https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Brief Communication: Antarctic sea ice loss brings observed trends into agreement with climate models

Caroline R. Holmes[1], Thomas J. Bracegirdle[1], Paul R. Holland[1], Julienne Stroeve[2,3,4], Jeremy Wilkinson[1]

[1]British Antarctic Survey, Cambridge, CB3 0ET, UK
[2]National Snow and Ice Data Center, Cooperative Institute for Research in Environmental Sciences, University of Colorado, Boulder, Colorado, USA
[3]Centre for Earth Observation Science, University of Manitoba, Winnipeg, Manitoba, Canada
[4]Earth Sciences, University College London, London, UK

Correspondence to: Caroline R. Holmes (calmes@bas.ac.uk)

Abstract. Most climate models do not reproduce the 1979-2014 increase in Antarctic sea ice cover. This was a contributing factor in successive Intergovernmental Panel on Climate Change (IPCC) reports allocating low confidence to model projections of sea ice over the 21st century. We show that recent rapid declines bring observed sea ice area trends into line with the models. This implies that projections of substantial future Antarctic sea ice loss may be more reliable than previously thought, with substantial implications for the evolution of the Southern Hemisphere climate.

1 Introduction

The early years of the twenty-first century revealed a puzzling conundrum in Antarctic sea ice (Turner and Comiso, 2017; National Academies of Sciences and Medicine, 2017). Observations of Antarctic sea ice extent (SIE) showed a small increase during the satellite era (which began in late 1978), with annual mean values reaching a maximum in 2014, but most climate models simulated SIE declines over the same period. Various studies examined possible reasons for this discrepancy (Turner and Comiso, 2017). Specifically, the community discussed whether it could be explained by internal variability masking the anthropogenic forced signal in observations (Gagné et al., 2015; Rosenblum and Eisenman, 2017; Roach et al., 2020) and the extent to which it revealed model deficiencies in sea ice processes (Fox-Kemper, 2021). Others suggested that trends in sea ice, particularly SIE, may not be a robust metric of model performance, particularly when the observational time series is too short to separate internal variability from anthropogenic forcing (Notz, 2014). Even so, the poorly understood discrepancy has been a contributing factor in a widespread lack of confidence in projections of 21st century Antarctic sea ice decline, and consequently in many aspects of projected climate change around Antarctica which are underpinned by projections of substantial sea ice decline (Bracegirdle et al., 2015; Bracegirdle et al., 2018)

Recently, Antarctic sea ice has exhibited a starkly different pattern of behaviour. Following the pre-2015 era of slightly increasing ice extent, rapid ice loss beginning in early 2015 culminated in a dramatic drop in spring 2016-17 (Turner et al.,

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

2017). This led to several years of record low SIE, which has been framed as a 'new sea ice state' (Purich and Doddridge, 2023). This situation shows no sign of abating, with further declines since 2021 leading to monthly-mean SIE records being broken in all but two months of 2023 so far (Siegert et al., 2023). The initial decline showed strong linkages to patterns of intrinsic atmospheric variability (Turner et al., 2017; Schlosser et al., 2018; Zhang et al., 2022) which have high internal variability on short (sub-annual) timescales. However, growing evidence of the contribution of warming in the subsurface ocean (Zhang et al., 2022; Purich and Doddridge, 2023), and the magnitude and spatial homogeneity of the sea ice reductions since 2016/17, point to more sustained declines.

In light of this sea ice loss, we re-consider whether trends simulated by the current generation of climate models, from the Coupled Model Intercomparison Project Phase 6 (CMIP6; (Eyring et al., 2016)) dataset, are consistent with observations. Whilst previous studies have considered trends to 2013 based on CMIP5 models (Rosenblum and Eisenman, 2017) and to 2018 based on CMIP6 (Roach et al., 2020), we might expect the situation to have changed for two reasons. First, being able to assess trends in longer timeseries (due to the longer observational record) potentially reduces the impact of short-term internal variability on trend calculations (Notz, 2014). Second, and more specifically, these data now include the recent years of observed rapid decline of sea ice, impacting long-term trends.

2. Data and Methods

2.1 Sea Ice Metric

Sea ice cover is calculated as either sea ice extent, SIE (the total area of all gridboxes where sea ice concentration SIC exceeds a 15% threshold), or sea ice area, SIA (the sum of gridbox areas multiplied by gridbox SIC). SIA has larger observational uncertainties, as it is more sensitive to differences in SIC. However, SIE is a non-linear measure and so can give misleading results when comparing models and observations or when calculating trends (Notz, 2014). Therefore, in contrast to some previous assessments but following community precedent (Roach et al., 2020), we assess SIA. SIA and SIE have similar trends (Figure B1).

2.2 Model Data

CMIP6 monthly SIA is obtained from the University of Hamburg (UHH) CMIP6 Sea Ice Area directory (https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/cmip6-sea-ice-area.html, accessed 2023-08-17) and aggregated into weighted annual means. This is supplemented by SIA for the two NorESM models, which are not available in the UHH dataset due to a bug in an earlier version of NorESM released SIA data. We merge historical simulations ending in December 2014 with the ssp585 forcing scenario run for 2015 to 2023. ssp585 indicates a global average radiative forcing of 8.5 W m^{-2} by 2100 (O'neill et al., 2016). This is a high-emissions forcing scenario; however, emissions scenarios have little bearing on model spread for the time period considered here. The resulting historical-ssp585 merger constitutes 188

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

ensemble members from 39 models (Table A1), each contributing between 1 and 57 members of an initial condition ensemble. Since only four models contain more than six members, we use a maximum of six members from each model to avoid weighting the results too heavily to models with large ensembles, leading to using 98 members from 39 models. The sensitivity of our results to this treatment of model ensembles and to the emission scenario is discussed in Appendix C.

2.3 Observational data

For an estimate of observed sea ice cover, NSIDC Sea Ice Index v3.0 SIA (Fetterer, 2017) is used, available from January 1979-September 2023. We investigate the role of observational uncertainty by also using other observational estimates for 1979-2019 from the UHH SIA dataset (Dörr, 2021). Data for missing months (December 1987-January 1988 for the Sea Ice Index v3.0) are infilled by interpolating between the same month in the previous and following year (Rosenblum and Eisenman, 2017). To allow analysis of 2023, we create synthetic extensions of the data for October-December 2023, by linearly decaying the anomaly from the 2017-2022 mean to zero by January 2024. A conservative alternative, of decaying instead to the high-valued 2008-2016 mean by January 2024, has negligible impact on the p-values shown in Fig. 1c, so our analysis is robust to this assumption.

2.4 Trend evaluation methodology

Our evaluation methodology is an extension of that previously used for CMIP5 (Rosenblum and Eisenman, 2017). Linear trends are calculated for all periods of at least 35 years overlapping with the satellite record (January 1979-September 2023) using the OLS method of the python statsmodels.api package. We calculate the mean and standard deviation of the model ensemble trends and use these to fit a Gaussian distribution, with cumulative distribution function F(X), to the distribution of modelled trends. To estimate the probability that a trend at least as large as observed would occur in the climate model population, the p-value for a one-tailed test is 1-F(x) where x is the observed trend.

3 Results

3.1 Trend evaluation

The recent decade of data has reduced the significant positive trend (Parkinson, 2019) in observed annual-mean and monthly SIA to near-zero (Fig. 1a and 1b, red lines; Figure B1). For some months the trend since 1979 is now weakly negative, and trends are insignificant in all months (Figure B1). Meanwhile, adding the further decade of data hardly changed the multi-model mean trend at all (Fig. 1a and 1b, blue lines). The mean trend remains strongly negative, although a few simulations have weakly positive trends. The simulated trends are less influenced by internal climate variability as more years are added, and therefore the standard deviation of the modelled trends for a fixed start year of 1979 decreases over time.

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

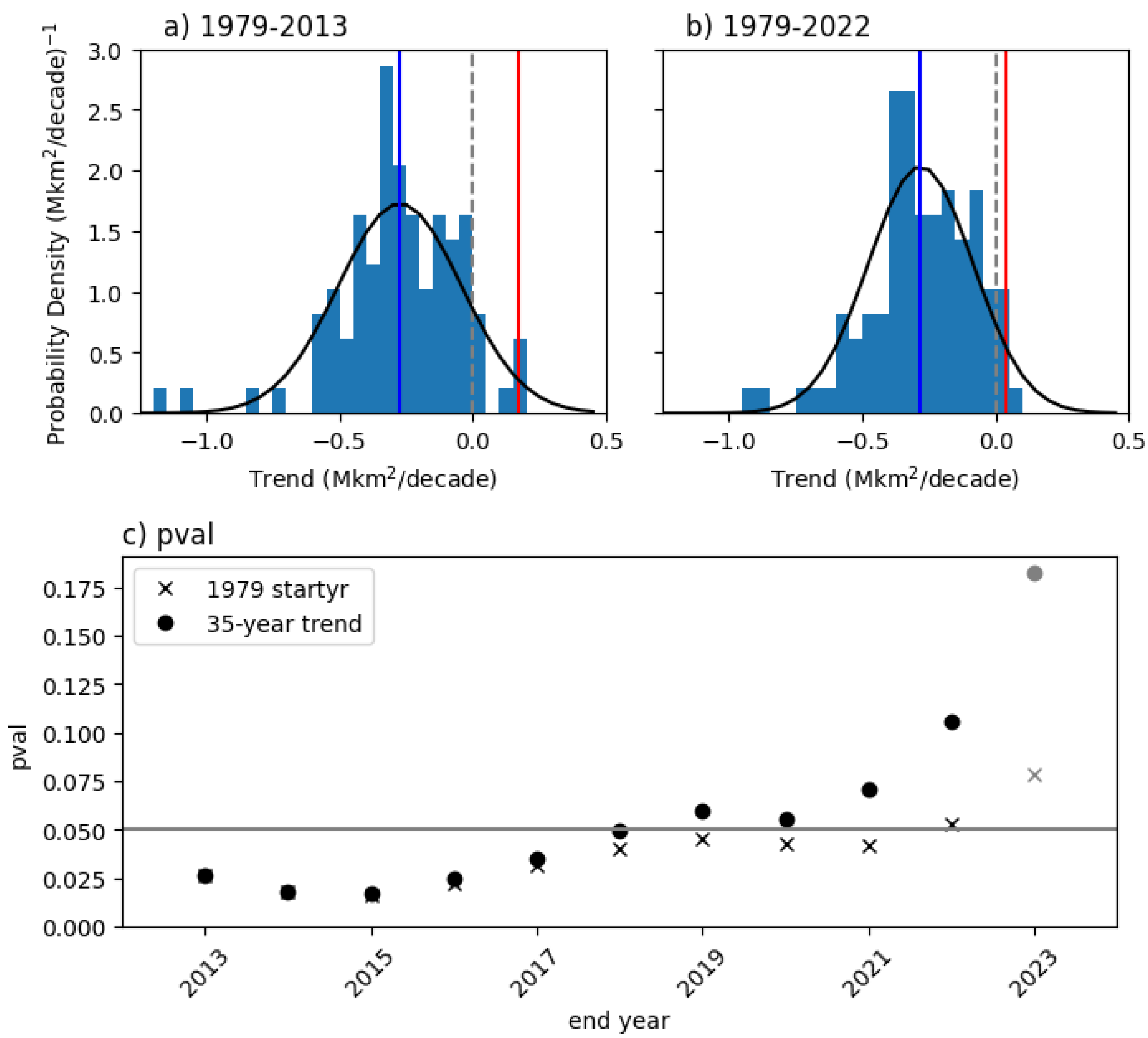

Figure 1 (a-b) Linear trends in annual mean SIA in satellite observations (red) and CMIP6 models (blue histogram) and Gaussian fit to CMIP6 distribution (black) for the periods (a) 1979-2013 and (b) 1979-2022. The dashed vertical line indicates zero trend and the blue line indicates the multi-model mean. (c) the probability of observing a trend at least as large as observed (a one-tailed test) under the null hypothesis that observations are taken from the same population as the CMIP6 multi-model ensemble, for varying end dates and either a fixed start date of 1979 as in panels a) and b) (crosses) or fixed trend length of 35 years (dots). The 2023 estimate (grey) is based on a synthetic extension of the observed time series (Section 2.3).

In light of these findings, we test the null hypothesis that observed sea ice trends are consistent with trends simulated across the CMIP6 multi-model ensemble, and consider how additional years of data affect the outcome of this test. We consider trends calculated with both a fixed start date (1979) and fixed duration (35 years) to aid our interpretation. Until 2018, the probability of a CMIP6 model trend matching or exceeding the observed trend is very low ($p<0.05$; Fig. 1c), so the null hypothesis is rejected. This can be interpreted as saying that the models' anthropogenic forced trend, as estimated by the

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

multi-model mean trend, is inconsistent with the observed trend, when allowing for modelled internal variability. Indeed, sensitivity experiments in climate models have also implied that the forced trend in models is too strong (Schneider and Deser, 2018). This test provides a clear result; the short time period of under forty years should allow for a generous range of modelled trends, but this range still fails to accommodate the observations.

From 2015, the probability of CMIP6 trends matching or exceeding the observed trend starts to increase, as the ice loss brings observations into line with the models (Figure 1c). However, if trends are calculated with a fixed 1979 start date, progressively lengthening the trend under consideration decreases the modelled trend standard deviation while hardly affecting the model mean trend (Figure C1). This makes it less likely the observed trends will fall within the distribution of modelled trends. Only in 2022 does the recent rapid decline in observations counteract this effect and finally bring observed trends into line with the models (null hypothesis not rejected at p=0.05; Figure 1c). In contrast, for 'fixed duration' trends, the standard deviation of modelled trends remains large, while the observed trend more rapidly declines and becomes negative due to the neglect of early low SIA years (Gagné et al., 2015; Schroeter et al., 2023) in addition to the inclusion of the recent low SIA years. Therefore, the null hypothesis is no longer rejected at p=0.05 as early as 2019 under this method.

4 Discussion and Conclusions

The results above show that, firstly, there is a mismatch between observed and modelled trends for the earlier period up to around 2018, as previously shown, suggesting that modelled anthropogenic trends are too strong relative to modelled variability during that period. However, secondly, our results show that the persistent low Antarctic SIA of 2022 and 2023 brings observed trends into line with the ensemble of modelled trends, permitting the interpretation that modelled forced trends and variability are realistic on 45-year timescales (the full length of the modern satellite record). This is an important conclusion, since these longer timescales are of greatest relevance to centennial projections of climate change, and to the attribution of anthropogenically forced change. Moreover, even trends on the shorter 35-year timescale fall within the model ensemble for the five most recent 35-year periods (Fig. 1c).

The fact that these assessments of model skill changed after the addition of recent years could be explained in several ways. Conceptually, for any time period there is a distribution of model trends and also a distribution of possible real trends that could have occurred (depending upon the evolution of internal climate variability). Each observed trend is a single realisation of the distribution of possible real trends. The observed trends with earlier end dates were outside the model trend distribution. Now, the latest observed trends fall within with the distribution of modelled trends. In other words, the observed trends previously lay within the region where the modelled and real trend distributions did not overlap, and now lie within the region where they do.

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

The modelled and real trend distributions will differ in their spread if the models have inaccurate variability, and in their mean if the models have an inaccurate anthropogenic forced trend. Therefore, inaccurate variability, particularly on multidecadal timescales, could explain the changing assessment of skill. Alternatively, it could be that the modelled anthropogenic trends are too strong (Schneider and Deser, 2018), or emerge too early. For example, this is consistent with the models under-estimating the 'two-timescale' response to stratospheric ozone forcing whereby increasing westerlies cause a sea ice increase on 'short' timescales and decline on 'long' timescales (Ferreira et al., 2015; Kostov et al., 2017). The latest observed ice decline may be finally revealing the influence of anthropogenic forcing, bringing the observations closer to the models, which have long predicted this decline.

The importance of our results is in showing that we can no longer rule out climate model simulations of Antarctic sea ice based on trends alone. There are many measures by which modelled sea ice may be assessed, including seasonal and interannual variability (Zunz et al., 2013), spatial patterns, and physical processes (Holmes et al., 2019), but models no longer fail the fundamental test of being able to simulate trends, at least over the 45-year modern satellite era. Future studies can therefore move on to more detailed model assessments, including representation of the recent rapid decline itself. The rapid declines are as yet short-lived, so an improved understanding of multi-decadal sea ice variability and its representation in climate models is critical for further interpreting these results. Further, processes lacking from models, such as increasing freshwater input from accelerating ice sheet melt (Swart et al., 2023), may provide further complications in the relative evolution of modelled and observed sea ice over the 21st century.

Our results have broad ramifications for future assessments of CMIP6 outputs. First, revising our confidence in the climate models has consequences for the attribution of historical climate changes. Secondly, we should now have some level of greater confidence in the strong projected declines in Antarctic sea ice under anthropogenic forcing (Roach et al., 2020), whereby ice becomes near-absent in summer (Holmes et al., 2022). This in turn will influence our understanding of the future evolution of all aspects of the Southern Hemisphere climate - including Southern Ocean heat and carbon uptake, circumpolar winds (Bracegirdle et al., 2018), and melting of the Antarctic Ice Sheet – and of marine ecosystem function; all of which underpins decisions about the mitigation of future greenhouse gas emissions and about ecosystem management.

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Appendix A: CMIP6 models

MODEL	1979-2013 MEAN TREND	1979-2023 MEAN TREND	1989-2023 MEAN TREND	N MEMBERS	N MEMBERS USED
ACCESS_CM2	-0.04944	-0.17256	-0.26519	1	1
ACCESS_ESM1	-0.15064	-0.09875	-0.10404	3	3
AWICM1	-0.40479	-0.47263	-0.41977	1	1
BCC_CSM2	0.19367	-0.44325	-0.80301	1	1
CAMS_CSM1	-0.06683	-0.09611	-0.22959	2	2
CESM2	-0.36938	-0.38158	-0.38781	3	3
CESM2_WACCM	-0.47352	-0.44688	-0.44569	3	3
CIESM	-0.25127	-0.26142	-0.26097	1	1
CMCC_CM2_SR5	-0.35578	-0.33006	-0.3281	1	1
CMCC_ESM2	-0.29702	-0.24735	-0.25416	1	1
CNRM_CM6	-0.36236	-0.37923	-0.37553	6	6
CNRM_CM6_1_HR	-0.44265	-0.58323	-0.95026	1	1
CANESM5	-0.38605	-0.35623	-0.37311	19	6
E3SM_1_1	-0.32336	-0.35954	-0.42247	1	1
ECEARTH3	-0.26721	-0.22236	-0.23598	57	6
ECEARTH3_CC	-0.23065	-0.12555	-0.14672	1	1
ECEARTH3_VEG	-0.14929	-0.19563	-0.27552	5	5
ECEARTH3_VEG_LR	-0.32488	-0.27959	-0.29294	1	1
FGOALS_F3L	-0.12234	-0.15916	-0.10923	1	1
FGOALS_G3	-0.27949	-0.2264	-0.13548	4	4
FIO_ESM	-0.31626	-0.34199	-0.33857	3	3
GFDL_CM4	-0.2228	-0.19287	-0.15889	1	1
GFDL_ESM4	-0.0388	-0.11107	-0.07462	1	1
GISS_E2_1_G	-0.13518	0.007827	0.062373	1	1
HADGEM3_GC31_LL	-0.5135	-0.60736	-0.67449	3	3
HADGEM3_GC31_MM	-0.31207	-0.31269	-0.36247	4	4
INM_CM4_8	-0.19266	-0.21018	-0.22783	1	1
INM_CM5_0	-0.23805	-0.23203	-0.2004	1	1

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

IPSL_CM6A_LR	-0.36316	-0.38448	-0.41388	6	6
KIOST_ESM	-0.25891	-0.2145	-0.15571	1	1
MIROC6	-0.01415	-0.01507	-0.00588	3	3
MIROC_ES2L	-0.07167	-0.08444	-0.10766	8	6
MPI_ESM1_2_HR	-0.27659	-0.27352	-0.35618	2	2
MPI_ESM1_2_LR	-0.10824	-0.07766	0.018626	30	6
MRI_ESM2	-0.3255	-0.3771	-0.43644	1	1
NESM3	-0.20193	-0.28303	-0.37379	2	2
NORESM2_LM	-0.09616	-0.08174	-0.10233	1	1
NORESM2_MM	-0.0136	-0.0771	-0.04084	1	1
UKESM1_0_LL	-0.72078	-0.66572	-0.65195	5	5

Table A1: The models available for the study, the number of ensemble members available and the number used, and the ensemble mean trend (across the ensemble members used only) for the period specified. NorESM values were calculated by the authors from SIC data; all other values were obtained from the CMIP6 SIA Directory made available by the University of Hamburg and methods are fully detailed there.

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Appendix B: Monthly trends

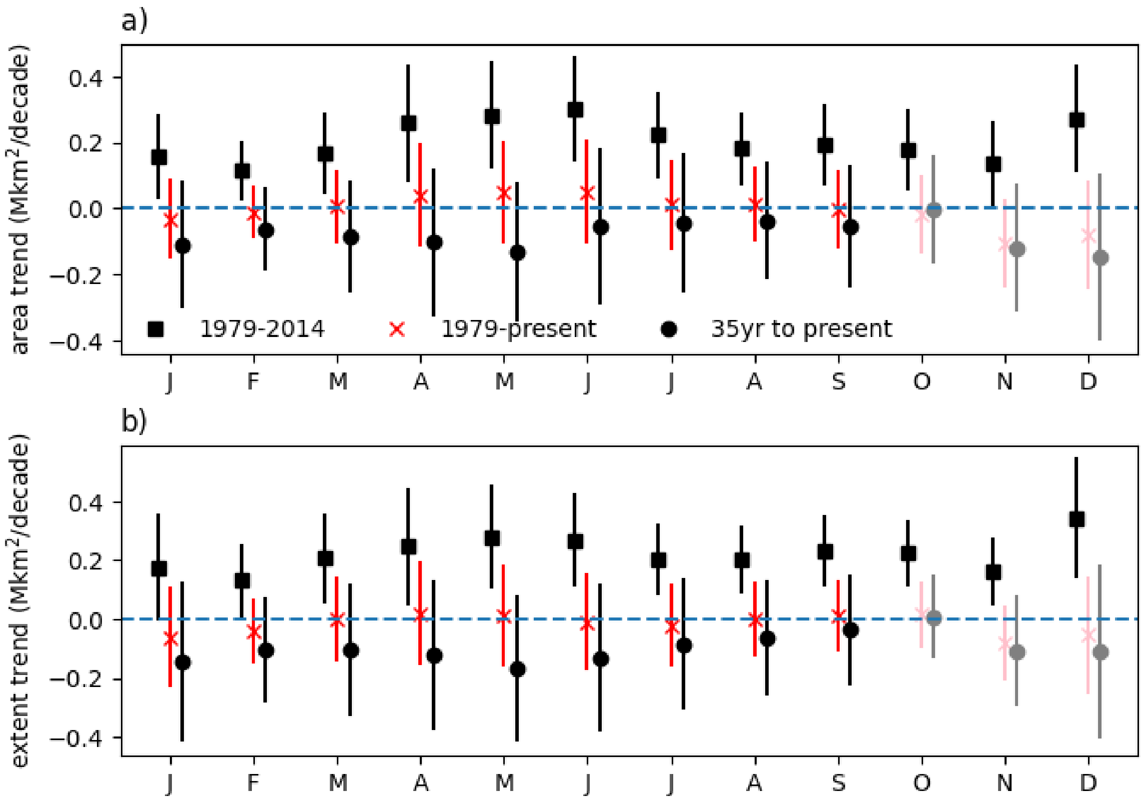

Table B2: Observed sea ice trends in individual months for (squares) 1979-2014, (crosses) full 45-year trend 1979-present, and (circles) 35-year trend to present. 'Present' indicates 2023 for January to September and 2022 for October to December; trends ending in 2022 are in a lighter colour for distinction. 1979-present trends are highlighted in shades of red as this period is the focus of the paper. a) Sea Ice Area, b) Sea Ice Extent. 5th-95th percentile uncertainties are indicated by vertical lines. Data are from the Sea Ice Index (see Methods).

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Appendix C: Sensitivity Tests

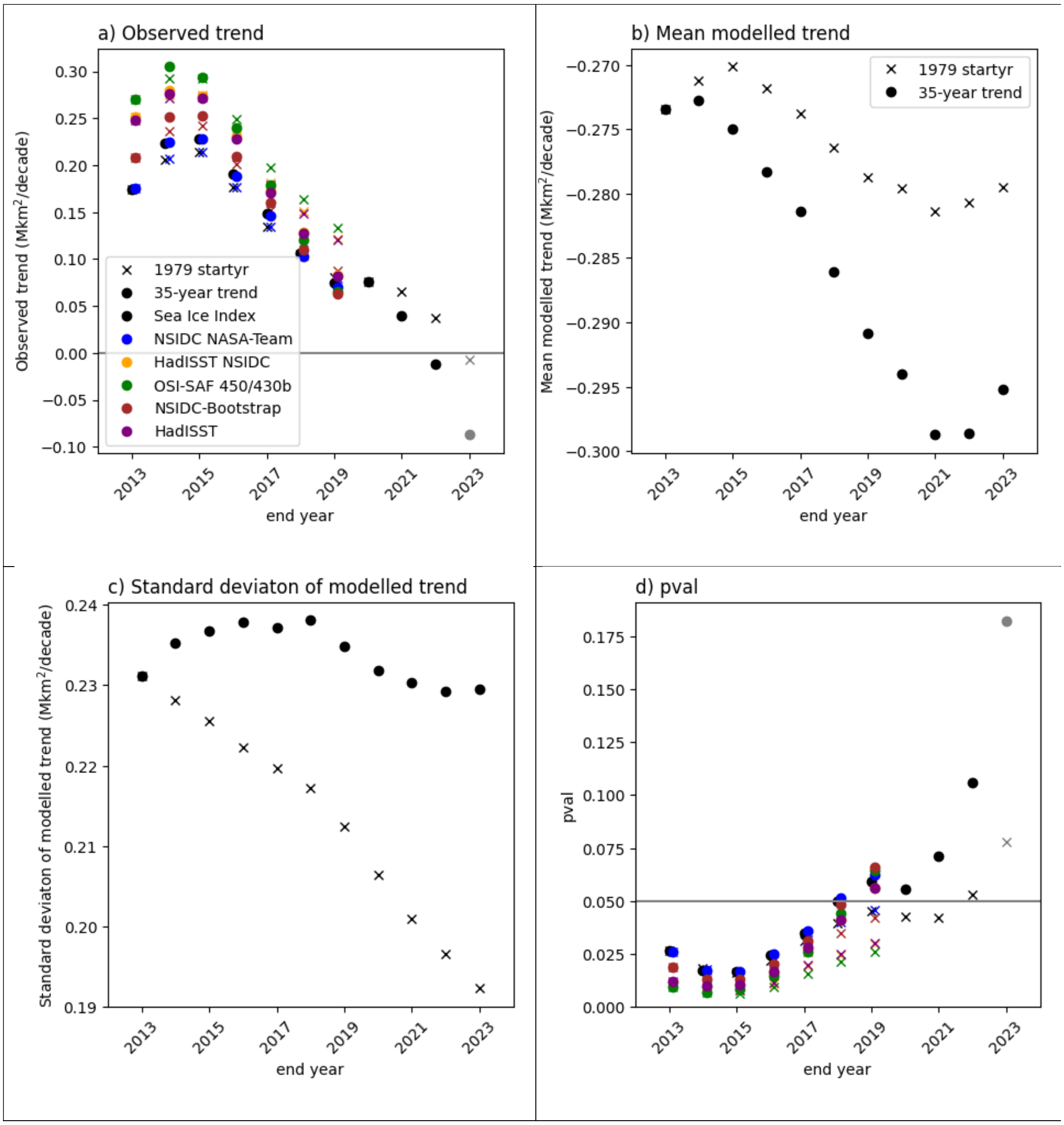

Figure C1: Contributions to the p-value shown in Figure 1c). a) observed trend; Sea Ice Index in black as in main text, other datasets as indicated. b) mean of modelled trends, c) standard deviation of modelled trends, d) p-value (as main text Figure 1 but with alternative observational estimates (Dörr, 2021)).

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Sensitivity to Observational Dataset

Observational uncertainty in SIA is particularly high prior to winter 1987 (not shown) due to missing SIC data. Trends in the other datasets, in particular OSI-SAF (Figure C1, green), are in general more strongly positive than those in the Sea Ice Index (Figure C1a). Therefore, for the '1979 start date' trends, these might exhibit consistency with model-simulated trends at later end dates than 2022 (Figure C1d, crosses); note that all datasets already display consistency for the 35-year trends ending in 2019 onwards (Figure C1d, dots).

Sensitivity to treatment of CMIP6 models

We also tested the sensitivity of our conclusions to our treatment of CMIP6 models. First, by including all ensemble members (instead of a maximum of six per model), we largely add simulations from models with weak negative average trends (Table A1) and so increase consistency with observations (not shown). However, the evolution with end year of the model-observation comparison (Fig. 1c) and the broad timings of threshold crossings are unchanged. Second, we tested sensitivity to using the weaker forcing scenario ssp245 instead of ssp585 for the extension of modelled trends after 2014. The effect of forcing scenario is small early in the 21st century (Hawkins and Sutton, 2012), so that any difference arising is due to internal variability or structural differences between the models with simulations available. For the overlapping subset of 147 model-realisation combinations, ssp245 has marginally stronger trends and so is slightly less consistent with observations. In contrast, using the full ssp245 ensemble (with all available members) means including a larger ensemble of MIROC6 than in the overlapping subset or in the ssp585 ensemble; MIROC6 implausibly has virtually no sea ice year-round (Shu et al., 2020) and therefore zero trends (Holmes et al., 2022) leading to weaker mean trends and slightly greater consistency with observations. In summary, these effects are small, and so our conclusions are robust to these sensitivity tests.

Code Availability

The code for calculating trends, performing the evaluation and preparing figures is available from the corresponding author on request.

Data Availability

Sea Ice Area from the CMIP6 models is available from the University of Hamburg (UHH) CMIP6 Sea Ice Area directory (https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/cmip6-sea-ice-area.html, accessed 2023-08-17). The NSIDC Sea Ice Index v3.0 SIA (Fetterer, 2017) is available from https://nsidc.org/arcticseaicenews/sea-ice-tools/. Other observational estimates of sea ice area (Dörr, 2021) are available from https://doi.org/10.25592/uhhfdm.8559.

Author Contributions

CRH, TJB and PRH conceived the study. CRH conducted the analysis and prepared the figures. All authors discussed the results and reviewed the manuscript.

https://doi.org/10.5194/egusphere-2023-2881
Preprint. Discussion started: 18 December 2023

Competing Interests

The authors declare they have no conflicts of interest.

Acknowledgements

All authors received funding from NERC grant DEFIANT (NE/W004739/1); JS also received funding from Canada 150 Research Chairs program (C150 grant no. 50296). The World Climate Research Programme's (WCRP) Working Group on Coupled Modelling, which is responsible for CMIP, and the climate modelling groups, are thanked for producing and making available their model output.

References

Bracegirdle, T. J., Hyder, P., and Holmes, C. R.: CMIP5 diversity in southern westerly jet projections related to historical sea ice area: Strong link to strengthening and weak link to shift, J. Climate, 31, 195-211, 2018.
Bracegirdle, T. J., Stephenson, D. B., Turner, J., and Phillips, T.: The importance of sea ice area biases in 21st century multimodel projections of Antarctic temperature and precipitation, Geophys. Res. Lett., 42, 10,832-810,839, 2015.
Dörr, J. N., Dirk; Kern, Stefan: UHH sea-ice area product, 1850-2019 (v2019_fv0.01) [dataset], https://doi.org/10.25592/uhhfdm.8559, , 2021.
Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., and Taylor, K. E.: Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization, Geosci. Model Dev., 9, 1937-1958, 2016.
Ferreira, D., Marshall, J., Bitz, C. M., Solomon, S., and Plumb, A.: Antarctic Ocean and sea ice response to ozone depletion: A two-time-scale problem, J. Climate, 28, 1206-1226, 2015.
Fetterer, F., K. Knowles, W. N. Meier, M. Savoie, and A. K. Windnagel.: Sea Ice Index, Version 3 Boulder, Colorado USA. National Snow and Ice Data Center. [dataset], https://doi.org/10.7265/N5K072F8, 2017.
Fox-Kemper, B., H.T. Hewitt, C. Xiao, G. Aðalgeirsdóttir, S.S. Drijfhout, T.L. Edwards, N.R. Golledge, M. Hemer, R.E. Kopp, G. Krinner, A. Mix, D. Notz, S. Nowicki, I.S. Nurhati, L. Ruiz, J.-B. Sallée, A.B.A. Slangen Y. Yu: Ocean, Cryosphere and Sea Level Change, 1211-1362, 10.1017/9781009157896.011., 2021.
Gagné, M. È., Gillett, N., and Fyfe, J.: Observed and simulated changes in Antarctic sea ice extent over the past 50 years, Geophys. Res. Lett., 42, 90-95, 2015.
Hawkins, E. and Sutton, R.: Time of emergence of climate signals, Geophys. Res. Lett., 39, 2012.
Holmes, C., Bracegirdle, T., and Holland, P.: Antarctic sea ice projections constrained by historical ice cover and future global temperature change, Geophys. Res. Lett., 49, e2021GL097413, 2022.
Holmes, C. R., Holland, P. R., and Bracegirdle, T. J.: Compensating biases and a noteworthy success in the CMIP5 representation of Antarctic sea ice processes, Geophys. Res. Lett., 46, 4299-4307, 2019.
Kostov, Y., Marshall, J., Hausmann, U., Armour, K. C., Ferreira, D., and Holland, M. M.: Fast and slow responses of Southern Ocean sea surface temperature to SAM in coupled climate models, Clim. Dynam., 48, 1595-1609, 2017.
National Academies of Sciences, E. and Medicine: Antarctic sea ice variability in the southern ocean-climate system: Proceedings of a workshop, 2017.
Notz, D.: Sea-ice extent and its trend provide limited metrics of model performance, Cryosphere, 8, 229-243, 2014.
O'Neill, B. C., Tebaldi, C., Van Vuuren, D. P., Eyring, V., Friedlingstein, P., Hurtt, G., Knutti, R., Kriegler, E., Lamarque, J.-F., and Lowe, J.: The scenario model intercomparison project (ScenarioMIP) for CMIP6, Geosci. Model Dev., 9, 3461-3482, 2016.
Parkinson, C. L.: A 40-y record reveals gradual Antarctic sea ice increases followed by decreases at rates far exceeding the rates seen in the Arctic, Proceedings of the National Academy of Sciences, 116, 14414-14423, 2019.
Purich, A. and Doddridge, E. W.: Record low Antarctic sea ice coverage indicates a new sea ice state, Communications Earth & Environment, 4, 314, 2023.

Roach, L. A., Dörr, J., Holmes, C. R., Massonnet, F., Blockley, E. W., Notz, D., Rackow, T., Raphael, M. N., O'Farrell, S. P., Bailey, D. A., and Bitz, C. M.: Antarctic Sea Ice Area in CMIP6, Geophys. Res. Lett., 47, e2019GL086729, https://doi.org/10.1029/2019GL086729, 2020.

Rosenblum, E. and Eisenman, I.: Sea ice trends in climate models only accurate in runs with biased global warming, J. Climate, 30, 6265-6278, 2017.

Schlosser, E., Haumann, F. A., and Raphael, M. N.: Atmospheric influences on the anomalous 2016 Antarctic sea ice decay, Cryosphere, 12, 1103-1119, 2018.

Schneider, D. P. and Deser, C.: Tropically driven and externally forced patterns of Antarctic sea ice change: Reconciling observed and modeled trends, Clim. Dynam., 50, 4599-4618, 2018.

Schroeter, S., O'Kane, T. J., and Sandery, P. A.: Antarctic sea ice regime shift associated with decreasing zonal symmetry in the Southern Annular Mode, Cryosphere, 17, 701-717, 2023.

Shu, Q., Wang, Q., Song, Z., Qiao, F., Zhao, J., Chu, M., and Li, X.: Assessment of sea ice extent in CMIP6 with comparison to observations and CMIP5, Geophys. Res. Lett., 47, e2020GL087965, 2020.

Siegert, M. J., Bentley, M. J., Atkinson, A., Bracegirdle, T. J., Convey, P., Davies, B., Downie, R., Hogg, A. E., Holmes, C., and Hughes, K. A.: Antarctic extreme events, Frontiers in Environmental Science, 11, 1229283, 2023.

Swart, N., Martin, T., Beadling, R., Chen, J.-J., England, M. H., Farneti, R., Griffies, S. M., Hatterman, T., Haumann, F. A., and Li, Q.: The Southern Ocean Freshwater release model experiments Initiative (SOFIA): Scientific objectives and experimental design, EGUsphere, 2023, 1-30, 2023.

Turner, J. and Comiso, J.: Solve Antarctica's sea-ice puzzle, Nature, 547, 275-277, 2017.

Turner, J., Phillips, T., Marshall, G. J., Hosking, J. S., Pope, J. O., Bracegirdle, T. J., and Deb, P.: Unprecedented springtime retreat of Antarctic sea ice in 2016, Geophys. Res. Lett., 44, 6868-6875, 2017.

Zhang, L., Delworth, T. L., Yang, X., Zeng, F., Lu, F., Morioka, Y., and Bushuk, M.: The relative role of the subsurface Southern Ocean in driving negative Antarctic Sea ice extent anomalies in 2016–2021, Communications Earth & Environment, 3, 302, 2022.

Zunz, V., Goosse, H., and Massonnet, F.: How does internal variability influence the ability of CMIP5 models to reproduce the recent trend in Southern Ocean sea ice extent?, Cryosphere, 7, 451-468, 2013.