# Supplement of From atmospheric water isotopes measurement to firn core interpretation in Adelie Land: A case study for isotope-enabled atmospheric models in Antarctica

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### Figure



**Figure S1:** Molecular mixing ratio (ppmv) calculated from Meteo France sensors (estimated by the Magnus-Tetens approximations from temperature, surface pressure and relative humidity), versus molecular mixing ratio measured by Picarro. Red line is the linear relationship estimated from hourly averaged data.



**Figure S2:** Humidity calibration curves showing the difference between the measured  $\delta^{18}O(\delta D)$  and the reference  $\delta^{18}O(\delta D)$  values (true value of the standard used: NEEM and FP5). The results of averaging measurements over 10-minutes. Measurements have been averaged over 10 minutes and error bars represent the standard deviation over this time period (data resolution is 1 second for 2018, 2019 and SDM measurements, 1 minutes for 2020 measurements).



**Figure S3:** Observed relationships between  $\delta^{18}O$  (‰) and humidity (ppmv) daily data according to the seasons. Red lines are the fitted linear relationships estimated from daily data (best linear fit in red).



**Figure S4:** Observed relationships between  $\delta^{18}$ O (‰) and 2-m temperature (°C) daily data according to the seasons (best linear fit in red: extrapolated over the -29°C – 2°C range).



**Figure S5:** Relationship between the water isotopic composition of 1) the vapor measured at DDU station (daily average) and 2) the vapor which would be at equilibrium with precipitation samples assuming isotopic equilibrium (event-basis sampling). The red line is the linear regression on all precipitation.



**Figure S6:** Meteorological and isotopic measurements at DDU at 6-hour resolution, in black. Panels from top to bottom: 1) 2m-temperature (°C) from Meteo France weather station, 2) humidity (ppmv) measured by the Picarro laser spectrometer; 3)  $\delta^{18}$ O (‰) in water vapor. Colored lines are ECHAM6-wiso first level outputs (6-hour resolution) at DDU closest grid cells (colors as in Fig. 1).



**Figure S7:** Comparison of humidity,  $\delta$ 18O and d-excess modeled by ECHAM6-wiso and measured at DDU station over the period 2019-2020. In black: Humidity and isotopic

measurements at DDU at daily resolution. Panels from top to bottom: 1) humidity (ppmv) measured by the Picarro laser spectrometer; 3)  $\delta^{18}O$  (‰) in water vapor. Green lines correspond to the ECHAM6-wiso combination of first level outputs (daily resolution) using isotopes optimisation (combination c), see main text).



**Figure S8:** Relationship between the measured isotopic composition of precipitation at DDU and the daily output from ECHAM6-wiso grid combination (see text). Red line is the linear regression.



**Figure S9:** Temporal evolution of the relationship between modeled and measured  $\delta^{18}$ O for the precipitation and the vapor water at DDU over the period 2019-2020. Panel a): Correlation (R<sup>2</sup>, yellow) and slope (pink) between modeled and measured  $\delta^{18}$ O in precipitation over a 3-month running window. Panel b):  $\delta^{18}$ O of precipitation from the ECHAM6-wiso model (violet) and from sample measurements (green dots with indication of the range of daily precipitation rate). Panel c): Correlation (R<sup>2</sup>, yellow) and slope (pink) between modeled and measured  $\delta^{18}$ O in the atmospheric vapor over a 3-month running window. Panel d):  $\delta^{18}$ O of the vapor obtained from the ECHAM6-wiso model (violet) and from the atmospheric vapor over a 3-month running window. Panel d):  $\delta^{18}$ O of the vapor obtained from the ECHAM6-wiso model (violet) and from measurements (green). Panel e): Temperature (red) and daily precipitation amount (blue) from EAR5 reanalyse.



**Figure S10:** Water stable isotopic composition of precipitation ( $\delta^{18}$ O in ‰) versus 2-meter temperature at DDU from measurements (left) and ECHAM6-wiso grid combination (over the 2019-2020 period). Linear fit in red. In ECHAM, we consider only precipitation with a total daily amount superior to 1kg.m-<sup>2</sup>. Note also that we used the daily precipitation amount as a weight to compute the linear relationship (relationship with unweighted values: slope: 0.2, intercept: -15).





**Figure S11**: a: moving standard deviation over 40 samples (approximately 5 years) of S1C1 firn core (black), ERA VFC (green) and ECHAM VFC (red). b: VFCs moving standard deviation after diffusion.



**Figure S12**: Relationship between the isotopic composition in vapor and in precipitation for measurement (left) and ECHAM6-wiso grid combination (right). Red lines are linear fits from daily means.

## Table

Grid cells	Temp (°C)		T2m (°C)	Hum (ppmv)		Q2m (ppmv)	δ <sup>18</sup> Ο (‰)		Temp Corr.		Hum corr		d18 corr		m.a.s.l.	mean topo (m.a.s.l.)
	mean	std	Mean	mean	std	Mean	mean	std	slope	R²	slope	R²	slope	R²		
#1	-12.3	8.0	-12.1	2332	1368	2389	-22.2	2.9	1.1	0.9	1.1	0.9	0.3	0.4	68	0
#3	-12.5	7.9	-12.2	2357	1406	2431	-21.9	3.0	1.1	0.9	1.1	0.9	0.3	0.4	72	0
Oceanic	-12.4	8.0	-12.2	2345	1386	2410	-22.1	2.9	1.1	0.9	1.1	0.9	0.3	0.4	70	0
#2	-18.1	8.1	-19.0	1195	1024	1115	-39.4	7.3	1.1	0.9	0.8	0.9	1.1	0.6	712	645
#4	-17.7	7.9	-18.6	1237	1079	1162	-38.6	7.1	1.1	0.9	0.8	0.9	1.1	0.6	633	516
Continental	-17.9	8.0	-18.8	1216	1050	1138	-39.8	7.1	1.1	0.9	0.8	0.9	1.1	0.6	673	581

**Table S1**: Comparison of ECHAM6-wiso outputs and data. Left section of the table: models outputs for the first level of grid cells as defined in Fig. 1; for comparison, model outputs at 2m available for temperature and humidity are also given. Middle section of the table: correlation coefficients between daily modelled outputs for each cell and data measured at DDU. Right section of the table: altitude of the grid cell center for the first level computed by ECHAM6-wiso.

ECHAM6-wiso grid cells (Fig. 1) of first level outputs results: Mean and standard deviation of meteorological parameters and isotopic composition (left section). For comparison, model outputs at 2m available for temperature and humidity are also given. Correlations between daily modeled outputs for each cell and data measured at DDU are displayed in the middle section of the table. The two right columns give the altitude of the center of the grid cell for the first level computed by ECHAM6-wiso and the altitude of the ground level at its coordinates. Bold lines correspond to combined (averaged) oceanic (#1 and #3) and continental (#2 and #4) cells.

		S1C1	ERA	5 tempera	ature	ECHAM6-wiso from precip.			
			wo diff.	diff.	diff. + strat.	wo diff.	diff.	diff. + strat.	
δ⁼0 (‰)	mean	-18.7	-20.8	-20.8	-20.8	-17.5	-17.5	-17.5	
	std	2.4	1.9	0.9	1.0	2.4	1.3	2.1	

**Table S2**: Mean, standard deviation of isotopic composition of S1C1 and VFC records built from ERA 5 temperature and ECHAM6-wiso precipitation (see text). Calculations were performed for each VFC for different configurations: 1) isotopic diffusion 2) after isotopic diffusion 3) after isotopic diffusion and addition of simulated stratigraphic noise (results presented are the averages of 40 draws of white noise simulation).

### Text

### Text S1:

The Matlab VFC scripts (described in Casado et al. (2020)) use as inputs the 2mtemperature and the precipitation amount to create layers of firn core at a density calculated through the Herron-Langway model (Herron and Langway, 1980) using forcing by surface temperature and accumulation. The temperature is converted in isotopic signal using a temperature to isotopes relationship (here estimated to be 0.44 %.° $C^{-1}$  at DDU). The total precipitation amount of ERA5 and ECHAM6-wiso are rescaled to match the mean annual amount of accumulation at the drilling site (21.8 ± 6.9 cm of w.e. year<sup>-1</sup>, Goursaud et al. (2017)). Here, the rescaling coefficient is 3.3 for ERA5 and 2.2 for ECHAM6-wiso (the total amount of snowfall in ERA5 and ECHAM6-wiso is larger than what has been actually accumulated at S1C1). VFC outputs are given with a resolution of 1 mm of w.e. Then, we resampled the VFCs using the density profile to match the resolution of S1C1 measurements (3cm of snow samples).

#### Text S2:

The effect of isotopic diffusion in the firn layers is estimated using the classical diffusion model from Johnsen et al. (2000) with addition of depth-dependent diffusion length (Laepple et al., 2018). We use Matlab VFC scripts as described in Casado et al. (2020). As expected, diffusion smooths the signal (Figure 6) and  $\delta^{18}$ O standard deviations in the VFCs become almost twice lower than in the S1C1 core (see Table S2 and Fig. S12).

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