

DUACS DT2021 reprocessed altimetry improves sea level retrieval in the coastal band of the European Seas

Antonio Sánchez-Román¹, M. Isabelle Pujol², Yannice Faugère², Ananda Pascual¹

¹Instituto Mediterráneo de Estudios Avanzados, C/Miquel Marqués, 21, 07190 Esporles, Spain

²Collecte Localisation Satellites, Parc Technologique du Canal, 8-10 rue Hermès, 31520 Ramonville-Saint-Agne, France

Correspondence to: Antonio Sánchez Román (asanchez@imedea.uib-csic.es)

Abstract. More than 29 years of altimeter data have been recently reprocessed by the multi satellite Data Unification and Altimeter Combination System (DUACS) and made available under the name of DT2021 processing through the Copernicus Marine Service (CMEMS) and the Copernicus Climate Change (C3S) Service. New standards have been applied and various geophysical correction parameters have been updated compared to the previous release in order to improve the product quality. This paper describes the assessment of this new release through the comparison of both *all satellites* and *two satellites* products with external in situ tide gauge measurements in the coastal areas of the European Seas for a time period spanning from 1 January 1993 to 31 May 2020. The aim is to quantify the improvements on the previous DT2018 processing version on the retrieval of sea level in the coastal zone.

The results confirmed that the CMEMS product in the new DT2021 processing version better solves the signal in the coastal band. The *all satellites* dataset showed a reduction of 3% in errors when compared with tide gauges and of 5% in the variance of the differences between the datasets compared to DT2018 reprocessing. Moreover, the *all satellites* dataset provided more accurate sea level measurements when comparing with tide gauges respect to the climatic *two satellites* dataset due to the better performance of the former for the assessment of higher than climatic frequency signals. On the contrary, the *two satellite* dataset is the most suitable product for the assessment of long-term sea level SSH trends in the coastal zone due to its larger stability to the detriment of the *all satellites* dataset.

1 Introduction

On December 2021, more than 29 years of Level 3 (L3) and Level 4 (L4) altimetry products were reprocessed, released and made freely available for users as the “DT2021” version (CMEMS-SL-QUID, 2022; C3S-PUG, 2022; Faugère et al., 2022) of the multi-satellite Data Unification and Altimeter Combination System (DUACS) products by the European Copernicus Program (<http://marine.copernicus.eu/>) substituting the former “DT2018” product version (Taburet et al., 2019) which is no longer available in the Copernicus Catalogue. Currently, two types of altimetric L4 gridded products generated by the DUACS production system are available: the so called *all satellites* global and regional (European Seas) gridded products disseminated via the Copernicus Marine Service (CMEMS) project (CMEMS-SL-QUID, 2022); and the *two satellites* global gridded products distributed via the Copernicus Climate Change Service (C3S) project (C3S-PUG, 2022). Currently, the *two satellites* products are also distributed via the CMEMS project. The *all satellites* products are dedicated to the retrieval of mesoscale signals on a global or regional scale whereas the *two satellites* ones are dedicated to monitoring the long-term evolution of sea level, thus being suitable for using in climate applications (Taburet et al., 2019).

The Level 2 altimeter standards used to compute sea level anomaly (SLA) in the CMEMS and C3S products are identical (CMEMS-SL-QUID, 2022), but the reference used to compute SLAs differs: CMEMS products use a mean profile of sea surface heights along the theoretical track of the satellites with a repetitive orbit, whilst C3S products use a mean sea surface (MSS) for all missions. In the latest release, new up-to-date standards have been applied and various geophysical correction

39 parameters have been updated compared to the previous DT2018 version (Table A1 in Appendix A). This provides both an
40 improved accuracy of SLA and lower regional sea level biases.

41 Namely, (i) a **new internal tide correction** that allows the prediction of the two main tidal constituents of both diurnal and
42 semidiurnal tidal frequencies has been applied. The solution proposed by Zaron (2019) is used (HRET 8.1 version). This
43 correction reduces the coherent signal characteristic of internal tide and provides a more precise reconstruction of mesoscale
44 eddies. The use of the internal tide correction induces a reduction of internal tide signature on along-track data improving the
45 precision of the resulting L4 gridded product (CMEMS-SL-QUID, 2022).

46 (ii) a **new MSS** for non-repetitive missions and recent missions consisting in a hybrid gridded MSS field made up of three
47 different gridded MSS models is used. Namely, SIO MSS model (Sandwell et al., 2017) is used in open ocean, CNES_CLS-
48 2015 model (Pujol et al., 2018) is used in coastal areas (distance to the coast lower than 20 km) and DTU15 model (Andersen
49 et al. 2016) is used in the Arctic region (latitude larger than 80 northern degrees). This hybrid solution contributes to reduce
50 the SLA errors at short wavelengths. A **new mean profile** (precise MSS along the altimeter tracks) is used for historical
51 repetitive missions (CMEMS products). New mean profiles were estimated along the historical repetitive tracks of
52 Topex/Poseidon/Jason, Topex/Poseidon/Jason-interleaved phase, ERS/Envisat/AltiKa, Sentinel-3A and GFO in consistency
53 with the different standards used in DT2021 version. This improves the SLA signal at long wavelengths.

54 (iii) a **new Mean Dynamic Topography** for the Global (Mulet et al., 2021), and the Mediterranean and Black Seas is applied
55 (Jousset et al., 2020,2022).

56 (iv) an **improved Long Wavelength Error (LWE) correction**, delivered in L3 product, has been computed as the final step
57 of multi-mission cross-calibration processing. Progress with respect to the previous version has been done by first estimating
58 the LWE correction with higher frequency along the different tracks (100 Km instead of 500 km used previously), then by
59 improving the interpolation methodology (Optimal Interpolation instead of Spline used previously) to retrieve the correction
60 on each along-track position. It is expected to remove local SLA residual biases between neighbouring altimeter tracks.

61 (v) and finally, the DT2021 products version includes an **upgraded mapping parameterisation** that contributes to improve
62 the mesoscale signal visible on L4 products. Namely, the spatial and temporal correlation scales are optimised improving the
63 reconstruction of the mesoscale signal, a more precise definition of the observation's errors computed with regard to the new
64 altimeter standards is provided, and finally, a more precise estimation and correction of LWE in the mapping process is applied
65 removing local SLA residual biases. A complete description of the different evolutions implemented in the DUACS DT2021
66 products version can be found in CMEMS-SL-QUID (2022).

67 The validation (quality check) of altimetry products is a key step in the data processing pipe to assess and characterise the
68 errors associated with the altimetry measurements. This issue is crucial in the coastal zone, where traditional altimetry have
69 been often unable to produce meaningful signals of sea level change due to the typically shallower water, bathymetric
70 gradients, and shoreline shapes, among others (Vignudelli et al., 2019; Sánchez-Román et al., 2020). Actually, global and
71 regional products from DT2021 and DT2018 reprocessings are not optimised for the coastal band promoting larger errors in
72 the retrieval of sea level with regard to the open ocean.

73 Nevertheless, the monitoring of sea level changes in coastal areas is an important societal issue (Pujol et al., 2023). Thus, most
74 of the efforts of the international community in the recent past have been focused on the research and development of
75 techniques for coastal altimetry, with substantial support from space agencies such as the European Space Agency (ESA), the
76 Centre National d'Études Spatiales (CNES), and other research institutions (Cipollini et al., 2017). Efforts of the coastal
77 altimetry community are aimed at extending the capabilities of current altimeters closer to the coastal zone. This includes the
78 application of improved geophysical corrections, data recovery strategies near the coast using new editing criteria, and high-
79 frequency along-track sampling associated with updated quality control procedures (Vignudelli et al., 2019). As a result,
80 regional altimeter products such as PISTACH (Mercier et al., 2010), X-TRACK (Roblou et al., 2011; Birol et al., 2017); X-

81 TRACK-ALES (Birol et al., 2021) and ESA EO4SIBS (Grégoire, 2021) focused on the coastal zone have been developed over
82 the last few years (Pujol et al., 2023). These products are disseminated to both the international scientific community and
83 society through regular specific coastal altimetry workshops.

84 Different metrics are used to assess the quality of altimetry data. They mainly consist in the analysis of the SLA field at
85 different steps of the processing; check consistency of the SLA along the tracks of different altimeters and between gridded
86 and along-track products; and comparisons with external in situ measurements (CMEMS-SL-QUID, 2022). In situ and
87 altimetric observations are complementary and are often assumed to observe the same signals (Wöppelmann and Marcos,
88 2016). In coastal areas, tide gauge measurements are commonly used. In Taburet et al. (2019), DUACS DT2018 L4 global
89 gridded products were assessed in the coastal areas through a comparison with monthly tide gauge measurements from the
90 Permanent Service for Mean Sea Level (PSMSL) Network (PSMSL, 2016). These authors reported a global reduction of 0.6%
91 in variance with respect to the previous processing (DUACS DT2014 dataset). Pascual et al. (2006, 2009) investigated the
92 consistency between previous versions of the altimeter L4 gridded products and tide gauge data from the PSMSL repository
93 in the coastal zone reporting mean square differences between the two datasets ranging between 30% and 90% in the European
94 coasts. More recently, Sánchez-Román et al. (2020) assessed the quality of DUACS L3 products in the coastal band of the
95 European Seas through comparison with independent tide gauge measurements. These authors reported a mean root mean
96 square (rms) difference between both datasets lower than 7 cm for the whole region, with mean values ranging around less
97 than 4 cm in the Mediterranean basin and around 10 cm for the North West European Shelf (NWS) area (see Fig. 2 in Sánchez-
98 Román et al., 2020 for the location of this region). The quality of the DUACS DT2021 product version has been also assessed
99 through the comparison with monthly tide gauge measurements from the Global Sea Level Observing System
100 (GLOSS)/Climate Variability and Predictability (CLIVAR) network. CMEMS-SL-QUID (2022) reports improved results
101 when using the latest reprocessing with a reduction in variance of the differences between altimetry and tide gauges ranging
102 between 0.2% and more than 5% of the tide gauge signal in the European coasts; with respect to the previous product version.

103 This paper focuses on improvements of the latest reprocessing of DUACS Delayed Time (DT) reanalysis (referred hereinafter
104 as DT2021) in the retrieval of sea level in the coastal band of the European Seas with respect to the previously available
105 reprocessed products (referred hereafter as DT2018). To do that, an intercomparison of L4 global altimetry gridded products
106 and in situ tide gauges located along the European coasts from the Copernicus Catalogue is conducted. The performance of
107 the DT2021 processing *all satellites* and *two satellites* versions on the sea level retrieval is also assessed. The paper is organized
108 as follows: the SLA data used, the tide gauge dataset, and the method for comparing altimeter and in situ measurements are
109 detailed in section 2. Section 3 describes the performance of the DT2021 processing product version in the retrieval of sea
110 level in the coastal band. Also, the improvements over the previous DT2018 processing product version is assessed. Finally,
111 the discussion and main conclusions are included in section 4.

112 **2 Materials and methods**

113 **2.1 Sea level anomaly data**

114 The DUACS reprocessed L4 global satellite SLA maps used in this study correspond to both the DT2021 (CMEMS-SL-QUID,
115 2022; C3S-PUG, 2022; Faugère et al., 2022) and DT2018 product (Taburet et al., 2019) versions. SLA gridded products cover
116 the global ocean with a spatial and temporal resolution of $\frac{1}{4}$ of a degree and 1 day, respectively. Two different SLA datasets
117 for each one of the DUACS product versions are considered: the *all satellites* L4 global gridded product disseminated via the
118 CMEMS and the *two satellites* L4 global gridded product distributed via the C3S and CMEMS. The first one is computed with
119 a satellite constellation including all the available altimeters at a given time (ranging from 2 to 7 over the period considered in

120 this study, see e.g. Fig. 1 in Coastal Altimetry Team, 2021; Morrow et al., 2023). As a consequence, the errors are not constant
121 in time since they depend on the number of satellites used. This product focuses on the mesoscale mapping capacity of the
122 altimeter data together with the stability of the overall dataset. The *two satellites* SLA dataset is obtained by merging a steady
123 number of altimeters (two) in the satellite constellation. Two satellites is the minimum requirement to retrieve mesoscale
124 signals in delayed time conditions. (Pascual et al., 2006; Dibarboure et al., 2011). This fact also promotes nearly consistent
125 errors during the whole time period (some variation of the error can occur related to changes of the *two satellites* constellation).
126 This product focuses on the stability of the global mean sea level (MSL), even if this implies potential reduction of the spatial
127 sampling of the ocean. The reader is referred to Fig. 1 in Sánchez-Román et al. (2020) for more information about the DUACS
128 procedure flowchart applied to the altimetry data and also to the processing of the tide gauge data used to compare with
129 altimetry (next section). The time period investigated common to both DT2021 and DT2018 reprocessings spans from 1
130 January 1993 to 31 May 2020 due to the presently availability of DUACS DT2018 products. A complete description of the
131 SLA datasets can be found in CMEMS-SL-QUID (2022).

132 **2.2 Tide gauge observations**

133 The sea level records used to compare with satellite altimetry were extracted from the Copernicus Catalogue
134 (www.marineinsitu.eu). The tide gauge stations located in the European Seas' domain were initially considered for this study.
135 Following the methodology described in Sánchez-Román et al. (2020), the quality flags of the tide gauge records were checked
136 in order to remove observations with no quality check, potentially and bad data, and changes in the vertical reference of the
137 tide gauge. Also, observations with values larger than three times the standard deviation of the time series were rejected as
138 they could not be representative of ocean sea level changes but local features (e.g., river discharge, Laíz et al., 2013). The final
139 dataset consists of 213 tide gauge stations (Fig. 1) with time series exhibiting between 90% and 100% of valid data. The
140 stations and their information are listed in Table B1 in Appendix B.

141 Before they can be compared with altimeter data, tide gauge measurements have to be processed to remove oceanographic
142 signals whose temporal periods are not resolved by altimetry, thus avoiding important aliasing errors (Vignudelli et al., 2019).
143 We applied the methodology described in Sánchez-Román et al. (2020). In the following we summarise the corrections applied
144 to the tide gauge records:

- 145 • **Correction of oceanic tidal effects** by filtering tidal components (mainly diurnal and semidiurnal tidal constituents).
146 The u-tide software (Codiga, 2011) is used. The annual and semiannual frequencies, mainly driven by steric effect,
147 are kept in the tidal residuals since they are included in the altimetry data.
- 148 • **Removal of the atmospherically induced sea level** caused by the action of atmospheric pressure and wind
149 (Dorandeu and Le Traon, 1999; Carrère and Lyard, 2003). The same Dynamic Atmospheric Correction (DAC) as
150 for altimetry is applied for the sake of consistency. The 6 hourly fields of this correction, available at the Archiving,
151 Validation and Interpretation of Satellite Oceanographic Data (AVISO) website, are used. For each tide gauge site,
152 the nearest grid point was selected and used to remove the atmospherically induced sea level from observations,
153 previously converted into 6 hourly records (Marcos et al., 2015).
- 154 • **Correction of vertical movements** associated with glacial isostatic adjustment (GIA). GIA was considered as the
155 only source of vertical land motions. Its effects were removed from the SSH records, previously averaged into daily
156 data, by using the Peltier mantle viscosity model (VM2) (Peltier, 1998, 2004).

157 2.3 Method for comparing altimeter and in situ tide gauge records

158 The comparison method of altimetry with tide gauges consisted of collocating both datasets in time and space. As a first step,
159 a 15-day low-pass Loess filter was applied to altimetry and tide gauge time series to remove the high frequencies that cannot
160 be resolved by the altimetric data (Pascual et al., 2009; Ballarotta et al., 2019; Sánchez-Román et al., 2020). Then, the
161 correlations between each tide gauge record and SLA time series corresponding to grid points within a radius of 1 degree
162 around the tide gauge site were computed and the most correlated altimetry point was chosen. Only long-term monitoring
163 stations with a lifetime of more than three years were used in order to allow statistical significance. Statistical analyses were
164 performed using all available data pairs (altimetry-tide gauge). The collocated altimeter and tide gauge measurements were
165 analysed in terms of the rms difference and variance of the time series. In addition, the robustness of the results was investigated
166 according to Sánchez-Román et al. (2017, 2020) using a bootstrap method (Efron and Tibshirani, 1986), which allows us to
167 estimate quantities related to a dataset by averaging estimates from multiple data samples. To do that, the dataset is iteratively
168 resampled with replacement. A total of 1.000 iterations were used to ensure that meaningful statistics such as standard deviation
169 could be calculated on the sample of estimated values, thus allowing us to assign measures of accuracy to sample estimates
170 (Sánchez-Román et al., 2020).

171 3 Results

172 3.1 Performance of DUACS DT2021 products in the retrieval of sea level in the coastal band

173 This section presents the statistics of the comparisons performed between the DUACS DT2021 *all satellites* and *two satellites*
174 datasets and the tide gauge observations from the Copernicus catalogue in the coastal region of the European Seas in terms of
175 errors (rms differences) and variance of the differences between the datasets. According to Sánchez-Román et al. (2020), the
176 bootstrapping technique was applied to gain an estimation of the standard errors of the differences between the datasets.

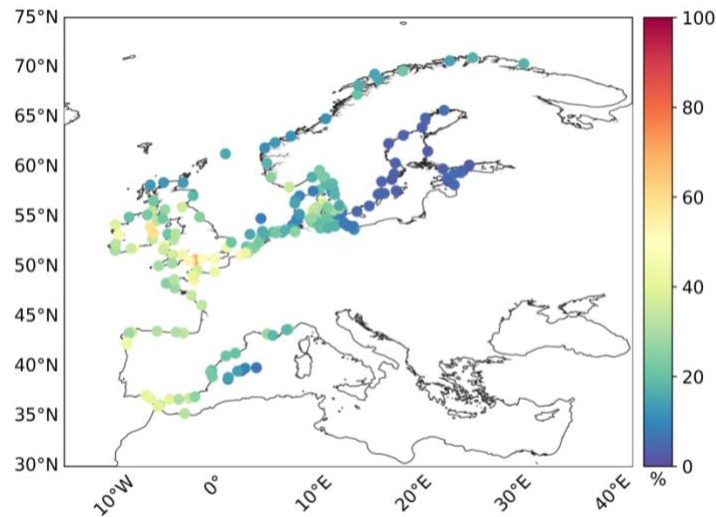
177 The mean value of the rms difference between the *all satellites* dataset and tide gauges is 4.11 cm, the variance of the
178 differences (altimetry–tide gauge) is 17 cm², and the mean distance between the location of the tide gauge and the
179 corresponding altimeter data with the highest correlation is 82 km (Table 1). These values raise to 4.35 cm, 19 cm², and 87
180 km, respectively, when using the *two satellites* dataset. The tide gauge stations (213 stations) common to both datasets were
181 used. Thus, the *all satellites* dataset reduces the rms differences with tide gauges in the European coasts by 5%, the variance
182 differences between the datasets by 10% and the mean distance between the most correlated altimetry point and tide gauges
183 by 6%. Also, the number of valid data pairs used to conduct the intercomparison enhanced by 0.2% when using the *all satellites*
184 dataset. This is due to the larger number of satellite missions used to generate this dataset, that provides lower errors in the
185 optimal interpolation procedure compared to the *two satellites* dataset.

186 **Table 1. Intercomparison of DUACS DT2021 satellite altimetry (ALT) and tide gauge (TG) data from the European coasts in terms**
187 **of the rms differences (cm) and variance (cm²) of the differences between the datasets. The number of tide gauge stations used in the**
188 **comparison, the mean distance between tide gauges and the most correlated gridded altimetry points, and the number of total data**
189 **pairs (altimetry-tide gauge) used in the computation are displayed. The common tide gauge stations for the *all satellites* and *two***
190 ***satellites* datasets were used. Values in parenthesis show the uncertainties (error bars) computed for the rms differences and variance**
191 **from the bootstrap method using 1.000 iterations. Finally, the improvement (%) of the *all satellites* dataset in comparison with tide**
192 **gauges in terms of lower rms differences, lower variance of the differences (altimetry-tide gauge), and lower mean distance between**
193 **the most correlated altimetry point and tide gauges with respect to the *two satellites* dataset is also displayed.**

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DUACS DT2021	<i>all satellites dataset</i>	<i>two satellites dataset</i>	<i>all satellites improvement</i>
rms diff. (cm)	4.11 (0.01)	4.35 (0.01)	5 %
var TG (cm ²)	89 (1)		
var ALT (cm ²)	81 (1)	79 (1)	
var TG-ALT (cm ²)	17 (1)	19 (1)	10 %
data pairs	1.163.588	1.161.315	0.2 %
stations	213		
Distance TG (km)	82	87	6 %

195 Fig. 1 shows the consistency between the DUACS DT2021 *all satellites* dataset and the tide gauge data computed from Eq.
196 (1) in Sánchez-Román et al. (2020). Consistency is expressed as the mean square differences between both datasets, computed
197 as the variance of the differences (altimetry–tide gauge), in terms of percentage of the tide gauge variance. Overall, mean
198 square differences lower than 5 % are observed in the central and eastern parts of the Baltic Sea, emphasising the precision of
199 the corrections applied to the altimeter data in the basin; whereas they reach values between 20% and 30% for stations located
200 in the connection region with the North Atlantic Ocean. The mean square differences are between 20% and 50% for most of
201 the stations located along the Atlantic shore, this including the Strait of Gibraltar area. Such large error could be related to
202 imprecisions of the correction applied (i.e. ocean tide and DAC) to the altimeter data (Pascual et al., 2008; Laíz et al., 2016;
203 Sánchez-Román et al., 2020), and also to both the larger spatiotemporal variability observed in this region (figure not shown),
204 and to a larger non tidal variance with respect to that found in the Baltic Sea (Von Schuckmann et al., 2018). Finally, the
205 Mediterranean and Norwegian Seas show mean square differences ranging between 15% and 30%, except for the Balearic
206 Islands (western Mediterranean) and the southwestern part of Norway where values between 5% and 15% are obtained. The
207 consistency between the DUACS DT2021 *two satellites* dataset and tide gauges (figure not shown) presents a quite similar
208 spatial pattern and results. These outcomes improve the ones reported in Sánchez-Román et al. (2020) from the
209 intercomparison conducted between Sentinel-3A L3 along track DUACS DT2018 dataset and tide gauge measurements in the
210 region computed over a period of two and a half years.



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212 **Figure 1. Location of the 213 tide gauges of the global product in the Copernicus catalogue along the European coasts and the**
213 **western Mediterranean Sea used to compare with altimetry data after applying the selection criteria described in the text. Colours**
214 **indicate the mean square differences between the tide gauge and altimetry sea level (DT2021 *all satellites* series). Units are the**
215 **percentage of the tide gauge variance.**

216 3.2 Improvement of DT2021 over DT2018 reprocessing

217 3.2.1 all satellites SLA dataset

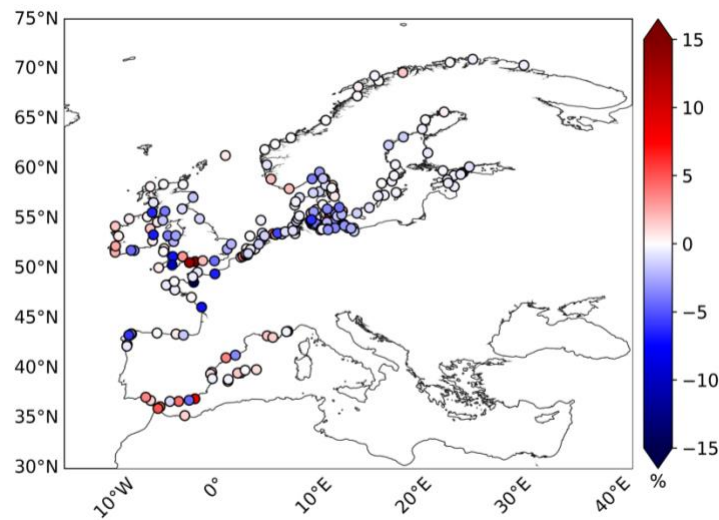
218 This section focuses on the statistics of the comparisons performed between the DUACS DT2021 and DT2018 reprocessing
 219 *all satellites* datasets and the tide gauge observations. The mean value of the rms difference between the DT2018 processing
 220 dataset and tide gauges is 4.22 cm, the variance of the differences (altimetry–tide gauge) is 18 cm², and the mean distance
 221 between the location of the tide gauge and the corresponding altimeter data with the highest correlation is 88 km (Table 2).

222 Overall, these values are larger than those reported in the previous section for the comparison using the DT2021 processing
 223 dataset (see Table 1). As a consequence, the DT2021 *all satellites* dataset reduces (i) the errors with tide gauges in the European
 224 coasts by 3%, (ii) the variance of the differences between the datasets by 5%, and (iii) the mean distance between the most
 225 correlated altimetry point and tide gauges by 7%. Also, the number of valid data pairs used to conduct the intercomparison is
 226 enhanced by 0.1% when using the DT2021 processing *all satellites* dataset. This highlights the impact of the new DUACS
 227 DT2021 reprocessing on the coastal areas, that provides more valid measurements, located closer to the tide gauge sites,
 228 compared to DT2018 reprocessing.

229 **Table 2. The same as Table 1 but for the intercomparison using the DUACS DT2018 reprocessing. The improvements (%) of the**
 230 **DUACS DT2021 reprocessing *all satellites* and *two satellites* SLA datasets with respect to the previous DT2018 reprocessing are also**
 231 **shown.**

DUACS DT2018	<i>all satellites</i> dataset	<i>two satellites</i> dataset	<i>all satellites</i> DT2021 improvement	<i>two satellites</i> DT2021 improvement
rms diff. (cm)	4,22 (0,01)	4,41 (0,01)	3 %	1 %
var TG (cm ²)		89 (1)		
var ALT (cm ²)	80 (1)	78 (1)		
var TG-ALT (cm ²)	18 (1)	19 (1)	5 %	no improvement
data pairs	1.162.231	1.161.349	0,1 %	no improvement
stations		213		
Distance TG (km)	88	90	7 %	3 %

232 The new standards and updated geophysical corrections applied to the DUACS DT2021 reprocessing compared to the previous
 233 DT2018 version have a direct impact on the observation of coastal ocean sea level in the gridded products. To characterise
 234 this impact, the difference between DT2021 and DT2018 consistency is shown in Fig. 2. The spatial distribution of the
 235 differences in consistency shows an overall better performance of the DT2021 reprocessing (blue colours) at the connection
 236 region between the Baltic Sea and the eastern North Atlantic Ocean and in most of the Atlantic shore, where an improvement
 237 larger than 15% is found for some tide gauge sites. A degradation of the DT2021 reprocessing is observed in most of the
 238 stations located in the western Mediterranean Sea and the southern coasts of Spain, including the Strait of Gibraltar area, and
 239 also in some stations located in the coasts of France, England and Ireland. On the other hand, discrepancies are hardly observed
 240 between the two reprocessings in the Baltic and Norwegian Seas.



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Figure 2. Spatial distribution of the differences (DT2021 minus DT2018) for the mean square differences between the tide gauge and altimetry sea level. Units are the percentage of the tidal variance. The SLA *all satellites* dataset has been used. Blue colours denote an improvement of the DUACS DT2021 reprocessing whilst red colours indicate its degradation with respect to the DT2018 reprocessing.

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3.2.2 two satellites SLA dataset

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We present here the statistics of the intercomparison between the climatic (two satellites) DT2021 and DT2018 processing and tide gauges. The mean value (Table 2) of the rms difference between the DT2018 processing dataset and tide gauges is 4.41 cm, the variance of the differences (altimetry–tide gauge) is 19 cm², and the mean distance between the location of the tide gauge and the corresponding altimeter data with the highest correlation is 90 km. If these results are compared with those reported above for the comparison using the DT2021 processing dataset (Table 1), it can be observed that the latter only improves the previous DT2018 reprocessing in terms of the errors with tide gauges, that are reduced by 3%, and the mean distance between the most correlated altimetry point and tide gauges, reduced by 7%; whereas the variance of the differences between the datasets and the number of valid data pairs used to conduct the intercomparison are quite similar. Such improvements are around 60% lower than those reported for the *all satellites* datasets. This fact is reflected in the spatial distribution of the differences between DT2021 and DT2018 consistency with tide gauges (figure not shown). A better performance of the DT2021 reprocessing is obtained at the connection region between the Baltic Sea and the eastern North Atlantic Ocean and in part of the Atlantic shore (coasts of United Kingdom and France). There is a degradation of the DT2021 reprocessing in most of the stations located in the western Mediterranean Sea and the southern coasts of Spain; and in some stations located in the coasts of France, England and Ireland. Also, negligible discrepancies between the two reprocessings are found in the Baltic Sea. This spatial pattern is quite similar to that obtained for the *all satellites* dataset described above. However, a degradation of the DT2021 reprocessing is observed in most of the stations located in both the NWS region (southern coasts of the North Sea) and the Norwegian Sea. This is a novelty with respect to the previous computation emphasising the overall poorer improvements of the DUACS DT2021 *two satellites* dataset over the previous reprocessing.

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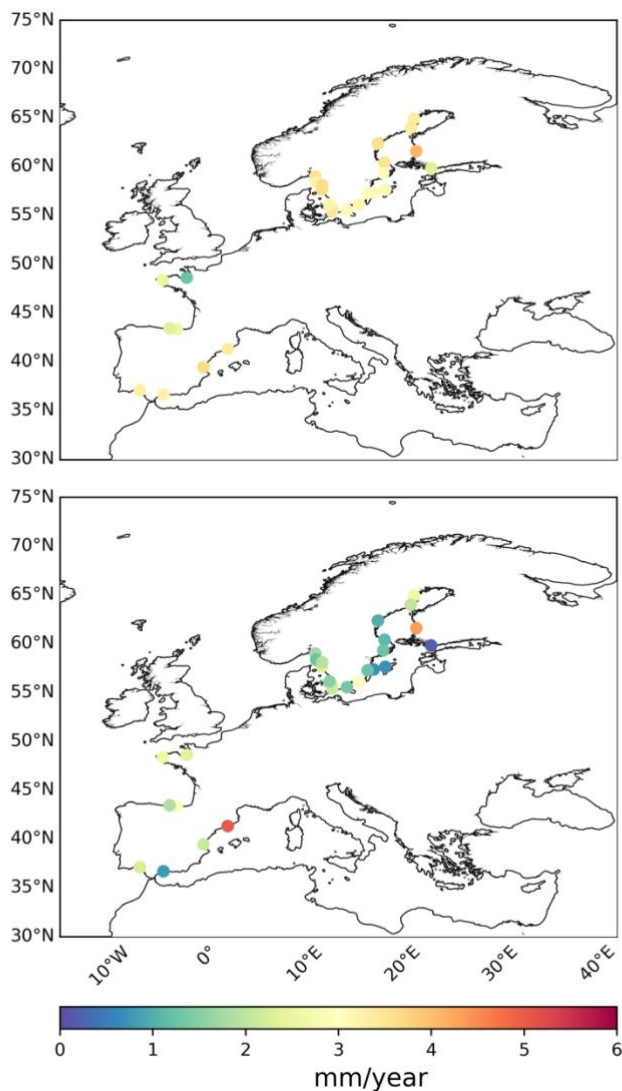
3.2 Performance of DT2021 reprocessing in monitoring the long term evolution of sea level

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The computation described above has been conducted by using all available information from the tide gauge dataset, thus including time series of different length spanning from few years to less than three decades (Table B1 of Appendix B). To assess the performance of DUACS DT2021 processing version in monitoring the long-term evolution of sea level in the coastal zone of the European Seas the analyses described above were repeated for a specific time period spanning 20 years: from 1 January 2000 to 31 December 2019. This time period has been chosen because of the largest number of available altimeter missions used to generate the *all satellites* SLA maps. Tide gauge time series with valid data within such time interval were

272 considered; this allowing the intercomparison altimetry–tide gauges for long-term time series with the same length. Moreover,
273 only tide gauge time series with at least 99% of valid data were used in order to allow the analysis of linear trends. This reduced
274 the original tide gauge dataset to a subset of 27 stations (Tables B1, B2 of Appendix B) mainly located in the northern half of
275 the Baltic Sea (70% of stations) with sparse stations distributed along the coasts of France and Spain (Fig. 3). This analysis
276 has also been conducted for the DUACS DT2018 reprocessing for comparison purposes.

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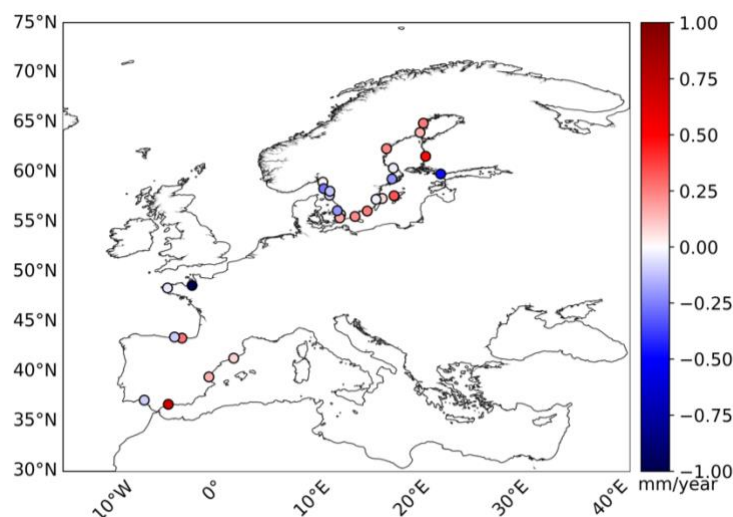
278

279 **Figure 3. Spatial distribution of linear trends (mm year^{-1}) for altimetry (upper panel) and tide gauges (lower panel) computed from**
280 **monthly averaged data for the 20 year time period spanning from 1 January 2000 to 31 December 2019. The *all satellites* dataset**
281 **from the DUACS DT2021 reprocessing has been used.**

282 Linear trends based on monthly observations at each tide gauge site (Fig. 3 and Table B2 of Appendix B) computed from
283 DUACS DT2021 *all satellites* dataset (upper panel) show a homogeneous spatial pattern with overall values varying from 2.30
284 to 4.10 mm year^{-1} in the Baltic and Mediterranean Seas and between 2.30 and 3.30 mm year^{-1} in the sparse stations located
285 along the North Atlantic European shore, except for the station of SaintMalo that presents a linear trend of 1.26 mm year^{-1} .
286 Linear trends computed from tide gauges (lower panel) exhibit a more heterogeneous spatial pattern with values ranging
287 between less than 1 mm year^{-1} for some stations located in the Baltic Sea, and 5.06 mm year^{-1} for the station of Barcelona
288 (western Mediterranean Sea). However, most of the tide gauge stations present trend values ranging from 1.30 to 3 mm year^{-1} .
289 ¹. These results provide further evidence, if needed, of the European Seas coastal sea level rise, including the westernmost part

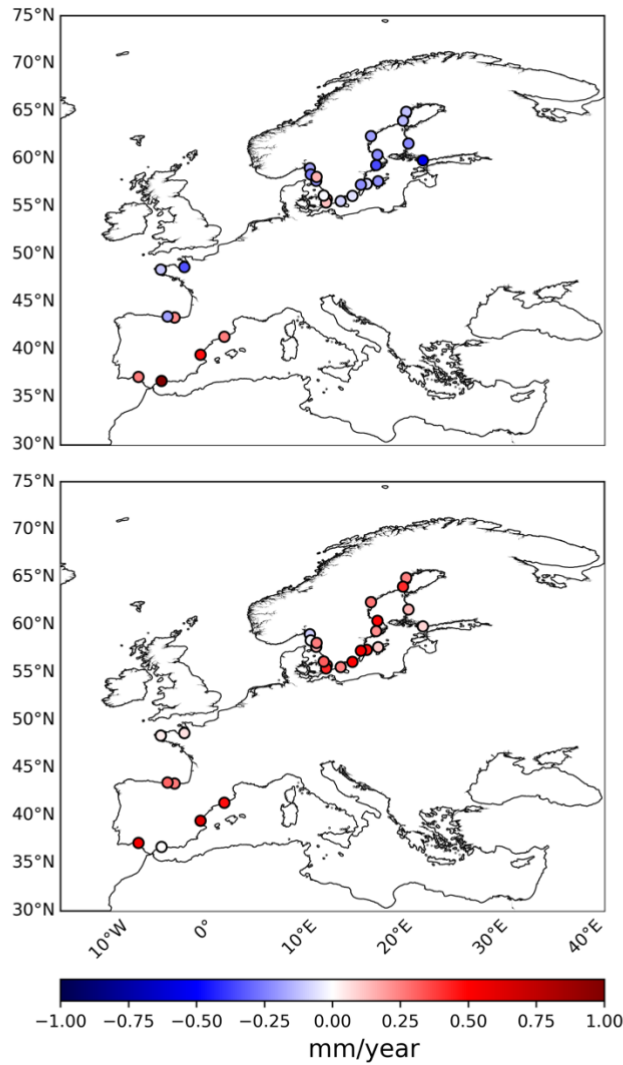
290 of the Mediterranean Sea. The differences in trends between the two datasets vary, in absolute values, between near 0 mm
291 year⁻¹ (Brest station, Atlantic French coast) to close to 2.60 mm year⁻¹ found in the station of Spikarna (Baltic Sea).

292 Linear trends computed from DUACS DT2021 *two satellites* dataset (figure not shown) exhibit a quite similar spatial pattern
293 with values ranging from 2.60 to 3.80 mm year⁻¹ in the Baltic and Mediterranean Seas; and between 2.40 and 3.40 mm year⁻¹
294 along the North Atlantic European coasts. However, some discrepancies between the two datasets are observed. These
295 differences, computed as *all satellites* minus *two satellites* datasets, are displayed in Fig. 4. Overall larger linear trends (up to
296 1 mm year⁻¹) were obtained for the *all satellites* dataset in the northernmost and central Baltic Sea as well as in the stations
297 located in the Mediterranean Sea whilst lower values of the same magnitude are mainly observed at the entrance of the Baltic
298 Sea and in most of the stations located along the North Atlantic European shore.



299
300 **Figure 4. Spatial distribution of the differences (*all satellites* minus *two satellites* datasets) for the linear trends (mm year⁻¹) from**
301 **altimetry computed from monthly averaged data for the 20 year time period spanning from 1 January 2000 to 31 December 2019.**
302 **The DUACS DT2021 processing version has been used. Blue (red) colours denote lower (larger) trends for the *all satellites* dataset.**

303 On the other hand, linear trends computed from the DT2018 reprocessing (figures not shown) exhibit a quite similar spatial
304 pattern than that reported for the DT2021 processing version with overall values ranging from 2.20 (2.40) to 4.35 (3.60) mm
305 year⁻¹ in the Baltic and Mediterranean Seas and between 2.40 (2.10) and 3.05 (2.85) mm year⁻¹ along the North Atlantic
306 European coasts for the *all satellites* (*two satellites*) dataset. Thus, hardly any differences in range are observed between the
307 *all satellites* dataset from the two reprocessing whereas these differences increase for the *two satellites* dataset with a lower
308 variability observed for the DT2018 reprocessing. This fact has an impact on the spatial distribution of the differences between
309 the two processing versions (Fig. 5).



310

311 **Figure 5. Spatial distribution of the differences (DT2021 minus DT2018 reprocessing) for linear trends (mm year^{-1}) for altimetry**
 312 **computed from the *all satellites* dataset (upper panel) and the *two satellites* dataset (lower panel). Monthly averaged data for the 20-**
 313 **year time period spanning from 1 January 2000 to 31 December 2019 has been used. Blue (red) colours denote lower (larger) trends**
 314 **for the DT2021 reprocessing.**

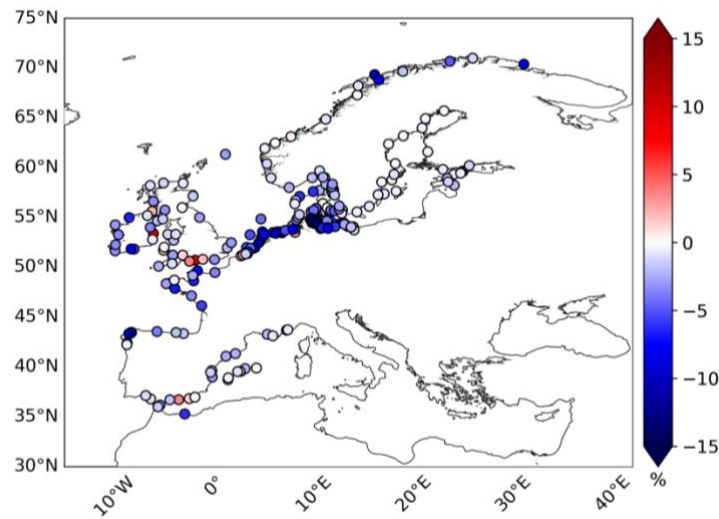
315 For the *all satellites* dataset (upper panel in Fig. 5), two different spatial patterns were observed with lower trends for the
 316 DT2021 reprocessing in the Baltic Sea basin and most of the stations located along the North Atlantic European coasts; whereas
 317 larger values are obtained for the tide gauge stations located in the western Mediterranean Sea and some sparse stations at the
 318 entrance of the Baltic Sea. On the contrary, the spatial distribution of the differences between the two reprocessing for the *two*
 319 *satellites* dataset (lower panel in Fig. 5) depicts a homogeneous spatial pattern with overall larger trends for the DT2021
 320 reprocessing except for the tide gauge station of Barseback located in the connection region between the Baltic Sea and the
 321 eastern North Atlantic Ocean (Table B2 in the Appendix B). Fig. 5 also reveals the differences between the two reprocessing,
 322 and for the two datasets, when comparing with linear trends from tide gauges: the *two satellites* dataset from the DT2021
 323 processing version presents larger differences with tide gauges with respect to the DT2018 reprocessing in the whole domain,
 324 whilst this is only observed for sparse stations along the North Atlantic shore and the stations located in the Mediterranean Sea
 325 for the *all satellites* dataset. Thus, closer results were obtained from the DT2021 *all satellites* product with respect to the former
 326 DT2018 processing version in most of the Baltic Sea region and the stations located along the North Atlantic European coast.

327 **4 Discussion and conclusions**

328 More than 29 years of DUACS Level 3 and Level 4 altimeter data have been recently reprocessed and delivered under the
329 name of DT2021 processing version through the Copernicus Marine Service and the Copernicus Climate Change Service.
330 The *all satellites* SLA products include all the available altimeter missions (ranging from 2 to 7 over the period considered in
331 this study), which makes the errors not constant in time since they depend on the number of satellites used. Maps from the *all*
332 *satellites* products provide the most accurate sea level estimation with the best spatial and temporal sampling of the ocean at
333 all times. The *two satellites* SLA dataset is obtained by merging a steady number of altimeters (two) in the satellite
334 constellation. This promotes consistent errors during the whole time period. Maps that include only two satellites are used to
335 compute the most homogeneous and stable sea level record over time and space. Thus, *two satellites* products are dedicated to
336 monitoring long term sea level evolution for climate applications and analysing ocean–climate indicators such as global and
337 regional MSL evolution (Taburet et al., 2019).

338 The new standards applied to the DT2021 version and the update of various geophysical correction parameters compared to
339 the previous release improved the *all satellites* product quality having a direct impact on the observation of coastal ocean sea
340 level in the gridded products. To achieve independent comparisons, SLA from altimetry in the coastal zone of the European
341 Seas were examined through comparison with in situ tide gauge measurements. Compared to the previous DT2018 version,
342 an improvement in the *all satellites* dataset was obtained, with a reduction of 3% in errors when compared with tide gauges
343 and of 5% in the variance of the differences between the datasets. The mean distance between the most correlated altimetry
344 point and tide gauges reduced by 7%. Also, the number of valid data pairs used to conduct the intercomparison enhanced by
345 0.1% when using the DT2021 processing. This highlights the impact of the new DUACS DT2021 version on the coastal areas,
346 that provides more valid measurements and located closer to the tide gauge sites, compared to DT2018 reprocessing. On the
347 other hand, almost no improvement of the DT2021 *two satellites* dataset over the previous reprocessing was found when using
348 all available information from the tide gauge dataset (time series of different length) in the computation: errors with tide gauges
349 were reduced by 1%, and the mean distance between the most correlated altimetry point and tide gauges was reduced by 3%.
350 The variance of the differences between the datasets and the number of valid data pairs used to conduct the intercomparison
351 were quite similar among the DT2021 and DT2018 processing versions. These improvements were around 60% lower than
352 those reported for the *all satellites* datasets. This fact could be explained by differences in the mapping parameters used for
353 the two products: DT2021 mapping parameters (i.e., spatial and temporal correlation scales, a priori errors on the
354 measurements) are evolved in CMEMS products (CMEMS QUID, 2022) with the objective to better retrieve mesoscale
355 signals, whilst no evolution of the mapping parameter was implemented in C3S DT2021 product (C3S PUG, 2022).

356 The quality assessment of DUACS DT2021 reprocessing revealed a better performance of the *all satellites* products in the
357 retrieval of SSH in the coastal zone with respect to the *two satellites* products for the time period investigated (27 years).
358 Namely, a reduction of 5% in errors with tide gauges and 10% in variance difference between altimetry and tide gauges was
359 obtained when using the *all satellites* dataset with respect to the *two satellites* product. This is because despite the larger
360 stability of the *two satellites* dataset, this product is optimised for climatic signal when analysing low frequency signals (SSH
361 trends). Thus, it is less performant for higher frequency signals. In this context (analysis of high frequency signals), the results
362 reported here show that the *all satellites* dataset should be considered for the analysis of long time series of SSH in the coastal
363 zone of the European Seas including all frequency signals. This can be clearly seen in Fig. 6 showing the differences (computed
364 as *all satellites* minus *two satellites* datasets) for consistency between altimetry and tide gauges.



365

366 **Figure 6. Spatial distribution of the differences (*all satellites* minus *two satellites* datasets) for the mean square differences between**
 367 **the tide gauge and altimetry sea level. Units are the percentage of the tidal variance. The DUACS DT2021 processing version has**
 368 **been used. Blue (red) colours denote an improvement (degradation) of the *all satellites* dataset.**

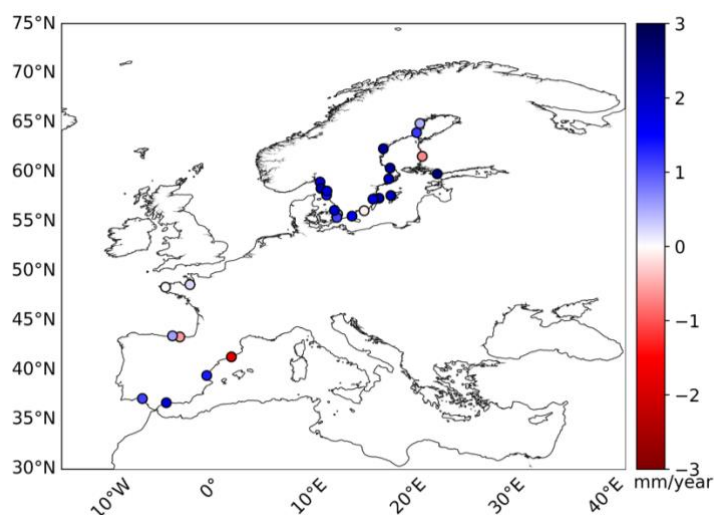
369 An overall better performance (blue colours) of the *all satellites* product with respect to the *two satellites* one was observed in
 370 the whole domain except in the Baltic Sea and the westernmost part of the Norwegian Sea, where similar results are obtained.
 371 The improvement is larger in most of the Atlantic shore, namely at the connection region between the Baltic Sea and the
 372 eastern North Atlantic Ocean, the NWS region and the northern Norwegian Sea, with a reduction in the variance difference
 373 between the two datasets larger than 15%. The Mediterranean Sea and the Strait of Gibraltar area show closer values between
 374 the two products with an improvement lower than 5%. These improvements could be explained by the better sampling of high
 375 frequency signal in the coastal zone in *all satellites* dataset due to the large number of altimeters available to generate the SLA
 376 maps compared to the *two satellites* maps. Improved mapping parameters for mesoscale (and thus high frequency) processes
 377 could also contribute. The observed degradation of the *all satellites* product with respect to the *two satellites* one at some tide
 378 gauge sites could be due to high-frequency local features badly captured by the *all satellites* product that translate in larger
 379 errors when comparing with tide gauges.

380 Linear trends based on monthly observations at each tide gauge site were computed to assess whether the DUACS DT2021
 381 release can be representative of the local sea level along the European coasts and western Mediterranean Sea. To do that, sea
 382 level linear trends for the period 2000-2019 were computed from both the *all satellites* and *two satellites* datasets. The analysis
 383 was repeated for the DT2018 reprocessing to have a term of comparison. A homogeneous spatial pattern with overall values
 384 ranging from 2.30 (2.40) to 4.10 (3.80) mm year⁻¹ was obtained for the *all satellites* (*two satellites*) dataset from the DT2021
 385 reprocessing. This promotes a mean trend for the whole domain of 3.14 (3.13) mm year⁻¹. These trends slightly differ from
 386 those computed from the tide gauge subset covering the 20 year time period, that show values ranging between less than 1 to
 387 5.06 mm year⁻¹; the mean trend for the whole domain is 1.96 mm/year.

388 Thus, trends computed from DT2021 products are on average around 1.2 mm year⁻¹ larger than those obtained from tide
 389 gauges. Similar overestimations in altimetry mean trends were reported by Agha-Karimi et al. (2021) in the Baltic Sea for
 390 datasets covering the time period spanning between 1993 and 2020. These discrepancies could be attributed to the
 391 heterogeneous distribution of both datasets and also the crustal land uplift due to postglacial rebound resulting from the last
 392 glacial age affecting the Baltic Basin, where most of the tide gauge stations are located. This translates in altimetry
 393 conventional measurements being not accurate enough in the coastal zone. On the other hand, when using the former DUACS
 394 DT2018 processing version slightly larger discrepancies with tide gauges were obtained for the *all satellites* dataset, with a

395 mean trend of 3.18 mm year⁻¹; whilst the *two satellites* product showed closer values to tide gauges with a mean linear trend
396 of 2.85 mm year⁻¹.

397 Overall, linear trend differences (altimetry – tide gauge) for the DT2021 reprocessing varying, in absolute value, from 0.16 to
398 2.57 mm year⁻¹, in an average of 1.43 mm year⁻¹ were obtained for the *all satellites* dataset. They varied from 0.03 to 2.65 mm
399 year⁻¹, in an average of 1.40 mm year⁻¹ for the *two satellites* dataset. These discrepancies are lower than 1.5 mm year⁻¹ in
400 average and corroborate the agreement and complementarity of the two techniques to measure sea level variability in the
401 coastal zone. They also emphasise a better performance of the C3S DT2021 dataset in the estimation of sea level linear trends
402 in the coastal zone. This was also corroborated by the computation conducted for the DT2018 reprocessing: lower differences
403 between tide gauge and altimetry trends computed from the *two satellites* dataset were obtained. Fig. 7 displays the spatial
404 distribution of the differences in trend computed as altimetry minus tide gauges for the *two satellites* dataset from the DT2021
405 reprocessing. An overall overestimation of trends from altimetry in the whole domain was obtained. On the contrary, three
406 tide gauge sites: Bilbao in the Atlantic Spanish coast, Pori at the eastern side of the Baltic Sea, and Barcelona in the western
407 Mediterranean Sea (Table B2 in Appendix B) showed a long-term sea level linear trend 0.58, 0.70 and 1.81 mm year⁻¹ larger,
408 respectively, than that found for the closest altimetry point with the largest correlation. The differences in trend could be
409 attributed to the aforementioned reasons rendering altimetry measurements being not accurate enough in the coastal zone. In
410 any case, the linear trends for the tide gauge of Barcelona described above are of the same order of magnitude than those
411 reported by Taibi and Haddad (2019) computed for the time period spanning from 1993 to 2015 (linear trend of 2.74 mm year⁻¹
412 ¹ for altimetry; 6.73 mm year⁻¹ for the tide gauge; trend difference of 3.99 mm year⁻¹), thus supporting the results obtained here.



413
414 **Figure 7. Spatial distribution of the differences in linear trends (mm year⁻¹) between the altimetry and tide gauge sea level computed**
415 **for the 20 year time period spanning from 1 January 2000 to 31 December 2019. The *two satellites* dataset from the DUACS DT2021**
416 **reprocessing has been used. Blue (red) colours denote a larger (lower) altimetry linear trend.**

417 The intercomparison conducted here between L4 gridded products from the new DUACS DT2021 release and the DT2018
418 version previously available; and tide gauges have demonstrated the better performance of the new DT2021 version in the
419 retrieval of sea level in the coastal zone of the European Seas. Furthermore, the *all satellites* dataset provided more accurate
420 sea level measurements when comparing with tide gauges respect to the climatic *two satellites* dataset due to the better
421 performance of the former for the assessment of higher than climatic frequency signals. On the opposite, when analysing linear
422 trends from 20-year long time series the *two satellite* dataset was the most suitable product for the assessment of long-term sea
423 level SSH trends in the coastal zone due to its larger stability to the detriment of the CMEMS *all satellites* dataset.

424 SLA and derived geostrophic velocities from altimeter data have been widely compared with in situ multiplatform
425 measurements by the coastal altimetry community in order to both validate altimetry measurements and demonstrate their
426 capabilities to monitor sea level and surface currents in the coastal zone. Heslop et al (2017) provided the first multiplatform
427 evaluation evolving data from the Sentinel-3A altimeter in the Balearic Sea (Western Mediterranean Sea). Their outcomes
428 demonstrated the capacity of this satellite mission to retrieve fine-scale oceanographic features of around 20 km of diameter
429 showing differences between along-track absolute dynamic topography (ADT) from altimetry and glider derived dynamic
430 height (DH) data along the satellite track of 1.23 cm. In the same region, Aulicino et al. (2018) compared along-track ADT
431 data from the SARAL-AltiKa mission with glider derived DH along two satellites tracks. They found a very similar spatial
432 pattern with differences ranging between 1.10 and 2.90 cm. Pascual et al., (2015) also conducted an assessment of
433 SARAL/AltiKa data in the coastal band through the comparison of along-track surface derived geostrophic velocities with
434 surface velocities from a coastal high-frequency (HF) radar system installed in the Ibiza Channel (Balearic Sea). These authors
435 found that the velocities derived from altimetry solved the general mesoscale features in the region with rms differences with
436 the in situ measurements of 13 cm s^{-1} .

437 The new Surface Water and Ocean Topography (SWOT) satellite mission, launched in December 2022, is considered to be
438 the next major breakthrough in satellite ocean observation (Morrow et al., 2023). The SWOT mission aims to provide SSH
439 measurements in two dimensions along a wide-swath altimeter track with an expected effective resolution down to
440 wavelengths of 15–30 km (Barceló-Llull et al., 2021). Thus, SWOT observations will fill the gap in our knowledge of the
441 15–150 km 2D SSH dynamics (Morrow et al., 2019) allowing, in some regions, the observation of the full range of mesoscale
442 features. The assessment of their impact on the large scale ocean circulation and climate system will be one of the major
443 challenges for the next decade.

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Appendix A

Table A1. Altimeter standard used in DT2021 release. Changes with the DT2018 reprocessing are highlighted in bold format.

	TP	Jason 1	Jason 2	Jason 3	ERS-1	ERS-2	ENVISAT	SARAL	Sentine 13A	Sentine 13B	Geosat FO	Cryosat 2	HY 2A	HY-2B
ORBIT	GSFC STD 18	POE-E	POE-F		Reaper		POE-E	POE-F	POE-F		GSFC	POE-F	POE-D	POE-F
IONOSPHERIC	Filtered dual-frequency altimeter range measurements (Ablain and Legeais, 2010); DORIS on Poseidon		Filtered dual-frequency altimeter range (Ablain and Legeais, 2010) from SSB C-band)	Filtered dual-frequency altimeter range measurements (Ablain and Legeais, 2010)	Reaper NIC09 model (Scharroo and Smith, 2010)	GIM (Iijima et al., 1999)	Filtered from L2; c>65: GIM (Iijima et al., 1999) corrected from 8mm bias	GIM (Iijima et al., 1999)	Filtered from L2		GIM (Iijima et al., 1999)			GIM (Iijima et al., 1999)
SEA STATE BIAS	Non parametric (Tran et al., 2010) on Topex; BM4 on Poseidon	2D Non parametric (Tran, 2015)	Non parametric (Tran et al., 2012)	Non parametric (Tran et al., 2012)	BM3 (Gaspar and Ogor, 1994)	Non parametric (Mertz et al., 2005)	2D Non parametric (Tran, 2017)	Non parametric (Tran, 2019)	Non parametric (Tran, 2012)		Non parametric (Tran, 2010)	2D Non parametric (baseline C) (Tran, 2018) Baseline C	Non parametric (Tran, 2012)	From L2 product
WET TROPOSPHERE	GPD+ (Ferrandes et al., 2015)	JMR (GDRE) radiometer	AMR radiometer		GPD+ (Ferrandes et al., 2015)		MWR radio meter reprocessed	Neural Network (5 entries) V4	MWR 3 radiometer		GFO Radiometer and ECMWF model	GPD+ (Fernandes Lázaro, 2016)	ECMWF model	ECMWF model
DRY TROPOSPHERE	ERA (1-hour) model based													
OCEAN TIDE	FES 2014 B (Carrere et al., 2016)													
POLE TIDE	(Desai et al., 2015); Mean Pole Location 2017 (Ries and Desai, 2017)													
SOLID TIDE	Elastic response to tidal potential (Cartwright and Tayler, 1971; Cartwright and Edden, 1973)													
INTERNAL TIDE	(Zaron, 2019) (HRETv8.1 tidal frequencies: M2, K1, S2, O1)													
MEAN DYNAMIC TOPOGRAPHY	MDT CNES_CLS_2018 (Mulet et al., 2021) merged with regional MDT CMEMS_2020 Mediterranean and Black Sea (Jousset and Mulet, 2020; Jousset et al., 2022)													

	TP	Jason 1	Jason 2	Jason 3	ERS-1	ERS-2	ENVISAT	SARAL	Sentine 13A	Sentine 13B	Geosat FO	Cryosat 2	HY 2A	HY-2B	
DYNAMICAL ATMOSPHERIC CORRECTION	TUGO High frequencies forced with analysed ERA 5 pressure and wind field + inverse barometer Low frequencies		TUGO HF forced with analysed ERA 5 pressure and wind field; and after 02/2016 MOG 2D HF force d with analysed ECMWF pressure and wind field + inverse barometer LF	MOG 2D HF forced with analysed ECMWF pressure and wind (Carrère and Lyard, 2003) (operational version 3.2.0) + inverse barometer LF	TUGO High frequencies forced with analysed ERA 5 pressure and wind field + inverse barometer Low frequencies			TUGO HF force d with analysed ERA 5 pressure and wind field; and after 02/2016 MOG 2D HF forced with analysed ECMWF pressure and wind field + inverse barometer LF	MOG2D High frequencies forced with analysed ECMWF pressure and wind field (Carrère and Lyard, 2003) (operational version 3.2.0) + inverse barometer Low frequencies		TUGO High frequencies forced with analysed ERA 5 pressure and wind field + inverse barometer Low frequencies		TUGO High frequencies forced with analysed ERA 5 pressure and wind field ; and after 02/2016 MOG 2D High frequencies forced with analysed ECMWF pressure and wind field + inverse barometer Low frequencies		MOG 2D High frequencies forced with analysed ECMWF pressure and wind field (Carrère and Lyard, 2003) (operational version 3.2.0) + inverse barometer Low frequencies
MEAN SEA SURFACE	Mean profile	Mean profile for repetitive orbit phases; Hybrid MSS (SIO,CNES/CLS15,DTU15) for geodetic/LOR phase	Mean profile	Mean profile for repetitive orbit phases ; Hybrid MSS (SIO, CNES/CLS15,DTU15) for geodetic phase		Mean profile	Mean profile for repetitive orbit phases; Composite MSS (SIO,CNES/CLS15,DTU15) for geodetic/drift phase		Mean profile	Hybrid MSS (SIO,CNES/CLS15,DTU15)	Hybrid (SIO,CNES/CLS15,DTU15)			MSS	

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Appendix B

Table B1. List of the 213 tide gauge records with their location and time period analysed. Bold stations indicate the tide gauge sites from the subset covering the 20-year period spanning from January 2000 to December 2019 listed in Table B2.

	Station name	Lon (°E)	Lat (°N)	Period analysed		Station name	Lon (°E)	Lat (°N)	Period analysed
1	Bagenkop	10.68	54.75	11/2006 - 05/2020	52	Ratan	20.90	63.99	01/1993 - 05/2020
2	Bandholm	11.48	54.83	01/2014 - 05/2020	53	Ringhals	12.11	57.25	01/1993 - 05/2020
3	Barhoeft	13.03	54.44	01/2011 - 05/2020	54	Rodby	11.35	54.65	01/2005 - 05/2020
4	Barseback	12.9	55.76	01/1993 - 05/2020	55	Rodvig	12.37	55.25	01/1993 - 05/2020
5	Bogense	10.08	55.57	01/2014 - 05/2020	56	Rohukula	23.42	58.90	12/2009 - 12/2019
6	Dragor	12.68	55.60	07/2011 - 05/2020	57	Roskilde	12.08	55.65	12/2011 - 05/2020
7	Drogden	12.71	55.54	01/1993 - 05/2020	58	Rostock	12.15	54.08	01/2011 - 05/2020
8	Eckernfoerde	9.84	54.47	01/2011 - 05/2020	59	Simrishamn	14.36	55.56	01/1993 - 05/2020
9	Faaborg	10.25	55.10	01/2014 - 05/2020	60	SjællandsOdde	11.37	55.97	01/1993 - 05/2020
10	Forsmark	18.21	60.41	01/1993 - 05/2020	61	Skagen	10.59	57.72	04/1993 - 09/2018
11	Fredericia	9.75	55.57	01/2005 - 05/2020	62	Skagsudde	19.01	63.19	10/1993 - 05/2020
12	Furuogrund	21.23	64.92	01/1993 - 05/2020	63	Skonor	12.83	55.42	01/1993 - 07/2018
13	Gedser	11.93	54.57	03/1993 - 05/2020	64	Smogen	11.22	58.35	01/1993 - 05/2020
14	GoteborgAgnesberg	12.01	57.79	01/2013 - 05/2020	65	Sonderborg	9.78	54.92	01/2014 - 05/2020
15	GoteborgEriksberg	11.91	57.70	01/2013 - 05/2020	66	Spikarna	17.53	62.36	01/1993 - 05/2020
16	GoteborgLarjeholm	12.01	57.77	01/2013 - 05/2020	67	Stenungsund	11.83	58.09	01/1993 - 05/2020
17	GoteborgTingstadstunneln	11.99	57.72	01/2013 - 05/2020	68	Stockholm	18.08	59.32	01/1993 - 05/2020
18	GoteborgTorshammen	11.79	57.68	01/1993 - 05/2020	69	Stralsund	13.10	54.32	01/2011 - 05/2020
19	Greifswald	13.45	54.09	01/2011 - 05/2020	70	Tallinn	24.76	59.44	11/2005 - 05/2020
20	Grena	10.93	56.41	01/1993 - 05/2020	71	TimmendorfPoel	11.38	53.99	01/2011 - 05/2020
21	Hanko	22.98	59.82	01/1993 - 05/2020	72	Travemuende	10.87	53.96	01/2005 - 05/2020
22	Heiligenhafen	11.01	54.37	01/2011 - 05/2020	73	Uddevala	11.89	58.35	12/2010 - 05/2020
23	Holbaek	11.72	55.72	12/2011 - 05/2020	74	Ueckermuende	14.07	53.75	01/2011 - 05/2020
24	Hov	10.27	55.92	12/2011 - 05/2020	75	Vedbaek	12.57	55.85	12/2011 - 05/2020
25	Juelsminde	10.02	55.72	12/1996 - 05/2020	76	Viken	12.58	56.14	01/1993 - 05/2020
26	Kalix	23.10	65.70	01/1993 - 05/2020	77	Virtsu	23.51	58.58	12/2009 - 05/2020
27	Kalkgrund	9.89	54.82	01/2011 - 05/2020	78	Visby	18.28	57.64	01/1993 - 05/2020
28	Kalvehave	12.17	55.00	01/2014 - 05/2020	79	Wisnar	11.46	53.90	01/2011 - 05/2020
29	Kappeln	9.94	54.66	01/2011 - 05/2020	80	Wolgast	13.77	54.04	01/2011 - 05/2020
30	Karrebaeksminde	11.65	55.18	01/2014 - 05/2020	81	BrestTG	-4.50	48.38	01/1993 - 05/2020
31	Kelnase	25.01	59.64	02/2017 - 05/2020	82	CherbourgTG	-1.64	49.65	01/1993 - 05/2020
32	KielHoltenau	10.16	54.37	01/2005 - 05/2020	83	ConcarneauTG	-3.91	47.87	06/1999 - 05/2020
33	KiellTG	10.27	54.50	01/2011 - 05/2020	84	LaRochelleTG	-1.23	46.15	10/1995 - 05/2020
34	Koege	12.20	55.45	01/2012 - 05/2020	85	LeConquetTG	-4.78	48.36	01/1993 - 05/2020
35	Koserow	14.00	54.06	11/2005 - 11/2019	86	LeHavreTG	0.11	49.48	01/1993 - 05/2020
36	KristinebergI	11.45	58.25	04/2012 - 05/2020	87	MarseilleTG	5.35	43.28	10/1998 - 05/2020
37	Kungsholmsfort	15.59	56.11	01/1993 - 05/2020	88	MonacoTG	7.42	43.73	04/1999 - 05/2020
38	Kungsvik	11.13	59.00	01/1993 - 05/2020	89	NiceTG	7.29	43.70	03/1998 - 05/2020
39	LandsortNorra	17.86	58.77	10/2004 - 05/2020	90	RoscoffTG	-3.97	48.72	01/1993 - 05/2020
40	Langballigau	9.65	54.82	01/2011 - 05/2020	91	SaintGildasTG	-2.25	47.14	02/1993 - 06/2017
41	Leppneeme	24.87	59.55	02/2017 - 05/2020	92	SaintMaloTG	-2.03	48.64	08/1993 - 04/2020
42	Luebeck	10.70	53.89	01/2011 - 05/2020	93	ToulonTG	5.91	43.12	01/1993 - 05/2020
43	Marviken	16.84	58.55	01/1993 - 09/2019	94	Aberdeen	-2.08	57.15	01/1993 - 05/2020
44	Munalaui	24.12	58.23	02/2016 - 05/2020	95	AlcudiaTG	3.14	39.83	09/2009 - 05/2020
45	Neustadt	10.81	54.10	01/2011 - 05/2020	96	AlgecirasTG	-5.40	36.18	07/2009 - 05/2020
46	OlandsNorraUdde	17.10	57.37	01/1993 - 05/2020	97	AlmeriaTG	-2.48	36.83	01/2006 - 05/2020
47	Onsala	11.92	57.39	06/2015 - 05/2020	98	Aranmore	-8.50	54.99	05/2008 - 05/2020
48	Oskarshamn	16.48	57.28	01/1993 - 05/2020	99	ArklowHarbur	-6.15	52.79	08/2003 - 05/2020
49	Paldiski	24.08	59.33	10/2006 - 05/2020	100	Ballycotton	-8.00	51.83	10/2010 - 05/2020
50	Pori	21.46	61.59	01/1993 - 05/2020	101	Ballyglass	-9.89	54.25	05/2008 - 04/2020
51	Porvoo	25.63	60.21	08/2014 - 05/2020	102	Bangor	-5.67	54.67	11/1994 - 05/2020
	Station name	Lon (°E)	Lat (°N)	Period analysed		Station name	Lon (°E)	Lat (°N)	Period analysed
103	BarcelonaTG	2.16	41.34	01/1993 - 05/2020	162	AlteWeserTG	8.13	53.86	01/2014 - 05/2020
104	Barmouth	-4.03	52.72	01/1993 - 05/2020	163	AndenesTG	16.13	69.33	01/2014 - 05/2020
105	BilbaoTG	-3.05	43.36	01/1993 - 05/2020	164	AWGTG	5.94	53.49	06/2015 - 05/2020
106	BonanzaTG	-6.34	36.80	01/1993 - 05/2020	165	BergenTG	5.32	60.40	01/2007 - 05/2020
107	Bournemouth	-1.87	50.71	06/1996 - 05/2020	166	BodoeTG	14.39	67.29	01/2007 - 05/2020
108	CarbonerasTG	-1.90	36.97	07/2013 - 05/2020	167	BorkumTG	6.75	53.56	01/2014 - 05/2020
109	Castletownbere	-9.90	51.65	12/2006 - 05/2020	168	Brouwershavensegat8TG	3.62	51.77	08/2014 - 12/2019
110	CorunaTG	-8.39	43.36	01/1993 - 05/2020	169	CadzandTG	3.38	51.38	08/2014 - 12/2019
111	Dundalk	-6.39	54.01	04/2008 - 01/2013	170	DenHelderTG	4.75	52.97	01/2014 - 12/2019
112	Felixstowe	1.35	51.97	01/1993 - 01/2011	171	EemshavenTG	6.84	53.46	08/2014 - 05/2020
113	Fenit	-9.86	52.27	01/2007 - 05/2020	172	EuroplatformTG	3.28	52.00	01/2014 - 12/2019
114	Ferrol2TG	-8.25	43.48	01/2007 - 05/2020	173	F3platformTG	4.72	54.85	08/2014 - 12/2019
115	FerrolTG	-8.33	43.46	01/2007 - 05/2020	174	HammerfestTG	23.68	70.66	01/2014 - 05/2020
116	Fishguard	-4.98	52.02	01/1993 - 05/2020	175	HanstholmTG	8.60	57.12	01/2015 - 05/2020
117	FormenteraTG	1.42	38.73	09/2009 - 05/2020	176	HarstadTG	16.55	68.80	01/2014 - 05/2020
118	GandiaTG	-0.15	38.99	07/2007 - 05/2020	177	HavnebyTG	8.57	55.09	01/2015 - 05/2020

119	GijonTG	-5.70	43.56	07/1995 - 05/2020	178	HelgeroaTG	9.86	59.00	01/2007 - 05/2020
120	Hinkley	-3.13	51.22	01/1993 - 05/2020	179	HelgolandTG	7.89	54.18	01/2014 - 05/2020
121	Holyhead	-4.62	53.32	02/2005 - 05/2020	180	HirtshalsTG	9.97	57.60	01/2015 - 05/2020
122	Howth	-6.07	53.39	10/2006 - 11/2019	181	HoekVanHollandTG	4.12	51.98	01/2014 - 12/2019
123	HuelvaTG	-6.83	37.13	09/1996 - 05/2020	182	HoernumTG	8.30	54.76	01/2014 - 05/2020
124	IbizaTG	1.45	38.91	01/2003 - 05/2020	183	HonningsvaagTG	25.97	70.98	01/2007 - 05/2020
125	Ilfracombe	-4.12	51.22	01/1993 - 05/2020	184	HuibertgatTG	6.40	53.57	06/2014 - 12/2019
126	Kinlochbervie	-5.05	58.46	01/1993 - 05/2020	185	IJmondstroompaalTG	4.52	52.46	08/2014 - 05/2020
127	LangosteiraTG	-8.53	43.35	01/2014 - 05/2020	186	K141TG	3.63	53.27	06/2015 - 05/2020
128	Leith	-3.18	55.99	01/1993 - 05/2020	187	KabelvaagTG	14.48	68.21	01/2007 - 05/2020
129	Llandudno	-3.82	53.31	05/2014 - 05/2020	188	KristiansundTG	7.73	63.11	01/2007 - 05/2020
130	Lowestoft	1.75	52.47	01/1993 - 05/2020	189	L91TG	4.87	53.57	06/2015 - 05/2020
131	MahonTG	4.27	39.89	10/2009 - 05/2020	190	LauwersoogTG	6.20	53.41	06/2015 - 12/2019
132	MalagaTG	-4.42	36.71	01/1993 - 05/2020	191	LichteilandGoeree1TG	3.67	51.93	01/2015 - 05/2020
133	MarinTG	-8.69	42.41	01/2010 - 05/2020	192	ListTG	8.44	55.02	01/2014 - 09/2018
134	MelillaTG	-2.92	35.29	10/2007 - 05/2020	193	MaloyTG	5.11	61.93	01/2007 - 05/2020
135	Milford	-5.05	51.72	01/1993 - 05/2020	194	MandoTG	8.58	55.28	01/2015 - 05/2020
136	Millport	-4.90	55.75	01/1993 - 05/2020	195	NieuwpoortTG	2.73	51.15	08/2014 - 05/2020
137	MotrilTG	-3.52	36.72	01/2010 - 05/2020	196	NorderneyTG	7.16	53.70	01/2014 - 05/2020
138	Newhaven	0.07	50.78	01/1993 - 05/2020	197	NorthCormorantTG	1.16	61.34	08/2014 - 05/2020
139	Newlyn	-5.53	50.10	01/1993 - 09/2018	198	OostendeTG	2.93	51.23	08/2014 - 05/2020
140	NorthShields	-1.43	55.00	01/1993 - 05/2020	199	OscarsborgTG	10.60	59.68	01/2007 - 05/2020
141	PalmadeMallorcaTG	2.64	39.56	09/2009 - 05/2020	200	RorvikTG	11.23	64.86	01/2007 - 05/2020
142	Plymouth	-4.19	50.37	01/1993 - 05/2020	201	StavangerTG	5.73	58.97	01/2014 - 05/2020
143	PortEllen	-6.19	55.63	01/1993 - 02/2011	202	ThyboronKystTG	8.21	56.71	01/2015 - 05/2020
144	Portpatrick	-5.12	54.84	01/1993 - 05/2020	203	TorsmindeKystTG	8.12	56.37	01/2015 - 05/2020
145	Portrush	-6.67	55.20	07/1995 - 05/2020	204	TregdeTG	7.55	58.01	01/2007 - 05/2020
146	Portsmouth	-1.11	50.80	01/1993 - 05/2020	205	TromsoeTG	18.96	69.65	01/2007 - 05/2020
147	RingaskiddyNMCI	-8.30	51.84	01/2012 - 05/2020	206	VardoeTG	31.10	70.37	01/2014 - 05/2020
148	RossaveelPier	-9.56	53.27	09/2020 - 05/2020	207	VikerTG	10.95	59.04	01/2007 - 05/2020
149	SaguntoTG	-0.21	39.63	07/2006 - 05/2020	208	VlakteVdRaanTG	3.24	51.50	08/2014 - 05/2020
150	SantanderTG	-3.79	43.46	01/1993 - 05/2020	209	VlielandHavenTG	5.09	53.30	08/2014 - 05/2020
151	StHelier	-2.12	49.18	01/1993 - 05/2020	210	WangeroogeTG	7.93	53.81	01/2014 - 05/2020
152	Stornoway	-6.38	58.22	01/1993 - 05/2020	211	WestkapelleTG	3.44	51.52	08/2014 - 05/2020
153	TarifaTG	-5.60	36.01	07/2009 - 05/2020	212	WilhelmshavenTG	8.15	53.51	01/2014 - 05/2020
154	TarragonaTG	1.21	41.08	05/2011 - 05/2020	213	ZeebruggeTG	3.20	51.35	08/2014 - 05/2020
155	Tobermory	-6.06	56.62	03/1993 - 05/2020					
156	ValenciaTG	-0.33	39.46	01/1993 - 05/2020					
157	VigoTG	-8.73	42.24	01/1993 - 05/2020					
158	Weymouth	-2.45	50.61	01/1993 - 05/2020					
159	Wick	-3.08	58.43	01/1993 - 05/2020					
160	ANDRATX	2.39	39.55	06/2011 - 05/2020					
161	AalesundTG	6.15	62.47	01/2007 - 05/2020					

Table B2. Tide gauge stations from the subset covering the 20-year period spanning from January 2000 to December 2019 located in the Baltic and Mediterranean Seas, and along the North Atlantic European shore. The location of the tide gauge sites, the linear trend (mm year⁻¹) computed from the DUACS DT2021 and DT2018 reprocessing *all satellites* and *two satellites* most correlated altimeter grid point to tide gauges, the tide gauges, and the mean trend value are displayed.

Station	Longitude (°E)	Latitude (°N)	Trend DT2021 all satellites (mm year ⁻¹)	Trend DT2021 two satellites (mm year ⁻¹)	Trend DT2018 all satellites (mm year ⁻¹)	Trend DT2018 two satellites (mm year ⁻¹)	Trend TG (mm year ⁻¹)
Barseback	12.90	55.76	3.22	3.26	2.99	2.88	2.60
Forsmark	18.21	60.41	3.49	3.53	3.71	3.03	1.12
Furuogrund	21.23	64.92	3.33	3.07	3.45	2.82	2.62
GoteborgTorshammen	11.79	57.68	3.62	3.70	3.84	3.56	2.31
Hanko	22.98	59.82	2.33	2.81	2.88	2.72	0.16
Kungsholmsfort	15.59	56.11	3.12	2.87	3.19	2.37	2.96
Kungsvik	11.13	59.00	3.52	3.50	3.70	3.60	1.72
OlandsNorraUdde	17.10	57.37	3.10	3.03	3.18	2.58	0.69
Oskarshamn	16.48	57.28	3.02	3.05	3.24	2.50	1.27
Pori	21.46	61.59	4.11	3.64	4.35	3.50	4.34
Ratan	20.90	63.99	3.34	3.19	3.48	2.77	2.02
Simrishamn	14.36	55.56	3.12	2.90	3.21	2.65	1.34
Skonor	12.83	55.42	3.43	3.26	3.33	2.83	2.14
Smogen	11.22	58.35	3.23	3.50	3.48	3.50	1.26
Spikarna	17.53	62.36	3.56	3.32	3.75	3.05	0.99
Stenungsund	11.83	58.09	3.63	3.75	3.48	3.49	1.93
Stockholm	18.08	59.32	3.02	3.23	3.37	3.01	1.26
Viken	12.58	56.14	3.21	3.42	3.22	3.10	1.51
Visby	18.28	57.64	3.13	2.80	3.35	2.70	0.73
Brest	-4.50	48.38	2.57	2.61	2.68	2.57	2.64
SaintMalo	-2.03	48.64	1.26	2.59	1.60	2.54	2.37
Bilbao	-3.05	43.36	2.63	2.36	2.40	2.10	2.94
Huelva	-6.83	37.13	3.30	3.39	3.05	2.85	2.27
Santander	-3.79	43.46	2.33	2.42	2.51	2.12	1.88
Barcelona	2.16	41.34	3.33	3.25	3.07	2.77	5.06
Malaga	-4.42	36.71	3.25	2.58	2.19	2.58	0.77
Valencia	-0.33	39.46	3.62	3.47	3.15	2.84	2.13
Mean value			3.14	3.13	3.18	2.85	1.96

Data availability

Altimetry datasets are available from the Copernicus Marine Service web portal (<https://resources.marine.copernicus.eu/products/>, last access: 15 July 2022). Tide gauge measurements are available from the Copernicus Marine INS-TAC data repository web portal (www.marineinsitu.eu, last access: 3 June 2022). Tide gauge data are provided by the following regional in situ data production centres: Puertos del Estado (Spain) for the Iberia-Biscay-Ireland region; HCMR (Greece) for the Mediterranean Sea; IMR (Norway) for the Arctic; SMHI (Sweden) for the Baltic Sea; BSH (Germany) for the North West Shelves region; Coriolis (France) for the global ocean. The ancillary data used to obtain the Dynamic Atmospheric Correction applied to the altimetry grid point closest to the tide gauge locations are available at the AVISO webpage: <https://www.aviso.altimetry.fr/en/> (last access: 16 May 2022).

Author contributions

Conceptualisation: Antonio Sánchez Román, M. Isabelle Pujol, Ananda Pascual and Yannice Faugère; altimetry data processing: M. Isabelle Pujol; tide gauge data processing: Antonio Sánchez Román; statistical analysis: Antonio Sánchez Román and Ananda Pascual; manuscript writing: Antonio Sánchez Román, with inputs from all co-authors. All authors have read and agreed to the published version of the manuscript.

Competing interests

The authors declare that they have no conflict of interest

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