



1	Measurement report:
2 3	Intra-annual Variability of Black/Brown Carbon and Its Interrelation with Meteorological Conditions over Gangtok, Sikkim
4 5 6	Pramod Kumar ¹ , Khushboo Sharma ¹ , Ankita Malu ² , Rajeev Rajak ² , Aparna Gupta ¹ , Bidyutjyoti Baruah ¹ , Shailesh Yadav ¹ , Thupstan Angchuk ¹ , Jayant Sharma ¹ , Rakesh Kumar Ranjan ^{1#} , Anil Kumar Misra ¹ , and Nishchal Wanjari ¹
7 8	¹ DST's Centre of excellence on Water Resources, Cryosphere and Climate Change Studies, Department of Geology, Sikkim University, Gangtok, Sikkim, India -737102
9	² Department of Geology, Sikkim University, Gangtok, Sikkim, India -737102
10	[#] Corresponding Author: rkranjan@cus.ac.in
11	
12	Abstract
13	Black carbon (BC) and brown carbon (BrC) have versatile nature, and they have apparent role
14	in the climate variability and changes. As the anthropogenic activity is surging, the BC and
15	BrC are also reportedly increasing. So, the monitoring of BC/BrC and observation of land use
16	land cover changes (LULCC) at regional level are necessary for the various interconnected
17	meteorological phenomenal changes. The current study investigates BC, BrC, CO ₂ , BC from
18	fossil fuels (BCff), BC from biomass burning (BCbb), LULCC, and their relationship to the
19	corresponding meteorological conditions over Gangtok in Sikkim Himalayan region. The
20	concentration of BC (BrC) 43.5 $\mu g/m^3$ (32.0 $\mu g/m^3)$ is found to be highest during the March-
21	2022 (April-2021). Surface pressure has been found to have a significant positive correlation
22	with BC, BCff, BCbb and BrC. The boundary layer is calmer and more stable when the surface
23	pressure is higher, which keeps contaminants deposited there. The wind, on the other hand,
24	appears to represent the dispersion of pollutants with a strong negative correlation. The fact
25	that all pollutants and precipitation have been shown to behave similarly points to moist
26	scavenging of the pollutants. Despite the dense cloud cover, it is clear that the area is not
27	receiving convective precipitation, implying that orographic precipitation is occurring over the
28	region. Most of Sikkim receives convective rain from May to September, indicating that the
29	region has significant convective activity contributed from the Bay of Bengal during monsoon
30	season. Furthermore, monsoon months have the lowest concentrations of BC, BCbb, BCff, and
31	BrC, suggesting the potential of convective rain (as rain out scavenging) to remove most of
32	the pollutants. Moreover, BC and BrC show positive radiative feedback.
33	Key words: Black carbon; Brown carbon; LULC; Sikkim Himalaya; Meteorology; Biomass

34 burning; Radiative forcing.





35 1.0 Introduction

Black carbon (BC), and brown carbon (BrC), are part of fine particulates air pollution that have 36 apparent role in the climate variability and changes. BC/BrC is a short-lived climate pollutant 37 with a lifetime of only days to weeks after release in the atmosphere (Pierrehumbert, 2014). 38 39 During this short period of time, BC/BrC can have significant direct and indirect impacts on 40 the climate, cryosphere, agriculture, and human health (Shindell et al., 2012). It consists of pure carbon in several interconnected forms. BC is formed through the incomplete combustion 41 of fossil fuels, biofuel, and biomass, and is one of the main types of particles in both 42 43 anthropogenic and naturally occurring soot (Bond et al., 2004). BrC in the atmosphere have 44 been attributed to burning of biomass and fossil fuels, biogenic release of fungi, plant debris, 45 and humic matter and multiphase reactions between the gas-phase, particulate, and cloud 46 microdroplet constituents in the atmosphere (Laskin et al., 2015). BC/BrC is transported from its source to many locations across the world (Ramanathan and Carmichael, 2008). The 47 released BC/BrC is vertically distributed and travels through the atmosphere according to wind 48 speed and direction, interacting with numerous components before sinking on the earth's 49 surface through wet or dry deposition. Its hygroscopic nature makes more susceptible to cloud 50 seeding and cloud formation process and so directly helps in precipitation mechanism in high 51 52 humid conditions (Stevens and Feingold, 2009). In addition, it absorbs both incoming and outgoing radiation, atmospheric BC/BrC modifies radiative forcing, disturbs atmospheric 53 stability, regional circulation and rainfall pattern, affects cloud albedo, material damage, 54 55 reduces agricultural productivity, degrades ecosystem and affects human health (Zhang et al., 56 2013). However, due to an insufficiency of observations, BrC is one of the least understood 57 and uncertain warming agents (Yue et al., 2022). Several studies have been carried out to examine the concentration of BC and BrC all over the world and in India as well (Reddy and 58 Venkataraman, 2002a, 2002b; Venkataraman et al., 2006; Park et al., 2010; Sloss, 2012; Helin 59 et al., 2021; 2020; Kumar et al., 2020; Watham et al., 2021; Bhat et al., 2022; Runa et al., 2022; 60 Yue et al., 2022; Kumar et al, 2018b). However, the overall worldwide BC emission is 61 estimated to be 4800-7200 Gg per year (Klimont et al., 2017). In 2001, India's total BC 62 63 emissions were projected to be 1343.78 Gg (Sloss, 2012). Residential fuel burning and transportation contributes maximum to the global anthropogenic BC emission (Helin et al., 64 2021). About 60 to 80% of residential fuels (coal and biomass) emissions are reported from 65 Asian and African countries, whereas approximately 70% of diesel engines emission is found 66 67 to be from Europe, North America, and Latin America.





68 On the other hand, emissions on the Indian subcontinent have increased by 40% since the year of 2000. According to Reddy and Venkataraman (2002a, 2002b), the estimated BC emissions 69 70 in India are fossil fuels, 100 Gg biofuel, 207 Gg open burning, and 39 Gg with a climatic forcing of +1.1 W/m², black carbon is the second-most significant human emission in the 71 72 current atmosphere (Sharma et al., 2022). BC concentration was measured by Zhao et al. 73 (2017) in the south-eastern Tibetan Plateau (TP). Daily mean BC loadings ranged from 57.7 74 to 5368.9 ng/m³ demonstrating a high BC burden even at free tropospheric altitudes (Zhao et 75 al., 2017). Black carbon (BC) deposition was estimated at the Nepal Climate Observatory -76 Pyramid (NCO-P) site in the Himalayan region during the pre-monsoon season (March-May). A total BC deposition rate of 2.89 μ g/m³/day was estimated, resulting in a total deposition of 77 266 μg/m³ for March–May (Yasunari et al., 2010). From the Indian perspective, several key 78 79 short-term incidents contribute to a rise in India's BC concentration from biomass burning and 80 other sources (Kumar et al., 2020). Burning agricultural waste (stubble) is widespread in India and several other nations. Many studies suggest increased BC in northern India, notably Indo-81 Gangetic plain (IGP) is the global absorbing aerosol hotspot (Venkataraman et al., 2006; 82 Ramanathan and Carmichael, 2008). In India post-monsoon paddy crop waste burning occurs 83 84 in the month October and November in north and northwest part of India (Venkataraman et 85 al., 2006). In the north-western Indo-Gangetic Plain (IGP) (especially- Punjab, Haryana, and western Uttar Pradesh), stubble burning is a popular practice (Venkataraman et al., 2006). 86 Long-distance transport of BC aerosols, mostly from Asia to the north Pacific and South 87 America to the southwest Atlantic, is often recognised as a significant factor in local 88 89 concentration (Evangelista et al., 2007). However, in India only local sources (89%) affects 90 BC concentrations (Zhang et al., 2013), as there aren't many movements of transboundary 91 aerosols contribution over the IGP (Kumar et al., 2018a; Kedia et at., 2014; Ramachandran and Rupakheti, 2022; Ramachandran et al., 2020). Both marine and continental air masses 92 93 contributed to total aerosol loading over middle-IGP (Kumar et al., 2017; Shukla et al., 2022). 94 Black carbon is a light-absorbing particle that are released into the atmosphere directly in the

Black carbon is a light-absorbing particle that are released into the atmosphere directly in the form of ultrafine (<0.1 µm) to fine particles (<2.5µm) (Gupta et al., 2017). BC is a good tracer for particle deposition as it is non-volatile, insoluble, and chemically inert, and it can also mix well with other aerosol species in the atmosphere (Kiran et al., 2018). As a result, BC deposition data are important not just for BC sinks but also for a broader understanding of aerosol deposition. BC emissions are mostly influenced by significant changes in the energy sector, fuel usage, industrial expansion, and an increase in the number of vehicles (Bisht et al., 2015). Residential fuel like wood, agricultural waste, and cow dung used for cooking and





102 biomass usage for home purposes are the primary sources of BC emissions (Venkataraman et 103 al., 2006). The Asian mainland is a substantial contributor to global BC emissions and has 104 been identified as a hotspot (Gupta et al., 2017). BC has a high absorption ability, accounting 105 for 90-95 percent of total atmospheric aerosol absorption (Hansen et al., 1984). It can absorb 106 solar energy in the visible-infrared band and warm the environment. In comparison to carbon 107 dioxide, BC has a much shorter life cycle in the atmosphere. As a result, mitigation or reduction 108 has a greater positive impact on the atmosphere (Kirchstetter et al., 2004; Takemura and 109 Suzuki, 2019). Changing land use land cover (LULC) has very significant impact on weather, 110 climate and aerosols (Mahmood et al., 2010). It is well stabilised fact that the LULC change 111 has direct relation with land surface temperature, vehicular emission and anthropogenic 112 activity (Aithal and MC, 2019). Which motivated the present study for the further analysis for 113 Sikkim region land use land cover change and its relation with temperature and BC/BrC for 114 the March 2021 to March 2022. The current study's objectives are to assess the intra-annual 115 variability of Black/Brown Carbon (BC/BrC) (diurnal/daily/monthly) during the study period 116 March-2021 to March-2022, as well as the interrelationship between meteorological conditions and BC/BrC, along with LULC change for three decades 2000, 2010, and 2020, and its 117 relationship with anthropogenic activity over Gangtok. 118

119 2.0 Study location

120 The Gangtok Municipal Corporation (GMC) has been selected for the present study on the 121 basis of its urban exposer and settlement change for three decades as well as congruently 122 temperature rise (figure S1). The sampling has been carried out at Pani House area in Gangtok, GMC, having longitude 88.609°E and latitude 27.323°N. Sikkim is surrounded by Nepal, 123 124 China and Bhutan from west, north and east respectively and consists of the trans and greater 125 Himalayan range. It has one of the most fragile forest covers. The Gangtok is densely 126 populated city and capital of state Sikkim which is situated in the East Sikkim district (see 127 figure 1a). The population of the Sikkim has been found to be increased as per Indian census 128 for three decades as this can be seen in table S1.

129 **3.0 Data and Methodology**

The real time sampling of BC was carried out from 10th March 2021 to 17th March 2022, at
Gangtok using the seven-channels dual spot Aethalometer (Model AE-33-7, Magee Scientific,
USA). The data was collected for the measurement of BC and BrC associated with particulate
matter having an aerodynamic diameter less than 2.5 µm (PM_{2.5}). The concentration of BC,
BrC, BCbb, and BCff have been estimated by Carbonaceous Aerosol Analysis Tools (CAAT)





135 software tool from the Magee Scientific Aethalometer model AE33 (Hansen and Schnell, 136 2005). The carbon dioxide (CO₂) was measured using a CO₂ sensor (Vaisala-GMP343) which 137 is attached to the aethalometer. The inlet of the aethalometer was mounted at a height of 15 m 138 above ground level. One of the main sources of uncertainty in using aerosol absorption 139 measurements to estimate BrC mass concentration is the fact that other species, such as black 140 carbon and dust, can also contribute to the measured absorption. This can lead to overestimation of BrC mass concentration, particularly in environments where these species 141 142 are also present. However, in the Sikkim region has one of the higher precipitation regions in 143 the world and negligible contribution of the dust pollution. Furthermore, there must be lesser 144 over/under estimation. Therefore, the present study used mass concentration.

A new data set of BC, BrC, Black Carbon from biomass burning (BCbb), Black Carbon from 145 fossil fuels (BCff), BrC, percentage contribution of biomass burning to BC (BB%) and CO₂ 146 147 has been generated over the unreported region of Sikkim Himalaya. The diurnal and monthly 148 data set of BC, BCbb, BCff, BrC, BB% and CO2 have been given in the details in 149 supplementary materials (Table S2 and S3). In addition to this the meteorological data has been selected ERA5 reanalysis for the study. LULC data has been taken from USGS earth 150 151 explorer of 2000 and 2010 landsat-5, 2020 landsat-8, and 2021 for sentinel-2 (Karra et al., 152 2021). LULC data has been chosen for the month of December to minimize the cloud cover. The details of the LULC calculation steps used are given in the supplementary section 153 154 (methodology \$1.3). The brief of the data set is discussed in the table 1.

155 **3.1 Estimation of BrC**

156 The Carbonaceous Aerosol Analysis Tools (CAAT) software tool from the Magee Scientific 157 Aethalometer model AE33 was utilized to estimate the concentrations of BC, BrC, BCbb, and 158 BCff. The absorption coefficients of BC and BrC were determined using the multi-wavelength 159 absorption coefficients provided by the aethalometer. The presence of BrC was identified by 160 observing the maximum light absorption between 370–590 nm, but its absorption may increase 161 significantly below this range depending on its composition. The attenuation of illumination 162 measured in this study using the aethalometer was attributed solely to the contribution of BC and BrC. It is believed that the absorption coefficient at 370 nm measured by the aethalometer 163 164 represents the combined absorption coefficients of BC and BrC, which is denoted as σBC + 165 BrC (370 nm). This assumption is similar to the model used in the multi-wavelength 166 absorbance analyzer (MWAA) approach for source allocation, as described in Massabò et al. (2015). Equation (3.13) was used to calculate the σBrC (370 nm) absorption coefficient 167





168(supplementary methodology S1), which involved subtracting the contribution of BC (σ BC169(370 nm)) from the observed absorption coefficient (σ BC + BrC (370 nm)).170 σ BrC(370nm) = σ BC + BrC(370 nm) - σ BC(370 nm)171The σ BC (370 nm), was calculated by applying the power-law fit to absorption data in the 590-172950 nm wavelength range provided in equation (1).

173
$$\sigma BC(\lambda) = \beta \lambda^{-AAE} BC$$
 Eq. (2)

The absorption angstrom exponent of BC is denoted as AAE_{BC} , with β being a constant value. 174 As BC is a significant contributor to light absorption at wavelengths beyond 590 nm, the 175 176 contribution of other aerosol species can be neglected, and the AAE_{BC} can be calculated using 177 equation (3.15) (supplementary methodology S1), as stated in Rathod and Sahu (2022). The AAE for both BC and BrC can be expressed as σ , and in this study, the AAE definition by 178 179 Moosmüller et al. (2011a) was used instead of the AAE specified for a wavelength pair. This value is determined by equation (2), which calculates the negative log-log slope of the 180 181 absorption spectrum at wavelength λ .

182
$$AAE = \frac{dln\sigma BC}{d \ln \lambda}$$
 Eq. (3)

Instead of the conventional approach where AAEBC is assumed to be 1, we utilized the AAEBC that was observed onsite to calculate $\sigma BC(\lambda)$. Equation (3.16) was employed to determine σBrC (370 nm) by substituting $\sigma BC(\lambda)$ at 370 nm, which was obtained using equation (3), into equation (3.13) (refer to supplementary methodology S1.1, S1.2, and figure S2 for details).

188 BrC(370 nm) = BC + BrC(370 nm) - (370 nm) -
$$AAE_{BC}$$
 Eq. (4)

To calculate $\sigma BrC(\lambda)$ at 470 nm and 520 nm, we can subtract the modelled BC from the measured absorption coefficients, in a similar manner. It is worth noting that the BrC absorption coefficients are very low at wavelengths beyond 590 nm, according to Rathod et al. (2017) and Rathod and Sahu (2022), hence they are not taken into account (supplementary methodology S1).

194 **3.2 Data Analysis**

LULC change also has a direct impact on vehicular emissions and other anthropogenic
activities. Urbanization, conceivably, can lead to increased vehicle traffic and emissions,
which can contribute to air pollution and climate change. Changes in land use can also affect





198 the amount and type of vegetation, which can influence the carbon cycle and the amount of 199 greenhouse gases in the atmosphere. The ERA-5 reanalysis data has been used for 200 meteorological analysis viz. wind pattern, precipitation, relative humidity, and temperature 201 (Hersbach et al., 2020). The hourly data has been taken for the analysis and then daily, monthly 202 and seasonal average has been computed for the study period over the Sikkim and surrounding 203 states for a better understanding the meteorological conditions influencing the BC, and BrC. 204 The total precipitation is computed as sum of the hourly data for a day to daily total 205 precipitation and further it was summed for monthly cumulative total precipitation using sum 206 formula as

207 Monthly Cumulative Total Precipitation =
$$\sum_{i=1}^{n} X$$
 Eq. (5)

Where, 'i' is the initial and 'n' the last date and X is hourly total precipitation taken from
ERA5. The wind circulation has been computed using u-component and v-component of wind
and the wind speed has been calculate as

211 Wind Speed =
$$\sqrt{u^2 + v^2}$$
 Eq. (6)

212 The temperature and relative humidity averaged have been computed using mean formula as

213
$$Average = \frac{\sum_{i}^{n} x}{n}$$
 Eq. (7)

Where, 'i' is initial and 'n' last date of the of variables such as temperature, relative humidity and wind components.

Let x and y be two real-valued random variables such that the correlation coefficient spearmen
Pearson can be calculated between the BC/BrC and meteorological parameters. The
Coefficient of Pearson Correlation (PCC) (Pearson, 1909; Benesty et al., 2009) as

219
$$PCC = \frac{n(\sum xy) - (\sum x)(\sum y)}{\sqrt{[n \sum x^2 - (\sum x)^2][n \sum y^2 - (\sum y)^2]}}$$
Eq. (8)

220 Where 'n' is the population size of the variables used for the study.

221 Table 1 contains additional information about the dataset, and a more detailed methodology

222 can be found in the supplementary section (S1).

223 4.0 Results and Discussions

224 The anthropogenic activities in Gangtok are drastically increased in last 20 years. As evident

from figure 1b, c and d, LULC has been changed since 2000 to 2020 over the Gangtok





226 municipal corporation (GMC). The population change and growth have also been observed 227 over the Sikkim (Table S1). LULC during year 2000 and 2010 evidently shows that most of 228 the fallow land has been built-up due to recent change in the policy of construction in Sikkim 229 suggesting urban settlement load over Gangtok is increased significantly. As a result, there is 230 a significant increase in built-up areas in GMC for last 20 years. The vegetation cover has also 231 reduced since 2000 to 2020 (figure 1b, c, and d). The rainfed water bodies are reducing from 232 the GMC. However, due to its seasonal nature, streams are lesser emerged in 2020. Which 233 perhaps shows the precipitation pattern alteration over GMC due to highly built-up sprawl. 234 The built-up extent has been sprawling and consuming the dense vegetation regions as well. 235 This increases the study region's urge to be acknowledged so that Sikkim's future policymakers 236 can consider the effects of rising anthropogenic activities. This anthropogenic activity leading 237 to heavy load on environmental over one of the cleanest states of India. Long-term 238 spatiotemporal variation of 2-meter air temperature justifies the LULC change and warming pattern over the Gangtok region (figure S1a, S1b, S1c, S1d, and S1e). The decadal warming 239 rate is varying from 0.25° to 0.45°C (figure S1e). Thereafter, BC and BrC over the Gangtok 240 has been measured to report the issue and get more attention to the scientific and local 241 community. The higher anthropogenic activity releases the higher amount of emission in the 242 243 name of development due to population load on the region (i.e., growth rate has been raised 244 from 12.89 to 13.05% in recent years) (Table S11). Diurnal variation of the BC, BrC, BC BCbb, BC BCff and CO₂ apparently show two peaks. BC, BCff and CO₂ have almost similar 245 246 time of peaks observed. The first peak is found during 8-10AM. And, the second peak is 247 observed during 8-10PM. However, BrC and BCbb have the peak concentration during 10-248 11AM (figure 2a). The same for meteorological conditions is observed and referred to figure 249 2b.

250 The daily timeseries of the BC, BCbb, BCff, BrC, BB% and CO₂ show the highest fluctuation during 20th to 30th March in both 2021 and 2022 years respectively. The maximum BC (BrC) 251 content was found in March 2022 (April-2021), at 43.5µg/m³ (32µg/m³). The lowest 252 fluctuation is observed during 15th May to 15th September 2021 (figure 3a). The intense peaks 253 of BC, BCff and CO₂ has been observed during 10th October to 15th November 2021 (figure 254 3a) that may be linked to the heavy tourist season of the state and indicating towards the traffic 255 overload in the Gangtok (Sharma et al, 2022). As, the meteorological conditions are also 256 favouring the similar circumstances to accumulate the pollutant during 10th October to 15th 257 258 November 2021 (figure 3b). The lowest surface pressure with minimum fluctuation and the 259 highest temperature and dewpoint temperature with minimum fluctuation is being noticed





during the 15th June to 20th September 2021 (figure 3b). BrC is found the highest with 260 maximum fluctuation during 10th January to 30th March that is pointing towards winter wood 261 262 burning for the subsistence as similar observed BCbb. The monthly variations of BC, BCbb, 263 BCff, BrC, BB% is discussed in figure 4a, and the highest value of standard deviation were 264 observed during March 2022 for BC, BCff, and April 2021 for BCbb, BrC, BB%. The CO₂ is 265 observed almost constant with a small value of standard deviation. The maximum 266 concentration of the BC, BCff is found in March 2022. However, BCbb and BrC were 267 measured highest in April 2021. The minimum concentration of the BrC was seen in the month 268 of August 2021 as the highest total precipitation month with high wind speed, temperature and 269 dewpoint temperature and relative humidity (figure 4b, S3 and S4).

270 The good significant correlation between BC and BCff suggested that the major contribution 271 of the BC is fossil fuel burning (Osborne et al, 2008). A strong significant correlation between 272 BCbb and BrC indicating that major contributor of BrC is biomass burning that can be justified 273 by BB% and BrC strong significant positive correlation coefficient (figure 5). A good 274 significant positive correlation between CO₂ and BC/BCff suggesting that fossil fuel burning 275 is one of the causes of CO_2 concentration or vis versa. Dewpoint temperature and CO_2 has 276 strong significant positive correlation coefficient suggesting to positive radiative forcing of the 277 CO₂. The Similar has been found for the temperature. BCbb/BrC and temperature has strong significant negative correlation suggesting the negative radiative nature of the BCbb/BrC 278 279 (figure S5). Moreover, net thermal/solar radiation (STR/SSR) and BC/BrC have significant positive correlation (figure 5, and S5). A strong significant positive correlation between 280 281 surface pressure and BC/BCff (BCbb/BrC) has been observed (figure 5). Higher the surface 282 pressure makes calm condition and stable boundary layer, which keeps the pollutants accumulated in the boundary layer (Bharali et al., 2019). However, the opposite has been 283 284 observed for the wind that indicates the dispersion of pollutants with strong negative correlation. The similar has been observed for the total precipitation and all the pollutants, 285 286 delineates to wet scavenging of the pollutants. The relative humidity is also showing the similar 287 result to the total precipitation with greater values of coefficient. The negative correlation 288 between total precipitation and surface pressure suggested that the rain fall over the region 289 mostly occurs in low pressure system that is causes due to the vertical rising of air parcel and 290 cause to condensation and precipitation. However, cloud condensation nuclei formation and 291 precipitation are prompted by aerosols (BC and BrC). Thereafter, BC and BrC have crucial 292 role in precipitation mechanism.





293 Total precipitation and wind circulation suggested that the study region is receiving 294 precipitation in entire month of the study period (i. e., most of the time rain form and sometimes 295 snow). As the maxima is observed during the month of August and minima during March 296 2022. The wind pattern delineates during the May to September 2021 the monsoon seasonal 297 strong effect (figure 6). And rest of the period the wind is conversing in the valley and 298 diverging from the mountain (figure 6). The strong wind and heavy rain fall suggested the 299 pollutant scavenging (rain out or wash out) that is why it is significant negatively correlated. 300 The relative humidity and temperature patten also justify the same as the temperature gradients 301 change from January to December and moisture content reduction in the atmosphere (figure 302 S6). The lowest in month of February is observed and temperature gradient getting steep from 303 the November (figure S6). The dewpoint temperature contour and surface pressure shading 304 match well suggesting that the surface pressure creates the dewpoint temperature gradient and 305 keep it sustained and stable atmospheric condition (figure S7). During month of June, it is very 306 peculiar that the dewpoint temperature contours are wide and very small gradient is observed 307 (figure 7). Which is pointing toward the warm condition during the June over entire Sikkim. Figure 7 discusses about the cloud cover and convective precipitation over the Sikkim. It is 308 309 clear from (figure 7a to d) the region is not receiving the much convective precipitation even 310 if there is huge cloud cover, which leads to a conclusion of orographic precipitation over the region (figure 7). However, the relative humidity is very high over the sampling site from 311 312 lower to upper middle level of the atmosphere during the study period (figure S3). During May 313 to September the convective rain is receiving most part of the Sikkim approved that the region 314 has high convective activity added from the Bay of Bengal as the monsoon season. Again, 315 from October to April the region is not receiving the convective rain even though there is 316 strong cloud cover pointing toward the orographic rainfall over entire Sikkim. That's making the Sikkim unique weather condition (figure S3 and S4). And, least concentration of BC, BCff, 317 318 BCbb and BrC is observed during the monsoon months supporting the convective rain (i.e., 319 rain out scavenging) of all pollutants. The BC and BrC have a significant positive correlation 320 with thermal and solar radiation, indicating positive radiative feedback. A stronger negative 321 correlation between CO2 and surface thermal radiation (STR) and surface solar radiation 322 (SSR) would have significant implications (figure 5). The negative correlation between CO2 323 and STR implies that as the concentration of CO2 in the atmosphere increases, the amount of 324 heat radiated from the Earth's surface into space decreases. This can lead to an increase in the 325 Gangtok's temperature, which can have various impacts on climate and weather as well (figure 326 S1, and 5). The negative correlation between CO2 and SSR implies that as the concentration





of CO2 in the atmosphere increases, the amount of solar radiation absorbed by the Earth's surface decreases (figure 5). Overall, a significant negative correlation between CO2 and STR/SSR would indicate a stronger influence of greenhouse gas concentrations on the surface's radiation balance and would have important implications for climate change as well as anomalous warming over the Gangtok region (figure S1).

332 5.0 Conclusions

In accordance with the LULC between 2000 and 2010, Sikkim's recent changes to its 333 development regulations have resulted in the majority of fallow land being consumed by 334 construction, which suggests that Gangtok's urban settlement load has increased significantly. 335 336 In addition, the LULC for 2020 depicts a booming built-up region over the GMC. Since 2000 337 to 2020, the vegetation cover has likewise decreased. However, due to the seasonal nature, 338 streams are lesser in 2020, indicating precipitation pattern variation over GMC. The areas 339 covered in dense vegetation are also being consumed by the expanding built-up area. The present study is the report of newly produced data BC and BrC for the fragile region of 340 341 Himalayas and relation with meteorological conditions. It has been observed that the 342 temperature over Gangtok is increasing as well. The peak concentration of BC/BrC has been found during October 2021 and March 2021 and 2022. The diurnal distribution of BC/BrC 343 344 suggests the two peaks in a day, first in the 8-10AM and second in 9-11PM. The meteorological 345 conditions for the same has been observed to be favourable to diurnal variation of BC/BrC 346 concentration. In the monthly variation of the BC/BrC is delineated that the peak concentration 347 of BC, BCbb, BCff, during March 2022. However, BrC and BB% have maximum concentration during April 2021. BB% and BrC as well as BB and carbon dioxide have a strong 348 349 significant positive correlation coefficient, which is evidence that biomass burning is a 350 substantial factor in the rise in carbon dioxide levels. In addition to this, there is a strong, 351 positive correlation between CO₂ and BC/BCff, indicating that burning fossil fuels is also one 352 of the causes of rising CO₂ levels. The net thermal radiation, net solar radiation and BC, BrC 353 relationship suggested that the BC, and BrC have positive radiative forcing. Furthermore, the 354 monsoon months show the lowest concentrations of BC, BCbb, BCff, BrC, and BB%, demonstrating the convective rain (i.e., rain out scavenging) ability to remove majority of 355 356 contaminants. Both the BC and BrC reveal evidence of positive radiative feedback.

357 Data Availability

Data is provided in the 'supplementary section' and for further detail knowledge about it canbe available from the corresponding author on the adequate request.





- 360 Data link for the data access:
- 361 https://docs.google.com/spreadsheets/d/1N4F_fT68syY6n0UIfA6nzI5o-
- 362 <u>8LUWjyFfk5NpfquRyg/edit?usp=sharing</u>
- 363 Conflict of Interest
- 364 None conflict of interest.

365 Authors Contribution

- 366 Dr. Pramod Kumar: conceptualization, drafting, writing, figures, and editing
- 367 Ms. Khushboo Sharma: sampling, data analysis and figures.
- 368 Ms. Ankita Malu: data analysis, figures, and editing
- 369 Mr. Rajeev Rajak: editing
- 370 Ms. Aparna Gupta: editing
- 371 Mr. Bidyutjyoti Baruah: editing
- 372 Mr. Jayant Sharma: sampling
- 373 Dr. Shailesh Yadav: editing, and mentoring
- 374 Dr. Thupstan Angchuk: editing, and mentoring
- 375 Dr. Rakesh Kumar Ranjan: conceptualization, data interpretation, mentoring, and editing.
- 376 Dr. Nishchal Wanjari: editing and mentoring.
- 377 Dr. Anil Kumar Misra: editing and mentoring.

378 Acknowledgements

- 379 Authors acknowledge to the Department of Science and Technology, Government of India,
- and host department "DST's Centre of Excellence (CoE), at Department of Geology, Sikkim
- 381 University, DST/CCP/CoE/186/2019 (G)," for the generation of BC/BrC data. We also
- acknowledge to free data sources used in the study as ERA5, USGS earth explorer. Authors
- appreciate freely available software such as R-studio, QGIS, CDO, and GrADS used for the
- analysis and visualization. We also acknowledge the anonymous persons whom so ever have
- helped and supported for the Black Carbon data collection.

386 **References**

Aithal, B. H., & MC, C. (2019). Assessing land surface temperature and land use change
through spatio-temporal analysis: a case study of select major cities of India. Arabian Journal
of Geosciences, 12(11), 1-16. <u>https://doi.org/10.1007/s12517-019-4547-1</u>

Benesty, J., Chen, J., Huang, Y., and Cohen, I. (2009). Pearson correlation coefficient. In Noise
reduction in speech processing (pp. 1-4). Springer, Berlin, Heidelberg.
<u>https://doi.org/10.1007/978-3-642-00296-0 5</u>





- Bharali, C., Nair, V. S., Chutia, L., & Babu, S. S. (2019). Modeling of the effects of wintertime
 aerosols on boundary layer properties over the Indo Gangetic Plain. Journal of Geophysical
- 395 Research: Atmospheres, 124(7), 4141-4157. https://doi.org/10.1029/2018JD029758
- Bhat, M. A., Romshoo, S. A., & Beig, G. (2022). Characteristics, source apportionment and
 long-range transport of black carbon at a high-altitude urban centre in the Kashmir valley,
 North-western Himalaya. Environmental Pollution, 305, 119295.
 https://doi.org/10.1016/j.envpol.2022.119295
- Bisht, D.S., Dumka, U.C., Kaskaoutis, D.G., Pipal, A.S., Srivastava, A.K., Soni, V.K., Attri,
 S.D., Sateesh, M. and Tiwari, S., (2015). Carbonaceous aerosols and pollutants over Delhi
 urban environment: temporal evolution, source apportionment and radiative forcing. Science
- 403 of the Total Environment, 521, 431-445. <u>https://doi.org/10.1016/j.scitotenv.2015.03.083</u>
- Bond, T. C., Streets, D. G., Yarber, K. F., Nelson, S. M., Woo, J. H., & Klimont, Z. (2004). A
 technology-based global inventory of black and organic carbon emissions from combustion.
 Journal of Geophysical Research: Atmospheres, 109(D14).
 https://doi.org/10.1029/2003JD003697
- 408 Evangelista, H., Maldonado, J., Godoi, R.H.M., Pereira, E.B., Koch, D., Tanizaki-Fonseca, K.,
- 409 Van Grieken, R., Sampaio, M., Setzer, A., Alencar, A. and Gonçalves, S.C. (2007). Sources
- 410 and transport of urban and biomass burning aerosol black carbon at the South–West Atlantic
- 411 Coast. Journal of Atmospheric Chemistry, 56(3), 225-238. <u>https://doi.org/10.1007/s10874-</u>
 412 006-9052-8
- Gupta, P., Singh, S. P., Jangid, A., & Kumar, R. (2017). Characterization of black carbon in
 the ambient air of Agra, India: Seasonal variation and meteorological influence. Advances in
 Atmospheric Sciences, 34(9), 1082-1094. https://doi.org/10.1007/s00376-017-6234-z
- Hansen, A. D. A., & Schnell, R. C. (2005). The aethalometer. Magee Scientific Company,
 Berkeley, California, USA, 7.
- Hansen, J., Lacis, A., Rind, D., Russell, G., Stone, P., Fung, I., Ruedy, R. and Lerner, J. (1984).
 Climate sensitivity: Analysis of feedback mechanisms. feedback, 1, 1-3.
- Helin, A., Virkkula, A., Backman, J., Pirjola, L., Sippula, O., Aakko-Saksa, P., Väätäinen, S.,
 Mylläri, F., Järvinen, A., Bloss, M. and Aurela, M. (2021). Variation of absorption Ångström
 exponent in aerosols from different emission sources. Journal of Geophysical Research:
 Atmospheres, 126(10), 2020JD034094. <u>https://doi.org/10.1029/2020JD034094</u>
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas,
 J., Peubey, C., Radu, R., Schepers, D. and Simmons, A. (2020). The ERA5 global reanalysis.
 Quarterly Journal of the Royal Meteorological Society, 146(730), 1999-2049.
 https://doi.org/10.1002/qj.3803
- Karra, K., Kontgis, C., Statman-Weil, Z., Mazzariello, J. C., Mathis, M., & Brumby, S. P.
 (2021). Global land use/land cover with Sentinel 2 and deep learning. In 2021 IEEE
 international geoscience and remote sensing symposium IGARSS (pp. 4704-4707). IEEE.
 https://doi.org/10.1109/IGARSS47720.2021.9553499
- Kedia, S., Ramachandran, S., Holben, B. N., & Tripathi, S. N. (2014). Quantification of aerosol
 type, and sources of aerosols over the Indo-Gangetic Plain. Atmospheric Environment, 98,
- 434 607-619. https://doi.org/10.1016/j.atmosenv.2014.09.022





- Kirchstetter, T. W., Novakov, T., & Hobbs, P. V. (2004). Evidence that the spectral dependence of light absorption by aerosols is affected by organic carbon. Journal of Geophysical Research: Atmospheres, 109(D21). <u>https://doi.org/10.1029/2004JD004999</u>
- Kiran, V. R., Talukdar, S., Ratnam, M. V., & Jayaraman, A. (2018). Long-term observations
 of black carbon aerosol over a rural location in southern peninsular India: Role of dynamics
 and meteorology. Atmospheric Environment, 189, 264-274.
 https://doi.org/10.1016/j.atmosenv.2018.06.020
- Klimont, Z., Kupiainen, K., Heyes, C., Purohit, P., Cofala, J., Rafaj, P., Borken-Kleefeld, J.
 and Schöpp, W. (2017). Global anthropogenic emissions of particulate matter including black
 carbon. Atmospheric Chemistry and Physics, 17(14), 8681-8723. <u>https://doi.org/10.5194/acp-17-8681-2017</u>
- Kumar, M., Parmar, K. S., Kumar, D. B., Mhawish, A., Broday, D. M., Mall, R. K., &
 Banerjee, T. (2018a). Long-term aerosol climatology over Indo-Gangetic Plain: Trend,
 prediction and potential source fields. Atmospheric environment, 180, 37-50.
 https://doi.org/10.1016/j.atmosenv.2018.02.027
- Kumar, M., Raju, M. P., Singh, R. S., & Banerjee, T. (2017). Impact of drought and normal
 monsoon scenarios on aerosol induced radiative forcing and atmospheric heating in Varanasi
 over middle Indo-Gangetic Plain. Journal of Aerosol Science, 113, 95-107.
 https://doi.org/10.1016/j.jaerosci.2017.07.016
- Kumar, P., Patton, A. P., Durant, J. L., & Frey, H. C. (2018b). A review of factors impacting
 exposure to PM2.5, ultrafine particles and black carbon in Asian transport microenvironments.
 Atmospheric environment, 187, 301-316. https://doi.org/10.1016/j.atmosenv.2018.05.046
- Kumar, R. R., Soni, V. K., & Jain, M. K. (2020). Evaluation of spatial and temporal
 heterogeneity of black carbon aerosol mass concentration over India using three year
 measurements from IMD BC observation network. Science of the Total Environment, 723,
 138060. https://doi.org/10.1016/j.scitotenv.2020.138060
- Laskin, A., Laskin, J., & Nizkorodov, S. A. (2015). Chemistry of atmospheric brown carbon.
 Chemical reviews, 115(10), 4335-4382. <u>https://doi.org/10.1021/cr5006167</u>
- Mahmood, R., Pielke Sr, R.A., Hubbard, K.G., Niyogi, D., Bonan, G., Lawrence, P., McNider,
 R., McAlpine, C., Etter, A., Gameda, S. and Qian, B. (2010). Impacts of land use/land cover
 change on climate and future research priorities. Bulletin of the American Meteorological
 Society, 91(1), 37-46. <u>https://doi.org/10.1175/2009BAMS2769.1</u>
- Massabò, D., Caponi, L., Bernardoni, V., Bove, M.C., Brotto, P., Calzolai, G., Cassola, F.,
 Chiari, M., Fedi, M.E., Fermo, P. and Giannoni, M. (2015). Multi-wavelength optical
 determination of black and brown carbon in atmospheric aerosols. Atmospheric Environment,
 108,1-12. <u>https://doi.org/10.1016/j.atmosenv.2015.02.058</u>
- 471 Moosmüller, H., Chakrabarty, R. K., Ehlers, K. M., & Arnott, W. P. (2011a). Absorption
- 472 Ångström coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical
 473 particles. Atmospheric Chemistry and Physics, 11(3), 1217-1225.
- 474 https://doi.org/10.1021/acs.estlett.8b00118
- 475 Osborne, S. R., Johnson, B. T., Haywood, J. M., Baran, A. J., Harrison, M. A. J., & McConnell,
- 476 C. L. (2008). Physical and optical properties of mineral dust aerosol during the Dust and





- 477 Biomass-burning Experiment. Journal of Geophysical Research: Atmospheres, 113(D23).
- 478 <u>https://doi.org/10.1029/2007JD009551</u>
- 479 Park, RJ, Kim, MJ, Jeong, JI, Youn, D., & Kim, S. (2010). A contribution of brown carbon
- 480 aerosol to the aerosol light absorption and its radiative forcing in East Asia. Atmospheric
- 481 Environment, 44 (11), 1414-1421. <u>https://doi.org/10.1016/j.atmosenv.2010.01.042</u>
- Pearson, K. (1909). Determination of the coefficient of correlation. Science, 30(757), 23-25.
 DOI:10.1126/science.30.757.23
- Pierrehumbert, R. T. (2014). Short-lived climate pollution. Annual Review of Earth and
 Planetary Sciences, 42, 341-379. DOI: 10.1146/annurev-earth-060313-054843
- Ramachandran, S., & Rupakheti, M. (2022). Trends in the types and absorption characteristics
 of ambient aerosols over the Indo-Gangetic Plain and North China Plain in last two decades.
 Science of The Total Environment, 831, 154867.
 https://doi.org/10.1016/j.scitotenv.2022.154867
- Ramachandran, S., Rupakheti, M., & Lawrence, M. G. (2020). Black carbon dominates theaerosol absorption over the Indo-Gangetic Plain and the Himalayan foothills. Environment
- 492 international, 142, 105814. <u>https://doi.org/10.1016/j.envint.2020.105814</u>
- Ramanathan, V., & Carmichael, G. (2008). Global and regional climate changes due to black
 carbon. Nature geoscience, 1(4), 221-227. <u>https://doi.org/10.1038/ngeo156</u>
- Rathod, T. D., & Sahu, S. K. (2022). Measurements of optical properties of black and brown
 carbon using multi-wavelength absorption technique at Mumbai, India. Journal of Earth
 System Science, 131(1), 32. <u>https://doi.org/10.1007/s12040-021-01774-0</u>
- Rathod, T., Sahu, S. K., Tiwari, M., Yousaf, A., Bhangare, R. C., & Pandit, G. G. (2017). Light
 absorbing properties of brown carbon generated from pyrolytic combustion of household
 biofuels. Aerosol and Air Quality Research, 17(1), 108-116.
 https://doi.org/10.4209/aaqr.2015.11.0639
- Reddy, M. S., & Venkataraman, C. (2002a). Inventory of aerosol and sulphur dioxide
 emissions from India: I—Fossil fuel combustion. Atmospheric Environment, 36(4), 677-697.
 https://doi.org/10.1016/S1352-2310(01)00463-0
- Reddy, M. S., & Venkataraman, C. (2002b). Inventory of aerosol and sulphur dioxide
 emissions from India. Part II—biomass combustion. Atmospheric Environment, 36(4), 699712. https://doi.org/10.1016/S1352-2310(01)00464-2
- Runa, F., Islam, M., Jeba, F., & Salam, A. (2022). Light absorption properties of brown carbon
 from biomass burning emissions. Environmental Science and Pollution Research, 29(14),
 21012-21022. <u>https://doi.org/10.1007/s11356-021-17220-z</u>
- 511 Sharma, K., Ranjan, R.K., Lohar, S., Sharma, J., Rajak, R., Gupta, A., Prakash, A. and Pandey,
- 512 A.K. (2022). Black Carbon Concentration during Spring Season at High Altitude Urban Center
- 513 in Eastern Himalayan Region of India. Asian Journal of Atmospheric Environment (AJAE),
- 514 16(1). https://doi.org/10.5572/ajae.2021.149
- 515 Shindell, D., Kuylenstierna, J.C., Vignati, E., van Dingenen, R., Amann, M., Klimont, Z.,
- 516 Anenberg, S.C., Muller, N., Janssens-Maenhout, G., Raes, F. and Schwartz, J. (2012).





- 517 Simultaneously mitigating near-term climate change and improving human health and food
 518 security. Science, 335(6065), 183-189. DOI: 10.1126/science.1210026
- Shukla, K. K., Sarangi, C., Attada, R., & Kumar, P. (2022). Characteristic dissimilarities
 during high aerosol loading days between western and eastern Indo-Gangetic Plain.
- Atmospheric Environment, 269, 118837. <u>https://doi.org/10.1016/j.atmosenv.2021.118837</u>
- 522 Sloss, L. (2012). Black carbon emissions in India. CCC/209. IEA Clean Coal Centre, London,523 38.
- Stevens, B., & Feingold, G. (2009). Untangling aerosol effects on clouds and precipitation in
 a buffered system. Nature, 461(7264), 607-613. <u>https://doi.org/10.1038/nature08281</u>
- Takemura, T., & Suzuki, K. (2019). Weak global warming mitigation by reducing black carbon
 emissions. Scientific reports, 9(1), 1-6. <u>https://doi.org/10.1038/s41598-019-41181-6</u>
- Venkataraman, C., Habib, G., Kadamba, D., Shrivastava, M., Leon, J.F., Crouzille, B.,
 Boucher, O. and Streets, D.G. (2006). Emissions from open biomass burning in India:
 Integrating the inventory approach with high-resolution Moderate Resolution Imaging
 Spectroradiometer (MODIS) active-fire and land cover data. Global biogeochemical cycles,
 20(2). <u>https://doi.org/10.1029/2005GB002547</u>
- Watham, T., Padalia, H., Srinet, R., Nandy, S., Verma, P. A., & Chauhan, P. (2021). Seasonal
 dynamics and impact factors of atmospheric CO2 concentration over subtropical forest
 canopies: observation from eddy covariance tower and OCO-2 satellite in Northwest
 Himalaya, India. Environmental Monitoring and Assessment, 193(2), 1-15.
 https://doi.org/10.1007/s10661-021-08896-4
- Yasunari, T., Bonasoni, P., Laj, P., Fujita, K., Vuillermoz, E., Marinoni, A., Cristofanelli, P.,
 Duchi, R., Tartari, G. and Lau, K.M. (2010). Estimated impact of black carbon deposition
 during pre-monsoon season from Nepal Climate Observatory–Pyramid data and snow albedo
 changes over Himalayan glaciers. Atmospheric Chemistry and Physics, 10(14), 6603-6615.
 https://doi.org/10.5194/acp-10-6603-2010
- Yue, S., Zhu, J., Chen, S., Xie, Q., Li, W., Li, L., Ren, H., Su, S., Li, P., Ma, H. and Fan, Y.
 (2022). Brown carbon from biomass burning imposes strong circum-Arctic warming. One
 Earth, 5(3), 293-304. <u>https://doi.org/10.1016/j.oneear.2022.02.006</u>
- Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z.,
 Zhao, Y., and Shen, Z. (2013). Chemical characterization and source apportionment of PM2.5
 in Beijing: seasonal perspective, Atmos. Chem. Phys., 13, 7053–7074,
 https://doi.org/10.5194/acp-13-7053-2013

















555

556 Figure 2. (a) The hourly observation of Black Carbon, Black Carbon through biomass burning, Black Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon 557 Dioxide (BC, BCbb, BCff, BrC, BB%, and CO₂, respectively) (The corresponding unit for BC, 558 BCbb, BCff, BrC: µg/m³; BB%: % and CO₂: ppm) for 16th March 2021 to 10th March 2022 559 over study location (lat:27.32; lon:88.61). The light colour shading refers $\pm \sigma$ standard 560 deviation for each variable. (b) Same as figure 2a, but for meteorological parameters as 561 562 dewpoint temperature (DewPT), temperature (Temp), surface pressure (SrfPres), windspeed, total precipitation (TP), and relative humidity (Rh) during 16th March 2021 to 10th March 2022. 563







Figure 3. (a) The daily mean of Black Carbon, Black Carbon through biomass burning, Black 566 Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon Dioxide 567 (BC, BCbb, BCff, BrC, BB%, and CO2, respectively) (The corresponding unit for BC, BCbb, 568 BCff, BrC: µg/m3; BB%: % and CO₂: ppm) for 16th March 2021 to 10th March 2022 over study 569 570 location (lat:27.32; lon:88.61). The light colour shading refers $\pm \sigma$ standard deviation for each 571 variable. (b) same as figure 3a, but for meteorological parameters as dewpoint temperature (DewPT), temperature (Temp), surface pressure (SrfPres), Windspeed, total precipitation (TP), 572 and relative humidity (Rh) during 1st January 2021 to 31st March 2022. 573







Figure 4. (a) The monthly mean of Black Carbon, Black Carbon through biomass burning, 576 577 Black Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon Dioxide (BC, BCbb, BCff, BrC, BB%, and CO2, respectively) (The corresponding unit for 578 BC, BCbb, BCff, BrC: µg/m3; BB%: % and CO2: ppm) for 16th March 2021 to 10th March 579 2022 over study location (lat:27.32; lon:88.61). The error bar shows $\pm \sigma$ standard deviation for 580 each variable. (b) Same as figure 4a, but for meteorological parameters as dewpoint 581 temperature (DewPT), temperature (Temp), surface pressure (SrfPres), windspeed, total 582 precipitation (TP), and relative humidity (Rh) during January 2021 to March 2022. 583

584







Figure 5. Correlation among BC, BCbb, BCff, BrC, BB%, CO₂ and, dewpoint temperature (DTmp), temperature (Temp), surface pressure (Press), Wind, total precipitation (TP), Relative humidity (Rh), net solar radiation (SSR), and net thermal radiation (STR). The (***) shows 99% significance, (**) shows 95% significance, (*) 90% significance and () shows no significance. The correlation coefficient values (-0.3 to -0.49) or (0.3 to 0.49) are considered as 'a good correlation', values \leq (-0.5) or \geq (0.5) considered as 'a strong correlation'.







593

Figure 6. Monthly total precipitation (cumulative) and wind circulation pattern during January
2021 to March 2022. The Shading shows precipitation pattern, and streamline shows wind
circulation.







598

Figure 7. Monthly convective rain and total cloud cover during January 2021 to March 2022.The shading shows convective rain pattern, and contour shows total cloud cover fraction.





602

List of Tables

Table 1. The details of datasets used for the present study.

604

Variables	Data sets	Years (Span)	Resolution		Source	Deference
variables			Temporal	Horizontal	Source	Kelerence
Black and Brown Carbon	Observation and analysis, data generated using Aethalometer AE33	March 2021- March 2022	Weekly	Point Location (Gangtok)	Original data generated	Present Study
Total precipitation Relative humidity Temperature (2 meter) Wind (surface wind) Surface pressure Dewpoint temperature Net solar	ERA5 (ECMWF)	2021 to 2022	Hourly	0.25 [°] * 0.25 [°]	ECMWF https://cds. climate.co pernicus.e u/cdsapp#! /dataset/re analysis- era5- single- levels?tab =form	Hersbach et al., 2020
and thermal radiation downward						
LULC	LandSat-5, LandSat-8 and earth explorer USGS	Decembe r 2000, Decembe r 2010, Decembe r 2020	2000, 2010, 2020	30m, 30m	earth explorer USGS. <u>https://eart</u> <u>hexplorer.</u> <u>usgs.gov/</u>	earth explorer USGS.
LULC	Sentinel-2 Esri Inc.	Decembe r 2021	2021	10 m	Esri Inc. https://ww w.arcgis.c om/home/i tem.html?i d=d3da5d d386d140 cf93fc9ec bf8da5e31	Karra et. al., 2021