

1 Marked recent declines in boron in Baltic Sea cod otoliths – a bellwether  
2 of incipient acidification in a vast hypoxic system?

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11 **Abstract:** Ocean acidification is spreading globally as a result of anthropogenic CO<sub>2</sub> emissions,  
12 but the Baltic Sea has until recently been thought to be relatively well-buffered by terrigenous  
13 inputs of alkalinity from its watershed. We discovered a 3- to 5-fold decline in boron (as B:Ca)  
14 in otoliths of eastern Baltic Sea cod (EBC) between the late 1990s and 2021. ~~B:Ca is positively~~  
15 ~~proportional to pH in carbonates, as B in the form of borate is taken up in the CaCO<sub>3</sub> matrix.~~  
16 Examining a time series of EBC otoliths, we found varying levels of B:Ca since the 1980s, with  
17 the most recent years at an all-time low during this period. This trend correlates with declines in  
18 pH and dissolved oxygen, but not with changes in salinity. We examined possible physiological  
19 influences on B:Ca by including a collection of ~~healthy~~ Icelandic cod as an out-group. Icelandic  
20 cod otoliths showed strongly positive correlations of B:Ca with physiologically regulated P:Ca;  
21 this was not the case for EBC. Finally, B:Ca in EBC otoliths is ~~anti-negatively~~ correlated to some  
22 extent with Mn:Mg, a proposed proxy for hypoxia exposure. This negative relationship is  
23 hypothesized to reflect the dual phenomena of hypoxia and acidification as a result of  
24 decomposition of large algal blooms. Taken together, the otolith biomarkers Mn:Mg and B:Ca  
25 suggest a general increase in both hypoxia and acidification within the Baltic intermediate and  
26 deep waters in the last decade reflected in cod otoliths.

27 **Keywords:** Baltic Sea cod, otolith chemistry, boron, pH, hypoxia

## 30 1 Introduction

31 Ocean acidification, a known result of greenhouse gas emissions of CO<sub>2</sub> loading into the  
32 oceans, is highly variable in its distribution in the world's oceans (Doney et al., 2009). This is  
33 true in shallow coastal seas such as the brackish Baltic Sea, where in some locations surface  
34 pCO<sub>2</sub> values in winter greatly exceed other coastal regions (Reusch et al., 2018), but buffering  
35 can occur from river inputs bearing proton acceptors weathered from rock and soil (Müller et al.,  
36 2016; Kuliński et al., 2017). Alkalinity, defined as the sum of acid neutralizing compounds, is a  
37 complicated mix in the Baltic Sea (Kuliński et al., 2022). In the northern Baltic, i.e., the  
38 Bothnian Sea, total alkalinity (A<sub>T</sub>) is low as it is largely driven by A<sub>T</sub>-poor riverine inputs from  
39 the Scandinavian Peninsula that is built mostly of granite. In the southern and central Baltic, A<sub>T</sub>  
40 is higher due to salt inputs from the North Sea, influence of A<sub>T</sub>-rich continental rivers, and an  
41 additional contribution from internal cycling of carbon as algal blooms decompose at anoxic  
42 conditions (Kuliński et al., 2022). Despite buffering from alkalinity sources, pH is highly  
43 variable in space and time with its recent tendency to decline at greater depths in many parts of  
44 the Baltic Sea—(Kulinski et al., 2022). ~~which is~~ This decline is associated with the ongoing  
45 eutrophication and higher vertical export of organic matter that leads also to worsening  
46 deoxygenation (Kuliński et al., 2022).

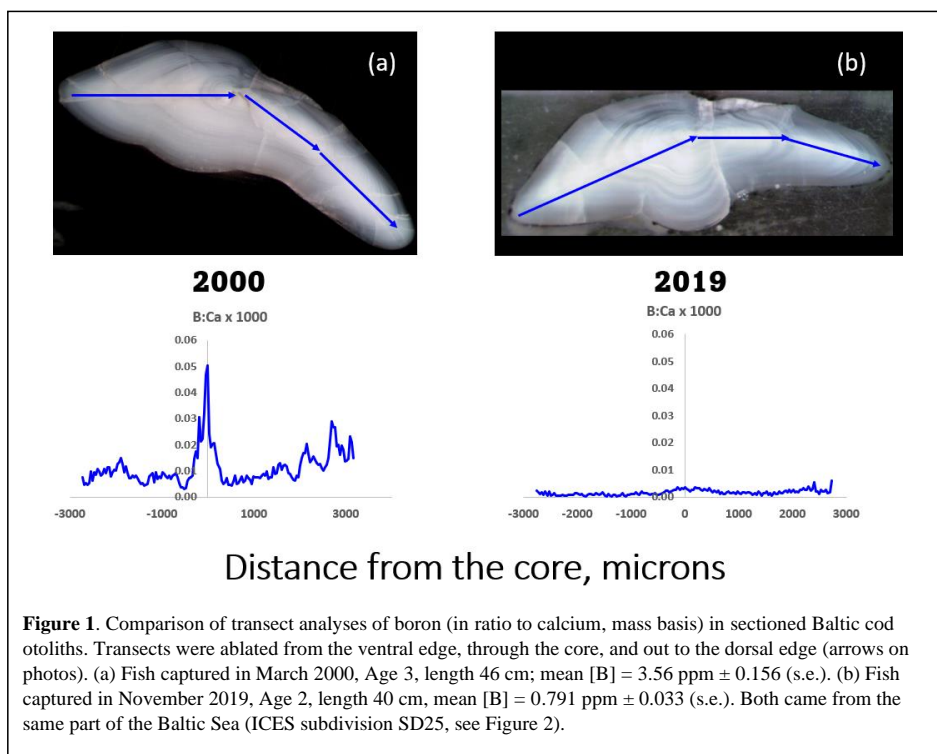
47 The population of the Eastern Baltic cod (*Gadus morhua*, hereafter referred to as EBC)  
48 has been severely impacted by a number of factors that include past overfishing, hypoxia,  
49 parasite infections, and seal predation (Eero et al., 2020), leading to reduced growth and age at  
50 maturity (Eero et al., 2016; ICES, 2019). Cod exposed to hypoxia were shown to have lower  
51 growth rates and worsened body condition—~~quantified by the Fulton K index of weight:length~~  
52 ratio (Limburg & Casini, 2018, 2019; Casini et al., 2016; Limburg & Casini, 2018).

53 Hypoxia ~~Recent studies that reported links between hypoxia exposure and body condition~~  
54 have relied on the measurement of trace elements in fish otoliths (ear-stones). ~~was quantified by~~  
55 ~~means of a proxy developed from analyzing manganese concentrations in otoliths (ear stones) of~~  
56 ~~fish (Limburg et al., 2015), in ratio to magnesium which corrects to some extent for growth~~  
57 ~~influences on Mn uptake (Limburg et al., 2018; Limburg & Casini, 2018).~~ Briefly, otoliths, the  
58 calcified structures that form part of the hearing/balance system in teleost fishes, are sectioned  
59 and then analyzed by ablating micro-transects along the major growth axis and analyzing by

60 mass spectrometry (see Methods). [Hypoxia exposure was quantified by means of a proxy](#)  
61 [developed from analyzing manganese concentrations in otoliths \(ear stones\) of fish \(Limburg et](#)  
62 [al., 2015\), in ratio to magnesium which corrects to some extent for growth influences on Mn](#)  
63 [uptake, since both are affected by growth but only manganese is affected directly by hypoxic](#)  
64 [conditions \(Limburg et al., 2018; Limburg & Casini, 2018\).](#) The resulting elemental data are  
65 lifetime concentration histories, being incorporated at the time of exposure. The use of otolith  
66 manganese to track hypoxia exposure is one of the emerging biomarkers in fisheries ecology  
67 (Reis-Santos et al., 2022).

68 Recently, with acquisition of more sensitive instrumentation, we began to experiment  
69 with quantifying elements having sub-ppm concentrations in cod otoliths, including the trace  
70 element boron. [Interestingly, preliminary results from this work showed elevated levels of boron](#)  
71 [in 2000 but not in 2019. Boron is noteworthy because it is an indicator of salinity. Specifically,](#)  
72 Boron in seawater generally correlates with salinity (Kuliński et al., 2017) predominantly in the  
73 form of weak boric acid ( $H_3BO_3$ ) at standard seawater salinity (35 PSU) and pH of 8. However,  
74 the dissociated anion borate  $B(OH)_4$  is positively, nonlinearly correlated with pH (Yu et al.,  
75 2007). Borate is the dominant form incorporated into carbonates such as limestones,  
76 foraminifera, and corals in proportion to pH (Yu et al., 2007; Hönisch et al., 2012; Levin et al.,  
77 2015) and is often confirmed by  $\delta^{11}B$  analysis, since it has a large, negative fractionation  
78 compared to boric acid. However, even bulk boron, as B:Ca ratios, has been found to correlate  
79 positively (and linearly) with pH in foraminifera (Yu et al., 2007), albeit in a species-specific  
80 manner.

81 We measured sub-ppm to low-ppm concentrations of B (ca. 0.5 – 5 ppm, see Methods  
82 section) in otoliths of EBC captured in 2000 that showed what appeared to be seasonal variations  
83 (example shown in Figure 1(a)), albeit not in synchrony with other elements that have proven  
84 useful to age cod (Heimbrand et al., 2020). It was therefore surprising when we began to analyze  
85 otoliths of EBC captured in 2019 and discovered greatly reduced concentrations (Figure 1(b)).  
86 The measurements were repeated for verification.



87 With such differences suggestive of change, we returned to an archive of otoliths used for  
88 another project (Limburg and Casini, 2018), as well as adding previously un-analyzed otoliths, to  
89 develop a time series of otolith B:Ca.

90 In this study, we explored the extent to which otolith B:Ca varied through time and  
91 whether its values were correlated with pH, salinity, or other otolith derived values that are  
92 proxies for environmental or physiological factors. In particular, otolith P:Ca is known to be

93 under physiological influence (Thomas et al., 2017; Heimbrand et al., 2020). In cod, higher  
94 values of P:Ca occur during the growing season and are thus hypothesized to be associated with  
95 growth and other activities. We selected this as well as our proxy for hypoxia exposure, otolith  
96 Mn:Mg, and to document any temporal trends that could be correlated with pH, salinity, or other  
97 environmental factors. We posed the following questions:

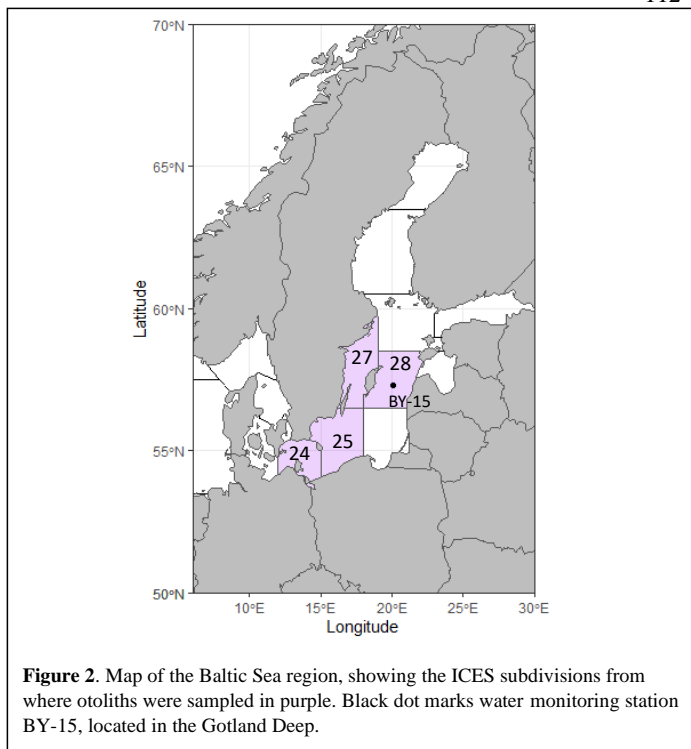
- 98
- 99 1. What are the temporal trends in otolith boron (as B:Ca ratios)?
  - 100 2. How do B:Ca ratios correlate to trends in water chemistry, particularly salinity and pH?
  - 101 3. If pH declines are associated with hypoxia, do we see a correlation of B:Ca with our hypoxia  
102 proxy, manganese:magnesium (Mn:Mg)?
  - 103 4. If there are any physiological influences, do we see a correlation of B:Ca with elemental ratios  
104 known to be under physiological control, e.g., phosphorus:calcium (P:Ca)?

105

## 106 **2 Materials and Methods**

### 107 **2.1 Otolith collections and microchemical analysis**

108 Otoliths of Baltic cod (N = 156) were obtained from both fishery-independent and  
109 fishery-dependent surveys conducted by the Swedish Fisheries Board and its successor, the  
110 Department of Aquatic Resources, Swedish University of Agricultural Sciences. Fish were  
111 collected from ICES sub-divisions (SD) 24, 25, 27, and 28 (Figure 2), spanning the period 1988-



2021. To construct the time series, we selected previously analyzed (but not for B:Ca) otoliths that were comparable in age range to those collected in 2019-2021. These All otoliths were measured, weighed, and subsequently embedded in epoxy (Struers EpoFix) and sectioned transversely to expose the core. They were smoothed with a succession of lapping films to 3 microns.

In addition, cod otoliths (N = 14) were obtained from fisheries-

131 independent surveys from Iceland (ICES SD 5a), as an out-group for studies of possible  
132 physiological correlations. This part of the North Atlantic does not have hypoxia or acidification  
133 issues in the deep waters. The Icelandic otoliths were prepared and analyzed in similar fashion.

134 Baltic cod were aged by examining seasonal patterns in otolith Mg:Ca and P:Ca,  
135 following Heimbrand et al. (2020). Icelandic cod were also aged by otolith chemistry, although  
136 several of them had been previously aged visually by age readers in Iceland. For the latter  
137 otoliths, the seasonal chemical patterns produced age estimations that matched exactly with the  
138 visual estimates.

139 Chemical analyses were performed by laser ablation inductively coupled plasma mass  
140 spectrometry (LA-ICP-MS) at the Analytical and Technical Services group at SUNY ESF.  
141 Otoliths were ablated with a 192-nm laser ablation unit (Teledyne Cetac Excite 2) along transects  
142 as illustrated in Figure 1; the ablated material was transported via an Ar-He carrier gas mixture  
143 into a Thermo iCAP TQ plasma mass spectrometer where isotopes were quantified ( $^{11}\text{B}$ ,  $^{25}\text{Mg}$ ,  
144  $^{31}\text{P}$ ,  $^{43}\text{Ca}$ ,  $^{55}\text{Mn}$ ,  $^{63}\text{Cu}$ ,  $^{66}\text{Zn}$ ,  $^{88}\text{Sr}$ ,  $^{127}\text{I}$ ,  $^{138}\text{Ba}$ , and  $^{208}\text{Pb}$ ). Daily performance of the system was  
145 optimized with NIST 612 glass standard and U.S. Geological Survey standards MACS-3,  
146 MAPS-4, and MAPS-5 were used as standards for calibration and drift correction. The MAPS  
147 (abbreviation for Micro-Analytical Phosphate Standard) standards, while having good matching  
148 to the otolith matrix, were slightly softer and were therefore ablated at a higher energy (25%  
149 power, 2.03 J/cm<sup>2</sup> fluence) than the otoliths (20% power, 1.62 J/cm<sup>2</sup>) to achieve an otolith  
150 calcium value around 38 weight percent. Transects were made with 110  $\mu\text{m}$ -diameter ablation  
151 spots at a speed of 7  $\mu\text{m}/\text{sec}$ . Although often edge-to-edge laser transects were made (to check  
152 for symmetry), only data collected on the dorsal (longer) axis were used in our analysis due to  
153 greater spatial resolution. Once data were worked up, annual mean values were parsed by the age  
154 determination technique in Heimbrand et al. (2020), i.e., averages within annual growth rings  
155 were calculated. These were matched to corresponding calendar years.

## 156 2.2 Water data

157 Water properties data (temperature (T, °C), salinity (S, psu), dissolved oxygen (DO,  
158 mL/L), pH, and total alkalinity ( $A_T$ , mmol/L)) were downloaded from the Swedish  
159 Meteorological and Hydrological Institute's database, SHARKWeb (<https://sharkweb.smhi.se/>).  
160 Station BY-15, Gotland Deep, was selected as being representative of the central Baltic (Figure  
161 2) and having one of the longest and most detailed timeseries. We selected depths that  
162 corresponded to where cod had been found in fishery-independent surveys (Casini et al., 2021).  
163 Between 1985-1995, we averaged water data from 30-60 meters, and from 1996 onward,  
164 between 40-75 meters. Annual mean water values were computed and matched to corresponding  
165 otolith chemistry data (annual means of Element:Ca data parsed to calendar years as described  
166 above).

## 167 168 2.3 Analysis

169 Visualization ([graphs](#)) and statistical analyses were performed in Excel, R ([base R](#) (R  
170 Core Team 2022) and [ggplot2](#) (Wickham, 2016)), and JMP version 17.0 (JMP 2022). Statistical  
171 analyses were separated into examination of relationships of B:Ca values to potential  
172 environmental drivers (primarily pH and S, but also  $A_T$ , DO, and T) and to the internal variables  
173 P:Ca and Mn:Mg. ~~P:Ca was found to correlate highly with B:Ca in Icelandic cod otoliths~~  
174 ~~(Figure 3(a)). Otolith P:Ca is known to be under physiological influence (Thomas et al., 2017;~~  
175 ~~Heimbrand et al., 2020). In cod, higher values of P:Ca occur during the growing season and are~~  
176 ~~thus hypothesized to be associated with growth and other activities. Otolith Mn:Mg, on the other~~  
177 ~~hand, is our proxy for hypoxia exposure (Limburg and Casini, 2018), and often correlated~~  
178 ~~negatively with B:Ca (Figure 3(b))—consistent with co-occurrence of reduced conditions in~~  
179 ~~water (and presence of bioavailable ionic  $Mn^{2+}$ ) and high  $pCO_2$  and thus low pH, both being the~~  
180 ~~effect of enhanced organic matter remineralization and/or poor ventilation of the water layers~~  
181 ~~occupied by cod. For the environmental drivers, graphical a PCA analysis was first performed to~~  
182 ~~examine relationships among them. Thereafter, B:Ca was regressed on the first principal~~  
183 ~~component eigenvector to test for its relationship to the main drivers. performed This was~~  
184 followed by mixed linear models ~~of environmental drivers on B:Ca and visualizations~~, with each  
185 fish's age (nested within fish ID) being treated as a random effect. ~~This analysis took account of~~  
186 ~~repeated measures within the same individual, and was intended to examine more closely the~~  
187 ~~effect of individual drivers.~~

188 For the internal variables, the P:Ca and Mn:Mg transect data were regressed on  
189 corresponding B:Ca for each fish, and the slopes, intercepts and  $R^2$  values were calculated. These  
190 were examined for trends by decade (1980s, 1990s, 2000s, and 2010s) as well as compared with  
191 the Icelandic out-group. Positive correlations with P:Ca would suggest a physiological influence  
192 on B:Ca incorporation, whereas negative correlations with Mn:Mg would suggest a linkage  
193 between B:Ca as an indicator of pH and our indicator of hypoxia exposure (cf. Cavole, 2021).  
194 ~~These were tested using ANOVA with slopes as the dependent variable and decades (for Baltic~~  
195 ~~Sea) and Iceland as grouping variables.~~

196

## 197 **3 Results**

### 198 **3.1. Trends in water variables**



199 A principal components analysis of the water variables [found that the first two](#)  
200 [components explained 92.9% of the variance \(72.4% by PC1 and 20.5% by PC2\). The analysis](#)  
201 [showed three groupings: \(1\) DO and pH that were opposite to \(2\) salinity and A<sub>T</sub>, respectively;](#)  
202 [and \(3\) temperature was on a separate axis \(Figure S1\). Linear regression of B:Ca on the first](#)  
203 [principal component yielded a negative relation \(B:Ca = 0.000485\(PC1\) + 0.00246, R<sup>2</sup> = 0.21, p](#)  
204 [< 0.0001\).](#)

205 Examination of the long-term time series of data from Station BY-15 shows that while  
206 temperature generally increased (Figure [S23\(a\)](#)), the other water variables showed very nonlinear  
207 patterns (Figures [S23\(b-e\)](#)). Alkalinity is in part a function of salinity; dividing A<sub>T</sub> by S produces  
208 a time series showing a dramatic state change around 1990 (Figure [S23\(f\)](#)), particularly in the  
209 water layers occupied by cod. Despite the increasing A<sub>T</sub>, pH has declined more or less  
210 monotonically at midwater depths (40-75 m, Figure [S23\(c\)](#)).

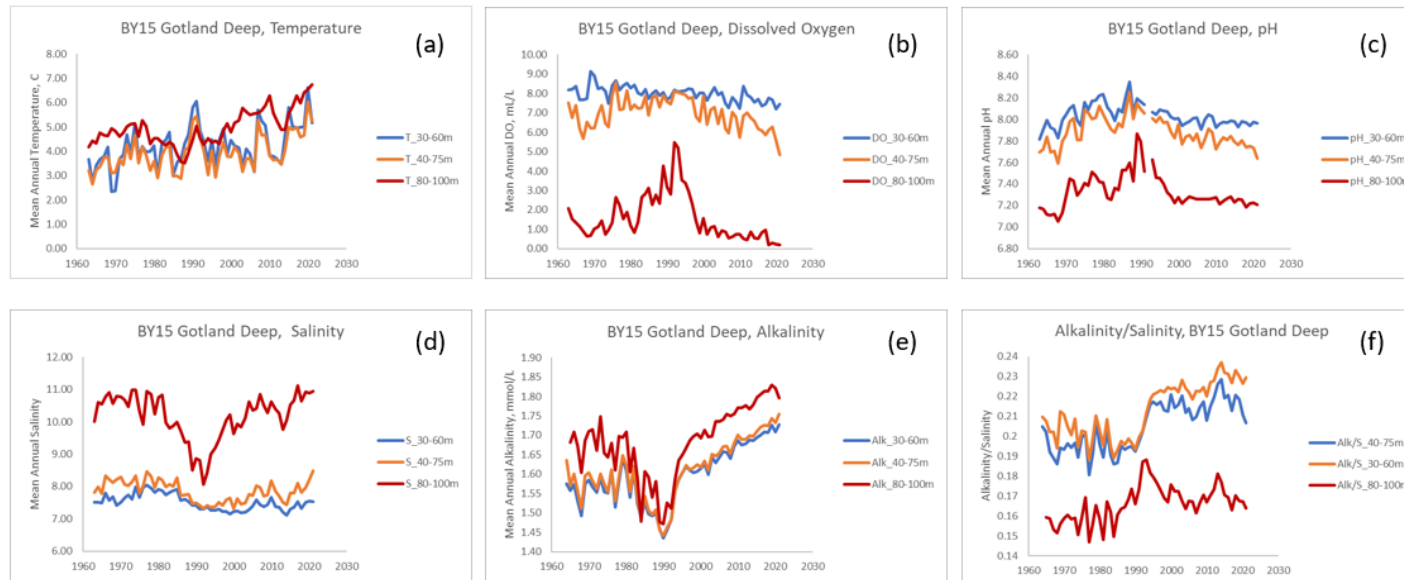
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### Station BY-15, Gotlandsdjupet 1963-2021



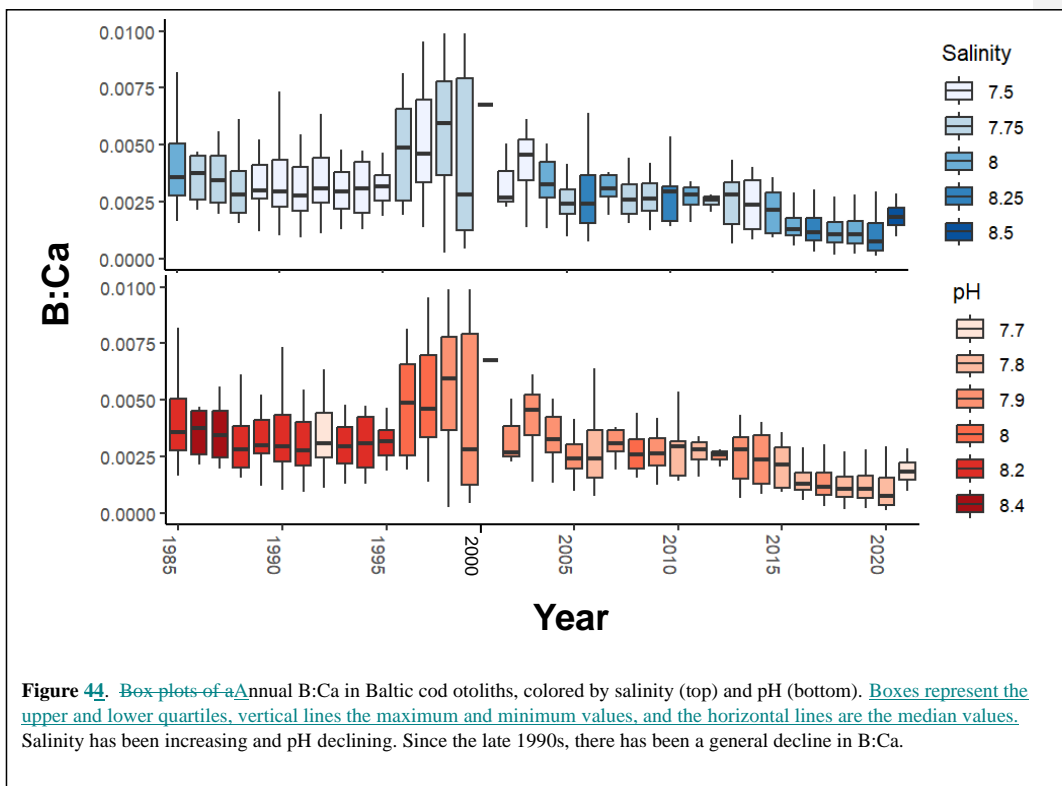
**Figure 3.** Time series of water data for three depth zones, Gotland Deep Station BY-15. (a) Temperature. (b) Dissolved oxygen. (c) pH. (d) Salinity. (e) Total alkalinity. (f) Alkalinity normalized to salinity.

214

215 **3.2 Trends in otolith B:Ca with salinity and pH**

216 Over the period 1985-2019, box plots of B:Ca by year show a pattern of increase toward  
 217 a maximum, albeit with great variation, in the late 1990s, followed by gradual decline (Figure  
 218 44). Visual comparisons with salinity and pH show that neither explain all the variation in B:Ca,  
 219 but since 2000, pH and B:Ca both trend downward whereas salinity has increased. The increase  
 220 in B:Ca in 2021 is from age-0, which is elevated relative to older ages (cf. Figure 1(a)).

221



222

223

224 **3.3. Effects of water parameters on otolith B:Ca**

225 **3.3.1. Univariate effects of water parameters on otolith B:Ca**

226 ~~Repeated~~ Following the PCA, ~~repeated~~ measures regressions of otolith B:Ca on each of  
227 the five water variables was conducted to examine effect sizes and significance each variable  
228 alone. ~~Repeated-measures~~ These regressions showed significant effects of each of the five water  
229 variables on otolith B:Ca (Table 1). A<sub>T</sub> had the largest effect and was negative (Table 1).  
230 Dissolved oxygen and pH both had positive slopes, and temperature and salinity both were  
231 negative. All univariate analyses were highly significant. Ages of individual fish were not  
232 significant as measured by Wald's p-value in restricted maximum likelihood (REML) analysis.

**Table 1.** Repeated measures univariate regressions of otolith B:Ca on water variables, ordered by goodness of fit (-2 log likelihood criterion; lowest is best).

Variable	Parameter slope	R <sup>2</sup>	df	F ratio	p	-2 Log Likelihood
Total Alkalinity	-9.42E-03	0.24	647	158.8	< 0.0001	-6358
Dissolved Oxygen	1.00E-03	0.21	615	142.2	< 0.0001	-6341
Salinity	-3.00E-03	0.22	565	133.5	< 0.0001	-6336
Temperature	-8.30E-04	0.13	639	58.7	< 0.0001	-6269
pH	5.10E-03	0.15	633	109.4	< 0.0001	-6189

233  
234  
235  
236 **3.3.2. Two-factor tests of water parameters on otolith B:Ca**

237 Because of ~~our~~ hypothesized positive relationships of salinity and pH on B:Ca, as well as  
238 the large effect size of A<sub>T</sub>, we examined pairwise (two-factor) models of these independent  
239 variables on otolith B:Ca (Table 2). This allowed us to examine the relative strength and  
240 direction of trends (slope parameters, Table 2) on B:Ca when testing two independent variables  
241 together. Aside from an overall model using all the data, these were filtered by decade to  
242 examine finer scale trends.

**Table 2.** Repeated measures 2-factor models of effects on otolith B:Ca, by decade. Numbers are estimated coefficients. Red = significant at  $p < 0.01$ . Blue = significant at  $p < 0.05$ . Negative 2 Log Likelihood is a metric of goodness-of-fit, with smaller numbers indicating better fit.

1. pH and Alkalinity:

Parameter	1980s	1990s	2000s	2010s	Overall
pH	0.00128	-0.00774	-0.00375	-0.00121	-0.00385
Alkalinity	0.00873	-0.00538	-0.01840	-0.02585	-0.01585
-2 Log Lik.	-399.2	-1231.8	-896	-3979.4	-6225.5

2. pH and Salinity:

Parameter	1980s	1990s	2000s	2010s	Overall
pH	0.00151	-0.00226	-0.00983	0.00378	0.00225
Salinity	0.00322	0.00365	-0.00347	-0.00049	-0.00208
-2 Log Lik.	397.9	-1231.6	-894.8	-3948.2	-6201.6

3. Alkalinity and Salinity:

Parameter	1980s	1990s	2000s	2010s	Overall
Alkalinity	-0.007269	0.002042	-0.01253	-0.027822	-0.006591
Salinity	0.003768	0.004802	-0.000171	0.000478	-0.001274
-2 Log Lik.	-401.4	-1357.9	-891.9	-3977.6	-6354.8

243  
 244 In both pairwise tests with pH, this parameter had positive slopes in the 1980s and 2010s  
 245 but negative slopes in the 1990s and 2000s. Alkalinity had mostly a negative relation to otolith  
 246 B:Ca. Salinity had a positive effect in the 1980s and 1990s but became negative in the 2000s and  
 247 2010s. Salinity overall had a negative effect on B:Ca, in contrast to expectations. We note the  
 248 strong, negative effect of pH in the 1990s when paired with  $A_T$ , and the strong, positive effect of  
 249 pH in the 2010s when paired with salinity. Age, when treated as a random variable nested within  
 250 individual fish (by fish ID), was again not significant. However, in a separate analysis (Figure  
 251 [S3S2](#)), age (not treated as a random variable) and otolith B:Ca were inversely related in the  
 252 1980s, 1990s, and 2000s, and had a weakly positive trend in the 2010s.

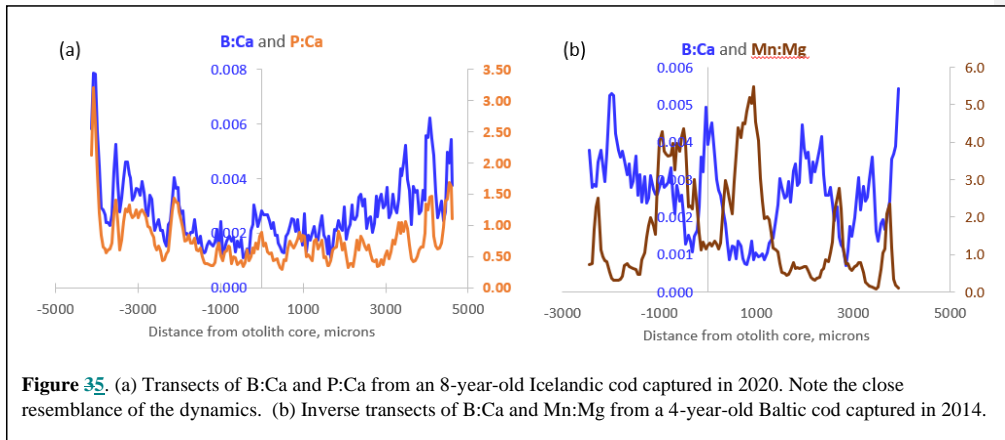
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256 **3.4. Internal parameters and otolith B:Ca**

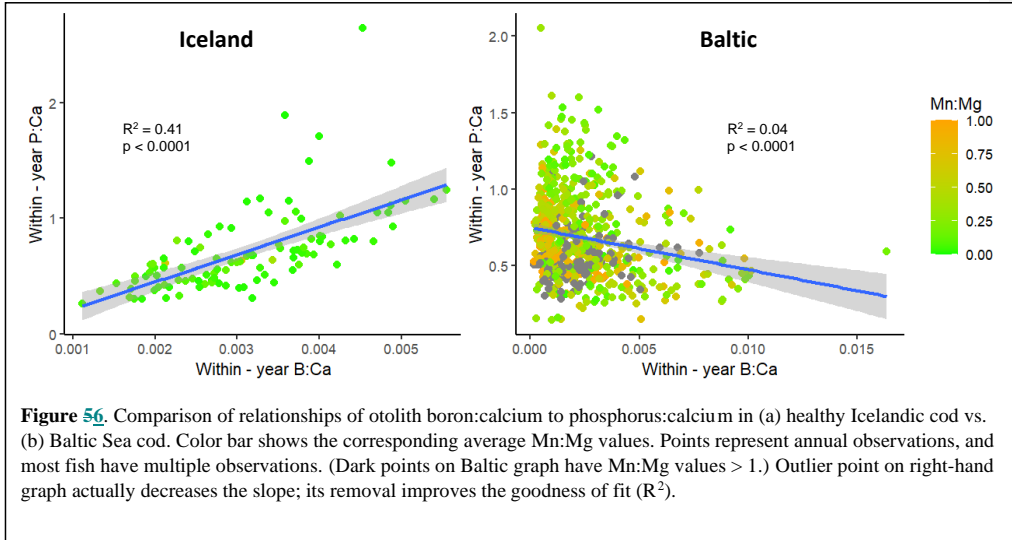
257 During exploratory examination of the data, P:Ca was found to correlate highly with  
258 B:Ca in Icelandic cod otoliths (Figure 5(a)). Otolith Mn:Mg, on the other hand, is our proxy for  
259 hypoxia exposure (Limburg and Casini, 2018), and often correlated negatively with B:Ca (Figure  
260 5(b)). This is consistent with co-occurrence of reduced conditions in water (and presence of  
261 bioavailable ionic Mn<sup>2+</sup>) and high pCO<sub>2</sub> and thus low pH, both being the effect of enhanced  
262 organic matter remineralization and/or poor ventilation of the water layers occupied by cod.



263 **Figure 35.** (a) Transects of B:Ca and P:Ca from an 8-year-old Icelandic cod captured in 2020. Note the close  
264 resemblance of the dynamics. (b) Inverse transects of B:Ca and Mn:Mg from a 4-year-old Baltic cod captured in 2014.

264 **3.4.1. P:Ca as a proxy for physiological status**

265 For this analysis, the Icelandic “out-group” of cod living in normoxic water were  
266 included. Annual average P:Ca was highly correlated with B:Ca in Icelandic cod otoliths, but  
267 inversely correlated in Baltic cod with a great deal of scatter (Figure 56). Although the overall  
268 trend for Baltic cod is negative, this hides trends that become more visible at a finer temporal  
269 scale. Thus, we examined the slopes and R<sup>2</sup> values of individual fish regressions of P:Ca on B:Ca  
270 (Figure S4S3). The 2010s showed the strongest trends when slopes were plotted on R<sup>2</sup> values,  
271 with Icelandic cod and the 2000s also trending positively, and weaker trends in the 1980s and  
272 1990s (Figure S4S3(a)). Box plots of R<sup>2</sup> values by decade and out-group showed wide variation  
273 (Figure S4S3(b)), and an analysis of variance found no statistical difference among Baltic fish by  
274 decade, but Icelandic otoliths had significantly higher R<sup>2</sup> values (p < 0.001).



276

277

### 278 3.4.2. Mn:Mg as a proxy for hypoxia exposure

279 We tested the hypothesis that hypoxia (indexed by increased otolith Mn:Mg) was coupled  
 280 to acidification (indexed by declining B:Ca). We used a similar analysis of individual fish  
 281 regressions of Mn:Mg, our hypoxia exposure proxy, on B:Ca, as was performed with P:Ca (see  
 282 [Methods](#)). This analysis led to somewhat inconclusive results (Figure [S5S4](#)). The decades of the  
 283 1980s, 1990s, and 2000s all had mostly positive slopes of Mn:Mg with B:Ca, and slopes  
 284 increased with increasing  $R^2$  values. Only the 2010s displayed a pattern consistent with the  
 285 hypothesis. Iceland showed no relationship of slopes to goodness of fit. Further examination of  
 286 the data included plotting trends of these slopes against fish age (Figure [S6S5](#)). This revealed  
 287 that younger EBC otoliths tended to have positive slopes, with slopes becoming more negative  
 288 with age, for the decades of the 1990s-2010s; no trend was observed in the 1980s, and Icelandic  
 289 cod had increasingly positive slopes as they aged (Figure S6). For the 2010s, it suggests that as  
 290 Baltic cod age they encounter more hypoxia perhaps as they use deeper habitats, and with it, a

291 decline in B:Ca that may indicate lower environmental pH. Overall, this analysis suggests that  
292 the decade of the 2010s may have had evidence of acidification driven by the same mechanisms  
293 that produced hypoxia.

294

#### 295 **4. Discussion**

296 This study was initiated because of finding a large discrepancy in otolith B:Ca ratios  
297 between fish captured in 2000 and others captured in 2019 – 2021. Archived otoliths from  
298 previous analyses were [then](#) used to create a time series of data from 1985-2021. Otolith B:Ca  
299 ratios were analyzed in relation to long-term records of water chemistry in a central location in  
300 the Baltic Sea to represent trends in environmental conditions. Ratios were also examined against  
301 an elemental ratio (P:Ca) determined in other studies to be controlled physiologically, as well as  
302 a ratio (Mn:Mg) that serves as a proxy for hypoxia exposure (Limburg and Casini, 2018).

303 ~~—— Boron is an essential micronutrient for plants, and is found in cyanobacteria and  
304 macroalgae (Howe, 1998). Its toxicity to fishes appears to be less than for invertebrates (Taylor  
305 et al., 1985; Howe, 1998) and is considered to be non-toxic in most aquatic systems.~~

306 Boron has been included in a number of otolith chemistry studies, mostly as part of a  
307 multivariate suite of tracers to determine provenance. Martino et al. (2017) studied boron as part  
308 of a suite of trace elements that might be affected by experimentally elevated pCO<sub>2</sub>, but found no  
309 effect. Cavole (2021) performed the first studies of otolith B:Ca as a potential proxy for ocean  
310 acidification, using wild fishes across a natural pH gradient (7.4 to 8.1). Cavole et al. (2023),  
311 examining potential tracers of warming, hypoxia, and acidification in deep-sea fishes, found  
312 wide variation in B:Ca but no consistent trends.

313 In this study, univariate regressions of otolith B:Ca on five water variables all were  
314 highly significant, [with dissolved oxygen and pH showing positive relationships while negative  
315 relationships were observed with alkalinity, temperature, and salinity](#) (Table 1). These variables  
316 are interrelated through ecological processes: warming temperatures and salinity in combination  
317 with nutrients produces massive algal blooms in the Baltic Sea, which as they die off and decay,  
318 consume oxygen and change the pH and alkalinity (Kuliński et al., 2022). In that analysis, total  
319 alkalinity (A<sub>T</sub>) showed the strongest effect, followed by dissolved oxygen and salinity.



320 Further pairwise analysis of  $A_T$ , pH, and salinity by decade showed varying patterns of  
321 effect sizes and significance. Importantly, when testing pH and salinity together, pH had a  
322 significant, positive effect on B:Ca and salinity a significant, negative effect. Kuliński et al.  
323 (2018), building on earlier studies, showed that dissolved boron is tightly coupled to salinity in  
324 the Baltic Sea, with an intercept  $> 0$  indicating watershed sources of B input. The decoupling of  
325 otolith B:Ca from salinity and strong relationship to pH and alkalinity in the 2010s suggests that  
326 environmental conditions moved into a different regime from the previous three decades.

327 Among the analyzed environmental variables ([Fig. S2Figure 3](#)), temperature shows a  
328 clear, positive trend starting from the 1980s. This is consistent with analyses performed by  
329 Lehmann et al. (2022) who found warming of 0.3-0.5 °C per decade in the whole water column  
330 in the central Baltic (station BY15) between 1979 and 2018. For other variables, trends are not  
331 that straightforward. For salinity, the variability reveals a stratification pattern typical for the  
332 central Baltic, with significantly higher S in the deep waters and a brackish surface water layer.  
333 However, in the 1980s and early 1990s there was a clear S drop in the deep waters being the  
334 result of a lower frequency of so-called Major Baltic Inflows that bring dense, saline water from  
335 the North Sea to the Baltic (Meier et al. 2022). This relaxed haline stratification in the central  
336 Baltic, which in turn led to better ventilation of deep waters and significantly raised O<sub>2</sub>  
337 availability ([Fig. S2Figure 3](#) (b) and (d); Lehmann et al., 2022) increased water volume for cod  
338 spawning and feeding. However, the 1980s decade was also the time when the maximum  
339 nutrient loads from land occurred, propelling eutrophication in the central Baltic. High organic  
340 matter production (and its export to bottom waters) co-occurred with the re-enhancement of  
341 water column stratification at the turn of the 20<sup>th</sup> and 21<sup>st</sup> centuries (increase in bottom S, [Fig.](#)  
342 [S2Figure 3](#) (d)). Both these factors led to a significant drop in O<sub>2</sub> in deeper waters ([Fig. S2Figure](#)  
343 [3](#) (b); Kuliński et al. (2022)), which limited the optimal water volume for cod expansion – a  
344 negative factor, that also coincided with high fishing exploitation.

345 Furthermore, all these changes together with a rising pressure from atmospheric CO<sub>2</sub>  
346 increase made fundamental changes in the marine CO<sub>2</sub> system and seawater pH. First, the  
347 relatively low pH in the 1960s was increasing by the 1990s as an initial effect of eutrophication  
348 ([Fig. S2Figure 3](#) (c)). Surprisingly, this trend was paralleled with a drop in  $A_T$  ([Fig. S2Figure 3](#)  
349 (e)), which is a counterintuitive phenomenon in marine CO<sub>2</sub> system studies. However, the

350 relatively constant  $A_T/S$  ratio (Fig. S2Figure 3 (f)) clearly suggests that this initial  $A_T$  decrease  
351 was an effect of salinity decrease. Already in the 1990s we can observe a [regime shiftchange in](#)  
352 [trends](#) and a clear decoupling of  $A_T$  from  $S$  as revealed by the marked  $A_T/S$  increase, suggesting  
353  $A_T$  increase being higher than observed  $S$  changes at that time (Fig. S2Figure 3 (d,e,f)). While in  
354 surface waters this  $A_T$  increase was likely due to climate-related (or  $CO_2$ -induced) enhancement  
355 of weathering on land and higher riverine  $A_T$  loads, the  $A_T$  increase observed in deep waters was  
356 probably an effect of hypoxic/anoxic  $A_T$  generation (Müller et al., 2016, Neumann et al., 2022).  
357 The latter is consistent with a large and sudden expansion of the hypoxic/anoxic areas in the  
358 central Baltic that occurred in the late 1990s (Meier et al., 2022) – a phenomenon that is also  
359 believed to be a root cause for a biogeochemical regime shift to the so-called vicious cycle – a  
360 self-supporting eutrophication mechanism (Kuliński et al., 2022). Due to these effects, pH after  
361 reaching its maximum in the late 1980s (effect of eutrophication), dropped somewhat and  
362 remained relatively constant in the 21<sup>st</sup> century expressing in that way the mutually cancelling  
363 effects of ocean acidification (effect of rising atmospheric  $CO_2$ ) and hypoxia/anoxia/weathering-  
364 driven alkalization.

365 As is the case for some other trace elements (Hüssy et al., 2021), in this study boron  
366 uptake into otoliths appears to have some unexplained physiological regulation. This conclusion  
367 is based on close correspondence between B:Ca and P:Ca in otoliths from Icelandic cod that are  
368 not stressed by hypoxia (Figure 34(a), Figure 56(a)). In contrast to the Icelandic cod, the  
369 relationship of P:Ca to B:Ca in Baltic cod otoliths was highly variable, with a negative trend in  
370 the aggregate (Figure 56(b)). Regressions of P:Ca on B:Ca conducted on individual fish revealed  
371 mostly positive slopes, with highest slopes observed in the 2010s (Figure S4S3(a)), but lower,  
372 more variable  $R^2$  values (Figure S4S3(b)). Although not statistically different,  $R^2$  values were  
373 lowest in the 1980s and next-lowest in the 2010s. Compared to the Icelandic cod, all Baltic  
374 otoliths showed less coupling between P:Ca and B:Ca.

375 The  $CO_2^-$  and  $A_T$ -induced pH changes have definitely modified borates' availability in  
376 water and likely the B:Ca ratio in the Baltic cod's otoliths (Fig. 45). However, as shown by the  
377 strong correlation of B:Ca and the physiologically regulated P:Ca in the healthy, normoxic  
378 Icelandic cod's otoliths, the B:Ca ratio may also contain built-in information about the overall  
379 physiological condition of fish (and not only the pH-dependency). This suggests that a complex

380 picture of B:Ca development in the EBC otoliths over the recent decades may be affected by  
381 multiple factors, not only by pH changes but also physiological stress caused by overall changes  
382 that have recently occurred in the Baltic (Eero et al., 2020). These include expansion of  
383 hypoxia/anoxia, eutrophication, and biogeochemical regime shift. Together with overfishing,  
384 these factors led to an overall collapse of the EBC population in the Baltic.

385 Examination of B:Ca in relation to an otolith index of hypoxia exposure, Mn:Mg,  
386 revealed mostly positive slopes of individual fish regressions, with the notable exceptions of  
387 Baltic cod from the 2010s and Icelandic cod (Figure [S5S4](#)). Limburg and Casini (2018) adopted  
388 the ratio Mn:Mg as a hypoxia exposure proxy rather than Mn:Ca in part to correct for growth  
389 influences on Mn (see Limburg et al., 2015) as otolith Mg is related to metabolism and growth  
390 (Limburg et al., 2018). The positive relationships of Mn:Mg to B:Ca cannot be fully explained,  
391 but the inverse relationships observed in the 2010s is consistent with the hypothesis that cod  
392 were exposed to both hypoxia and lower pH (and also higher alkalinity). Some of the positive  
393 slopes were due to young ages of fish; we observed more coupling of Mn:Mg and B:Ca in young  
394 fish, and we hypothesize that at younger ages the growth effect on otolith Mn is not fully  
395 removed by dividing by Mg. Baltic cod had increasingly negative slopes of the Mn:Mg vs. B:Ca  
396 regressions as they aged (Figure [S6S5](#)), with the 2010s displaying the most negative slopes.

397

## 398 **5. Conclusions**

399 Baltic cod showed variable patterns of otolith B:Ca over a 35-year period, with maximum  
400 values in the late 1990s, and a gradual decline in B:Ca thereafter, to an all-time low in 2019-20.  
401 These trends are imperfectly related to water chemistry data, but showed strong correlations that  
402 varied by decade. Otolith B:Ca ratios and relationships in the 2010s were suggestive of  
403 environmental changes in pH and total alkalinity that could be coupled to deoxygenation.  
404 Physiological influences observed in a healthy Icelandic population were diminished in Baltic  
405 cod.

406 Undoubtedly, some of the relationships found could be due to the choice of samples, and  
407 a more focused study might select otoliths from areas where trends in acidification are clear.  
408 Future work could also examine otoliths from other species, as well as determine, if possible, the

409 form of boron (borate or boric acid) taken up by otoliths. Complex interactions notwithstanding,  
410 we suggest that B:Ca in cod's otoliths can be a prospective variable in the palette of important  
411 and available tools to look into the environmental changes through the lens of otolith chemistry.

412

413 *Code/Data availability.* The codes and data are available upon request to the corresponding  
414 *author.*

415 *Author contributions:* KL designed and carried out the study. YH contributed data and figures.  
416 *All authors contributed text.*

417 *Competing interests.* The authors declare that they have no conflict of interest.

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420

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