

1 **Measurement report:**

2 **Intra-annual Variability of Black/Brown Carbon and Its Interrelation with**
3 **Meteorological Conditions over Gangtok, Sikkim**

4 Pramod Kumar¹, Khushboo Sharma¹, Ankita Malu², Rajeev Rajak², Aparna Gupta¹,
5 Bidyutjyoti Baruah¹, Shailesh Yadav¹, Thupstan Angchuk¹, Jayant Sharma¹, Rakesh Kumar
6 Ranjan^{1#}, Anil Kumar Misra¹, and Nishchal Wanjari¹

7 ¹DST's Centre of excellence on Water Resources, Cryosphere and Climate Change Studies,
8 Department of Geology, Sikkim University, Gangtok, Sikkim, India -737102

9 ²Department of Geology, Sikkim University, Gangtok, Sikkim, India -737102

10 #Corresponding Author: rkranjan@cus.ac.in

11
12 **Abstract**

13 Black carbon (BC) and brown carbon (BrC) have versatile natures, and they have an apparent
14 role in climate variability and changes. As the anthropogenic activity is surging, the BC and
15 BrC are also reportedly increasing. So, the monitoring of BC/BrC and observation of land use
16 land cover changes (LULCC) at a regional level are necessary for the various interconnected
17 meteorological phenomena changes. The current study investigates BC, BrC, CO₂, BC from
18 fossil fuels (BC_{ff}), BC from biomass burning (BC_{bb}), LULCC, and their relationship to the
19 corresponding meteorological conditions over Gangtok in the Sikkim Himalayan region. The
20 concentration of BC (BrC) 43.5 µg/m³ (32.0 µg/m³) was found to be highest during the March-
21 2022 (April-2021). Surface pressure exhibits a significant positive correlation with BC, BC_{ff},
22 BC_{bb}, and BrC. Higher surface pressure results in a calmer and more stable boundary layer,
23 which effectively retains deposited contaminants. Conversely, the wind appears to facilitate
24 the dispersion of pollutants, showing a strong negative correlation. The fact that all pollutants
25 and precipitation have been shown to behave similarly points to moist scavenging of the
26 pollutants. Despite the dense cloud cover, it is clear that the area is not receiving convective
27 precipitation, implying that orographic precipitation is occurring over the region. Most of
28 Sikkim receives convective rain from May to September, indicating that the region has
29 significant convective activity contributed from the Bay of Bengal during the monsoon season.
30 Furthermore, monsoon months have the lowest concentrations of BC, BC_{bb}, BC_{ff}, and BrC,
31 suggesting the potential of convective rain (as rain out scavenging) to remove most of the
32 pollutants. Moreover, BC and BrC show positive radiative feedback.

33 **Keywords:** Black carbon; Brown carbon; LULC; Sikkim Himalaya; Meteorology; Biomass
34 burning; Radiative forcing.

35 **1.0 Introduction**

36 Black carbon (BC), and brown carbon (BrC), are part of fine particulates in air pollution that
37 have a deceptive role in climate variability and changes. BC/BrC is a short-lived climate
38 pollutant with a lifetime of only days to weeks after release in the atmosphere (Pierrehumbert,
39 2014). During this short period of time, BC/BrC can have significant direct and indirect
40 impacts on the climate, cryosphere, agriculture, and human health (Shindell et al., 2012). It
41 consists of pure carbon in several interconnected forms. BC is formed through the incomplete
42 combustion of fossil fuels, biofuel, and biomass, and is one of the main types of particles in
43 both anthropogenic and naturally occurring soot (Bond et al., 2004). BrC in the atmosphere
44 has been attributed to the burning of biomass and fossil fuels, the biogenic release of fungi,
45 plant debris, and humic matter, and multiphase reactions between the gas-phase, particulate,
46 and cloud microdroplet constituents in the atmosphere (Laskin et al., 2015). BC/BrC is
47 transported from its source to many locations across the world (Ramanathan and Carmichael,
48 2008). The BC/BrC released into the atmosphere exhibits vertical distribution and follows the
49 prevailing wind speed and direction. It engages with various atmospheric components before
50 eventually settling on the Earth's surface through either wet or dry deposition processes. Its
51 hygroscopic properties render it more prone to cloud seeding and cloud formation, thereby
52 contributing directly to the precipitation mechanism in regions with high humidity (Stevens
53 and Feingold, 2009). In addition, it absorbs both incoming and outgoing radiation, atmospheric
54 BC/BrC modifies radiative forcing, disturbs atmospheric stability, regional circulation, and
55 rainfall pattern, affects cloud albedo, material damage, reduces agricultural productivity,
56 degrades ecosystem, and affects human health (Zhang et al., 2013). However, due to an
57 insufficiency of observations, BrC is one of the least understood and uncertain warming agents
58 (Yue et al., 2022). Numerous studies have been conducted to analyze the global distribution
59 of BC and BrC, including research focused on these species within India as well (Reddy and
60 Venkataraman, 2002a, 2002b; Venkataraman et al., 2006; Park et al., 2010; Sloss, 2012; Helin
61 et al., 2021; 2020; Kumar et al., 2020a; Watham et al., 2021; Bhat et al., 2022; Runa et al.,
62 2022; Yue et al., 2022; Kumar et al, 2018b). However, the overall worldwide BC emission is
63 estimated to be 4800-7200 Gg per year (Klimont et al., 2017). In 2001, India's total BC
64 emissions were projected to be 1343.78 Gg (Sloss, 2012). Residential fuel burning and
65 transportation contribute maximum to the global anthropogenic BC emission (Helin et al.,
66 2021). About 60 to 80% of residential fuels (coal and biomass) emissions are reported from
67 Asian and African countries, whereas approximately 70% of diesel engine emissions are found

68 to be from Europe, North America, and Latin America (Johnson et al., 2019; Ayompe et al.,
69 2021; Adeeyo et al., 2022; Sun et al., 2022).

70 On the other hand, emissions on the Indian subcontinent have increased by 40% since the year
71 2000 (Kurokawa and Ohara, 2020; Sun et al., 2022). According to Reddy and Venkataraman
72 (2002a, 2002b), the estimated BC emissions in India are fossil fuels, 100 Gg biofuel, 207 Gg
73 open burning, and 39 Gg with a climatic forcing of $+1.1 \text{ W/m}^2$, black carbon is the second-
74 most significant human emission in the current atmosphere (Sharma et al., 2022). BC
75 concentration was measured by Zhao et al. (2017) in the south-eastern Tibetan Plateau (TP).
76 Daily mean BC loadings ranged from 57.7 to 5368.9 ng/m^3 demonstrating a high BC burden
77 even at free tropospheric altitudes (Zhao et al., 2017). Black carbon (BC) deposition was
78 estimated at the Nepal Climate Observatory - Pyramid (NCO-P) site in the Himalayan region
79 during the pre-monsoon season (March-May). A total BC deposition rate of $2.89 \mu\text{g/m}^3/\text{day}$
80 was estimated, resulting in a total deposition of $266 \mu\text{g/m}^3$ for March–May (Yasunari et al.,
81 2010). From the Indian perspective, several key short-term incidents contribute to a rise in
82 India's BC concentration from biomass burning and other sources (Kumar et al., 2020a).
83 Burning agricultural waste (stubble) is widespread in India and several other nations. Many
84 studies suggest that increased BC in northern India, notably the Indo-Gangetic Plain (IGP) is
85 the global absorbing aerosol hotspot (Venkataraman et al., 2006; Ramanathan and Carmichael,
86 2008). In India, post-monsoon paddy crop waste burning occurs in the months of October and
87 November in the north and northwest parts of India (Venkataraman et al., 2006). In the north-
88 western Indo-Gangetic Plain (IGP) (especially- Punjab, Haryana, and western Uttar Pradesh),
89 stubble burning is a popular practice (Venkataraman et al., 2006). Long-distance transport of
90 BC aerosols, mostly from Asia to the North Pacific and South America to the southwest
91 Atlantic, is often recognized as a significant factor in local concentration (Evangelista et al.,
92 2007). However, in India, only local sources (89%) affect BC concentrations (Zhang et al.,
93 2013), as there aren't many movements of transboundary aerosols contribution over the IGP
94 (Kumar et al., 2018a; Kedia et al., 2014; Ramachandran and Rupakheti, 2022; Ramachandran
95 et al., 2020). Both marine and continental air masses contributed to total aerosol loading over
96 middle-IGP (Kumar et al., 2017; Shukla et al., 2022).

97 Black carbon is a light-absorbing particle that is released into the atmosphere directly in the
98 form of ultrafine ($<0.1 \mu\text{m}$) to fine particles ($<2.5 \mu\text{m}$) (Gupta et al., 2017). BC is a good tracer
99 for particle deposition as it is non-volatile, insoluble, and chemically inert, and it can also mix
100 well with other aerosol species in the atmosphere (Kiran et al., 2018). As a result, BC
101 deposition data are important not just for BC sinks but also for a broader understanding of

102 aerosol deposition. BC emissions are mostly influenced by significant changes in the energy
103 sector, fuel usage, industrial expansion, and an increase in the number of vehicles (Bisht et al.,
104 2015). Residential fuels like wood, agricultural waste, and cow dung used for cooking and
105 biomass usage for home purposes are the primary sources of BC emissions (Venkataraman et
106 al., 2006). The Asian mainland is a substantial contributor to global BC emissions and has
107 been identified as a hotspot (Gupta et al., 2017). BC has a high absorption ability, accounting
108 for 90-95 percent of total atmospheric aerosol absorption (Hansen et al., 1984). It can absorb
109 solar energy in the visible-infrared band and warm the environment. In comparison to carbon
110 dioxide, BC has a much shorter life cycle in the atmosphere. As a result, mitigation or reduction
111 has a greater positive impact on the atmosphere (Kirchstetter et al., 2004; Takemura and
112 Suzuki, 2019). Changing land use land cover (LULC) has a very significant impact on weather,
113 climate, and aerosols (Mahmood et al., 2010). It is well-established fact that the LULC change
114 has a direct relation with land surface temperature, vehicular emission, and anthropogenic
115 activity (Aithal and MC, 2019). This motivated the present study for further analysis of Sikkim
116 region land use land cover change and its relation with temperature and BC/BrC for March
117 2021 to March 2022. The current study's objectives are to assess the intra-annual variability of
118 Black/Brown Carbon (BC/BrC) (diurnal/daily/monthly) during the study period March-2021
119 to March-2022, as well as the interrelationship between meteorological conditions and
120 BC/BrC, along with LULC change for three decades 2000, 2010, and 2020, and its relationship
121 with anthropogenic activity over Gangtok.

122 **2.0 Study location**

123 The Gangtok Municipal Corporation (GMC) has been selected for the present study on the
124 basis of its urban exposure and settlement change for three decades as well as congruently
125 temperature rise (Figure S1). The sampling was carried out at the Pani House area in Gangtok,
126 GMC, having a longitude of 88.609°E and a latitude of 27.323°N. Sikkim is surrounded by
127 Nepal, China, and Bhutan from west, north, and east respectively, and consists of the trans and
128 greater Himalayan range. Moreover, Sikkim has one of the most fragile forest covers.
129 However, Gangtok is a densely populated city and capital of the state of Sikkim which is
130 situated in the East Sikkim district (see Figure 1a). The population of Sikkim has been found
131 to have increased as per the Indian census for three decades as can be seen in table S1.

132 **3.0 Data and Methodology**

133 The real-time sampling of BC was carried out from 10th March 2021 to 17th March 2022, at
134 Gangtok using the seven-channel dual spot Aethalometer (Model AE-33-7, Magee Scientific,

135 USA). The Aethalometer AE-33 is an aerosol instrument with a detection limit of <0.005
136 $\mu\text{g}/\text{m}^3$ for a 1-hour period and a measuring range of 0.01 to $100 \mu\text{g}/\text{m}^3$. It has a programmable
137 measuring frequency of 1 second or 1 minute and a programmable flow rate of 2 to 5 lpm. The
138 data was collected for the measurement of BC and BrC associated with particulate matter
139 having an aerodynamic diameter of less than $2.5 \mu\text{m}$ ($\text{PM}_{2.5}$). The concentration of BC, BrC,
140 BC_{bb} , and BC_{ff} have been estimated by the Carbonaceous Aerosol Analysis Tools (CAAT)
141 software tool from the Magee Scientific Aethalometer model AE33 (Hansen and Schnell,
142 2005). The carbon dioxide (CO_2) was measured using a CO_2 sensor (Vaisala-GMP343) which
143 is attached to the aethalometer. The inlet of the aethalometer was mounted at a height of 15 m
144 above ground level. One of the main sources of uncertainty in utilizing aerosol absorption
145 measurements to estimate the BrC absorption coefficient at 370 nm is the potential contribution
146 of other species, such as black carbon and dust, to the measured absorption. This can result in
147 an overestimation of BrC mass concentration, especially in environments where these species
148 coexist. However, the Sikkim region stands out for having one of the highest precipitation
149 levels globally and minimal dust pollution contribution. Consequently, there is likely to be less
150 over or underestimation. Therefore, in this study, mass concentration was employed to address
151 these uncertainties.

152 A new data set of BC, BrC, Black Carbon from biomass burning (BC_{bb}), Black Carbon from
153 fossil fuels (BC_{ff}), the percentage contribution of biomass burning to BC (BB%) and CO_2 has
154 been generated over the unreported region of Sikkim Himalaya. The diurnal and monthly data
155 sets of BC, BC_{bb} , BC_{ff} , BrC, BB%, and CO_2 have been given in the details in supplementary
156 materials (Table S2 and S3). In addition to this, the meteorological data has been selected for
157 ERA5 reanalysis for the study. LULC data has been taken from USGS earth explorers of 2000
158 and 2010 Landsat-5, 2020 Landsat-8, and 2021 for Sentinel-2 (Karra et al., 2021). LULC data
159 has been chosen for the month of December to minimize the cloud cover. The details of the
160 LULC calculation steps used are given in the supplementary section (methodology S1.3). The
161 brief of the data set is discussed in the table 1.

162 **3.1 Estimation of BrC**

163 The Carbonaceous Aerosol Analysis Tools (CAAT) software tool from the Magee Scientific
164 Aethalometer model AE33 was utilized to estimate the concentrations of BC, BrC, BC_{bb} , and
165 BC_{ff} . The absorption coefficients of BC and BrC were determined using the multi-wavelength
166 absorption coefficients provided by the aethalometer. The presence of BrC was identified by
167 observing the maximum light absorption between 370–590 nm, but its absorption may increase

168 significantly below this range depending on its composition. The attenuation of illumination
 169 measured in this study using the aethalometer was attributed solely to the contribution of BC
 170 and BrC. It is believed that the absorption coefficient at 370 nm measured by the aethalometer
 171 represents the combined absorption coefficients of BC and BrC, which is denoted as σ_{BC+BrC}
 172 (370 nm). This assumption is similar to the model used in the multi-wavelength absorbance
 173 analyzer (MWAA) approach for source allocation, as described in Massabò et al. (2015).
 174 Equation (1) was used to calculate the σ_{BrC} (370 nm) absorption coefficient (supplementary
 175 methodology S1), which involved subtracting the contribution of BC (σ_{BC} (370 nm)) from the
 176 observed absorption coefficient (σ_{BC+BrC} (370 nm)).

$$177 \quad \sigma_{BrC}(370 \text{ nm}) = \sigma_{BC+BrC}(370 \text{ nm}) - \sigma_{BC}(370 \text{ nm}) \quad \text{Eq. (1)}$$

178 The σ_{BC} (370 nm), was calculated by applying the power-law fit to absorption data in the 590-
 179 950 nm wavelength range provided in equation (1).

$$180 \quad \sigma_{BC}(\lambda) = \beta \lambda^{-AAE_{BC}} \quad \text{Eq. (2)}$$

181 The absorption angstrom exponent of BC is denoted as AAE_{BC} , with β being a constant value.
 182 As BC is a significant contributor to light absorption at wavelengths beyond 590 nm, the
 183 contribution of other aerosol species can be neglected, and the AAE_{BC} can be calculated using
 184 equation (3), as stated in Rathod and Sahu (2022). The AAE for both BC and BrC can be
 185 expressed as σ , and in this study, the AAE definition by Moosmüller et al. (2011a) was used
 186 instead of the AAE specified for a wavelength pair. This value is determined by equation (3),
 187 which calculates the negative log-log slope of the absorption spectrum at wavelength λ .

$$188 \quad AAE_{BC} = - \frac{d \ln \sigma_{BC}}{d \ln \lambda} \quad \text{Eq. (3)}$$

189 Instead of the conventional approach where AAE_{BC} is assumed to be 1, we utilized the AAE_{BC}
 190 that was observed onsite to calculate $\sigma_{BC}(\lambda)$. Equation (4) was employed to determine σ_{BrC}
 191 (370 nm) by substituting $\sigma_{BC}(\lambda)$ at 370 nm, which was obtained using equation (2) (Wang et
 192 al., 2020), into equation (4) (refer to supplementary methodology S1.1, S1.2, and Figure S2
 193 for details).

$$194 \quad \sigma_{BrC}(370 \text{ nm}) = \sigma_{BC+BrC}(370 \text{ nm}) - \beta(370 \text{ nm})^{-AAE_{BC}} \quad \text{Eq. (4)}$$

195 To calculate $\sigma_{BrC}(\lambda)$ at 470 nm and 520 nm, we can subtract the modelled BC from the
 196 measured absorption coefficients, in a similar manner. It is worth noting that the BrC
 197 absorption coefficients are very low at wavelengths beyond 590 nm (Wang et al., 2020),

198 according to Rathod et al. (2017) and Rathod and Sahu (2022), hence they are not taken into
199 account (supplementary methodology S1).

200 **3.2 Data Analysis**

201 LULC change also has a direct impact on vehicular emissions and other anthropogenic
202 activities. Urbanization, conceivably, can lead to increased vehicle traffic and emissions,
203 which can contribute to air pollution and climate change. Changes in land use can also affect
204 the amount and type of vegetation, which can influence the carbon cycle and the amount of
205 greenhouse gases in the atmosphere. The ERA-5 reanalysis data has been used for
206 meteorological analysis viz. wind pattern, precipitation, relative humidity, and temperature
207 (Hersbach et al., 2020). The hourly data has been taken for the analysis and then the daily,
208 monthly, and seasonal average has been computed for the study period over the Sikkim and
209 surrounding states for a better understanding of the meteorological conditions influencing the
210 BC, and BrC. The total precipitation is computed as a sum of the hourly data for a day to daily
211 total precipitation and further, it was summed for monthly cumulative total precipitation using
212 the sum formula as

$$213 \quad \text{Monthly Cumulative Total Precipitation} = \sum_i^n X \quad \text{Eq. (5)}$$

214 Where 'i' is the initial 'n' the last date and X is the hourly total precipitation taken from ERA5.
215 The wind circulation has been computed using the u-component and v-component of wind and
216 the wind speed has been calculated as

$$217 \quad \text{Wind Speed} = \sqrt{u^2 + v^2} \quad \text{Eq. (6)}$$

218 The temperature and relative humidity averaged have been computed using the mean formula
219 as

$$220 \quad \text{Average} = \frac{\sum_i^n X}{n} \quad \text{Eq. (7)}$$

221 Where, 'i' is the initial and 'n' last date of the variables such as temperature, relative
222 humidity, and wind components.

223 Let x and y be two real-valued random variables such that the correlation coefficient Spearman
224 Pearson can be calculated between the BC/BrC and meteorological parameters. The
225 Coefficient of Pearson Correlation (PCC) (Pearson, 1909; Benesty et al., 2009) as

$$PCC = \frac{n(\sum xy) - (\sum x)(\sum y)}{\sqrt{[n\sum x^2 - (\sum x)^2][n\sum y^2 - (\sum y)^2]}} \quad \text{Eq. (8)}$$

227 Where ‘n’ is the population size of the variables used for the study.

228 Table 1 contains additional information about the dataset, and a more detailed methodology
229 can be found in the supplementary section (S1).

230 **4.0 Results and Discussions**

231 The anthropogenic activities in Gangtok have drastically increased in the last 20 years. As
232 evident from Figures 1b, c, and d, LULC has been changed from 2000 to 2020 over the
233 Gangtok Municipal Corporation (GMC). Population change and growth have also been
234 observed in the Sikkim (Table S1). LULC during the years 2000 and 2010 evidently shows
235 that most of the fallow land has been built up due to a recent change in the policy of
236 construction in Sikkim suggesting urban settlement load over Gangtok has increased
237 significantly. As a result, there is a significant increase in built-up areas in GMC for the last
238 20 years. The vegetation cover has also reduced from 2000 to 2020 (Figure 1b, c, and d). The
239 rainfed water bodies are reducing from the GMC. However, due to its seasonal nature, streams
240 are lesser emerged in 2020 which perhaps shows the precipitation pattern alteration over GMC
241 due to the highly built-up sprawl. The built-up extent has been sprawling and consuming the
242 dense vegetation regions as well. This increases the study region's urge to be acknowledged so
243 that Sikkim's future policymakers can consider the effects of rising anthropogenic activities.
244 This anthropogenic activity leads to a heavy load on the environment over one of the cleanest
245 states of India. Long-term spatiotemporal variation of 2-meter air temperature justifies the
246 LULC change and warming pattern (Xiao-lei et al., 2022) over the Gangtok region (Figure
247 S1a, S1b, S1c, S1d, and S1e). The decadal warming rate is varying from 0.25° to 0.45°C
248 (Figure S1e). Thereafter, BC and BrC over the Gangtok have been measured to report the issue
249 and get more attention to the scientific and local community. The higher anthropogenic activity
250 releases a higher amount of emission in the name of development due to the population load
251 on the region (Shaddick et al., 2020) (i.e., the growth rate has been raised from 12.89 to 13.05%
252 in recent years) (Table S1). Diurnal variation of the BC, BrC, BC_{bb}, BC_{ff}, and CO₂ show two
253 peaks. BC, BC_{ff}, and CO₂ have almost similar time of peaks observed. The first peak is found
254 during 8-10 AM. And, the second peak is observed during 8-10 PM. However, BrC and BC_{bb}
255 have the peak concentration during 10-11 AM and 6-8 PM (Figure 2a), suggesting the peak
256 biomass burning time over the region. The meteorological conditions are observed as low

257 dewpoint, low temperature, high surface pressure, low wind speed, and high relative humidity
258 to the corresponding 8-10 AM, while the opposite is found in 8-10 PM referred to Figure 2b.

259 The daily time series of the BC, BC_{bb}, BC_{ff}, BrC, BB%, and CO₂ show the highest fluctuation
260 from 20th to 30th March in both 2021 and 2022 years respectively. The maximum BC (BrC)
261 content was found in March 2022 (April-2021), at 43.5 $\mu\text{g}/\text{m}^3$ (32 $\mu\text{g}/\text{m}^3$). The lowest
262 fluctuation is observed from 15th May to 15th September 2021 (Figure 3a). The intense peaks
263 of BC, BC_{ff}, and CO₂ were observed from 10th October to 15th November 2021 (Figure 3a)
264 which may be linked to the heavy tourist season of the state and indicate the traffic overload
265 in the Gangtok (Sharma et al, 2022). The meteorological conditions also favour similar
266 circumstances to accumulate the pollutant from 10th October to 15th November 2021 (Figure
267 3b). The lowest surface pressure with minimum fluctuation and the highest temperature and
268 dewpoint temperature with minimum fluctuation was noticed from the 15th June to 20th
269 September 2021 (Figure 3b). BrC is found to be the highest with significant variability from
270 the 10th of January to the 30th of March, pointing to winter wood burning for livelihood, which
271 is also supported by BC_{bb}. The monthly variations of BC, BC_{bb}, BC_{ff}, BrC, and BB% are
272 discussed in Figure 4a, and the highest value of standard deviation was observed during March
273 2022 for BC, BC_{ff}, and April 2021 for BC_{bb}, BrC, and BB%. The CO₂ is observed almost
274 constant with a small value of standard deviation. The maximum concentration of the BC, BC_{ff}
275 is found in March 2022. However, BC_{bb} and BrC were measured highest in April 2021. This
276 is probably inferring to high tourist season (i.e., vehicular emission) as well as random wood
277 burning at higher altitude regions surrounding the Gangtok. The minimum concentration of
278 the BrC was seen in the month of August 2021 as the highest total precipitation month with
279 high wind speed, temperature dewpoint temperature, and relative humidity (Figure 4b, S3, and
280 S4) (Rana et al., 2023).

281 The good correlation between BC and BC_{ff} showed that the primary source of BC is fossil fuel
282 combustion (Osborne et al, 2008; Jung et al., 2021). A significant correlation between BC_{bb}
283 and BrC indicates that biomass burning is a major contributor to BrC (Prabhu et al., 2020),
284 which is supported by the BB% and BrC (Figure 5). The positive correlation between CO₂ and
285 BC/BC_{ff} suggests that fossil fuel burning is influencing the CO₂ concentration (Rana et al.,
286 2023). Dewpoint temperature and CO₂ have a significant positive correlation suggesting
287 positive radiative forcing of CO₂ (Huang et al., 2017; Stjern et al., 2023). A similar relationship
288 has also been observed for temperature. BC_{bb}/BrC and temperature have a significant negative
289 correlation suggesting the negative radiative nature of the BC_{bb}/BrC (Figure S5). Moreover,
290 net thermal/solar radiation (STR/SSR) and BC/BrC have a significant positive correlation

291 (Figure 5, and S5) (Liu et al., 2020). A significant positive correlation between surface pressure
292 and BC/BC_{ff} (BC_{bb}/BrC) has been observed (Figure 5). Higher surface pressure creates calm
293 conditions and a stable boundary layer, which keeps the pollutants accumulated in the
294 boundary layer (Igarashi et al., 1988; Lee et al., 1995; Bharali et al., 2019; Liu et al., 2021).
295 However, the opposite has been observed for the wind indicating the dispersion of pollutants
296 with a strong negative correlation. A similar relationship has been observed between total
297 precipitation and all the pollutants, indicating the process of wet scavenging of pollutants (Yoo
298 et al., 2014; Ohata et al., 2016; Ge et al., 2021; Wu et al., 2022). The relative humidity is also
299 showing a similar result to the total precipitation with greater values of coefficient. The
300 negative correlation between total precipitation and surface pressure suggests that the rain falls
301 over the region mostly occurs in a low-pressure system that is caused due to the vertical rising
302 of an air parcel and causes condensation and precipitation (Johnson and Hamilton, 1988;
303 Sarkar, 2018). Aerosols, including black carbon (BC) and absorbing organic aerosol (brown
304 carbon, BrC), play a vital role as condensation nuclei for cloud-droplet growth, and a fraction
305 of mineral particles initiate the freezing of supercooled cloud droplets, leading to the release
306 of precipitation in the form of snow, hail, and rain (Mason, 1999). However, cloud
307 condensation nuclei formation and precipitation are prompted by primary aerosols, secondary
308 aerosols (such as nitrate, and sulfate), and BC/BrC (Ohata et al., 2016; Liu et al., 2020; Moteki,
309 2023). Moreover, BC particles are mainly hydrophobic and less efficient as CCN compared to
310 more hydrophilic particles; they can still act as CCN under certain conditions. These conditions
311 include the size and mixing state of the particles, as well as the atmospheric conditions such
312 as relative humidity and temperature (Ohata et al., 2016; Moteki, 2023; Liu et al., 2020). The
313 conditions required for BC particles to efficiently play the role of CCN depend on several
314 factors, including their size, mixing state, and atmospheric conditions (Moteki, 2023; Liu et
315 al., 2020). For example, smaller BC particles are more efficient as CCN than larger ones
316 (Moteki, 2023). The mixing state of BC particles also plays a role, as externally mixed BC
317 particles are less efficient as CCN than internally mixed ones (Liu et al., 2020). Atmospheric
318 conditions such as relative humidity and temperature also affect the efficiency of BC particles
319 as CCN (Moteki, 2023). For example, higher relative humidity and lower temperatures can
320 increase the efficiency of BC particles as CCN (Moteki, 2023). Additionally, relative humidity
321 over the study region is very high during the entire year with the favourable temperature.
322 Thereafter, BC and BrC have a crucial role in the precipitation mechanism (Zhu et al., 2021;
323 Li et al., 2023a) over the study region. Total precipitation and wind circulation indicated that
324 the study region received precipitation throughout each month of the study period (i.e., most

325 of the time in the form of rain and occasionally snow). Hence the maximum is observed in
326 August and the minimum in March 2022. The wind pattern illustrates the monsoon seasonal
327 strong influence from May to September 2021 (Figure 6). The wind converges in the valley
328 and diverges from the mountain for the rest of the period (figure 6). Because the strong wind
329 and heavy rainfall indicated pollution scavenging (rain out or wash out), it is significantly
330 negatively correlated as TP vs BC_{bb}; TP vs BC_{ff}; TP vs BrC (Figure 5).

331 The relative humidity and temperature follow the same pattern when the temperature gradients
332 change from January to December, resulting in a decrease in moisture content in the
333 atmosphere (Figure S6). The lowest in the month of February is observed and the temperature
334 gradient gets steep from November (Figure S6). The dewpoint temperature contour and surface
335 pressure shading match well suggesting that the surface pressure creates the dewpoint
336 temperature gradient and keeps it sustained and stable atmospheric condition (Jung et al., 2023)
337 (Figure S7). During the month of June, it is very peculiar that the dewpoint temperature
338 contours are wide and a very small gradient is observed (Figure 7). This points toward the
339 warm conditions during the June over entire Sikkim. The cloud cover and convective
340 precipitation over Sikkim are discussed in Figure 7. It is clear from (Figures 7a to d) that the
341 region is not receiving much convective precipitation even if there is huge cloud cover, which
342 leads to a conclusion of orographic precipitation over the region (Figure 7). However, the
343 relative humidity is very high over the sampling site from the lower to upper middle level of
344 the atmosphere during the study period (Figure S3). Most of Sikkim receives convective rain
345 from May to September, which indicates that the region has strong convective activity added
346 from the Bay of Bengal during the monsoon season (Rahman et al., 2012; Kumar et al., 2020b;
347 Kakkar et al., 2022; Biswas and Bhattacharya, 2023). Again, from October to April, the region
348 does not receive convective rain even though there is strong cloud cover pointing toward the
349 orographic rainfall over the entire Sikkim (Kumar and Sharma, 2023). That's making the
350 Sikkim unique weather conditions (Figures S3 and S4). And, the least concentration of BC,
351 BC_{ff}, BC_{bb}, and BrC is observed during the monsoon months. This observation supports the
352 convective rain, as rain out scavenging, of all pollutants (Liu et al., 2020; Moteki, 2023).
353 During the monsoon season, the region experiences high convective activity, which is added
354 from the Bay of Bengal (Brooks et al., 2019; Liu et al., 2020; Moteki, 2023; Sankar et al.,
355 2023). Convective rain is an effective process for removing air pollutants from the atmosphere
356 (Liu et al., 2020; Moteki, 2023). Wet removal of BC and BrC occurs via cloud particle
357 formation and subsequent conversion to precipitation or impaction processes with
358 hydrometeors below clouds during precipitation (Liu et al., 2020; Moteki, 2023; Sankar et al.,

2023). The BC and BrC have a significant positive correlation with thermal and solar radiation, indicating positive radiative feedback (Zhang et al., 2020; Wang et al., 2021; Li et al., 2023a). A stronger negative correlation between CO₂ and surface thermal radiation (STR) and surface solar radiation (SSR) would have significant implications (Figure 5). The negative correlation between CO₂ and STR implies that as the concentration of CO₂ in the atmosphere increases, the amount of heat radiating from the Earth's surface into space decreases (Zhang et al., 2020). This can lead to an increase in the Gangtok's temperature, which can have various impacts on climate and weather as well (Figures S1, and 5). The negative correlation between CO₂ and SSR implies that as the concentration of CO₂ in the atmosphere increases, the amount of solar radiation absorbed by the Earth's surface decreases (Davis, 2017; Zhang et al., 2020; Li et al., 2023b) (Figure 5). Overall, a significant negative correlation between CO₂ and STR/SSR would indicate a stronger influence of greenhouse gas concentrations on the surface's radiation balance (Chiodo et al., 2018) and would have important implications for climate change as well as anomalous warming over the Gangtok region (Figure S1).

5.0 Conclusions

In accordance with the LULC between 2000 and 2010, Sikkim's recent changes to its development regulations have resulted in the majority of fallow land being consumed by construction, which suggests that Gangtok's urban settlement load has increased significantly. In addition, the LULC for 2020 depicts a booming built-up region over the GMC. From 2000 to 2020, the vegetation cover has likewise decreased. However, due to the seasonal nature, streams are lesser in 2020, indicating precipitation pattern variation over GMC. The areas covered in dense vegetation are also being consumed by the expanding built-up area. The present study is the report of newly produced data BC and BrC for the fragile region of the Himalayas and its relation with meteorological conditions. It has been observed that the temperature over Gangtok is increasing as well. The peak concentration of BC/BrC has been found during October 2021, March 2021, and 2022. The diurnal distribution of BC/BrC suggests the two peaks in a day, first at 8-10 AM and second at 9-11 PM. The meteorological conditions for the same have been observed to be favourable to diurnal variation of BC/BrC concentration. The monthly variation of the BC/BrC delineated the peak concentration of BC, BC_{bb}, and BC_{ff}, during March 2022. However, BrC and BB% have maximum concentration during April 2021. BB% and BrC as well as BB and carbon dioxide have a strong significant positive correlation coefficient, which is evidence that biomass burning is a substantial factor in the rise in carbon dioxide levels. In addition to this, there is a strong, positive correlation between CO₂ and BC/BC_{ff}, indicating that burning fossil fuels is also one of the causes of

393 rising CO₂ levels. The net thermal radiation, net solar radiation, and BC, BrC relationship
394 suggested that BC and BrC have positive radiative forcing. Furthermore, the monsoon months
395 show the lowest concentrations of BC, BC_{bb}, BC_{ff}, BrC, and BB%, demonstrating the
396 convective rain (i.e., rain out scavenging) ability to remove a majority of contaminants. Both
397 the BC and BrC reveal evidence of positive radiative feedback. BC particles in the atmosphere
398 have a strong ability to absorb solar radiation, and their lifetime depends on atmospheric
399 transport, aging, and wet scavenging processes. Organic aerosols, including BrC, can undergo
400 photochemical aging, affecting their ability to act as cloud condensation nuclei (CCN). The
401 effective density of BC is a crucial factor in evaluating its climate effect, and variations in BC
402 density can lead to uncertainties in predicting CCN number concentration.

403 **Data Availability**

404 Data is provided in the ‘supplementary section’ and for further detail knowledge about it can
405 be available from the corresponding author on the adequate request.

406 Data link for the data access:

407 [https://docs.google.com/spreadsheets/d/1N4F_ft68syY6n0UIfA6nzI5o-](https://docs.google.com/spreadsheets/d/1N4F_ft68syY6n0UIfA6nzI5o-8LUWjyFfk5NpfquRyg/edit?usp=sharing)
408 [8LUWjyFfk5NpfquRyg/edit?usp=sharing](https://docs.google.com/spreadsheets/d/1N4F_ft68syY6n0UIfA6nzI5o-8LUWjyFfk5NpfquRyg/edit?usp=sharing)

409 **Conflict of Interest**

410 None conflict of interest.

411 **Authors Contribution**

412 Dr. Pramod Kumar: conceptualization, drafting, writing, figures, and editing

413 Ms. Khushboo Sharma: sampling, data analysis, and figures.

414 Ms. Ankita Malu: data analysis, figures, and editing

415 Mr. Rajeev Rajak: editing

416 Ms. Aparna Gupta: editing

417 Mr. Bidyutjyoti Baruah: editing

418 Mr. Jayant Sharma: sampling

419 Dr. Shailesh Yadav: editing, and mentoring

420 Dr. Thupstan Angchuk: editing, and mentoring

421 Dr. Rakesh Kumar Ranjan: conceptualization, data interpretation, mentoring, and editing.

422 Dr. Nishchal Wanjari: editing and mentoring.

423 Dr. Anil Kumar Misra: editing and mentoring.

424 **Acknowledgments**

425 Authors acknowledge to the Department of Science and Technology, Government of India,
426 and host department “DST’s Centre of Excellence (CoE), at Department of Geology, Sikkim
427 University, DST/CCP/CoE/186/2019 (G),” for the generation of BC/BrC data. We also
428 acknowledge to free data sources used in the study as ERA5, and USGS earth explorer.
429 Authors appreciate freely available software such as R-studio, QGIS, CDO, and GrADS used
430 for the analysis and visualization. We also acknowledge Anirud Rai, Kuldeep Dutta, Abhinav
431 Tiwari, Richard Rai, and the anonymous persons who so ever have helped and supported the
432 Black Carbon data collection.

433 **References**

- 434 Adeeyo, R.O., Edokpayi, J.N., Volenzo, T.E., Odiyo, J.O. and Piketh, S.J.: Determinants of
435 solid fuel use and emission risks among households: insights from Limpopo, South Africa.
436 *Toxics*, 10(2), p.67, <https://doi.org/10.3390/toxics10020067>, 2022.
- 437 Aithal, B. H., & MC, C.: Assessing land surface temperature and land use change through
438 spatio-temporal analysis: a case study of select major cities of India. *Arabian Journal of*
439 *Geosciences*, 12(11), 1-16, <https://doi.org/10.1007/s12517-019-4547-1>, 2019.
- 440 Ayompe, L.M., Davis, S.J. and Egoh, B.N.: Trends and drivers of African fossil fuel CO2
441 emissions 1990–2017. *Environmental Research Letters*, 15(12), p.124039, DOI 10.1088/1748-
442 9326/abc64f, 2021.
- 443 Benesty, J., Chen, J., Huang, Y., and Cohen, I.: Pearson correlation coefficient. In *Noise*
444 *reduction in speech processing* (pp. 1-4). Springer, Berlin, Heidelberg.
445 https://doi.org/10.1007/978-3-642-00296-0_5, 2009.
- 446 Bharali, C., Nair, V. S., Chutia, L., & Babu, S. S.: Modeling of the effects of wintertime
447 aerosols on boundary layer properties over the Indo Gangetic Plain. *Journal of Geophysical*
448 *Research: Atmospheres*, 124(7), 4141-4157, <https://doi.org/10.1029/2018JD029758>, 2019.
- 449 Bhat, M. A., Romshoo, S. A., & Beig, G.: Characteristics, source apportionment and long-
450 range transport of black carbon at a high-altitude urban centre in the Kashmir valley, North-
451 western Himalaya. *Environmental Pollution*, 305, 119295,
452 <https://doi.org/10.1016/j.envpol.2022.119295>, 2022.
- 453 Bisht, D.S., Dumka, U.C., Kaskaoutis, D.G., Pipal, A.S., Srivastava, A.K., Soni, V.K., Attri,
454 S.D., Sateesh, M. and Tiwari, S.: Carbonaceous aerosols and pollutants over Delhi urban
455 environment: temporal evolution, source apportionment and radiative forcing. *Science of the*
456 *Total Environment*, 521, 431-445, <https://doi.org/10.1016/j.scitotenv.2015.03.083>, 2015.
- 457 Biswas, J. and Bhattacharya, S.: Future changes in monsoon extreme climate indices over the
458 Sikkim Himalayas and West Bengal. *Dynamics of Atmospheres and Oceans*, 101, p.101346.
459 <https://doi.org/10.1016/j.dynatmoce.2022.101346>, 2023.
- 460 Bond, T. C., Streets, D. G., Yarber, K. F., Nelson, S. M., Woo, J. H., & Klimont, Z.: A
461 technology-based global inventory of black and organic carbon emissions from combustion.

462 Journal of Geophysical Research: Atmospheres, 109(D14),
463 <https://doi.org/10.1029/2003JD003697>, 2004.

464 Brooks, J., Liu, D., Allan, J.D., Williams, P.I., Haywood, J., Highwood, E.J., Kompalli, S.K.,
465 Babu, S.S., Satheesh, S.K., Turner, A.G. and Coe, H.: Black carbon physical and optical
466 properties across northern India during pre-monsoon and monsoon seasons. *Atmospheric
467 Chemistry and Physics*, 19(20), pp.13079-13096. <https://doi.org/10.5194/acp-19-13079-2019>,
468 2019.

469 Chiodo, G., Polvani, L.M., Marsh, D.R., Stenke, A., Ball, W., Rozanov, E., Muthers, S. and
470 Tsigaridis, K.: The response of the ozone layer to quadrupled CO₂ concentrations. *Journal of
471 Climate*, 31(10), pp.3893-3907, doi: 10.1175/jcli-d-19-0086.1, 2018.

472 Davis, W.J.: The relationship between atmospheric carbon dioxide concentration and global
473 temperature for the last 425 million years. *Climate*, 5(4), p.76,
474 <https://doi.org/10.3390/cli5040076>, 2017.

475 Evangelista, H., Maldonado, J., Godoi, R.H.M., Pereira, E.B., Koch, D., Tanizaki-Fonseca, K.,
476 Van Grieken, R., Sampaio, M., Setzer, A., Alencar, A. and Gonçalves, S.C.: Sources and
477 transport of urban and biomass burning aerosol black carbon at the South–West Atlantic Coast.
478 *Journal of Atmospheric Chemistry*, 56(3), 225-238, [https://doi.org/10.1007/s10874-006-
479 9052-8](https://doi.org/10.1007/s10874-006-9052-8), 2007.

480 Ge, B., Xu, D., Wild, O., Yao, X., Wang, J., Chen, X., Tan, Q., Pan, X. and Wang, Z.: Inter-
481 annual variations of wet deposition in Beijing from 2014–2017: implications of below-cloud
482 scavenging of inorganic aerosols. *Atmospheric Chemistry and Physics*, 21(12), pp.9441-9454,
483 <https://doi.org/10.5194/acp-21-9441-2021>, 2021.

484 Gupta, P., Singh, S. P., Jangid, A., & Kumar, R.: Characterization of black carbon in the
485 ambient air of Agra, India: Seasonal variation and meteorological influence. *Advances in
486 Atmospheric Sciences*, 34(9), 1082-1094, <https://doi.org/10.1007/s00376-017-6234-z>, 2017.

487 Hansen, A. D. A., & Schnell, R. C.: *The aethalometer*. Magee Scientific Company, Berkeley,
488 California, USA, 7., 2005.

489 Hansen, J., Lacis, A., Rind, D., Russell, G., Stone, P., Fung, I., Ruedy, R. and Lerner, J.:
490 *Climate sensitivity: Analysis of feedback mechanisms*. feedback, 1, 1-3, 1984.

491 Helin, A., Virkkula, A., Backman, J., Pirjola, L., Sippula, O., Aakko-Saksa, P., Väättäinen, S.,
492 Mylläri, F., Järvinen, A., Bloss, M. and Aurela, M.: Variation of absorption Ångström
493 exponent in aerosols from different emission sources. *Journal of Geophysical Research:*
494 *Atmospheres*, 126(10), 2020JD034094, <https://doi.org/10.1029/2020JD034094>, 2021.

495 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas,
496 J., Peubey, C., Radu, R., Schepers, D. and Simmons, A.: The ERA5 global reanalysis.
497 *Quarterly Journal of the Royal Meteorological Society*, 146(730), 1999-2049,
498 <https://doi.org/10.1002/qj.3803>, 2020.

499 Huang, Y., Xia, Y. and Tan, X.: On the pattern of CO₂ radiative forcing and poleward energy
500 transport. *Journal of Geophysical Research: Atmospheres*, 122(20), pp.10-578,
501 <https://doi.org/10.1002/2017JD027221>, 2017.

- 502 Igarashi, S., Sasaki, H. & Honda, M.: Influence of pressure gradient upon boundary layer
503 stability and transition. *Acta Mechanica* 73, 187–198, <https://doi.org/10.1007/BF01177038>,
504 1988.
- 505 Johnson, R.H. and Hamilton, P.J.: The relationship of surface pressure features to the
506 precipitation and airflow structure of an intense midlatitude squall line. *Monthly Weather*
507 *Review*, 116(7), pp.1444-1473, [https://doi.org/10.1175/1520-0493\(1988\)116<1444:TROSPF>2.0.CO;2](https://doi.org/10.1175/1520-0493(1988)116<1444:TROSPF>2.0.CO;2) , 1988.
- 509 Johnson, M.A., Garland, C.R., Jagoe, K., Edwards, R., Ndemere, J., Weyant, C., Patel, A.,
510 Kithinji, J., Wasirwa, E., Nguyen, T. and Khoi, D.D., (2019). In-home emissions performance
511 of cookstoves in Asia and Africa. *Atmosphere*, 10(5), p.290.
512 <https://doi.org/10.3390/atmos10050290>
- 513 Jung, K.H., Goodwin, K.E., Perzanowski, M.S., Chillrud, S.N., Perera, F.P., Miller, R.L. and
514 Lovinsky-Desir, S.: Personal exposure to black carbon at school and levels of Fractional
515 Exhaled nitric Oxide in New York city. *Environmental Health Perspectives*, 129(9), p.097005,
516 <https://doi.org/10.1289/EHP8985>, 2021.
- 517 Jung, C.H., Lee, H.M., Park, D., Yoon, Y.J., Choi, Y., Um, J., Lee, S.S., Lee, J.Y. and Kim,
518 Y.P.: Parameterization of below-cloud scavenging for polydisperse fine mode aerosols as a
519 function of rain intensity. *Journal of Environmental Sciences*, 132, pp.43-55,
520 <https://doi.org/10.1016/j.jes.2022.07.031>, 2023.
- 521 Karra, K., Kontgis, C., Statman-Weil, Z., Mazzariello, J. C., Mathis, M., & Brumby, S. P.:
522 Global land use/land cover with Sentinel 2 and deep learning. In 2021 IEEE international
523 geoscience and remote sensing symposium IGARSS (pp. 4704-4707). IEEE,
524 <https://doi.org/10.1109/IGARSS47720.2021.9553499> , 2021.
- 525 Kedia, S., Ramachandran, S., Holben, B. N., & Tripathi, S. N.: Quantification of aerosol type,
526 and sources of aerosols over the Indo-Gangetic Plain. *Atmospheric Environment*, 98, 607-619,
527 <https://doi.org/10.1016/j.atmosenv.2014.09.022>, 2014.
- 528 Kirchstetter, T. W., Novakov, T., & Hobbs, P. V.: Evidence that the spectral dependence of
529 light absorption by aerosols is affected by organic carbon. *Journal of Geophysical Research:*
530 *Atmospheres*, 109(D21), <https://doi.org/10.1029/2004JD004999>, 2004.
- 531 Kiran, V. R., Talukdar, S., Ratnam, M. V., & Jayaraman, A.: Long-term observations of black
532 carbon aerosol over a rural location in southern peninsular India: Role of dynamics and
533 meteorology. *Atmospheric Environment*, 189, 264-274,
534 <https://doi.org/10.1016/j.atmosenv.2018.06.020>, 2018.
- 535 Kakkar, A., Rai, P.K., Mishra, V.N. and Singh, P.: Decadal trend analysis of rainfall patterns
536 of past 115 years & its impact on Sikkim, India. *Remote Sensing Applications: Society and*
537 *Environment*, 26, p.100738, <https://doi.org/10.1016/j.rsase.2022.100738>, 2022.
- 538 Klimont, Z., Kupiainen, K., Heyes, C., Purohit, P., Cofala, J., Rafaj, P., Borken-Kleefeld, J.
539 and Schöpp, W.: Global anthropogenic emissions of particulate matter including black carbon.
540 *Atmospheric Chemistry and Physics*, 17(14), 8681-8723, <https://doi.org/10.5194/acp-17-8681-2017>, 2017.
- 542 Kumar, M., Parmar, K. S., Kumar, D. B., Mhawish, A., Broday, D. M., Mall, R. K., &
543 Banerjee, T.: Long-term aerosol climatology over Indo-Gangetic Plain: Trend, prediction and

544 potential source fields. *Atmospheric environment*, 180, 37-50,
545 <https://doi.org/10.1016/j.atmosenv.2018.02.027>, 2018a.

546 Kumar, M., Raju, M. P., Singh, R. S., & Banerjee, T.: Impact of drought and normal monsoon
547 scenarios on aerosol induced radiative forcing and atmospheric heating in Varanasi over
548 middle Indo-Gangetic Plain. *Journal of Aerosol Science*, 113, 95-107,
549 <https://doi.org/10.1016/j.jaerosci.2017.07.016>, 2017.

550 Kumar, P., Patton, A. P., Durant, J. L., & Frey, H. C.: A review of factors impacting exposure
551 to PM_{2.5}, ultrafine particles and black carbon in Asian transport microenvironments.
552 *Atmospheric environment*, 187, 301-316, <https://doi.org/10.1016/j.atmosenv.2018.05.046>,
553 2018b.

554 Kumar, R. R., Soni, V. K., & Jain, M. K.: Evaluation of spatial and temporal heterogeneity of
555 black carbon aerosol mass concentration over India using three year measurements from IMD
556 BC observation network. *Science of the Total Environment*, 723, 138060,
557 <https://doi.org/10.1016/j.scitotenv.2020.138060>, 2020a.

558 Kumar, P., Sharma, M.C., Saini, R. and Singh, G.K.: Climatic variability at Gangtok and
559 Tadong weather observatories in Sikkim, India, during 1961–2017. *Scientific reports*, 10(1),
560 p.15177, <https://doi.org/10.1038/s41598-020-71163-y>, 2020b.

561 Kumar, P. and Sharma, M.C.: Frontal changes in medium-sized glaciers in Sikkim, India
562 during 1988–2018: Insights for glacier-climate synthesis over the Himalaya. *Iscience*, 26(10),
563 DOI: 10.1016/j.isci.2023.107789, 2023.

564 Kurokawa, J. and Ohara, T.: Long-term historical trends in air pollutant emissions in Asia:
565 Regional Emission inventory in ASia (REAS) version 3. *Atmospheric Chemistry and Physics*,
566 20(21), pp.12761-12793, <https://doi.org/10.5194/acp-20-12761-2020>, 2020.

567 Laskin, A., Laskin, J., & Nizkorodov, S. A.: Chemistry of atmospheric brown carbon.
568 *Chemical reviews*, 115(10), 4335-4382, <https://doi.org/10.1021/cr5006167>, 2015.

569 Lee, T., Fisher, M., & Schwarz, W.: Investigation of the effects of a compliant surface on
570 boundary-layer stability. *Journal of Fluid Mechanics*, 288, 37-58,
571 doi:10.1017/S0022112095001054, 1995.

572 Li, S., Zhang, H., Wang, Z., Chen, Y.: Advances in the Research on Brown Carbon Aerosols:
573 Its Concentrations, Radiative Forcing, and Effects on Climate. *Aerosol Air Qual. Res.* 23,
574 220336, <https://doi.org/10.4209/aaqr.220336>, 2023a.

575 Lin, J., Guo, Y., Li, J., Shao, M. and Yao, P.: Spatial and temporal characteristics of carbon
576 emission and sequestration of terrestrial ecosystems and their driving factors in mainland
577 China—a case study of 352 prefectural administrative districts. *Frontiers in Ecology and*
578 *Evolution*, 11, p.1169427, <https://doi.org/10.3389/fevo.2023.1169427>, 2023b.

579 Liu, D., He, C., Schwarz, J.P. and Wang, X.: Lifecycle of light-absorbing carbonaceous
580 aerosols in the atmosphere. *NPJ Climate and Atmospheric Science*, 3(1), p.40,
581 <https://doi.org/10.1038/s41612-020-00145-8>, 2020.

582 Liu, C., Huang, J., Tao, X., Deng, L., Fang, X., Liu, Y., Luo, L., Zhang, Z., Xiao, H.W. and
583 Xiao, H.Y.: An observational study of the boundary-layer entrainment and impact of aerosol

584 radiative effect under aerosol-polluted conditions. *Atmospheric Research*, 250, p.105348,
585 <https://doi.org/10.1016/j.atmosres.2020.105348>, 2021.

586 Mahmood, R., Pielke Sr, R.A., Hubbard, K.G., Niyogi, D., Bonan, G., Lawrence, P., McNider,
587 R., McAlpine, C., Etter, A., Gameda, S. and Qian, B.: Impacts of land use/land cover change
588 on climate and future research priorities. *Bulletin of the American Meteorological Society*,
589 91(1), 37-46, <https://doi.org/10.1175/2009BAMS2769.1>, 2010.

590 Mason, J.: The role of aerosols in cloud physics and climate change. *Science Progress*, 82(3),
591 185-207, <https://doi.org/10.1177/003685049908200301>, 1999.

592 Massabò, D., Caponi, L., Bernardoni, V., Bove, M.C., Brotto, P., Calzolari, G., Cassola, F.,
593 Chiari, M., Fedi, M.E., Fermo, P. and Giannoni, M.: Multi-wavelength optical determination
594 of black and brown carbon in atmospheric aerosols. *Atmospheric Environment*, 108,1-12,
595 <https://doi.org/10.1016/j.atmosenv.2015.02.058>, 2015.

596 Moosmüller, H., Chakrabarty, R. K., Ehlers, K. M., & Arnott, W. P.: Absorption Ångström
597 coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical particles.
598 *Atmospheric Chemistry and Physics*, 11(3), 1217-1225,
599 <https://doi.org/10.1021/acs.estlett.8b00118>, 2011a.

600 Moteki, N.: Climate-relevant properties of black carbon aerosols revealed by in situ
601 measurements: a review. *Progress in Earth and Planetary Science*, 10(1), pp.1-16,
602 <https://doi.org/10.1186/s40645-023-00544-4>, 2023.

603 Ohata, S., Moteki, N., Mori, T., Koike, M. and Kondo, Y.: A key process controlling the wet
604 removal of aerosols: new observational evidence. *Scientific reports*, 6(1), p.34113,
605 <https://doi.org/10.1038/srep34113>, 2016.

606 Osborne, S. R., Johnson, B. T., Haywood, J. M., Baran, A. J., Harrison, M. A. J., & McConnell,
607 C. L.: Physical and optical properties of mineral dust aerosol during the Dust and Biomass-
608 burning Experiment. *Journal of Geophysical Research: Atmospheres*, 113(D23),
609 <https://doi.org/10.1029/2007JD009551>, 2008.

610 Park, RJ, Kim, MJ, Jeong, JI, Youn, D., & Kim, S.: A contribution of brown carbon aerosol to
611 the aerosol light absorption and its radiative forcing in East Asia. *Atmospheric Environment*,
612 44 (11), 1414-1421, <https://doi.org/10.1016/j.atmosenv.2010.01.042> , 2010.

613 Pearson, K.: Determination of the coefficient of correlation. *Science*, 30(757), 23-25,
614 [DOI:10.1126/science.30.757.23](https://doi.org/10.1126/science.30.757.23), 1909.

615 Pierrehumbert, R. T.: Short-lived climate pollution. *Annual Review of Earth and Planetary*
616 *Sciences*, 42, 341-379, [DOI: 10.1146/annurev-earth-060313-054843](https://doi.org/10.1146/annurev-earth-060313-054843), 2014.

617 Prabhu, V., Soni, A., Madhwal, S., Gupta, A., Sundriyal, S., Shridhar, V., Sreekanth, V. and
618 Mahapatra, P.S.: Black carbon and biomass burning associated high pollution episodes
619 observed at Doon valley in the foothills of the Himalayas. *Atmospheric Research*, 243,
620 p.105001, <https://doi.org/10.1016/j.atmosres.2020.105001>, 2020.

621 Rahman, H., Karuppaiyan, R., Senapati, P.C., Ngachan, S.V. and Kumar, A.: An analysis of
622 past three decade weather phenomenon in the mid-hills of Sikkim and strategies for mitigating
623 possible impact of climate change on agriculture. *Climate change in Sikkim: Patterns, impacts*

624 and initiatives, pp.1-18, [http://sikkimforest.gov.in/climate-change-in-sikkim/2-chapter-](http://sikkimforest.gov.in/climate-change-in-sikkim/2-chapter-An%20analysis%20of%20past%20three%20decade%20weather.pdf)
625 [An%20analysis%20of%20past%20three%20decade%20weather.pdf](http://sikkimforest.gov.in/climate-change-in-sikkim/2-chapter-An%20analysis%20of%20past%20three%20decade%20weather.pdf) , 2012.

626 Ramachandran, S., & Rupakheti, M.: Trends in the types and absorption characteristics of
627 ambient aerosols over the Indo-Gangetic Plain and North China Plain in last two decades.
628 Science of The Total Environment, 831, 154867,
629 <https://doi.org/10.1016/j.scitotenv.2022.154867>, 2022.

630 Ramachandran, S., Rupakheti, M., & Lawrence, M. G.: Black carbon dominates the aerosol
631 absorption over the Indo-Gangetic Plain and the Himalayan foothills. Environment
632 international, 142, 105814, <https://doi.org/10.1016/j.envint.2020.105814>, 2020.

633 Ramanathan, V., & Carmichael, G.: Global and regional climate changes due to black carbon.
634 Nature geoscience, 1(4), 221-227, <https://doi.org/10.1038/ngeo156>, 2008.

635 Rana, A., Rawat, P. and Sarkar, S.: Sources, transport pathways and radiative effects of BC
636 aerosol during 2018–2020 at a receptor site in the eastern Indo-Gangetic Plain. Atmospheric
637 Environment, p.119900, <https://doi.org/10.1016/j.atmosenv.2023.119900>, 2023.

638 Rathod, T. D., & Sahu, S. K.: Measurements of optical properties of black and brown carbon
639 using multi-wavelength absorption technique at Mumbai, India. Journal of Earth System
640 Science, 131(1), 32, <https://doi.org/10.1007/s12040-021-01774-0>, 2022.

641 Rathod, T., Sahu, S. K., Tiwari, M., Yousaf, A., Bhangare, R. C., & Pandit, G. G.: Light
642 absorbing properties of brown carbon generated from pyrolytic combustion of household
643 biofuels. Aerosol and Air Quality Research, 17(1), 108-116,
644 <https://doi.org/10.4209/aaqr.2015.11.0639>, 2017.

645 Reddy, M. S., & Venkataraman, C.: Inventory of aerosol and sulphur dioxide emissions from
646 India: I—Fossil fuel combustion. Atmospheric Environment, 36(4), 677-697,
647 [https://doi.org/10.1016/S1352-2310\(01\)00463-0](https://doi.org/10.1016/S1352-2310(01)00463-0), 2002a.

648 Reddy, M. S., & Venkataraman, C.: Inventory of aerosol and sulphur dioxide emissions from
649 India. Part II—biomass combustion. Atmospheric Environment, 36(4), 699-712,
650 [https://doi.org/10.1016/S1352-2310\(01\)00464-2](https://doi.org/10.1016/S1352-2310(01)00464-2), 2002b.

651 Runa, F., Islam, M., Jeba, F., & Salam, A.: Light absorption properties of brown carbon from
652 biomass burning emissions. Environmental Science and Pollution Research, 29(14), 21012-
653 21022, <https://doi.org/10.1007/s11356-021-17220-z>, 2022.

654 Sankar, T.K., Ambade, B., Mahato, D.K., Kumar, A. and Jangde, R.: Anthropogenic fine
655 aerosol and black carbon distribution over urban environment. Journal of Umm Al-Qura
656 University for Applied Sciences, pp.1-10, <https://doi.org/10.1007/s43994-023-00055-4>, 2023.

657 Sarkar, A.: A generalized relationship between atmospheric pressure and precipitation
658 associated with a passing weather system. MAUSAM, 69(1), pp.133-140, DOI:
659 10.54302/mausam.v69i1.242, 2018.

660 Sharma, K., Ranjan, R.K., Lohar, S., Sharma, J., Rajak, R., Gupta, A., Prakash, A. and Pandey,
661 A.K.: Black Carbon Concentration during Spring Season at High Altitude Urban Center in
662 Eastern Himalayan Region of India. Asian Journal of Atmospheric Environment (AJAE),
663 16(1), <https://doi.org/10.5572/ajae.2021.149>, 2022.

664 Shindell, D., Kuylenstierna, J.C., Vignati, E., van Dingenen, R., Amann, M., Klimont, Z.,
665 Anenberg, S.C., Muller, N., Janssens-Maenhout, G., Raes, F. and Schwartz, J.: Simultaneously
666 mitigating near-term climate change and improving human health and food security. *Science*,
667 335(6065), 183-189, DOI: [10.1126/science.1210026](https://doi.org/10.1126/science.1210026), 2012.

668 Shaddick, G., Thomas, M.L., Mudu, P., Ruggeri, G. and Gumy, S.: Half the world's population
669 are exposed to increasing air pollution. *NPJ Climate and Atmospheric Science*, 3(1), p.23,
670 <https://doi.org/10.1038/s41612-020-0124-2>, 2020.

671 Shukla, K. K., Sarangi, C., Attada, R., & Kumar, P.: Characteristic dissimilarities during high
672 aerosol loading days between western and eastern Indo-Gangetic Plain. *Atmospheric*
673 *Environment*, 269, 118837, <https://doi.org/10.1016/j.atmosenv.2021.118837>, 2022.

674 Sloss, L.: Black carbon emissions in India. CCC/209. IEA Clean Coal Centre, London, 38,
675 (2012).

676 Stevens, B., & Feingold, G.: Untangling aerosol effects on clouds and precipitation in a
677 buffered system. *Nature*, 461(7264), 607-613, <https://doi.org/10.1038/nature08281>, 2009.

678 Stjern, C.W., Forster, P.M., Jia, H., Jouan, C., Kasoar, M.R., Myhre, G., Olivié, D., Quaas, J.,
679 Samset, B.H., Sand, M. and Takemura, T.: The Time Scales of Climate Responses to Carbon
680 Dioxide and Aerosols. *Journal of Climate*, 36(11), pp.3537-3551,
681 <https://doi.org/10.1175/JCLI-D-22-0513.1>, 2023.

682 Sun, Y., Hao, Q., Cui, C., Shan, Y., Zhao, W., Wang, D., Zhang, Z. and Guan, D.: Emission
683 accounting and drivers in East African countries. *Applied Energy*, 312, p.118805,
684 <https://doi.org/10.1016/j.apenergy.2022.118805>, 2022.

685 Takemura, T., & Suzuki, K.: Weak global warming mitigation by reducing black carbon
686 emissions. *Scientific reports*, 9(1), 1-6, <https://doi.org/10.1038/s41598-019-41181-6>, 2019.

687 Venkataraman, C., Habib, G., Kadamba, D., Shrivastava, M., Leon, J.F., Crouzille, B.,
688 Boucher, O. and Streets, D.G.: Emissions from open biomass burning in India: Integrating the
689 inventory approach with high-resolution Moderate Resolution Imaging Spectroradiometer
690 (MODIS) active-fire and land cover data. *Global biogeochemical cycles*, 20(2),
691 <https://doi.org/10.1029/2005GB002547>, 2006.

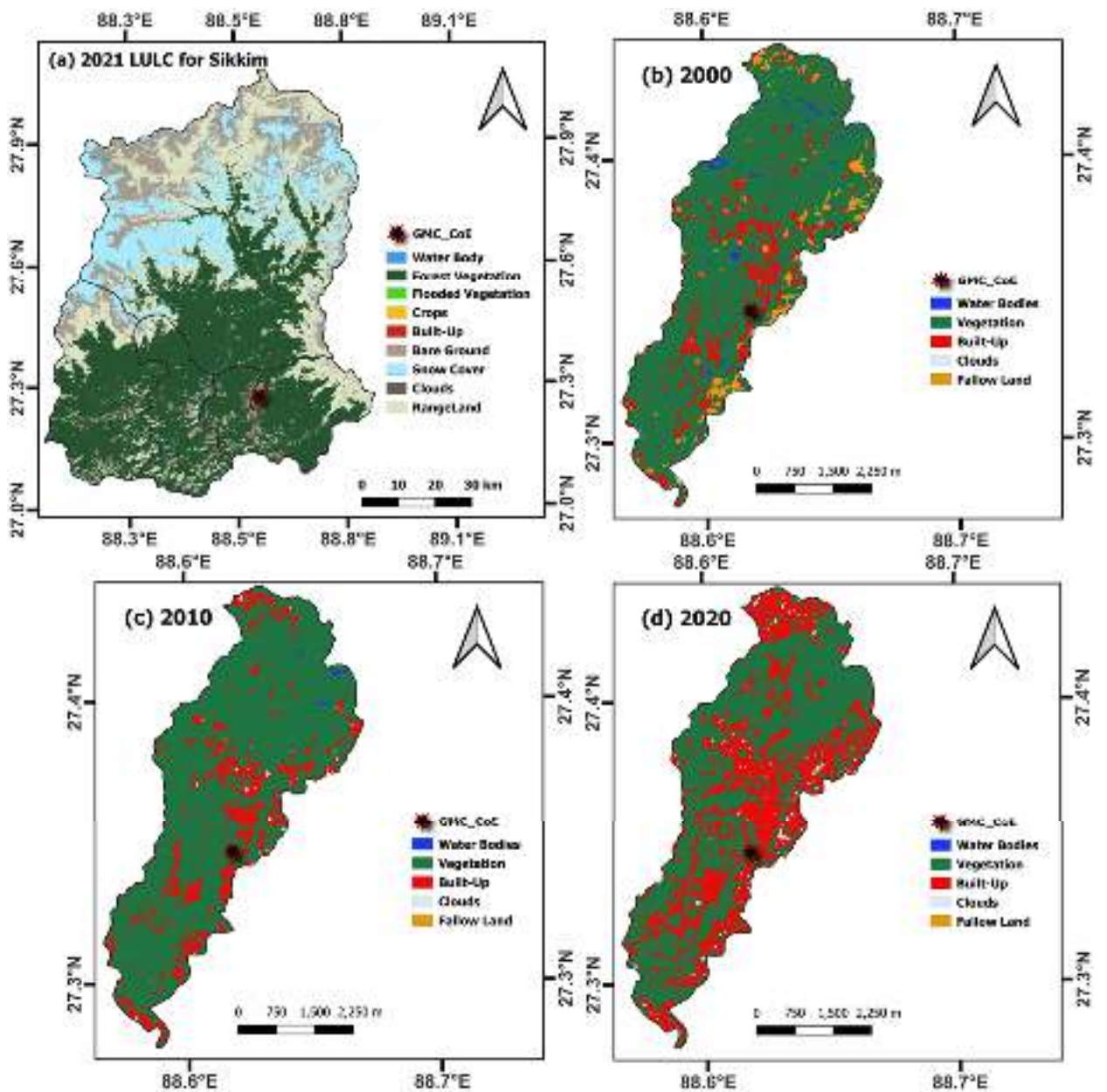
692 Watham, T., Padalia, H., Srinet, R., Nandy, S., Verma, P. A., & Chauhan, P.: Seasonal
693 dynamics and impact factors of atmospheric CO₂ concentration over subtropical forest
694 canopies: observation from eddy covariance tower and OCO-2 satellite in Northwest
695 Himalaya, India. *Environmental Monitoring and Assessment*, 193(2), 1-15,
696 <https://doi.org/10.1007/s10661-021-08896-4>, 2021.

697 Wang, Q., Liu, H., Ye, J., Tian, J., Zhang, T., Zhang, Y., Liu, S. and Cao, J.: Estimating
698 Absorption Ångström Exponent of Black Carbon Aerosol by Coupling Multiwavelength
699 Absorption with Chemical Composition. *Environmental Science & Technology Letters*, 8(2),
700 pp.121-127, <https://doi.org/10.1021/acs.estlett.0c00829>, 2020.

701 Wang, L., Jin, W., Sun, J., Zhi, G., Li, Z., Zhang, Y., Guo, S., He, J. and Zhao, C.: Seasonal
702 features of brown carbon in northern China: Implications for BrC emission control.
703 *Atmospheric Research*, 257, p.105610, <https://doi.org/10.1016/j.atmosres.2021.105610>,
704 2021.

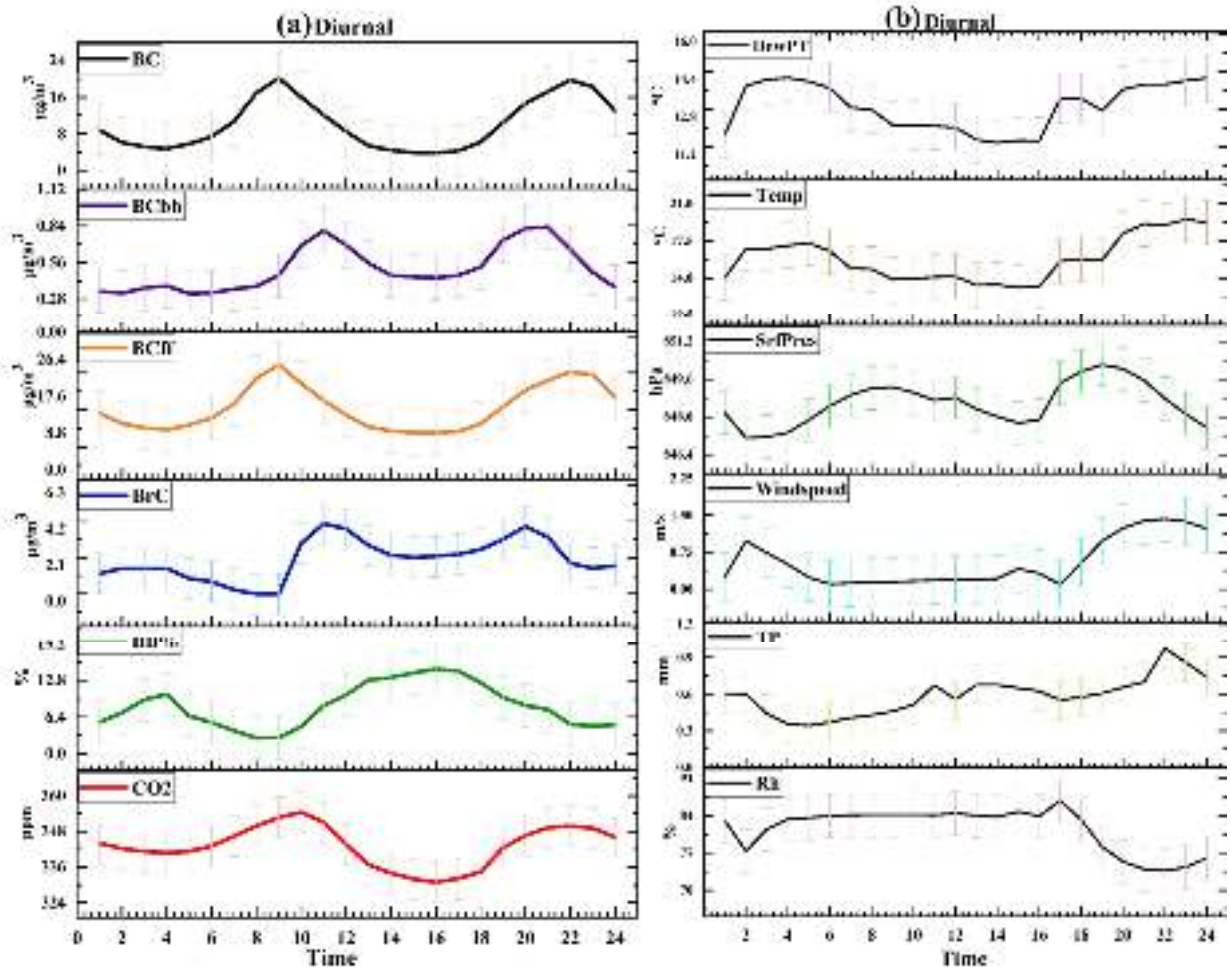
- 705 Wu, Y., Wang, Y., Zhou, Y., Liu, X., Tang, Y., Wang, Y., Zhang, R. and Li, Z.: The wet
706 scavenging of air pollutants through artificial precipitation enhancement: A case study in the
707 Yangtze River Delta. *Frontiers in Environmental Science*, 10, p.1027902,
708 <https://doi.org/10.3389/fenvs.2022.1027902>, 2022.
- 709 Xiao-lei, C. H. U., L. U. Zhong, W. E. I. Dan, and L. E. I. Guo-ping.: Effects of land use/cover
710 change on temporal and spatial variability of precipitation and temperature in the Songnen
711 Plain of China. *Journal of Integrative Agriculture* 21, no. 1: 235, doi: 10.1016/S2095-
712 3119(20)63495-5, 2022.
- 713 Yasunari, T., Bonasoni, P., Laj, P., Fujita, K., Vuillermoz, E., Marinoni, A., Cristofanelli, P.,
714 Duchi, R., Tartari, G. and Lau, K.M.: Estimated impact of black carbon deposition during pre-
715 monsoon season from Nepal Climate Observatory–Pyramid data and snow albedo changes
716 over Himalayan glaciers. *Atmospheric Chemistry and Physics*, 10(14), 6603-6615,
717 <https://doi.org/10.5194/acp-10-6603-2010>, 2010.
- 718 Yoo, J.M., Lee, Y.R., Kim, D., Jeong, M.J., Stockwell, W.R., Kundu, P.K., Oh, S.M., Shin,
719 D.B. and Lee, S.J.: New indices for wet scavenging of air pollutants (O₃, CO, NO₂, SO₂, and
720 PM₁₀) by summertime rain. *Atmospheric Environment*, 82, pp.226-237,
721 <https://doi.org/10.1016/j.atmosenv.2013.10.022>, 2014.
- 722 Yue, S., Zhu, J., Chen, S., Xie, Q., Li, W., Li, L., Ren, H., Su, S., Li, P., Ma, H. and Fan, Y.:
723 Brown carbon from biomass burning imposes strong circum-Arctic warming. *One Earth*, 5(3),
724 293-304, <https://doi.org/10.1016/j.oneear.2022.02.006>, 2022.
- 725 Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z.,
726 Zhao, Y., and Shen, Z.: Chemical characterization and source apportionment of PM_{2.5} in
727 Beijing: seasonal perspective, *Atmos. Chem. Phys.*, 13, 7053–7074,
728 <https://doi.org/10.5194/acp-13-7053-2013>, 2013.
- 729 Zhang, A., Wang, Y., Zhang, Y., Weber, R.J., Song, Y., Ke, Z. and Zou, Y.: Modeling the
730 global radiative effect of brown carbon: a potentially larger heating source in the tropical free
731 troposphere than black carbon. *Atmospheric Chemistry and Physics*, 20(4), pp.1901-1920,
732 <https://doi.org/10.5194/acp-20-1901-2020>, 2020.
- 733 Zhu, C.S., Qu, Y., Huang, H., Chen, J., Dai, W.T., Huang, R.J. and Cao, J.J.: Black carbon and
734 secondary brown carbon, the dominant light absorption and direct radiative forcing
735 contributors of the atmospheric aerosols over the Tibetan Plateau. *Geophysical research letters*,
736 48(11), p.e2021GL092524, <https://doi.org/10.1029/2021GL092524>, 2021.

List of Figures



738

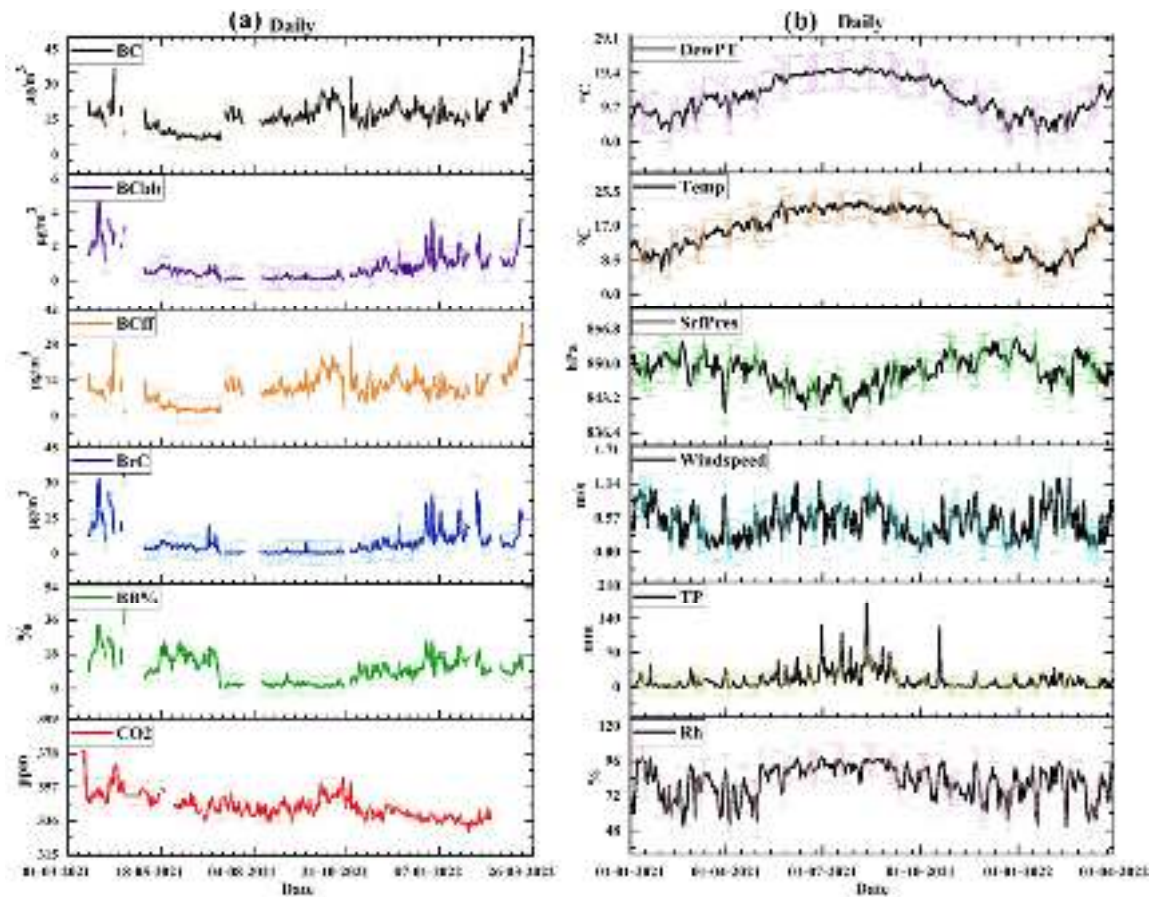
739 Figure 1. The study location and land use land cover for 2000, 2010, 2020, and 2021 for
740 December over Gangtok and Sikkim region using Landsat-5, Landsat-8, and Sentinel-2 data
741 sets.



742

743 Figure 2. (a) The hourly observation of Black Carbon, Black Carbon through biomass burning,
 744 Black Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon
 745 Dioxide (BC, BC_{bb}, BC_{ff}, BrC, BB%, and CO₂, respectively) (The corresponding unit for BC,
 746 BC_{bb}, BC_{ff}, BrC: $\mu\text{g}/\text{m}^3$; BB%: % and CO₂: ppm) for 16th March 2021 to 10th March 2022 over
 747 study location (lat:27.32; lon:88.61). The light colour shading refers to $\pm\sigma$ standard deviation
 748 for each variable. (b) Same as Figure 2a, but for meteorological parameters such as dewpoint
 749 temperature (DewPT), temperature (Temp), surface pressure (SrfPres), windspeed, total
 750 precipitation (TP), and relative humidity (Rh) from 16th March 2021 to 10th March 2022.

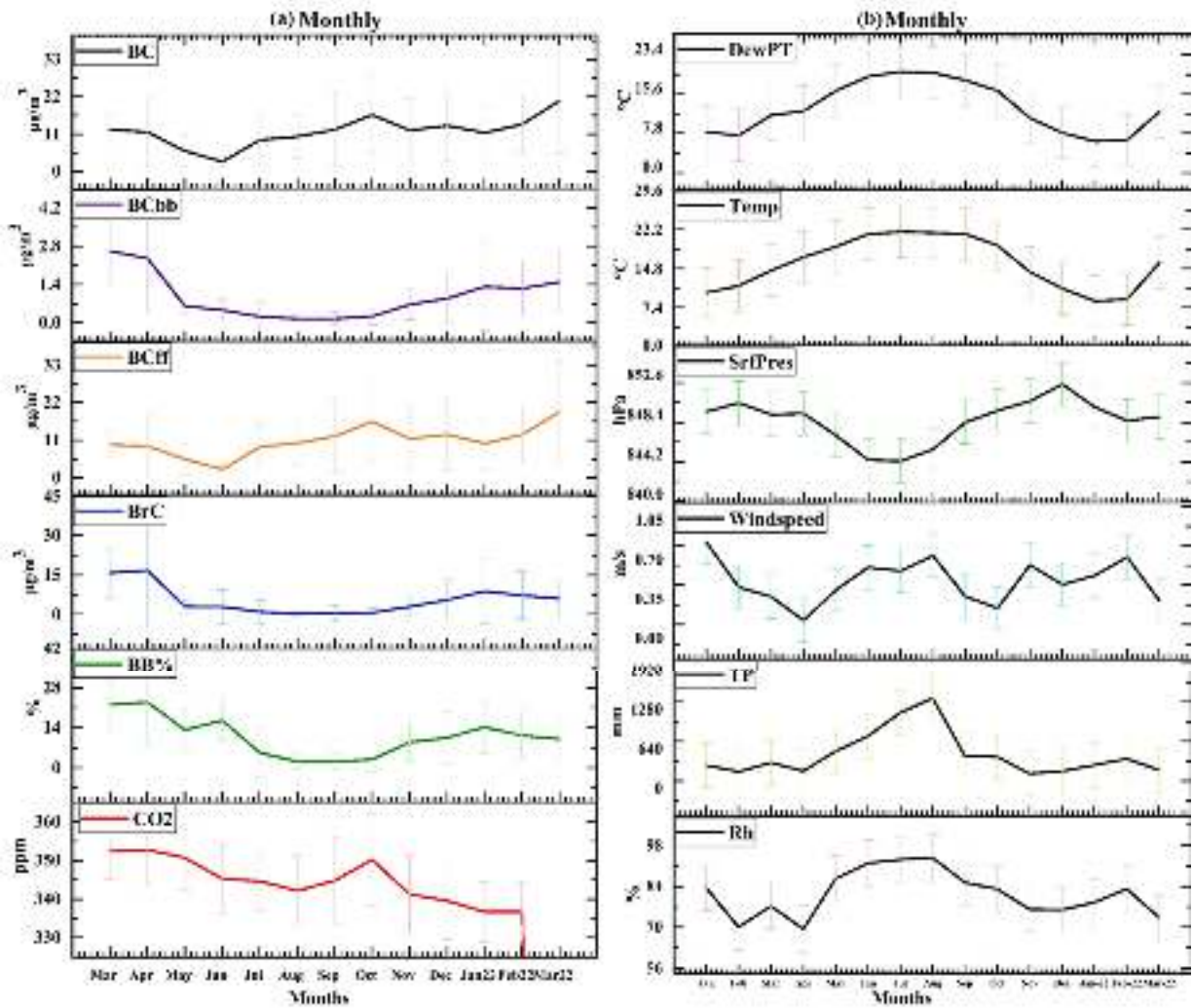
751



752

753 Figure 3. (a) The daily mean of Black Carbon, Black Carbon through biomass burning, Black
 754 Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon Dioxide
 755 (BC, BC_{bb}, BC_{ff}, BrC, BB%, and CO₂, respectively) (The corresponding unit for BC, BC_{bb},
 756 BC_{ff}, BrC: $\mu\text{g}/\text{m}^3$; BB%: % and CO₂: ppm) for 16th March 2021 to 10th March 2022 over study
 757 location (lat:27.32; lon:88.61). The light colour shading refers to $\pm\sigma$ standard deviation for
 758 each variable. (b) same as Figure 3a, but for meteorological parameters such as dewpoint
 759 temperature (DewPT), temperature (Temp), surface pressure (SrfPres), Windspeed, total
 760 precipitation (TP), and relative humidity (Rh) from 1st January 2021 to 31st March 2022.

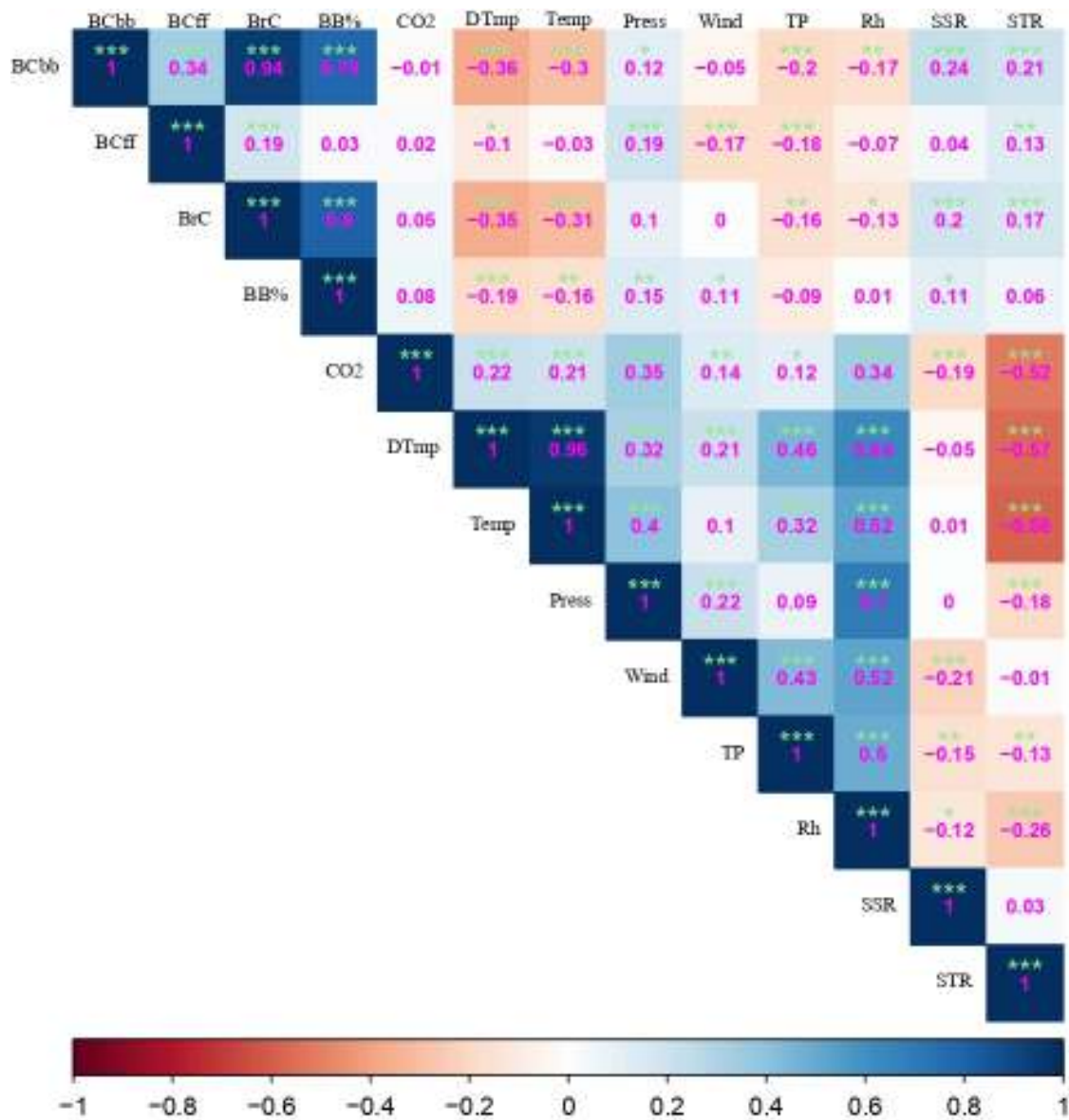
761



762

763 Figure 4. (a) The monthly mean of Black Carbon, Black Carbon through biomass burning,
 764 Black Carbon through fossil fuel, Brown Carbon, Biomass Burning percentage and Carbon
 765 Dioxide (BC, BC_{bb}, BC_{ff}, BrC, BB%, and CO₂, respectively) (The corresponding unit for BC,
 766 BC_{bb}, BC_{ff}, BrC: $\mu\text{g}/\text{m}^3$; BB%: % and CO₂: ppm) for 16th March 2021 to 10th March 2022 over
 767 study location (lat:27.32; lon:88.61). The error bar shows $\pm\sigma$ standard deviation for each
 768 variable. (b) Same as Figure 4a, but for meteorological parameters such as dewpoint
 769 temperature (DewPT), temperature (Temp), surface pressure (SrfPres), windspeed, total
 770 precipitation (TP), and relative humidity (Rh) during January 2021 to March 2022.

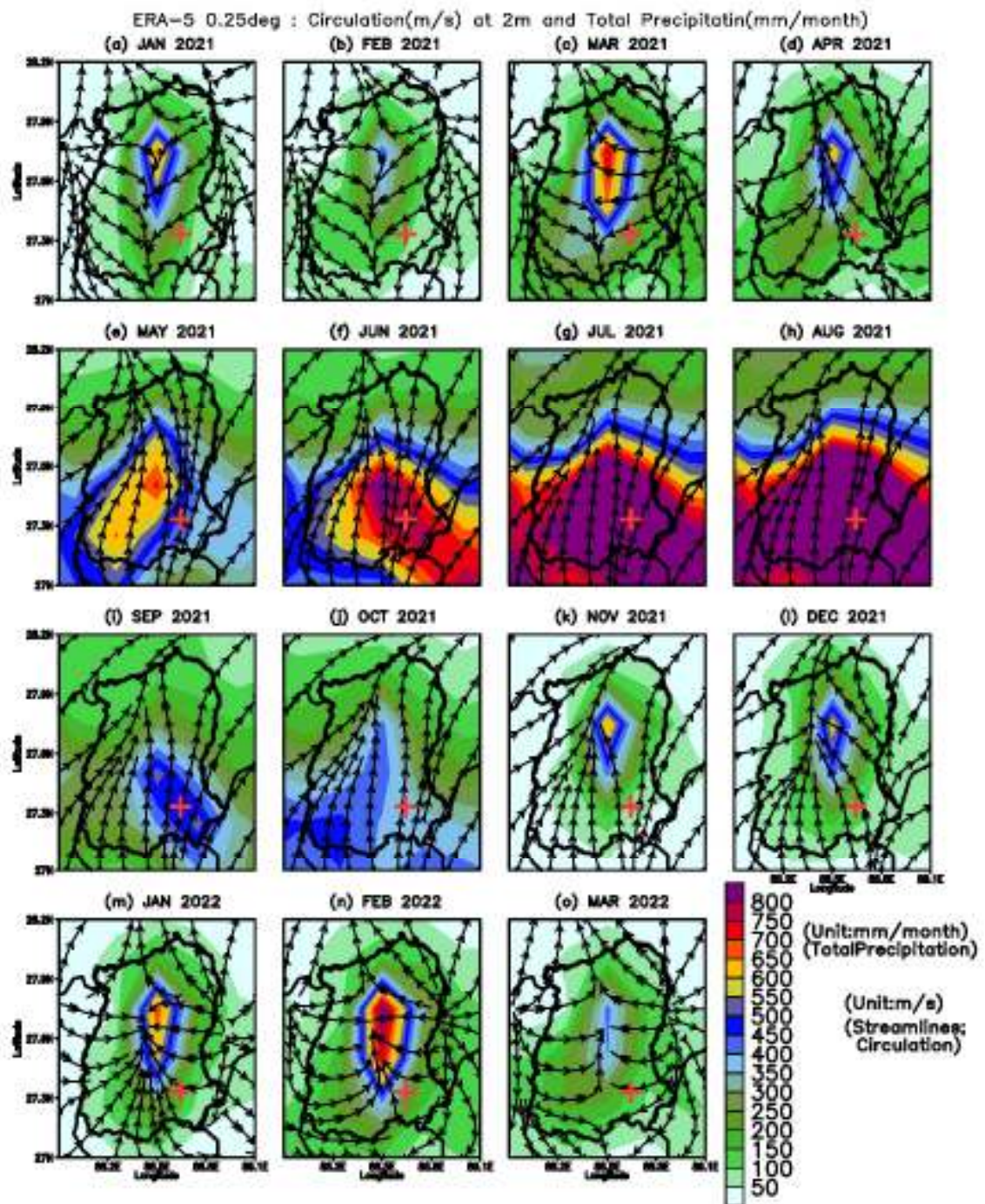
771



772

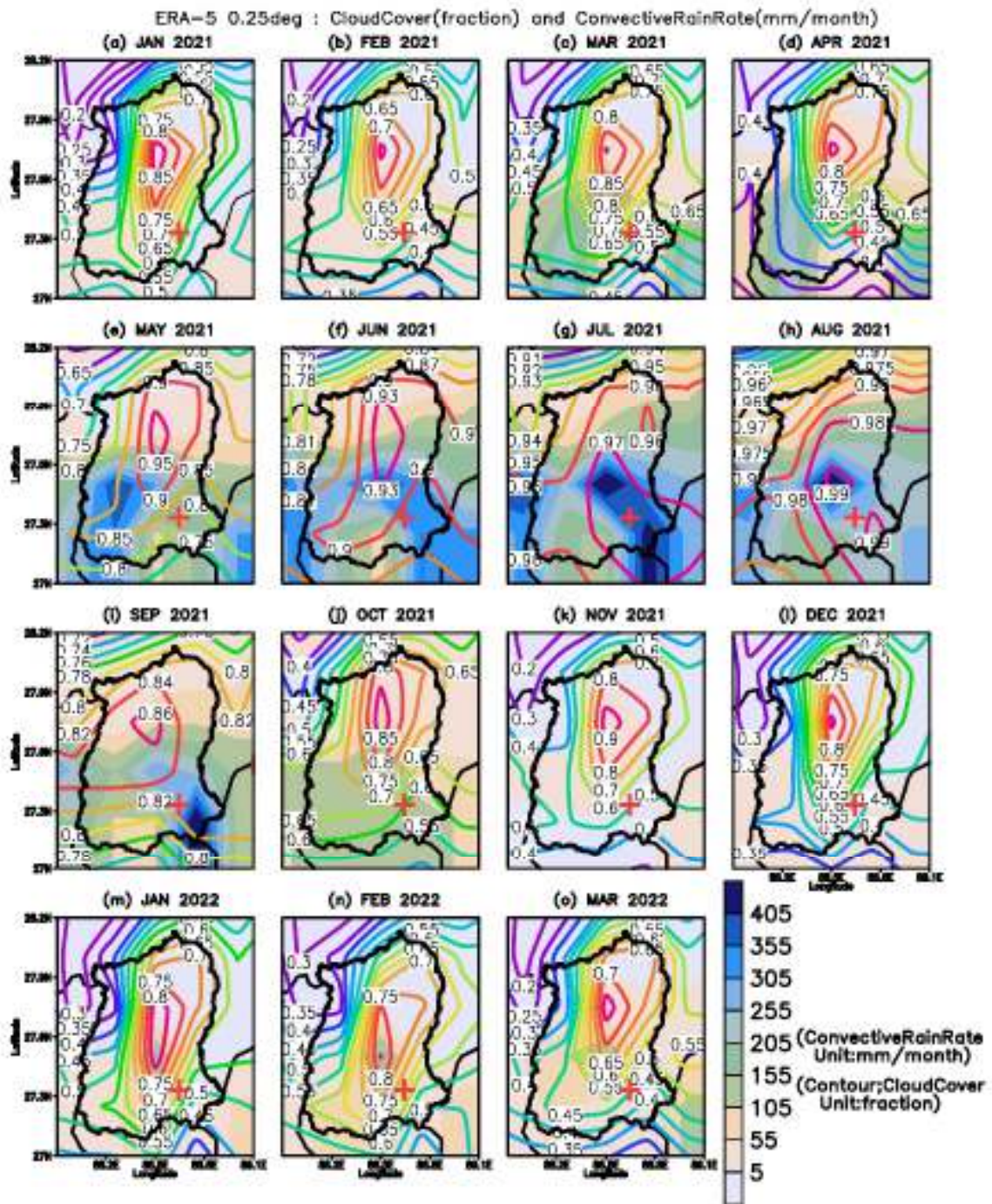
773 Figure 5. Correlation among BC, BC_{bb}, BC_{ff}, BrC, BB%, CO₂ and, dewpoint temperature
 774 (DTmp), temperature (Temp), surface pressure (Press), Wind, total precipitation (TP), Relative
 775 humidity (Rh), net solar radiation (SSR), and net thermal radiation (STR). The (***) shows
 776 99% significance, (**) shows 95% significance, (*) 90% significance, and () shows no
 777 significance. The correlation coefficient values (-0.3 to -0.49) or (0.3 to 0.49) are considered
 778 ‘a good correlation’, and values \leq (-0.5) or \geq (0.5) are considered “a strong correlation”.

779



780 Figure 6. Monthly total precipitation (cumulative) and wind circulation pattern during January
 781 2021 to March 2022. The shading shows precipitation patterns, and the streamline shows wind
 782 circulation. The (+) mark is a representation of the sampling location.

783



784 Figure 7. Monthly convective rain and total cloud cover during January 2021 to March 2022.
 785 The shading shows a convective rain pattern, and the contour shows a total cloud cover
 786 fraction. The (+) mark is a representation of the sampling location.

787

789 Table 1. The details of datasets used for the present study.

790

Variables	Data sets	Years (Span)	Resolution		Source	Reference
			Temporal	Horizontal		
Black and Brown Carbon	Observation and analysis, data generated using Aethalometer AE33	March 2021-March 2022	Weekly	Point Location (Gangtok)	Original data generated	Present Study
Total precipitation	ERA5 (ECMWF)	2021 to 2022	Hourly	0.25° * 0.25°	ECMWF https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-single-levels?tab=form	Hersbach et al., 2020
Relative humidity						
Temperature (2 meter)						
Wind (surface wind)						
Surface pressure						
Dewpoint temperature						
Net solar, and thermal radiation downward						
LULC	LandSat-5, LandSat-8 and earth explorer USGS	December 2000, December 2010, December 2020	2000, 2010, 2020	30m, 30m	earth explorer USGS. https://earthexplorer.usgs.gov/	earth explorer USGS.
LULC	Sentinel-2 Esri Inc.	December 2021	2021	10 m	Esri Inc. https://www.arcgis.com/home/item.html?id=d3da5dd386d140cf93fc9ecbf8da5e31	Karra et al., 2021

791