Intensive photochemical oxidation in the marine atmosphere: Evidence from direct radical measurements

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Abstract: Comprehensive observations of hydroxyl (OH) and hydroperoxy (HO$_2$) radicals were conducted in October 2019 at a coastal continental site in the Pearl River Delta (YMK site, 22.55°N, 114.60°E). The daily maximum OH and HO$_2$ concentrations were $(4.7–9.5) \times 10^6$ cm^{-3} and $(4.2–8.1) \times 10^8$ cm^{-3}, respectively. The synchronized air mass transport from the northern cities and the South China Sea exerted a time-varying influence on atmospheric oxidation. Under a typical ocean-atmosphere (OCM), reasonable measurement model agreement was achieved for both OH and HO$_2$ using a 0-D chemical box model incorporating the regional atmospheric chemistry mechanism version 2-Leuven isoprene mechanism (RACM2-LIM1), with daily averages of 4.5×10^6 cm^{-3} and 4.9×10^8 cm^{-3}, respectively. Land mass (LAM) influence promoted more active photochemical processes, with daily averages of 7.1×10^6 cm^{-3} and 5.2×10^8 cm^{-3} for OH and HO$_2$, respectively. Heterogeneous uptake had certain effects on HOx chemistry, but the influence of the halogen mechanism was limited by NOx level. Intensive photochemistry occurred after precursor accumulation, allowing local net ozone production comparable with surrounding suburban environments (5.52 ppb/h during the LAM period). Rapid oxidation process was accompanied by a higher diurnal nitrous acid (HONO) concentration (> 400 ppt). After a sensitivity test, HONO-related chemistry elevated the ozone production rate by 33% and 39% during the LAM and OCM periods, respectively. The nitric acid (P(HNO$_3$)) and sulfuric acid (P(H$_2$SO$_4$)) formation rates also increased simultaneously (~43% and ~48% for LAM and OCM sectors, respectively). Without the HONO constraint, simulated O$_3$ decreased from ~75 ppb to a global background (~35 ppb), and daytime HONO concentrations were reduced to a low level (~70 ppt). For coastal cities, the particularity of the HONO chemistry tends to influence the ozone-sensitive system and eventually magnifies the background ozone. Therefore, the promotion of oxidation by elevated precursors deserves a lot of attention when aiding pollution mitigation policies.

Keywords: FAGE-LIF; OH and HO$_2$ radicals; Atmospheric oxidation; Marine boundary layer; Precursors;

1 Introduction

The marine boundary layer (MBL) occupies 71% of the planetary boundary layer, is a massive active carbon sink on Earth, and plays an irreplaceable role in coping with global climate change (Stone et al., 2012; Woodward-Massey et al., 2022a; Liu et al., 2022a). As a typical background atmosphere on the Earth, the MBL is equivalent to a natural smog chamber with limited anthropogenic emissions and is characterized by low NOx (the sum of nitric oxide (NO) and nitrogen dioxide (NO_2)) and non-methane hydrocarbons (NMHCs) under a layer of clean air (Woodward-Massey et al., 2022a). The lifetime of OH radical, a key oxidant, is on the order of a few hundred milliseconds (Fuchs et al., 2012). Due to the scarcity of oxidation precursors, including nitrous acid (HONO), formaldehyde (HCHO), and NMHCs, the reaction between O^1D and water vapor generally dominates the radical initiation pathway in the marine environment. For example, in a tropical boundary layer observation experiment (reactive halogens in the marine boundary layer, RHaMBLe), ozone photolysis was found to account for 70% of the OH radical source based on the master chemical mechanism (MCM) (Whalley et al., 2010). The vital role of ozone photolysis is contrasting with typical polluted and semi-polluted areas investigated in a series of field campaigns, in which the propagation routes were found to dominate the radical source (Yang et al., 2021a; Tan et al., 2019a). Therefore, studying the radical chemistry in the MBL provides a valuable opportunity to test the current understanding of atmospheric oxidation mechanisms in a natural setting.

Since the earliest observations off the coast of northern Norfolk in the Weybourne Atmospheric Observatory Summer Experiment in June 1995 (WAOSE95), more observations and simulations of radical chemistry in the MBL environment have been conducted using ground-based, airborne, and shipborne instruments (Qi et al., 2007; Kanaya et al., 2002; Kanaya et al., 2001; Mallik et al., 2018; Woodward-Massey et al., 2022b; Carpenter et al., 2011; Grenfell et al., 1999; Brauers et al., 2001; Whalley et al., 2010). Most field measurements have yielded well-reproduced OH and HO_2 concentration profiles via chemical mechanisms, with differences of within ~20%.

However, the base model is not sufficient to describe the radical chemistry in some exceptional cases, especially in regard to the HO_2 radical. Considering the practical association between halogen (Cl, Br, and I) chemistry and heterogeneous chemistry in marine new particle formation, particularly the involvement of heterogeneous iodine-organic chemistry, exploring the synchronous influence of these mechanisms on HOx (OH and HO_2) radical chemistry in the MBL region is a worthy endeavor (Xu et al., 2022; Huang et al., 2022). The mixing of air masses of continental and marine origins can lead to more variability in radical concentrations. During seasonal measurements of both OH and HO_2 in the Atlantic Ocean, variance analysis indicated that around 70% of the variance of OH and HO_2 was due to diurnal behavior (in the form of photolysis frequency), while the remaining variance was attributed to long-term seasonal cycles (in the form of the changes in O_3, CO and air mass contribution) (Vaughan et al., 2012).

The Chinese economy has undergone rapid development in recent years, and the co-occurrence of primary and secondary regional pollution has become a severe problem (Lu et al., 2019; Liu et al., 2022c). The interactions between air pollutants from upwind cities, shipping vessels, and other anthropogenic emissions lead to precursor accumulation (Sun et al., 2020; Zeren et al., 2022). The background ozone concentration in key regions of China has increased year by year, highlighting the significant influence of anthropogenic activities on the atmospheric oxidation in background regions in China (Wang et al., 2009; Chen et al., 2022). However, little research has been dedicated to the radical chemistry and oxidation mechanism in regions with both coastal and continental features. To fill this research gap, in this study, a field campaign was conducted on photochemistry in the MBL at a coastal site in the Pearl River Delta. The OH and HO_2 radicals associated with other related species were measured in October 2019, and the radical-related oxidation process was identified to determine the photochemical efficiency in the marine atmosphere.

2 Materials and methods

2.1 Site description

As shown in Fig. 1(a), this observation campaign lasted for 11 days from October 18 to October 28, 2019, in Yangmeikeng (YMK, 22.55°N, 114.60°E), a coastal site in Shenzhen, Pearl River Delta. As the core city of the Greater Bay Area, Shenzhen is bordered by Dongguan to the north, Huizhou to the east, and Hong Kong to the south. The YMK site is on the Dapeng Peninsula, to the southeast of Shenzhen, between Mirs Bay and Daya Bay. As it is adjacent to the port of Hong Kong, precursors from ship emissions may influence the atmospheric chemistry. The site is a part of Shenzhen Ecological Monitoring Center station, approximately 35 m above sea level, and the sea is approximately 150 m to the east. No apparent local emissions exist, and the surrounding forest is lush (Fig. 1(b)). Previous literatures reported the monoterpene concentration in the YMK site, with a daily mean of 0.187 ppb (Zhu et al., 2021). Aboundant biogenic emissions will likely influence the local chemistry. In addition to anthropogenic and vegetation emissions, the site is also affected by the synchronization of plumes from northern cities and the South China Sea (Niu et al., 2022; Xia et al., 2021). Due to its significant time-varying pollution characteristics, this area is an ideal site for studying the effects of plume transport on atmospheric oxidation.

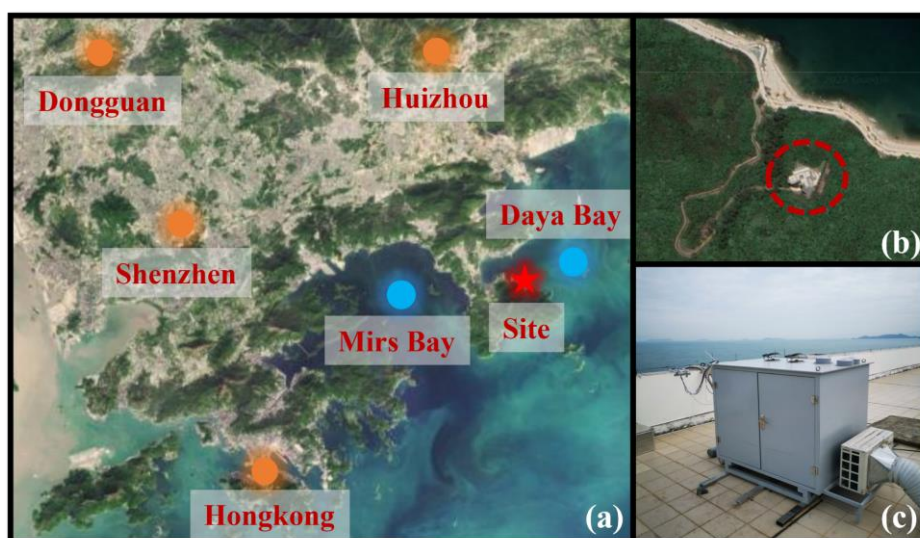



Fig. 1. Details of the observation site**(a)** The location of the measurement site and surrounding cities. The satellite map data is extracted from Google Earth.**(b)** Th close shot of the measurement site location.**(c)** The actual image for the LIF-Box.

Using the hybrid single-particle Lagrangian integrated trajectory (HYSPLIT)

model, the 24-h backward trajectories on special days were obtained. In Fig. S1, the red, blue, and green trajectories represent the results at altitudes of 100, 500, and 1000 m above ground level, respectively. Two typical transportation pathways dominated the air parcels. One originated from the northern megacities in the Pearl River Delta (defined as the land mass, LAM), especially on October 18, 19, and 27. In contrast, a clean air mass from the east or northeast was mainly transported to the observation site from the ocean (defined as the ocean mass, OCM), with representative episodes on October 22, 25, and 26.

2.2 Instrumentation

2.2.1 HOx radical measurements

The OH and HO_2 radicals were measured via laser-induced fluorescence (LIF). The OH radical can be directly measured by exciting the fluorescence using a 308-nm laser. HO_2 is converted into the OH radical via chemical transformation and then detected in the form of OH radical. The self-developed instrument, the Anhui Institute of Optics Fine Mechanics-LIF (AIOFM-LIF), was used to conduct the measurements (Zhang et al., 2022a; Wang et al., 2021; Wang et al., 2019). This system has been used in key regions of China, including the Yangtze River Delta, Pearl River Delta, and Chengdu-Chongqing region, and achieved good performance in a comparison experiment with a LIF system jointly developed by Forschungszentrum Jülich and Peking University (PKU-LIF) (Zhang et al., 2022b).

The system and detection interference process have been described in detail in previous studies (Zhang et al., 2022a). Briefly, the system consists of a laser output module, a radical detection module, and a control and data acquisition module. These modules are integrated into a sampling box with constant temperature and humidity control (Fig. 1(c)). The laser output module is a union of an Nd:yttrium-aluminum-garnet (YAG) solid-state laser, a 532-nm laser output, and a tunable dye laser. The radical detection module utilized a single pass laser configuration, and the laser beam was amplified to a diameter of 8 mm. OH and HO_2 fluorescence cells are combined in parallel and share a common axial optical path. The 308-nm laser is introduced into

162 the HO_2 cell first and then into the OH cell via an 8-m fiber. To maintain the detection

163 efficiency, the power in the OH fluorescence cell should be at least 15 mW. In the

164 detection process, a set of lenses was deployed and positioned in front of the

165 microchannel plate detector (MCP) to boost the fluorescence collection capacity. Each

166 MCP detector contains a timing control instrument to optimize the signal-to-noise

167 ratio (SNR) of the fluorescence detection. Efficient ambient air sampling was

168 achieved using an aluminum nozzle (0.4 mm orifice), and the pressure in the chamber

169 was maintained at 400 Pa via a vortex vacuum pump (XDS35i, Edwards) to reduce

170 fluorescence quenching.

171   A wavelength modulation for the background measurement that periodically

172 switches from an on-resonant state to a non-resonant state has been widely used to

173 obtain spectral zero. Due to the synchronous reaction at 308nm, wavelength

174 modulation is not applicable to ozone photolysis interference. Through laboratory

175 experiments, at 20 mW laser energy, every 1% water vapor concentration and 50 ppb

176 ozone concentration can generate a 2.5×10^5 cm^{-3} OH concentration. The results in

177 this paper have subtracted the ozone photolysis interference (Fig. S2). In terms of

178 system design, the AIOFM-LIF system incorporates a short-length inlet design to

179 minimize interferences from ozonolysis and other unknown factors (the distance from

180 radical sampling to flourescence excitation is ~150 mm). An OH measurement

181 comparison with an interference-free instrument, PKU-LIF, was conducted in a real

182 atmosphere in a previous study (Zhang et al., 2022b). The ozonolysis interference on

183 the measurement consistency of both systems was excluded under high-VOCs

184 condition. Overall, the key parameters related to ozonolysis reactions (O_3、alkenes、

185 isoprene and NOx) in YMK was similar to that during the intercomparison

186 experiment, implies that the chemical conditions do not favor the generation of

187 potential interference to OH measurement (Table S1). For HO_2 measurement, the NO

188 gas was mixed with 2% in N_2 to achieve HO_2-to-OH conversion. NO was passed

189 through a ferrous sulfate filter to remove impurities (NO_2, HONO, and so on) before

190 being injected into the detection cell. The NO concentration (~1.6×10^{12} cm^{-3})

191   corresponding to a conversion efficiency of ~15% was selected to avoid $RO_2 \rightarrow HO_2$

192   interference (especially from RO_2 radicals derived from long-chain alkanes ($C \geq 3$),

193   alkenes, and aromatic hydrocarbons). Previous study denoted that the percentage

194   interference from alkene-derived RO_2 under these operating conditions was no more

195   than 5% (Wang et al., 2021).

196   A standard HOx radical source was used to complete the calibration of the

197   detection sensitivity (Wang et al., 2020). The radical source is based on the

198   simultaneous photolysis of H_2O/O_2 by a 185 nm mercury lamp. Humidified air flow is

199   introduced to produce equal amounts of OH and HO_2 radicals after passing the

200   photolysis region. The flow remained in a laminar condition with a maximum flow

201   rate of 20 SLM (standard liters per minute). As the luminous flux in photolysis region

202   is difficult to accurately measure, the linearly correlation between ozone concentration

203   and 185 nm light flux was established. Ozone concentration in the flow tube was

204   measured by a home-made Cavity Ring Down Spectrometer (CRDS, and the

205   detection limit is 15 ppt@30 s, 1σ). Mercury lamp intensity is adjusted to establish.

206   The instrument was calibrated every 1 or 2 days (except for shutdown during rainy

207   periods), and the sensitivity used for the data processing was an average of all of the

208   calibration results. In the YMK campaign, the humidity varied between 40 – 80% (Fig.

209   S3). In order to test different atmospheric conditions, both low (~40%) and high

210   (~70%) levels of water vapor were selected to produce OH and HO_2 radicals for

211   calibration, and the corresponding HOx concentration obtained from the standard

212   source was 1.0×10^9 cm^{-3} and 1.8×10^9 cm^{-3}, respectively (Zhang et al., 2022b).

213   Considering the system uncertainty and calibration uncertainty, the detection

214   limits of the OH and HO_2 radicals were 3.3×10^5 cm^{-3} and 1.1×10^6 cm^{-3} (60 s, 1σ),

215   respectively. At a typical laser power of 15 mW, the measurement accuracy for OH

216   and HO_2 measurement was 13% and 17%, respectively.

217   **2.2.2 Supporting measurements**

218   In addition to measuring the HOx radicals, an extensive suite of relevant species

219   was also measured close to the LIF instrument to improve the analysis of the radical

220   photochemistry. Detailed information about the measurement instrument is presented

221   in Table S2, including the meteorological parameters (wind speed (WS), wind

222   direction (WD), temperature (T), relative humidity (RH), pressure (P), and solar

223   radiation (J-values)) and chemical parameters (ozone (O_3), carbon monoxide (CO),

224   sulfur dioxide (SO_2), HONO, NO, NO_2, HCHO, NMHCs, and particulate matter

225   ($PM_{2.5}$)). HONO measurement was conducted using a commercial Long-Path

226   Absorption Photometer (LOPAP). The LOPAP method utilizes two absorption tubes

227   in series for differential correction, which effectively eliminates the influence of

228   known interfering substances such as NO_2 and N_2O_5, offering an advantage over

229   traditional wet chemistry methods. Zero air measurements were taken every 8 hours

230   for a duration of 20 minutes to correct for instrument baseline fluctuations. This

231   method has been extensively tested for its suitability in detecting HONO in complex

232   atmospheric conditions, as demonstrated in previous studies by (Yang et al., 2022a;

233   Yang et al., 2021b; Wang et al., 2023). Eight measured photolysis rates (j(NO_2),

234   j(H_2O_2), j(HCHO), j(HONO), j(NO_2), j(NO_3), j(O1D)) were used as model constraints.

235   In addition to HCHO, other volatile organic compounds (VOCs) were detected using

236   a gas chromatograph coupled with a flame ionization detector and mass spectrometer

237   (GC-FID-MS). Ninety-nine types of VOCs, including C_2–C_{11} alkanes, C_2–C_6 alkenes,

238   C_6–C_{10} aromatics, halohydrocarbons, and some oxygenated VOCs (OVOCs), were

239   observed using the GC-FID-MS at a 1-h time interval. Only isoprene was considered

240   as a representative of biogenic VOCs (BVOCs). All of the instruments were located

241   close to the roof of the monitoring building, nearly 12 m above the ground to ensure

242   that all of the pollutants were located in a homogeneous air mass.

2.3 Model description

244   A 0-D chemical box model incorporating a condensed mechanism, the regional

245   atmospheric chemistry mechanism version 2-Leuven isoprene mechanism (RACM2-

246   LIM1), was used to simulate the radical concentrations and the generation of ozone

247   (Stockwell et al., 1997; Griffith et al., 2013; Tan et al., 2017). The meteorological

248   parameters, pollutants, and precursor concentrations mentioned in Section 2.2.2 were

input into the model as boundary conditions. All of the constraints were unified to a temporal resolution of 15 min through averaging or linear interpolation. The overall average during the observations was substituted for large areas of missing data due to instrument maintenance or failure. Three days of data were entered in advance as the spin-up period, and a synchronized time-dependent dataset was eventually generated. The hydrogen (H_2) and methane (CH_4) concentrations were set to fixed values of 550 ppb and 1900 ppb, respectively. The physical losses of species due to processes such as deposition, convection, and advection were approximately replaced by an 18 h atmospheric lifetime, corresponding to first order loss rate of ~1.5 cm/s (by assuming a boundary layer height of about 1 km). The sensitivity analysis shows that when the lifetime changes within 8 – 24 hours, the values differed less than 5% for both OH, HO_2, k_{OH} (Fig. S4). According to the measurement accuracy, the simulation accuracy of the model for the OH and HO_2 radicals was 50% (Zhang et al., 2022a).

In addition, another steady-state calculation method (PSS) can also be used to estimate the concentrations of OH and HO_2 radicals (Eq. (1)(2), (Woodward-Massey et al., 2022a; Slater et al., 2020)). Since the k_{OH} and RO_2 concentrations were not obtained in this observation, simulated values are used as substitutes. Other radical and reactive intermediates are actual values that measured from the instruments in Table S2.

$$[OH]_{PSS} = \frac{j_{HONO}[HONO] + \varphi_{OH}j(O^1D)[O_3] + k_{HO_2+NO}[NO][HO_2]}{k_{OH}} \quad (1)$$

$$[HO_2]_{PSS} = \frac{k_{CO+OH}[CO][OH] + j_{HCHO}[HCHO] + k_{RO_2+NO}[NO][RO_2]}{k_{HO_2+NO}[NO]} \quad (2)$$

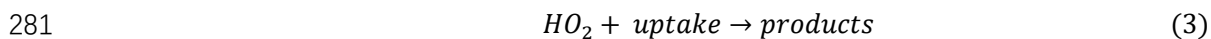
Considering the environmental characteristics of the MBL, the gas-phase mechanisms for bromine (Br) and iodine (I) were introduced into the base model to diagnose the impacts of the reactive bromine chemistry. The details of the mechanisms involved are listed in Tables S3 and S4. The halogen species were not available in the YMK site, so the typical levels of BrO and IO concentration in MBL site was used as a reference value (average daytime concentration of ~5 ppt) (Xia et al., 2022; Bloss et al., 2010; Whalley et al., 2010).

The heterogeneous uptake of HO_2 is considered to play an important role in the

278   MBL region (Whalley et al., 2010; Zou et al., 2022; Woodward-Massey et al., 2022a).

279   In order to assess the impact of HO_2 uptake on HOx radical chemistry, we

280   incorporated HO_2 uptake reaction into the base model (Eq. (3) - (5)).

$$HO_2 + uptake \rightarrow products \tag{3}$$

$$k_{HO_2+uptake} = \frac{\gamma \times ASA \times v_{HO_2}}{k_{HO_2+NO}[NO]} \tag{4}$$

$$v_{HO_2} = \sqrt{\frac{8 \times R \times T}{0.033 \times \Pi}} \tag{5}$$

284   Here, ASA represents the aerosol surface area [μm^2 cm^{-3}], which can be estimated

285   as 20 times the $PM_{2.5}$ concentration [$\mu g/cm^3$]. v_{HO_2} [cm^{-1}] can be calculated using Eq.

286   (5), where T and R represent the temperature and gas constant, respectively. The

287   heterogeneous uptake coefficien (γ) for HO_2 usually has high uncertainty, with typical

288   values ranging from 0 to 1 (Song et al., 2021). In this study, we set γ to 0.08 to

289   evaluate the influence of HO_2 uptake on radical concentrations.

3 Results

3.1 Meteorological and chemical parameters

292   Fig. S3 presents the time series of the main meteorological parameters and

293   pollutants during the observation period at the YMK site. Except for on 2 days,

294   October 26 and 28, the meteorological characteristics of the other days were generally

295   stable. The daily maximum T, RH, and J-values did not vary significantly. The

296   suitable temperature (20–30°C) and humidity (40–80%) conditions promoted the

297   stable oxidation of the diurnal photochemistry. The peak $j(O^1D)$ value was

298   approximately 2.0×10^{-5} s^{-1}, exhibiting the typical characteristics of intense light

299   radiation in autumn in the Pearl River Delta region (Yang et al., 2022a; Tan et al.,

300   2022).

301   As typical marine air components, the concentrations of NOx, CO, $PM_{2.5}$, and

302   other pollutants were lower than those detected in other observation campaigns in

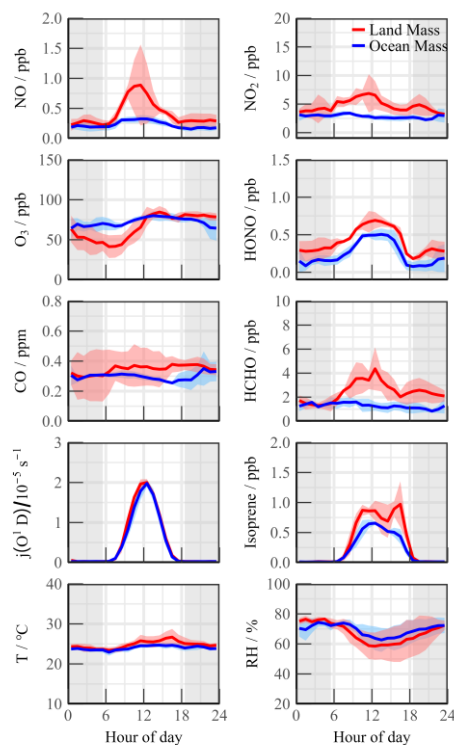
303   both urban and suburban areas in the Pearl River Delta region (Tan et al., 2019b; Lu et

304   al., 2012; Yang et al., 2022b). Serval observation campaigns have discovered the

305 relationship between wind direction and radical chemistry (Lu et al., 2012; Fuchs et

306 al., 2017; Niu et al., 2022). Although there was no apparent wind speed condition, the

307 dominant air mass still influenced the pollutant concentrations due to the particularity

308 of the marine site.

309
310 **Fig. 2.** Mean diurnal profiles of measured trace gases parameters during Land mass and Ocean mass episodes. The
311 coloured shadows denote the 25 and 75% percentiles. The grey areas denote nighttime.

312 During the OCM period, the NOx and HCHO concentrations exhibited relatively

313 clean characteristics that were consistent with those previously observations in open

314 ocean (RHaMBLe, SOS, CHABLIS and ALBATROSS, Table 1). Isoprene, a

315 representative BVOC, achieved a diurnal concentration of 0.58 ± 0.06 ppb, indicated

316 slightly local emissions could have impacted the concentrations of the precursor

317 species even in OCM sector. The ozone concentration in the YMK site was always at

318 the critical value of the updated Class I standard (GB3095-2012, average hourly O_3 of

319 81 ppb at 25°C and 1013 kPa). The occurrence of fewer emissions reduced the

320 titration effect, resulting in the ozone exhibiting no apparent diurnal trend on some of

321 the dates and a high background value at night (78.1 ± 7.6 ppb).

322 As a coastal site, chemical conditions could be influenced by local land emissions

323 depending on the wind direction. Compared with the OCM period, the meteorological

conditions (T, RH, and J-values) changed slightly during the LAM episode, but the pollutants were accumulated due to the transport of the plume from the northern cities (Fig. 2). The CO and $PM_{2.5}$ concentrations exhibited good consistency and even mild pollution features ((0.36 ± 0.12 ppm) and (37.70 ± 7.91 μg/m^3), respectively), reflecting the influence of human activities. Both NO and NO_2 peaked at around 10:00, exhibiting prominent pollution characteristics. HONO exhibited a distribution with high daytime (0.66 ± 0.08 ppb) and low nighttime (0.33 ± 0.09 ppb) concentrations. This unique distribution of HONO has been observed in remote environments in several previous observation campaigns (Jiang et al., 2022; Crilley et al., 2021). High HONO concentration in the daytime will affect the chemical composition of the atmosphere and the secondary pollution generation.

Table 1. Summary of radical concentrations and related species concentrations at MBL. All data are listed as the average in noontime (10:00~15:00).

Campaign	Location	Category	Date	OH (10^6cm^{-3})	HO$_2$ (10^8cm^{-3})	HCHO (ppb)	HONO (ppb)	NOx (ppb)	O$_3$ (ppb)	Reference
WAOSE95	Weybourne, UK	Coastal	1995 (Jun)	5.0	-	1.50	0.10	<2.0	40.0	(Grenfell et al., 1999)
ALBATROSS	Atlantic Ocean	Open ocean	1996 (Oct-Nov)	7.0	-	0.50	-	-	25.0	(Brauers et al., 2001)
EASE96	Mace Head, Ireland	Coastal	1996 (Jul-Aug)	2.3	2.6	-	-	~1.0	45.0	(Carslaw et al., 1999)
EASE97	Mace Head, Ireland	Coastal	1997 (Apr-May)	1.8	1.0	0.70	-	0.95	46.0	(Creasey et al., 2002)
ORION99	Okinawa Island, Japan	Coastal	1999 (Aug)	4.0	4.3	-	0.20	6.3	23.0	(Kanaya et al., 2001)
RISOTTO	Rishiri Island, Japan	Coastal	2000 (June)	7.4	3.1	-	-	0.45	-	(Kanaya et al., 2002)
RISFEX	Rishiri Island, Japan	Coastal	2003 (Aug)	2.7	1.5	-	-	0.2	28.0	(Qi et al., 2007)
CHABLIS	Antarctica	Open ocean	2005 (Jan-Feb)	1.0	1.1	0.12	0.007	0.02	7.0	(Bloss et al., 2010)
RHaMBLe	Atlantic Ocean	Open ocean	2007 (May-Jun)	9.0	6.0	0.30	-	0.014	35.0	(Whalley et al., 2010)
SOS	Cape Verde, Atlantic Ocean	Open ocean	2009 (Jun; Sep)	9.0	4.0	1.9	-	0.050	40.0	(Carpenter et al., 2011)
CYPHEX	Cyprus, Mediterranean	Coastal	2014 (Jul)	5.8	6.3	~1.0	~0.080	<1.0	69.0	(Mallik et al., 2018)
ICOZA (NW-SE)	North Norfolk, UK	Coastal	2015 (Jul)	3.0	1.4	0.9	0.052	2.0	39.0	(Woodward-Massey et al., 2022a)
ICOZA (SW)	North Norfolk, UK	Coastal	2015 (Jul)	4.1	1.0	1.1	0.097	3.0	31.0	(Woodward-Massey et al., 2022a)
HT	Hok Tsui, China	Coastal	2020 (Oct-Nov)	4.9	-	1.0	0.15	~4.0	65.0	(Zou et al., 2022)
YMK (Land Mass)	Shenzhen, China	Coastal	2019 (Oct)	7.1	5.2	3.4	0.66	6.4	75.6	This work

YMK (Ocean Mass)	Shenzhen, China	Coastal	2019 (Oct)	4.5	4.9	1.2	0.48	3.0	78.1	This work

337

The detailed information for VOCs species during the YMK campaign has been added in the Table S5. The daily maximum NMHC concentration peaked at 27.81 ± 9.91 ppb, and the maximum value of ~40 ppb occurred on October 27. Local biological emissions significantly affected the NMHC composition of the site, and isoprene achieved a noon maximum of 0.82 ± 0.16 ppb. Neither anthropogenic alkenes (2.21 ± 0.94 ppb) nor aromatic (1.31 ± 0.25 ppb) hydrocarbons were abundant, and OVOCs accounted for approximately 50% of the total. As a photochemical indicator, formaldehyde peaked at ~4 to ~8 ppb during the LAM episode, suggesting a more vigorous oxidation process. The HONO concentration was 6.8 times higher than the SW scenario in the ICOZA observation (a pollution period dominated by a southwest wind direction), while the HCHO concentration was 3.1 times higher. (Woodward-Massey et al., 2022a). The abundance of oxidation precursors (HONO, HCHO, O_3, and NMHCs) reflected the unique atmospheric conditions in the marine environment in China, which originated from the complex atmospheric pollution.

3.2 HOx radical concentrations and modeled OH reactivity

Fig. 3(a)(b) shows the time series of the simulated and observed OH and HO_2 radical concentrations during the observation campaign. The time series of the simulated OH reactivity (k_{OH}) is presented in Fig. 3(c). The observed OH and HO_2 radicals exhibited significant diurnal trends. The daily maximum OH and HO_2 values were $(4.7–9.5) \times 10^6 \, \mathrm{cm}^{-3}$ and $(4.2–8.1) \times 10^8 \, \mathrm{cm}^{-3}$, respectively. The peak k_{OH} value was commonly less than 10 s^{-1}. Due to human activities, the simulated k_{OH} reached more than ~15 s^{-1} on some days. The radical concentrations and reactivity exhibited similar trends, which differed from reports on urban and semi-urban areas where inorganic species (NOx and CO) were the dominant controllers of k_{OH} (Zhang et al., 2022a; Tan et al., 2019b; Lou et al., 2010). The k_{OVOCs} was separated into $k_{OVOCs(Obs)}$ and $k_{OVOCs(Model)}$ (Fig. 3(c)). Specifically, $k_{OVOCs(Obs)}$ includes the observed species such as formaldehyde (HCHO), acetaldehyde (ACD), higher aldehydes (ALD),

365 acetone (ACT), ketones (KET), and oxidation products of isoprene (MACR and

366 MVK). The model-generated intermediates, such as glyoxal, methylglyoxal,

367 methylethyl ketone, and methanol, are categorized as $k_{OVOCs(Model)}$. Approximately 50%

368 of the total k_{OVOCs} are represented by unconstrained species ($k_{OVOCs(Model)}$), which

369 contribute a daily k_{OH} of 1.39 s^{-1}. Overall, the observed OH and HO$_2$ concentrations

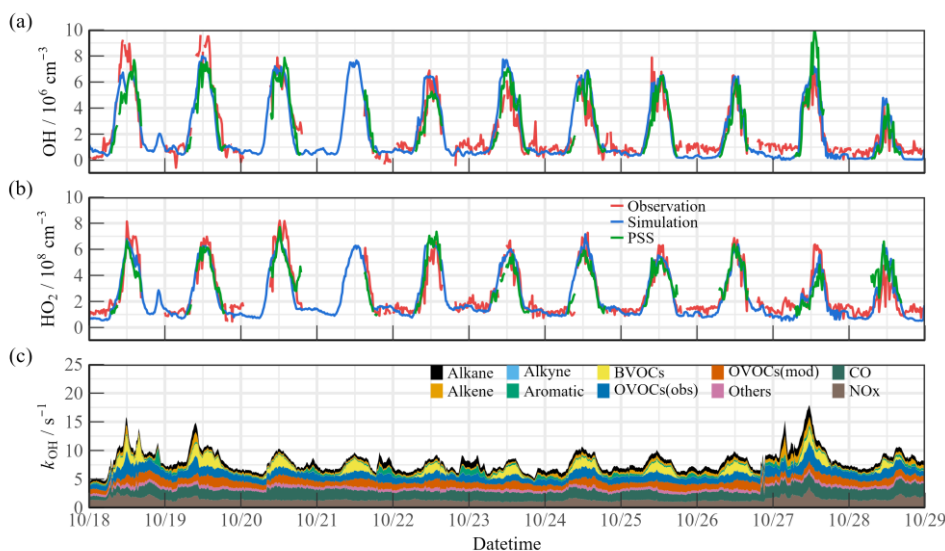
370 were both well reproduced by the base model incorporating the RACM2-LIM1

371 mechanism. The observed OH was underestimated only on the first days, and a slight

372 model overestimation happened on October 23&24. PSS calculation showed good

373 agreement with the base model, providing evidence of the balance of radical internal

374 consistency in the daytime. It should be noted that the OH reactivity of unmeasured

375 VOCs may be underestimated due to the lumped groups in RACM2 mechanism.

376
377 **Fig. 3.** Timeseries of the observed and modelled parameters for OH, HO$_2$ and k_{OH} during the observation
378 period. **(a)** OH, **(b)** HO$_2$, **(c)** k_{OH}.

379   The air mass transport of the precursors induced photochemistry accumulation,

380 which was then reflected in the changes in the oxidation progress. It is worth

381 comparing the concentrations and reactivities of the radicals by classifying the

382 predominant air mass (Fig. 4). During the OCM period, the observed OH and HO$_2$

383 radicals could be reflected by the base chemical mechanism, with daily averages of

384 4.5×10^6 cm^{-3} and 4.9×10^8 cm^{-3}, respectively. Compared to other campaigns (Table

385 1), the observed maximum values were within reasonable ranges (OH: $2 - 9 \times 10^6$

386 cm^{-3}; HO$_2$: $1 - 6 \times 10^8$ cm^{-3}). Despite low NOx levels during the OCM period, the

387 HO$_2$ radical was not overestimated using the base model, which was dissimilar to

388 many MBL observations (Bloss et al., 2010). However, both the OH and HO$_2$ radical

389 concentrations reached higher levels during the LAM-dominant period, indicating a

390 more active photochemical process (Section 4.1). The diel averages for the OH and

391 HO$_2$ radicals were 7.1×10^6 cm^{-3} and 5.2×10^8 cm^{-3}, respectively, which were

392 notably higher than the levels reported in the ICOZA observations (Woodward-

393 Massey et al., 2022a). The base scenario underestimated both the OH and HO$_2$

394 concentrations between 10:00 and 15:00, and the observation-to-model ratio was

395 greater than 1.2. The calculated daily maximum total OH reactivity was 8.8 s^{-1}, and

396 nearly 70% of the reactivity was accounted for by the organic species, among which

397 the OVOCs were the largest contributor (30.6%). The anthropogenic alkanes, alkenes,

398 and aromatic hydrocarbons contributed less than 10% to the reactivity. Compared

399 with the OCM-dominant episode, the higher reactivity during the LAM period

400 indicated the occurrence of efficient recycling during the ROx (the sum of OH, HO$_2$,

401 and RO$_2$) propagation (12.4 s^{-1} vs. 8.8 s^{-1}). The higher contributions of the BVOCs

402 (only isoprene was considered, 15.6%) and OVOCs (30.2%) to the reactivity reflected

403 the diverse composition of the VOCs in the forest environment. Under enhanced

404 photochemistry, the calculated OH reactivity could be an underestimation of the total

405 OH reactivity, so a missing OH source may be masked. As a representative of the

406 OVOCs, HCHO reflects the photochemical level to a certain extent. As shown in Fig.

407 S5, a solid positive dependence between the OH$_{obs}$-to-OH$_{mod}$ ratio and HCHO was

408 observed (the daytime data were restricted according to j(O^{1}D) $> 5 \times 10^{-6}$ s^{-1}). With

409 the increase of photochemical intensity, the ratio between the observed and simulated

410 OH radical showed an obvious mismatch. Obtaining the full magnitude of the radical-

411 related parameters is necessary to compensate for the discrepancy in the concentration

412 closure experiments.

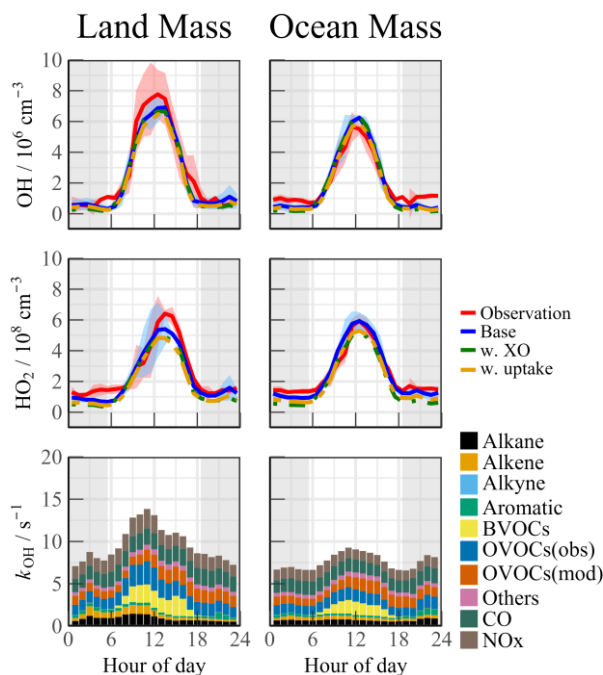



Fig. 4. Median diurnal profiles of the observed and modelled OH, HO$_2$, k_{OH} during LAM and OCM episodes. The coloured shadows for OH and HO$_2$ radicals denote the 25 and 75% percentiles. The grey areas denote nighttime.

Halogen species have been recognized as potent oxidizers that can boost photochemistry (Xia et al., 2022; Peng et al., 2021). A sensitivity test was performed by imposing BrO and IO into the base model to diagnose the impact of the halogen chemistry on the troposphere chemistry. The concentration of BrO and IO is set to ~5 ppt, which is a typical level in MBL site (Xia et al., 2022; Bloss et al., 2010; Whalley et al., 2010). The details of the mechanisms involved are listed in Tables S3 and S4. In this scenario (Fig. 4, green line). The daytime concentration of HO$_2$ radical decreased by 8.5% and 13.3% during the LAM and OCM periods, respectively, compared to the base model. However, there was no significant change in the concentration of OH radicals (<3%). Traditionally, it is believed that the inclusion of halogen chemistry leads to higher modeled OH concentrations and lower modeled HO$_2$ concentrations. Therefore, the lack of an increase in OH concentration with the introduction of the halogen mechanism at the YMK site calls for further investigation (Fig. S6). By modifying the NO concentration in different levels (Scenario 1: [NO] $\times$ 150%, Scenario 2: base, Scenario 3: [NO]$\times$20%, Scenario 4: [NO]$\times$10%), the response of HOx radicals to the halogen mechanism varied under different NO levels. As the

constrained NO increased from 30 ppt to 500 ppt, the reduction in HO$_2$ radicals due to the Br and I mechanisms ranged between 10% and 20%. At elevated NOx levels, reactions between halogen radicals and NOx occurred, inhibiting the formation of OH radicals. In Scenario 1, the OH concentration even decreased by 3.5% when introducing the halogen mechanism. When NO concentration was constrained around 30 ppt (Scenario 4), similar to those obtained in RHaMBLe/CYPHEX campaigns, the modelled OH concentration increased by 14.4%, while the HO$_2$ concentration decreased by approximately 20.8% (Whalley et al., 2010; Bloss et al., 2010). Therefore, the sensitivity of OH radicals to the halogen mechanism in the YMK region is primarily limited by the local NOx concentration level.

Alough the modelled and measured HO$_2$ showed good agreement,the effect of HO$_2$ heterogeneous processes on the chemistry of HOx radicals is also worth exploring. The inclusion of heterogeneous processes ($\gamma = 0.08$) did reduce the modelled HO$_2$ concentration for ~10% during both LAM and OCM periods (Fig. 4, yellow line). This reduced agreement between observation and simulation emphasizes the presence of a missing HO$_2$ source in the base model.

4 Discussion

4.1 Experimental radical budget balance

4.1.1 OH radical

A process-oriented experiment was conducted to investigate the photochemistry progress from a budget balance perspective (Woodward-Massey et al., 2022b; Tan et al., 2019b; Yang et al., 2021a). The OH was in a photostationary steady state due to its short lifetime. The total OH removal rate was directly quantified from the union of the OH concentration and the reactivity (Eq.(6)):

$$D(OH) = [OH] \times k_{OH}. \tag{6}$$

The total production rate of the OH radical was the sum of the primary sources (O$_3$/HONO photolysis and ozonolysis reactions) and secondary sources (HO$_2$ + NO) (Eq.(7)):

$$461 \qquad P(OH) = j_{HONO}[HONO] + \varphi_{OH}j(O^1D)[O_3] + \Sigma i \left\{ \varphi_{OH}^i k_{Alkenes+O_3}^i [Alkenes][O_3] \right\}$$

$$462 \qquad + (k_{HO_2+NO}[NO] + k_{HO_2+O_3}[O_3])[HO_2] . \tag{7}$$

Here, φ_{OH} and φ_{OH}^i represent the OH yields in the O_3 photolysis and alkene ozonolysis processes, respectively.

The diel profiles of the experimental OH budget during the LAM and OCM periods are shown in Fig. 5. Both the observed OH and HO_2 radicals were introduced into the budget calculations. Because k_{OH} was not measured during the observation experiment, the simulated value was used to analyze the removal rate. Therefore, D(OH) should be considered a lower limit as it uses calculated rather than measured k_{OH} (Yang et al., 2022b). During the OCM period, the HO_2 + NO reaction accounted for ~50% of the OH yield. The maximum of 6.6 ppb/h occurred at around 12:00. The photolysis reactions could increase the daytime contributions of HONO and O_3 to 1.52 ppb/h and 0.84 ppb/h, respectively (10:00–15:00). The contribution of the non-photolytic radical source (ozonolysis reactions) was almost negligible.

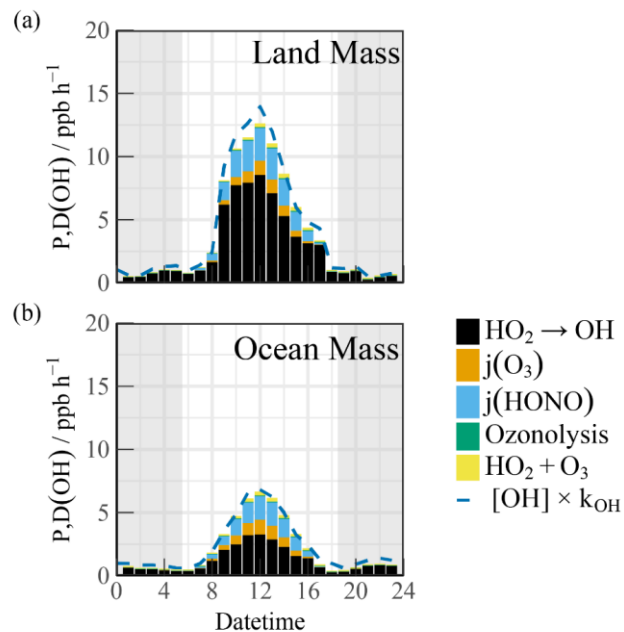



Fig. 5. The diurnal profiles of the experimental OH budget during **(a)** LAM and **(b)** OCM episodes. The blue line denotes the OH destruction rate([OH]×k_{OH}). The grey areas denote nighttime.

Compared with other marine observations, the calculated OH generation rate was approximately twice that reported in the ICOZA and five times that obtained in the RHaMBLe campaigns (Woodward-Massey et al., 2022b; Whalley et al., 2010).

During the LAM period, the OH generation rate reached a maximum of 12.6 ppb/h, accompanied by a secondary source contribution of 67% (from the reaction between HO_2 and NO) during the daytime, which was close to several observations related to polluted plumes (Woodward-Massey et al., 2022b; Tan et al., 2019b; Lu et al., 2012; Yang et al., 2022b). When the simulated k_{OH} was introduced into the experimental budgets, the difference between P(OH) and D(OH) was less than 2 ppb/h.

4.1.2 Total ROx radicals

The budget analysis of the HO_2 and RO_2 radicals could not be performed well due to the lack of RO_2 radical observation data. The diurnal profiles of the ROx production (P(ROx)) and termination rate (L(ROx)) for the different air masses are shown in Fig. 6. The P(ROx) could reach 3.36 ppb /h with an ocean plume. HONO photolysis controlled nearly half of the primary sources (45.7%), and the daily distribution was consistent with that of solar radiation. The ozone-related contributions from photolysis and ozonolysis were approximately 46.6% (25.1% from photolysis and 11.5% from ozonolysis,respectively). The remaining contribution was from the photolysis of carbonyls (HCHO and OVOCs) (15.0%). The anthropogenic contribution to the radical chemistry was not ignorable, and the ROx source in this observation was exponentially higher than that in other MBL observations (Woodward-Massey et al., 2022b; Stone et al., 2012; Whalley et al., 2010; Mallik et al., 2018). The P(ROx) of the LAM was close to that in Shenzhen (~4 ppb/h) but was significantly lower than that in Yufa (~7 ppb/h) and the BackGarden (~11 ppb/h) (Tan et al., 2019b; Lu et al., 2012; Yang et al., 2022b). The reactions between ROx and NOx and self-combination were the main pathways of radical termination (~70%). The contribution of the formation of peroxynitrate to the L(ROx) could not be ignored in the daytime.

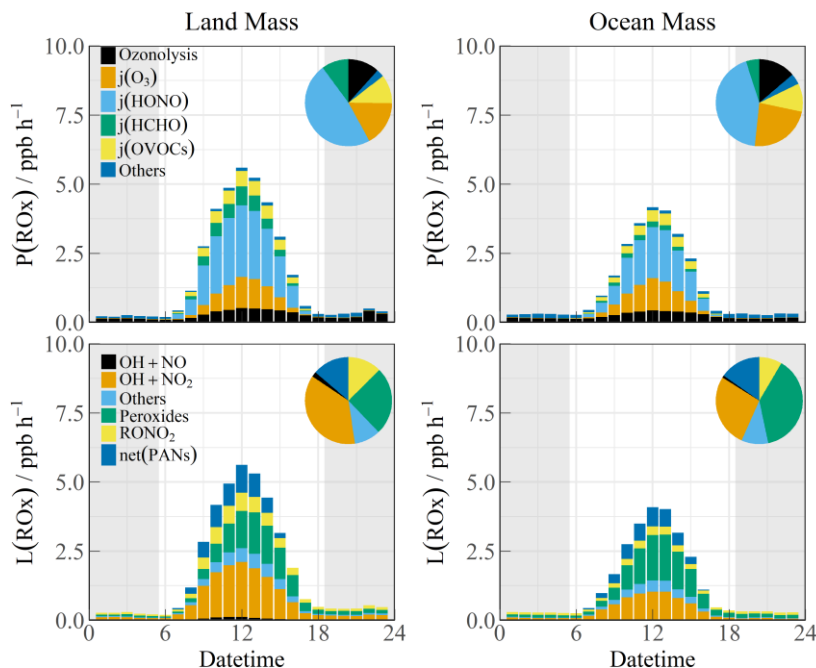



Fig. 6. The diurnal profiles of ROx budget during Land mass and Ocean mass episodes. The pie chart denotes proportions in different parts during the daytime (10:00-15:00). The grey areas denote nighttime.

The high daytime HONO concentrations observed at the YMK site is a notable phenomenon. Due to the high HONO concentration during the daytime, the photolysis reaction made daytime contributions of 1.52 ppb/h and 2.19 ppb/h during the OCM and LAM periods, respectively. As the only known gas-phase source, OH + NO accounted for a negligible proportion of the HONO production rate. Given the location of the site, HONO from cruise ship emissions is a possible component of the primary anthropogenic source (Sun et al., 2020). Other active tropospheric HONO sources (heterogeneous reactions with NO_2 and $p(NO_3^-)$ photolysis) are worthy of consideration and significantly contribute to the atmospheric oxidation in the MBL area (Zhu et al., 2022; Crilley et al., 2021).

4.2 Local ozone production rate

Peroxy radical chemistry is the essential photochemical source of tropospheric ozone (F(Ox), Eq.(8)):

$$F(O_x) = k_{HO_2+NO}[NO][HO_2] + \sum_i (1 - \alpha_i)k_{RO_2^i+NO}[NO]RO_2^i \tag{8}$$

NO reacts with HO_2 and RO_2 radicals to form NO_2, and then, photolysis occurs to form O_3 under solar radiation. NO_2 and ozone are the two sides of the oxidation reservoir. The effect of local emissions on the photodynamic equilibrium can be avoided by characterizing the photochemical production of the total oxidants (Tan et

al., 2019b). α_i represents the side generation ratio of organic nitrate, which also affects the quantum yield of NO_2 (Tan et al., 2018). Ox is mainly photochemically removed through ozone photolysis, ozonolysis, radical chain propagation (OH/HO_2 + O_3), and chain termination (OH + NO_2) reactions in the troposphere (D(Ox), Eq.(9)):

$$D(O_x) = \varphi_{OH} j(O^1 D)[O_3] + \Sigma i \left\{ k^i_{Alkenes+O_3}[Alkenes][O_3] \right\} + (k_{O_3+OH}[OH] +$$

$$k_{O_3+HO_2}[HO_2])[O_3] + k_{OH+NO_2}[OH][NO_2] \quad (9)$$

The net formation rate P(O_x) can be calculated by subtracting D(Ox) from F(Ox) (Eq.(10)):

$$P(O_x) = F(O_x) - D(O_x) \quad (10)$$

The simulated RO_2 radical concentration was introduced into the F(Ox) calculation. The diurnal variations in the ozone generation in the different air masses are shown in Fig. 7. The contribution of the HO_2 radical to F(Ox) was approximately 60%. The RO_2 radicals consisted of various types such as methyl peroxy (MO_2), acetyl peroxy radicals (ACO_3/RCO_3), and other radicals derived from alkanes (ALKAP), alkenes (ALKEP), and isoprene (ISOP), which accounted for an additional 40% of the F(Ox). On a daytime basis, the maximum F(Ox) reached 7.4 ppb/h at around 12:00 in the OCM period, while a persistent-high value (maximum of 12.5 ppb/h at 10:00–14:00) occurred in the LAM period. A vast amount of Ox was consumed in the nitric acid (OH + NO_2) formation pathways, i.e., higher than the ozonolysis removal. The daily averaged ozone production rates were 5.52 and 2.76 ppb/h during the LAM and OCM periods, respectively.

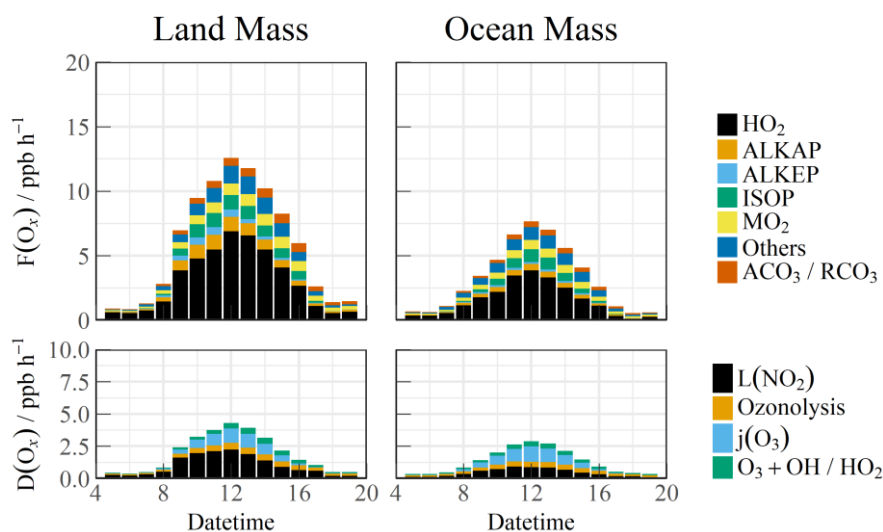



Fig. 7. The diurnal profiles of the speciation F(O_X) and D(O_x) during Land mass and Ocean mass episodes. The data were calculated by the measured OH and HO_2 and modelled RO_2 radicals.

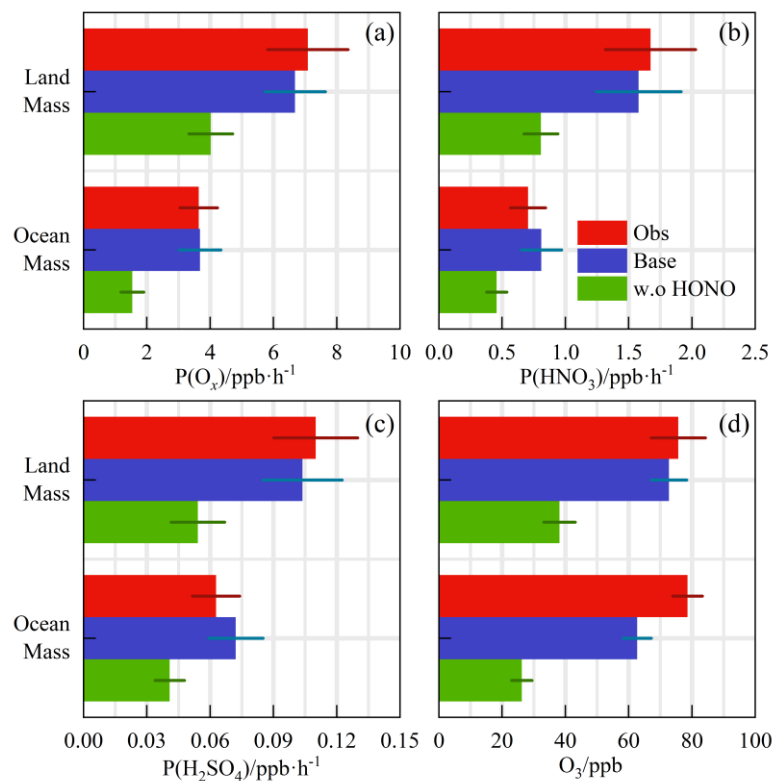
4.3 Relationship between precursors and oxidation rates

Despite the low level of human activities, oxidation precursors have an extended lifetime in the stable atmosphere of coastal areas. Intensive photochemical reactions occur after the accumulation of precursors, resulting in local net ozone production comparable to that in the surrounding suburban environments (Zeren et al., 2022). Simultaneous observations of both urban and coastal settings in Shenzhen have indicated that the oxidation rates are comparable (Xia et al., 2021). The coupling of precursor transport and local photochemical processes in marine areas makes it meaningful to explore secondary pollution generation (Fig. 8(a), (b), and (c)). No obvious radical source was missing during the LAM and OCM periods, and the oxidation level was that expected from the base model. On a daytime basis, the mean diurnal profile of the P(Ox) reached ~7 ppb/h in the LAM period, and the average nitric acid (P(HNO$_3$)) and sulfuric acid (P(H$_2$SO$_4$)) production rates were ~1.6 and ~0.11 ppb/h, respectively. The P(HNO$_3$) production rate was similar to the average of observations in the Pearl River Delta region (~1.3 ppb/h), while that of the P(H$_2$SO$_4$) was only half the average level (~0.24 ppb/h) (Lu et al., 2013; Tan et al., 2019b; Yang et al., 2022b). During the OCM period, the characteristics of the ocean air mass alleviated the photochemical process, and the production rates of the secondary pollutants decreased by approximately half and were close to the average levels in winter (Ma et al., 2019).

Contrary to numerous ocean observations, in the YMK site, intensive oxidation was accompanied by a high diurnal HONO level (higher than 400 ppt) (Fig. 9). The ozone levels were consistent with the Grade I air quality standard and far exceeded the global background concentration (~40 ppb). Daytime photolysis reactions of HONO contributed 1.52 ppb/h and 2.19 ppb/h to P(ROx) during the OCM and LAM periods, respectively, which were much higher than the values in several megacities during the photochemically polluted season (Tan et al., 2019a). Given the significance of HONO photolysis in driving atmospheric chemistry, a sensitivity test was conducted without constraints on HONO (i.e., w.o HONO) to specifically quantify the

580 contribution of HONO-induced secondary pollution. Only the homogeneous reaction

581 (OH + NO) participated in the formation of HONO in the default mode without

582 HONO input (Liu et al., 2022b).

583

584 **Fig. 8.** The calculated reaction rates based on the observed concentrations for Land mass and Ocean mass
585 episodes **(a)** P(Ox), **(b)** P(HNO$_3$), **(c)** P(H$_2$SO$_4$). **(d)** The observed and modelled O$_3$ concentration with a first-order
586 loss term. The deposition process was equivalent to a lifetime of 15 hours to all species. All the rates and
587 concentration are averaged for the daytime period between 10:00 and 15:00.

588 The modelled OH, HO$_2$ and RO$_2$ change when the model was unconstrained to

589 HONO were shown in Fig. S7. After evaluation, in LAM and OCM sectors,

590 concentration changes for OH were 46.9% and 43.2%, for HO$_2$ were 38.3% and

591 34.3%, for RO$_2$ were 43.7% and 39.0%, respectively. The P(Ox) was found to be 33%

592 and 39% lower during the LAM and OCM periods, respectively. The nitric acid

593 (P(HNO$_3$)) and sulfuric acid (P(H$_2$SO$_4$)) formation rates also increased

594 simultaneously (~43% and ~48% for LAM and OCM sectors, respectively). The

595 sensitivity test identified the privileged role of the HONO-related mechanisms in the

596 OH chemistry, which resulted in a correlation between the efficient radical recycling

597 and secondary pollution.

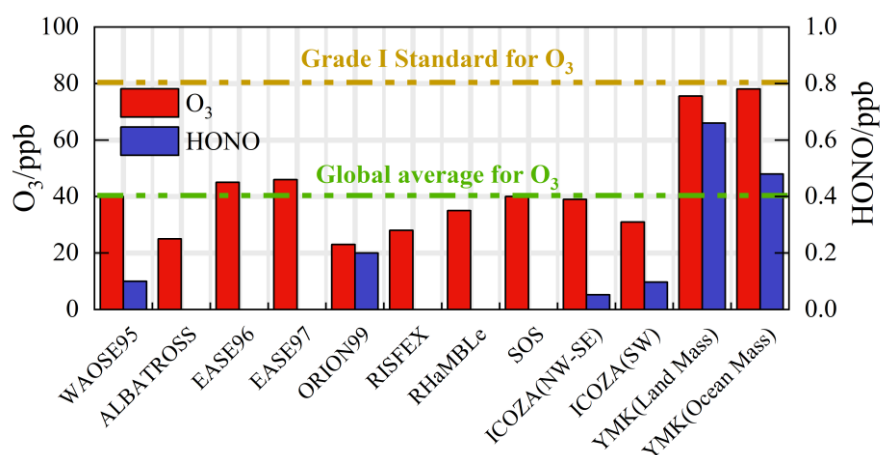



Fig. 9. Summary of both ozone and HONO concentrations in previous marine observations. The concentrations are averaged for the daytime period between 10:00 and 15:00.

A time-dependent box model was used to test the association between the HONO chemistry and the local ozone generation (Fig. 8(d)). In order to isolate the O_3 photochemical production, the impacts of vertical entrainment and horizontal advection were in general ignored. On the basis of the base scenario run, constraints of the observed ozone and NO concentrations were removed to predict ozone. The observed and modelled O_3 concentrations in Fig. 8(d) are averaged for the daytime period between 10:00 and 15:00. The observed diurnal ozone concentrations were 75.7 ppb and 78.6 ppb during the LAM and OCM periods, respectively. The daytime ozone was well reproduced by the time-dependent box model, and the deviation of the simulation was less than 20% (Fig. 8(d)). After removing the HONO constraint, the simulated ozone concentrations were 38.2 and 26.3 ppb, i.e., 48% and 58% lower, during the LAM and OCM periods, respectively. Simulated O_3 decreased from ~75 ppb to a global background, and daytime HONO concentration were reduced to a low level (~70 ppt) (Woodward-Massey et al., 2022a; Zhu et al., 2022; Xia et al., 2022). The elevated daytime HONO had an additional effect on the oxidation in the background atmosphere. For coastal cities, the particularity of the HONO chemistry in the MBL tends to influence the ozone-sensitive system and eventually magnifies the ozone background. Therefore, the promotion of oxidation by elevated precursor concentrations is worth considering when formulating emission reduction policies. In regions where HONO concentrations are elevated, the sources of HONO would need to be identified to aid pollution mitigation policies.

5 Conclusions

Comprehensive observations of HOx radicals and other relevant species were conducted in October 2019 at a coastal site in the Pearl River Delta (the YMK site, 22.55°N, 114.60°E). The overall air pollutants exhibited typical coastal features due to the scarce anthropogenic emissions. The daily maximum OH and HO_2 concentrations were $(4.7–9.5) \times 10^6$ cm^{-3} and $(4.2–8.1) \times 10^8$ cm^{-3}, respectively. The base RACM2-LIM1 model satisfactorily reproduced both the observed OH and HO_2 radical concentrations, but a slight overestimation of the OH radical occurred. The daily maximum calculated total OH reactivity was 9.9 s^{-1}, and nearly 70% of the reactivity was contributed by organic species.

In addition to anthropogenic and vegetation emissions, the synchronized air mass transport from the northern cities and the South China Sea exerted a time-varying influence on radical photochemistry and atmospheric oxidation. During the OCM period, the observed OH and HO_2 radical concentrations could be reflected by the base chemical mechanism, with daily average values of 4.5×10^6 cm^{-3} and 4.9×10^8 cm^{-3}, respectively.

In the episode that was dominated by ocean mass, the HO_2 + NO reaction accounted for ~50% of the primary OH yield. A higher OH generation rate was found(12.6 ppb/h) during the LAM period, and the secondary source accounted for 67% of the total, which was similar to several observations in polluted plumes. Reactions between ROx and NOx and self-combination were the main pathways of radical termination (~70%), and the contribution of peroxynitrate formation to the L(ROx) could not be ignored in the daytime.

Intensive photochemical reactions occur after the accumulation of precursors, resulting in local net ozone production comparable to that in the surrounding suburban environments. The daily average ozone production rates were 5.52 and 2.76 ppb/h in the LAM and OCM periods, respectively. The rapid oxidation process was accompanied by a higher diurnal HONO concentration (higher than 400 ppt). A non-HONO-constrained sensitivity test was performed to quantify the HONO-induced

651 contribution to secondary pollution. After evaluation, the P(Ox) values were 33% and
652 39% lower during the LAM and OCM periods, respectively. The nitric acid (P(HNO$_3$))
653 and sulfuric acid (P(H$_2$SO$_4$)) formation rates also increased simultaneously (~43%
654 and ~48% for LAM and OCM sectors, respectively). Simulated O$_3$ decreased from
655 ~75 ppb to a global background, and daytime HONO concentration were reduced to a
656 low level (~70 ppt). For coastal cities, the particularity of the HONO chemistry in the
657 MBL tends to influence the ozone-sensitive system and eventually magnifies the
658 ozone background. Therefore, the promotion of oxidation by elevated precursor
659 concentrations is worth considering when formulating emission reduction policies.

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Data availability

666 The data used in this study are available from the corresponding author upon request
667 (rzhu@aiofm.ac.cn).

Author contributions

669 WQ Liu, PH Xie, RZ Hu contributed to the conception of this study. GX Zhang and
670 RZ Hu performed the data analyses and manuscript writing. All authors contributed to
671 measurements, discussed results, and commented on the paper.

Competing interests

673 The contact author has declared that none of the authors has any competing interests.
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