Subsurface floats in the Filchner Trough provide first direct under-ice tracks of the circulation on shelf

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Abstract.
Bottom water formation in the Weddell Sea and mass loss from the Filchner-Ronne Ice Shelf are tightly linked by the

supply of Warm Deep Water to the continental shelf. Heavy sea ice cover and icebergs restrict ship access and upper ocean
measurements by moorings, compelling us to try new sampling methods. We present results from the first dedicated under-sea-
5 ice float experiment tracking circulation on the continental shelf of the eastern Weddell Sea. Seven Apex profiling floats were
deployed in 2017 at three different locations of the eastern Weddell Sea continental shelf, targeting the inflowing modified
Warm Deep Water (mWDW) as well as the outflowing Ice Shelf Water (ISW). The floats capture a warm mWDW regime with
southward inflow over the eastern continental shelf and a cold ISW regime with a recirculation of ISW in the Filchner Trough
throughout the four years of observations. We provide the first Lagrangian in situ confirmation that the mWDW flowing onto
10 the continental shelf follows two pathways: the eastern flank of the Filchner Trough and ~~via~~ a small trough on the shallow shelf
farther east. In the present circulation regime, this warm water is blocked from reaching the ice shelf cavity due to the presence
of the thick ISW layer inside the Filchner Trough. The floats' trajectories and hydrography reveal the dynamically active front,
flow reversal, and eddying motion between these two water masses along the eastern flank of the Filchner Trough.

1 Introduction

15 The ocean circulation at the Antarctic margin controls the amount of heat that can be transported toward ice shelves surrounding
Antarctica (Thompson et al., 2018). Ice shelves are the floating extensions of the Antarctic Ice Sheet and help stabilize the ice
sheet's flow toward the ocean (Dupont and Alley, 2005; Reese et al., 2018). By modulating ocean heat reaching ice shelf
cavities, the ocean circulation over the Antarctic continental shelves has a direct impact on the Antarctic Ice Sheet's mass
balance and therefore global sea-level rise (Joughin et al., 2012; DeConto and Pollard, 2016). In turn, modulation in the

20 freshwater input to the ocean through ocean-driven basal melting of the ice shelves affects water mass formation and ocean

circulation, and in particular dense water production feeding the global overturning circulation (Lago and England, 2019;

1
 Moorman et al., 2020). Understanding the ocean circulation over the Antarctic continental shelf, and how it varies, therefore
represents a key endeavour in our understanding of global climate.

On the Antarctic continent, the southern Weddell Sea is a key region for the formation of the Antarctic Bottom Water that
25 fills the bottom of the world oceans (Foldvik and Gammelsrød, 1988; Meredith et al., 2013). This region hosts the Filchner-

Ronne Ice Shelf (FRIS), Antarctica's largest ice shelf in volume (Fox et al., 1994). Its underlying cavity is categorized as a
"cold and dense" cavity, associated with relatively low basal melt rates and dense water masses (Nicholls et al., 2009). This ice
shelf regime is distinguished from "warm cavities" characterized by relatively large basal melt rates (Paolo et al., 2015; Rignot
et al., 2019) and "cold and fresh cavities" associated with relatively low basal melt and light water masses (Hattermann et al.,
30 2014; Silvano et al., 2016). These different regimes are maintained by ocean-cryosphere feedback involving the ice shelves,
sea ice, and ocean circulation (Jacobs et al., 1992; Thompson et al., 2018).

The FRIS in the Weddell Sea has had a relatively stable mass balance over past decades compared to warm cavity ice shelves
west of the Antarctic Peninsula (Rignot et al., 2019); it is thought to be relatively protected from the warming influence of the
Circumpolar Deep Water (CDW) of the Southern Ocean (Schmidtko et al., 2014) as a result of the presence of the Weddell
35 Gyre system (Vernet et al., 2019), but also because of the local frontal dynamics on the continental slope associated with the
cold and dense water over the shelf (Thompson et al., 2018). The circulation and hydrography on the shelf is dominated by the
production of High Salinity Shelf Water (HSSW, with temperatures at the surface freezing point ~-1.9oC) that is formed by sea
ice formation on the shelf, specifically within coastal polynyas (Haid and Timmermann, 2013). This dense water mass enters
the ice shelf cavity at its western entrance and transforms to a slightly colder and fresher Ice Shelf Water (ISW; temperatures
40 <-1.9oC) through interaction with the ice shelf while it circulates inside the cavity. ISW exits the cavity through the Filchner
Trough that forms a relatively shallow channel across the eastern continental shelf from the shelf break towards the ice shelf
(Nicholls et al., 2009). Below about 500 m the trough is a semi-closed basin retaining ISW behind its sill. While observations
have shown that the ISW flows northward along the eastern flank of the Filchner Trough (Darelius et al., 2014; Ryan et al.,
2017), there is also evidence that the ISW seasonally exits the cavity over the western flank (Darelius and Sallée, 2018). Once
45 the ISW overflows the continental shelf across the sill of the Filchner Trough (Darelius et al., 2009; Foldvik et al., 2004) it
contributes to the formation of Weddell Sea Deep Water and Weddell Sea Bottom Water, the precursors of Antarctic Bottom
Water. Recent studies have shown a reduction in volume of the Weddell Sea Bottom Water linked to

changes in dense water
production on the continental shelf (Zhou et al., 2023). While past work has revealed the processes involved in the formation
and modification of ISW, detailed information about its pathways and residence times are still lacking.
50 The existence of a pool of dense water over the southern Weddell Sea continental shelf is thought to create a barrier for ~~the~~
the advection of relatively warm water off the continental shelf toward the FRIS cavity. The warm water originates from the ~~CDW~~
CDW that upwells in the Southern Ocean; in the Weddell Sea, it is entrained into the Weddell Gyre system and transformed ~~into~~
into Warm Deep Water (WDW), a fresher and cooler version of CDW (Ryan et al., 2016). Conductivity-temperature-depth
(CTD)
sections and mooring time~~-~~ series have been only sporadically obtained in the area. These observations have suggested
55 that
55 there are two main pathways for WDW to enter the continental shelf and subsequently flow onto the
continental shelf and
subsequently southward toward the ice shelf as
a yet cooler and fresher version of the WDW, referred to as modified WDW ~~(mWDW): an eastern pathway,~~
~~along the eastern~~

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2
(mWDW): an eastern pathway, along the eastern flank of the Filchner Trough, and a western pathway, in the Central Trough
west of the Berkner Bank (Fig. 1; Nicholls et al.,
2009, 2008; Ryan et al., 2017). The western pathway is associated with
the largest heat content (Nicholls et al., 2008), and
about half of the heat ~~at the shelf break~~ that actually reaches the FRIS edge (Davis et al.,
60 2022). CTD observations have also
60 indirectly suggested that mWDW could enter the continental shelf by a third pathway
flowing through the "Small Trough" (see
Fig. 1), directly east of the Filchner Trough, but this pathway has never been directly
measured (Janout et al., 2021) Most of the heat advected
southward onto the continental shelf through the Filchner Trough
is thought to recirculate northward or to be mixed away
before reaching the FRIS cavity (Dinniman et al., 2011; Darelius
et al., 2016). The flow of mWDW into the Filchner Trough
occurs seasonally during austral summer and is linked to the wind
65 variability (Årthun et al., 2012; Daae et al., 2017; Ryan
65 et al., 2017). In 2017 the mWDW layer over the eastern flank of the
Filchner Trough was warmer than in previous years and
persisted for a longer time period (Ryan et al., 2020), but it is unclear
how far south the warm water reached.

Thus far, mWDW has ~~only~~ been observed at the ice front near 79∘S only once in 2013 (Darelius et al.,
2016). However, the
region has
been suggested to be prone to a potential self-reinforcing change in the ocean circulation that could abruptly shift
70 its cavity
regime, turning it from a cold and dense cavity into a warm cavity with important global consequences for deep

water formation

70 and sea level rise (Hellmer et al., 2012, 2017; Hazel and Stewart, 2020). As a result, efforts to understand
the controls on the
flux of warm water and heat transport rate to the FRIS cavity have received much attention in recent years (Årthun et al., 2012;
Darelius et al., 2016; Ryan et al., 2017; Daae et al., 2017; Ryan et al., 2020; Daae et al., 2020).

Over recent decades, the flow of warm water onto the continental shelf has been sampled using ship-based observations
75 (Conductivity-Temperature-Depth (CTD) casts), CTD-instrumented marine mammals, and mooring observations. Ship-based
75
observations are only possible during summer and are mostly concentrated in regions lightly covered by sea- ice during this
period. Marine mammal observations depend on seal behaviour (Labrousse et al., 2021), so cannot be targeted toward specific
currents or water- masses; also, they do not usually cover the entire water column and have larger measurement and positioning
uncertainties than ship-based observations (Siegelman et al., 2019). Mooring observations over the continental shelf (e.g. Ryan
80 et al., 2017; Darelius et al., 2014) sample at fixed points throughout the year, but possible damage by drifting icebergs leaves
80 instruments in the upper part of the water column vulnerable and rarely instrumented. Autonomous instruments such as Argo
floats have not been deployed on the southern Weddell Sea continental shelf. The various strategies used to observe the ocean
over the continental shelf complement each other, but important gaps can be identified in the present observational system
in this challenging ice-covered part of the ocean: a dearth of wintertime observations, a lack of observations of water-mass
85 mixing, and the lack of reliable information on circulation routes and time scales.
85
In this study, we employ an innovative approach for sampling water mass characteristics and circulation on the Weddell Sea
continental shelf with the aim of partly addressing this observational gap. We present and use, for the first time in this region,
observations from Lagrangian profiling floats. Seven Lagrangian floats were deployed on the continental shelf in February
2017. Three RAFOS sound sources were deployed specifically for this experiment, with RAFOS hydrophones on the floats
90 being used to provideinfer their locations every six hours. Although the ocean environmental conditions of the Weddell Sea con-
90 tinental
shelf make float deployments very risky for the instruments, due to the large area covered by multi-year ice and the
presence

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of ice shelf cavities where instruments can get lost, all of the floats survived at least one year and up to four years,

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providing
a unique set of observations. A similarThe approach was applied to determine float tracks undertaken in the present study is similar to that taken in a recent study in a different

region, the Dotson Ice Shelf in
the Amundsen Sea (Girton et al., 2019) at distances up to about 100km.

95 Our observations on the eastern Weddell Sea shelf allow us to describe the circulation in the Filchner Trough region and
95 the water mass pathways based on the float trajectories and the hydrography profiles, some of which are taken in previously
unexplored areas that are typically covered with sea ice. We are able to examine the Lagrangian pathways of the mWDW
toward the ice shelf, its interaction with the ISW, and the residence time of the ISW within the Filchner Trough. Selected
mooring observations and ship-based measurements from the same time period complementing the float data are presented.
100 This combined dataset allow us to present the first direct measurements of pathway and transit times for outflowing ISW
100 and inflowing mWDW. In particular, we demonstrate the existence of a mWDW pathway through a sSmall tTrough which was
previously suggested, but was never been directly measured. We also present the first direct observations of a pool of ISW with
very long
residence times of several years in the deeper part of the Filchner Trough.

2 Data and Methods

105 2.1 Autonomous RAFOS-enabled profiling floats

105 In February 2017, seven autonomous profiling Apex floats were deployed from the R/V James Clark Ross (cruise number
JR16004; Sallee, 2018) at three different locations on the Weddell Sea continental shelf (Fig. 1): three floats on the continental
slope upstream of the Filchner Trough (~26.02°W, 74.05°S); two at the eastern part of the Filchner Trough sill (~30.07°W,
74.08°S); and two farther south in the deeper part of the Filchner Trough (~32.07°W, 76.05°S). The floats were deployed in
110 open waters, before the seasonal advance of sea ice. Each of the floats carried a SeaBird SBE41 sensor to measure conductivity
110 (C), temperature (T), and pressure (P), as well as a hydrophone to record the timed sound source signals (Table 1). The RAFOS
protocol (Rossby et al., 1986) uses the timed arrivals of signals from the sound sources to determine the position. All floats
together provided in total 1106 CTD profiles over a time span of up to four years, across several hundreds of kilometers between
the Filchner and Brunt ice shelves. Six of these seven floats sampled a total of 244 profiles located over the continental shelf
115 between February 2017 and April 2018; one of the floats (P12681; Fig. 3) circulated northward after being deployed and did
115 not sample the shelf regime and is not discussed further.
Profiles of absolute salinity (SA), conservative temperature (Θ), potential density referenced to the surface (σ0), and pressure
(P) were derived from the observed profiles using the revised seawater equation of state (McDougall and Barker, 2011). The
profiling frequency was adjusted between summer and winter months: the floats profiled once per day in summer, and then

120 once every five days the rest of the year. The floats are programmed to start their (ascending) profile from 1800 dbar, which is

120 generally deeper than the continental shelf. This means that in most cases they descended to the sea floor before starting the

profile. Between profiles, the floats drifted at a programmed parking depth, and took CTD measurements at this depth every

six hours. The parking depth chosen depended on the deployment positions of the floats: the two southernmost floats, deployed

inside the Filchner Trough, were programmed to drift at 250 dbar to reduce the risk of drifting inside ice shelf cavities; the other

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125 floats were programmed to drift at 400 dbar to target typical depths where mWDW is found on the Weddell Sea continental

125 shelf (Årthun et al., 2012; Nicholls et al., 2009; Foldvik et al., 1985).

When a float surfaced, its positioning data and other recorded measurements were telemetered via the Iridium satellite

network. If a float were to get to the surface when sea ice was present, its sensors and antennas would likely be damaged. To

reduce this risk, each float implemented a "sea ice avoidance" logic to predict the presence of sea ice from the observed ocean

130 characteristics measured during ascent (Klatt et al., 2007; Wong and Riser, 2011; Porter et al., 2019; Silvano et al., 2019). If

130 the measured in situ temperature, averaged between 20–40 m, was less than -1.69∘C, the float halted its ascent, descended to

its parking depth and stored its data to transmit them at the next successful surfacing.

2.2 Positioning of the floats

The Filchner Trough is a region covered by sea ice during most of the year and by multi-year ice in some parts, preventing

135 floats from surfacing and acquiring GPS positions for their profiles. In order to locate the hydrographic profiles under ice and

135

to obtain high-resolution Lagrangian trajectories of the floats, three sound sources were moored in the eastern continental shelf

(74.02∘ S, 28.07∘ W; 74.85∘ S, 30.38∘ W; 75.39∘ S, 28.64∘ W; Fig. 1). Each source transmitted ~~with~~ a frequency modulated signal

~~signal~~ four times a day with a carrier of about 260 Hz, and frequency increasing linearly by 2.5 Hz during 80s seconds (Rossby et al.,

1986). The floats were programmed with listening windows consistent with the sound source transmission times, and saved

140 the six highest correlations between the transmitted and expected sound source signal, as well as the six times of arrival

140 (TOA)

of the sound signal corresponding to each of the six correlations. The stored set of six TOAs and maximum correlations were

~~were~~ then transmitted via the Iridium satellite network when the float eventually surfaced and used to triangulate the position of the

floats during their drift.

The TOAs are corrected for three main biases. First, the offset between the start of the float listening

window and the time
of ~~of~~

145 sound transmission for each source is removed from the received TOAs; for each listening window, the float started listening

~~145~~

one minute before the first source transmission, 21 minutes before the second source, and 41 minutes before the third source

transmission. Second, the TOAs were corrected for the linear drift of the float internal clock. Third, TOAs were adjusted to

account for the delay in activating the memory card that records them. The card's activation is delayed by a margin of 10 to ~~24~~

24 seconds after the initiation of the float's listening window, introducing a temporal offset within the TOA measurements.

150 Each

TOA was then quality checked using the corresponding maximum correlation, re-confirmed by eye, and outliers and unrealistic

~~150~~ records were removed, to eventually form a clean and validated time-series of TOAs for each float. The time-series

of TOAs was then

smoothed with a loess filter (with a tri-cube weight function), removing frequencies higher than once per day to produce a

better estimate of the daily mean position used here.

The corrected TOAs were converted into distances between the float and the sound source using a fixed sound speed in

155 water

of 1455 m s−1, derived from CTD measurements. Given that the float's hydrophones sampled every 0.3075 s during their

~~155~~

listening window, the minimum achievable spatial resolution of the triangulation was 477 m. Distances from the three sources

were then converted into geographical position by minimizing the error of the triangulation procedure. The geographical

position was ~~comput~~inferred for all time steps in which three sound source signals passing all quality controls were received. When

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only two sound source signals passed the quality control, computation of geographical position was attempted, but used only

160 if the calculated position was consistent with the position before and after.

~~160~~

The computed positions were then compared with the GPS positions for all periods when both were available. The differ-

ences between the GPS and the TOA-derived position ranged between 7 and 17 km, and this difference is considered here

as a

~~as a~~ rough estimate of the absolute error in positioning. The seven trajectories are displayed in Fig. 3 ~~, with their smoothed~~

~~trajectories obtained from the 24-hour low-pass filtered raw TOA~~.

2.3 Mooring and ship-based CTD data

165 In addition to the Lagrangian floats, in the present study we use the Eulerian time-series acquired at two mooring sites. One

of the two moorings (M30.5W) was located at 76.09◦ S, 30.47◦ W (Fig. 1) from 22 January 2016 to 04

February 2018 (Ryan
et al., 2020) at 445 m depth. It was deployed during the Polarstern cruise PS96 (Schröder, 2016) and recovered during the
Polarstern cruise PS111 (Janout et al., 2019; Schröder, 2018). Temperature was recorded at 111, 97, 73, 48 and 33 m above
bottom with additional salinity and pressure measurements at the shallowest and the deepest depths, and velocity at 18 m above

170 bottom. The other mooring (P5) was located farther north at 75.39° S, 28.64° W at 437 m depth (Fig. 1). It was deployed on 09
February 2017 during the cruise of the float deployment and recovered on 09 March 2021 during the Polarstern cruise PS124
(Hellmer, 2020). Temperature was recorded at six different depths (7, 30, 55, 75, 104, 114 m above bottom) and velocity at
three depths (6, 53, 102 m above bottom). Here we use the temperature and velocity measurements at 382 m and 384 m depth,
respectively, which are the closest measurements to the float drift depth.

175 We also use the hydrographic data from the ship-based CTD casts of the float deployment cruise to calculate the thickness
of the ISW layer. We calculate the ISW thickness at each profile (both from the ship-based CTD casts and the float profiles)
as the total thickness of all points where $\Theta < -1.9$°C. We only include layers that are located in the bottom layer (i.e. with the
lowermost measurement of $\Theta < -1.9$°C at <50 dbar from the bottom), to exclude measurements of Winter Water that may also
be colder than the surface freezing point but located higher up in the water column.

180 3 Results

3.1 Float trajectories

The profiling floats deployed on the Weddell Sea continental shelf drifted and profiled for a time period of three months to
four years, providing hydrography and drift data over the eastern portion of sea-ice-covered continental shelf. With the use of
three sound sources, we were able to recover the float locations with a sampling interval of six hours and of up to one year
185 duration.
The maximum distance from the sound sources for which good quality TOAs were obtained was ~350 km (Fig. 2).
This is a
significantly shorter range than in open ocean experiments with sources transmitting in the SOFAR channel (typically
800-1000 m
depth; Rossby and Webb, 1970; Wong and Riser, 2011; Rossby et al., 1993), likely because of acoustic refraction
coupled with

rough reflecting interfaces that attenuate the signal. Our experiment deliberately avoided targeting flow under the
ice shelf to
focus on the dynamics of the ISW and mWDW, and their interactions, on the continental shelf.

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190 Two floats deployed in Feb. 2017 on the continental slope (P12682 and P12677; Fig. 3) followed the isobaths westward and
then turned south onto the continental shelf at the opening to a small canyon cutting across the continental shelf between the
Filchner Trough and Brünt Ice Shelf. We will here refer to this small canyon as the Small Trough (Fig. 1). One of the two floats
(P12682) moved southward through the Small Trough over the shallow eastern plateau until it reached the Filchner Trough at
∼76.5◦ S in early May. On May 23, its trajectory then sharply veered from a southward direction to a northward flow along
195 the lower eastern flank of the Filchner Trough. It continued profiling within the Filchner Trough, but without information on
its position (Fig. 4c). The other float (P12677) started off in a similar direction, but it stopped recording on May 20, when it
was still located within the Small Trough on its way towards the Filchner Trough.
The two floats deployed Feb. 2017 on the Filchner Trough Sill (P12686 and P12703, Fig. 3) were deployed only 6 km and
7 hours apart, but followed very distinct routes. One of the two (P12703) started off northward towards the continental slope,
200 but turned southward on February 28 and continued above the eastern plateau towards the Filchner Trough. It reached ∼76.7◦S
on May 1, where it abruptly turned northward again at a similar location to where P12682 changed direction. In contrast, the
second float (P12686) started off on a southward track close to the upper eastern flank of the Filchner Trough but changed
its direction abruptly at ∼75.5◦ S, where it drifted back northward along the lower eastern flank of the Filchner Trough. It
continued westward along the Filchner Trough sill, leaving the Filchner Trough to follow isobaths over the continental slope.
205 All four of these floats (P12682, P12677, P12686 and P12703) had southward and relatively straight trajectories at different
locations so long as they remained over the eastern shallow plateau, or upper eastern flank, above 400-500 m bathymetry depth.
As soon as the floats reached deeper bathymetry in the Filchner Trough, their directions changed and they ultimately drifted
back northward.
To capture the flow of the ISW within the Filchner Trough, two additional floats were deployed directly inside the Filchner
210 Trough, at ∼76.5◦ S (P12679 and P12684; Fig. 3). One of these two floats (P12679) could only be positioned for one month,
during which it followed an eddying pattern close to its deployment position; however, it was able to profile the water column
for more than a year, which it spent entirely within the Filchner Trough as indicated by the sampled ISW layer (Fig. 5c,d).
The second float (P12684) went northward toward 75.4◦ S following the eastern slope of the Filchner Trough before flowing
back southward to 76.5◦ S. After several months of missing RAFOS positions, it reappeared at the western side of the Filchner
215 Trough up to the Berkner Bank (39◦ W), providing Lagrangian positions more than one year after its deployment. Compared
to the persistent southward trajectories on the eastern plateau, the trajectories of these two floats that are parked at 200 m
(shallower than the eastern plateau floats) within the Filchner Trough show a strong eddying pattern and a recirculation within
the Filchner Trough.

3.2 Modified Warm Deep Water and Ice Shelf Water regimes

220 The two groups of floats identified above appear to document two distinct regimes associated with two different water masses
characterizing the eastern Weddell Sea continental shelf (Fig. 3b). While the first group of four floats (12677, 12682, 12686 ~~and~~

~~7~~

~~12703) are associated with an eastern shallow plateau regime of mWDW~~
and 12703) are associated with a regime of circulation of mWDW on the eastern shallow continental shelf, the second group
of two floats (12679 and 12684) is
associated with a Filchner Trough regime of ISW circulation.

7
To further highlight these two regimes, we investigate the temperature of the water column in the layers that host mWDW
225 and ISW (Fig 6). mWDW is defined as the water mass of temperature greater than -1.7∘C in the layer of potential density

σ0 = 27.75 ± 0.05 kg m−3, and ISW as the water mass of temperature lower than the surface freezing point (∼-1.9∘C) in the
bottom ocean layer (Nicholls et al., 2009). Based on these definitions, we describe below (i) the circulation and timescale of
advection of mWDW on the eastern plateau; (ii) the interplay between mWDW and ISW on the eastern plateau; and (iii) the
residence time of ISW in the Filchner Trough.
230 The four floats drifting southward on the eastern shallow plateau are associated with temperatures greater than -1.6∘C (Fig
~~(Fig~~ 6a), suggesting that the region is flooded with southward-flowing mWDW. Based on the float trajectories, the southward
mWDW transport ~~is most~~appears to be more efficient on the shallow eastern plateau, compared to a pathway closer to the Filchner
Trough. Indeed,
float P12686, which has the southward trajectory closest to the Filchner Trough, does not reach far south but
gets quickly
entrained into a northward flow that advects the float out of the Filchner Trough and then westward along the
235 continental slope.
~~235~~ In contrast, the two floats (P12682 and P12703) that were drifting farther east closer to the coastline were
transported much
farther south (100-200 km). Their trajectories also indicate that at least two main pathways occurred in 2017 for the advection
of mWDW from the continental slope onto the continental shelf (Fig. 1): the Small Trough (P12682) and the northeastern
corner of the Filchner Trough (P12703). It takes the floats about two months to reach the interior Filchner Trough
from the
shelf break: 64 days through the Small Trough and 54 days along the eastern flank of the Filchner Trough from the
240 shelf break
~~240~~ at 74.6∘S (the two floats' northernmost common latitude) to reach 76.5∘S (the two floats' southernmost common
latitude).

The notion that the Small Through provides a conduit for mWDW from the continental slope to the Filchner Trough
is further
supported by the velocity and temperature time series taken at the sound source mooring P5 positioned there (Fig 8).
The float
passed within 8 km of this mooring and the trajectory derived from the mooring velocities at the approximate float
drift depth
agrees well with the float trajectory. These mooring observations show a south-southwestward flow throughout the
245 whole year
245 with temperatures in the mWDW range from February to June.

Interestingly, neither of these two floats went farther south than ~76.7◦ S, where the eastern shallow plateau narrows.
Instead, they became entrained in a northward flow. This turning point of the floats also coincides with the southernmost extent
of the warm water in the climatology of the temperature at about 400 m depth shown in Fig. 1. We suggest that in this region poses
there is a bathymetric constraint on the southward flow: the narrowing shelf forces the flow off the relatively flat, shallow
250 isobaths of
250 the plateau and onto the eastern flank of the Filchner Trough that is dominated by a northward flow and colder
water masses. At this
point, the floats' abrupt turn is accompanied by eddying motion suggestive of mixing with ISW along the eastern flank (Fig.
6a). One of the floats (P12682) crossed a clear front between the mWDW and the ISW in the beginning of June when entering
the Filchner Trough (Fig. 4c), which coincides with the abrupt switch from a southward flow, within the mWDW layer, to a
northward flow within the ISW layer (Fig. 4d).
255 The strong horizontal shear on the eastern side of the Filchner Trough, inferred from the opposing float directions over
the eastern flank of the Filchner Trough, is likely
associated with instabilities and the resulting mixing of water masses.

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Unfortunately, the set of Lagrangian trajectories does
not provide enough spatial and temporal sampling to address mixing or
instability questions fully here. The trajectories of the

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floats show, however, some eddying motion at along the eastern flank of
the Filchner Trough, which might be an indication of
instabilities and eddy activity, hence turbulent mixing between mWDW
260 and ISW.

260 The temperature of the water lying at the bottom of the continental shelf indicates the presence of an ISW layer (Fig. 6b)
that is observed at all profiles within the Filchner Trough, but also at the bottom of the eastern shallow plateau at around 76◦S.
The ISW layer reaches to a depth of about 250 m within the area sampled by the floats inside the Filchner Trough (Fig. 5; see
Section 2.3 for details on how ISW thickness is computed). While the ISW layer here is up to 1000 m thick

at the deepest part
265 of the trough close to the ice front, the ISW layer on the shallower shelf is much thinner (Fig. 7a). The float profiles inside the
265 Filchner Trough also show that the ISW layer persists all year round, with only small variations in depth and temperature (Fig.
5). The ISW layer on the eastern plateau is sampled by float P12703 during a short period in early April (Fig. 4e), shortly before
it passed the mooring M30.5W that provides information on the time variability of the ISW layer in this shallower region (Fig.
7b). The mooring temperature records show that the presence of an ISW layer was seasonally variable: ISW was present at the
270 bottom of the mooring site during austral summer (December to March; Fig. 7b) before the continental shelf got flooded by
270 mWDW during autumn and winter (March to October). After the ISW has been replaced by mWDW, there were still several
intrusions of ISW onto the continental shelf, during which the float passed the mooring site. The seasonal interplay between
the ISW and the mWDW on the shallow eastern plateau has previously been described in more detail in Ryan et al. (2017).
Our floats show that the mWDW is advected from the continental slope through the Small Trough and along the eastern flank
275 of the Filchner Trough with~~in~~ a timescale of about 2 months.

275 The two floats deployed on the eastern flank of the Filchner Trough (P12679 and 12684) drifted at the upper limit of the thick
ISW layer (drift depth of about 250 dbar) that fills the Filchner Trough (Fig. 5). The two floats circulated inside the Filchner
Trough for one and four years, respectively. Based on the ~~existing~~available positions, they circulated as far south as 76.64∘ S, and the
northernmost position is the last GPS position sent out by 12684 in January 2021 at 75.18∘ S. Despite the missing locations
280 during large parts of the sampling period, the temperature measurements show the presence of a thick ISW throughout their
280 whole measurement period. The depths of the profiles suggest that both floats moved from the shallow eastern flank of the
Filchner Trough to the central or even western part of the trough, collecting profiles up to 984 m depth. For both floats, the
temperatures within the ISW layer show a transition towards colder temperatures of the ISW during austral summer 2017/18
compared to the previous summer.
285
This result is in agreement with the change from the Berkner Mode towards the Ronne Mode described in Janout et al.
285 (2021)~~; and~~ Hattermann et al. (2021), ~~where~~in which the ISW temperature changes based on the formation site of High Salinity Shelf
Water. It is striking that neither of the two floats was advected northward out of the Filchner Trough in one year (for float
P12679) or even four years (for float P12684), but instead recirculated within the Trough. This is indicated by both the position
of the floats and, when not available, by the hydrography and depth of the water-column sampled by the floats (Fig 5). The
290 other floats, however, which were advected northward along the eastern flank of the Filchner Trough, seem to have left the
290 continental shelf eventually, as indicated by the trajectory of float P12686 that continued westward along the continental shelf

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west of Filchner Trough and by the hydrography of the two floats P12682 and P12703 that ~~got~~were entrained in the northward ~~flow~~

flow of ISW at about ~76.7◦S. The latter two floats started sampling a deep layer of warm water in austral summer 2017/2018 ~~with~~

9

with a profile depth of more than 1600 m which can only be located over the deep ocean off the continental shelf. We conclude

295 that the flow within the deeper part of the Filchner Trough is recirculating within the trough, keeping the ISW layer trapped

295 for several years, while the ISW closer to the eastern flank of the Filchner Trough is advected northward and across the sill

relatively quickly (within one year).

4 Discussion

We deployed seven autonomous floats to document the circulation of two key water masses on the southern Weddell Sea

300 continental shelf. The year-long Lagrangian trajectories within this sea -ice -covered region, as well as the up to four year-long

300 full-depth CTD profiles along the floats' trajectories, shed new light on pathways and timescales associated with mWDW inflow to the continental shelf, on the circulation of ISW within the Filchner Trough, and on the mixing between the two water masses.

The floats sampled the main inflow pathways of mWDW from the continental slope toward the continental shelf. Consistent

305 with previous work (Darelius et al., 2016; Årthun et al., 2012), our observations describe a southward flow of mWDW into

305 the

Filchner Trough across its sill at ~30◦ W. In addition, we demonstrate that mWDW is also advected southward through ~~the~~

the Small Trough on the eastern shallower continental shelf, at ~28◦ W (Fig 1). This alternative passage onto the continental ~~shelf~~

shelf was previously hypothesised but never ~~demonstrat~~directly observed as a possible gateway for mWDW (e.g. Ryan et al., 2020).
Other evidence

supporting this notion of an inflow pathway through the Small Trough is provided by current and temperature

measurements

310 taken over one year by the mooring P5 located in the Small Trough. Also, consistent with this flow, Labrousse

310 et al. (2021)

described the presence of a large pool of mWDW in 2017 over the entire continental shelf east of the Filchner

Trough, found

as far east as 28◦W, although their observations were slightly later in the year, in June and July. This collection of

observations

thus suggests that both the northeastern corner of the Filchner Trough and the Small Trough support a southward ~~mWDW~~

mWDW pathway in summer. Limitations on our observational period des not permit us to infer interannual variability, but we
note that
315 2017, the year in which we acquired most of our float observations, was described by Ryan et al. (2020) as a year of
315 particularly
warm inflow of mWDW. They proposed that the heaving of the Antarctic Slope Front was stronger in 2017 than in
other years,
allowing for warmer, saltier mWDW to intrude on the continental shelf and for the mWDW inflow to last longer.

All of the floats presented here that capture the southward flow of mWDW consistently depict the eventual entrainment of
southward flowing mWDW into a northward flow along the eastern side of the Filchner Trough. Such a turning point in the
320 trajectories was also found in idealized modelling results from the Filchner Trough region by Daae et al. (2017), where in
320 several experiments the southward flowing current along the eastern flank of the trough only reaches about 50 km into the
trough before it gets drawn out of the trough again by the northward flow of ISW. This blocking of the mWDW inflow by the
presence of ISW is consistent with the absence of mWDW close to FRIS that was earlier described from in situ observations
(Darelius et al., 2016). Here, we provide a novel view on the southward extent of the mWDW pathway: mWDW entering the

10
325 continental shelf farther east from the Filchner Trough is able to penetrate a greater distance to the south toward the ice shelf.
325 Hence, the Small Trough may be of particular importance for the impact of warm water on the shelf. At the same time, our
10
observations underline the importance of the ISW presence and its northward advection on blocking mWDW access to the
FRIS cavity. Any change in ISW production, which directly relates to the sea-ice driven production of HSSW (Hattermann
et al., 2021; Janout et al., 2021), would therefore provide a feedback on changes in mWDW access to the FRIS cavity, and
therefore on the heat flux to the base of the ice shelf.
330 While our observations indicate that the northward ISW pathway on the eastern side of the Filchner Trough blocks mWDW
access to the FRIS cavity, the two floats that sampled ISW in the deeper part of the trough showed convolutedrecirculating trajectories and
and long residence time-scales of one year (P12679) and up to—possibly more than—four years (P12684), respectively. Such a
a long residence time was so far unknown and could only be revealed with under-ice float data. Although the reconstitution of
of trajectories was relatively poor for those two floats, their hydrography observations and irregular available positionings (either
335 through

335 RAFOS or GPS) indicate that ISW in the deeper part of the depression stays in the region. The outflow of ISW is
therefore
likely confined to specific pathways, topographically linked to the flanks of the Filchner Trough. The pathway

of
ISW outflow
along the eastern flank of the Filchner Trough is clearly demonstrated by the four floats that first sampled the
southward flow of
mWDW before getting entrained in a northward flow of ISW as soon as they reached over deeper isobaths within the Filchner
Trough. A northward pathway of ISW along the western flank is not observed by the floats, despite their
340 presence on the

~~340~~ western flank. However, we cannot exclude an outflow of ISW on the western side of the trough, along which
ISW has been
observed to exit the FRIS cavity seasonally (Darelius and Sallée, 2018).
The strong horizontal shear on the eastern side of the Filchner Trough, inferred from the opposing float directions over
the eastern flank of the Filchner Trough, is likely associated with instabilities and the resulting mixing of water masses.
Unfortunately, the set of Lagrangian trajectories does not provide enough spatial and temporal sampling to address mixing or
345 instability questions fully here. The trajectories of the floats show, however, some eddying motion ~~at~~
along the eastern flank of
the Filchner Trough, which might be an indication of instabilities and eddy activity, hence turbulent mixing, between mWDW
and ISW.

5 Conclusions

The first Lagrangian ~~float~~ observations ~~were acquired~~ over the Filchner–Ronne continental shelf were acquired from 2017 onward. This dataset
350 provides year-round observations, combining full depth profiles and high resolution (6-hourly) Lagrangian trajectories under

sea ice by acoustic positioning. Two mWDW inflow gateways are documented: ~~the~~an eastern ~~most~~ passage, the Small Trough, at
~28◦W, that was previously hypothesised to be a pathway but never before observed as a mWDW passage; ~~the second passage~~and a western
~~is~~passage in the northeastern corner of the Filchner Trough, at ~30◦W. Both gateways are important for the seasonal mWDW
flooding
of the shallow continental shelf east of the Filchner Trough.
355 Our observations corroborate the importance of the exchanges between ISW and mWDW on the continental shelf, with the
~~the~~ presence of a thick ISW layer within the Filchner Trough blocking mWDW access to the FRIS cavity, by consistently
entraining
and mixing mWDW within an efficient ISW northward flow over the eastern slope of the Filchner Trough. Within

~~11~~
the Trough
itself, the circulation of ISW appears to be more convoluted and is associated with significant recirculation, with
a recirculation

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time- scale of up to at least four years, as indicated by one of our floats. The results from the deployment of 360 a set of autonomous
360 floats on the ice-covered Filchner-Ronne sector of the Antarctic continental shelf is demonstrated here to be a promising tool

for the observation of year-round hydrography and circulation of the Antarctic continental shelves, in combination with other
observational strategies.

Author contributions. JBS and LV conceived and designed the analysis. AM developed the algorithm for float positioning under the super-
vision of JBS and AL. NS and EP contributed to the visualisation of data on all the figures. AL and PL were instrumental in acquiring high
365 quality observations, especially around the deployment and programming of sound sources and floats. All authors contributed to the writing
of the manuscript.

Data availability. All Flfloats observations used in this paper are available here : https://doi.org/10.5281/zenodo.10353500

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Table 1. List and characteristics of all floats deployed in 2017.

Float	Deployment Location	Deployment time (UTC)	Last CTD profile with RAFOS positions	Parking depth [m]	Number of profiles	Number of profiles
P12677	26.188°W, 74.443°S	10/02/2017 06:29	31/03/2020	400	31	137
P12679	32.742°W, 76.465°S	18/02/2017 14:16	24/03/2018	250	49	220
P12681	26.163°W, 74.461°S	10/02/2017 05:44	11/01/2021	400	38	127
P12682	26.183°W, 74.453°S	10/02/2017 06:06	07/01/2021	400	29	90
P12684	32.718°W, 76.473°S	18/02/2017 14:47	14/01/2021	250	75	192
P12686	30.714°W, 74.849°S	12/02/2017 22:26	25/01/2018	400		
P12703	30.646°W, 74.853°S	12/02/2017 22:48	30/10/2020	400		

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Figure 1. a) Map of the eastern Weddell Sea continental shelf with a schematic of the circulation. The arrows show the characteristic
pathways of WDW (red), mWDW (dashed red) and ISW (blue). The colored background is the mean

conservative temperature (Θ) between
360 and 420 m, from Jourdain et al. (2020). Bathymetric contours from IBSCO (Arndt et al., 2013) are drawn every 100 m above 1000 m
depth and every 1000 m ~~thence, with~~ in regions deeper than 1000 m. The 500 m isobath is shown in bold as reference. Grounded ice is shown as gray areas and ice shelves as white areas. The black diamonds represent the positions of the sound source moorings, the cyan crosses the positions of the two moorings used in the study, and the green dots the launching sites of the floats. The map is presented in the Lambert projection. In this area, 1∘ longitude corresponds to about 30 km. The inset in the lower right shows the location within Antarctica.

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Figure 2. Histogram of the ~~maximum~~ distance between the floats and the sound sources. Distances of the floats to the sound source are used for all the positions recovered with two or more Time of Arrival (TOA) records.

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Figure 3. (a) Drifting positions of the seven floats deployed in 2017. The smoothed positions computed from the RAFOS system are shown
in color (squares), with the deployment location highlighted with a large circle of corresponding color. Note that P12682 is deployed at the
same location as P12677. The crosses along the trajectories represent the locations of vertical profiles used in (b). The colored background
represents the bathymetry according to the colorbar, with contours as in Fig. 1. (b) Θ-S diagram of all profiles taken by the floats, color
coded by the water depth at the location of the profile ~~inferred from the~~ estimated as the maximum profile depth). Contours show the density anomaly and the black line represent the surface freezing point (Tf).

21
 The black circles (mooring) and green (float) circles along the trajectories mark the first day of each month according to the
500 labels on the side in corresponding colors, and the first measurement dates are marked with big circles according to the legend.

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Figure 4. (a,c,e,g) Depth-time plot of conservative temperature for the floats that sampled mWDW on the shelf, Floats 12677, 12682, 12703
and 12686. The white contours represent the drift depth of the floats between the profiles. The cyan contours are the -1.9oC isotherm ~~for the~~
,
corresponding to the definition of ISW, and the red contours are the 27.75 kg m−3 isopycna ~~for the~~,
corresponding to the definition of mWDW. The triangles at the upper x-axis show the timing of the profiles, where those with existent RAFOS position are colored black. The last available position is marked with a vertical black line. (b,d,f,h) show the RAFOS trajectories of the floats on top of the bathymetry (as in Fig. 1), with the drift positions in gray and the profile positions in black, corresponding to the black triangles in (a,c,e,g). The deployment location is marked with a while circle. The colored circles are the existing positions closest to the 1st of each month, corresponding to the colored triangles at the lower x-axis in (a,c,e,g). Although the CTD timeseries are up to four years, we only show the first year in which we have ~~RAFOS positions and when the floats are still over the continental shelf.~~
Figure 5. Same as Fig. 4, but for the floats that sampled ISW within the Filchner Trough, Floats 12684 (a,b) and 12679 (c,d).
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Figure 6. a) Conservative temperature (Θ; colored squares) of the float profiles at the 27.75 kg m−3 ± 0.05 kg m−3 density, showing the Θ
in the mWDW density range. b) Bottom Θ of the float profiles (colored squares). The grey lines in a) and b) are the RAFOS drifting positions
of the floats. The colored background represents the bathymetry with contours as in Fig. 1

Figure 7. a) ISW layer thickness from the float profiles (squares) and the ship-based CTD casts during the deployment cruise in 2017
(triangles) over the Weddell Sea continental shelf. The colored filling of the symbols show the thickness of the ISW layer (only if any ISW
present), taken as the bottom water mass below the surface freezing point (-1.9◦C). The colored background represents the bathymetry with
contours as in Fig. 1). b) Hovmöller of the conservative temperature (Θ) of the mooring M30.5W (see its position in panel a)). The black
contour represents the -1.9◦C contour for ISW and the gray contours is the -1.7◦C contour for mWDW. The red triangle shows the timing
when the float P12703 passes closest to the mooring location. The black diamonds show the depths of the temperature sensors.

Figure 8. Progressive vector diagram of the current at 430 m depth measured at the sound source mooring P5 located in the Small Trough
(thick line; see mooring location in Fig. 1) and of the trajectory of float P12682 that passes the mooring location (thin line). The displacements
are centered around the time when the float passed the mooring for better comparison. The color shading denotes conservative temperature
(Θ). The black ~~(mooring) and green (float)~~ circles along the <u>mooring</u> trajectories mark the first day of each month according to the labels on the side in corresponding colors~~, and t~~. <u>Similarly, the first day of each month along the float trajectory is shown as green circles. T</u>he first measurement dates are marked with big circles according to the legend.