



1    Marked recent declines in boron in Baltic Sea cod otoliths – a bellwether  
2                                   of incipient acidification in a vast hypoxic system?

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4                                   Karin E. Limburg<sup>1,2</sup>, Yvette Heimbrand<sup>2</sup>, Karol Kuliński<sup>3</sup>

5    <sup>1</sup>Department of Environmental Biology, State University of New York College of Environmental  
6                                   Science and Forestry, Syracuse, NY USA

7    <sup>2</sup>Department of Aquatic Resources, Swedish University of Aquatic Sciences, Uppsala, Sweden

8    <sup>3</sup>Department of Marine Chemistry and Biochemistry, Institute of Oceanology of the Polish  
9                                   Academy of Sciences, Sopot, Poland

10    *Correspondence to:* Karin E. Limburg (klimburg@esf.edu)

11    **Abstract:** Ocean acidification is spreading globally as a result of anthropogenic CO<sub>2</sub> emissions,  
12    but the Baltic Sea has until recently been thought to be relatively well-buffered by terrigenous  
13    inputs of alkalinity from its watershed. We discovered a 3- to 5-fold decline in boron (as B:Ca)  
14    in otoliths of eastern Baltic Sea cod (EBC) between the late 1990s and 2021. B:Ca is positively  
15    proportional to pH in carbonates, as B in the form of borate is taken up in the CaCO<sub>3</sub> matrix.  
16    Examining a time series of EBC otoliths, we found varying levels of B:Ca since the 1980s, with  
17    the most recent years at an all-time low during this period. This trend correlates with declines in  
18    pH and dissolved oxygen, but not with changes in salinity. We examined possible physiological  
19    influences on B:Ca by including a collection of healthy Icelandic cod as an out-group. Icelandic  
20    cod otoliths showed strongly positive correlations of B:Ca with physiologically regulated P:Ca;  
21    this was not the case for EBC. Finally, B:Ca in EBC otoliths is anti-correlated to some extent  
22    with Mn:Mg, a proposed proxy for hypoxia exposure. This negative relationship is hypothesized  
23    to reflect the dual phenomena of hypoxia and acidification as a result of decomposition of large  
24    algal blooms. Taken together, the otolith biomarkers Mn:Mg and B:Ca suggest a general increase  
25    in both hypoxia and acidification within the Baltic intermediate and deep waters in the last  
26    decade reflected in cod otoliths.

27    **Keywords:** Baltic Sea cod, otolith chemistry, boron, pH, hypoxia

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## 30 **1 Introduction**

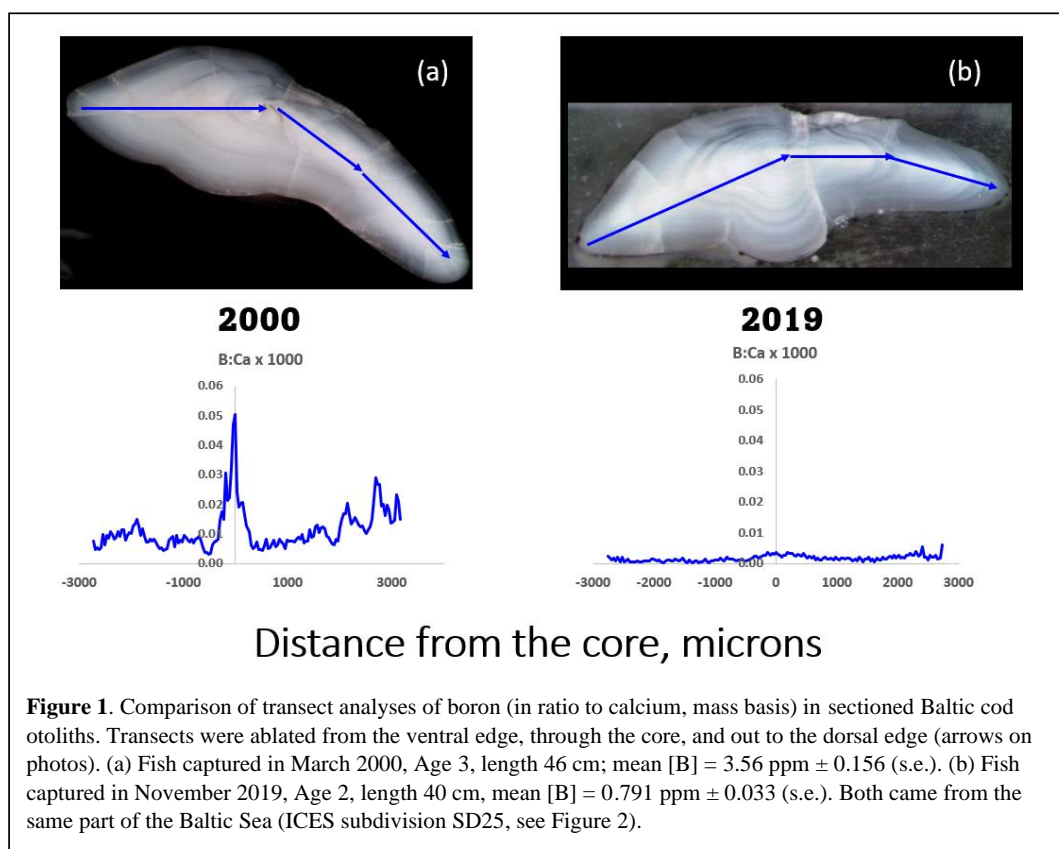
31 Ocean acidification, a known result of greenhouse gas emissions of CO<sub>2</sub> loading into the  
32 oceans, is highly variable in its distribution in the world's oceans (Doney et al., 2009). This is  
33 true in shallow coastal seas such as the brackish Baltic Sea, where in some locations surface  
34 pCO<sub>2</sub> values in winter greatly exceed other coastal regions (Reusch et al., 2018), but buffering  
35 can occur from river inputs bearing proton acceptors weathered from rock and soil (Müller et al.,  
36 2016; Kuliński et al., 2017). Alkalinity, defined as the sum of acid neutralizing compounds, is a  
37 complicated mix in the Baltic Sea (Kuliński et al., 2022). In the northern Baltic, i.e., the  
38 Bothnian Sea, total alkalinity ( $A_T$ ) is low as it is largely driven by  $A_T$ -poor riverine inputs from  
39 the Scandinavian Peninsula that is built mostly of granite. In the southern and central Baltic,  $A_T$   
40 is higher due to salt inputs from the North Sea, influence of  $A_T$ -rich continental rivers, and an  
41 additional contribution from internal cycling of carbon as algal blooms decompose at anoxic  
42 conditions (Kuliński et al., 2022). Despite buffering from alkalinity sources, pH is highly  
43 variable in space and time with its recent tendency to decline at greater depths in many parts of  
44 the Baltic Sea, which is associated with the ongoing eutrophication and higher vertical export of  
45 organic matter that leads also to worsening deoxygenation (Kuliński et al., 2022).

46 The population of the Eastern Baltic cod (*Gadus morhua*, hereafter referred to as EBC)  
47 has been severely impacted by a number of factors that include past overfishing, hypoxia,  
48 parasite infections, and seal predation (Eero et al., 2020), leading to reduced growth and age at  
49 maturity (Eero et al., 2016; ICES, 2019). Cod exposed to hypoxia were shown to have lower  
50 growth rates and worsened condition (Limburg & Casini, 2018, 2019). Hypoxia exposure was  
51 quantified by means of a proxy developed from analyzing manganese concentrations in otoliths  
52 (ear-stones) of fish (Limburg et al., 2015), in ratio to magnesium which corrects to some extent  
53 for growth influences on Mn uptake (Limburg et al., 2018; Limburg & Casini, 2018). Briefly,  
54 otoliths, the calcified structures that form part of the hearing/balance system in teleost fishes, are  
55 sectioned and then analyzed by ablating micro-transsects along the major growth axis and  
56 analyzing by mass spectrometry (see Methods). The resulting elemental data are lifetime  
57 concentration histories, being incorporated at the time of exposure. The use of otolith  
58 manganese to track hypoxia exposure is one of the emerging biomarkers in fisheries ecology  
59 (Reis-Santos et al., 2022).



60 Recently, with acquisition of more sensitive instrumentation, we began to experiment  
61 with quantifying elements having sub-ppm concentrations in cod otoliths, including the trace  
62 element boron. Boron in seawater generally correlates with salinity (Kuliński et al., 2017)  
63 predominantly in the form of weak boric acid ( $H_3BO_3$ ) at standard seawater salinity (35 PSU)  
64 and pH of 8. However, the dissociated anion borate  $B(OH)_4^-$  is positively, nonlinearly correlated  
65 with pH (Yu et al., 2007). Borate is the dominant form incorporated into carbonates such as  
66 limestones, foraminifera, and corals in proportion to pH (Yu et al., 2007; Hönisch et al., 2012;  
67 Levin et al., 2015) and is often confirmed by  $\delta^{11}B$  analysis, since it has a large, negative  
68 fractionation compared to boric acid. However, even bulk boron, as B:Ca ratios, has been found  
69 to correlate positively (and linearly) with pH in foraminifera (Yu et al., 2007), albeit in a species-  
70 specific manner.

71 We measured sub-ppm to low-ppm concentrations of B (ca. 0.5 – 5 ppm, see Methods  
72 section) in otoliths of EBC captured in 2000 that showed what appeared to be seasonal variations





73 (example shown in Figure 1(a)), albeit not in synchrony with other elements that have proven  
74 useful to age cod (Heimbrand et al., 2020). It was therefore surprising when we began to analyze  
75 otoliths of EBC captured in 2019 and discovered greatly reduced concentrations (Figure 1(b)).  
76 The measurements were repeated for verification.

77 With such differences suggestive of change, we returned to an archive of otoliths used for  
78 another project (Limburg and Casini, 2018), as well as adding previously un-analyzed otoliths, to  
79 develop a time series of otolith B:Ca to document any temporal trends that could be correlated  
80 with pH, salinity, or other environmental factors. We posed the following questions:

81

- 82 1. What are the temporal trends in otolith boron (as B:Ca ratios)?
- 83 2. How do B:Ca ratios correlate to trends in water chemistry, particularly salinity and pH?
- 84 3. If pH declines are associated with hypoxia, do we see a correlation of B:Ca with our hypoxia  
85 proxy, manganese:magnesium (Mn:Mg)?
- 86 4. If there are any physiological influences, do we see a correlation of B:Ca with elemental ratios  
87 known to be under physiological control, e.g., phosphorus:calcium (P:Ca)?

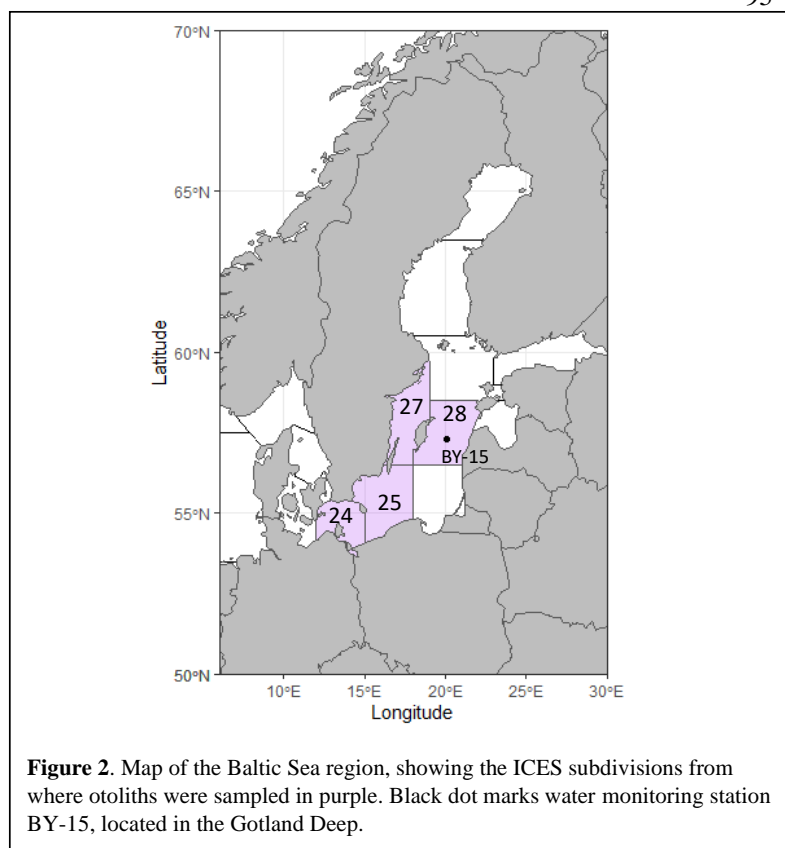
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## 89 **2 Materials and Methods**

### 90 **2.1 Otolith collections and microchemical analysis**



91 Otoliths of Baltic cod (N = 156) were obtained from both fishery-independent and  
92 fishery-dependent surveys conducted by the Swedish Fisheries Board and its successor, the  
93 Department of Aquatic Resources, Swedish University of Agricultural Sciences, collected from  
94 ICES sub-divisions (SD) 24, 25, 27, and 28 (Figure 2). These were measured, weighed, and



95 subsequently embedded in epoxy (Struers EpoFix) and sectioned transversely to expose the core. They were smoothed with a succession of lapping films to 3 microns.

In addition, cod otoliths (N = 14) were obtained from fisheries-independent surveys from Iceland (ICES SD 5a), as an out-group for studies of possible physiological correlations. This part of the North Atlantic does not have hypoxia or acidification issues in the deep waters. The Icelandic

114 otoliths were prepared and analyzed in similar fashion.

115 Chemical analyses were performed by laser ablation inductively coupled plasma mass  
116 spectrometry (LA-ICP-MS) at the Analytical and Technical Services group at SUNY ESF.  
117 Otoliths were ablated with a 192-nm laser ablation unit (Teledyne Cetac Excite 2) along transects  
118 as illustrated in Figure 1; the ablated material was transported via an Ar-He carrier gas mixture  
119 into a Thermo iCAP TQ plasma mass spectrometer where isotopes were quantified ( $^{11}\text{B}$ ,  $^{25}\text{Mg}$ ,  
120  $^{31}\text{P}$ ,  $^{43}\text{Ca}$ ,  $^{55}\text{Mn}$ ,  $^{63}\text{Cu}$ ,  $^{66}\text{Zn}$ ,  $^{88}\text{Sr}$ ,  $^{127}\text{I}$ ,  $^{138}\text{Ba}$ , and  $^{208}\text{Pb}$ ). Daily performance of the system was  
121 optimized with NIST 612 glass standard and U.S. Geological Survey standards MACS-3,



122 MAPS-4, and MAPS-5 were used as standards for calibration and drift correction. The MAPS  
123 (abbreviation for Micro-Analytical Phosphate Standard) standards, while having good matching  
124 to the otolith matrix, were slightly softer and were therefore ablated at a higher energy (25%  
125 power, 2.03 J/cm<sup>2</sup> fluence) than the otoliths (20% power, 1.62 J/cm<sup>2</sup>) to achieve an otolith  
126 calcium value around 38 weight percent. Transects were made with 110 µm-diameter ablation  
127 spots at a speed of 7 µm/sec. Once data were worked up, annual mean values were parsed by the  
128 age determination technique in Heimbrand et al. (2020), i.e., averages within annual growth rings  
129 were calculated. These were matched to corresponding calendar years.

## 130 **2.2 Water data**

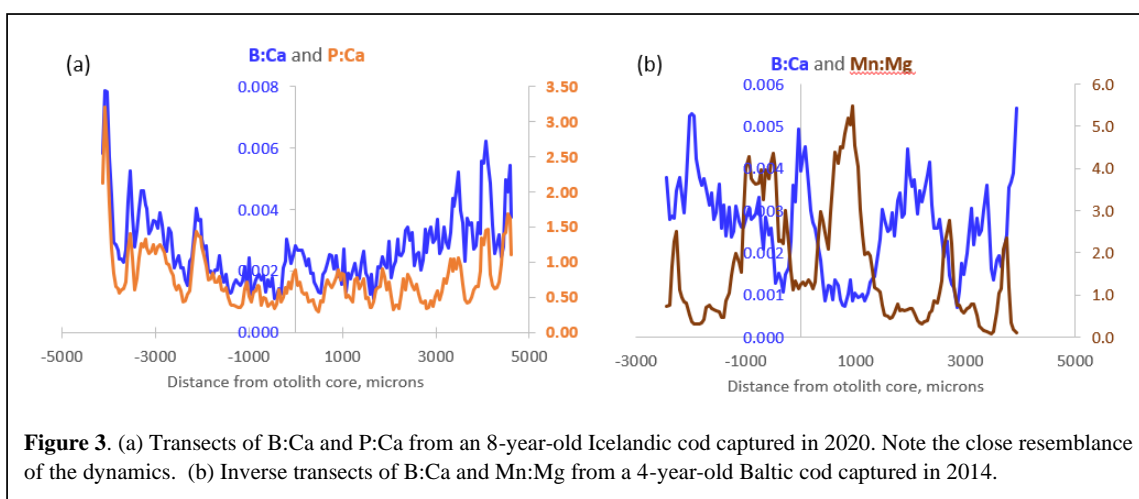
131 Water properties data (temperature (T, °C), salinity (S, psu), dissolved oxygen (DO,  
132 mL/L), pH, and total alkalinity (A<sub>T</sub>, mmol/L)) were downloaded from the Swedish  
133 Meteorological and Hydrological Institute's database, SHARKWeb (<https://sharkweb.smhi.se/>).  
134 Station BY-15, Gotland Deep, was selected as being representative of the central Baltic (Figure  
135 2) and having one of the longest and most detailed timeseries. We selected depths that  
136 corresponded to where cod had been found in fishery-independent surveys (Casini et al., 2021).  
137 Between 1985-1995, we averaged water data from 30-60 meters, and from 1996 onward,  
138 between 40-75 meters. Annual mean water values were computed and matched to corresponding  
139 otolith chemistry data (annual means of Element:Ca data parsed to calendar years as described  
140 above).

## 141 **2.3 Analysis**

142 Visualization and statistical analyses were performed in Excel, R (R Core Team 2022),  
143 and JMP version 17.0 (JMP 2022). Statistical analyses were separated into examination of  
144 relationships of B:Ca values to potential environmental drivers (primarily pH and S, but also A<sub>T</sub>,  
145 DO, and T) and to the internal variables P:Ca and Mn:Mg. P:Ca was found to correlate highly  
146 with B:Ca in Icelandic cod otoliths (Figure 3(a)). Otolith P:Ca is known to be under  
147 physiological influence (Thomas et al., 2017; Heimbrand et al., 2020). In cod, higher values of  
148 P:Ca occur during the growing season and are thus hypothesized to be associated with growth  
149 and other activities. Otolith Mn:Mg, on the other hand, is our proxy for hypoxia exposure  
150 (Limburg and Casini, 2018), and often correlated negatively with B:Ca (Figure 3(b)) - consistent  
151 with co-occurrence of reduced conditions in water (and presence of bioavailable ionic Mn<sup>2+</sup>) and



152 high  $p\text{CO}_2$  and thus low pH, both being the effect of enhanced organic matter remineralization  
153 and/or poor ventilation of the water layers occupied by cod. For the environmental drivers,  
154 graphical analysis was performed followed by mixed linear models and visualizations, with each  
155 fish's age (nested within fish ID) being treated as a random effect. For the internal variables, the  
156 P:Ca and Mn:Mg transect data were regressed on corresponding B:Ca for each fish, and the



**Figure 3.** (a) Transects of B:Ca and P:Ca from an 8-year-old Icelandic cod captured in 2020. Note the close resemblance of the dynamics. (b) Inverse transects of B:Ca and Mn:Mg from a 4-year-old Baltic cod captured in 2014.

157 slopes, intercepts and  $R^2$  values were calculated. These were examined for trends by decade  
158 (1980s, 1990s, 2000s, and 2010s) as well as compared with the Icelandic out-group. Positive  
159 correlations with P:Ca would suggest a physiological influence on B:Ca incorporation, whereas  
160 negative correlations with Mn:Mg would suggest a linkage between B:Ca as an indicator of pH  
161 and our indicator of hypoxia exposure (cf. Cavole, 2021).

162

### 163 **3 Results**

#### 164 **3.1. Trends in water variables**

165 A principal components analysis of the water variables showed three groupings: (1) DO  
166 and pH that were opposite to (2) salinity and  $A_T$ , and (3) temperature was on a separate axis  
167 (Figure S1). Examination of the long-term time series of data from Station BY-15 shows that  
168 while temperature generally increased (Figure S2(a)), the other water variables showed very  
169 nonlinear patterns (Figures S2(b-e)). Alkalinity is in part a function of salinity; dividing  $A_T$  by S  
170 produces a time series showing a dramatic state change around 1990 (Figure S2(f)), particularly

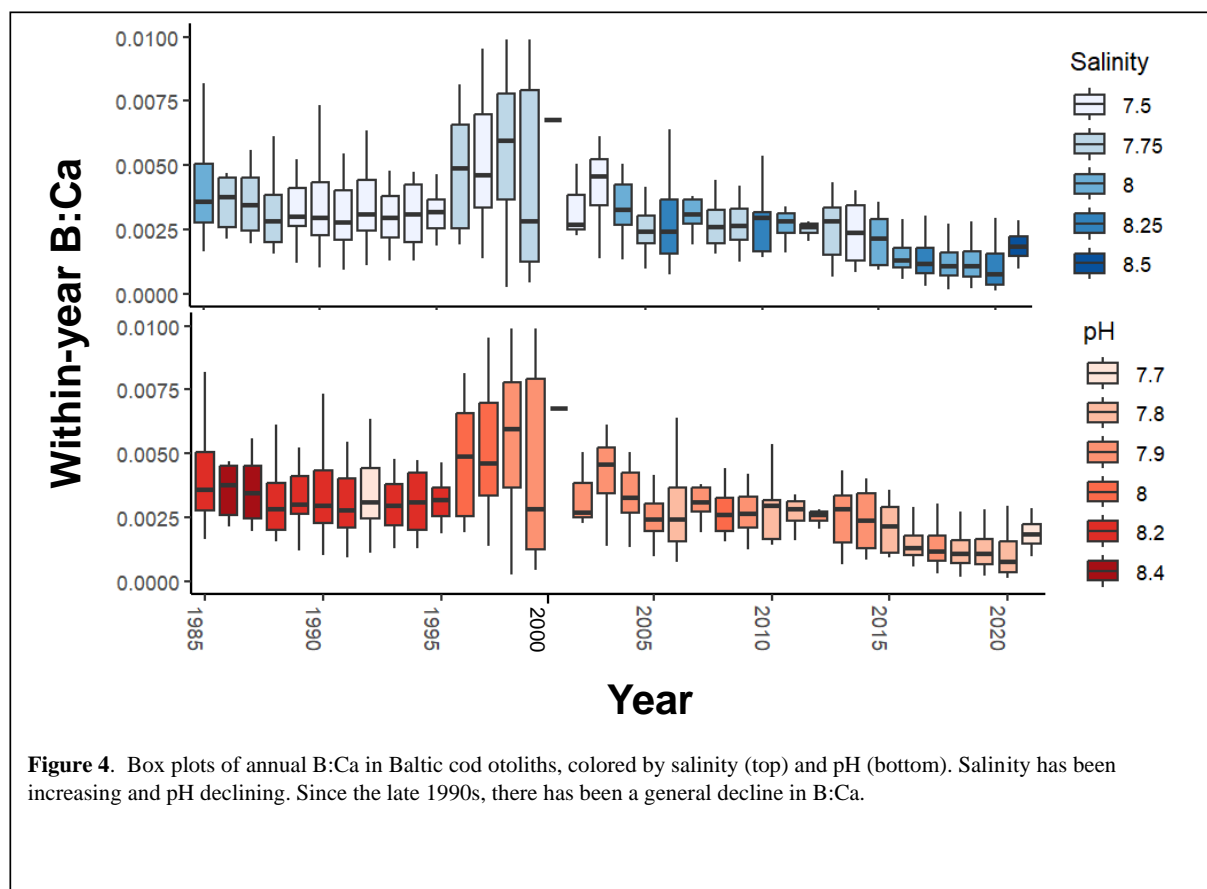


171 in the water layers occupied by cod. Despite the increasing  $A_T$ , pH has declined more or less  
172 monotonically at midwater depths (40-75 m, Figure S2(c)).

### 173 3.2 Trends in otolith B:Ca with salinity and pH

174 Over the period 1985-2019, box plots of B:Ca by year show a pattern of increase toward  
175 a maximum, albeit with great variation, in the late 1990s, followed by gradual decline (Figure 4).  
176 Visual comparisons with salinity and pH show that neither explain all the variation in B:Ca, but  
177 since 2000, pH and B:Ca both trend downward whereas salinity has increased. The increase in  
178 B:Ca in 2021 is from age-0, which is elevated relative to older ages (cf. Figure 1(a)).

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182 **3.3. Effects of water parameters on otolith B:Ca**

183 **3.3.1. Univariate effects of water parameters on otolith B:Ca**

184 Repeated measures regressions of otolith B:Ca on each of the five water variables was  
 185 conducted to examine effect sizes and significance. Repeated measures regressions showed  
 186 significant effects of each of the five water variables on otolith B:Ca (Table 1).  $A_T$  had the  
 187 largest effect and was negative (Table 1). Dissolved oxygen and pH both had positive slopes, and  
 188 temperature and salinity both were negative. All univariate analyses were highly significant.  
 189 Ages of individual fish were not significant as measured by Wald's p-value in restricted  
 190 maximum likelihood (REML) analysis.

**Table 1.** Repeated measures univariate regressions of otolith B:Ca on water variables, ordered by goodness of fit (-2 log likelihood criterion; lowest is best).

Variable	Parameter slope	R <sup>2</sup>	df	F ratio	p	-2 Log Likelihood
Total Alkalinity	-9.42E-03	0.24	647	158.8	< 0.0001	-6358
Dissolved Oxygen	1.00E-03	0.21	615	142.2	< 0.0001	-6341
Salinity	-3.00E-03	0.22	565	133.5	< 0.0001	-6336
Temperature	-8.30E-04	0.13	639	58.7	< 0.0001	-6269
pH	5.10E-03	0.15	633	109.4	< 0.0001	-6189

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192 **3.3.2. Two-factor tests of water parameters on otolith B:Ca**

193 Because of our hypothesized positive relationships of salinity and pH on B:Ca, as well as  
 194 the large effect size of  $A_T$ , we examined pairwise models of these independent variables on  
 195 otolith B:Ca (Table 2). Aside from an overall model, these were filtered by decade to examine  
 196 finer scale trends.

**Table 2.** Repeated measures 2-factor models of effects on otolith B:Ca, by decade. Numbers are estimated coefficients. Red = significant at  $p < 0.01$ . Blue = significant at  $p < 0.05$ . Negative 2 Log Likelihood is a metric of goodness-of-fit, with smaller numbers indicating better fit.

Parameter	1980s	1990s	2000s	2010s	Overall
1. pH and Alkalinity:					
pH	0.00128	-0.00774	-0.00375	-0.00121	-0.00385



Alkalinity	0.00873	-0.00538	-0.01840	-0.02585	-0.01585
-2 Log Lik.	-399.2	-1231.8	-896	-3979.4	-6225.5
2. pH and Salinity:					
<b>Parameter</b>	<b>1980s</b>	<b>1990s</b>	<b>2000s</b>	<b>2010s</b>	<b>Overall</b>
pH	0.00151	-0.00226	-0.00983	0.00378	0.00225
Salinity	0.00322	0.00365	-0.00347	-0.00049	-0.00208
-2 Log Lik.	397.9	-1231.6	-894.8	-3948.2	-6201.6
3. Alkalinity and Salinity:					
<b>Parameter</b>	<b>1980s</b>	<b>1990s</b>	<b>2000s</b>	<b>2010s</b>	<b>Overall</b>
Alkalinity	-0.007269	0.002042	-0.01253	-0.027822	-0.006591
Salinity	0.003768	0.004802	-0.000171	0.000478	-0.001274
-2 Log Lik.	-401.4	-1357.9	-891.9	-3977.6	-6354.8

197

198 In both pairwise tests with pH, this parameter had positive slopes in the 1980s and 2010s  
 199 but negative slopes in the 1990s and 2000s. Alkalinity had mostly a negative relation to otolith  
 200 B:Ca. Salinity had a positive effect in the 1980s and 1990s but became negative in the 2000s and  
 201 2010s. Salinity overall had a negative effect on B:Ca, in contrast to expectations. We note the  
 202 strong, negative effect of pH in the 1990s when paired with  $A_T$ , and the strong, positive effect of  
 203 pH in the 2010s when paired with salinity. Age, when treated as a random variable nested within  
 204 individual fish (by fish ID), was again not significant. However, in a separate analysis (Figure  
 205 S3), age (not treated as a random variable) and otolith B:Ca were inversely related in the 1980s,  
 206 1990s, and 2000s, and had a weakly positive trend in the 2010s.

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### 208 3.4. Internal parameters and otolith B:Ca

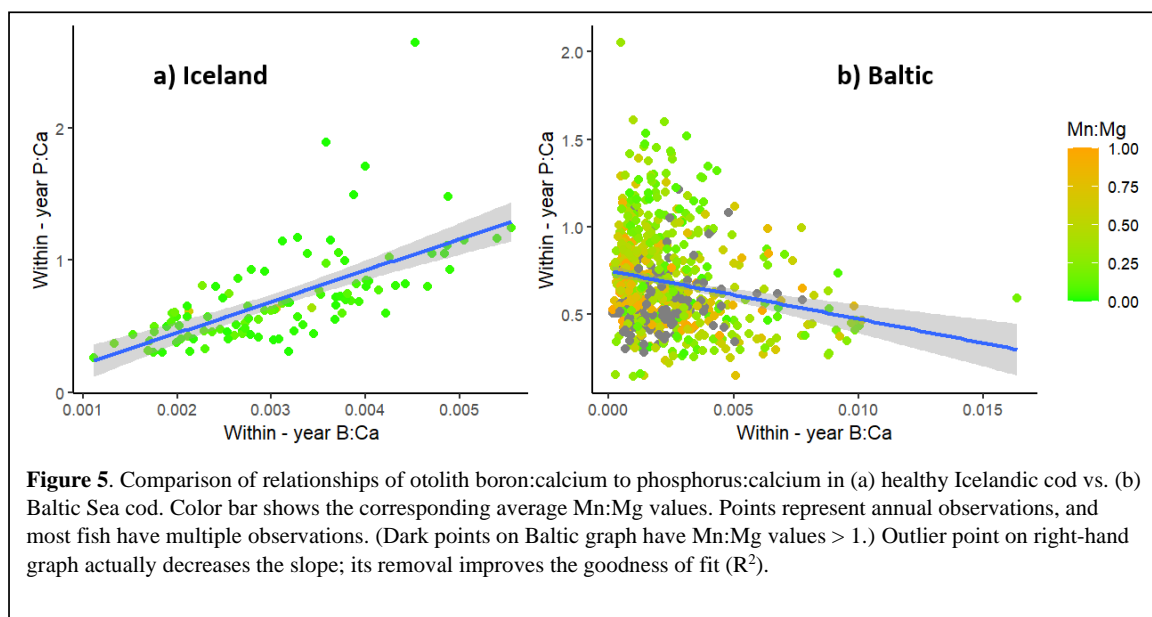
#### 209 3.4.1. P:Ca as a proxy for physiological status

210 For this analysis, the Icelandic “out-group” of cod living in normoxic water were  
 211 included. Annual average P:Ca was highly correlated with B:Ca in Icelandic cod otoliths, but  
 212 inversely correlated in Baltic cod with a great deal of scatter (Figure 5). Although the overall  
 213 trend for Baltic cod is negative, this hides trends that become more visible at a finer temporal  
 214 scale. Thus, we examined the slopes and  $R^2$  values of individual fish regressions of P:Ca on B:Ca



215 (Figure S4). The 2010s showed the strongest trends when slopes were plotted on  $R^2$  values, with  
216 Icelandic cod and the 2000s also trending positively, and weaker trends in the 1980s and 1990s  
217 (Figure S4(a)). Box plots of  $R^2$  values by decade and out-group showed wide variation (Figure  
218 S4(b)), and an analysis of variance found no statistical difference among Baltic fish by decade,  
219 but Icelandic otoliths had significantly higher  $R^2$  values ( $p < 0.001$ ).

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### 223 3.4.2. Mn:Mg as a proxy for hypoxia exposure

224 We tested the hypothesis that hypoxia (indexed by increased otolith Mn:Mg) was coupled  
225 to acidification (indexed by declining B:Ca). We used a similar analysis of individual fish  
226 regressions of Mn:Mg, our hypoxia exposure proxy, on B:Ca, as was performed with P:Ca. This  
227 analysis led to somewhat inconclusive results (Figure S5). The decades of the 1980s, 1990s, and  
228 2000s all had mostly positive slopes of Mn:Mg with B:Ca, and slopes increased with increasing  
229  $R^2$  values. Only the 2010s displayed a pattern consistent with the hypothesis. Iceland showed no  
230 relationship of slopes to goodness of fit. Further examination of the data included plotting trends



231 of these slopes against fish age (Figure S6). This revealed that younger EBC otoliths tended to  
232 have positive slopes, with slopes becoming more negative with age, for the decades of the 1990s-  
233 2010s; no trend was observed in the 1980s, and Icelandic cod had increasingly positive slopes as  
234 they aged (Figure S6). For the 2010s, it suggests that as Baltic cod age they encounter more  
235 hypoxia perhaps as they use deeper habitats, and with it, a decline in B:Ca that may indicate  
236 lower environmental pH. Overall, this analysis suggests that the decade of the 2010s may have  
237 had evidence of acidification driven by the same mechanisms that produced hypoxia.

238

#### 239 **4. Discussion**

240 This study was initiated because of finding a large discrepancy in otolith B:Ca ratios  
241 between fish captured in 2000 and others captured in 2019 – 2021. Archived otoliths from  
242 previous analyses were used to create a time series of data from 1985-2021. Otolith B:Ca ratios  
243 were analyzed in relation to long-term records of water chemistry in a central location in the  
244 Baltic Sea to represent trends in environmental conditions. Ratios were also examined against an  
245 elemental ratio (P:Ca) determined in other studies to be controlled physiologically, as well as a  
246 ratio (Mn:Mg) that serves as a proxy for hypoxia exposure (Limburg and Casini, 2018).

247 Boron is an essential micronutrient for plants, and is found in cyanobacteria and  
248 macroalgae (Howe, 1998). Its toxicity to fishes appears to be less than for invertebrates (Taylor  
249 et al., 1985; Howe, 1998) and is considered to be non-toxic in most aquatic systems.

250 Boron has been included in a number of otolith chemistry studies, mostly as part of a  
251 multivariate suite of tracers to determine provenance. Martino et al. (2017) studied boron as part  
252 of a suite of trace elements that might be affected by experimentally elevated pCO<sub>2</sub>, but found no  
253 effect. Cavole (2021) performed the first studies of otolith B:Ca as a potential proxy for ocean  
254 acidification, using wild fishes across a natural pH gradient (7.4 to 8.1). Cavole et al. (2023),  
255 examining potential tracers of warming, hypoxia, and acidification in deep-sea fishes, found  
256 wide variation in B:Ca but no consistent trends.

257 In this study, univariate regressions of otolith B:Ca on five water variables all were  
258 highly significant (Table 1). These variables are interrelated through ecological processes:  
259 warming temperatures and salinity in combination with nutrients produces massive algal blooms



260 in the Baltic Sea, which as they die off and decay, consume oxygen and change the pH and  
261 alkalinity (Kuliński et al., 2022). In that analysis, total alkalinity ( $A_T$ ) showed the strongest  
262 effect, followed by dissolved oxygen and salinity.

263 Further pairwise analysis of  $A_T$ , pH, and salinity by decade showed varying patterns of  
264 effect sizes and significance. Importantly, when testing pH and salinity together, pH had a  
265 significant, positive effect on B:Ca and salinity a significant, negative effect. Kuliński et al.  
266 (2018), building on earlier studies, showed that dissolved boron is tightly coupled to salinity in  
267 the Baltic Sea, with an intercept  $> 0$  indicating watershed sources of B input. The decoupling of  
268 otolith B:Ca from salinity and strong relationship to pH and alkalinity in the 2010s suggests that  
269 environmental conditions moved into a different regime from the previous three decades.

270 Among the analyzed environmental variables (Fig. S2), temperature shows a clear,  
271 positive trend starting from the 1980s. This is consistent with analyses performed by Lehmann et  
272 al. (2022) who found warming of 0.3-0.5 °C per decade in the whole water column in the central  
273 Baltic (station BY15) between 1979 and 2018. For other variables, trends are not that  
274 straightforward. For salinity, the variability reveals a stratification pattern typical for the central  
275 Baltic, with significantly higher S in the deep waters and a brackish surface water layer.  
276 However, in the 1980s and early 1990s there was a clear S drop in the deep waters being the  
277 result of a lower frequency of so-called Major Baltic Inflows that bring dense, saline water from  
278 the North Sea to the Baltic (Meier et al. 2022). This relaxed haline stratification in the central  
279 Baltic, which in turn led to better ventilation of deep waters and significantly raised  $O_2$   
280 availability (Fig. S2 (b) and (d); Lehmann et al., 2022) increased water volume for cod spawning  
281 and feeding. However, the 1980s decade was also the time when the maximum nutrient loads  
282 from land occurred, propelling eutrophication in the central Baltic. High organic matter  
283 production (and its export to bottom waters) co-occurred with the re-enhancement of water  
284 column stratification at the turn of the 20<sup>th</sup> and 21<sup>st</sup> centuries (increase in bottom S, Fig. S2 (d)).  
285 Both these factors led to a significant drop in  $O_2$  in deeper waters (Fig. S2 (b); Kuliński et al.  
286 (2022)), which limited the optimal water volume for cod expansion – a negative factor, that also  
287 coincided with high fishing exploitation.

288 Furthermore, all these changes together with a rising pressure from atmospheric  $CO_2$   
289 increase made fundamental changes in the marine  $CO_2$  system and seawater pH. First, the



290 relatively low pH in the 1960s was increasing by the 1990s as an initial effect of eutrophication  
291 (Fig. S2 (c)). Surprisingly, this trend was paralleled with a drop in  $A_T$  (Fig. S2 (e)), which is a  
292 counterintuitive phenomenon in marine  $CO_2$  system studies. However, the relatively constant  
293  $A_T/S$  ratio (Fig. S2 (f)) clearly suggests that this initial  $A_T$  decrease was an effect of salinity  
294 decrease. Already in the 1990s we can observe a regime shift and a clear decoupling of  $A_T$  from  
295  $S$  as revealed by the marked  $A_T/S$  increase, suggesting  $A_T$  increase being higher than observed  $S$   
296 changes at that time (Fig. S2 (d,e,f)). While in surface waters this  $A_T$  increase was likely due to  
297 climate-related (or  $CO_2$ -induced) enhancement of weathering on land and higher riverine  $A_T$   
298 loads, the  $A_T$  increase observed in deep waters was probably an effect of hypoxic/anoxic  $A_T$   
299 generation (Müller et al., 2016, Neumann et al., 2022). The latter is consistent with a large and  
300 sudden expansion of the hypoxic/anoxic areas in the central Baltic that occurred in the late 1990s  
301 (Meier et al., 2022) – a phenomenon that is also believed to be a root cause for a biogeochemical  
302 regime shift to the so-called vicious cycle – a self-supporting eutrophication mechanism  
303 (Kuliński et al., 2022). Due to these effects, pH after reaching its maximum in the late 1980s  
304 (effect of eutrophication), dropped somewhat and remained relatively constant in the 21<sup>st</sup> century  
305 expressing in that way the mutually cancelling effects of ocean acidification (effect of rising  
306 atmospheric  $CO_2$ ) and hypoxia/anoxia/weathering-driven alkalinization.

307 As is the case for some other trace elements (Hüssy et al., 2021), in this study boron  
308 uptake into otoliths appears to have some unexplained physiological regulation. This conclusion  
309 is based on close correspondence between B:Ca and P:Ca in otoliths from Icelandic cod that are  
310 not stressed by hypoxia (Figure 3(a), Figure 5(a)). In contrast to the Icelandic cod, the  
311 relationship of P:Ca to B:Ca in Baltic cod otoliths was highly variable, with a negative trend in  
312 the aggregate (Figure 5(b)). Regressions of P:Ca on B:Ca conducted on individual fish revealed  
313 mostly positive slopes, with highest slopes observed in the 2010s (Figure S4(a)), but lower, more  
314 variable  $R^2$  values (Figure S4(b)). Although not statistically different,  $R^2$  values were lowest in  
315 the 1980s and next-lowest in the 2010s. Compared to the Icelandic cod, all Baltic otoliths  
316 showed less coupling between P:Ca and B:Ca.

317 The  $CO_2^-$  and  $A_T$ -induced pH changes have definitely modified borates' availability in  
318 water and likely the B:Ca ratio in the Baltic cod's otoliths (Fig. 4). However, as shown by the  
319 strong correlation of B:Ca and the physiologically regulated P:Ca in the healthy, normoxic



320 Icelandic cod's otoliths, the B:Ca ratio may also contain built-in information about the overall  
321 physiological condition of fish (and not only the pH-dependency). This suggests that a complex  
322 picture of B:Ca development in the EBC otoliths over the recent decades may be affected by  
323 multiple factors, not only by pH changes but also physiological stress caused by overall changes  
324 that have recently occurred in the Baltic (Eero et al., 2020). These include expansion of  
325 hypoxia/anoxia, eutrophication, and biogeochemical regime shift. Together with overfishing,  
326 these factors led to an overall collapse of the EBC population in the Baltic.

327 Examination of B:Ca in relation to an otolith index of hypoxia exposure, Mn:Mg,  
328 revealed mostly positive slopes of individual fish regressions, with the notable exceptions of  
329 Baltic cod from the 2010s and Icelandic cod (Figure S5). Limburg and Casini (2018) adopted  
330 the ratio Mn:Mg as a hypoxia exposure proxy rather than Mn:Ca in part to correct for growth  
331 influences on Mn (see Limburg et al., 2015) as otolith Mg is related to metabolism and growth  
332 (Limburg et al., 2018). The positive relationships of Mn:Mg to B:Ca cannot be fully explained,  
333 but the inverse relationships observed in the 2010s is consistent with the hypothesis that cod  
334 were exposed to both hypoxia and lower pH (and also higher alkalinity). Some of the positive  
335 slopes were due to young ages of fish; we observed more coupling of Mn:Mg and B:Ca in young  
336 fish, and we hypothesize that at younger ages the growth effect on otolith Mn is not fully  
337 removed by dividing by Mg. Baltic cod had increasingly negative slopes of the Mn:Mg vs. B:Ca  
338 regressions as they aged (Figure S6), with the 2010s displaying the most negative slopes.

339

## 340 **5. Conclusions**

341 Baltic cod showed variable patterns of otolith B:Ca over a 35-year period, with maximum  
342 values in the late 1990s, and a gradual decline in B:Ca thereafter, to an all-time low in 2019-20.  
343 These trends are imperfectly related to water chemistry data, but showed strong correlations that  
344 varied by decade. Otolith B:Ca ratios and relationships in the 2010s were suggestive of  
345 environmental changes in pH and total alkalinity that could be coupled to deoxygenation.  
346 Physiological influences observed in a healthy Icelandic population were diminished in Baltic  
347 cod.



348           Undoubtedly, some of the relationships found could be due to the choice of samples, and  
349 a more focused study might select otoliths from areas where trends in acidification are clear.  
350 Future work could also examine otoliths from other species, as well as determine, if possible, the  
351 form of boron (borate or boric acid) taken up by otoliths. Complex interactions notwithstanding,  
352 we suggest that B:Ca in cod's otoliths can be a prospective variable in the palette of important  
353 and available tools to look into the environmental changes through the lens of otolith chemistry.

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