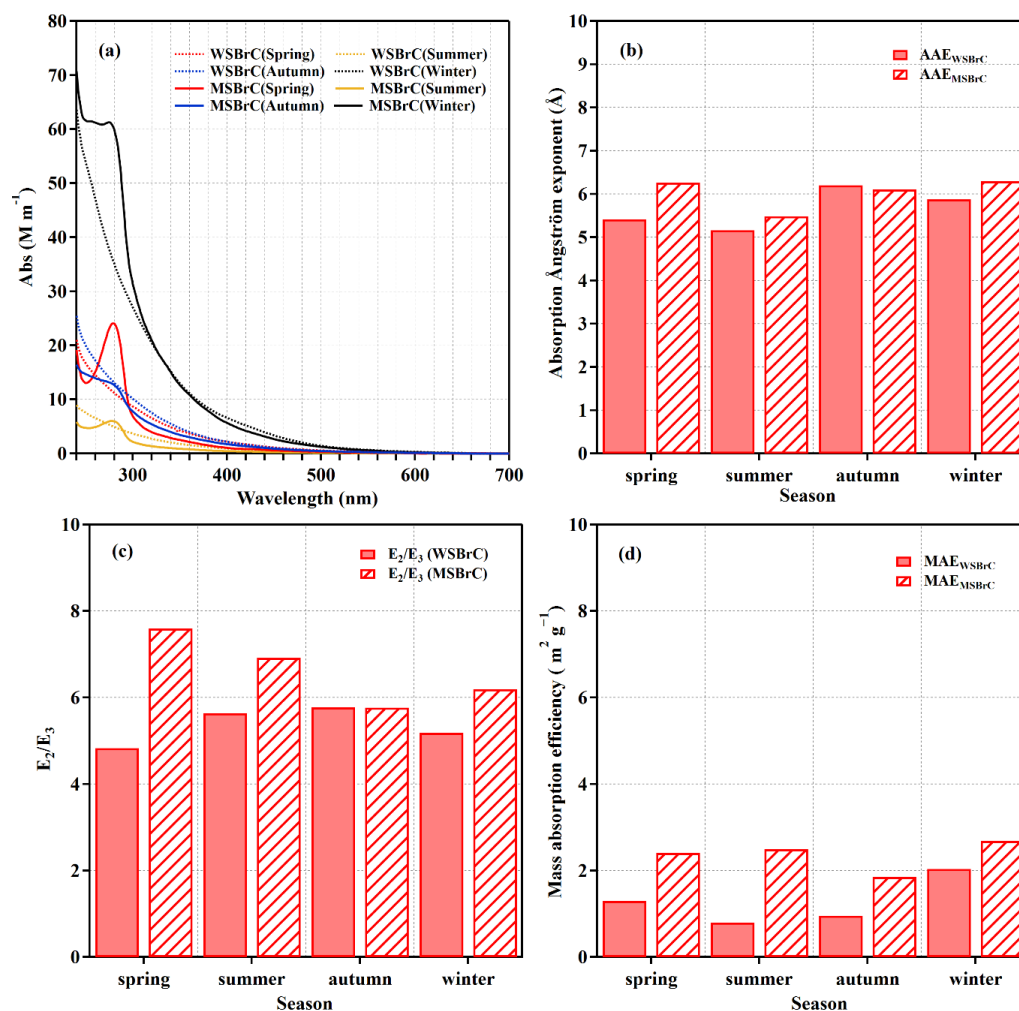
288





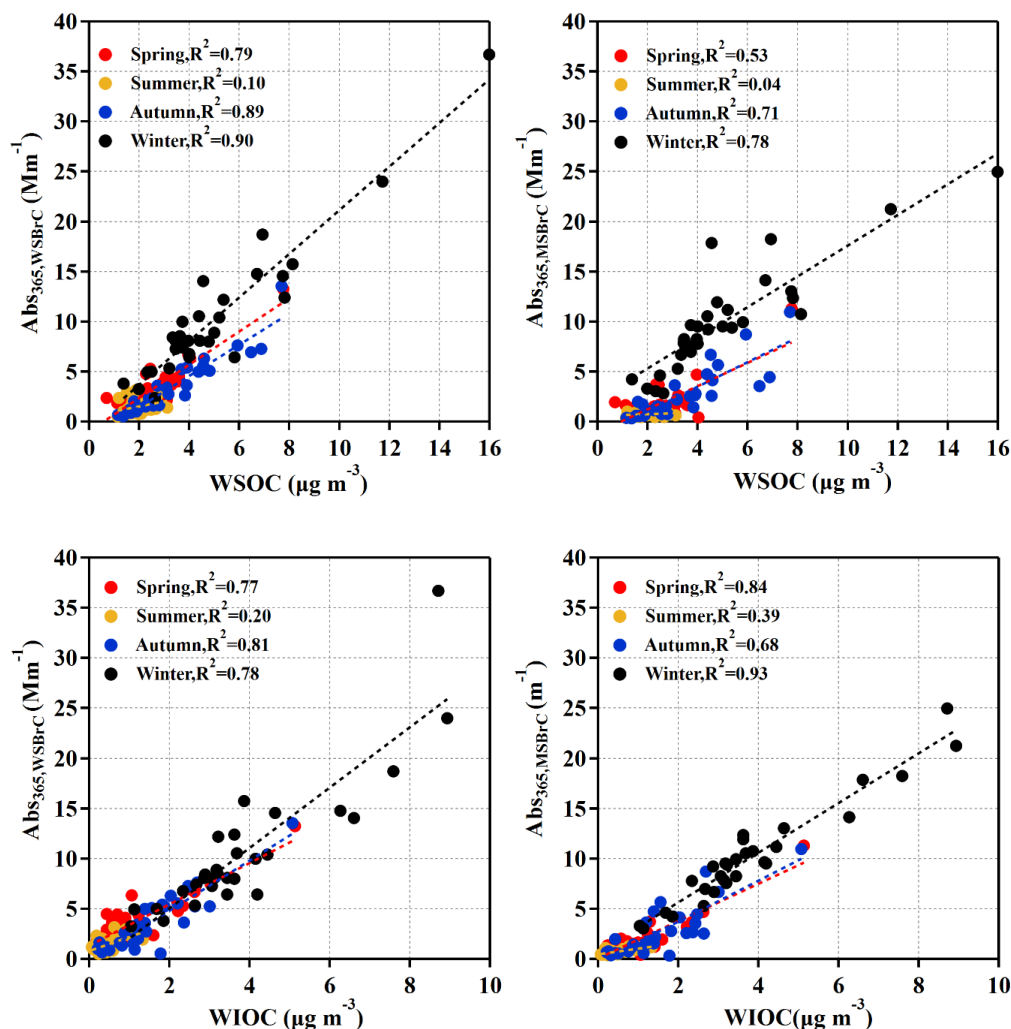
**Table 1.** Mass concentrations of WSOC, WIOC and absorbance efficiency of WSBrC and WI-MSBrC in PM<sub>2.5</sub> from Tianjin, North China.

	Spring		Summer		Autumn		Winter		Annual	
	Range	Ave ± SD	Range	Ave ± SD	Range	Ave ± SD	Range	Ave ± SD	Range	Ave ± SD
WSOC ( $\mu\text{g m}^{-3}$ )	0.69–4.03	2.48±0.82	1.14–3.12	1.88±0.53	1.16–7.68	3.45±1.71	1.37–16.0	5.06±2.99	0.69–16.0	3.25 ± 2.18
WIOC ( $\mu\text{g m}^{-3}$ )	0.23–2.62	0.88 ± 0.63	0.00–1.33	0.43±0.32	0.21–5.07	1.55±1.04	0.00–8.93	3.74±2.09	0.00–8.93	1.68 ± 1.77
	Concentrations									
	Optical parameters									
Abs <sub>366</sub> (M m <sup>-1</sup> )	0.66–13.3	3.45±2.29	0.49–3.16	1.47±0.77	0.55–13.5	3.71±2.83	2.35–36.7	10.4±6.76	0.49–36.7	4.74±5.10
MAE <sub>366</sub> (m <sup>2</sup> g <sup>-1</sup> )	0.52–3.41	1.31±0.55	0.38–1.98	0.80±0.44	0.40–1.76	0.96±0.33	0.90–3.08	2.04±0.46	0.38–3.41	1.28±0.66
WSBrC	3.85–7.57	5.42±0.74	3.90–6.88	5.17±0.83	5.12–7.99	6.21±0.65	4.50–7.39	5.88±0.58	3.85–7.99	5.66±0.82
rC	1.13–1.63	1.37±0.09	1.16–1.49	1.31±0.07	1.36–1.61	1.47±0.07	1.29–1.44	1.37±0.02	1.13–1.63	1.38±0.09
BIX	0.82–1.24	1.01±0.11	0.79–1.04	0.91±0.06	0.83–1.26	1.06±0.08	1.03–1.39	1.20±0.08	0.79–1.39	1.05±0.13
HIX	1.84–3.76	2.76±0.47	2.47–3.98	3.12±0.44	2.11–4.17	3.11±0.51	1.72–3.72	2.47±0.43	1.72–4.17	2.87±0.53
k <sub>366</sub>	0.023–0.149	0.057±0.024	0.017–0.086	0.035±0.020	0.018–0.077	0.042±0.015	0.039–0.134	0.089±0.021	0.017–0.149	0.056±0.029
SFE <sub>300–600</sub> (W g <sup>-1</sup> )	0.62–2.71	1.46±0.52	0.60–2.99	1.21±0.67	0.81–5.13	1.99±0.84	1.40–4.76	3.12±0.71	0.60–5.13	1.95±1.02
SFE <sub>300–700</sub> (W g <sup>-1</sup> )	0.98–6.36	3.39±1.42	1.22–10.5	3.68±2.58	1.48–12.5	5.12±2.17	3.75–13.1	7.60±2.17	0.98–13.1	4.97±2.71
Abs <sub>366</sub> (M m <sup>-1</sup> )	0.44–11.3	1.99±1.95	0.40–1.26	0.74±0.25	0.32–11.0	2.83±2.51	2.85–25.0	10.0±5.13	0.32–25.0	3.87±4.69
MAE <sub>366</sub> (m <sup>2</sup> g <sup>-1</sup> )	0.42–5.81	2.41±1.28	0.89–7.05	2.50±1.78	0.18–4.70	1.86±1.02	2.01–3.42	2.69±0.36	0.18–7.05	2.36±1.26
WI-MSBrC	3.94–8.38	6.27±0.90	4.27–9.19	5.49±1.26	2.08–12.9	6.11±1.86	5.49–6.76	6.30±0.27	2.08–12.9	6.06±1.23
rC	1.29–1.77	1.51±0.11	1.34–1.92	1.58±0.12	1.48–1.73	1.57±0.06	1.61–2.24	1.73±0.11	1.29–2.24	1.60±0.13
BIX	0.94–1.76	1.23±0.18	0.92–1.65	1.32±0.18	0.83–1.36	1.05±0.14	1.20–1.62	1.43±0.09	0.83–1.76	1.26±0.21
HIX	0.11–1.26	0.42±0.28	0.11–0.49	0.25±0.08	0.30–2.38	1.23±0.61	0.62–1.79	1.33±0.30	0.11–2.38	0.81±0.60
k <sub>366</sub>	0.018–0.253	0.105±0.057	0.039–0.307	0.109±0.079	0.008–0.205	0.081±0.045	0.0870–1.49	0.117±0.016	0.0010–3.07	0.077±0.064
SFE <sub>300–600</sub> (W g <sup>-1</sup> )	0.64–8.84	3.61±1.91	0.60–2.99	1.21±0.67	0.75–7.01	2.98±1.52	3.04–5.29	4.13±0.57	0.64–8.84	2.98±1.70
SFE <sub>300–700</sub> (W g <sup>-1</sup> )	2.48–21.8	8.70±5.03	1.22–10.5	3.68±2.58	0.92–51.3	8.69±9.23	7.06–11.7	9.36±4.51	0.92–51.3	7.58±5.75



290

291 **Figure 2.** Seasonal averages of (a) absorption spectra in the wavelength range of 240–700 nm  
 292 plotted on a linear scale, (b) absorption Ångström exponent (AAE), (c) ratio of MAE at 250 nm to  
 293 that at 365 nm ( $E_2/E_3$ ) and (d) mass absorption efficiency (MAE:  $m^2 g^{-1}$ ) of WSBrC and WI-  
 294 MSBrC in PM<sub>2.5</sub> from Tianjin, North China.



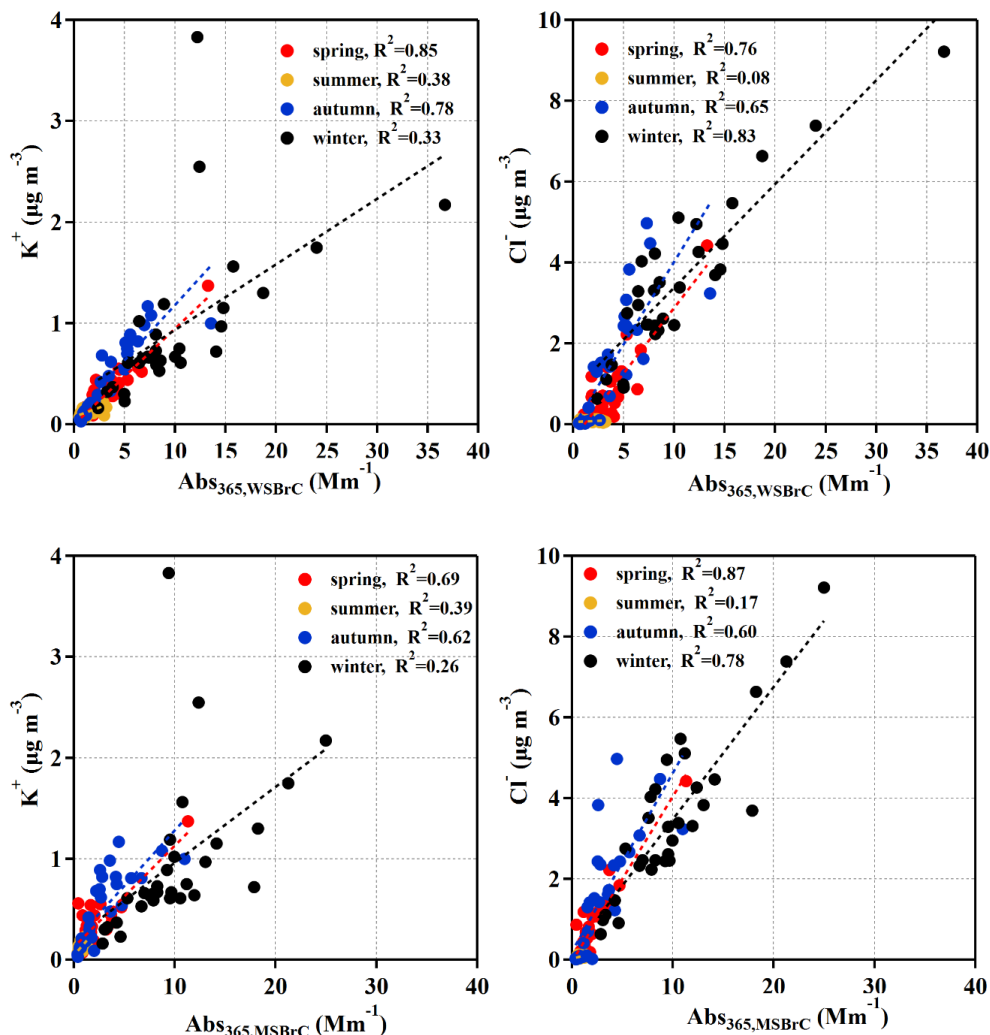
295

296

297 **Figure 3.** Scatter plots of Abs<sub>365, WSBrc</sub> and Abs<sub>365, WI-MSBrc</sub> with WSBrc and WI-MSBrc in  
 298 PM<sub>2.5</sub> from Tianjin in each season during 2018–2019. The WSOC and WIOC data is obtained  
 299 from (Dong et al., 2023).

300

301 As shown in Fig. 4, considerable positive correlations were found between Abs<sub>365</sub> and K<sup>+</sup> and  
 302 Cl<sup>-</sup> in autumn and spring, indicating that biomass burning was a major source of BrC in those  
 303 seasons. In addition, the correlation between WSOC chromophores and K<sup>+</sup> in autumn was stronger  
 304 than that between methanol-soluble photogenic groups, which again confirmed that most of the  
 305 chromophores generated by biomass burning were water-soluble photogenic groups. However, the  
 306 correlation between WI-MSBrc and Cl<sup>-</sup> was stronger than that between WSBrc and Cl<sup>-</sup> in spring,  
 307 which might be due to enhanced contribution of a large number of BrC chromophores, which were  
 308 insoluble in water, from dust in spring.



309

310

311 **Figure 4.** Scatter plots of  $\text{Abs}_{365, \text{WSBrC}}$  and  $\text{Abs}_{365, \text{MSBrC}}$  with  $K^+$  and  $\text{Cl}^-$  in  $\text{PM}_{2.5}$  from Tianjin  
 312 in each season during 2018–2019. The concentration of  $K^+$  and  $\text{Cl}^-$  from (Dong et al., 2023).  
 313

314 3.1.2 Absorption Ångström exponent (AAE)

315 The magnitude of the AAE can reflect the difference in BrC sources and atmospheric  
 316 chemical processes (Lack et al., 2013). It has been reported that the AAE of light-absorbing organic  
 317 species (i.e., BrC) is much larger than that of soot (BC). The AAE was found to be between 2 and  
 318 4 for the particles containing both soot and BrC. AAE value of particulate matter was closely  
 319 related to its chemical composition, mixing state, particle size and other factors.



320 It is important to note that the solvent extractant light absorption characteristics of organic  
321 components do not be affected by particle size and chemical composition of aerosols. The AAE  
322 value of the extract mainly depends on the types of absorbable components in the extract. Hoffer  
323 et al. (2006) isolated humic-like substances (HULIS) from particulate matter emitted by  
324 combustion substances by combining water extraction with exchange column, and measured its  
325 AAE value as 7 (Hoffer et al., 2006). As shown in Fig. 2b, the seasonal average AAE of WSBrc  
326 varied slightly between 3.85 and 7.99 with an average of 5.66, which was comparable to those  
327 reported from New Delhi, Beijing and the outflow region of northern China (Lesworth et al., 2010).  
328 The AAE of WSBrc in Tianjin was also similar to that (range, 6–8) reported in the off-line  
329 particulate matter samples collected from the southeastern United States (Hecobian et al., 2010)  
330 and downtown Atlanta (Liu et al., 2013). The AAE of WI-MSBrC ( $6.06 \pm 1.23$ ) was comparable  
331 with that of WSBrc, consistent with a previous study in urban Beijing during winter and Xi'an,  
332 China. It has been reported that the AAE of BrC is increased with the increase in polarity (Chen et  
333 al., 2016a). The higher AAE of the BrC implies that the OA is mostly polar in Tianjin.

### 334 3.1.3 Mass absorption efficiency (MAE: $\text{m}^2 \text{g}^{-1}$ )

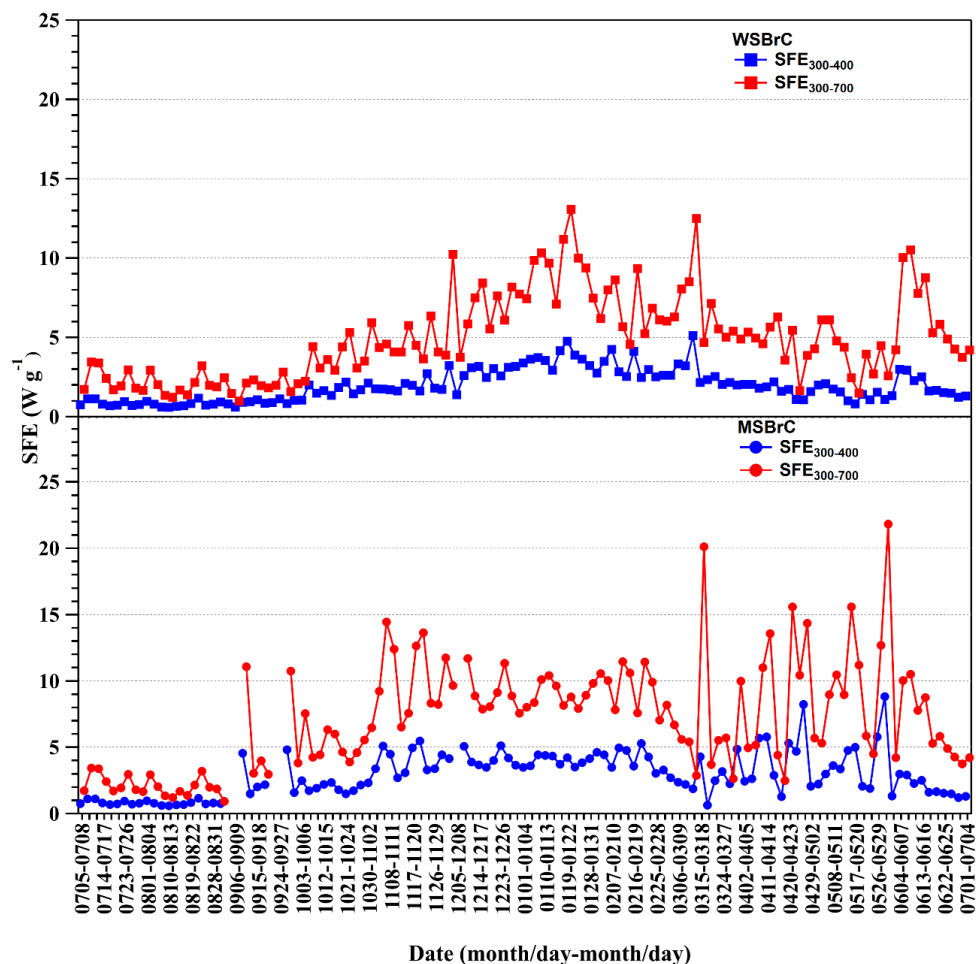
335 The average  $\text{MAE}_{365}$  of both WSBrc and WI-MSBrC were higher in winter (1.28 and 2.36  
336  $\text{m}^2 \text{g}^{-1}$ , respectively). It is interesting to note that minimum value of  $\text{MAE}_{365}$  of WSBrc appeared  
337 in summer ( $0.80 \pm 0.44$ ), which was opposite to that of WI-MSBrC, the least value was appeared  
338 in autumn ( $1.86 \pm 1.02$ ), which was similar with that reported in Xi'an (Li et al., 2020b). Such  
339 large seasonal differences indicated that the BrC sources might be different in each season. During  
340 winter, contributions of aerosols from coal and biomass burning were significantly higher than  
341 other seasons due to increased residential heating activities. The lower  $\text{MAE}_{365}$  values observed  
342 in summer and autumn may be attributed to biogenic sources and/or aged secondary BrC. Whereas  
343 the largest  $\text{MAE}_{365}$  appeared in cold period, which result in severe air pollution in the cold period.

### 344 3.2 Direct radiative forcing of BrC

345 Radiative forcing efficiency is calculated by integrating wavelengths from 300 nm to 700 nm.  
346 In this study,  $\text{SFE}_{300-400}$  was integrated to estimate the radiative forcing efficiency of BrC, because  
347 the BrC strongly absorbs light in the UV range. The temporal variations of SFE in different  
348 wavelength ranges in the two solvents were shown in Fig. 5.  $\text{SFE}_{300-400}$  and  $\text{SFE}_{300-700}$  showed  
349 similar seasonal trend in both the solvents, which was the same as the seasonal variation of the  $k$ .  
350 In WSBrc, it was 61% and 52% larger in winter than those in summer, respectively, indicating  
351 that BrC abundance and strong light absorption capacity in winter led to a significant increase in  
352 direct radiative forcing of the BrC. The integrated average SFE for 300–400 nm ( $\text{SFE}_{300-400}$ ) of  
353 WSBrc and WI-MSBrC were  $1.95 \pm 1.02$  and  $2.98 \pm 1.70$ , respectively. The average  $\text{SFE}_{300-700}$   
354 of both WSBrc ( $4.97 \pm 2.71$ ) and WI-MSBrC ( $7.58 \pm 5.75$ ) were about 2.5 times larger than that  
355 of the  $\text{SFE}_{300-400}$ . The SFE of WSBrc (0.98–13.1) and WI-MSBrC (0.92–51.3) in the range of  
356 300–700 nm varied widely, which was much larger than the maximum SFE obtained in the range  
357 of 300–400 nm. Furthermore,  $\text{SFE}_{300-400}$  accounted for 40% of  $\text{SFE}_{300-700}$  in the two extractants,  
358 which are similar to that reported in Tianjin by Deng et al. (2022), which indicates that the light  
359 absorption by BrC in UV range play a significant role in the radiative forcing. The imaginary part  
360 ( $k$ ) is a significant parameter indicting the direct radiative forcing of aerosols in climate model.  
361 The  $k$  of WSBrc and WI-MSBrC in Tianjin showed the similar variation tendency in seasons, with



362 the largest value in winter with the averages of 0.089 and 0.117, respectively. The values are shown  
 363 in Table 1 and the  $k$  of WI-MSBrC was larger than WSBrc in all seasons, which indicate that most  
 364 of light-absorbing chromophores are soluble in MeOH.



365

366 **Figure 5.** Temporal variations in SFE of WSBrc and WI-MSBrC from 300–400nm and 300–  
 367 700nm in PM<sub>2.5</sub> from Tianjin.

368

369 3.3 Fluorescence characteristics of BrC

370 3.3.1 Fluorescent properties

371 The fluorescent properties of WSBrc and WI-MSBrC are presented in Table 1. The  
 372 composition of chromophores in WSBrc and WI-MSBrC was analyzed by their fluorescence  
 373 spectrum. By comparing the fluorescence intensity of chromophores in WSBrc and WI-MSBrC  
 374 (Fig. 6), it was found that the proportion of water-soluble chromophores in Tianjin PM<sub>2.5</sub> was





375 higher in autumn (35%), followed by summer (29%), winter (22%) and spring (22%). Previous  
376 studies have shown that biomass combustion produces more water-soluble chromophores  
377 (Budisulistiorini et al., 2017; Lin et al., 2016; Song et al., 2018). The relative content of water-  
378 soluble chromophores in Tianjin samples was higher in autumn than that in winter, suggesting that  
379 the BrC might mainly derived from biomass butning (BB) in autumn. Furthermore, the relative  
380 proportion of water-soluble chromophores in SOA is increased with the increase in oxidation  
381 degree (Updyke et al., 2012; Wong et al., 2017) and can reach up to 70% (Chen et al., 2020). The  
382 content of SOA in Tianjin was relatively high due to enhanced aging in summer. Therefore, the  
383 content of water-soluble chromophore in BrC in summer was higher than that in spring and winter.  
384 We also caculated the humification index (HIX) and fluorecence index (FI), which further  
385 supported that the water-soluble chromophores of BrC were significantly influenced by primary  
386 emissions in autumn and the aging in summer.

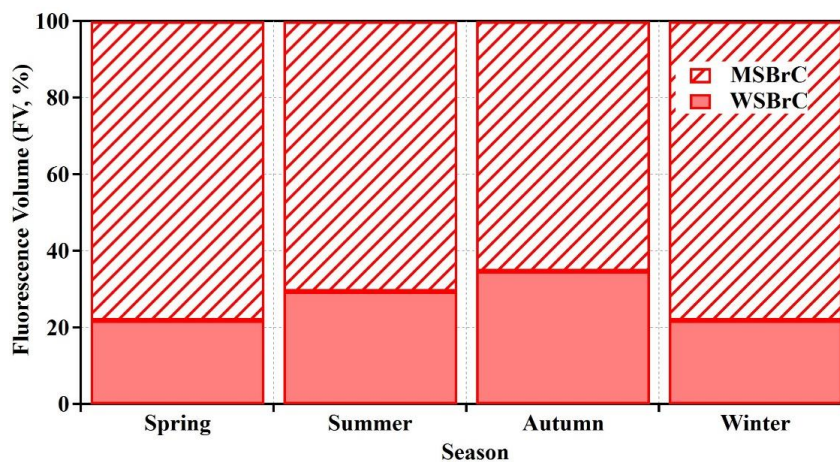
387 The fluorecence spectrum of WSBrC was similar to that of HULIS, and the humification  
388 index (HIX) is measured to reflect the degree of humification of the OA. It has been established  
389 that HIX and aromaticity show a consistent change law, that is, the value of HIX is higher, the  
390 substance has a high polycondensation degree (Deng et al., 2022). The average HIX of WSBrC  
391 and WI-MSBrC were  $2.87 \pm 0.53$  (1.72–4.17) and  $0.81 \pm 0.60$  (0.11–2.38) respectively, which  
392 indicate that the BrC in Tianjin might be either humified or aromatic, more apparently soluble in  
393 water. Besides, the temporal variation of HIX of WSBrC was completely different to that of WI-  
394 MSBrC. The higher molecular weight and aromatic organic compounds contribute more to  
395 WSBrC in summer and autumn while the contents of WI-MSBrC (winter > autumn > spring >  
396 summer) were opposite. This phenomenon was comparable to that in Nanjing (WSOC:  $7.07 \pm 2.41$ ;  
397 MSOC:  $4.84 \pm 2.47$ ) (Xie et al., 2020). This phenomenon demonstrates that the influence of  
398 molecular composition on changes in optical properties. Interestingly, it has been reported that  
399 aging processes and HIX have a significant relation (Deng et al., 2022). HIX of WSBrC confirms  
400 the BrC was significantly produced from aromatic compounds and subjected for significant  
401 atmospheric aging in summer.

402 Fluorecence index (FI) plays an important role in exploring the source and aging of OA and  
403 attracted much attention in recent times (Xie et al., 2020; Gao yan and Zhang, 2018; Qin et al.,  
404 2018; Deng et al., 2022). On the other hand, FI has been considered as an indicator to assess the  
405 terrestrially derived fulvic acids contribution to OA, the FI lower than 1.4 is associated with higher  
406 aromaticity (Gao yan and Zhang, 2018). The FI of WSBrC ranged from 1.13 to 1.63 with an  
407 average of 1.38, which was comparable to that reported from Lanzhou, China (1.2 in summer and  
408 1.7 in winter) (Gao yan and Zhang, 2018). Both FI and BIX of WSBrC in autumn have higher  
409 values, indicating that the BrC was mainly derived from terrestrial organic matter that should have  
410 largely consist of aromatic compounds. On the other hand, the variation of FI was related to  
411 photobleaching. With the increase of aromatic substances during the aging of OA, FI decreased  
412 with light strength. In contrast, the FI of WI-MSBrC ranged from 1.29 to 2.24 (ave. 1.60), with the  
413 lowest in spring ( $1.51 \pm 0.11$ ) and the highest in winter ( $1.73 \pm 0.11$ ), which were higher than that  
414 of WSBrC. The variation of HIX and FI of WSBrC and WI-MSBrC indicated that the fluorecence  
415 substance contains mainly aromatic species and significantly subjected for the aging in summer,  
416 whereas in autumn, that must have mainly derived from terrestrial organics, which were highly  
417 water-soluble.

418 The BIX showed an obvious seasonal variations and the average value was  $1.20 \pm 0.08$  for  
419 WSBrC and  $1.43 \pm 0.09$  for WI-MSBrC in Tianjin PM<sub>2.5</sub>, which were higher than those reported  
420 in the forest environment ( $1.01 \pm 0.23$ ) (Zhao et al., 2019). Furthermore, the BIX of WSBrC in

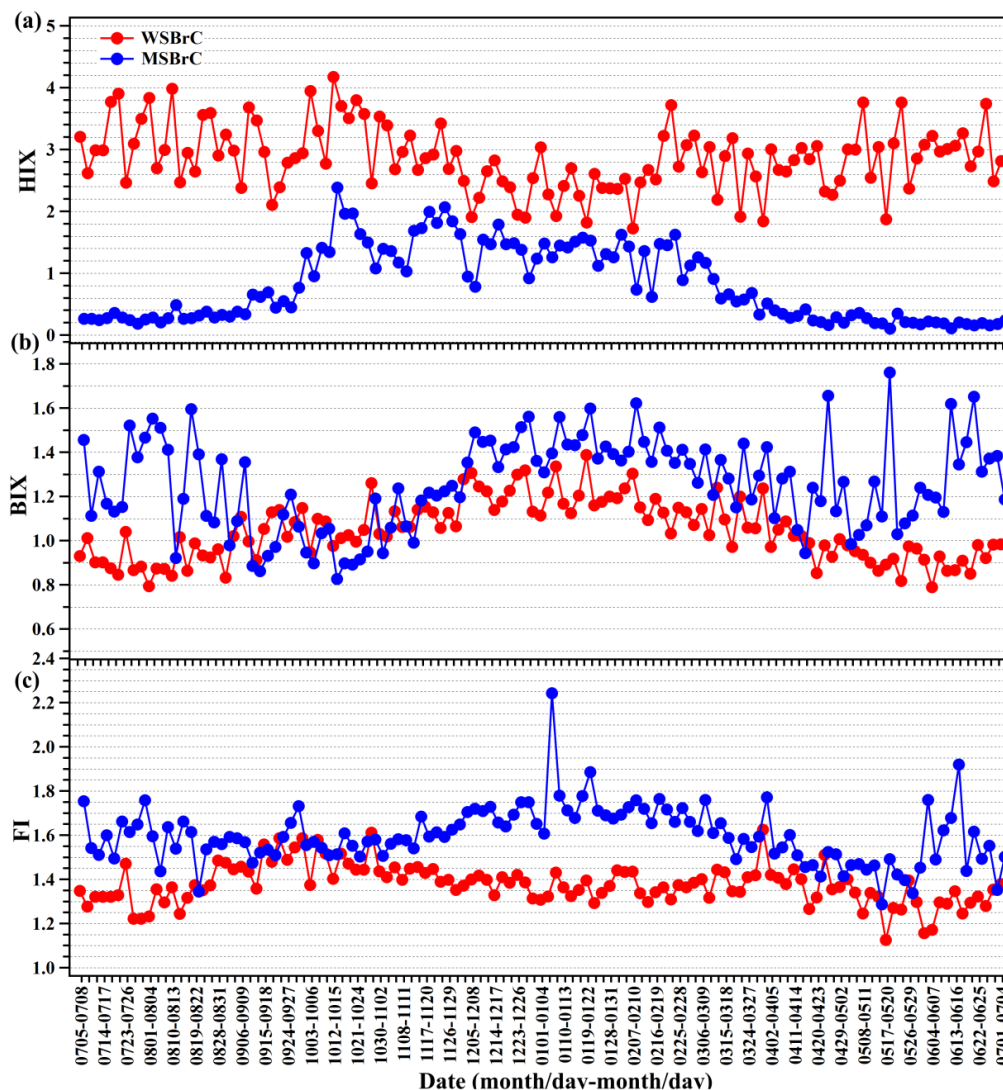


421 Tianjin PM<sub>2.5</sub> was  $1.20 \pm 0.08$  (1.03–1.39) in winter,  $1.06 \pm 0.08$  (0.83–1.26) in autumn,  $1.01 \pm$   
422  $0.11$  (0.82–1.24) in spring and  $0.91 \pm 0.06$  (0.79–1.04) in summer. Whereas the BIX of WI-  
423 MSBrC peaked in winter ( $1.43 \pm 0.09$ ) and showed the minimum value d in autumn ( $1.05 \pm 0.14$ ).  
424 The slightly rise in summer indicated that biological OM was more soluble in MeOH, which was  
425 different from BrC contributed by biological matter in autumn. Such results indicate that the  
426 chromophoros generated by primary sources in autumn were mostly soluble in water, while the  
427 chromophoros generated in summer were mostly water-insoluble, which was closely related to  
428 their molecular structure. Fig. 7 shows the seasonal variations of HIX, BIX and FI in WSBrc and  
429 WI-MSBrC.



430

431 **Figure 6.** Relative contributions of the fluorescent volumes of the WSBrc and WI-MSBrC in  
432 PM<sub>2.5</sub> from Tianjin.



433

434 **Figure 7.** Temporal variations in light absorption and fluorescence properties of BrC in PM<sub>2.5</sub>  
 435 Tianjin: (a) HIX, (b) BIX, and (c) FI.

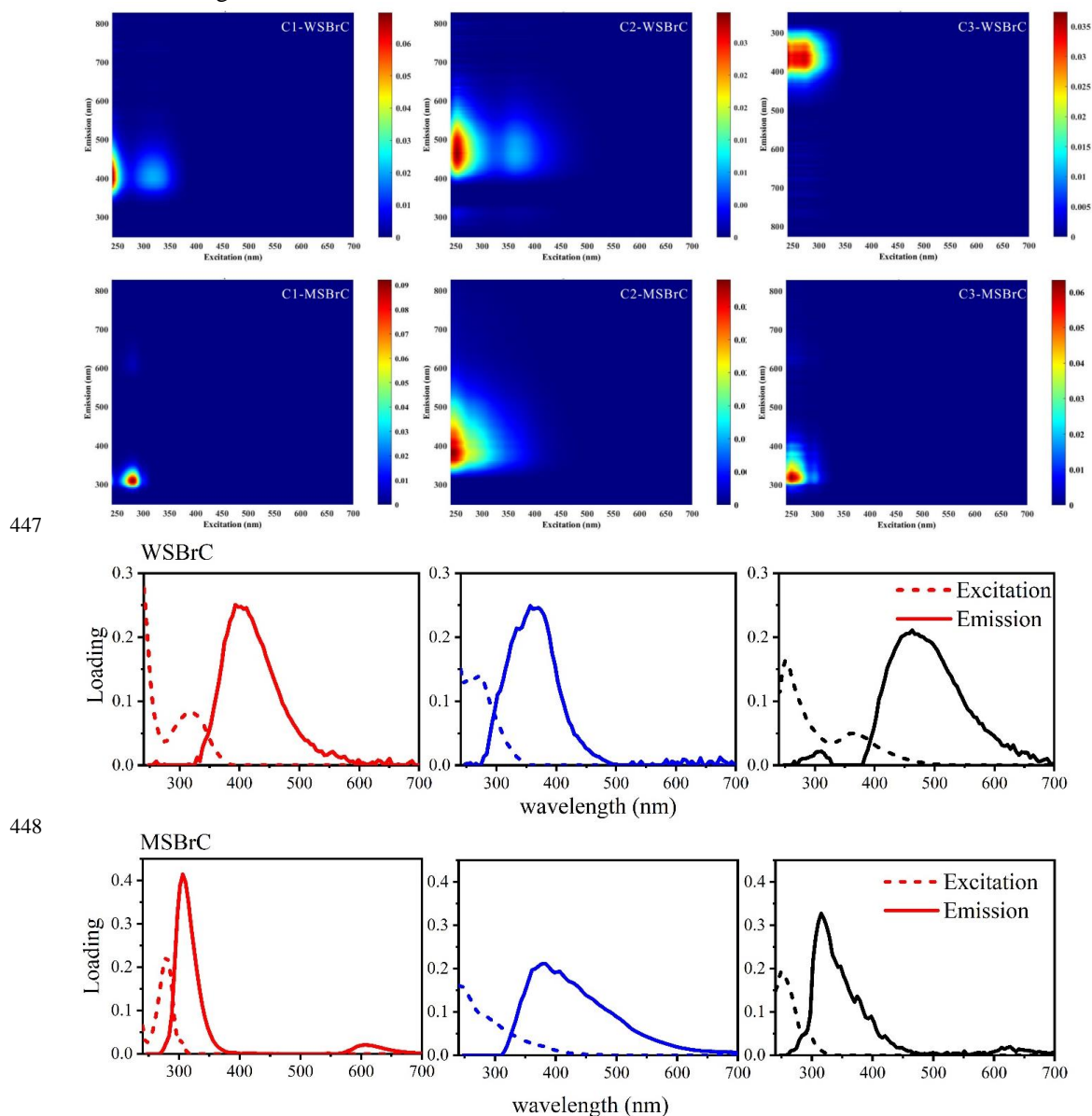
436

437 3.3.2 Chromophore identification

438 It has been reported that chromophores with different excitation emission wavelengths can  
 439 distinguish the types and sources of chromophores, but the types and sources of a large number of  
 440 chromophores have not been determined due to their complex chemical composition and sources.  
 441 Here, we successfully separated several fluorescence components from the EEM data using the  
 442 parallel factor analysis (PARAFAC) method, and the results are shown in Fig. 8. The fact of the



443 value of core consistency is close to 100 in PARAFAC indicates that the more the individual  
 444 components that are analyzed together, the more they make up 100% of the mixture, with no  
 445 unexplained residues. The core consistency of BrC extract in water was 89%, but the value can  
 446 reach 95% together with that in MeOH.



448  
 449  
 450 **Figure 8.** Three-dimensional excitation-emission matrix of three fluorescent components with  
 451 emission and excitation spectra of each fluorescent component at peak emission and excitation  
 452 wavelengths in WSBrc (above) and WI-MSBrc (below) obtained by PARAFAC model analysis.



453 A total of three types of chromophores, two with fluorescence characteristics similar to those  
 454 of humic-like substances (less oxygenated HULIS and highly oxygenated HULIS) and one with  
 455 fluorescence characteristics similar to those of protein compounds (PLOM), were identified in  
 456 Tianjin PM<sub>2.5</sub> by PARAFAC for EEMs. The types of chromophores obtained in this study together  
 457 with those from the literature are summarized in Table 2. Chromophore C1 in WSB<sub>r</sub>C has a  
 458 primary fluorescence peak at excitation/emission (Ex/Em): <240/393 nm, and a secondary  
 459 fluorescence peak at Ex/Em: 318/393 nm. C1 can be classified as a humus-like chromophore  
 460 because the bimodal distribution of the fluorescence spectrum is usually associated with humic-  
 461 like substances (HULIS). The emission wavelength of C1 was closer to the UV region than that of  
 462 the second peak of C2 in WSB<sub>r</sub>C, indicating the existence of a large number of aromatic substances,  
 463 conjugate systems and nonlinear ring systems (Deng et al., 2022). C2 (Ex/Em ~251, 363 nm/462  
 464 nm) was identified as a common HULIS in aerosols, with higher oxidation, aromatization,  
 465 molecular weight, conjugation, and unsaturation due to its larger emission wavelength (Wen et al.,  
 466 2021). The molecular weight of the fluorescent chromophore as well as its degree of conjugation  
 467 tend to increase with the excitation wavelength, and such increase in size and the conjugation  
 468 degree may be attributed to the presence of highly aromatic conjugated structures containing  
 469 heteroatoms (Chen et al., 2019). Compared to C1 and C2, C3 also contains two peaks, with shorter  
 470 wavelengths (<380 nm) emission peak, which is usually associated with PLOM such as tryptophan  
 471 and tyrosine, with low aromatic properties and small molecular size.

472  
 473 **Table 2.** Description and wavelength positions of PARAFAC components in this study and other  
 474 reports. (PLOM = protein compounds; HULIS = humic-like substances)

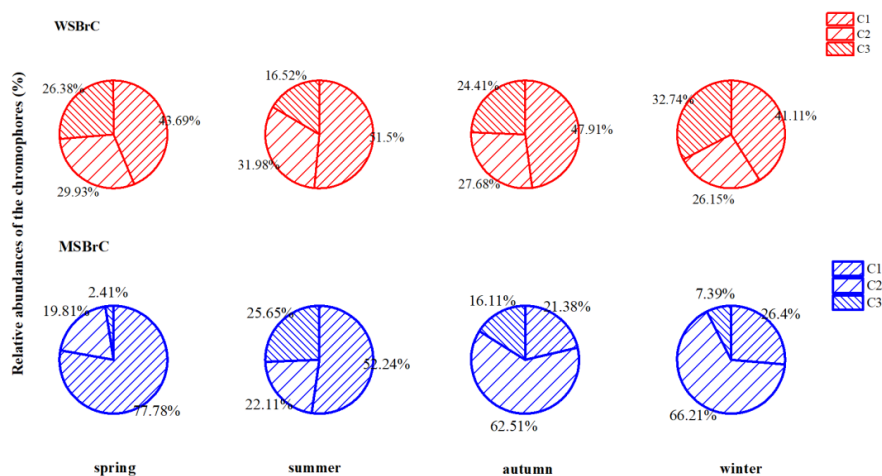
Category	Components	Ex(nm)	Em(nm)	Substances	References
WSB <sub>r</sub> C	C1	<240, 318	393	low-oxygenated HULIS	this study
	C2	251, 363	462	high-oxygenated HULIS	
	C3	<240, 271	356.3	PLOM, such as tryptophan and tyrosine	
WI-MSB <sub>r</sub> C	C1	<240, 279	306	PLOM, tyrosine-like	this study
	C2	<240	379	uncertain	
Water-soluble BrC	C3	251, 294	315	PLOM, tryptophan-like	(Deng et al., 2022)
	C1	250, 315	396	low-oxygenated HULIS	
	C2	250	465	highly-oxygenated HULIS	
	C3	250	385	low-oxygenated HULIS	
	C4	250	340	PLOM, tryptophan-like	
WSOC	C5	275	305	PLOM, tyrosine-like	(Wen et al., 2021)
	C1	240, 315	393	low-oxygenated HULIS	
	C2	245, 360	476	highly-oxygenated HULIS	
	C3	<240, 290	361	PLOM, such as tryptophan and tyrosine	
	C4	275	311	PLOM, tyrosine-like	
WSM and MSM	C1	255	415	HULIS-1 component	(Chen et al., 2019)
	C2	220	340	tryptophan-like component	
	C3	255	385	HULIS-2 component	
	C4	210	300	tyrosine-like component	
	C5	250	355	amino acid-like component	
WSOC	C1	245	410	HULIS, photodegradation of macromolecules	(Xie et al., 2020)
	C2	235	398	HULIS, aromatic and saturated compounds were presented	
	C3	250, 360	466	humic-like chromophores, more aromatic and consisted of more unsaturated compounds produced by condensation reactions	
	C4	250, 285	432	terrestrial humic-like chromophore	
	C5	<235	430	terrestrial humic-like substance, photochemical product	
MSOC	C6	275	408	low oxidation humic-like	this study
	C7	235, 275	372	protein-like chromophore	
	C8	260, 310	364	protein-like (tryptophan-like), may be related to PAHs	





475 However, MeOH-soluble chromophore C1 might be tyrosine-like substance. C2 could be  
 476 HULIS or PLOM, it's not quite certain because its emission wavelength <380nm fits the profile of  
 477 PLOM, but it is also close to the emission wavelength of HULIS. While C3 is a tryptophan-like  
 478 substance, which was reported to contain low aromatic and small molecular weight. In general,  
 479 phenols contribute significantly to C3 chromophore as they are the products of incomplete  
 480 pyrolysis of lignin and cellulose and are used as indicators of biomass burning (Wen et al., 2021).  
 481 The water-insoluble chromophores of all samples in this study can be classified as PLOM. This  
 482 indicates that the chromophores of protein substances mainly dissolve in solvents with high  
 483 polarity. As shown in Fig. 9, the water-soluble extracts contained more HULIS. In contrast, the  
 484 MeOH extracts contained more PLOM chromophores than those in the water-soluble extracts.

485 Surprisingly, according to the excitation emission wavelength, we classified the fluorescence  
 486 component of WI-MSBrC substance as PLOM, but the correlation between their fluorescence  
 487 intensity and BIX ( $R^2 = 0.06$ ,  $p < 0.05$ ) was very small, far lower than that of WSBrc substance  
 488 and BIX ( $R^2 = 0.18$ ,  $p < 0.05$ ). On the contrary, the correlation between their fluorescence intensity  
 489 and HIX ( $R^2 = 0.54$ ,  $p < 0.05$ ) was much higher than that of WSBrc ( $R^2 = 0.01$ ,  $p < 0.05$ ). Although  
 490 PLOM may be associated with some polycyclic aromatic hydrocarbons (PAHs) or phenols from  
 491 fossil fuel combustion and biomass burning, especially in urban aerosols, the correlation is  
 492 puzzling.  
 493



494

495 **Figure 9.** Relative abundances of the chromophores of the WSBrc and WI-MSBrC in PM<sub>2.5</sub>  
 496 from Tianjin.  
 497

498 Fig. 9 displays the average relative contributions of the fluorescent components of WSBrc  
 499 and WI-MSBrC during different periods. On average, the humic-like chromophores together  
 500 contributed more than 60% to the fluorescence intensity in WSBrc, suggesting that humic-like  
 501 chromophores played a dominant role in fluorescence properties of WSBrc in Tianjin. Generally,  
 502 the low-oxygenated chromophores C1 made considerable contributions in each season. C2, as  
 503 highly oxygenated HULIS, has a greater relative contribution in summer, which might be due to  
 504 the strong solar radiation in summer, which made some HULIS with little oxygen photodegraded  
 505 and form highly oxygenated HULIS through a series of oxidation reactions. In contrast, in WI-



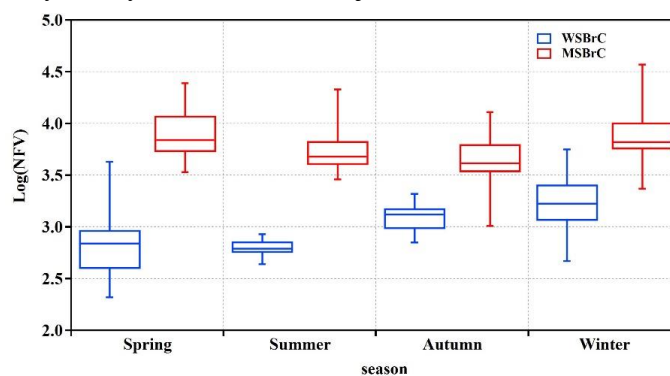


506 MSBrC, the average contribution of PLOM to fluorescence intensity was higher than 70% in  
 507 spring (80.2%) and summer (77.9%), but C2 component dominated in winter and autumn. This  
 508 indicated that biological activities increased in spring and summer and the relative abundance of  
 509 bioaerosols was higher during that period.

### 510 3.4 Potential sources of BrC

511 The types of chromophores present in aerosols are numerous and their sources are complex.  
 512 A chromophore may originate from a single source or may be contributed by multiple sources. To  
 513 explore the potential sources of BrC, correlations of FV with chemical components and light  
 514 absorption of PM<sub>2.5</sub> were studied. The total FVs of WSB<sub>r</sub>C and WI-MSB<sub>r</sub>C with SOC showed a  
 515 significant correlation in autumn ( $R^2 = 0.80$ ,  $p < 0.05$ ) and spring ( $R^2 = 0.52$ ,  $p < 0.05$ ). Furthermore,  
 516 the correlation between FVs and EC in each season was insignificant. Such relations suggest that  
 517 the secondary formation processes contributed more in autumn and spring. A good correlation  
 518 between FVs and Abs<sub>365</sub> of WSB<sub>r</sub>C and WI-MSB<sub>r</sub>C was found in all seasons, except in winter,  
 519 which indicate that most of light-absorbing materials would also have fluorescence character.

520 The relative contents of different chromophores in different polar extracts were also different.  
 521 The results showed that the overall optical properties of the different samples were different (Fig.  
 522 10). Recently, it has been reported that the aerosols derived from biomass burning and coal  
 523 combustion exhibit the highest NFV values, while SOA show the lowest NFV values (Chen et al.,  
 524 2020). NFV in all samples studied in Tianjin during 2018–2019 was very similar to that of POA  
 525 and higher than that of SOA. Such result reveal that the chromophores in the Tianjin PM<sub>2.5</sub> might  
 526 mainly be derived from a primary combustion sources. In addition, the NFVs of the Tianjin PM<sub>2.5</sub>  
 527 were higher in winter than in summer, which is likely and can be attributed to the photolysis of  
 528 chromophores in summer. In addition, NFV in MeOH-soluble OC was much higher than that in  
 529 WSB<sub>r</sub>C, which indicated that chromophores were abundant in WI-MSB<sub>r</sub>C than in the WSB<sub>r</sub>C.  
 530 Extraction of BrC by variety of solvents is a subject of our future research.



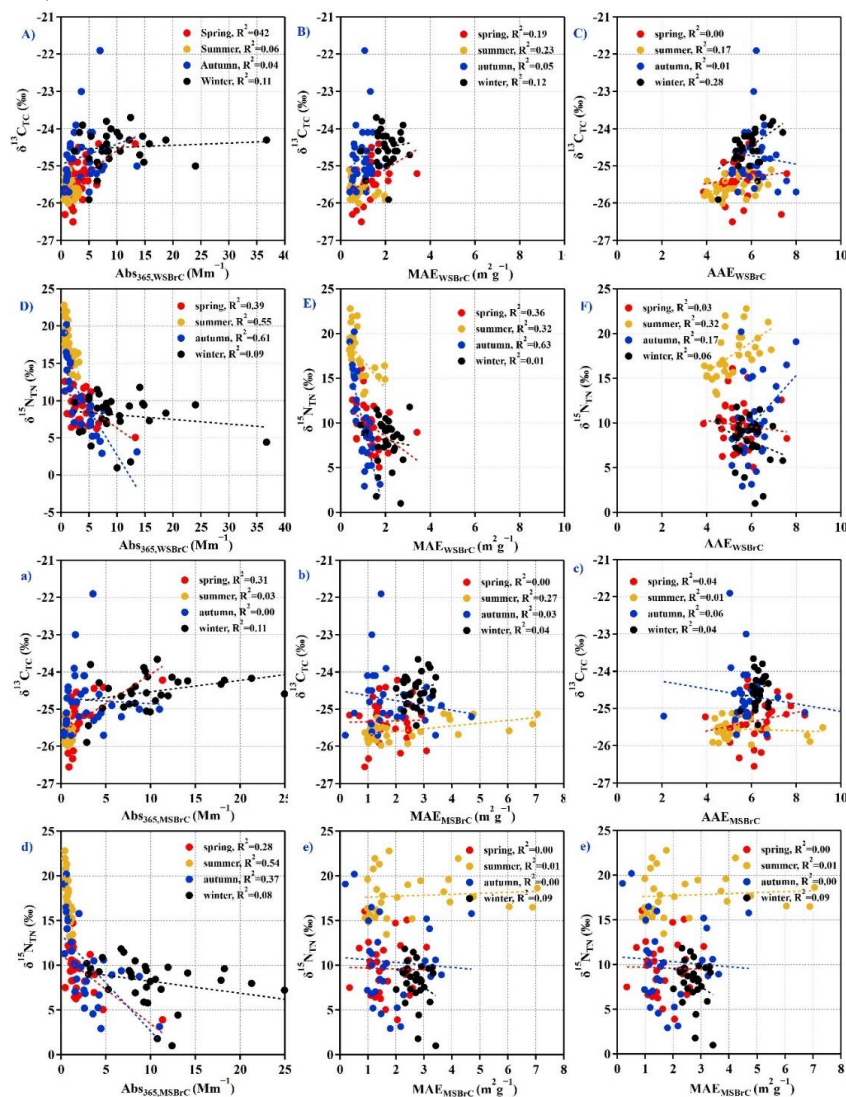
531

532 **Figure 10.** The normalized fluorescence volumes (NFVs) of the WSB<sub>r</sub>C and WI-MSB<sub>r</sub>C of  
 533 PM<sub>2.5</sub> from Tianjin, North China.

534 In order to further identify the sources of the chromophores, the correlation between  $\delta^{13}\text{C}_{\text{TC}}$   
 535 as well as  $\delta^{15}\text{N}_{\text{TN}}$  and the optical parameters of BrC were analyzed. As shown in Fig.11,  $\delta^{15}\text{N}_{\text{TN}}$  in  
 536 Tianjin PM<sub>2.5</sub> showed a negative correlation with Abs<sub>365</sub>, WSB<sub>r</sub>C in autumn ( $R^2 = 0.61$ ) and summer  
 537 ( $R^2 = 0.55$ ), whereas the  $\delta^{13}\text{C}_{\text{TC}}$  showed a significant relation neither with Abs<sub>365</sub>, WSB<sub>r</sub>C nor with  
 538 Abs<sub>365</sub>, WI-MSB<sub>r</sub>C in all seasons. Such relations indicated that some of N-containing substances



539 derived from biomass burning emissions, biological emissions and subsequent aging of aerosols  
 540 in summer and autumn might contain BrC chromophores and were soluble in water (Satish and  
 541 Rastogi, 2019).



542

543 **Figure 11.** Scatter plots between  $\delta^{13}\text{C}_{\text{TTC}}/\delta^{15}\text{N}_{\text{TN}}$  and optical parameters (Abs<sub>365</sub>, AAE, MAE) in  
 544 WSBrc and WI-MSBrc in PM<sub>2.5</sub> from Tianjin. The  $\delta^{13}\text{C}_{\text{TTC}}/\delta^{15}\text{N}_{\text{TN}}$  data is obtained from (Dong  
 545 et al., 2023).

546 **4. Summary and Conclusions**

547 This study presents the temporal variations in light absorption and fluorescence properties of  
 548 water-soluble BrC (WSBrc) and the water-insoluble but MeOH-soluble BrC (WI-MSBrc) in



549 PM<sub>2.5</sub> collected from Tianjin, North China during July 5, 2018 – July 5, 2019. Based on correlation  
550 between BrC and aerosol chemical composition, the possible sources of BrC were  
551 comprehensively analyzed. Light absorption properties of WSBrc and WI-MSBrC in Tianjin were  
552 investigated and found to be distinct from season to season, which was lower in spring and summer,  
553 compared with that in autumn and winter. The AAE of WI-MSBrC was comparable with that of  
554 WSBrc and indicated that Tianjin PM<sub>2.5</sub> contains more polar BrC. The Mass absorption efficiency  
555 of WSBrc and WI-MSBrC (MAE<sub>365</sub>) exhibited distinct seasonal variations, which was higher in  
556 winter and lower in summer and autumn. Biologically derived or secondary BrC might be one of  
557 the reasons for the lower MAE<sub>365</sub> values in summer and autumn. The light absorption of BrC in  
558 the range of 300–400 nm to the light absorption of WSBrc and WI-MSBrC in whole range (300–  
559 700 nm) was close to 40%, indicating that BrC plays an important role in climate warming and  
560 atmospheric photochemical reactions. In order to better assess its impact on climate, further work  
561 is needed through observations and laboratory simulations of BrC generation and aging.

562 In addition, based on PARAFAC analysis model, EEM data were comprehensively analyzed  
563 to compare the types and abundance of different color clusters in different aerosol samples, and a  
564 portion of source of BrC chromophore was determined by different excitation-emission  
565 wavelengths. In this study, BrC chromophore was divided into three categories: low-oxygenated  
566 HULIS, high-oxygenated HULIS and protein-like compound (PLOM). The high-oxygenated  
567 HULIS was more reactive than other categories in the atmospheric photooxidation. By comparing  
568 the fluorescence fluxes of different polar extraction solutions, it was found that WI-MSBrC  
569 substances contributed more than half of the fluorescence, indicating that there were more polar  
570 BrC substances in Tianjin aerosol, which attributed from anthropogenic emissions. The correlation  
571 between BrC optical properties and aerosol chemical composition indicated that fossil fuel  
572 combustion significantly contributed to BrC content in winter, while primary biological emission  
573 and aging reaction significantly contributed to the BrC content in summer. These results illustrated  
574 the light absorption properties of BrC in metropolis aerosols and emphasized its significant  
575 contribution to radiative forcing.

#### 576 **Declaration of competing interest**

577 The authors declare no competing interest in this paper.

#### 578 **Data Availability Statement**

579 The data used in this study can be found online at <https://doi.org/10.5281/zenodo.7316371> (Dong  
580 et al., 2022), and at <https://doi.org/10.5281/zenodo.5140861> (Dong et al., 2021).

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585 to calculate the SFE.



586 **Author contribution**

587 ZD and CMP conceptualized this study. ZD and PL conducted the sampling. ZD conducted the  
588 chemical analyses, interpreted the data and wrote the manuscript. CMP supervised the research  
589 and acquired the funding for this study. XZ, ZXY and ZXJ administrated the project. CMP, ZX,  
590 DJ, PF and CQL contributed in discussing the results and review and editing the manuscript.

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