



**Errors associated with calculating the gross nitrification rates in forested
catchments using the triple oxygen isotopic composition ($\Delta^{17}\text{O}$) of stream nitrate**

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1 **Abstract**

2 A novel method for quantifying the gross nitrification rate (GNR) in each forested
3 catchment using the triple oxygen isotopic composition ($\Delta^{17}\text{O}$) of stream nitrate eluted
4 from the catchment has been proposed and applied in several recent studies. However,
5 the equations used in the calculations include the approximation that the $\Delta^{17}\text{O}$ value
6 of nitrate metabolized through either assimilation or denitrification within the forested
7 soil is equal to the $\Delta^{17}\text{O}$ value of nitrate in the stream. The GNR estimated from the
8 $\Delta^{17}\text{O}$ value of stream nitrate was more than six times the actual GNR in our simulated
9 calculation for a forested catchment where the nitrate in the soil exhibited $\Delta^{17}\text{O}$ values
10 larger than those in the stream while showing a decreasing trend with increasing
11 depths until that of stream nitrate at the bottom. As most of the reported soil nitrate in
12 forested catchments from past studies showed $\Delta^{17}\text{O}$ values higher than those of the
13 stream nitrate eluted from each catchment, we concluded that the GNR estimated
14 from the $\Delta^{17}\text{O}$ value of stream nitrate in the forested catchments was, to some extent,
15 an overestimate of the actual GNR.

16

17 **1 Introduction**

18 Nitrate (NO_3^-) is one of the important nitrogen nutrients for primary production in
19 forested ecosystems. Nitrification is the microbial process that produces NO_3^- in each
20 forested ecosystem. Thus, quantifying the nitrification rate can assist in the evaluation
21 of the present and future states of each forest ecosystem. While the net nitrification
22 rate can be estimated from the increase in NO_3^- concentration during a certain period,



23 the gross nitrification rate (GNR), which includes the net nitrification rate and the
24 metabolic rate of nitrate (e.g., assimilated by plants or decomposed through
25 denitrification), reflects the internal N cycling better than the net nitrification rate
26 (Bengtsson et al., 2003), especially in forested ecosystems, where the net nitrification
27 rate is negligible (Stark and Hart, 1997) while the metabolic rate is significant so that
28 the GNR often exceeds the net nitrification rate by order of magnitude (Verchot et al.,
29 2001).

30 Recently, several studies have successfully estimated GNR in water environments
31 such as lakes, using the $\Delta^{17}\text{O}$ values of NO_3^- as a conserved tracer of the mixing ratio
32 between the atmospheric nitrate ($\text{NO}_3^-_{\text{atm}}$) deposited into the water environment and
33 the remineralized nitrate ($\text{NO}_3^-_{\text{re}}$) produced through nitrification therein (Tsunogai et
34 al., 2011, 2018). Although $\text{NO}_3^-_{\text{re}}$ always shows the $\Delta^{17}\text{O}$ values close to 0 ‰ because
35 its oxygen atoms derive from either terrestrial O_2 or H_2O through nitrification,
36 $\text{NO}_3^-_{\text{atm}}$ always displays an anomalous enrichment in ^{17}O with $\Delta^{17}\text{O}$ values being
37 approximately $+26 \pm 3$ ‰ in Japan (Tsunogai et al., 2010, 2016) because of oxygen
38 transfers from atmospheric ozone (Michalski et al., 2003; Nelson et al., 2018).
39 Additionally, $\Delta^{17}\text{O}$ is almost stable during “mass-dependent” isotope fractionation
40 processes (Michalski et al., 2004; Tsunogai et al., 2016). Therefore, regardless of the
41 partial metabolism through denitrification or assimilation after deposition in a water
42 column, $\Delta^{17}\text{O}$ can be used as a conserved tracer of $\text{NO}_3^-_{\text{atm}}$ to calculate the mixing



43 ratio of $\text{NO}_3^-_{\text{atm}}$ to total NO_3^- ($\text{NO}_3^-_{\text{atm}}/\text{NO}_3^-_{\text{total}}$) in a water column using the
44 following equation:

$$45 \quad [\text{NO}_3^-_{\text{atm}}]/[\text{NO}_3^-_{\text{total}}] = [\text{NO}_3^-_{\text{atm}}]/([\text{NO}_3^-_{\text{re}}] + [\text{NO}_3^-_{\text{atm}}]) = \Delta^{17}\text{O}/\Delta^{17}\text{O}_{\text{atm}} \quad (1)$$

46 where $\Delta^{17}\text{O}_{\text{atm}}$ and $\Delta^{17}\text{O}$ denote the $\Delta^{17}\text{O}$ values of $\text{NO}_3^-_{\text{atm}}$ and NO_3^- dissolved in
47 each water environment, respectively. Using both the $\text{NO}_3^-_{\text{atm}}/\text{NO}_3^-_{\text{total}}$ ratio estimated
48 from the $\Delta^{17}\text{O}$ value of NO_3^- in a lake water column and the deposition rate of
49 $\text{NO}_3^-_{\text{atm}}$ into the lake, the GNR has been successfully estimated (Tsunogai et al., 2011,
50 2018).

51 In addition to application in water environments, the $\Delta^{17}\text{O}$ method has also been
52 applied to forested catchments for GNR determination (Fang et al., 2015; Hattori et
53 al., 2019; Huang et al., 2020). By using the deposition flux of $\text{NO}_3^-_{\text{atm}}$ into the
54 catchment as well as the elution flux of both unprocessed $\text{NO}_3^-_{\text{atm}}$ and $\text{NO}_3^-_{\text{re}}$ from the
55 stream, which can be determined from the $\Delta^{17}\text{O}$ values of NO_3^- in stream water eluted
56 from the catchment, the GNR in each forested catchment has been estimated in a
57 manner similar to the estimation for the water environments (Fang et al., 2015).
58 Contrary to water environments, where the $\Delta^{17}\text{O}$ values of NO_3^- within the water
59 layers are homogeneous and can be measured easily, it is often difficult to determine
60 the $\Delta^{17}\text{O}$ values of the NO_3^- metabolized in soil layers. Consequently, past studies
61 approximated the values to be equal to the $\Delta^{17}\text{O}$ value of stream NO_3^- eluted from
62 each forested catchment without actual observation (Fang et al., 2015, Hattori et al.,
63 2019, Huang et al., 2020). However, such an approximation should be conducted with



64 extreme caution, as the $\Delta^{17}\text{O}$ values of soil NO_3^- are not always equal to those of the
65 stream (Hattori et al., 2019, Rose, 2014, Osaka et al., 2010). To clarify the details of
66 the approximation along with its impact on the final estimated GNR, we present an
67 accurate relationship between the $\Delta^{17}\text{O}$ of soil NO_3^- and GNR, starting from the basic
68 isotope mass balance equations. Then, we present the GNR estimated for a forested
69 catchment in which the $\Delta^{17}\text{O}$ values of NO_3^- in soil are measured. Finally, we
70 compare the GNR estimated in this study with the GNR estimated from the $\Delta^{17}\text{O}$
71 values of stream NO_3^- .

72

73 **2 Calculation**

74 The total mass balance equation of NO_3^- including the GNR in each catchment can
75 be expressed as follows:

$$76 \text{NO}_3^-_{\text{deposition}} + \text{GNR} = \text{NO}_3^-_{\text{leaching}} + \text{NO}_3^-_{\text{uptake}} + \text{GDR} \quad (2)$$

77 where $\text{NO}_3^-_{\text{deposition}}$, GNR, $\text{NO}_3^-_{\text{leaching}}$, $\text{NO}_3^-_{\text{uptake}}$, and GDR denote the deposition flux
78 of NO_3^- into each catchment, gross nitrification rate in each catchment, leaching flux
79 of NO_3^- from each catchment, uptake rate of NO_3^- in each catchment, and gross
80 denitrification rate in each catchment, respectively.

81 The isotope mass balance for each $\Delta^{17}\text{O}$ value of NO_3^- in the catchment can also be
82 calculated using the same method:

$$83 \text{NO}_3^-_{\text{deposition}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} + \text{GNR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}} = \text{NO}_3^-_{\text{leaching}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}} + \text{NO}_3^-_{\text{uptake}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}} + \text{GDR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}} \quad (3)$$



85 where $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}}$, and
86 $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}}$ denote the $\Delta^{17}\text{O}$ value of NO_3^- atm deposited into each
87 catchment, that of NO_3^- re produced through nitrification, that of NO_3^- eluted from
88 each catchment, that of NO_3^- assimilated by plants and other organisms in each
89 catchment, and that of NO_3^- decomposed through denitrification in each catchment,
90 respectively.

91 If the $\Delta^{17}\text{O}$ values of NO_3^- in the forested soil layers, where the NO_3^- was
92 metabolized through either assimilation (by plants and other organisms) or
93 denitrification, are equal to the $\Delta^{17}\text{O}$ value of NO_3^- in the stream, Eq. 4 can be
94 expressed as follows:

$$95 \quad \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}} = \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}} = \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}} \quad (4)$$

96 Consequently, by combining Eqs. 3 and 4, we could obtain the following
97 relationship:

$$98 \quad \text{NO}_3^-_{\text{deposition}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} + \text{GNR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}} = (\text{NO}_3^-_{\text{leaching}} + \text{NO}_3^-_{\text{uptak}} \\ 99 \quad \text{e} + \text{GDR}) \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}} \quad (5)$$

100 We could estimate the GNR using Eq. 6 obtained from Eqs. 2 and 5 because we can
101 approximate the $\Delta^{17}\text{O}$ values of NO_3^- re produced through nitrification
102 ($\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}}$) to be 0 (Michalski et al., 2003; Tsunogai et al., 2010):

$$103 \quad \text{GNR} = \text{NO}_3^-_{\text{deposition}} \times (\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} - \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}) / \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}} \quad (6)$$



104 Eq. 6 corresponds to the equation used in previous studies for quantifying the GNR
105 in each forested catchment (Eq. 4 in Fang et al., 2015; Eq. 8 in Hattori et al., 2019;
106 Eq. 4 Huang et al., 2020).

107

108 **3 Results and Discussion**

109 The $\Delta^{17}\text{O}$ values of NO_3^- in forested soil layers should be equal to those of NO_3^- in
110 the stream, as presented in Eq. 4 to obtain Eq. 6. While the number of simultaneous
111 observations of the oxygen isotopes of NO_3^- in both the soil and stream in a given
112 forested catchment is limited (Hattori et al., 2019, Osaka et al., 2010, Rose, 2014), the
113 limited observations show that the oxygen isotopic ratios of soil NO_3^- are mostly
114 higher than those of stream NO_3^- . For example, Hattori et al. (2019) reported that
115 more than 60 % of the soil exhibited $\Delta^{17}\text{O}$ values significantly higher than those of
116 stream NO_3^- determined simultaneously ($\Delta^{17}\text{O} = +1$ to $+3$ ‰). In addition, they found
117 a decreasing $\Delta^{17}\text{O}$ trend in soil NO_3^- with depth, declining from greater than $+20$ ‰ at
118 the surface to less than $+3$ ‰ at depths of 25–90 cm from the surface. A similar
119 decreasing trend in the vertical distribution had been found in $\delta^{18}\text{O}$ in another forested
120 catchment, from greater than $+35$ ‰ at the surface soil to less than $+10$ ‰ at depths of
121 30–50 cm from the soil surface (Osaka et al., 2010). Besides, most of the soil NO_3^-
122 also exhibited $\delta^{18}\text{O}$ values higher than those in the stream (Osaka et al., 2010).
123 Furthermore, Rose (2014) monitored the horizontal distribution of $\Delta^{17}\text{O}$ of soil NO_3^-
124 by randomly setting 15 tension-free lysimeters at depths of 0–10 cm in a 39 ha



125 forested catchment, reporting $\Delta^{17}\text{O}$ values significantly higher in soil NO_3^- ($+9.1 \pm$
126 5.8 ‰ on average) than in the stream NO_3^- ($+0.5 \text{ ‰}$ on average) eluted from the
127 forested catchment. As most of the root biomass is concentrated in the top 10 cm of
128 the soil in forested catchments (Jackson et al., 1996), most uptake reactions should
129 occur in that top 10 cm of soil. Consequently, the significant difference in the $\Delta^{17}\text{O}$
130 values between soil NO_3^- and stream NO_3^- , particularly in the surface soil layers,
131 imply that the estimated GNRs in the forested catchment obtained from Eq. 6 were
132 inaccurate.

133 To demonstrate the impact of the differences in $\Delta^{17}\text{O}$ between soil NO_3^- and stream
134 NO_3^- on the GNR, along with presenting the problems associated with the
135 approximation to obtain Eq. 6, we estimated the GNR for two simulated forested
136 soils—one with NO_3^- showing a decreasing trend in $\Delta^{17}\text{O}$ down to the $\Delta^{17}\text{O}$ of stream
137 NO_3^- (heterogeneous soil) (Fig. 1a and 1b) and one with NO_3^- showing the same
138 $\Delta^{17}\text{O}$ values as those of stream NO_3^- (homogeneous soil) (Fig. 2a and 2b). With
139 Hattori et al. (2019) reporting the NO_3^- deposition as $7.0 \text{ kg of N ha}^{-1} \text{ y}^{-1}$, NO_3^- leaching as
140 $2.6 \text{ kg of N ha}^{-1} \text{ y}^{-1}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}}$ as $+28.0 \text{ ‰}$, and $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ as $+2.2 \text{ ‰}$ in
141 their forested catchment study, we adopted the same values in our calculation.

142 We divided the soils in the heterogeneous forest soils into 10 layers in the vertical
143 direction, simulating the soils observed by Hattori et al. (2019), in which the $\Delta^{17}\text{O}$
144 values of NO_3^- gradually decreased with increasing depth, varying from $+28.0$ to
145 $+2.2 \text{ ‰}$ with a rate of decrease of $+2.58 \text{ ‰}$ for each step (Fig. 1b). Similarly, we



146 assumed a gradual decrease with increasing depth in the leaching flux of NO_3^- , i.e.,
147 from 7 to 2.6 kg of N $\text{ha}^{-1} \text{y}^{-1}$ with a rate of decrease of 0.44 kg of N $\text{ha}^{-1} \text{y}^{-1}$ for each
148 step (Fig. 1c). In the homogeneous forest soils, we also divided the forested soils into
149 10 layers in the vertical direction. The change with depth in the leaching flux of NO_3^-
150 was the same as that in the heterogeneous soils (Fig. 2c), whereas the $\Delta^{17}\text{O}$ values of
151 NO_3^- were constant at +2.2 ‰ in the soil layers (Fig. 2b).

152 Applying the total mass balance and isotope mass balance of NO_3^- shown in Eqs. 2
153 and 3 to each layer, we estimated both the GNR (Figs. 1e and 2e) and total metabolic
154 rate of NO_3^- (GDR + uptake) (Figs. 1d and 2d) in each layer assuming the following:
155 (1) $\Delta^{17}\text{O}$ values of NO_3^- are constant in each layer, (2) vertical flow of NO_3^- in the
156 soil layers proceed downward from the surface to the water layer with a uniform
157 residence time in each layer, and (3) the GNR and metabolic rate of NO_3^- (GDR +
158 uptake) is zero in the water layer (layers beyond the no. 10 soil layer). Then, by
159 integrating the GNR determined for each layer, we estimated the total GNR in each
160 forested catchment.

161 Although the GNR estimated for the catchment with the homogeneous $\Delta^{17}\text{O}$ values
162 in soil NO_3^- was 83.6 kg of N $\text{ha}^{-1} \text{y}^{-1}$, exactly equal to that estimated by Hattori et al.
163 (2019) using Eq. 6 (Fig. 2e), the total GNR was a much smaller 13.0 kg of N $\text{ha}^{-1} \text{y}^{-1}$,
164 simulated for the catchment with the heterogeneous $\Delta^{17}\text{O}$ values in soil NO_3^- (Fig.
165 1e). Consequently, we conclude the following: (1) past studies estimating the GNR
166 using Eq. 6 approximated the $\Delta^{17}\text{O}$ value of soil NO_3^- was homogeneous and always



167 equal to that of stream NO_3^- mathematically and (2) the differences between the $\Delta^{17}\text{O}$
168 values of the soil NO_3^- metabolized in a forested catchment and that of stream NO_3^-
169 resulted in a significant deviation in the GNR estimated using Eq. 6 from the actual
170 GNR.

171 Note that the linear variation in the leaching flux and $\Delta^{17}\text{O}$ values of soil NO_3^- used
172 in the simulated calculations (Fig. 1) is just one of many possible variations in the
173 forested catchments. It is impossible to decide whether the linear variation was
174 realistic until the downward water flux, along with the concentration and $\Delta^{17}\text{O}$ values
175 of NO_3^- , is determined for each soil layer. However, the simultaneous observations of
176 the oxygen isotopes of soil NO_3^- and stream NO_3^- (Hattori et al., 2019; Osaka et al.,
177 2010; Nakagawa et al., 2018; Rose, 2014) imply that the approximation of the $\Delta^{17}\text{O}$
178 values of the soil NO_3^- metabolized through assimilation or denitrification to be
179 always equal to the $\Delta^{17}\text{O}$ value of stream NO_3^- , shown in Fig. 2b, is unrealistic.

180 By combining the mass balance and isotope mass balance shown in Eqs. 2 and 3,
181 Eq. 7 can be obtained to accurately estimate the GNR:

$$\begin{aligned} 182 \text{ GNR} = & \text{NO}_3^-_{\text{leaching}} - \text{NO}_3^-_{\text{deposition}} + (\text{NO}_3^-_{\text{deposition}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} - \\ 183 & \text{NO}_3^-_{\text{leaching}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}) / \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{soil}} \end{aligned} \quad (7)$$

184 where $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{soil}}$ denotes the $\Delta^{17}\text{O}$ values of NO_3^- in forested soil, from which
185 the NO_3^- was metabolized through either assimilation or denitrification. As most of
186 the soil NO_3^- measured to date exhibit $\Delta^{17}\text{O}$ values higher than those of the stream
187 NO_3^- eluted from each catchment (Hattori et al., 2019, Rose, 2014), the GNR



188 estimated from stream NO_3^- using Eq. 6 is higher than the GNR estimated from soil
189 NO_3^- using Eq. 7, to some extent. In other words, the GNR estimated from Eq. 6
190 overestimated the GNR in each forested catchment to some extent.

191 If we estimated the downward water flux at each soil layer, together with the NO_3^-
192 concentration and $\Delta^{17}\text{O}$ value of NO_3^- in each soil layer using a tension-free lysimeter
193 (Inoue et al., 2021), we could estimate the vertical change in the leaching flux of
194 NO_3^- for each soil layer along with the $\Delta^{17}\text{O}$ value of soil NO_3^- in each layer. Then,
195 applying Eq. (7) in each layer, we can more accurately estimate the GNR for the
196 forested catchment by integrating the GNR estimated for each soil layer together with
197 a more accurate metabolic rate of NO_3^- (GDR + uptake) of the forested catchment.
198 However, without such an observation of the distribution of the $\Delta^{17}\text{O}$ value of NO_3^- , it
199 is difficult to assume that the $\Delta^{17}\text{O}$ values of soil NO_3^- are always equal to those of
200 stream NO_3^- ; thus, the GNR should be reported with errors in which the possible
201 variations in the $\Delta^{17}\text{O}$ values of soil NO_3^- are considered.

202

203 **4 Conclusion**

204 Past studies have proposed the $\Delta^{17}\text{O}$ method to determine the GNR in each forested
205 catchment. The equations used in the calculation presuppose that the $\Delta^{17}\text{O}$ values of
206 NO_3^- in forested soils are homogeneous and equal to those of NO_3^- in the stream;
207 however, in reality, the values are often heterogeneous and do not always equal to
208 those corresponding to the stream. It is essential to clarify/verify the $\Delta^{17}\text{O}$ values of



209 NO_3^- in the forested soils and stream before applying the stream $\text{NO}_3^- \Delta^{17}\text{O}$ values to
210 estimate the GNR.

211

212 *Data availability.* All data are presented in the Supplement.

213

214 *Author contributions.* WD, UT, and FN designed the study. WD and UT performed
215 data analysis and wrote the paper.

216

217 *Competing interests.* The authors declare that they have no conflict of interest.

218

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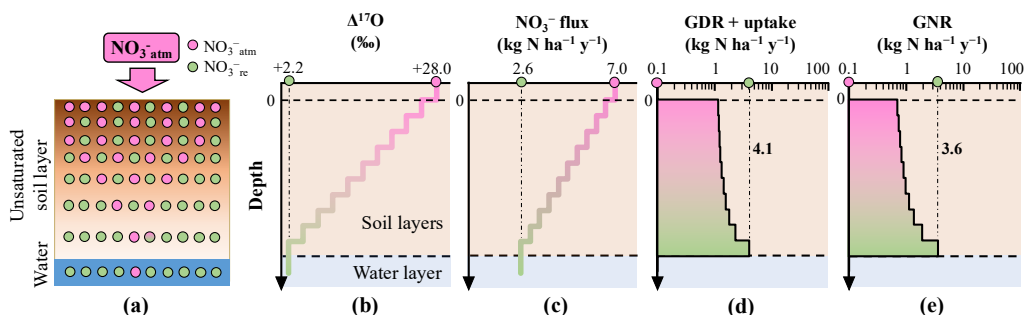


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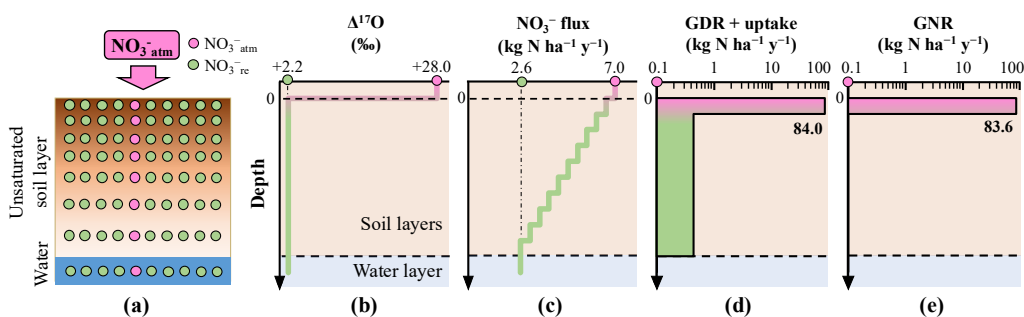
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304 **Figure 1.** Distribution of $\text{NO}_3^-_{\text{atm}}$ in the simulated forested soil where the distribution
 305 of the $\Delta^{17}\text{O}$ values of NO_3^- is heterogeneous (a). Vertical distribution of the following
 306 parameters in the forested soil: the simulated $\Delta^{17}\text{O}$ values of NO_3^- (b), simulated
 307 leaching flux of NO_3^- (c), estimated NO_3^- consumption rate (GDR + uptake) (d), and
 308 estimated GNR (e).

309



310 **Figure 2.** Distribution of $\text{NO}_3^-_{\text{atm}}$ in the simulated forested soil where the distribution
 311 of the $\Delta^{17}\text{O}$ values of NO_3^- is homogeneous (a). Vertical distribution of the following
 312 parameters in the forested soil: the simulated $\Delta^{17}\text{O}$ values of NO_3^- (b), simulated
 313 leaching flux of NO_3^- (c), estimated NO_3^- consumption rate (GDR + uptake) (d), and
 314 estimated GNR (e).