

BiasErrors associated with calculating the gross nitrification rates in forested catchments using the triple oxygen isotopic composition ($\Delta^{17}\text{O}$) of stream nitrate

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1 Abstract

2 A novel method for quantifying the gross nitrification rate (GNR) in ~~each~~ forested
3 catchments using the triple oxygen isotopic composition ($\Delta^{17}\text{O}$) of stream nitrate
4 ~~eluted~~ leached from the catchments has been proposed and applied in ~~several~~ recent
5 studies. However, the equations used in the calculations include the approximation
6 that the $\Delta^{17}\text{O}$ value of nitrate ~~metabolized~~ consumed through ~~either~~ assimilation or
7 denitrification ~~with~~ in the forested soil is equal to the $\Delta^{17}\text{O}$ value of stream nitrate ~~in~~
8 ~~the stream~~. The GNR estimated from the $\Delta^{17}\text{O}$ value of stream nitrate was more than
9 six times the ~~actual~~ GNR in our simulated calculation for a forested catchment where
10 the soil nitrate ~~in the soil~~ had exhibited $\Delta^{17}\text{O}$ values ~~larger~~ higher than those ~~in~~ the
11 stream nitrate. The $\Delta^{17}\text{O}$ values of the soil nitrate decreased ~~while showing a~~
12 ~~decreasing trend~~ with an increase in depth to ~~increasing depths until~~ that of the stream
13 nitrate at the bottom. ~~As most~~ Most of the reported soil nitrate in forested catchments
14 from past studies showed $\Delta^{17}\text{O}$ values higher than those of the stream nitrate
15 ~~eluted~~ leached from ~~the each~~ catchments. Thus, we concluded that the GNR estimated
16 from the $\Delta^{17}\text{O}$ value of stream nitrate in the forested catchments was, to ~~some~~ an
17 extent, an overestimate of the actual GNR.

18

19 1 Introduction

20 Nitrate (NO_3^-) is ~~one of the~~ an important nitrogen nutrients for primary production
21 in forested ecosystems. Nitrification is the microbial process that produces NO_3^- in
22 ~~each~~ forested ecosystems. Thus, quantifying the nitrification rate can assist in the

23 evaluation of the present and future states of ~~each~~-forested ecosystems. ~~While~~†The net
24 nitrification rate can be estimated from ~~the~~-an increase in NO_3^- concentration during a
25 certain period. ~~;~~ However, the gross nitrification rate (GNR), ~~which includes the~~ (net
26 nitrification rate ~~and the~~ + metabolic-consumption rate of nitrate (e.g., ~~that~~ assimilated
27 by plants or decomposed through denitrification)), reflects the internal N cycling
28 better than the net nitrification rate (Bengtsson et al., 2003), especially in forested
29 ecosystems. ~~;~~ ~~where~~ While/Although the net nitrification rate is often negligible (Stark
30 and Hart, 1997) ~~while~~ the ~~metabolic-consumption~~ rate is significant in forested
31 ecosystems, ~~so that~~ such the GNR often exceeds the net nitrification rate by several
32 orders of magnitude ~~;~~ (Verchot et al., 2001).

33 Recently, several studies have successfully estimated the GNR in water
34 environments, such as lakes, using the $\Delta^{17}\text{O}$ values of NO_3^- , as a ~~conserved~~-
35 conservative tracer of the mixing ratio between ~~the~~-atmospheric nitrate ($\text{NO}_3^-_{\text{atm}}$)
36 ~~deposited into the water environment~~ and ~~the remineralized~~ biologically produced
37 nitrate ($\text{NO}_3^-_{\text{bio}}$) ~~produced through nitrification therein~~ (Tsunogai et al., 2011, 2018).
38 The $\text{NO}_3^-_{\text{atm}}$ is deposited in the water environment, and the $\text{NO}_3^-_{\text{bio}}$ is produced
39 through nitrification. Although ~~The~~ $\text{NO}_3^-_{\text{bio}}$ always shows the $\Delta^{17}\text{O}$ values close to
40 0 ‰ because its oxygen atoms ~~derive~~ are derived from either terrestrial O_2 or H_2O
41 through nitrification. ~~In contrast~~ Contrarily, ~~the~~ $\text{NO}_3^-_{\text{atm}}$ always displays an
42 anomalous enrichment in ^{17}O with $\Delta^{17}\text{O}$ values being approximately $+26 \pm 3$ ‰ in
43 Japan (Tsunogai et al., 2010, 2016; Ding et al., 2022, 2023) because of oxygen

44 transfers from atmospheric ozone (Michalski et al., 2003; Nelson et al., 2018).
 45 Additionally, $\Delta^{17}\text{O}$ is almost stable during “mass-dependent” isotope fractionation
 46 processes (Michalski et al., 2004; Tsunogai et al., 2016). This is because possible
 47 variations in the $\delta^{17}\text{O}$ and $\delta^{18}\text{O}$ values during the processes of biogeochemical isotope
 48 fractionation follow the relation of $\delta^{17}\text{O} \approx 0.5 \delta^{18}\text{O}$, which cancels out the variations
 49 in the $\Delta^{17}\text{O}$ value. ~~Therefore~~ Thus, regardless of the partial ~~metabolism~~ consumption
 50 through denitrification or assimilation after deposition in a water column, the $\Delta^{17}\text{O}$
 51 can be used as a conservative tracer of $\text{NO}_3^-_{\text{atm}}$ to calculate the mixing ratio of
 52 $\text{NO}_3^-_{\text{atm}}$ to total NO_3^- ($\text{NO}_3^-_{\text{atm}}/\text{NO}_3^-_{\text{total}}$) in a water column using the following
 53 equation:

$$54 \quad [\text{NO}_3^-_{\text{atm}}]/[\text{NO}_3^-_{\text{total}}] = [\text{NO}_3^-_{\text{atm}}]/([\text{NO}_3^-_{\text{bio}}] + [\text{NO}_3^-_{\text{atm}}]) = \Delta^{17}\text{O}/\Delta^{17}\text{O}_{\text{atm}} \quad \text{---} \quad (1)$$

55 where the $\Delta^{17}\text{O}_{\text{atm}}$ and $\Delta^{17}\text{O}$ denote the $\Delta^{17}\text{O}$ values of $\text{NO}_3^-_{\text{atm}}$ and NO_3^- dissolved in
 56 ~~each~~ the water environment, respectively. Using ~~both~~ the $\text{NO}_3^-_{\text{atm}}/\text{NO}_3^-_{\text{total}}$ ratio
 57 estimated from the $\Delta^{17}\text{O}$ value of NO_3^- in a lake water column and the deposition rate
 58 of $\text{NO}_3^-_{\text{atm}}$ into the lake, the GNR ~~has been~~ was successfully estimated. This is because
 59 the partial consumption of NO_3^- has little influence on the $\text{NO}_3^-_{\text{atm}}/\text{NO}_3^-_{\text{total}}$ ratio in
 60 the lake water column (Tsunogai et al., 2011, 2018).

61 In addition to application in water environments, the $\Delta^{17}\text{O}$ method has ~~also~~ been
 62 applied to forested catchments ~~for to determine their~~ GNR ~~determination~~ (Fang et al.,
 63 2015; Hattori et al., 2019; Huang et al., 2020). ~~By u~~ Using the deposition flux of
 64 $\text{NO}_3^-_{\text{atm}}$ into the catchment ~~as well as~~ and the ~~elution~~ leaching flux of ~~both~~ unprocessed

65 $\text{NO}_3^-_{\text{atm}}$ and $\text{NO}_3^-_{\text{bio}}$ from ~~the~~ streams, ~~which can be determined from the $\Delta^{17}\text{O}$~~
66 ~~values of NO_3^- in stream water eluted from the catchment,~~ the GNR in ~~each~~ a the
67 forested catchment ~~has been~~ was estimated ~~in a manner similar~~ similarly to the
68 estimation for ~~the~~ water environments (Fang et al., 2015). Contrary to water
69 environments, where the $\Delta^{17}\text{O}$ values of NO_3^- ~~with~~ in the water layers are
70 homogeneous in the water column due to the active vertical mixing of water – and can
71 be measured easily, it is often difficult to determine the $\Delta^{17}\text{O}$ values of ~~the~~ NO_3^-
72 ~~metabolized~~ consumed in soil layers. Consequently, past studies have approximated
73 the values to be equal to the $\Delta^{17}\text{O}$ value of stream NO_3^- ~~eluted~~ leached from ~~each~~
74 forested catchments s without actual observation (Fang et al., 2015, Hattori et al., 2019,
75 Huang et al., 2020). However, such an approximation should be conducted with
76 extreme caution, as the $\Delta^{17}\text{O}$ values of soil NO_3^- are not always equal to those of ~~the~~
77 streams s – (Hattori et al., 2019, Rose, 2014, Nakagawa et al., 2018 ~~Osaka et al.,~~
78 2010). To clarify the details of the approximation ~~along with~~ and its impact on the
79 final estimated GNR, we present an accurate relationship between the $\Delta^{17}\text{O}$ of soil
80 NO_3^- and the GNR, ~~starting from the~~ using basic isotope mass balance equations.
81 Then, we present the ~~GNR-estimated~~ GNR for a forested catchment ~~in which~~
82 ~~the~~ whose $\Delta^{17}\text{O}$ values of soil NO_3^- ~~in soil are~~ were measured. Finally, we compared d
83 the GNR estimated in this study with the GNR estimated from the $\Delta^{17}\text{O}$ values of
84 stream NO_3^- .

85

86 2 Calculation

87 The total mass balance equation of NO_3^- including the GNR in ~~each~~ catchments
88 can be expressed as follows:

$$89 \text{NO}_3^- \text{deposition} + \text{GNR} = \text{NO}_3^- \text{leaching} + \text{NO}_3^- \text{uptake} + \text{GDR} \quad (2)$$

90 where $\text{NO}_3^- \text{deposition}$, GNR, $\text{NO}_3^- \text{leaching}$, $\text{NO}_3^- \text{uptake}$, and GDR denote the deposition flux
91 of NO_3^- into the ~~each~~ catchment, ~~gross nitrification rate~~ GNR in the ~~each~~ catchment,
92 leaching flux of NO_3^- from the ~~each~~ catchment, uptake rate of NO_3^- in the ~~each~~—
93 catchment, and gross denitrification rate in the ~~each~~ catchment, respectively.

94 The isotope mass balance for each $\Delta^{17}\text{O}$ value of NO_3^- in the catchment can also ~~be~~
95 be expressed using a ~~the similar equation~~ calculated using the same method:

$$96 \text{NO}_3^- \text{deposition} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} + \text{GNR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}} = \text{NO}_3^- \text{leaching} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}} + \text{NO}_3^- \text{uptake} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}} + \text{GDR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}} \quad (3)$$

98 where $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}}$, and
99 $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}}$ denote the $\Delta^{17}\text{O}$ value of NO_3^- deposited into the ~~each~~—
100 catchment, that of the NO_3^- ~~bio~~ produced through nitrification, that of the NO_3^-
101 ~~eluted~~ leached from the ~~each~~ catchment, that of the NO_3^- assimilated by plants and
102 other organisms in the ~~each~~ catchment, and that of the NO_3^- decomposed through
103 denitrification in the ~~each~~ catchment, respectively.

104 If the $\Delta^{17}\text{O}$ values of the NO_3^- in the forested soil layers, where the NO_3^- was
105 ~~metabolized~~ consumed through ~~either~~ assimilation ~~(by plants and other organisms)~~ or
106 denitrification, are equal to the $\Delta^{17}\text{O}$ values of NO_3^- in the stream, we could obtain
107 Eq. 4 ~~can be expressed as follows~~:

108 $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{uptake}} = \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{denitrification}} = \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ (4)

109 Consequently, by combining Eqs. 3 and 4, we could obtain the following

110 relationship:

111 $\text{NO}_3^-_{\text{deposition}} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} + \text{GNR} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}} = (\text{NO}_3^-_{\text{leaching}} + \text{NO}_3^-_{\text{uptake}}$
 112 $e + \text{GDR}) \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ (5)

113 We could estimate the GNR using Eq. 6 obtained from Eqs. 2 and 5 because we can

114 approximate the $\Delta^{17}\text{O}$ values of NO_3^- ~~bio~~ produced through nitrification

115 ($\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{nitrification}}$) to ~~be~~ 0 (Michalski et al., 2003; Tsunogai et al., 2010):

116 $\text{GNR} = \text{NO}_3^-_{\text{deposition}} \times (\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} - \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}) / \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ (6)

117 Eq. 6 corresponds to the equations used in previous studies ~~for to~~ quantifying the

118 GNR in ~~the each~~ forested catchments (Eq. 4 in Fang et al., 2015; Eq. 8 in Hattori et

119 al., 2019; Eq. 4 ~~in~~ Huang et al., 2020).

120

121 3 Results and Discussion

122 The $\Delta^{17}\text{O}$ values of NO_3^- in forested soil layers should be equal to those of stream

123 NO_3^- ~~in the stream~~, as presented in Eq. 4 to obtain Eq. 6. While the number of

124 simultaneous observations of the oxygen isotopes of NO_3^- in both the soil and stream

125 in a given forested catchment is limited (Hattori et al., 2019, Osaka et al., 2010, Rose,

126 2014, Nakagawa et al., 2018), the ~~limited~~ observations show that the oxygen isotopic

127 ratios of soil NO_3^- are mostly higher than those of stream NO_3^- . Differ from water

128 environments, vertical mixing of water/soil is difficult in forested soil, so the $\Delta^{17}\text{O}$

129 values of soil NO₃⁻ are often heterogeneous. For example, Hattori et al. (2019)
130 reported that ~~more than~~over 60 % of the soil exhibited $\Delta^{17}\text{O}$ values significantly
131 higher than those of stream NO₃⁻ determined simultaneously ($\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ $\Delta^{17}\text{O}$
132 = +1 to +3 ‰). In addition, they found a decreasing $\Delta^{17}\text{O}$ trend in soil NO₃⁻ with
133 depth, declining from greater than +20 ‰ at the surface to less than +3 ‰ at depths of
134 25–90 cm from the surface. A similar ~~decreasing~~ trend in the vertical distribution ~~had-~~
135 ~~been found~~was observed in $\delta^{18}\text{O}$ in another forested catchment, from ~~greater-~~
136 ~~than~~above +35 ‰ at the surface soil to less than +10 ‰ at depths of 30–50 cm from
137 the soil surface (Osaka et al., 2010). ~~Besides~~In addition, most of the soil NO₃⁻ also
138 exhibited $\delta^{18}\text{O}$ values higher than those ~~in-of~~ the stream NO₃⁻ (Osaka et al., 2010).
139 ~~Furthermore~~, Rose (2014) monitored the horizontal distribution of the $\Delta^{17}\text{O}$ of soil
140 NO₃⁻ by randomly setting 15 tension-free lysimeters at depths of 0–10 cm in a 39 ha
141 forested catchment, ~~reporting~~ $\Delta^{17}\text{O}$ values. They reported significantly higher $\Delta^{17}\text{O}$
142 values in soil NO₃⁻ ($+9.1 \pm 5.8$ ‰ on average) than ~~in-the~~those of stream NO₃⁻
143 ($+0.5$ ‰ on average) ~~eluted~~leached from the forested catchment. As most ~~of-the~~fine
144 roots and root biomass isare concentrated in the top 10 cm of the soil in forested
145 catchments (Jackson et al., 1996; Li et al., 2020), most uptake reactions should occur
146 in that top 10 cm of soil. Consequently, the significant difference in the $\Delta^{17}\text{O}$ values
147 between soil NO₃⁻ and stream NO₃⁻, particularly in ~~the~~ surface soil layers, implies
148 that the estimated GNRs in ~~the~~ forested catchments s obtained from Eq. 6 were
149 inaccurate.

150 To demonstrate the impact of the differences in the $\Delta^{17}\text{O}$ values of ~~between~~ soil
151 NO_3^- and stream NO_3^- on the GNR and present, ~~along with presenting~~ the problems
152 associated with the approximation to obtain Eq. 6, we estimated the GNR of ~~for~~ two
153 simulated forested soils. The $\Delta^{17}\text{O}$ of the first soil ~~—one~~ with NO_3^- decreased
154 showing a decreasing trend in $\Delta^{17}\text{O}$ down to the $\Delta^{17}\text{O}$ of the stream NO_3^-
155 (heterogeneous soil) (Figs. 1a and 1b). The second soil ~~and one~~ with NO_3^- showing
156 the same $\Delta^{17}\text{O}$ values as those of the stream NO_3^- (homogeneous soil) (Figs. 2a and
157 2b). With Because Hattori et al. (2019) reported ing the NO_3^- deposition as 7.0 kg of N
158 $\text{ha}^{-1} \text{y}^{-1}$, NO_3^- leaching as 2.6 kg of N $\text{ha}^{-1} \text{y}^{-1}$, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}}$ as +28.0 ‰, and
159 $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$ as +2.2 ‰ in ~~the~~ forested catchment they studied. We adopted
160 the same values in ~~our~~ the present calculation to simulate the same forested soil. All
161 the symbols (e.g., GNR) used in this study ~~here were in accordance~~ consistent with
162 those of Hattori et al. (2019) as well.

163 To estimate the GNR, ~~w~~ We divided ~~the soils in the~~ heterogeneous forest soils into
164 10 layers (i.e., 10 steps) in the vertical direction, ~~simulating the soils observed by~~
165 ~~Hattori et al. (2019), in which~~ The $\Delta^{17}\text{O}$ values of NO_3^- gradually decreased with
166 ~~increasing~~ an increase in depth, varying from +28.0 to +2.2 ‰ with a rate of
167 ~~—decrease of~~ +2.58 ‰ for each step (Fig. 1b). Similarly, we assumed a gradual
168 decrease with an increasing in depth in the leaching flux of NO_3^- , ~~i.e.,~~ (from 7 to 2.6
169 kg of N $\text{ha}^{-1} \text{y}^{-1}$ at ~~with~~ a rate of ~~decrease of~~ —0.44 kg of N $\text{ha}^{-1} \text{y}^{-1}$ for each per step)
170 (Fig. 1c). which This simulated the gradual net consumption of NO_3^- in accordance

171 with water flow in forested soils. ~~In~~†The homogeneous ~~forest soils,~~ wase also divided
172 ~~the forested soils~~ into 10 layers in the vertical direction. The change ~~with depth~~ in the
173 leaching flux of NO_3^- with depth was the same as that in the heterogeneous soils (Fig.
174 2c), whereas the $\Delta^{17}\text{O}$ values of NO_3^- were constant at +2.2 ‰ in the ~~soil~~ layers (Fig.
175 2b).

176 Applying the total mass balance and isotope mass balance of NO_3^- shown in Eqs. 2
177 and 3 to each layer, we estimated ~~both~~ the GNR (Figs. 1e and 2e) and total
178 ~~consumption~~metabolic rate of NO_3^- (GDR + uptake) (Figs. 1d and 2d) in each layer.
179 We assumeding the following ~~s:~~ (1) The $\Delta^{17}\text{O}$ values of NO_3^- ~~were~~are constant in
180 each layer.; (2) The vertical flow of NO_3^- in the soil layers proceededed downward from
181 the surface to the ~~water~~ final layer (No. 10) with a uniform residence time in each
182 layer.; ~~and~~ Finally, (3) the GNR and ~~consumption~~metabolic rate of NO_3^- (GDR +
183 uptake) wasis zero in the ~~soil~~ water layer (~~layers beyond the no. 10 soil~~ final layer) ~~–~~
184 (No. 10). ~~Thereafter,~~ by integrating the GNR determined for each layer, we
185 estimated the total GNR in ~~each~~ the forested catchment.

186 Although the GNR estimated for the catchment with the homogeneous $\Delta^{17}\text{O}$ values
187 in soil NO_3^- (~~homogeneous soil~~) was 83.6 kg of N $\text{ha}^{-1} \text{y}^{-1}$ (Fig. 2e), exactly equal to
188 that estimated by Hattori et al. (2019) using Eq. 6 (~~Fig. 2e~~), the ~~total~~ GNR estimated
189 for the catchment with the heterogeneous $\Delta^{17}\text{O}$ values in soil NO_3^- (~~heterogeneous~~
190 soil) was considerably lower ~~a much smaller~~ (13.0 kg of N $\text{ha}^{-1} \text{y}^{-1}$; Fig. 1e), while
191 the same parameters with the homogeneous $\Delta^{17}\text{O}$ values in soil NO_3^- were used for

192 NO_3^- deposition, NO_3^- leaching, $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}}$, and $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}$. simulated for the
 193 catchment with the heterogeneous $\Delta^{17}\text{O}$ values in soil NO_3^- (Fig. 1e). Even if we
 194 increased the number of the layers in the forested soils was increased into 20, 30, 50,
 195 100, and 1000 to enhance the precision of the GNR simulated for the catchment with
 196 the heterogeneous $\Delta^{17}\text{O}$ values in soil NO_3^- , the GNR was 11.4, 11.0, 10.5, 10.3, and
 197 10.1 kg of N $\text{ha}^{-1} \text{y}^{-1}$, respectively. Consequently, we concluded the following: (1)
 198 Past studies estimated the GNR using Eq. 6 did not approximated the $\Delta^{17}\text{O}$ values of
 199 soil NO_3^- consumed in the forested catchments were homogeneous and always
 200 equal to that of stream NO_3^- mathematically and. (2) The differences between the
 201 $\Delta^{17}\text{O}$ values of the soil NO_3^- metabolized consumed in a forested catchment and that
 202 of stream NO_3^- resulted in a significant deviation in the GNR estimated using Eq. 6
 203 from the actual GNR.

204 By combining the mass balance and isotope mass balance shown in Eqs. 2 and 3,
 205 Eq. 7 can be obtained as the equation to estimate the GNR accurately:

$$206 \text{GNR} = \text{NO}_3^- \text{leaching} - \text{NO}_3^- \text{deposition} + (\text{NO}_3^- \text{deposition} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} -$$

$$207 \text{NO}_3^- \text{leaching} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}) / \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{soil}} \quad (7)$$

208 where $\Delta^{17}\text{O}(\text{NO}_3^-)_{\text{soil}}$ denotes the “average” $\Delta^{17}\text{O}$ value of NO_3^- consumed through
 209 either assimilation or denitrification in the forested catchment. As most of the soil
 210 NO_3^- measured to date exhibited $\Delta^{17}\text{O}$ values higher than those of the stream NO_3^-
 211 leached from the catchments (Hattori et al., 2019, Rose, 2014). Consequently, the
 212 GNR estimated from stream NO_3^- using Eq. 6 was higher than exceeded the GNR

213 ~~estimated from soil NO₃⁻ using Eq. 7, to an~~ ~~some extent. In other words~~ ~~Therefore, the~~
 214 ~~GNR estimated from Eq. 6 was overestimated~~ ~~the GNR in each forested catchment to~~
 215 ~~some~~ ~~an extent.~~

216 ~~Note that~~ ~~†~~ ~~The linear variation in the leaching flux and Δ¹⁷O values of soil NO₃⁻~~
 217 ~~used in the simulated calculations (Fig. 1) is just one of many possible variations in~~
 218 ~~the~~ ~~forested catchments. It is impossible to~~ ~~decide~~ ~~determine~~ ~~whether the linear~~
 219 ~~variation was realistic~~ ~~or not~~ ~~until the downward water flux, along with the~~
 220 ~~concentration and Δ¹⁷O values of NO₃⁻, was~~ ~~is~~ ~~determined for each soil layer.~~

221 However, the simultaneous observations of the oxygen isotopes of soil NO₃⁻ and
 222 stream NO₃⁻ (Hattori et al., 2019; Osaka et al., 2010; Nakagawa et al., 2018; Rose,
 223 2014) ~~implied~~ ~~y~~ ~~that the approximation of the Δ¹⁷O values of~~ ~~the~~ ~~soil NO₃⁻~~
 224 ~~metabolized through assimilation or denitrification to be always equal to the Δ¹⁷O~~
 225 ~~value~~ ~~that~~ ~~of the~~ ~~stream NO₃⁻, shown in (Fig. 2b), was~~ ~~is~~ ~~unrealistic.~~

226 ~~By combining the mass balance and isotope mass balance shown in Eqs. 2 and 3,~~
 227 ~~Eq. 7 can be obtained to accurately estimate the GNR:~~

$$228 \text{GNR} = \text{NO}_3^- \text{leaching} - \text{NO}_3^- \text{deposition} + (\text{NO}_3^- \text{deposition} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{atm}} -$$

$$229 \text{NO}_3^- \text{leaching} \times \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{stream}}) / \Delta^{17}\text{O}(\text{NO}_3^-)_{\text{soil}} \quad (7)$$

230 ~~where Δ¹⁷O(NO₃⁻)_{soil} denotes the Δ¹⁷O values of NO₃⁻ in forested soil, from which~~
 231 ~~the NO₃⁻ was metabolized through either assimilation or denitrification. As most of~~
 232 ~~the soil NO₃⁻ measured to date exhibit Δ¹⁷O values higher than those of the stream~~
 233 ~~NO₃⁻ eluted from each catchment (Hattori et al., 2019, Rose, 2014), the GNR~~

234 ~~estimated from stream NO_3^- using Eq. 6 is higher than the GNR estimated from soil-~~
235 ~~NO_3^- using Eq. 7, to some extent. In other words, the GNR estimated from Eq. 6-~~
236 ~~overestimated the GNR in each forested catchment to some extent.~~

237 If we estimated the downward water flux at each soil layer, ~~together~~ with the NO_3^-
238 concentration and $\Delta^{17}\text{O}$ value of NO_3^- in each soil layer ~~through the methods such as~~
239 ~~those~~ using, e.g., a tension-free lysimeter (Inoue et al., 2021), we could estimate the
240 vertical change in the leaching flux of NO_3^- for each soil layer along with the $\Delta^{17}\text{O}$
241 ~~value~~ of soil NO_3^- ~~in each layer~~. Thereafter, applying Eq. (7) ~~to~~ each layer, we can
242 ~~more accurately~~ more accurately estimate the GNR for the forested catchment ~~more~~
243 ~~accurately~~, by integrating the GNR estimated for each soil layer ~~together~~ with a more
244 accurate NO_3^- consumption metabolic rate of NO_3^- (GDR + uptake) of the forested
245 catchment. ~~However, w~~ Without such an observation of the distribution ~~and the~~
246 leaching flux of ~~the $\Delta^{17}\text{O}$ value of NO_3^- , together with the $\Delta^{17}\text{O}$ values in forest soil,~~
247 ~~it is difficult to assume that the $\Delta^{17}\text{O}$ values of soil NO_3^- are always equal to those of~~
248 ~~stream NO_3^- ; thus,~~ the GNR ~~estimated by~~ using Eq. (6), assuming that the $\Delta^{17}\text{O}$
249 values of soil NO_3^- are always equal to those of stream NO_3^- , should be reported with
250 significant errors in which the possible variations in the $\Delta^{17}\text{O}$ values of soil NO_3^- are
251 considered.

252

253 **4 Conclusion**

254 Past studies have proposed the $\Delta^{17}\text{O}$ method ~~to determine~~for determining the GNR
255 in ~~each~~-forested catchments. The equations used in the calculation presuppose that the
256 $\Delta^{17}\text{O}$ values of NO_3^- consumed in forested soils are homogeneous and equal to those
257 of ~~NO_3^- in the stream~~ NO_3^- . However, in reality, the values are often heterogeneous
258 and do not always equal ~~to~~ those corresponding to of the stream. It is essential to
259 clarify/verify the $\Delta^{17}\text{O}$ values of NO_3^- in ~~the~~ forested soils and streams before
260 applying the stream NO_3^- - $\Delta^{17}\text{O}$ values of stream NO_3^- to estimate the GNR.

261

262 *Data availability.* All data are presented in the Supplement.

263

264 *Author contributions.* WD, UT, and FN designed the study. WD and UT performed
265 data analysis and wrote the paper.

266

267 *Competing interests.* The authors declare that they have no conflict of interest.

268

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282

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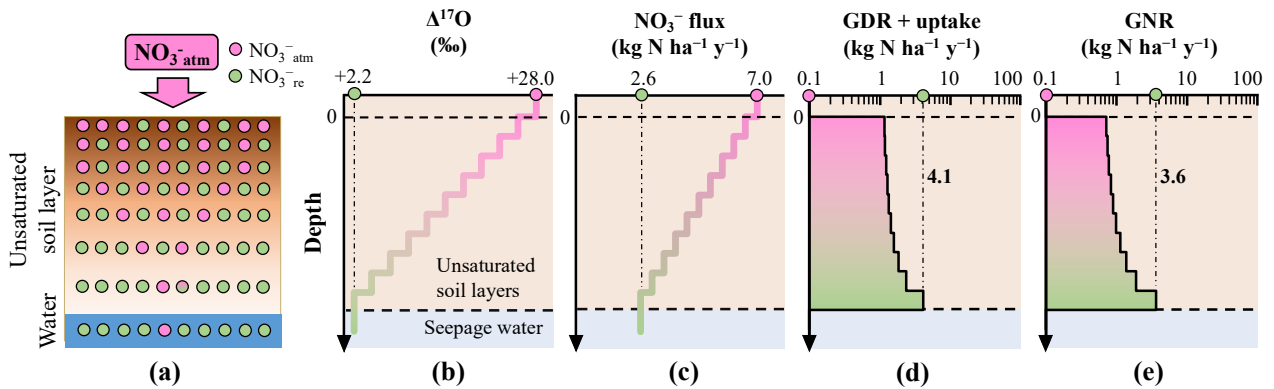
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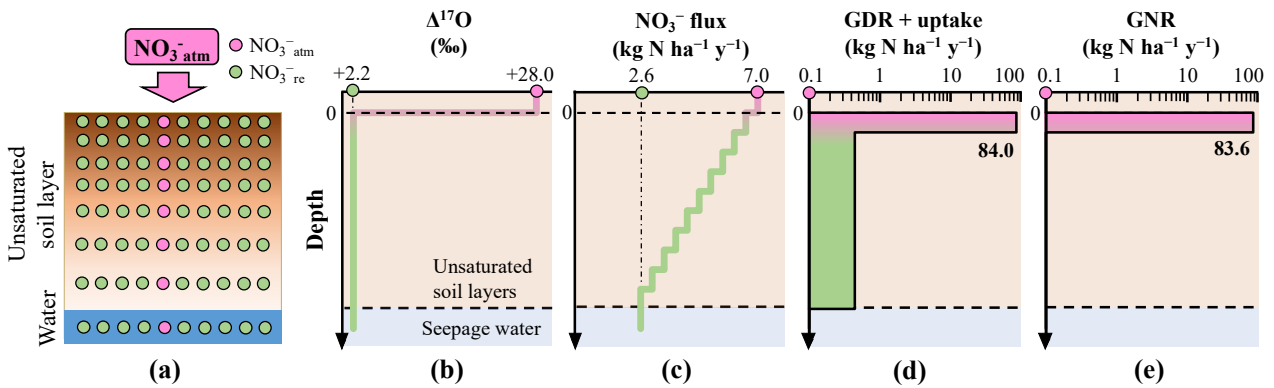
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367 **Figure. 1.** Distribution of $\text{NO}_3^-_{\text{atm}}$ in the simulated forested soil where the with
 368 heterogeneous distribution of the $\Delta^{17}\text{O}$ values of NO_3^- is heterogeneous (a). Vertical
 369 distribution of the following parameters in the forested soil: the simulated $\Delta^{17}\text{O}$ values
 370 of NO_3^- (b), simulated leaching flux of NO_3^- (c), estimated NO_3^- consumption rate
 371 (GDR + uptake) (d), and estimated GNR (e).

372



373 **Figure. 2.** Distribution of $\text{NO}_3^-_{\text{atm}}$ in the simulated forested soil with homogeneous
 374 where the distribution of the $\Delta^{17}\text{O}$ values of NO_3^- is homogeneous (a). Vertical
 375 distribution of the following parameters in the forested soil: the simulated $\Delta^{17}\text{O}$ values
 376 of NO_3^- (b), simulated leaching flux of NO_3^- (c), estimated NO_3^- consumption rate
 377 (GDR + uptake) (d), and estimated GNR (e).