

1 Relevance of near-surface soil moisture vs. terrestrial water storage

2 for global vegetation functioning

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14 **Abstract.** Soil water availability is an essential prerequisite for vegetation functioning. Vegetation takes up water from varying
15 soil depths depending on the characteristics of their rooting system and soil moisture availability across depth. The depth of
16 vegetation water uptake- is largely unknown across large spatial scales as a consequence of sparse ground measurements. At
17 the same time, emerging satellite-derived observations of vegetation functioning, surface soil moisture and terrestrial water
18 storage, present an opportunity to assess the depth of vegetation water uptake globally. In this study, we characterise vegetation
19 functioning through the Near-Infrared Reflectance of Vegetation (NIRv), and compare its relation to (i) near-surface soil
20 moisture from ESA-CCI and (ii) total water storage from GRACE at the monthly time scale during the growing season. The
21 relationships are quantified through partial correlations to mitigate the influence of confounding factors such as energy and
22 other water-related variables. We find that vegetation functioning is generally more strongly related to near-surface soil
23 moisture, particularly in semi-arid regions and areas with low tree cover. In contrast, in regions with high tree cover and in
24 arid regions, the correlation with terrestrial water storage is comparable to or even higher than with near-surface soil moisture,
25 indicating that trees can and do make use of their deeper rooting systems to access deeper soil moisture, similar to vegetation
26 in arid regions. At the same time we note that this comparison is hampered by different noise levels in these satellite data
27 streams. In line with this, an attribution analysis that examines the relative importance of these soil water storages for vegetation
28 reveals that they are controlled by (i) water availability influenced by the climate and (ii) vegetation type reflecting adaptation
29 of ecosystems to local water resources. Next to variations in space, the vegetation water uptake depth also varies in time.
30 During dry periods, the relative importance of terrestrial water storage increases, highlighting the relevance of deeper water
31 resources during rain-scarce periods. Overall, the synergistic exploitation of state-of-the-art satellite data products to

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32 disentangle the relevance of near-surface vs. terrestrial water storage for vegetation functioning can inform the representation
33 of vegetation-water interactions in land surface models to support more accurate climate change projections.

34 1. Introduction

35 The regulation of water, energy, and biogeochemical cycling between land and atmosphere is primarily dependent on
36 vegetation. In addition, global vegetation provides essential ecosystem services such as food production and uptake of some
37 of the anthropogenic carbon dioxide emissions (Keenan & Williams, 2018). (Keenan and Williams, 2018). Vegetation growth
38 depends on nutrient, water and energy availability. As a result, on a global scale, there are regions with energy or water limited
39 vegetation functioning (Orth, 2021). In energy-limited regions, the functioning of vegetation is controlled by radiation and
40 temperature, as they often lack sunny and warm conditions but have ample soil moisture. In contrast, soil moisture becomes
41 critical for vegetation growth in water-limited regions. Plant photosynthesis involves opening the stomata for the uptake of
42 CO₂, while at the same time water is lost through transpiration. However, in water-limited conditions, plants can reduce the
43 stomatal opening to avoid water loss, leading to a decrease in photosynthesis. Hence, variations in soil moisture are likely to
44 affect vegetation functioning in water-limited conditions. Moreover, climate change has led to an expanded water limitation
45 on vegetation (Denissen et al., 2022) and increased vegetation sensitivity to soil moisture (Li et al., 2022). For these reasons,
46 it is essential to better understand the dependence of vegetation functioning on soil moisture to comprehend their coping
47 mechanisms during drought to predict the future of global water, energy, and carbon cycles.

48
49 Plants extract water from varying soil depths based on the positioning of their roots and the availability of soil moisture; and
50 nutrients. In general, the plant water uptake depth further differs spatially across different climate regimes and vegetation
51 types, and temporally between seasons. Vegetation in arid regions is more susceptible to fluctuations in near-surface soil
52 moisture compared to vegetation in humid regions (Xie et al., 2019). Grasses, which generally have shorter roots than trees
53 and shrubs, are more reliant on near-surface moisture than deeper moisture (Schenk & Jackson, 2002). Grasses, which generally
54 have shorter roots than trees and shrubs, are more reliant on near-surface moisture than deeper moisture (Schenk and Jackson,
55 2002). Further, root water uptake profiles vary within individual plant types according to above-ground biomass and age, with
56 larger and older trees having deeper roots capable of extracting water from deeper soil layers (Schenk & Jackson, 2002; Tao
57 et al., 2021)(Schenk and Jackson, 2002; Tao et al., 2021). Additionally, within similar climate regimes, plant water uptake
58 varies across topographic positions. Upland and lowland roots tend to be shallower, making vegetation more reliant on near-
59 surface soil moisture, while roots go deeper in steep terrain between these landscapes to access both surface and deep moisture
60 (Fan et al., 2017).

61
62 Though spatial variations of plant water uptake depths across vegetation types and climate regimes, and temporal shift during
63 dry-months, are widely studied at point scale, inadequate deep soil moisture records pose a major obstacle to study vegetation

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64 root water uptake at a global scale. Microwave remote sensing allows to infer near-surface soil moisture dynamics globally.
65 However, such data may not fully represent root zone soil moisture as microwaves can only propagate through the top few
66 centimetres in soil (Capehart & Carlson, 1997). Land surface models provide an alternative source of global soil moisture data
67 across depths, but they are subject to uncertainties arising from meteorological data, inaccurate knowledge of soil and
68 vegetation characteristics, and the representation of complex processes such as photosynthesis, infiltration, and evaporation.
69 While microwaves penetrate only the top few centimeters and do not cover the entire soil moisture profile, they represent
70 larger depths of moisture variation, providing valuable insights into root zone soil moisture (Feldman et al., 2023). Land
71 surface models provide an alternative source of global soil moisture data across depths, but they are subject to uncertainties
72 arising from meteorological data, inaccurate knowledge of soil and vegetation characteristics, and the representation of
73 complex processes such as photosynthesis, infiltration, and evaporation (Koster et al., 2009; Seneviratne et al., 2010). Hence,
74 some studies have employed reanalysis-based soil moisture estimates, to investigate the relationship between vegetation and
75 soil moisture at the global (Li et al., 2021; Miguez-Macho & Fan, 2021); but those are likely to be impacted by model
76 assumptions affecting soil moisture dynamics, particularly for deeper layers where less observational constraints are available.
77 Hence, some studies have employed reanalysis-based soil moisture estimates, to investigate the relationship between
78 vegetation and soil moisture at the global scale ((Li et al., 2021; Miguez-Macho and Fan, 2021); but those are likely to be
79 impacted by model assumptions affecting soil moisture dynamics, particularly for deeper layers where less observational
80 constraints are available. Thus, studying vegetation interactions with the entire water column, including near-surface and deep
81 soil moisture, at a global scale using exclusively observation-based dataset is imperative to enhance the understanding of
82 relevance of near-surface and deep soil moisture for vegetation functioning.

83
84 The Gravity Recovery and Climate Experiment (GRACE) satellite mission, launched in 2002, provides total water storage
85 (TWS) anomalies observations at the global scale and offers a unique opportunity to investigate the relationship between
86 vegetation and the total water column. Furthermore, the inter-annual carbon dioxide growth rate in the atmosphere has been
87 found to be well correlated with the total water storage anomalies on a global scale, underlining the relevance of total water
88 column for vegetation functioning (Humphrey et al., 2018). The Gravity Recovery and Climate Experiment (GRACE) satellite
89 mission, launched in 2002, provides total water storage (TWS) anomalies observations at the global scale. The TWS captures
90 not only soil water but also snow and ice, canopy water, surface water and groundwater. Its depth of representation is therefore
91 difficult to physically quantify, and that is why we studies TWS anomalies. Nevertheless, they seem to be related to variations
92 of overall water availability (near-surface + deep soil moisture) for vegetation (Yang et al., 2014). The inter-annual carbon
93 dioxide growth rate in the atmosphere, for example, has been found to be well correlated with the total water storage anomalies
94 on a global scale, indicating the relevance of total water column for vegetation functioning (Humphrey et al., 2018). The TWS
95 captures not only soil water but also snow and ice, canopy water, surface water, and groundwater. In this study, we assume
96 that TWS anomalies can be used to estimate the variation of overall water availability (near-surface + deep soil moisture) for

97 vegetation under (i) snow-free conditions, and assuming that (ii) water storage variations in lakes or groundwater are negligible
 98 at the monthly time scale, (iii) and canopy water storage is much smaller than soil water storage and hence also negligible.
 99

100 This study focuses on understanding the relevance of near-surface vs. total water storage for vegetation functioning on a global
 101 scale using observation-based datasets, thereby inferring vegetation's large-scale water uptake depth from observation-based
 102 datasets. For this purpose, we utilise TWS and near-surface soil moisture and correlate them with vegetation functioning,
 103 represented by Near-Infrared Reflectance of Vegetation (NIRv). In particular, we analyse (1) what is the relevance of near-
 104 surface soil moisture vs. the terrestrial water storage for vegetation functioning?, (2) how does the importance of near-surface
 105 soil moisture vs. terrestrial water storage change during dry months? and (3) how do climatic, vegetation, and topographic
 106 characteristics explain the variability in the relevance of near-surface vs. terrestrial water storage for vegetation functioning?

107 2. Data and Methodology

108 **Table 1: Table summarising all the datasets.**

<u>Datasets</u>	<u>Variables</u>	<u>Source</u>	<u>Spatial</u> <u>Resolution</u>	<u>Temporal</u> <u>Resolution</u>	<u>Temporal</u> <u>Coverage</u>	<u>References</u>
Vegetation Functioning	Near Infrared Reflectance of Vegetation (NIRv)	MODIS/MOD13C1 v061	0.05 degree	<u>16 daily</u>	<u>2000 - present</u>	(Badgley et al., 2017)
	Solar Induced Chlorophyll Fluorescence (SIF)	GOME-2	0.5 degree	<u>16 daily</u>	<u>2007 - 2018</u>	(Köhler et al., 2015)
Soil Water Storage	Near-surface soil moisture (SSM)	ESA-CCI v04.4	0.25 degree	<u>Daily</u>	<u>1978 - 2022</u>	(Dorigo et al., 2017)
	Total Water Storage (TWS) Anomalies	GRACE	0.5 degree	<u>(Landerer & Swenson, 2012) Monthly</u>	<u>2002 - present</u>	<u>(Landerer and Swenson, 2012)</u>
Meteorological	Air Temperature (T _a)	ERA-5	<u>0.25 degree</u>	<u>Hourly</u>	<u>1940 - present</u>	(Hersbach et al., 2020)

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	Precipitation (P)					
	Net Radiation (R _n)					
	Dew point Temperature (T _d)					
Climatological	Aridity Index	Global Aridity Index and Potential Evapotranspiration Database - Version 3	30 arc seconds	Static	1970-2000	(Zomer et al., 2022)
Vegetation and Land cover class	Tree cover fraction	VFC5KYR	0.05 degree	(Hansen, Matthew & Song, Xiao-Peng, 2018)	1982 - 2016	(Hansen, Matthew and Song, Xiao-Peng, 2018)
	Land cover data	ESA-CCI	0.05 degree, 300 m	Yearly	1992 - 2018	ESA. Land Cover CCI Product User Guide Version 2. Tech. Rep. (2017)
Topographical data	Elevation	Earthenv	1 km	(Amatulli et al., 2018) Static		(Amatulli et al., 2018)
	Slope					
Soil data	Fraction of sand	FAO	0.05 degree	(Reynolds et al., 2000) Static		(Reynolds et al., 2000)
	Fraction of clay					
Irrigation	Percentage of Irrigated area	HID	5 arcmin	(Siebert et al., 2015) Yearly	1990 - 2005	(Siebert et al., 2015)

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110 2.1 Data

111 2.1.1 Vegetation Functioning:

112 In our study, vegetation functioning is characterised by satellite measurements of Near-Infrared Reflectance of vegetation
 113 (NIRv) and Solar Induced Fluorescence (SIF) (Table 1). NIRv is the product of near-infrared reflectance and the normalised
 114 difference vegetation index (NDVI) and represents the vegetation structure and vegetation greenness (Badgley et al., 2017).

115 The NIRv data is available at a high spatial resolution of 0.05°, and the original 16-day data was aggregated to the monthly

116 NIRv data. SIF is directly related to the photosynthetic activity of plants because the excess energy from sunlight, that triggers
117 the light reaction during photosynthesis, is dissipated by leaf as chlorophyll fluorescence (Mohammed et al., 2019). SIF data
118 is derived from the Global Ozone Monitoring Experiment (GOME-2), because GOME-2 provides relatively reliable data over
119 a long period (2007-2018). The 0.5° spatial and 16-day temporal resolution SIF data is processed into monthly data as described
120 by (Köhler et al., 2015).

121
122 The high spatial resolution of NIRv allows for a detailed study of the correlation of vegetation functioning with soil water
123 availability. Therefore, we performed the main analyses using NIRv data. However, SIF is more sensitive to drought stress
124 than NIRv (Qiu et al., 2022). Therefore, we perform additional analyses with SIF to show that the relationships hold for a
125 different and more direct indicator of vegetation functioning.

126 2.1.2 Soil Water Storage

127 This study includes two different measures of soil water availability. The near-surface soil moisture (SSM) provides an
128 estimate of water availability in the top layer of the soil, while the Terrestrial Water Storage (TWS) Anomaly provides an
129 estimate of the overall water column of the soil. The SSM data is derived from the European Space Agency (ESA) Climate
130 Change Initiative Program (CCI), which combines active and passive satellite microwave measurements to provide reliable
131 estimates of SSM (Dorigo et al., 2017). The ESA CCI soil moisture data, at a daily temporal resolution, was aggregated to
132 monthly temporal resolution. ~~The TWS Anomaly data is derived from the GRACE mission, which measures changes in the~~
133 ~~Earth's gravity field (Landerer & Swenson, 2012). Here, we use the JPL-Mascons product of TWS Anomalies which is~~
134 ~~available at a 0.5° spatial and monthly temporal resolution. The TWS Anomaly data is derived from the GRACE mission,~~
135 ~~which measures changes in the Earth's gravity field (Landerer and Swenson, 2012). Here, we use the JPL-Mascons product of~~
136 ~~TWS Anomalies which is available at a 0.5° spatial and monthly temporal resolution (Watkins et al., 2015)~~

137 2.1.3 Meteorological Data

138 Employed climate variables include monthly air temperature (T_a), 2m dew point temperature (T_d), precipitation (P), and net
139 radiation (R_n) from the ERA5 reanalysis products at a 0.25° spatial resolution. ~~The~~The vapor pressure deficit (vpd) is calculated
140 from T_a and T_d . Further, the aridity index is calculated from the ratio between the long-term mean R_n (mm y^{-1}) (1 MJ/sq.m/day
141 $= 0.408 \text{ mm/day}$) and P (mm y^{-1}) for each grid cell (Budyko, 1974). We opted for this formulation as it offers a direct estimation
142 of aridity and water (energy) constraints on vegetation. This eliminates the necessity to navigate through various formulations
143 utilized for calculating potential evapotranspiration. However, we conducted additional validations of our results using the
144 Global Aridity Index dataset (Zomer et al., 2022) based upon the FAO Penman-Monteith Reference Evapotranspiration
145 equation. The use of the Global Aridity Index did not change the results of our study (Section 3.4). In addition, the mean and
146 standard deviation of the climate variables are calculated and incorporated in the attribution analysis (Section 2.2.3).

2.1.4 Vegetation, soil, and topography data

To evaluate the resulting correlation of vegetation functioning and water storages with respect to vegetation characteristics, we employ the tree cover fraction data from the AVHRR vegetation continuous fields products (VCF5KYR, <https://lpdaac.usgs.gov/products/vcf5kyrv001/>) (Hansen, Matthew & Song, Xiao-Peng, 2018)(Hansen, Matthew and Song, Xiao-Peng, 2018). For this purpose, the mean of tree cover fraction for the years between 2007 and 2016 is calculated.

Topographical variables such as elevation and slope are incorporated along with other ~~climatological~~ meteorological variables to determine the relative contribution of different variables to the correlation between vegetation functioning and water storage. Topographic data at a 5 km resolution were downloaded from the EarthEnv. These data are calculated based on the 250 m GMTED dataset, and compared against the 90 m SRTM 4.1 dev dataset. The data were resampled to a coarser resolution of 5 km using various aggregation techniques, details of which are in (Amatulli et al., 2018). Furthermore, for each grid cell, the fraction of sand and clay in soil (Reynolds et al., 2000)(Reynolds et al., 2000) along with the percentage of irrigated area (Siebert et al., 2015)(Siebert et al., 2015) were considered in attribution analysis.

2.2 Methodology

2.2.1 Data pre-processing

A flowchart of the data pre-processing and analyses is presented in **Figure S1**. The time period of analysis is from 2007 to 2018 constrained by the concurrent availability of all involved datasets. All the analyses were performed in monthly temporal resolution and at 0.05° spatial resolution (for NIRv) and 0.5° spatial resolution (for SIF). The SSM and TWS data were initially available at 0.25° and 0.5° resolution, but were disaggregated or aggregated to 0.05° or 0.5° degrees, depending on the spatial resolution of the analysis performed, based on the assumption that the soil water storage anomalies are representative over larger areas. Also, the meteorological data and vegetation, soil, and topographic data were resampled into the same resolution. After aggregating all the datasets to 0.05° resolution, the monthly anomalies were calculated by subtracting the long term mean monthly cycle and by removing linear trends. A SIF threshold was applied in each grid cell to filter out non-growing season data. For this purpose, we filtered out all the months from 2007-2018 when the mean-monthly SIF value was below the threshold of 0.2 mW/m²/sr/nm. We apply an additional temperature threshold ($T_a > 5^\circ\text{C}$) to remove the months with frozen soil and snow cover, similar to (Li et al., 2021)(Li et al., 2021). Last, all months with missing soil water storage or vegetation functioning records were excluded.

175 **2.2.2 Calculate the relevance of near-surface (SSM) soil moisture and terrestrial water storage (TWS) for vegetation**
176 **functioning**

177 We calculated the ~~spearman~~Spearman correlation between vegetation functioning (NIRv) and soil water storages (SSM and
178 TWS) for each grid cell during growing season months when observations for at least 40 months were available. The cutoff of
179 40 months was implemented to guarantee a substantial number of observations for growing-season months in each grid cell.
180 This consideration assumes that the minimum number of growing-season months varies from 3 to 4 months per year globally.
181 In addition to soil moisture, also air temperature (T_a) and net radiation (R_n) affect the vegetation functioning. Moreover, SSM
182 (soil moisture) and TWS (total water storage) demonstrate a notable correlation, as illustrated in Figure S2, signifying the
183 presence of mutual information. To ~~focus~~exclusively ~~on~~examine the ~~effects~~individual impacts of each water
184 availability~~storage~~ variable on vegetation functioning and disentangle mutual information from other water variables, we
185 ~~corrected~~accounted for the confounding effects of ~~T_a and R_n~~ , by This entailed computing the partial correlation between NIRv
186 and water storages (SSM or TWS), while controlling for ~~T_a and R_n~~ , R_n , and the other water storage variable (TWS or SSM).
187 Since we focus on understanding the role of soil moisture on vegetation functioning, which is primarily critical in water-limited
188 conditions, we removed the grids cells with insignificant ($p < 0.05$) and negative partial correlations from our analysis. Such
189 negative partial correlations may hint at vegetation's converse effect on soil moisture (when increasing vegetation activity
190 depletes the soil moisture) and a negative correlation could occur in the grid cells where water limits vegetation productivity
191 through oxygen limitation (Ohta et al., 2014).

192
193 It is important to note that we chose not to apply a significance criterion in analyzing the partial correlation between NIRv and
194 water storages. When controlling for both water storage (TWS or SSM) and energy variables (T_a and R_n) in the partial
195 correlation (NIRv~SSM or TWS), a limited number of grid cells demonstrate significant correlation globally, given the high
196 correlation between SSM and TWS (Figure S2). This poses challenges for drawing global inferences on vegetation water
197 uptake. However, our overarching goal is to discern variations in the partial correlation of NIRv with water storages across
198 differing climate-vegetation gradients and how it changes from the growing season to dry months, rather than confirming
199 specific statistical thresholds. For this, we want to maintain a sufficient amount of grid cells necessary for making global
200 inferences. However, to ensure that our results are not affected by the significance criterion, we conducted additional analyses
201 considering only grid cells with a significant partial correlation (though a very small number compared to the total grid cells
202 available for each AI-TC class globally), as described in section 3.4.

203
204 The impact of all pre-processing steps on the number of grid cells included in this study is illustrated in Figure S3. Generally,
205 our filtering procedures enable us to concentrate primarily on water-limited regions, as they effectively remove a substantial
206 number of grid cells from the wet regions globally.

207 To analyse how the importance of SSM and TWS changes during dry months, we specifically selected the months
208 with~~characterized by~~ the lowest 10% SSM for each grid cell~~-,~~ representing the driest conditions within the growing-season

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209 months. The partial correlations between NIRv and water storages, $r(\text{NIRv} \sim \text{SSM})$ and $r(\text{NIRv} \sim \text{TWS})$ were calculated
210 separately for dry months. To focus on vegetation response to similar extent of dryness spatially, only grid cells with greater
211 than 100 monthly observations were considered for the dry months analysis. In addition, only the grid cells which had
212 significant positive partial correlation in growing season months were included for the dry months analysis.
213

214 After computing the partial correlations, we grouped the grid cells by aridity and tree cover ~~or land cover classes and calculated~~
215 ~~the mean correlation, for each aridity tree cover class with sufficient number of observations for both growing season and dry~~
216 ~~months. This classes, which~~ allowed us to analyse the evolution of correlations and the difference between the partial correlation
217 across aridity ~~and vegetation classes tree cover classes. Afterwards, we employed bootstrapping with 1000 repetitions to~~
218 ~~compute the bootstrap means and confidence interval using a full bootstrapping methodology (resampling with replacement~~
219 ~~from the original data) for each aridity-tree cover class with sufficient number of observations for both growing season and~~
220 ~~dry months.~~
221

222 Moreover, to test the robustness of the results, we did additional partial correlation analyses, for which we correlated the SIF
223 (instead of NIRv) with SSM and TWS. The analyses with SIF were performed at a spatial resolution of 0.5° , at which SIF data
224 was available.

225 2.2.3 Attribution Analysis

226 We used a random forest model to understand the spatial variability in the relevance of SSM versus TWS for NIRv. Random
227 forest is a nonparametric based regression algorithm which does not require any statistical assumptions on the predictor and
228 target variables which makes it particularly useful for detecting the nonlinear relationship (Breiman, 2001). Given potential
229 nonlinear impacts of various factors (climate, soil types, vegetation) on the relationship between moisture storages and
230 vegetation functioning, this study employed the random forest method to assess the relative contributions of these variables.
231

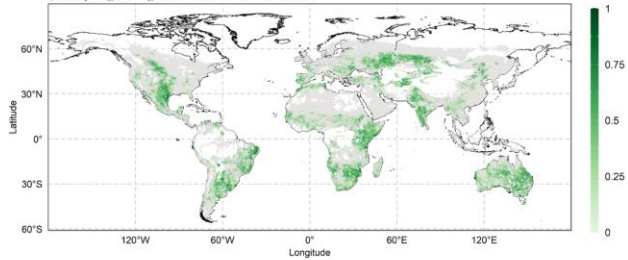
232 In our study, ~~1315~~ predictors were included in the random forest model based on their potential physical relevance to the target
233 variable, which is the difference in correlation between SSM and TWS with NIRv in growing season months. These predictors
234 included mean and standard deviation of climate variables (T_a , R_a , P and P_{vpd}), aridity index, topographical variables
235 (elevation and slope), vegetation variable (tree cover), soil-related variables (fraction of clay and sand), and percentage of
236 irrigated areas for each grid cell. We calculated the mean and standard deviation of the climate variables only during the
237 ~~growing-season months (, as determined for the months with $\text{SIF}_{\text{mean-monthly}} < 0.2 \text{ mW/m}^2/\text{sr/nm}$ were excluded).subsequent~~
238 ~~partial correlation analysis.~~ Furthermore, only the grid cells ~~having significant and exhibiting~~ positive partial correlation
239 between NIRv and SSM as well as NIRv and TWS during growing season-months were included in the random forest analysis.
240 For training a random forest model, we used the “xgboost” package in R (~~Chen & Guestrin, 2016~~); (Chen and Guestrin, 2016).
241

242 We further incorporate SHAP (SHapley Additive exPlanations) values for interpreting the predictions of the random forest
243 model (Lundberg et al., 2020). The SHAP value for a feature is the average difference in prediction of the model when that
244 feature is included compared to when it is excluded, over all possible combinations of features. By calculating SHAP values
245 for each feature in the model, we identified which features were most important in explaining the spatial variability in the
246 relevance of SSM versus TWS. For calculating the SHAP values, we employed “SHAPforxgboost” package in R.

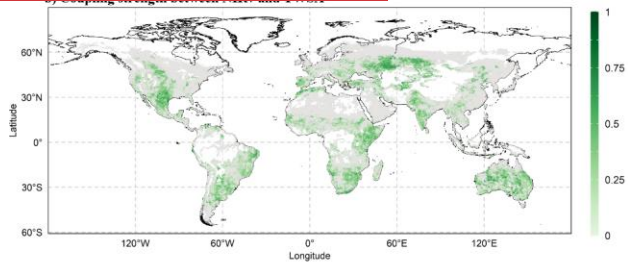
247 **3. Results and Discussion**

248 **3.1 Coupling of vegetation functioning with surface soil moisture and total water storage in the growing season**

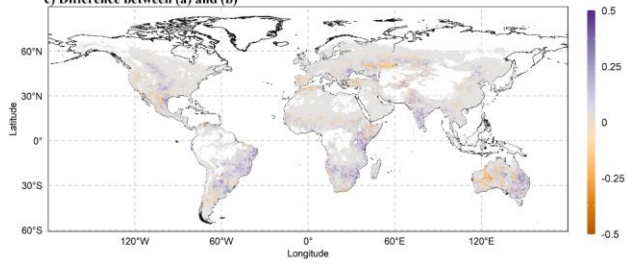
a) Coupling strength between NIRv and SSM

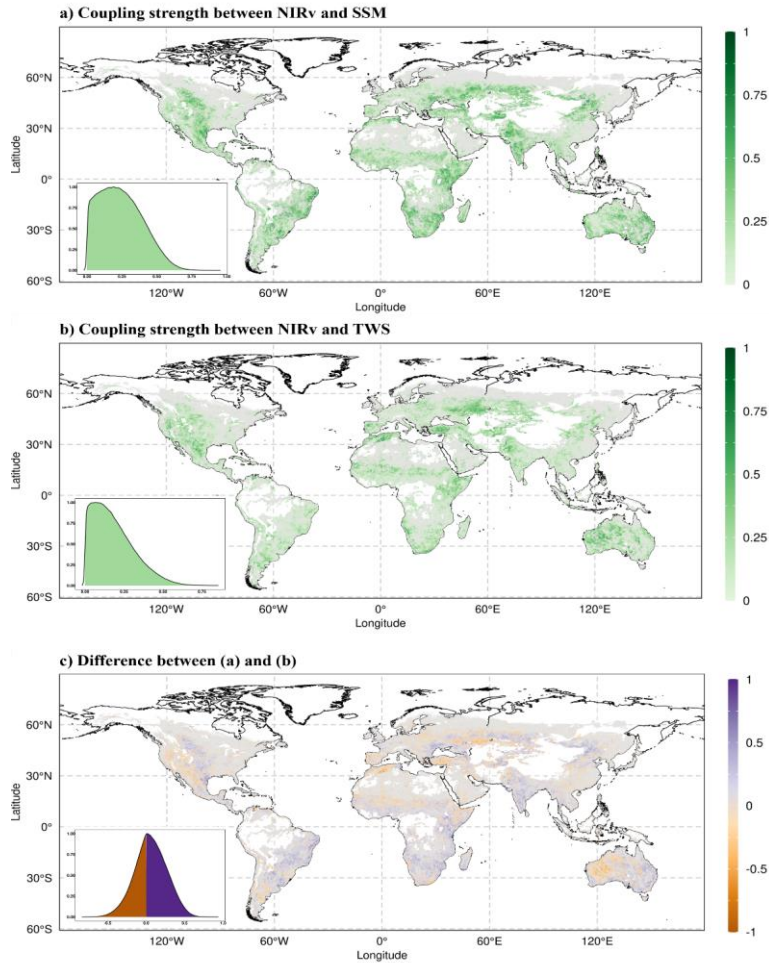


~~b) Coupling strength between NIRv and TWSA~~



c) Difference between (a) and (b)





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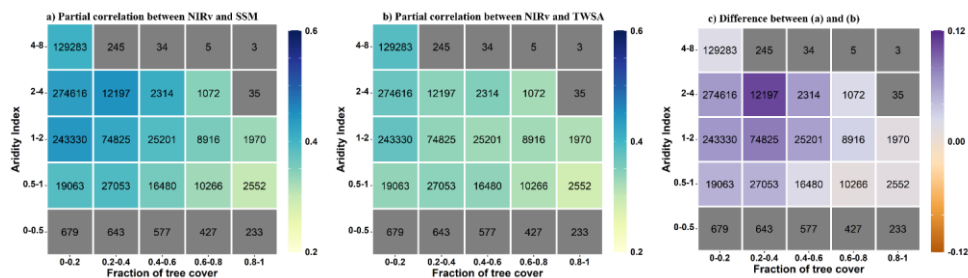
Figure 1: Coupling strength between vegetation functioning (NIRv) and (a) near-surface soil moisture (SSM), and (b) total water storage (TWS) during the growing season months. Monthly anomalies of all variables are used to calculate the partial correlation. (c) Difference between (a) and (b). The purple colour in (c) indicates the greater partial correlation of NIRv with SSM compared to the partial correlation of NIRv with TWS while orange colour indicates the opposite. Grid cells with **significant ($p < 0.05$)** and positive relationships for both correlations (a) and (b) are shown in (c) with blueish and orange colours. Light grey colour indicates **insignificant and/or** negative partial correlations between NIRv and water storage. **The absence of color within the land boundary**

257 signifies inadequate observational data for precise computation of the partial correlation. Each inset in the respective maps
 258 illustrates the probability distribution function (pdf) of the correlations.

259 The partial correlation of NIRv with near-surface soil moisture varies globally during growing-season months (**Figure 1a**).
 260 NIRv demonstrates stronger correlation with near-surface soil moisture within semi-arid climates, Central North America,
 261 South America, regions in South Africa and Australia. The correlation is stronger in Southern Europe and the Mediterranean
 262 region compared to central and Northern Europe. The correlation gradient from the hot and dry Mediterranean region to wet
 263 and cold Northern Europe corresponds to the gradient of water-limited ecosystems to energy-limited ecosystems obtained in
 264 other studies (Denissen et al., 2022; Teuling et al., 2009).

265
 266 The global correlation of NIRv with TWS follows a similar pattern as with SSM (**Figure 1b**) in growing-season months. The
 267 correlation of NIRv with TWS is higher in drier central northern America and Australia compared to other regions. The
 268 similarities in the correlation of NIRv with SSM and TWS are expected because the monthly anomalies of SSM and TWS are
 269 highly correlated during growing season months in most of our study area (**Figure S2**).

270
 271 The difference between the partial correlation of NIRv with SSM and TWS (**Figure 1c**) indicates that the NIRv correlates
 272 stronger with TWS in Western America, Southern Europe, and arid regions of Australia compared to other regions globally
 273 during growing-season months. In South America and Southern Africa, however, the NIRv shows a stronger correlation with
 274 SSM. It is difficult to determine which Although we control for the effect of soil water storage (SSM or TWS) is more critical
 275 when computing partial correlation to discern the relative importance for vegetation, because the near-surface soil moisture is
 276 included it should be noted that the varying noise levels inherent in the measurements of TWS, and both these datasets have
 277 very different noise levels might impact our results.



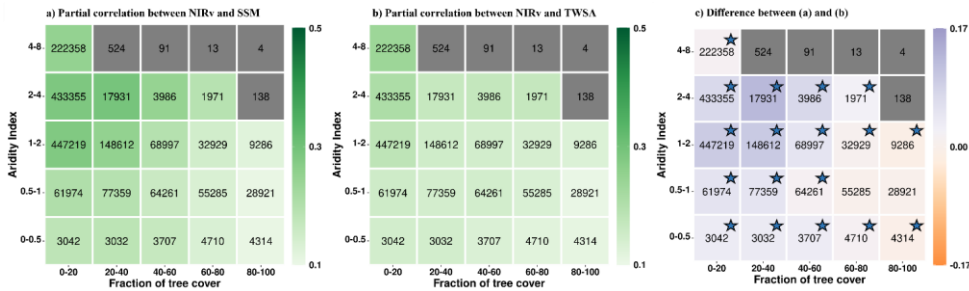


Figure 2: Summarising the coupling strengths of vegetation functioning (NIRv) with (a) near-surface soil moisture (SSM) and (b) terrestrial water storage (TWS) in the growing season-months across climate (aridity index) and vegetation regimes (fraction of tree cover). (c) shows the difference between (a) and (b). Numbers within the boxes denote the number of grid cells for each aridity-tree cover class. Aridity-tree cover classes containing less than 1000 grid cells are shown in grey. The colour bar indicates the mean partial correlation for each class, computed from bootstrapping. The asterisk in figure (c) signifies that the 95% confidence interval (lower and upper) shares the consistent sign (+/-) in the difference of partial correlation. Only grid cells with significant ($p < 0.05$) and positive relationships are considered.

Next, we analyse the partial correlation between NIRv and soil water storages across different aridity and tree cover fraction classes during growing season months. For this, we group the grid cells into different aridity and tree cover fraction classes and then do bootstrapping to compute mean partial correlation and the 95 percent confidence intervals for each class with more than 1000 grid cells. We find that the partial correlation of NIRv with SSM (Figure 2a) increases with increasing aridity for aridity index (0.5-4). This can be attributed to the intensification of water stress on vegetation under increasingly arid conditions, resulting in a stronger correlation between NIRv and SSM. However, for a further increase in aridity (4-8), the strength of the correlation of NIRv with SSM declines. This is due to a low soil moisture availability and low temporal variability under extremely arid conditions (Figure S7S4). The pattern of increasing correlation along aridity index is also observed in the partial correlation between NIRv and TWS. (Figure 2b).

Furthermore, the correlation of NIRv with SSM decreases for higher tree cover fractions (Figure 2a). However, such a gradient along tree cover fraction is less pronounced in the partial correlation of the NIRv with TWS (Figure 2b). This overall depicts that the coupling of vegetation functioning with SSM is generally higher for non-forested areas compared to forested areas while this gradient is less pronounced in the case of TWS.

It is difficult to conclude which soil moisture storage (SSM or TWS) is more important for a certain aridity-vegetation class because it inherently includes the difference in noise levels associated with SSM and TWS. However, we can compare the

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Commented [1]: Not sure if I understand correctly: for one box you can not conclude the importance of SSM or TWS but for the gradient change across boxes you can.

Repeating my previous comment, the result is not only about the observational noises but also the natural variability of two water objectives. But here the writing is more like that results are not explainable because they are born from noises. So you may need some references if they indicate any quantification of observational noises.

I'm also wondering if controlling TWS in a partial correlation of NIRv and SSM make some senses, compared to direct calculating the coefficient difference in lessening the noise issue.

Commented [2]: Yes the variability do play role, that's why we add Figure S7 for that.

Commented [3]: you may need to add the background about different noise level for ESA SM and GRACE

Commented [4]: Yes, and I am not sure how we do that? However, we have heatmaps with sd of SSM and TWSA as well as NIRv (Figure S7)

307 evolution of the gradient along tree cover or aridity index and assert how the relevance of SSM and TWS changes with varying
308 tree cover or aridity index. Taking this into account, we find that NIRv correlates more strongly with near-surface soil moisture
309 compared to terrestrial water storage in semi-arid regions with low tree cover (Figure 2c), suggesting that the vegetation
310 preferentially takes up water from SSM whenever available to meet its transpiration demand. This might be due to lower
311 energy expenditure on root water uptake, abundant nutrients and reduced chance of root water logging in the near-surface soil
312 moisture (Andrew Feldman et al., 2022; Schenk & Jackson, 2002; Tao et al., 2021). Conversely, the correlation between the
313 NIRv and TWS in arid areas (AI 4-8) and regions with a high fraction of tree cover is equivalent to or greater than that of
314 SSM, suggesting that trees can utilise their extensive root systems to access deeper soil moisture, as observed in arid vegetation.
315 This is consistent with previous studies reporting that the vegetation dependence on sub-surface soil moisture is higher in arid
316 and seasonal-arid climates (Miguez-Macho & Fan, 2021).

317
318 Though the difference in inherent noise levels associated with SSM and TWS impacts partial correlation analysis, we can
319 compare the evolution of the gradient along tree cover or aridity index and assert how the relevance of SSM and TWS changes
320 with varying tree cover or aridity index, assuming that the noise levels are similar across varying AI-TC classes. Taking this
321 into account, we find that NIRv correlates more strongly with near-surface soil moisture compared to terrestrial water storage
322 in semi-arid regions with low tree cover (Figure 2c), suggesting that the vegetation preferentially takes up water from SSM
323 whenever available to meet its transpiration demand. This might be due to lower energy expenditure on root water uptake,
324 abundant nutrients and reduced chance of root water logging in the near-surface soil moisture (Feldman et al., 2023; Schenk
325 and Jackson, 2002; Tao et al., 2021). Conversely, the correlation between the NIRv and TWS in arid areas (AI 4-8) and regions
326 with a high fraction of tree cover is equivalent to or greater than that of SSM, suggesting that trees can utilise their extensive
327 root systems to access deeper soil moisture, as observed in arid vegetation. This is consistent with previous studies reporting
328 that the vegetation dependence on sub-surface soil moisture is higher in arid and seasonal-arid climates (Miguez-Macho and
329 Fan, 2021). However, in certain regions with higher tree cover in humid areas, specifically with AI 0.5-1, such conclusions
330 cannot be confidently drawn statistically. The reason is that the confidence intervals for the difference in partial correlation of
331 NIRv with SSM and TWS fluctuate between positive (indicating greater relevance of SSM) and negative (indicating greater
332 relevance of TWS) values (Figure 2c).

333
334 Note that while our analysis focuses on regions with water-controlled vegetation as denoted by significantly positive
335 correlations between NIRv and the considered soil water storages, some of these grid cells are located in comparatively wet
336 climate regimes with aridity index values between 0.5 and 1 (Figure 2). This highlights the relevance of non-climatic factors
337 such as soil and vegetation types or topography in determining vegetation-water relationships in addition to the climate regime.
338 Next to this, in Figure 2c it seems that the relevance of terrestrial water storage is comparatively higher in wet climate (aridity
339 0.5-1) than in transitional climate regimes (aridity 1-2) as shown with the smaller correlation differences. This, however, is
340 probably not the case and simply a reflection of reduced variability in surface soil moisture (Figure S7S4).

Commented [5]: please indicate which aridity box are you mentioning? as i check, the box of tree 0.8-1 has too little aridity groups to suggest aridity higher TWS importance, and for tree 0.6-0.8 with higher aridity there are stronger importance for SSM.

Commented [6]: Done

Commented [7]: also another point relevant to the abstract and here, when you mention arid regions and regions with high tree fraction, you mean the intersection of them or not. i'm easily confused on it.

Commented [8]: Only to the combination of these two, right? Maybe make it a little clearer (orange seem to be only 3 grid cells am i right?)

Commented [9]: But your findings only show that for high tree cover so is it really in line with the reference then?

Commented [10]: Or are you analyzing here 2b solely, so that without looking at SSM but TWS only that the arid pixels show a stronger correlation compared to the less-arid ones in 2b?

Commented [11]: Not based on 2c.

341
342
343
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345

To ascertain that our results are not impacted by outliers, we analysed the heatmaps with 10th and 90th percentile correlation values for each aridity-vegetation class, instead of the mean correlation value (Figure S3). This shows consistent patterns of the partial correlation of NIRv with soil water storages as in Figure 2 and indicates that the gradients with tree cover and aridity are valid throughout the entire dataset.

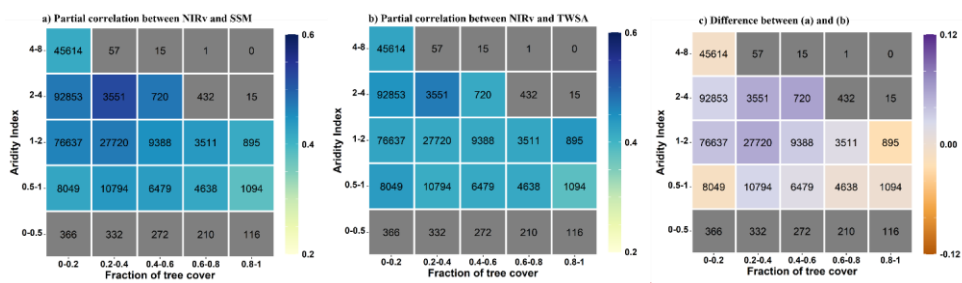
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3.2 Coupling of vegetation functioning with surface soil moisture and total water storage in dry months

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The correlation between NIRv and soil water storage increases during dry months (Figure 3a,b) compared to growing season months (Figure 2a,b). This increase is consistent for both SSM and TWS and across all tree cover fractions and aridity classes. This is because the water limitation on vegetation increases in dry months and so does the vegetation's sensitivity to the moisture. During the dry months, the correlation with near-surface soil moisture tends to rise, but the correlation with terrestrial water storage increases even more significantly (Figure 3c). This hints/indicates the relevance of deeper water resources during periods of scarce rainfall. The partial correlation maps (Figure S4S5) also reveal that NIRv's correlation with TWS increases more than its correlation with SSM for most grid cells.

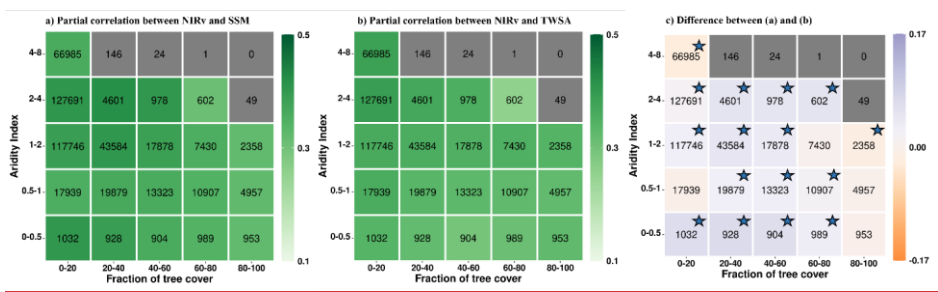
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355
356

Figure 3: Similar to Figure 2, but only considering the 10% driest months in each grid-cell.

357



Commented [12]: nice check!
Commented [13]: Thanks
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Commented [14]: yes sounds like the groundwater or really deep soil moisture supply vegetation water demand a lot during dry periods. falling into a topic of this paper but focusing on extremes and models (Mu, M. et al. Exploring how groundwater buffers the influence of heatwaves on vegetation function during multi-year droughts. Earth Syst. Dyn. 12, 919–938 (2021)). nice to see the consistence.
Commented [15]: Are you suggesting to cite this?

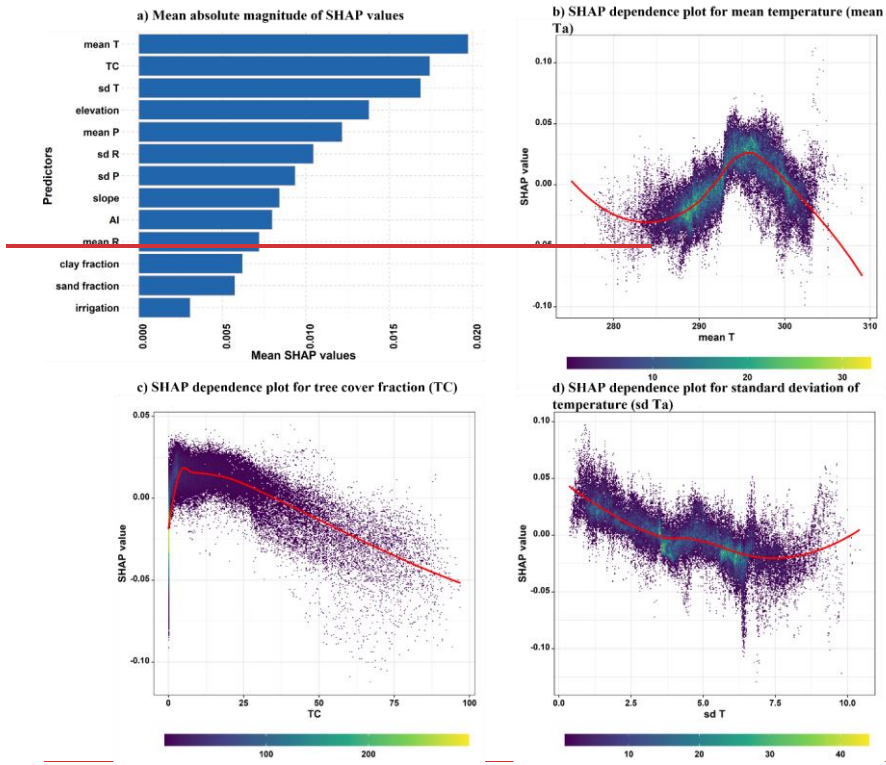
358 Figure 3: Summarising the coupling strengths of vegetation functioning (NIRv) with (a) near-surface soil moisture (SSM) and (b)
359 terrestrial water storage (TWS) in the 10% driest months in each grid-cell across climate (aridity index) and vegetation regimes
360 (fraction of tree cover). (c) shows the difference between (a) and (b). Numbers within the boxes denote the number of grid cells for
361 each aridity-tree cover class. Aridity-tree cover classes containing less than 1000 grid cells are shown in grey. The color bar denotes
362 the mean partial correlation for each class, computed from bootstrapping. The asterisk in figure (c) signifies that the 95% confidence
363 interval (lower and upper) shares the consistent sign (+/-) in the difference of partial correlation. Only grid cells with positive partial
364 correlation are considered.

365 During dry months, the number of analysed grid cells (**Figure 3**) is lower compared to all growing season months (**Figure 2**).
366 We performed a reanalysis of the correlation patterns within aridity-tree cover classes by selecting only those grid cells that
367 displayed ~~significant and~~ positive partial correlation between NIRv and soil water storages during both the dry months and the
368 growing season months. The results demonstrate that the previously observed patterns remain valid, thereby eliminating the
369 impact of the differing numbers of grid cells analysed. (**Figure ~~S5~~-S6**).

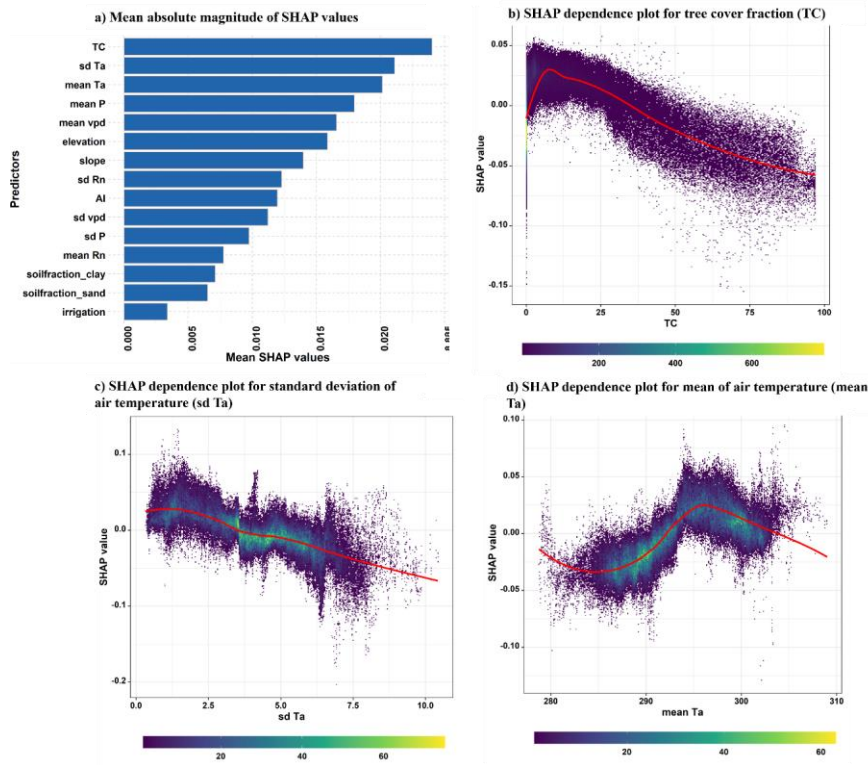
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3.3 Climate, vegetation, and topographic controls on the relevance of surface soil moisture vs. total water storage on vegetation



373



374

375 **Figure 4:** (a) Global feature importance based on the mean absolute magnitude of the SHAP values. The higher the mean SHAP
 376 values, the greater the predictor's relevance. (b-d) Evaluation of SHAP values (=contributions to the correlation difference
 377 illustrated in Figure 1c) against predictor values for the 3 most relevant predictors mean temperature during the growing season
 378 months (mean T_a), tree cover fraction (TC), and variability of temperature during the growing season months (sd T_a), and mean
 379 temperature (mean T_a) during the growing season months. The colour indicates the density of data points. For plotting (b), (c) and
 380 (d), only 10 percent random samples of the whole dataset are utilised.

381

382 We use a random forest model to understand the spatial variability in the relevance of SSM versus TWS for NIRv. The model
 383 was trained with 1315 climatic, vegetation, and topographic predictors against the target variable which is the difference of
 384 the partial correlations of NIRv with SSM and TWS during growing season-months ($R^2 = 0.6459$, see **methods section 2.2.3**).

385

386 The mean absolute SHAP value plot shows that the tree cover and the climate variables (mean and standard deviation of T_a)
 and tree cover are most important variables for explaining the spatial variability in the relative importance of SSM vs. TWS
 for vegetation functioning (**Figure 4a**). This overall highlights that the relative importance of SSM vs. TWS for the vegetation

387

387 is broadly controlled by ~~climate, influencing water availability and~~ vegetation type, reflecting the local adaptation of ecosystem
388 ~~and climate, influencing water availability~~ (Stocker et al., 2023).

389 ~~The relative importance of SSM and TWS varies non-linearly with the mean growing season temperature (Figure 4b). TWS~~
390 ~~tends to be more crucial for vegetation functioning in areas with low (approximately below 20°C) or high (above 27°C)~~
391 ~~growing season temperatures, while SSM has greater importance in regions with moderate growing season air~~
392 ~~temperatures. One possible explanation for this trend is that high temperatures induce a strong atmospheric water demand that dries near-~~
393 ~~surface soil layers, which leads vegetation to increase water extraction from deep soils. In contrast, SSM is more available~~
394 ~~during growing seasons characterised by moderate temperatures. Regions that experience relatively cold growing season~~
395 ~~temperatures exhibit stronger temperature and weather variability that may contribute to longer dry periods and, thus,~~
396 ~~emphasises the importance of deeper soil moisture for vegetation functioning. However, it should be noted that our findings~~
397 ~~regarding the relevance of TWS at high temperatures must be interpreted with caution due to the exclusion of most tropical~~
398 ~~forest regions from our analysis (Figure S6). Trec. As a result, most warm regions are dry, and there are only a few hot and wet~~
399 ~~regions included in our training data.~~

400
401 ~~In addition to mean growing season T_g ,~~ tree cover fraction is an important factor in determining the relevance of SSM and
402 TWS for vegetation functioning (Figure 4c). Regions with a high tree cover are more dependent on TWS, as trees generally
403 have deeper root systems that allow them to adjust water uptake between different depths (Tao et al., 2021). Grasslands on the
404 other hand have shallow roots that are more susceptible to surface soil moisture variations (Yang et al., 2014).

405
406 ~~Similarly, the relative importance of SSM and TWS varies non-linearly with the mean growing season temperature (Figure~~
407 ~~4b). TWS tends to be more crucial for vegetation functioning in areas with low (approximately below 20°C) or high (above~~
408 ~~27°C) growing season temperatures, while SSM has greater importance in regions with moderate growing season air~~
409 ~~temperatures. One possible explanation for this trend is that high temperatures induce a strong atmospheric water demand that~~
410 ~~dries near-surface soil layers, which leads vegetation to increase water extraction from deep soils. This observation is further~~
411 ~~underscored by the analogous pattern observed in the SHAP dependence plot for vpd, which accentuates atmospheric water~~
412 ~~demand (Figure S8). In contrast, SSM is more available during growing season months characterised by moderate~~
413 ~~temperatures. We hypothesize that the regions that experience relatively cold growing season temperatures exhibit stronger~~
414 ~~temperature and weather variability that may contribute to longer dry periods and, thus, emphasises the importance of deeper~~
415 ~~soil moisture for vegetation functioning. However, it should be noted that our findings regarding the relevance of TWS at high~~
416 ~~temperatures must be interpreted with caution due to the exclusion of most tropical forest regions from our analysis (Figure~~
417 ~~S7). As a result, most warm regions are dry, and there are only a few hot and wet regions included in our training data.~~

Commented [16]: so why don't you put vpd in explaining the spatial ssm and tws patterns?

Commented [17]: Hi Wantong, For now, we do not add vpd currently, because we thought T largely reflects vpd and will overall not change the conclusion. However, we have kept it on our list to go through it when there are other revisions needed. Thanks for raising the issue.

Commented [18]: Reference?

Commented [19]: This is our "possible explanation".

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420 Not only the mean of the growing season temperature, but also its variability is crucial for explaining the significance of SSM
421 and TWS for vegetation functioning (**Figure 4d**). A higher temporal variability in temperature increases the importance of
422 TWS for vegetation. [This is because atmospheric water demand scales with temperature. Hence, higher variability in
423 temperature implies more peaks in related atmospheric water demand which is a stronger incentive for plants to access] deeper
424 water storages which are more often available to meet the vegetation's transpiration demand.

425
426 **Figure S8** illustrates the effect of the other six important predictors on the model output. Apart from climatological parameters
427 (mean P, mean vpd, variability in R_n and P, and aridity index), elevation and slope explain part of the variability in the
428 relevance of SSM vs. TWS for NIRv. Although the reasons for increasing relevance of TWS for vegetation functioning at
429 higher elevation remain unclear, it may be due to elevation's strong correlation with other climatic variables such as T_a and P_a .

430
431 Several local studies identified other relevant factors that determine root water uptake depth such as forest stand age and tree
432 height, competition, root hydraulic architecture, and tree species (Zhu et al., 2022; Quijano et al., 2012; Stahl et al., 2013,
433 Gessler et al., 2021; Liu et al., 2021). For example, young trees more easily increase their root activity in the shallow or deep
434 soil dependent on soil moisture than mature trees (Zhu et al., 2022; Drake et al., 2011). These variables were not included in
435 our attribution analysis, because they are not available at global scale.

436 3.4 Robustness Tests

437 In the aforementioned analysis, we included grid cells exhibiting both positive partial correlations, whether significant or non-
438 significant. Upon further examination, we specifically assessed the evolution of partial correlation between NIRv and water
439 storages, considering only grid cells with significant partial correlation ($p < 0.05$). The observed patterns along the aridity-tree
440 cover gradient remained similar during growing season months. This suggests the robustness of our results to the choice of the
441 statistical significance criterion, albeit with a substantial reduction in the number of globally available grid cells when
442 considering only significant partial correlation (**Figure S9**).

443
444 Furthermore, to ensure that our results are robust to variations in the threshold for Solar-Induced Fluorescence (SIF) used to
445 define growing season months, we conducted additional analyses with a different SIF threshold. Instead of filtering out all
446 months from 2007-2018 when the mean-monthly SIF value was below the threshold of 0.2 mW/m²/sr/nm, we utilized a
447 threshold of 0.5 mW/m²/sr/nm. Elevating the SIF threshold implies the exclusion of additional months characterized by lower
448 vegetation activity for the partial correlation analysis. However, it is essential to note that this threshold does not seem to affect
449 the number of globally available grid cells during growing season months and hence patterns along AI-TC classes are similar.
450 Instead, it specifically influences the selection of dry months and hence the number of grid cells available for the analysis
451 during dry months. . Nevertheless, even with the elevated SIF threshold for defining growing season months, the observed

Commented [20]: yes you certainly know the importance of vpd in regulating the vegetation water uptake from the sub-surface by biophysical processes. and vpd is not equal to temperature. why don't you directly count it in? correct me if i'm wrong.

Commented [21]: Hi wantong, Please see above reply on vpd. We will take into account "vpd" when revision is done. Thanks

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452 patterns along aridity-tree cover (AI-TC) classes remain largely consistent with the results obtained in our main analyses
453 (Figure S10).

454
455 Although NIRv can largely reflect vegetation functioning (Badgley et al., 2017), we repeat our analysis with SIF, which is an
456 alternative and independent indicator for vegetation functioning and shows a near-linear relationship with gross primary
457 productivity at the ecosystem level (Guanter et al., 2012)(Guanter et al., 2012). However, SIF is only available at a coarse
458 resolution of 0.5 degree. The partial correlations, $r(\text{SIF}\sim\text{SSM})$ and $r(\text{SIF}\sim\text{TWS})$ largely agree with the pattern of $r(\text{NIRv}\sim\text{SSM})$
459 and $r(\text{NIRv}\sim\text{TWS})$ across varying aridity index and tree cover classes (Figure S9);S11. This suggests that our overall
460 conclusion on the relevance of SSM or TWS for vegetation functioning is robust across different indicators of vegetation
461 productivity.

462
463 ~~The partial correlation of NIRv with TWS is confounded by the presence of SSM within TWS, which makes it challenging to~~
464 ~~determine the relative importance of SSM and TWS for vegetation functioning. To address this issue, we re-calculated the~~
465 ~~partial correlation of NIRv with TWS while additionally controlling for SSM (next to T_e and R_n) during growing season~~
466 ~~months. With this additional control variable, we observed fewer grid cells with positive and significant correlations compared~~
467 ~~to the analysis without controlling for SSM. Additionally, the magnitude of the partial correlation of NIRv with TWS slightly~~
468 ~~decreased in most grid cells when controlling for SSM (Figure S10). Nevertheless, we still observed the decreasing relevance~~
469 ~~of SSM and increasing relevance of TWS along an increasing tree cover fraction. Similar gradient across the aridity index is~~
470 ~~also observed in this analysis controlling for SSM. Thus, we conclude that our findings hold even after controlling for the~~
471 ~~effect of SSM in TWS. Additionally, we tested if our results are robust when the aridity index is calculated based on the FAO~~
472 ~~Penman-Monteith Reference Evapotranspiration equation, for which we applied aridity classification based on UNEP 1997~~
473 ~~guidelines. Our results confirm the findings of Section 3.1 and Figure 2 that as aridity increases, the correlation of NIRv with~~
474 ~~Soil Surface Moisture (SSM) and Total Water Storage (TWS) intensifies. Moreover, in hyper arid regions ($\text{AI} < 0.03$) the~~
475 ~~correlation with TWS surpasses that with SSM (Figure S12). They also confirm that regions with higher tree Cover (TC)~~
476 ~~fraction correlates more strongly with TWS compared to SSM. Thus, the choice of aridity index formulation does not alter our~~
477 ~~main conclusions.~~

479 **4. Summary and Conclusions**

480 In this study we compare the relevance of near-surface soil moisture and of terrestrial water storage for vegetation functioning
481 across the globe. We find that in semi-arid regions and regions with low tree cover, vegetation preferentially utilises the water
482 from shallow soil, which is related to continuous availability of near-surface water availability and lack of deep rooting systems
483 respectively. The stronger correlation of NIRv with SSM than TWS is supported by site-level studies that find a higher root

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484 [water uptake of surface soil moisture \(Brinkmann et al., 2019, Gessler et al., 2021, Deseano Diaz et al., 2023; Kulmatiski and](#)
485 [Beard, 2013\)](#), also when deeper water is available. Some local studies however find a higher root water uptake from deeper
486 [layers \(Zhu et al., 2022\)](#).

487
488 By contrast, in mostly forested regions and in relatively dry climate regimes, the correlation with terrestrial water storage is
489 comparable or higher than with near-surface soil moisture, indicating that trees and vegetation in arid regions use their deep
490 root systems to access deeper soil moisture. [Point-scale studies also found a different water uptake depth for trees and grasses](#)
491 [in for example savanna ecosystems \(Kulmatiski et al., 2010\), and a different water uptake depth for tree species \(Kahmen et](#)
492 [al., 2022\)](#). [Liu et al. \(2021\) showed for example that for a karst forest in Southwest China, evergreen species rely mostly on](#)
493 [water sources from the 0-30 cm layer, while deciduous species extracted most water from the 30-70 cm layer.](#)

494
495 We also find that vegetation's preferential water uptake depth changes over time. During particularly dry months, the relative
496 importance of terrestrial water storage is higher, highlighting the importance of deep water resources during periods of low
497 soil water availability. This is in line with previous studies showing changes in vegetation's water uptake depth during drought
498 periods at small spatial scales where accessing water in deeper soil layers helps plants to alleviate water stress and maintain
499 transpiration (Migliavacca et al., 2009; Tao et al., 2021).

500
501 [Our global results are supported by site-scale studies that find that, during drought, the deeper roots play a more active role in](#)
502 [water extraction \(Stahl et al., 2013, Volkmann et al., 2016; Tao et al., 2021\)](#). In some studies however, the increase of deep
503 [water uptake is only relative: the absolute uptake of deep water does not increase, but the uptake of shallow water decreases](#)
504 [\(Brinkmann et al., 2019, Gessler et al., 2021, Rasmussen et al., 2020; Kühnhammer et al., 2023\)](#). This means that the uptake
505 [of deeper soil layers cannot compensate for the loss of water uptake from the dry topsoil. Contrary to trees, grasses do not shift](#)
506 [their uptake depth \(Deseano Diaz et al., 2023\), or even extract water from the most shallow soils \(Prechsl et al., 2015,](#)
507 [Kulmatiski and Beard, 2013\)](#).

508
509 Furthermore, we show that the spatial variability of the importance of near-surface soil moisture vs. terrestrial water storage
510 for vegetation functioning is influenced by ~~temperature and the fraction of tree cover~~ [and mean and standard deviation of air](#)
511 [temperature](#). This emphasises the role of climate in determining shallow vs. deep soil water resources, and the role of vegetation
512 in adapting to different soil water availability patterns.

513
514 Vegetation functioning and soil water storages are generally coupled in both directions, i.e. while soil moisture availability
515 affects vegetation functioning (positive coupling), this in turn also affects soil moisture through transpiration (negative
516 coupling). As our study focuses on water-controlled vegetation we only consider positive couplings and filter out grid cells

517 with negative correlations. Future research may consider the relevance of soil moisture across depths for the positive coupling
518 regions.

519

520 Overall, our analysis illustrates that satellite-based data can be used for belowground analysis at large spatial scales thanks to
521 the fact that satellite retrievals can assess soil water storage dynamics across depths and because vegetation in water-controlled
522 areas can be used as an indicator of soil water dynamics. Such novel ways to improve our understanding of belowground water
523 dynamics is necessary and valuable as respective in-situ observations are scarce and of limited representativeness for larger
524 areas, particularly given the typical spatial heterogeneity of soils and vegetation. Our results can further inform a better
525 representation of belowground processes in global models in order to support more accurate projections of future changes in
526 climate, water resources, and ecosystem services.

527 **Data availability**

528 The monthly SIF data is available from [https://www.gfz-potsdam.de/sektion/fernerkundungund-](https://www.gfz-potsdam.de/sektion/fernerkundungund-geoinformatik/projekte/global-monitoring-of-vegetation-fluorescence-globfluo/daten)
529 [geoinformatik/projekte/global-monitoring-of-vegetation-fluorescence-globfluo/daten](https://www.gfz-potsdam.de/sektion/fernerkundungund-geoinformatik/projekte/global-monitoring-of-vegetation-fluorescence-globfluo/daten). The NIRv was calculated from the red
530 and near-infrared reflectance obtained from the MOD13C1 v006 product (<https://lpdaac.usgs.gov/products/mod13c1v061/>).
531 The ESA-CCI soil moisture can be accessed through <https://esa-soilmoisture-cci.org/> and Terrestrial Water Storage Anomaly
532 data can be accessed through https://podaac.jpl.nasa.gov/dataset/TELLUS_GRACGRFO_MASCON_CRI_GRID_RL06_V2.
533 The ERA5 climate variables are available from <https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era5> . Tree
534 cover fraction data is available from the AVHRR vegetation continuous fields products
535 <https://lpdaac.usgs.gov/products/vcf5kyrv001/>, land cover data is available from <https://www.esa-landcover-cci.org/>, and
536 topographic data is available via <https://www.earthenv.org/topography>. Similarly, the irrigation fraction data could be accessed
537 from <https://mygeohub.org/publications/8> .

538 **Competing Interests**

539 The contact author has declared that none of the authors has any competing interests.

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