



## Assessment of the impacts of cloud chemistry on surface SO<sub>2</sub> and sulfate levels in typical regions of China

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### Abstract

A regional online chemical weather model WRF/ CUACE (China Meteorological Administration Unified Atmospheric Chemistry Environment) is used to assess the contributions of cloud chemistry to the SO<sub>2</sub> and sulfate levels in typical regions in China. By comparing with several time series of in-situ cloud chemical observations on Mountain Tai in Shandong Province of China, the CUACE cloud chemistry scheme is found to well reproduce the cloud processing the consumption of H<sub>2</sub>O<sub>2</sub>, O<sub>3</sub> and SO<sub>2</sub> and the increase of sulfate, and consequently is used in the regional assessment for a heavy pollution episode and monthly average in December 2016. During cloud availability in heavy pollution episode, the sulfate production increases 60-95% and SO<sub>2</sub> reduces over 80%. And the cloud chemistry mainly affects the middle and lower troposphere below 5 km as well as within the boundary layer, and contributes significantly to SO<sub>2</sub> reduction and sulfate increase in east-central China. Among the four typical contaminated regions in China, the Sichuan Basin (SCB) is mostly affected by the cloud chemistry, with the average SO<sub>2</sub> abatement about 1-15 ppb and sulfate increase about 10-70 µg/m<sup>3</sup>, followed by Yangtze River Delta (YRD) where SO<sub>2</sub> abatement is about 1-3 ppb and sulfate increase is about 10-30 µg/m<sup>3</sup>. However, the cloud chemistry contribution to Pearl River Delta (PRD) and North China Plain (NCP) are not significant and weaker than other two regions due to lighter pollution and less water vapor, respectively. In addition, the average contribution of cloud chemistry during the pollution period is distinctly greater than that for all December. This study provides a way to analyze the

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over-estimate phenomenon of SO<sub>2</sub> in many chemical transport models.

35 Keywords: SO<sub>2</sub>, sulfate, cloud chemistry, WRF/CUACE

## 1 Introduction

Aerosols interact with radiation and clouds, directly or indirectly affecting the atmospheric radiation balance and precipitation, which in turn affects weather and climate (Twomey et al., 1984; Twomey, 1991; Charlson et al., 1992; Ramanathan et al., 2001; Pye et al., 2020). Moreover, large amounts of aerosols dispersed in the atmosphere can reduce  
40 visibility and deteriorate air quality (Molina, 2002), which is harmful to human health and ecosystem (Xie et al., 2019; Sielski et al., 2021).

In addition to direct emissions, aerosols are mostly produced secondarily through the oxidation of precursor gases, and one of the important processes is the transformation in clouds. Global cloud coverage of about 21% to 95% provides an adequate environment for cloud chemistry processes (Kotarba, 2020; Ravishankara, 1997). As about 90% of the clouds  
45 formed in the atmosphere evaporate without deposition or forming the precipitation, large fractions of aerosols formed in them can then re-enter the atmosphere (Caffrey et al., 2001; Harris et al., 2013; Lelieveld et al., 1992). Globally, sulfate production from SO<sub>2</sub> oxidation accounts for about 80% of total sulfate, and more than half of it is produced in clouds (Hung et al., 2018; Faloon et al., 2010; Guo et al., 2012). Ge et al. (2021) found that cloud chemistry processes reduced SO<sub>2</sub> concentrations by 0 ~ 50% in most of east-central China in all seasons. Li (2011) found that the average SO<sub>4</sub><sup>2-</sup> concentration  
50 in cloud water accounted for 53.8% of the total aerosol concentration at a Mount site. Li et al. (2020) also found that cloud processes effectively reduced atmospheric O<sub>3</sub> and SO<sub>2</sub> concentrations by an average of 19.7% and 71.2%, respectively, at Mount Tai.

Multiphase oxidation of sulfate aerosols from SO<sub>2</sub> in aerosol particles in high humidity environment is one of the main causes of explosive growth of particulate matter in East Asia haze (Guo et al., 2014; Cheng et al., 2016; Song et al., 2019).  
55 From observations and laboratory works, there are four main pathways for this kind of oxidation of SO<sub>2</sub> which are H<sub>2</sub>O<sub>2</sub>, O<sub>3</sub>, NO<sub>2</sub>, and transition metal ions (TMIs) (Iibusuki and Takeuchi, 1987; Martin et al., 1991; Alexander et al., 2009; Harris et al., 2013; Cheng et al., 2016; Wang et al., 2016). Additional pathways of organic peroxides (ROOH) (Yao et al., 2019; Wang et al., 2019; Ye et al., 2018; Dovrou et al., 2019, photolysis products of nitrate (pNO<sub>3</sub><sup>-</sup>) (Gen et al., 2019a; 2019b), and excited triplet states of photosensitizer molecules (T\*) (Wang et al., 2020) have also been found recently to be important for  
60 multiphase oxidation of sulfur dioxide during very heavy hazy days. Unfortunately there are still much uncertainties and gaps



to put all of those pathways into model applications from observational and laboratory studies (Pye et al., 2020; Ravishankara, 1997; Liu et al., 2021). Several regional and global models have tried to include only two, O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub>, in-cloud oxidant in cloud chemistry mechanisms (Park et al., 2004; Tie et al., 2005; Salzen et al., 2000; Chapman et al., 2009; Leighton and Ivanova, 2008). A very few models can simulate the pathway of NO<sub>2</sub>, TMI<sub>s</sub> of Fe or Mn ions (Ge et al., 2021; Binkowski and Roselle, 2003; Chang et al., 1987; Terrenoire et al., 2015; Menut et al., 2013).

There has been very serious air pollution in central-east China where four heavy pollution zones of North China Plain (NCP), Yangtze River Delta (YRD), Sichuan Basin (SCB) and Pearl River Delta (PRD) are located (Yao et al., 2021; Zhang et al., 2012). Although many global and regional models have contained sulfate formation mechanisms in cloud chemistry, few models have assessed its contribution, especially the lack of detailed assessment of regional cloud chemistry on sulfate and SO<sub>2</sub> in China and those four typical pollution regions. Regional chemical models have reported the over-estimate of SO<sub>2</sub> (Buchard et al., 2014; He et al., 2015; Wei et al., 2019; Sha et al., 2019; Georgiou et al., 2018). The inadequate inclusion or lack of cloud chemistry of SO<sub>2</sub> consumption simulations is one of the main causes (Ge et al., 2021). Therefore, there is a very important need to quantify the contribution of cloud chemistry in these typical regions in central-east China when it comes to heavily polluted weather to get a better understand of multi-dimensional pollution interactions, especially between the upper layer and the surface.

This study is intended to use an on-line coupled chemical weather platform of CMA, WRF/CUACE, to analyze and evaluate the SO<sub>2</sub> in-cloud oxidation process in the four pollution regions in China, with two objectives: (1) evaluating the cloud chemistry scheme in WRF/CUACE by the in-situ cloud chemistry observations at Mount Tai in summers of 2015 and 2018; and (2) quantifying the contributions of cloud chemistry to the SO<sub>2</sub> and sulfate changes in a typical winter pollution month of December 2016 with a very long lasting heavy pollution episode. It is aimed to establish a system to assess the relative contribution of cloud chemistry to SO<sub>2</sub> oxidation pathways and sulfate productions to other clear-sky processes.

## 2 Model description and Methodology

### 2.1 Cloud chemistry in WRF/CUACE

WRF/CUACE is an on-line coupled chemical transport model under the WRF frame work with a comprehensive chemical module - CUACE, which is developed at CMA with a sectional aerosol physics, gas chemistry, aerosol-cloud interactions and thermodynamic equilibrium (Zhou et al., 2012; Zhou et al., 2016; Gong et al., 2003; Gong and Zhang, 2008; Zhang et al., 2021). There are seven types of aerosols, i.e. black carbon, organic carbon, sulfate, nitrate, ammonium, soil dust, and sea salt, and more than 60 gaseous species. The aerosol size spectrum is divided into 12 bins with fixed boundaries of



0.005-0.01, 0.01-0.02, 0.02-0.04, 0.04-0.08, 0.08-0.16, 0.16-0.32, 0.32-0.64, 0.64-1.28, 1.28-2.56, 2.56-5.12, 5.12-10.24 and  
 90 10.24-20.48  $\mu\text{m}$ . The system can simulate  $\text{PM}_{10}$ ,  $\text{PM}_{2.5}$ ,  $\text{O}_3$  and visibility. A complete heterogeneous chemistry module has  
 been built in CUAC for nine gas-to-particle heterogeneous reactions including  $\text{SO}_2$  to sulfate in CUACE (Zhou et al., 2021,  
 Zhang et al., 2021). The cloud chemistry mechanism in CUACE considers the pathways of multiphase oxidation of  $\text{SO}_2$  by  
 $\text{H}_2\text{O}_2$  and  $\text{O}_3$  in both stratocumulus and convective clouds (Gong et al., 2003; Von Salzen et al., 2000). The transport and  
 chemical effects of sulfur in convective clouds are calculated based on a convective cloud model by WRF. Within the cloudy  
 95 part of a grid box, the first-order rate constant (in  $\text{s}^{-1}$ ) of S(IV) oxidation is given by the following expression:

$$F = \left| \frac{1}{C_{\text{S(IV)}}} \frac{dC_{\text{S(IV)}}}{dt} \right| = F_1 C_{\text{O}_3} + F_2 C_{\text{H}_2\text{O}_2} \quad (1)$$

where  $C_{\text{S(IV)}}$  is the total concentration of S(IV) (gas phase plus dissolved),  $C_{\text{O}_3}$  is the total concentration of  $\text{O}_3$ , and  $C_{\text{H}_2\text{O}_2}$  is  
 the total concentration of hydrogen peroxide.

The effective rate constants  $F_1$  and  $F_2$  are given by the following expressions:

$$100 \quad F_1 = R_{\text{O}_3} f_1 \quad (2)$$

$$F_2 = R_{\text{H}_2\text{O}_2} f_2 \quad (3)$$

The reaction rate constants  $R_{\text{O}_3}$  and  $R_{\text{H}_2\text{O}_2}$  refer to Maahs (1983) and Martin et al. (1984):

$$R_{\text{O}_3} = \left\{ 4.4 \times 10^{11} \exp(-4131/T) + 2.61 \times 10^3 \exp(-966/T) [H^+] \right\}^{-1} (\text{Ms})^{-1} \quad (4)$$

$$R_{\text{H}_2\text{O}_2} = 8 \times 10^4 \exp[-3650(1/T - 1/298)] \left\{ 0.1 + [H^+] \right\}^{-1} (\text{Ms})^{-1} \quad (5)$$

105 In Equations (2) and (3), the factors  $f_1$  and  $f_2$  represent the partitioning of the substance between the aqueous and gas  
 phases and are determined by the Henry's law coefficients.

$$f_1 = \gamma f_{\text{SO}_2} f_{\text{O}_3} K_S \bar{K}_{\text{HO}} \quad (6)$$

$$f_2 = \gamma f_{\text{SO}_2} f_{\text{H}_2\text{O}_2} \bar{K}_{\text{HS}} \bar{K}_{\text{HP}} \quad (7)$$

where  $\gamma$  is the dimensionless volume fraction of liquid water in the cloud. The parameters  $f_{\text{SO}_2}$ ,  $f_{\text{O}_3}$  and  $f_{\text{H}_2\text{O}_2}$  are the  
 110 proportions of individual substances in the gas phase. They are also calculated from the dimensionless Henry's law constant  
 and  $\gamma$ .

$$f_{\text{SO}_2} = \left( 1 + \gamma \bar{K}_{\text{HS}} K_S \right)^{-1} \quad (8)$$



$$f_{O_3} = (1 + \gamma \bar{K}_{HO})^{-1} \quad (9)$$

$$f_{H_2O_2} = (1 + \gamma \bar{K}_{HP})^{-1} \quad (10)$$

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$$K_S = \bar{K}_{HS} \left( 1 + \frac{K_{1S}}{[H^+]} + \frac{K_{1S}K_{2S}}{[H^+]^2} \right) \quad (11)$$

The Henry's law constants used in (6) to (8) are listed in table 1. The equilibrium constant  $K_{HS}$  in Table 1 is set to be  $1.23 \times 10^{-3}$  M/atm in CUACE which is the same to that in Von et al (2000) but is  $10^3$  times lower than that in Leighton et al (1990).

120 In order to consider the dependence of the oxidation rates on the pH, the  $H^+$  concentration is calculated from ions balance.

$$[H^+] + [NH_4^+] = [OH^-] + 2[SO_4^{2-}] + 2[SO_3^{2-}] + [HSO_3^-] + [NO_3^-] + [HCO_3^-] \quad (12)$$

From Eqs. (1) ~ (12), CUACE can simulate the oxidation rates of  $SO_2$  by  $H_2O_2$  and  $O_3$  mainly in the liquid and gaseous environment in both stratocumulus and convective clouds in three-dimensional way.

## 125 2.2 Assessment criteria

Three variables, RTCLD, DT, and RT, are defined to assess the effect of the cloud chemistry mechanism on  $SO_2$  and sulfate. The RTCLD is defined as the ratio of changes in chemical species concentrations with and without the cloud chemistry mechanism.

$$RTCLD(i) = 1 - \frac{BECLD(i)}{AFCLD(i)} \quad (13)$$

130 Where  $i$  denotes the chemical component of  $SO_2$ ,  $O_3$ ,  $H_2O_2$ , and sulfate, hereafter. The BECLD denotes the concentration of component  $i$  before cloud chemistry, and the AFCLD after the cloud chemistry.

The DT indicates the difference in concentration of substance  $i$  with (CLD) and without (CCLD) cloud chemistry.

$$DT(i) = CLD(i) - CCLD(i) \quad (14)$$

The RT represents the concentration ratio change of the substance  $i$  with and without cloud chemistry:

$$135 \quad RT(i) = 1 - \frac{CCLD(i)}{CLD(i)} \quad (15)$$



## 2.3 Methodology

### 2.3.1 Model Evaluation – Case 1

Mount Tai with an altitude of about 1500 m, located in central Shandong Province, is the high point of the North China Plain. As the Riguan Peak of Mount Tai is far from pollution sources, and the water vapor conditions in summer favors very much cloud formations, it is an ideal observation site for cloud chemistry observation (Li et al., 2017; Li et al., 2020a; Li et al., 2020b). The observed concentrations of SO<sub>2</sub>, O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub> in cloudy conditions from June 19 to July 30, 2015 and from June 20 to July 30, 2018 with time interval of 1 h are obtained to evaluate the cloud chemistry scheme in WRF/CUACE (Li et al., 2017; Li et al., 2020a; Li et al., 2020b).

WRF/CUACE is set up with two-domain nesting for the evaluation, with the Riguan Peak as the central point (Fig. 1a). The horizontal resolution of outer domain (O) is 9 km with 100×104, and of the inner domain (I) is 3 km with 88×94 (Fig. 1a). There are 32 vertical layers with the top pressure of 100 hPa.

### 2.3.2 Simulations of Regional Characteristics – Case 2

In order to assess the regional contribution of cloud chemistry to SO<sub>2</sub> and sulfate in CUACE, December 2016 is selected with a widespread heavy pollution episode occurred in North and East China from Dec. 16 to 21 and covering NCP, YRD and SCB with the highest hourly fine particulate matter (PM<sub>2.5</sub>) concentration exceeding 1100 µg m<sup>-3</sup> (Yuan and Ma, 2017). The simulation region is set up as shown in Figure 1b also with two-level domain nesting. The outer domain (O) covers Central and East Asia with a horizontal resolution of 54 km and a grid of 139×112. The inner domain (I) covers most of China on the eastern side of the Qinghai-Tibet Plateau with a horizontal resolution of 18 km and a grid of 157×166. The vertical layer number of the model is the same as that in the Case 1.

Since the cloud water is the reaction pool of cloud chemistry, whether the simulation of cloud water is reasonable or not is directly related to the effectiveness of cloud chemistry. Both the cloud water and rainwater from WRF are coupled to the cloud chemistry module and main physics configurations are listed in Table 2.

## 2.4 Meteorological, environmental and Satellite Data

For both cases, the meteorological initial and boundary conditions for WRF/CUACE are from NCEP (National Centers for Environmental Prediction) FNL global reanalysis at a resolution of 1° with 6-h interval. The chemical lateral boundary conditions are from NOAA (National Oceanic and Atmospheric Administration) Meteorological Laboratory Regional Oxidant Model (NALROM) (Liu et al., 1996). The model is run in a restart way with a 5-day spin-up.



FY-2G cloud image data of CMA (China Meteorological Administration) with an 1 h interval is used to evaluate the cloud in both cases. Routine observations in 3 h interval from 23 meteorological stations of CMA and hourly pollution  
165 observations from 132 stations of the China General Environmental Monitoring Station are used to evaluate the meteorological fields and pollutants for December 2016. This includes Beijing, Tianjin, Tangshan, Shijiazhuang, Baoding, Hengshui and Jinan in the North China Plain, Shanghai, Nanjing, Hangzhou, Hefei, Suzhou, Yangzhou and Nantong in the Yangtze River Delta, Guangzhou, Dongguan, Zhuhai and Zhongshan in the Pearl River Delta, and Chengdu, Yibin, Nanchong, Deyang and Zigong in the Sichuan Basin. Meteorological elements include 2 m air temperature, 2 m relative  
170 humidity, and 10 m wind speed.

The MEIC (Multi-resolution Emission Inventory for China) inventory, at a resolution of  $0.25^\circ$ , is used as the anthropogenic emissions with the species of sulfur dioxide ( $\text{SO}_2$ ), nitrogen oxides ( $\text{NO}_x$ ), carbon monoxide (CO), ammonia ( $\text{NH}_3$ ), black carbon (BC), organic carbon (OC), non-methane volatile organic compounds (NMVOCs),  $\text{PM}_{2.5}$  and  $\text{PM}_{10}$  by five sectors of power, industry, transportation, residential, and agriculture (Li et al.,2017; Zheng et al.,2018). The emission  
175 base year of 2015 and 2017 are used for Case-1, of 2016 for Case-2, respectively.

### 3 Results and Discussions

#### 3.1 Evaluation of the cloud chemistry mechanism

In order to verify the cloud chemistry mechanism in WRF/CUACE, the simulations are compared with the observations at Mount Tai. By analyzing the satellite cloud images in and around Mount Tai and matching with the available observed data,  
180 two time periods with clouds from June 19 to July 30, 2015 and June 20 to July 30, 2018 are selected for the comparisons, defined as "cloud process-1" (CP-1) and "cloud process-2" (CP-2), respectively. The statistics of correlation coefficients (R), relative average deviation (RAD), and normalized mean deviation (NMB) between hourly simulated and observed  $\text{SO}_2$ ,  $\text{O}_3$ ,  $\text{H}_2\text{O}_2$  and sulfate are shown in Table 3. The results show that the RADs of  $\text{SO}_2$ ,  $\text{O}_3$ ,  $\text{H}_2\text{O}_2$  and sulfate in CP-1 and CP-2 are in the range of -35 % ~ 78 %, and the R for the four species are 0.34, 0.33 and 0.61 and 0.32 for CP-1, and 0.47, 0.40, 0.06 and  
185 0.54 for CP-2, respectively. The simulated and observed averages are within the same magnitude and close to each other. Among them, the simulated and observed averages of  $\text{SO}_2$  are very close in both CP-1 and CP-2, with RAD about -3% and -6%. The analysis reveals that although the correlation of  $\text{H}_2\text{O}_2$  in CP-2 is only 0.06, the RAD is 18%, the NMB is -19.6%, and the simulated mean value of  $\text{H}_2\text{O}_2$  is closer to the observed mean value compared to CP-1. For sulfate, the simulated correlation is good with R of 0.32 and 0.54, but the model underestimates sulfate concentrations with NMB of -71 % and -  
190 59.4 %. This indicates that although the model is able to simulate the trend of sulfate concentrations, the simulated



concentrations are lower than the observed values. Therefore, the underestimation of sulfate may be due to the incompleteness of other cloud chemistry mechanisms or the bias of liquid water content in the cloud physics. The above results show that SO<sub>2</sub>, sulfate, and two important oxidants, O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub>, are reasonably simulated in both time periods.

It is also found that when cloud chemistry occurs, SO<sub>2</sub> concentration values range from 0-2.5 ppbv, mostly less than 1 ppbv in CP-2, O<sub>3</sub> in the range of 25-125 ppbv, H<sub>2</sub>O<sub>2</sub> in the range of 0-100 μM (Fig. 2), and sulfate in the range of 0-100 μg/m<sup>3</sup>. All four species are agreed within a factor two of observed concentrations, while the sulfate is obvious underestimated in CP-1. In addition, CP-2 shows the observed concentration of H<sub>2</sub>O<sub>2</sub> is increased, compared to CP-1. The SO<sub>2</sub> concentrations are small for both time periods. Shen et al. (2012) found mean SO<sub>2</sub> concentrations of 8.1 ppbv and 6.4 ppbv for the Mount Tai site in summer 2007 and 2008. Li et al. (2020) observed that the yearly mean summer SO<sub>2</sub> concentrations of 1.3 ± 1.1 ppbv, 0.3 ± 0.2 ppbv, and 0.4 ± 0.2 ppbv in 2014, 2015, and 2018, respectively. The average O<sub>3</sub> concentrations are 50.3 ± 10.9 ppbv, 73.1 ± 19.7 ppbv and 66.3 ± 18.3 ppbv, respectively. The mean H<sub>2</sub>O<sub>2</sub> concentrations are 11.6 ± 13.1 μM, 22.8 ± 13.1 μM, and 45.1 ± 29.5 μM, respectively. Besides, H<sub>2</sub>O<sub>2</sub> concentrations are 0.55 ± 0.67 ppbv in 2007 (Ren et al., 2009), 0.93 ± 1.01 ppbv in 2018 (Ye et al., 2021), 2.05 ± 1.20 ppbv in 2019 (Ye et al., 2021). The results of our simulations are consistent with the trends of other observational studies, which all indicate clearly cloud chemistry processes consume SO<sub>2</sub>, H<sub>2</sub>O<sub>2</sub> and O<sub>3</sub> to produce sulfate. The results of our simulations are consistent with the trends of other observational studies, which all indicate clearly an increase in the atmospheric oxidation and a decrease in the SO<sub>2</sub> concentration.

Figure 3 shows the satellite cloud images, simulated column cloud and simulated liquid water content at 8:00 LST on June 24, and 8:00 LST on June 25 in CP-1. The results indicate that not only the simulated column cloud but also the cloud liquid water distribution simulated at the two times are consistent with the satellite observations.

Figure 4 shows the RTCLD of SO<sub>2</sub> and simulated liquid water contents at 2:00 and 8:00 LST on June 24, and 2:00 and 8:00 LST on June 25 in CP-1. The results indicate that the RTCLD of SO<sub>2</sub> distribution simulated at the four times are consistent with the cloud liquid water, so for the distribution of RTCLD of SO<sub>2</sub> with a reduction of more than 80% within the cloud region of Mount Tai and most of Shandong province, which is consistent with the cloud chemistry observation studies by Li (2020).

In summary, the cloud chemistry mechanism in WRF/CUACE is reasonable to reproduce the cloud chemistry for the gaseous pollutant SO<sub>2</sub>, sulfate and the important oxidants of O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub>. The model not only simulates the concentrations of the four species, but also the SO<sub>2</sub> decreasing trend and O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub> increasing trends with year.





### 3.2 Assessment of the cloud chemistry impacts on regional SO<sub>2</sub> and sulfate

220 This session will further assess the contribution of cloud chemistry for the four main pollution regions of NCP, YRD, PRD, and SCB (Fig. 1) in China for the whole December of 2016 and a pollution episode occurred during this period (Dec. 16-21) as Case 2.

#### 3.2.1 Meteorological evaluation

225 As the driving force of air pollution and cloud chemistry, the meteorological conditions simulated in the four major pollution regions (Table 4) are evaluated with observations. The correlations of all elements with observations are very close in both the whole December and the heavy pollution in four regions, indicating that the model performed well both in December and the heavy pollution episode. The temperature correlation is the best in December, followed by humidity and then wind speed which is consistent with previous researches (Zhou et al., 2012; Wang et al., 2015; Gao et al., 2016). The RMSEs of wind speeds all range from 1.03 to 1.50 m s<sup>-1</sup>, falling within the criteria (less than 2 m s<sup>-1</sup>) to define “good” model performance in stagnant weather proposed by Emery et al. (Emery et al., 2001). The RSME of wind speed during heavy pollution is smaller than that of December, which indicates that the model can well performance the static wind even though it is very small.

235 Figure 5 shows the satellite cloud image, simulated column cloud and simulated liquid water content for the maturation and dissipation stages of the episode. The satellite image shows that the cloud coverage region is mainly in the southwest including SCB on the 19<sup>th</sup>, almost covers most of eastern China including NCP, YRD, PRD and SCB on the 20<sup>th</sup> and the 21<sup>st</sup>, and then move eastward outside of China on the 22<sup>nd</sup> (Fig.4 a1-d1). The simulated clouds fit well with the satellite images (Fig.4 a2-d2). Since the cloud chemistry occurs in the cloud area, the cloud chemistry mainly affects three regions of NCP, YRD and SCB of the four pollution regions as PRD has relatively much less cloud. The column liquid water distribution also moves from west to east as the episode develops (Fig.5 a3-d3), but is located more southern part of eastern China than that of the cloud. In SCB and YRD, there is more abundant in liquid water content, which can reach more than 500 g m<sup>-2</sup>, while in PRD There is slightly less, up to 100 g m<sup>-2</sup>, and in NCP there is the least liquid water content, 0.1 g m<sup>-2</sup> mostly due to the dry environment and partly due to the overestimated temperature and underestimated humidity in Table 4.

245 The above analysis shows that the model basically reproduces the meteorological field in December and heavy pollution periods, which provides a better meteorological background basis for the effective simulation of pollution as well as cloud chemistry.



### 3.2.2 Pollutants Evaluation

The simulated hourly PM<sub>2.5</sub>, O<sub>3</sub> and SO<sub>2</sub> concentrations in four regions are also compared with the observations (Table 5). The simulations are all within a factor of two of the observations (figure omitted), and the mean values of the three pollutants simulated in the four regions are very close to the observations, indicating that model captures well the variability of PM<sub>2.5</sub>, O<sub>3</sub> and SO<sub>2</sub> concentrations during both December 2016 and the heavy pollution episode. In detail, O<sub>3</sub> correlates well with observations in all four regions for the whole month but not so well in the episode. PM<sub>2.5</sub> correlates better at PRD and YRD during the episode, but not so well in other regions and other two time periods. It is much less underestimated at NCP and SCB. Compared to the whole month, SO<sub>2</sub> is better simulated at NCP, YRD, and SCB during the pollution period. The following part of this paper will focus on assessing the effects of cloud chemical processes.

### 3.2.3 Assessment of regional contributions

The regional impacts of cloud chemical processes on surface SO<sub>2</sub> and sulfate are analyzed for whole December and for a pollution episode of Dec. 16-21. The pollution episode is investigated with respect to the maturity stage (21:00 LST on Dec. 20 - 17:00 LST on Dec. 21) and to the dissipation stage (2:00 LST on Dec. 22) for the four pollution regions NCP, YRD, PRD, and SCB.

The average impact of cloud chemistry on surface SO<sub>2</sub> and sulfate in December (Fig. 6, DT(SO<sub>2</sub>) and DT (sulfate)) is simulated. It is found that SO<sub>2</sub> declination in December is concentrated mostly in the central-eastern part of China, by an average of 0.1-1 ppb in most regions. SO<sub>2</sub> concentrations are reduced by 0.5-3.0 ppb in most part of NCP, YRD, PRD and SCB regions. Among them, there is a relatively stronger center by declining 3.0-10.0 ppb in SCB. Correspondingly, sulfate growth is mainly centering in SCB, with the maximum center up to 20-50 µg/m<sup>3</sup>. Sulfate concentrations are increased by 10-20 µg/m<sup>3</sup> in most part of NCP, YRD and PRD, and increased 5-10 µg/m<sup>3</sup> in others.

The spatial distribution of cloud chemistry contribution to SO<sub>2</sub> and sulfate during the whole mature stage of the heavy pollution is analyzed (Fig. 7, DT(SO<sub>2</sub>) and DT(sulfate)) on the 19<sup>th</sup> to the 21<sup>st</sup>. It shows that SO<sub>2</sub> decreases most significantly in SCB, exceeding 1-3 ppb in most area, to 3-10 ppb in the central region. It also reaches 1-3 ppb in most parts of YRD, PRD and NCP, while the smallest decrease is below 1 ppb in the northern part of NCP. Meanwhile, in terms of regional distribution, the regions of increasing sulfate concentration and decreasing sulfur dioxide concentration are correlated, but not identical. Sulfate concentration increases by more than 10 µg/m<sup>3</sup> from the southern part of NCP to central China and from most of the eastern part of SCB to the YRD region. Sulfate concentration increases by more than 20 µg/m<sup>3</sup> in SCB and up to more than 50 µg/m<sup>3</sup> in the central region, by 20-30 µg/m<sup>3</sup> in the YRD central region, and by 5-20 µg/m<sup>3</sup> in the PRD region.



In summary, comparing the contribution of cloud chemistry in the whole December with the pollution maturity stage of  
275 19-21, it shows that cloud chemistry in heavy pollution weather for SO<sub>2</sub> depletion and sulfate increase is mainly concentrated  
in the central-eastern part of China, and the four major pollution regions are also more obvious. However, SO<sub>2</sub> consumption  
and sulfate increase are not consistent, which is not only influenced by the local SO<sub>2</sub> concentration, but also by the cloud  
amount. Therefore, for SCB, which is less polluted and has much more clouds than NCP, the impact of cloud chemistry on  
sulfate and its precursor SO<sub>2</sub> is always the most significant in the SCB, both during heavy pollution and averaged over the  
280 whole December.

Exploring details into the pollution episode, it is found that the cloud chemistry influence is mainly on SCB and YRD at  
21:00 LST on Dec. 20. The sulfate concentration increases (RT(sulfate)) by about 20-100 µg/m<sup>3</sup> in most parts of SCB, and  
the highest is about 150-225 µg/m<sup>3</sup> in its southwest (Fig. 8a). The increase is about 10-40 µg/m<sup>3</sup> in most parts of YRD, and  
the highest increase can reach 40-100 µg/m<sup>3</sup> in the southeast near Hangzhou Bay. At 21<sup>st</sup> 17:00 LST, it shows that the  
285 pollution episode has moved eastward and the cloud chemistry process has a stronger impact on NCP than that of the  
previous time (Fig. 8b). The increase center is located in Shandong Province, with up to 100-225 µg/m<sup>3</sup>, and with about 10-60  
µg/m<sup>3</sup> in most of NCP. In PRD, it has been influenced the least by the cloud process with sulfate concentration increase of 10-  
40 µg/m<sup>3</sup>. At 22<sup>nd</sup> 12 LST, although the episode has gradually dissipated, it still had an impact on SCB and YRD with sulfate  
increase about 10-40 µg/m<sup>3</sup> in most of these four regions (Fig. 8c).

290 Above all, the contribution of cloud chemistry to surface sulfate during this pollution process is the highest in the SCB,  
followed by the NCP and YRD, with maximum concentration increases of 225 µg/m<sup>3</sup>, 100 µg/m<sup>3</sup>, and 40 µg/m<sup>3</sup>, respectively,  
and sulfate increases of 10-40 µg/m<sup>3</sup> in most regions of these three areas.

Further analysis of the simulation characteristics of cloud chemistry on all the regions during the pollution maturity  
stage, i.e., the cloud process from Dec. 19 to 21 (Table 6), and the whole month of December (Table 7), is investigated.  
295 Compared with CCLD, R of SO<sub>2</sub> in CLD increases by 0.6, 0.14, and 0.1 in YRD, SCB, and NCP, respectively, and the  
overestimation in NCP and PRD has been corrected during the pollution episode. R also increases by 0.1, 0.03 and 0.04 in  
YRD, SCB and NCP for the whole December, respectively.

For the performance of PM<sub>2.5</sub> of CLD, it decreases the underestimation in NCP and SCB with decreasing NMB and Rs  
remain almost the same with that of CCLD in the four pollution regions during the pollution episode and the whole December.  
300 There are little difference between the statistics results of R, NMB and RMSE with some positive improve for  
underestimation in NCP and SCB. R improves by 0.1 and NMB with cloud chemistry decreases from 39.25 % to 13.75 %.  
Moreover, the statistical results of all stations (SUM in Table 8) shows that the NMB of SO<sub>2</sub> decreased from 39.25% to



13.75% and the NMB of sulfate increased from -40.75% to -31.62% during the pollution period after the addition of cloud chemistry simulation, so the simulation bias of both SO<sub>2</sub> and sulfate decreased during the pollution period.

305

### 3.3 Site evaluation of cloud chemistry

Representative sites of Beijing, Nanjing, Guangzhou and Chengdu at NCP, YRD, PRD and SCB are selected to quantify the impact of cloud chemistry during the pollution episode. The net depletion ratio of SO<sub>2</sub> column concentration (RT(SO<sub>2</sub>)) during cloud chemistry is shown in Figure 9. It is found that SO<sub>2</sub> column concentration reduction maintained mostly a high value of over 60%, even to 80% sometimes, in Chengdu from 16<sup>th</sup> to 21<sup>st</sup>. In Nanjing, the SO<sub>2</sub> level is reduced by about 20-50% during the time from 17<sup>th</sup> to 19<sup>th</sup> and up to 80% from 20<sup>th</sup> to 21<sup>st</sup> when the episode matures there. The changes of SO<sub>2</sub> in these two cities are consistent with the changes in cloud and liquid cloud water content distributions during the pollution period in Figure 3. The SO<sub>2</sub> reduction in Beijing and Guangzhou is consistently maintained at around 40% during the time from 17<sup>th</sup> to 21<sup>th</sup>. The lower oxidative transformation is related to the lower liquid water content in Beijing, while in Guangzhou it is attributed to the combination of low pollution levels and low cloud water content. Figure 3 shows that Chengdu maintained abundant water vapor conditions from 17<sup>th</sup> to 21<sup>st</sup>, and so does Nanjing from 20<sup>th</sup> to 21<sup>st</sup>. However, the ambient water vapor content is quite low in Guangzhou and Beijing throughout the process, so the SO<sub>2</sub> oxidation is much lower than that of Chengdu and Nanjing. In conclusion, the cloud chemistry process can lead to SO<sub>2</sub> column concentration consumption share of more than 60% when cloud water content is abundant, which is also consistent with the observations of Mount Tai by Li (2020).

The impact of cloud chemistry on surface SO<sub>2</sub> and sulfate in four sites are also shown in Figure 10. The overall trend shows that the peak and valley regions of surface SO<sub>2</sub> consumption and sulfate increase are coincident. The cloud chemical processes make the surface SO<sub>2</sub> oxidation vary greatly between cities in different regions (Fig. 10a). From the 19<sup>th</sup> to the 21<sup>st</sup>, the percentage of surface SO<sub>2</sub> consumption can reach more than 90% in Chengdu and Nanjing, while it is below 30% in Beijing and Guangzhou, and does not reach 40% until the 22<sup>nd</sup>. Although the percentage of surface SO<sub>2</sub> consumption varies greatly, the increase in the percentage of sulfate does not vary so much between cities. From 16<sup>th</sup> to 22<sup>nd</sup>, the increase in surface sulfate in the four cities ranges from 60-95% (Figure 10b), and the sulfate increase rate interval contains the results summarized by Turnock et al. (2019).

Figure 11 is the variation of vertical profiles of sulfate increase by cloud chemistry at the four times at 12:00 LST on 20 Dec., at 04:00 LST on 21 Dec., at 04:00 and 12:00 LST on 22 Dec. in Beijing, Nanjing, Chengdu and Guangzhou. It shows



that the sulfate produced by cloud chemistry during this pollution process is concentrated mostly below 5 km in the troposphere, especially under 2 km. Again, less sulfate has been produced in Beijing in vertical than that of others by cloud chemistry.

#### 4. Summary and conclusions

335 The cloud chemistry mechanism in WRF/CUACE has been assessed by using the in-situ cloud chemistry observations of SO<sub>2</sub>, O<sub>3</sub>, and H<sub>2</sub>O<sub>2</sub> from Mount Tai in June-July 2015 and 2018. The results show that the mechanism well captures the cloud processes for the oxidation of SO<sub>2</sub>, reducing SO<sub>2</sub> by more than 80 % during the cloudy phase of heavy pollution, which is in good agreement with the observations.

The cloud chemistry contributions to the changes of SO<sub>2</sub> and sulfate in NCP, YRD, PRD and SCB regions are assessed  
340 by WRF/CUACE. During the pollution episode from 19 to 21, all four pollution areas are significantly affected by cloud chemistry, with SCB being the most obvious, with surface SO<sub>2</sub> reduction averaging up to 1-3 ppb and reaching 3-10 ppb in the high value areas, and surface sulfate increasing 10-30 µg/m<sup>3</sup> on average, with a maximum of more than 70 µg/m<sup>3</sup>. Most areas in NCP, YRD and PRD had an average SO<sub>2</sub> reduction of 0.5-3 ppb and sulfate increase of 5-30 µg/m<sup>3</sup>. Meanwhile, the Beijing area of NCP had the least impact among the four regions. Although the average impact of cloud chemistry is much  
345 weaker in the NCP with less water vapor there in December, the contribution of the southern part of NCP during heavy pollution time is still significant and cannot be ignored. In PRD, the contribution of cloud chemistry is weaker than other regions due to lighter pollution, although there are lots of clouds with abounded liquid water there. In addition, surface sulfate increases by 60-95 % in Beijing, Nanjing, Chengdu and Guangzhou in December during heavy pollution, similar to that of  
350 pollution period is significantly greater than that for all December. Vertically, the results of the pollution process for the 2016 winter heavy pollution episode showed that the cloud chemistry influence is mainly in the middle and lower troposphere below 2 km for four representative cities. Generally, the cloud chemistry improved the model performance by reducing the SO<sub>2</sub> overestimates and enhancing the correlations with observations for both SO<sub>2</sub> and sulfate.

This paper has been focused on the cloud chemical mechanism evaluation, and assessed the contribution of cloud  
355 chemistry to SO<sub>2</sub> and sulfate changes. In the future, more mechanisms should be added to improve the cloud chemistry mechanism in CUACE to more accurately simulate SO<sub>2</sub> and sulfate and other aerosol components such as nitrate, ammonium, carbonate, and organic aerosols.



#### **Code/data availability**

All source code and data can be accessed by contacting the corresponding authors Sunling Gong ([gongsl@cma.gov.cn](mailto:gongsl@cma.gov.cn)).

#### 360 **Authors contribution**

CZ and SG put forward the ideas and formulated overarching research goals. JL carried them out and wrote the manuscript with suggestions from all authors. LZ and JZ participated in the scientific interpretation and discussion. JC was assisted with data acquisition and processing. All authors contributed to the discussion and improvement of the manuscript.

#### **Competing interests**

365 The authors declare that they have no conflict of interest.

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370 **References**

- Alexander, B., Park, R. J., Jacob, D. J., and Gong, S.: Transition metal-catalyzed oxidation of atmospheric sulfur: Global implications for the sulfur budget, *Journal of Geophysical Research*, 114, <https://doi.org/10.1029/2008jd010486>, 2009.
- Buchard, V., da Silva, A. M., Colarco, P., Krotkov, N., Dickerson, R. R., Stehr, J. W., Mount, G., Spinei, E., Arkinson, H. L., and He, H.: Evaluation of GEOS-5 sulfur dioxide simulations during the Frostburg, MD 2010 field campaign, *Atmos. Chem. Phys.*, 14, 1929-1941, <https://doi.org/10.5194/acp-14-1929-2014>, 2014.
- 375 Binkowski, F. S., and Roselle, S. J.: Models - 3 Community Multiscale Air Quality (CMAQ) model aerosol component 1. Model description, *Journal of Geophysical Research: Atmospheres*, 108, <https://doi.org/10.1029/2001jd001409>, 2003.
- Caffrey, P., Hoppel, W., Frick, G., Pasternack, L., Fitzgerald, J., Hegg, D., Gao, S., Leitch, R., Shantz, N., Albrechtinski, T., and Ambrusko, J.: In-cloud oxidation of SO<sub>2</sub> by O<sub>3</sub> and H<sub>2</sub>O<sub>2</sub>: Cloud chamber measurements and modeling of particle growth, *Journal of Geophysical Research: Atmospheres*, 106, 27587-27601, <https://doi.org/10.1029/2000jd900844>, 2001.
- 380 Chang, J. S., Brost, R. A., Isaksen, I. S. A., Madronich, S., Middleton, P., Stockwell, W. R., and Walcek, C. J.: A three-dimensional eulerian acid deposition model physical concepts and formulation, *journal of geophysical research*, 92, 14,681-614,700, <https://doi.org/10.1029/jd092id12p14681>, 1987.
- Chapman, E. G., Gustafson, W. I., Easter, R. C., Barnard, J. C., and Fast, J. D.: Coupling aerosol-cloud-radiative processes in the WRF-Chem model: Investigating the radiative impact of elevated point sources, *Atmos. Chem. Phys.*, 9, 945 - 964, <https://doi.org/10.5194/acp-9-945-2009>.
- 385 Charlson, R., J., Schwartz, and S., E.: Climate forcing by anthropogenic aerosols, *Science*, 255, 423-423, <https://doi.org/10.1126/science.255.5043.423>, 1992.
- Emery, C., Tai, E., and Yarwood, G.: Enhanced meteorological modeling and performance evaluation for two Texas ozone episodes, 2001.
- 390 Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He, K., and Carmichael, G.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China, *Science Advances*, <https://doi.org/10.1126/sciadv.1601530>, 2016.
- Dovrou, E., Rivera-Rios, J. C., Bates, K. H., and Keutsch, F. N.: Sulfate Formation via Cloud Processing from Isoprene Hydroxyl Hydroperoxides (ISOPOOH), *Environ Sci Technol*, 53, 12476-12484, <https://doi.org/10.1021/acs.est.9b04645>, 2019.
- 395 Ervens, B.: Modeling the Processing of Aerosol and Trace Gases in Clouds and Fogs, *Chemical Reviews*, 115, 4157-4198, <https://doi.org/10.1021/cr5005887>, 2015.



- Faloona, I.: Sulfur processing in the marine atmospheric boundary layer: A review and critical assessment of modeling  
400 uncertainties. *Atmospheric Environment*, 43(18), 2841-2854, <https://doi.org/10.1016/j.atmosenv.2009.02.043>, 2009.
- Faloona, I., Conley, S. A., Blomquist, B., Clarke, A. D., Kapustin, V., Howell, S., Lenschow, D. H., and Bandy, A. R.: Sulfur  
dioxide in the tropical marine boundary layer: dry deposition and heterogeneous oxidation observed during the Pacific  
Atmospheric Sulfur Experiment, *Journal of Atmospheric Chemistry*, 63, 13-32, <https://doi.org/10.1007/s10874-010-9155-0>,  
2010.
- 405 Gao, M., Carmichael, G. R., Wang, Y., Ji, D., Liu, Z., and Wang, Z.: Improving simulations of sulfate aerosols during winter  
haze over Northern China: the impacts of heterogeneous oxidation by NO<sub>2</sub>, *Frontiers of Environmental Science &  
Engineering*, 10, 11, <https://doi.org/10.1007/s11783-016-0878-2>, 2016.
- Ge, W., Liu, J., Yi, K., Xu, J., Zhang, Y., Hu, X., Ma, J., Wang, X., Wan, Y., Hu, J., Zhang, Z., Wang, X., and Tao, S.:  
Influence of atmospheric in-cloud aqueous-phase chemistry on global simulation of SO<sub>2</sub> in CESM2, *Atmos. Chem. Phys.*,  
410 <https://doi.org/10.5194/acp-2021-406>, 2021.
- Georgiou, G. K., Christoudias, T., Proestos, Y., Kushta, J., Hadjinicolaou, P., and Lelieveld, J.: Air quality modelling in the  
summer over the eastern Mediterranean using WRF-Chem: chemistry and aerosol mechanism intercomparison, *Atmos. Chem.  
Phys.*, 18, 1555-1571, <https://doi.org/10.5194/acp-18-1555-2018>, 2018.
- Gen, M.; Zhang, R.; Huang, D. D.; Li, Y.; Chan, C. K.: Heterogeneous SO<sub>2</sub> Oxidation in Sulfate Formation by Photolysis of  
415 Particulate Nitrate, *Environ. Sci. Technol*, 6 (2), 86–91, <https://doi.org/10.1021/acs.estlett.8b00681>, 2019a.
- Gen, M., Zhang, R., Huang, D. D., Li, Y., and Chan, C. K.: Heterogeneous Oxidation of SO(2) in Sulfate Production during  
Nitrate Photolysis at 300 nm: Effect of pH, Relative Humidity, Irradiation Intensity, and the Presence of Organic Compounds,  
*Environ Sci Technol*, 53, 8757-8766, <https://doi.org/10.1021/acs.est.9b01623>, 2019b.
- Gong, S., and Zhang, X.: CUACE/Dust – an integrated system of observation and modeling systems for operational dust  
420 forecasting in Asia, *Atmospheric Chemistry and Physics*, 8, 2333-2340, <https://doi.org/10.5194/acp-8-2333-2008>, 2008.
- Gong, S. L., Barrie, L. A., Blanchet, J.-P., von Salzen, K., Lohmann, U., Lesins, G., Spacek, L., Zhang, L. M., Girard, E., Lin,  
H., Leaitch, R., Leighton, H., Chylek, P., and Huang, P.: Canadian Aerosol Module: A size-segregated simulation of  
atmospheric aerosol processes for climate and air quality models 1. Module development, *Journal of Geophysical Research*,  
108, <https://doi.org/10.1029/2001jd002002>, 2003.
- 425 Guo, J., Wang, Y., Shen, X., Wang, Z., Lee, T., Wang, X., Li, P., Sun, M., Jeffrey, L., Collett, J., Wang, W., and Wang, T.:  
Characterization of cloud water chemistry at Mount Tai, China: Seasonal variation, anthropogenic impact, and cloud  
processing, *Atmospheric Environment*, 60, 467-476, <https://doi.org/10.1016/j.atmosenv.2012.07.016>, 2012.





- Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China, *Proc Natl Acad Sci U S A*, 111, 17373-17378, 430 <https://doi.org/10.1073/pnas.1419604111>, 2014.
- Harris, E., Sinha, B., van Pinxteren, D., Tilgner, A., Fomba, K. W., Schneider, J., Roth, A., Gnauk, T., Fahlbusch, B., Mertes, S., Lee, T., Collett, J., Foley, S., Borrmann, S., Hoppe, P., and Herrmann, H.: Enhanced role of transition metal ion catalysis during in-cloud oxidation of SO<sub>2</sub>, *Science*, 340, 727-730, <https://doi.org/10.1126/science.1230911>, 2013.
- He, J., Zhang, Y., Glotfelty, T., He, R., Bennartz, R., Rausch, J., and Sartelet, K.: Decadal simulation and comprehensive 435 evaluation of CESM/CAM5.1 with advanced chemistry, aerosol microphysics, and aerosol-cloud interactions, *Journal of Advances in Modeling Earth Systems*, 7, 110-141, <https://doi.org/10.1002/2014ms000360>, 2015.
- Hung, H. M., Hsu, M. N., and Hoffmann, M. R.: Quantification of SO<sub>2</sub> Oxidation on Interfacial Surfaces of Acidic Micro-Droplets: Implication for Ambient Sulfate Formation, *Environ Sci Technol*, 52, 9079-9086, <https://doi.org/10.1021/acs.est.8b01391>, 2018.
- 440 Kotarba, A. Z.: Calibration of global MODIS cloud amount using CALIOP cloud profiles, *Atmospheric Measurement Techniques*, 13, 4995-5012, <https://doi.org/10.5194/amt-13-4995-2020>, 2020.
- Iibusuki, T., and Takeuchi, K.: Sulfur dioxide oxidation by oxygen catalyzed by mixtures of manganese(II) and iron(III) in aqueous solutions at environmental reaction conditions, *Atmospheric Environment*, 21, 1555-1560, [https://doi.org/10.1016/0004-6981\(87\)90317-9](https://doi.org/10.1016/0004-6981(87)90317-9), 1987.
- 445 Leighton, H. G., and Ivanova, I. T.: Aerosol–Cloud Interactions in a Mesoscale Model. Part I: Sensitivity to Activation and Collision–Coalescence, *Journal of the Atmospheric Sciences*, 65, 289-308, <https://doi.org/10.1175/2007jas2207.1>, 2008.
- Lelieveld, Jos, Heintzenberg, and Jost: Sulfate cooling effect on climate through in-cloud oxidation of anthropogenic SO<sub>2</sub>, *Science*, <https://doi.org/10.1126/science.258.5079.117>, 1992.
- Li, J. R.: Microphysical Characteristics and S(IV) Multiphase Chemical Reaction Mechanism of Orographic Clouds, Ph.D. 450 thesis, Fudan University, 127pp, 2020.
- Li, J., Wang, X., Chen, J., Zhu, C., Li, W., Li, C., Liu, L., Xu, C., Wen, L., Xue, L., Wang, W., Ding, A., and Herrmann, H.: Chemical composition and droplet size distribution of cloud at the summit of Mount Tai, China, *Atmos. Chem. Phys.*, 17, 9885-9896, <https://doi.org/10.5194/acp-17-9885-2017>, 2017.
- Li, J., Zhu, C., Chen, H., Fu, H., Xiao, H., Wang, X., Herrmann, H., and Chen, J.: A More Important Role for the Ozone - 455 S(IV) Oxidation Pathway Due to Decreasing Acidity in Clouds, *Journal of Geophysical Research: Atmospheres*, 125, <https://doi.org/10.1029/2020jd033220>, 2020a.



- Li, J., Zhu, C., Chen, H., Zhao, D., Xue, L., Wang, X., Li, H., Liu, P., Liu, J., Zhang, C., Mu, Y., Zhang, W., Zhang, L., Herrmann, H., Li, K., Liu, M., and Chen, J.: The evolution of cloud and aerosol microphysics at the summit of Mt. Tai, China, *Atmos. Chem. Phys.*, 20, 13735-13751, <https://doi.org/10.5194/acp-20-13735-2020>, 2020b.
- 460 Li, M., Liu, H., Geng, G., Hong, C., Liu, F., Song, Y., Tong, D., Zheng, B., Cui, H., Man, H., Zhang, Q., and He, K.: Anthropogenic emission inventories in China: a review, *Natl. Sci. Rev.*, 4, 834-866, <https://doi.org/10.1093/nsr/nwx150>, 2017.
- Liu, S. C., McKeen, S. A., Hsie, E.-Y., Lin, X., Kelly, K. K., Bradshaw, J. D., Sandholm, S. T., Browell, E. V., Gregory, G. L., Sachse, G. W., Bandy, A. R., Thornton, D. C., Blake, D. R., Rowland, F. S., Newell, R., Heikes, B. G., Singh, H., and R. W. Talbot: Model study of tropospheric trace species distributions during PEM-WEST A, *J. Geophys. Res.*, 101, 465 <https://doi.org/0148-0227/96/95JD-02277505.00>, 1996.
- Li, P. F.: Fog Water Chemistry and Fog-Haze Transformation in Shanghai, Fudan University, Ph.D. thesis, 145pp.
- Liu, T., Chan, A. W. H., and Abbatt, J. P. D.: Multiphase Oxidation of Sulfur Dioxide in Aerosol Particles: Implications for Sulfate Formation in Polluted Environments, *Environ Sci Technol*, 55, 4227-4242, <https://doi.org/10.1021/acs.est.0c06496>, 2021.
- 470 Maahs, H. G.: Kinetics and Mechanism of the Oxidation of S(IV) by Ozone in Aqueous Solution With Particular Reference to SO<sub>2</sub> Conversion in Nonurban Tropospheric Clouds, *Journal of Geophysical Research Oceans*, 88, 10721-10732, <https://doi.org/10.1029/JC088iC15p10721>, 1983.
- Martin, G. M., D. W. Johnson, and A. Spice, The measurement and parameterization of effective radius in warm stratocumulus cloud, *J. Atmos. Sci.*, 51, 1823-1842, [https://doi.org/10.1175/1520-0469\(1994\)051<1823:TMAPOE>2.0.CO;2](https://doi.org/10.1175/1520-0469(1994)051<1823:TMAPOE>2.0.CO;2), 475 1994.
- Martin, L. R., and Good, T. W.: Catalyzed oxidation of sulfur dioxide in solution: The iron-manganese synergism, *Atmospheric Environment*, 25, 2395-2399, [https://doi.org/10.1016/0960-1686\(91\)90113-L](https://doi.org/10.1016/0960-1686(91)90113-L), 1991.
- Menut, L., Bessagnet, B., Khvorostyanov, D., Beekmann, M., Colette, A., Coll, I., Curci, G., Foret, G., Hodzic, A., Mailler, S., Meleux, F., Monge, J. L., Pison, I., Turquety, S., Valari, M., Vautard, R., and Vivanco, M. G.: Regional atmospheric 480 composition modeling with CHIMERE, *Geoscientific Model Development*, 6, 981-1028, <https://doi.org/10.5194/gmdd-6-203-2013>, 2013.
- Molina: Air Quality in the Mexico Megacity: An Integrated Assessment, Alliance for Global Sustainability Bookseries, <https://doi.org/10.1007/978-94-010-0454-1>, 2002.
- Park, R. J., and Jacob, D. J.: Sources of carbonaceous aerosols over the United States and implications for natural visibility, 485 *Journal of Geophysical Research*, 108, <https://doi.org/10.1029/2002jd003190>, 2003.



- Pye, H. O. T., Nenes, A., Alexander, B., Ault, A. P., Barth, M. C., Clegg, S. L., Jr., J. L. C., Fahey, K. M., Hennigan, C. J., Herrmann, H., Kanakidou, M., Kelly, J. T., Ku, I.-T., McNeill, V. F., Riemer, N., Schaefer, T., Shi, G., Tilgner, A., Walker, J. T., Wang, T., Weber, R., Xing, J., Zaveri, R. A., and Zuend, a. A.: Havala O. T. Pye The acidity of atmospheric particles and clouds, *Atmos. Chem. Phys.*, 20, 4809 – 4888, <https://doi.org/10.5194/acp-20-4809-2020>, 2020.
- 490 Ramanathan, V., Crutzen, P. J., Kiehl, J. T., and Rosenfeld, D.: Aerosols, climate, and the hydrological cycle, *Science*, 294, 2119-2124, [10.1126/science.1064034](https://doi.org/10.1126/science.1064034), 2001.
- Ravishankara, A. R.: Heterogeneous and Multiphase Chemistry in the Troposphere, *Science*, 276, 1058-1065, <https://doi.org/10.1126/science.276.5315.1058>, 1997.
- Ren, Y., Ding, A., Wang, T., Shen, X., Guo, J., Zhang, J., Wang, Y., Xu, P., Wang, X., and Gao, J.: Measurement of gas-phase  
495 total peroxides at the summit of Mount Tai in China, *Atmospheric Environment*, 43, 1702-1711, <https://doi.org/10.1016/j.atmosenv.2008.12.020>, 2009.
- Sha, T., Ma, X., Jia, H., Tian, R., Chang, Y., Cao, F., and Zhang, Y.: Aerosol chemical component: Simulations with WRF-Chem and comparison with observations in Nanjing, *Atmospheric Environment*, 218, <https://doi.org/10.1016/j.atmosenv.2019.116982>, 2019.
- 500 Shen, X., Lee, T., Guo, J., Wang, X., Li, P., Xu, P., Wang, Y., Ren, Y., Wang, W., Wang, T., Li, Y., Carn, S. A., and Collett, J. L.: Aqueous phase sulfate production in clouds in eastern China, *Atmospheric Environment*, 62, 502-511, <https://doi.org/10.1016/j.atmosenv.2012.07.079>, 2012.
- Shimadera, H., Kondo, A., Shrestha, K. L., Kaga, A., and Inoue, Y.: Annual sulfur deposition through fog, wet and dry deposition in the Kinki Region of Japan - ScienceDirect, *Atmospheric Environment*, 45, 6299-6308, <https://doi.org/10.1016/j.atmosenv.2011.08.055>, 2011.
- 505 Sielski, J., Kazirod-Wolski, K., Jozwiak, M. A., and Jozwiak, M.: The influence of air pollution by PM<sub>2.5</sub>, PM<sub>10</sub> and associated heavy metals on the parameters of out-of-hospital cardiac arrest, *Sci Total Environ*, 788, 147541, <https://doi.org/10.1016/j.scitotenv.2021.147541>, 2021.
- Song, S., Nenes, A., Gao, M., Zhang, Y., Liu, P., Shao, J., Ye, D., Xu, W., Lei, L., Sun, Y., Liu, B., Wang, S., and McElroy, M.  
510 B.: Thermodynamic Modeling Suggests Declines in Water Uptake and Acidity of Inorganic Aerosols in Beijing Winter Haze Events during 2014/2015–2018/2019, *Environmental Science & Technology Letters*, 6, 752-760, <https://doi.org/10.1021/acs.estlett.9b00621>, 2019.
- Terrenoire, E., Bessagnet, B., Rouil, L., Tognet, F., Pirovano, G., Létinois, L., Beauchamp, M., Colette, A., Thunis, P., Amann, M., and Menuet, L.: High-resolution air quality simulation over Europe with the chemistry transport model CHIMERE,



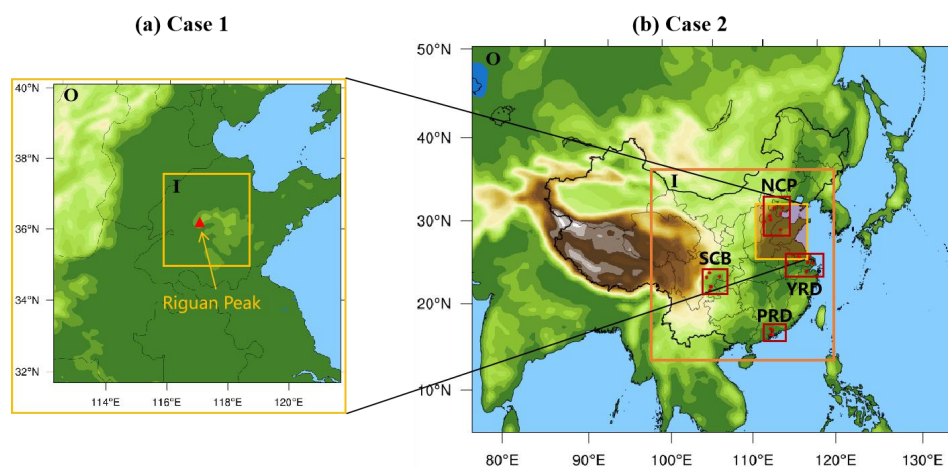
- 515 Geoscientific Model Development, 8, 21-42, <https://doi.org/10.5194/gmd-8-21-2015>, 2015.
- Tie, X.: Assessment of the global impact of aerosols on tropospheric oxidants, *Journal of Geophysical Research*, 110, <https://doi.org/10.1029/2004jd005359>, 2005.
- Tremblay, A., and Leighton, H.: A Three-Dimensional Cloud Chemistry Model, *Journal of Climate & Applied Meteorology*, 25, 652-671, [https://doi.org/10.1016/1352-2310\(96\)00063-5](https://doi.org/10.1016/1352-2310(96)00063-5), 1986.
- 520 Turnock, S. T., Mann, G. W., Woodhouse, M. T., Dalvi, M., O'Connor, F. M., Carslaw, K. S., and Sprackle, D. V.: The Impact of Changes in Cloud Water pH on Aerosol Radiative Forcing, *Geophysical Research Letters*, 46, 4039 - 4048, <https://doi.org/10.1029/2019GL082067>, 2019.
- Twomey: Aerosols, clouds and radiation, *Atmospheric Environment part A: general Topics*, 25, 2435-2442, [https://doi.org/10.1016/0960-1686\(91\)90159-5](https://doi.org/10.1016/0960-1686(91)90159-5), 1991.
- 525 Twomey, S. A., Piepgrass, M., and Wolfe, T. L.: An assessment of the impact of pollution on the global cloud albedo, *Tellus B: Chemical and Physical Meteorology*, 36B, 356-366, <https://doi.org/10.1111/j.1600-0889.1984.tb00254.x>, 1984.
- von Salzen, K., Leighton, H. G., Ariya, P. A., Barrie, L. A., Gong, S. L., Blanchet, J. P., Spacek, L., Lohmann, U., and Kleinman, L. I.: Sensitivity of sulphate aerosol size distributions and CCN concentrations over North America to SO<sub>x</sub> emissions and H<sub>2</sub>O<sub>2</sub> concentrations, *Journal of Geophysical Research: Atmospheres*, 105, 9741-9765, <https://doi.org/10.1029/2000jd900027>, 2000.
- 530 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S., Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P., Marrero-Ortiz, W., Secret, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y., Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation from London Fog to Chinese haze, *Proc Natl Acad Sci U S A*, 113, 13630-13635, <https://doi.org/10.1073/pnas.1616540113>, 2016.
- 535 Wang, H., Shi, G. Y., Zhang, X. Y., Gong, S. L., Tan, S. C., Chen, B., Che, H. Z., and Li, T.: Mesoscale modelling study of the interactions between aerosols and PBL meteorology during a haze episode in China Jing - Jin - Ji and its near surrounding region - Part 2: Aerosols' radiative feedback effects, *Atmos. Chem. Phys.*, 15, 3277-3287, <https://doi.org/10.5194/acp-15-3257-2015>, 2015.
- 540 Wang, J., Li, J., Ye, J., Zhao, J., Wu, Y., Hu, J., Liu, D., Nie, D., Shen, F., Huang, X., Huang, D. D., Ji, D., Sun, X., Xu, W., Guo, J., Song, S., Qin, Y., Liu, P., Turner, J. R., Lee, H. C., Hwang, S., Liao, H., Martin, S. T., Zhang, Q., Chen, M., Sun, Y., Ge, X., and Jacob, D. J.: Fast sulfate formation from oxidation of SO<sub>2</sub> by NO<sub>2</sub> and HONO observed in Beijing haze, *Nature Communications*, 11, 10.1038/s41467-020-16683-x, 2020.



- Wang, S., Zhou, S., Tao, Y., Tsui, W. G., Ye, J., Yu, J. Z., Murphy, J. G., McNeill, V. F., Abbatt, J. P. D., and Chan, A. W. H.:  
545 Organic Peroxides and Sulfur Dioxide in Aerosol: Source of Particulate Sulfate, *Environ Sci Technol*, 53, 10695-10704,  
<https://doi.org/10.1021/acs.est.9b02591>, 2019.
- Wei, Y., Chen, X., Chen, H., Li, J., Wang, Z., Yang, W., Ge, B., Du, H., Hao, J., Wang, W., Li, J., Sun, Y., and Huang, H.:  
IAP-AACM v1.0: a global to regional evaluation of the atmospheric chemistry model in CAS-ESM, *Atmos. Chem. Phys.*, 19,  
8269-8296, <https://doi.org/10.5194/acp-19-8269-2019>, 2019.
- 550 Xie, Y., Dai, H., Zhang, Y., Wu, Y., Hanaoka, T., and Masui, T.: Comparison of health and economic impacts of PM<sub>2.5</sub> and  
ozone pollution in China, *Environ Int*, 130, 104881, <https://doi.org/10.1016/j.envint.2019.05.075>, 2019.
- Yao, S., Wang, Q., Zhang, J., and Zhang, R.: Characteristics of Aerosol and Effect of Aerosol-Radiation-Feedback in Handan,  
an Industrialized and Polluted City in China in Haze Episodes, *Atmosphere*, 12, 670, <https://doi.org/10.3390/atmos12060670>,  
2021.
- 555 Yao, M., Zhao, Y., Hu, M., Huang, D., Wang, Y., Yu, J. Z., and Yan, N.: Multiphase Reactions between Secondary Organic  
Aerosol and Sulfur Dioxide: Kinetics and Contributions to Sulfate Formation and Aerosol Aging, *Environmental Science &  
Technology Letters*, 6, 768-774, <https://doi.org/10.1021/acs.estlett.9b00657>, 2019.
- Ye, C., Xue, C., Zhang, C., Ma, Z., Liu, P., Zhang, Y., Liu, C., Zhao, X., Zhang, W., He, X., Song, Y., Liu, J., Wang, W., Sui,  
B., Cui, R., Yang, X., Mei, R., Chen, J., and Mu, Y.: Atmospheric Hydrogen Peroxide H<sub>2</sub>O<sub>2</sub> at the Foot and Summit of Mt  
560 Tai Variations Sources, *Journal of Geophysical Research: Atmospheres*, 126, <https://doi.org/10.1029/2020JD033975>, 2021.
- Ye, J., Abbatt, J. P. D., and Chan, A. W. H.: Novel pathway of SO<sub>2</sub> oxidation in the atmosphere: reactions with monoterpene  
ozonolysis intermediates and secondary organic aerosol, *Atmos. Chem. Phys.*, 18, 5549-5565, <https://doi.org/10.5194/acp-18-5549-2018>, 2018.
- Yuan, D. M., Ma, X. H.: The severe haze process in 16 - 21 December 2016 and associated atmospheric circulation  
565 anomalies [J]. *Climatic and Environmental Research (in Chinese)*, 22 (6): 757 - 764, <https://doi.org/10.3878/j.issn.1006-9585.2017.17029>, 2017.
- Zhang, L., Gong, S., Zhao, T. L., Zhou, C. H., and Zhang, X. Y.: Development of WRF/CUACE v1.0 model and its  
preliminary application in simulating air quality in China, *Geoscientific Model Development*, <https://doi.org/10.5194/gmd-14-703-2021>, 2021.
- 570 Zhang, X. Y., Wang, Y. Q., Niu, T., Zhang, X. C., Gong, S. L., Zhang, Y. M., and Sun, J. Y.: Atmospheric aerosol  
compositions in China: spatial/temporal variability, chemical signature, regional haze distribution and comparisons with  
global aerosols, *Atmos. Chem. Phys.*, 12, 779-799, <https://doi.org/10.5194/acp-12-779-2012>, 2012.



- Zheng, B., Tong, D., Li, M., Liu, F., Hong, C., Geng, G., Li, H., Li, X., Peng, L., Qi, J., Yan, L., Zhang, Y., Zhao, H., Zheng, Y., He, K., and Zhang, Q.: Trends in China's anthropogenic emissions since 2010 as the consequence of clean air actions, *Atmos. Chem. Phys.*, 18, 14095-14111, <https://doi.org/10.5194/acp-18-14095-2018>, 2018.
- Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.: Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol formation during the January 2013 haze episode in North China, *Atmos. Chem. Phys.*, 14, 2031-2049, <https://doi.org/10.5194/acp-15-2031-2015>, 2015.
- Zhou, C. H., Gong, S., Zhang, X. Y., Liu, H. L., Xue, M., Cao, G. L., An, X. Q., and Che, H. Z.: Towards the improvements of simulating the chemical and optical properties of Chinese aerosols using an online coupled model - CUACE/Aero, *Chemical and physical and meteorology*, <https://doi.org/10.3402/tellusb.v64i0.18965>, 2012.
- Zhou, C. H., Zhang, X. Y., Gong, S., Wang, Y. Q., and Xue, M.: Improving aerosol interaction with clouds and precipitation in a regional chemical weather modeling system, *Atmospheric Chemistry and Physics*, 16, 145-160, <https://doi.org/10.5194/acp-16-145-2016>, 2016.
- Zhou Y., Gong S., Zhou C., Zhang L., He J., Wang Y., Ji D., Feng J., Mo J., Ke H.: A new parameterization of uptake coefficients for heterogeneous reactions on multi-component atmospheric aerosols, *Science of the Total Environment*, 781, <https://doi.org/10.1016/j.scitotenv.2021.146372>, 2021.



**Figure 1. Model nesting domains and target regions: (a) domain for the Case 1. The red triangle is the Mount Tai site, (b) nesting domains for regional assessment for Case 2. Red dots are where the surface observations of air pollutants are. The target four regions are NCP for the North China Plain, YRD for the Yangtze River Delta , PRD for the Pearl River Delta and SCB for the Sichuan Basin.**

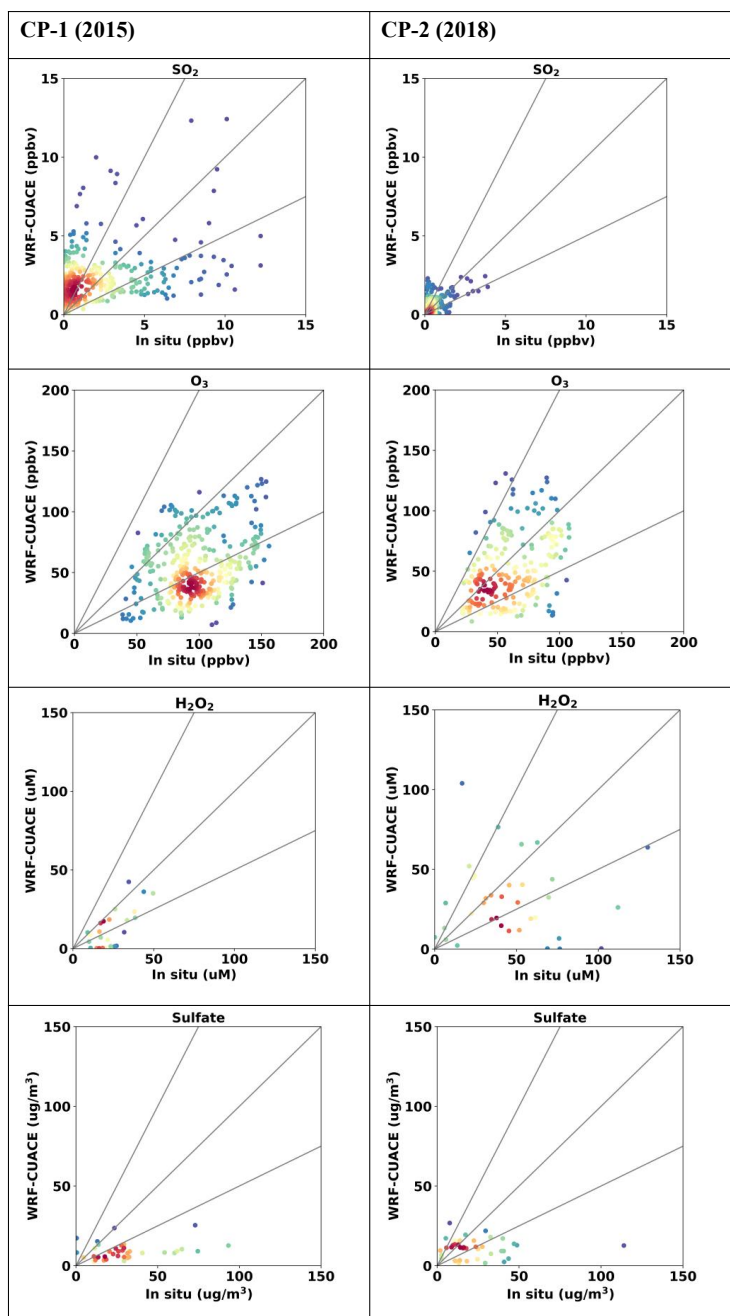
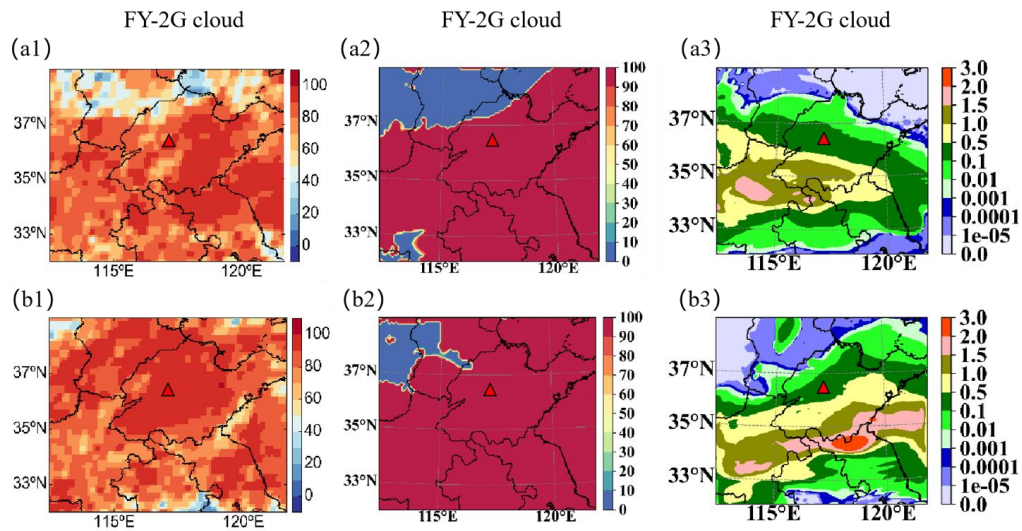


Figure 2. Scatter plots of hourly SO<sub>2</sub>, O<sub>3</sub>, H<sub>2</sub>O<sub>2</sub> and sulfate concentrations between WRF/CUACE and in situ observations at Mount Tai.





**Figure 3.** (a1, b1): the cloud image of FY2G, (a2, b2): the cloud fraction by WRF/CUACE. (a3, b3): the liquid water content by WRF/CUACE (Units:  $\text{g m}^{-2}$ ). (a) is for 8:00 LST on 24 June, (b) is for 8:00 LST on 25 June.

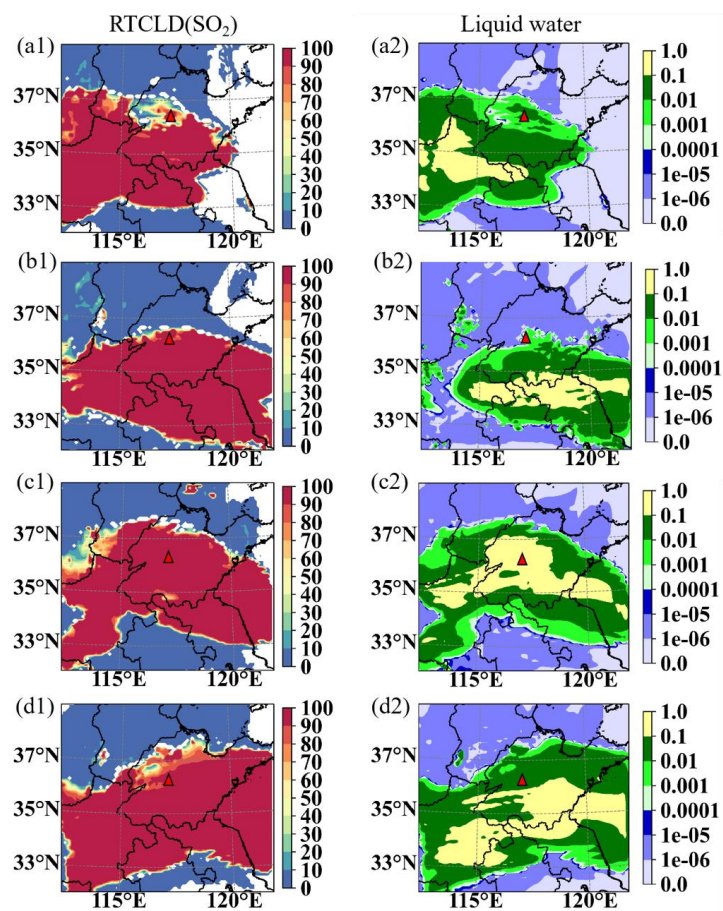
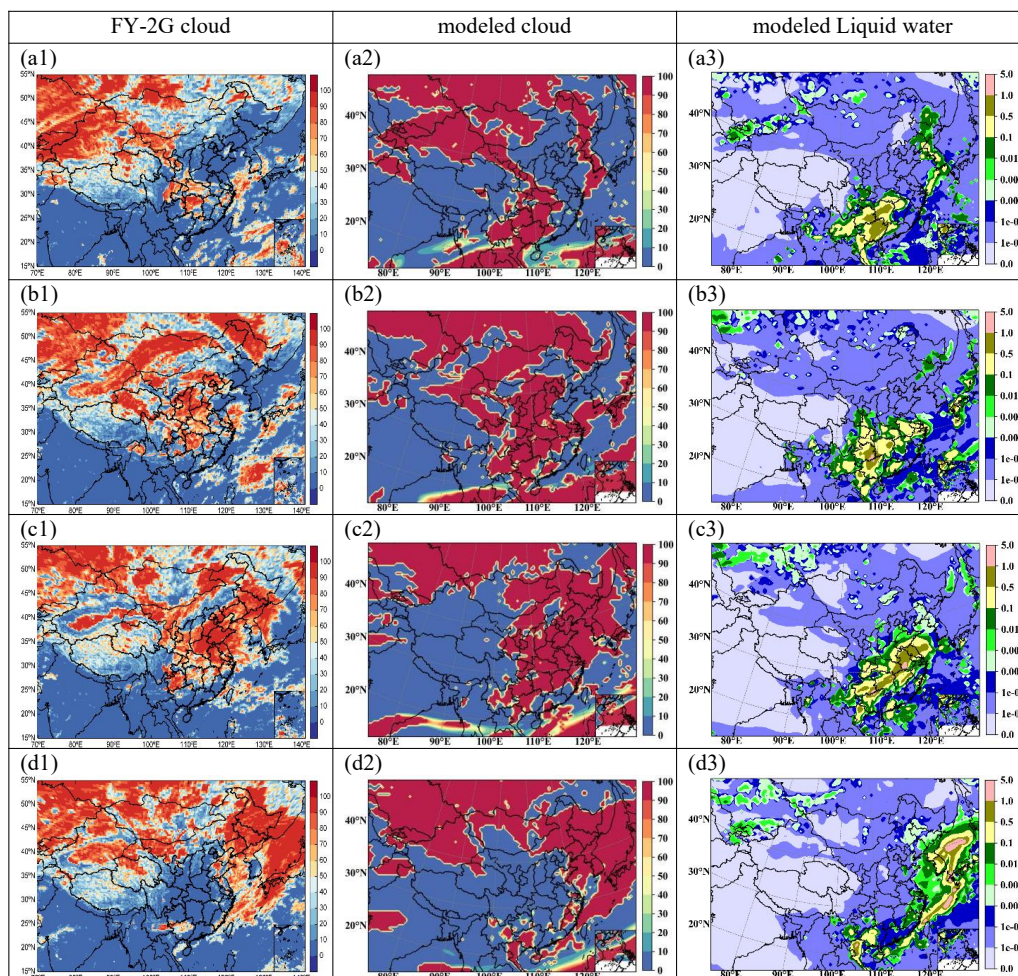
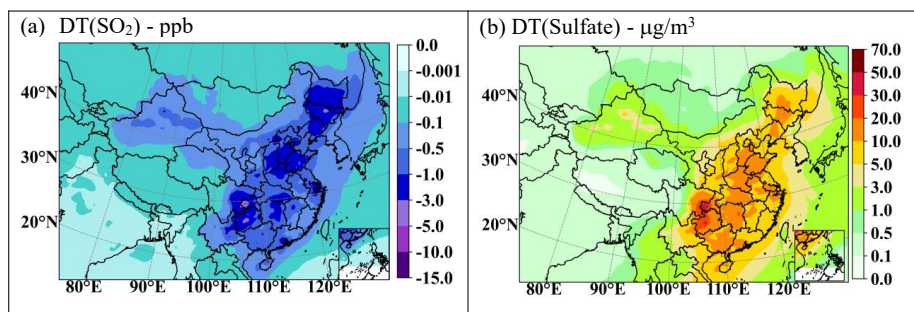


Figure 4. The percentage of  $\text{SO}_2$  oxidized (a1, b1, c1 and d1) (%) and the liquid water content (a2, b2, c2 and d2) (g/kg) by WRF/CUACE, where (a) is for 2:00 LST on 24 June, (b) is for 8:00 LST on 24 June, (c) is for 2:00 LST on 25 June and (d) is for 8:00 LST on 25 June.



**Figure 5.** (a1, b1, c1, d1): the cloud image of FY-2G (%); (a2, b2, c2, d2): the column cloud of WRF/CUACE (%); (a3, b3, c3, d3): the liquid water content of WRF/CUACE ( $\text{g}/\text{m}^2$ ). (a) is for 8:00 LST on 19 Dec., (b) is for 8:00 LST on 20 Dec., (c) is for 8:00 LST on 21 Dec., and (d) is for 8:00 LST on 22 Dec.



**Figure 6.** The mean  $\text{SO}_2$  concentration decreased (a) and sulfate concentration increased (b) by cloud chemistry in Dec. 2016.

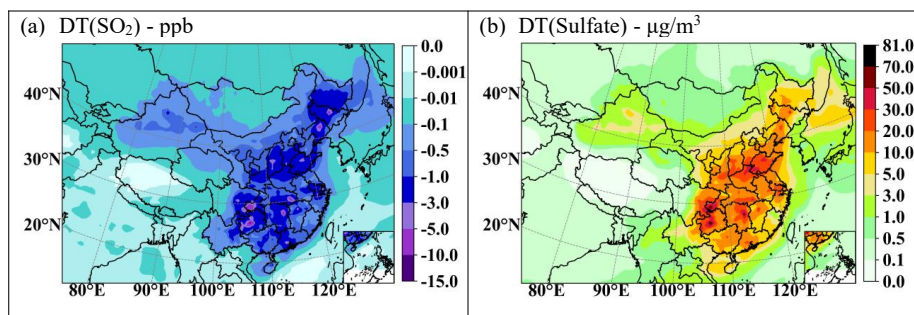


Figure 7. The mean SO<sub>2</sub> concentration decreased (a) and sulfate concentration increased (b) by cloud chemistry from 19-21 Dec., 2016.

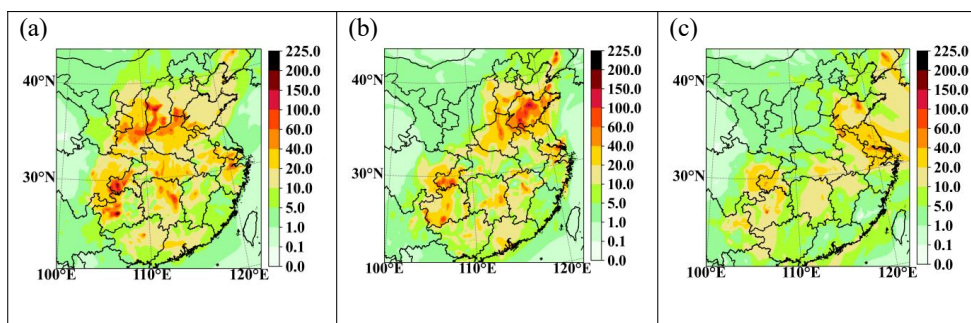


Figure 8. The differences in surface sulfate concentrations between with and without cloud chemistry at 21:00 on 20 Dec. (a), at 17:00 on 21 Dec. (b), and at 12:00 on 22 Dec. (c) ( $\mu\text{g}/\text{m}^3$ ).

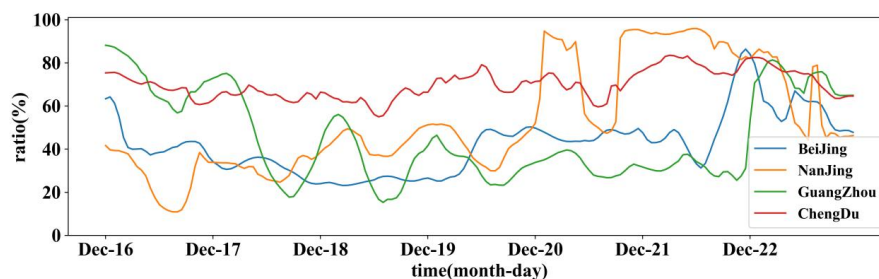


Figure 9. The percentage of SO<sub>2</sub> column concentration reduced due to oxidation in cloud chemistry.

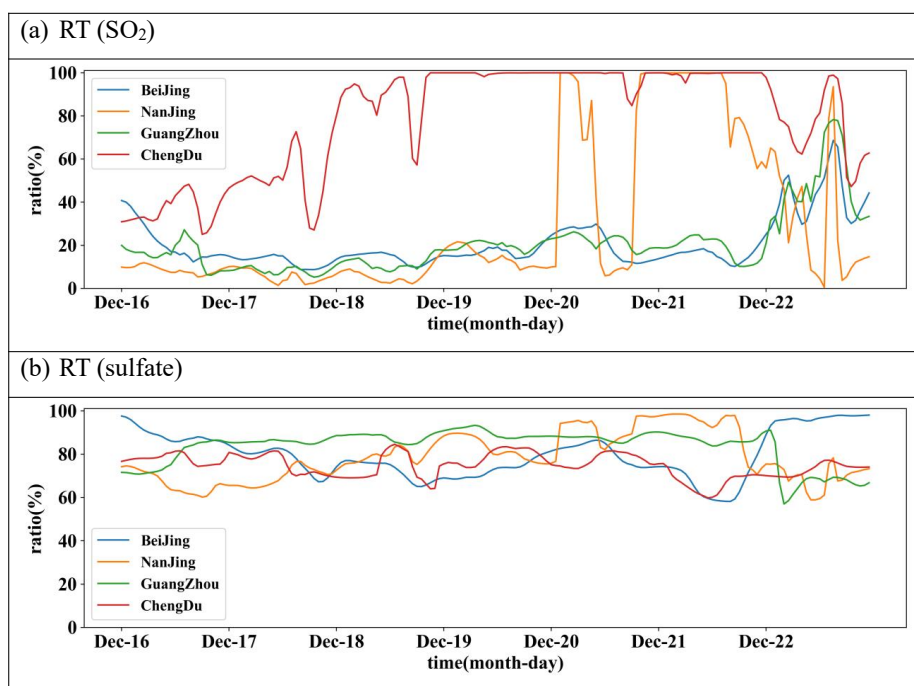


Figure 10. The percentage of surface SO<sub>2</sub> reduced (a) and the surface sulfate increased (b) influenced by cloud chemistry.

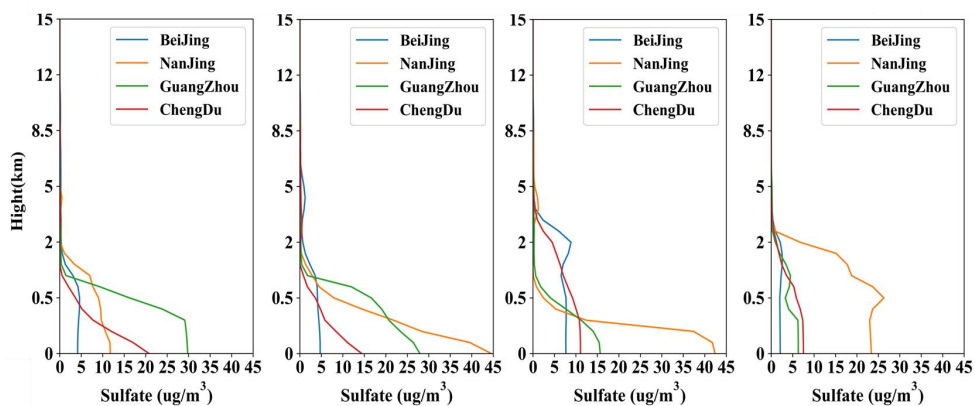


Figure 11. Vertical profiles of sulfate concentration difference DT at 12:00 on 20 Dec., at 21:00 on 20 Dec., at 17:00 on 21 Dec., and at 12:00 on 22 Dec.



**Table 1. Equilibrium Constants for the Parameterization of the Cloud Chemistry in CUACE.**

Equilibrium Relation	Constant Expression	Equilibrium Constant		
		K(298)	a	Unit
$SO_2(g) + H_2O(aq) \leftrightarrow SO_2(aq)$	$K_{HS} = \frac{[SO_2(aq)]}{[SO_2(g)]}$	1.23	3120	$\frac{M}{atm}$
$SO_2(aq) \leftrightarrow H^+ + HSO_3^-$	$K_{1S} = \frac{[H^+][HSO_3^-]}{[SO_2(aq)]}$	$1.7 \times 10^{-2}$	2090	M
$HSO_3^- \leftrightarrow H^+ + SO_3^{2-}$	$K_{2S} = \frac{[H^+][SO_3^{2-}]}{[HSO_3^-]}$	$6.0 \times 10^{-8}$	1120	M
$O_3(g) + H_2O(aq) \leftrightarrow O_3(aq)$	$K_{HO} = \frac{[O_3(aq)]}{[O_3(g)]}$	$1.15 \times 10^{-2}$	2560	$\frac{M}{atm}$
$H_2O_2(g) + H_2O(aq) \leftrightarrow H_2O_2(aq)$	$K_{HP} = \frac{[H_2O_2(aq)]}{[H_2O_2(g)]}$	$9.7 \times 10^4$	6600	$\frac{M}{atm}$

**Table 2. physics parameterization schemes in WRF.**

Physical management	Parameterization	References
microphysics scheme	Lin	Lin et al. (1983)
shortwave radiation	Goddard	Chou and Suarez (1994)
longwave radiation	RRTM	Mlawer et al. (1997)
land surface scheme	Noah	Chen and Dudhia (2001)
boundary layer scheme	MYJ	Janjić (1994)
cumulus scheme	Grell-3D	Grell (1993)

**Table 3. Statistics for SO<sub>2</sub>, O<sub>3</sub>, H<sub>2</sub>O<sub>2</sub> and sulfate in cloud chemistry at Mount Tai site.**

	Obs.	Mod		R	RAD(%)	NMB(%)	
CP-1	SO <sub>2</sub> (ppbv)	2.16	2.32	SO <sub>2</sub>	0.34	-3.4	7.1
	O <sub>3</sub> (ppbv)	97.79	55.27	O <sub>3</sub>	0.33	27.8	-43.5
	H <sub>2</sub> O <sub>2</sub> (uM)	26.52	16.83	H <sub>2</sub> O <sub>2</sub>	0.78	22.4	-36.6
	Sulfate(μg/m <sup>3</sup> )	31.68	9.20	Sulfate	0.32	55.0	-71.0
CP-2	SO <sub>2</sub> (ppbv)	0.56	0.63	SO <sub>2</sub>	0.47	-6.1	12.9
	O <sub>3</sub> (ppbv)	60.68	50.97	O <sub>3</sub>	0.40	8.7	-16.0
	H <sub>2</sub> O <sub>2</sub> (uM)	46.92	32.36	H <sub>2</sub> O <sub>2</sub>	0.06	18.4	-29.6
	Sulfate(μg/m <sup>3</sup> )	28.08	11.40	Sulfate	0.54	42.2	-59.4



**Table 4. Statistical metrics for meteorology in NCP, YRD, PRD and SCB for Dec. 16-21 and the whole Dec. in 2016.**

		Obs.		Mod			R		NMB(%)	
		16-21	Dec.	16-21	Dec.		16-21	Dec.	16-21	Dec.
NCP	T2(°C)	0.99	1.11	2.83	2.05	T2	0.70	0.84	187.3	84.9
	RH2(%)	78.83	68.31	52.30	48.75	RH2	0.54	0.64	-33.7	-28.6
	WS10(m/s)	1.47	1.69	1.67	2.16	WS10	0.49	0.54	14.1	27.5
YRD	T2(°C)	9.24	7.99	9.51	8.40	T2	0.94	0.96	2.9	5.1
	RH2(%)	79.22	75.62	73.81	72.98	RH2	0.86	0.85	-6.8	-3.5
	WS10(m/s)	2.16	2.27	2.78	2.99	WS10	0.74	0.76	28.7	31.9
PRD	T2(°C)	18.31	17.27	18.98	17.93	T2	0.93	0.92	3.6	3.8
	RH2(%)	72.19	70.43	64.31	65.35	RH2	0.76	0.68	-10.9	-7.2
	WS10(m/s)	1.78	2.36	2.02	3.24	WS10	0.67	0.72	13.6	37.1
SCB	T2(°C)	10.18	9.70	10.46	9.99	T2	0.74	0.75	2.8	3.1
	RH2(%)	81.63	79.92	74.09	71.26	RH2	0.66	0.60	-9.2	-10.8
	WS10(m/s)	1.09	1.27	1.63	1.92	WS10	0.49	0.36	49.1	50.5

**Table 5. Statistical metrics for hourly SO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> in NCP, YRD, PRD and SCB for Dec. 16-21 and the whole Dec. in 2016.**

		Obs.(µg/m <sup>3</sup> )		Mod(µg/m <sup>3</sup> )		R		NMB(%)	
		16-21	Dec.	16-21	Dec.	16-21	Dec.	16-21	Dec.
NCP	SO <sub>2</sub>	41.99	61.45	49.98	39.99	0.60	0.48	46.3	-15.6
	O <sub>3</sub>	8.75	7.41	7.42	10.89	0.47	0.60	-15.3	-32.4
	PM <sub>2.5</sub>	351.33	182.08	194.75	110.79	0.30	0.62	-48.2	-35.7
YRD	SO <sub>2</sub>	21.81	16.30	15.82	14.90	0.61	0.45	-25.3	-32.8
	O <sub>3</sub>	31.25	14.37	9.25	22.11	0.33	0.68	-54.0	-45.5
	PM <sub>2.5</sub>	70.26	82.90	119.14	84.23	0.70	0.73	18.0	19.3
PRD	SO <sub>2</sub>	13.64	24.01	24.02	16.98	0.32	0.39	76.1	11.9
	O <sub>3</sub>	45.74	56.27	56.51	57.36	0.84	0.80	23.0	13.9
	PM <sub>2.5</sub>	55.66	83.56	83.78	77.54	0.84	0.39	50.1	51.5
SCB	SO <sub>2</sub>	19.98	10.03	12.43	8.80	0.49	0.19	-49.8	-47.4
	O <sub>3</sub>	21.97	49.01	45.27	54.19	0.20	0.47	123.1	97.4
	PM <sub>2.5</sub>	135.56	90.99	117.03	70.99	0.27	0.28	-32.9	-29.1



**Table 6. Statistical metrics for hourly SO<sub>2</sub> and PM<sub>2.5</sub> in NCP, YRD, PRD and SCB for Dec. 19-21, respectively.**

19-21 Dec.		SO <sub>2</sub>				PM <sub>2.5</sub> (µg/m <sup>3</sup> )			
		NCP	YRD	PRD	SCB	NCP	YRD	PRD	SCB
Obs.(µg/m <sup>3</sup> )		41.99	21.81	13.64	19.98	351.33	70.26	55.66	135.56
Mod (µg/m <sup>3</sup> )	CLD	49.98	15.82	24.02	12.43	194.75	119.14	83.78	117.03
	CCLD	65.61	25.68	30.81	15.84	166.60	63.82	77.00	72.55
R	CLD	0.60	0.61	0.32	0.49	0.30	0.70	0.84	0.27
	CCLD	0.61	0.55	0.31	0.34	0.31	0.77	0.83	0.30
NMB(%)	CLD	46.4	-25.3	76.1	-49.8	-48.2	18.0	50.1	-32.9
	CCLD	56.3	17.7	125.9	-20.7	-52.6	-9.2	38.3	-46.5

**Table 7. Statistical metrics for hourly SO<sub>2</sub> and PM<sub>2.5</sub> in NCP, YRD, PRD and SCB for Dec., respectively.**

Dec.		SO <sub>2</sub>				PM <sub>2.5</sub>			
		NCP	YRD	PRD	SCB	NCP	YRD	PRD	SCB
Obs.(µg/m <sup>3</sup> )		47.36	22.17	15.18	16.72	172.18	70.59	51.18	100.11
Mod (µg/m <sup>3</sup> )	CLD	39.99	14.90	16.98	8.80	110.79	84.23	77.54	70.99
	CCLD	44.68	19.51	20.71	13.59	93.58	71.81	70.11	55.91
R	CLD	0.48	0.45	0.39	0.19	0.62	0.73	0.39	0.28
	CCLD	0.50	0.35	0.34	0.16	0.65	0.77	0.38	0.33
NMB(%)	CLD	-15.6	-32.8	11.9	-47.4	-35.7	19.3	51.5	-29.1
	CCLD	-5.7	-12.0	36.4	-18.8	-45.7	1.7	37.0	-44.2

**Table 8. Statistical metrics for hourly SO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> in all selected sites for Dec. 16-21 and the whole Dec..**

SUM		19-21		Dec.	
		SO <sub>2</sub>	PM <sub>2.5</sub>	SO <sub>2</sub>	PM <sub>2.5</sub>
Obs.(µg/m <sup>3</sup> )		25.68	167.95	27.01	32.19
Mod (µg/m <sup>3</sup> )	CLD	29.21	114.85	21.35	87.56
	CCLD	35.75	99.51	25.85	74.05
R	CLD	0.66	0.67	0.64	0.56
	CCLD	0.65	0.70	0.63	0.57
NMB (%)	CLD	13.8	-31.6	-21.0	-16.3
	CCLD	39.3	-40.8	-4.3	-29.2