



1 Efficient droplet activation of ambient black carbon particles in sub-

2 urban environment

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18 Abstract

19 The cloud condensation nuclei (CCN) activity of black carbon particles (BC) importantly determines their 20 impacts on cloud microphysics and atmospheric lifetime. This process is crucially determined by the amount of more hygroscopic coating materials BC can acquire during aging process. It remains a challenge for ambient 21 22 measurements to capture this process and link this with CCN activity of BC. Here by using coupled 23 measurements of size-resolved number concentrations of CCN at configured water supersaturation (SS) and BC-24 containing particles (BCc), we are able to directly monitor the droplet activation diameter (D50) and activation 25 fraction of BCc ($F_{act,BC}$) in ambient environment. The number concentration of BCc was found to peak at diameter 180-210 nm after acquiring coatings, larger than that for all particles (50-150 nm). This led to that the 26 initially smaller BC can be enlarged and become larger and more hygroscopic, even having higher CCN activities 27 28 than other particles. The $F_{act,BC}$ increased from 42% to 69% in number and from 67% to 85% in mass, as SS 29 increased from 0.1% to 0.2%, but tended to reach a plateau when SS>0.2%. Notably, Fact, BC and D50 linearly 30 correlated with equivalent photochemical age, at a rate of +2% per hour and -3nm per hour, respectively. The 31 results suggest BCc from anthropogenic sources can readily serve as CCN at a relatively low SS, and over half of the BC population can be activated in a few hours. This means the surface-sourced BC can efficiently incorporate 32

33 into clouds and may exert important indirect radiative impacts.

34 Keywords: black carbon; cloud condensation nuclei; droplet activation fraction; photochemical age





36 Introduction

Black carbon (BC), a significant anthropogenic pollutant, generate from the incomplete combustion of fossil fuels and profoundly influences the climate through direct (Ramanathan and Carmichael, 2008; Bond et al., 2013), indirect (Koch et al., 2011), and semi-direct (Koch and Del Genio, 2010) effects. The climate response to BC is sensitive to its lifetime (Hansen et al., 1997), which remains a considerable uncertainty in model studies due to spatial and temporal variations (Park et al., 2005; Koch et al., 2009; Lund et al., 2017).

42 Wet scavenging, the primary removal mechanism of atmospheric BC (Jacobson, 2010), that BC particles serve as 43 cloud condensation nucleation (CCN) activated into the cloud droplets and removed through precipitation. 44 Although freshly emitted BC particles are generally hydrophobic, they become more hygroscopic through 45 atmospheric processing, acquiring soluble materials like sulfates, nitrates, and secondary organic aerosol 46 compounds, or by coagulating with other particles (Dusek et al., 2006; Tritscher et al., 2011). The increased hygroscopicity enhances CCN activation efficiency of BC (Weingartner et al., 1995), yet there is significant 47 48 variability in the literature regarding the degree to which BC serves as CCN (Wu et al., 2019; Rose et al., 2011), 49 due to the diversity in properties of BC resulting from its various sources and atmospheric aging process 50 (Henning et al., 2012; Dalirian et al., 2018).

The κ-Köhler theory (Petters and Kreidenweis, 2007), combined with the ZSR mixing rule (Pruppacher and Klett, 51 52 1997; Henning et al., 2010), has been successful in characterizing the hygroscopicity of internally mixed BC, 53 assuming a spherical core-shell structure. Chamber studies where the physic-chemical conditions and the initial emitted particles are well controlled indicate that the hygroscopicity of BC is determined by the fraction and 54 55 hygroscopicity of the coating material (Snider et al., 2010; Friebel et al., 2019). A relatively thin soluble coating 56 like 2nm corresponding to about 5% of the total volume for 250nm particles are enough to make hydrophobic BC CCN active under typical atmospheric supersaturation (Dalirian et al., 2018). The hygroscopic properties of the 57 58 coating materials can vary significantly; notably, increase of the hygroscopicity of BC induced by secondary 59 inorganic aerosols (e.g., ammonium nitrate) is more pronounced than that caused by organic aerosols (Kuwata et al., 2007; Liu et al., 2013). However, laboratory experiments often fail to replicate the complexity of atmospheric 60 61 conditions, comprehensive field measurements are need to capture the real-time evolution of the CCN properties 62 of BC.

63 Despite the advancements in understanding the physicochemical transformations of BC through field observation 64 (Liu et al., 2017), quantifying its CCN activity in the ambient conditions remains challenging due to the 65 variability in its sources, composition, mixing state, and atmospheric processes (Zhang et al., 2008). Direct field 66 observations indicate that number fraction of 50%-90% BC can be activated into cloud or fog, dependent on actual ambient supersaturation (Motos et al., 2019a; Motos et al., 2019b;). More than 50% of BC mass can been 67 68 removed through the wetting removing process during the transport from surface to planetary boundary layer 69 (PBL), as evidenced in studies in East Asia polluted area (Kondo et al., 2016; Liu et al., 2020) in east Asian 70 polluted area, but it's difficult to know true ambient supersaturation in the in-situ cloud observation. However, 71 true ambient supersaturation is challenging to determine in in-situ cloud observations.

72 The hygroscopic of BC particles can be derived from the measured mixed state of BC by size resolved soot 73 particle aerosol mass spectrometer (SP-AMS) (Onasch et al., 2012; Wu et al., 2019), and the hygroscopicity 74 tandem differential mobility (HTDMA) instrument under sub saturation condition, however, only limited 75 atmospheric CCN efficiency studies of BC particles were conducted under the water supersaturation condition. 76 Continuous-flow cloud condensation nuclei counter (CCNC) can measure the CCN number concentration of the 77 bulk aerosol under well-control supersaturation condition (Roberts and Nenes, 2005), but it cannot distinct how 78 many parts of BC are activated as the number proportion of BC account for about 10% of the ambient aerosol. 79 Recently, Hu et al. (2021a) developed an novel method to direct observe the number activation ratios of BC





under certain supersaturation, based on the size resolved CCNC and BC measurements, revealing potential
 overestimation of hygroscopicity of BC based on bulk chemical composition.

82 In this study, continuous observation of the CCN activation and hygroscopic properties of BC under water

83 supersaturation (0.1%, 0.2%, and 0.3%) were conducted during springtime in a sub-urban site in Beijing. The

84 variation of the hygroscopicity of BC during different pollution levels was investigated according to classification

of four different air masses clusters. The equivalent photochemical age (t_{age}) , calculated based on the Proton-

86 Transfer-Reactor Time-of-Flight Mass Spectrometer, was used to characterize the evolution of the CCN

87 activation properties of BC under ambient atmospheric aging.

88 2. EXPERIMENTAL AND METHODS

89 2.1. Measurements site and Instruments Setup

The filed campaign was conducted for more than 1 month from 29th Apr. to 2nd Jun. 2020 at Beijing Cloud Laboratory and Observational Utilities Deployment Base (CLOUD Base) (Tian et al., 2022), which is approximately 65Km to the northeast of the central Beijing, with the main local emissions are from industrial, traffic, and residential, which represent a typical sub-urban environment. Higher pollution condition at this site was correlate with pollutant regional transport from the west and south pollutant area, while clean condition was attribute to the pollutant dilution by the northwest clean air (Hu et al., 2022).

The instruments setup was shown in Figure 1a. The bulk and size-resolved physic-chemical and hygroscopicity properties of aerosol were simultaneously measured during the field campaign. A cyclone PM_{2.5} cut-off inlet (Model: SCC1.829, BGI Inc., USA) was used to remove the coarse mode particles (larger than 2.5µm), and the air was dried by a Nafion tube before entering the instruments. In addition to particle measurement, volatile organic compounds (VOCs) were measured by a Proton-Transfer-Reactor Time-of-Flight Mass Spectrometer (PTR-TOF-MS 8000, Ionicon Analytik GmbH Innsbruck, Austria) (Fig. 1a).

102 2.2 Aerosol Chemical Components Measurements and Hygroscopic Growth Factor Calculation

103 The mass concentration of non-refractory submicron aerosol species including organic aerosol (OA), chloride 104 (Chl⁻), ammonium (NH₄⁺), nitrate (NO₃⁻) and sulfate (SO₄⁻) was measured by the Aerodyne high-resolution time-105 of-flight AMS (HR-ToF-AMS, Aerodyne Research Inc., USA). The principle and instruction of HR-ToF-AMS 106 was already detailed in previous publications (Jayne et al., 2000; Drewnick et al., 2005). The HR-ToF-AMS was 107 calibrated before and after the experiment, and relative ionization efficiencies of NH₄⁺ (4.0) and SO₄²⁻ (1.2) were 108 calibrated using pure ammonium nitrate and ammonium sulfate, respectively, while default relative ionization 109 efficiency values were used for OA (1.4), NO₃⁻ (1.1), and Chl⁻ (1.3) (Middlebrook et al., 2012).

110 The measurement of individual refractory black carbon (rBC)-containing particles (BCc) and BC-free particles 111 (BCf) was conducted using the Single Particle Soot Photometer (SP2, DMT Inc., USA). The SP2 employs an 112 intense λ =1064 nm Nd:YAG laser to heat BCc until incandescence, with the mass and diameter (Dc) of rBC 113 quantified based on the intensity of the incandescence signal. Prior to the experiment, the incandescence signal

114 was calibrated using Aquadag standards (Acheson Inc., USA) and was corrected for ambient BC using a factor of

- 115 0.75 (Laborde et al., 2012).
- 116 The measurement of individual refractory BC (rBC)-containing particles (BCc) and BC-free particles (BCf) was
- 117 conducted using the Single Particle Soot Photometer (SP2, DMT Inc., USA). The SP2 use intense 1064nm
- 118 Nd:YAG laser beam to heat BCc to vaporize, with the mass and diameter (Dc) of rBC quantified based on the
- 119 intensity of the incandescence signal. The incandescence signal was calibrated using Aquadag standard (Acheson
- 120 Inc. USA) and was corrected for ambient BC using a factor of 0.75 before the experiment (Laborde et al., 2012).
- Particles exhibiting only scattering signal were classified as BCf, with their diameters determined based on intensity of scattering signal, calibrated by polystyrene latex spheres (PSLs). Due to the distortion of the
 - 4



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123 scattering signal caused by the vaporization of the coating material of BCc, the leading edge only (LEO) fitting 124 method was employed to derive the entire size of BCc (Gao et al., 2007; Liu et al., 2014). As the optical diameter 125 of BCc is largely independent of the morphology of BCc, the measured optical diameter by SP2 is assumed to be 126 equal to the volume equivalent diameter (D_{ve}) (Moteki et al., 2010; Hu et al., 2021b). Consequently, the relative 127 bulk volume of coating to BC ($V_{coating}/V_{BC}$) was calculated as the cubic ratio of diameter of D_{ve} and Dc (Liu et al., 128 2017):

$$VR = \frac{V_{coating}}{V_{BC}} = \frac{\sum_i D_{ve}^3}{\sum_i D_c^3} - 1$$
(1)

130 where D_{ve} is volume equivalent diameter of the entire BCc, and D_c is the rBC core diameter for the i^{th} single BCc, 131 respectively.

132 The hygroscopic growth factor (κ) which depend on the chemical properties (i.e. molecular weight, dissociate, 133 and osmotic coefficient, etc.) can be used to predict the activation properties of particles as CCN under certain 134 supersaturation (SS) (Petters and Kreidenweis, 2007). The hygroscopicity growth factor (κ) for all-particles (κ_{all}) 135 were calculated applying Zdanovskii-Stokes-Robinson (ZRS) mixing rule for the volume contribution from 136 different chemical compositions(Mcdonald, 1953). The expression for κ_{all} is:

137
$$\kappa_{all} = \sum_i \varepsilon_i \, \kappa_i \tag{2}$$

138 where κ_i and ε_i represents the hygroscopic growth factor and volume fraction for each chemical component (i.e. 139 (NH₄)₂SO4, NH₄HSO₄, NH₄NO3, and BC etc.).

140 The κ of single BCc (κ_{BC}) is depended on the volume fraction and κ of the coating material, and calculated as:

141
$$\kappa_{BCc} = \kappa_{coating} \times \left(\frac{VR}{1+VR}\right) \quad (3)$$

142 where the κ_{coating} was chosen as the κ_{all} which also represent the average κ of particles, and $V_{\text{coating}}/V_{BC}$ is 143 volume ratios of coating material over the BC, which reflect the coating thickness of BCc particles. The κ_{BC} was 144 test successfully to predict the activation of BCc as CCN in the field in-situ cloud observation (Motos et al., 145 2019a).

146 2.3 Size Resolved CCN Activity of All-particles and BCc

147 Atmospheric polydisperse aerosols were classified into monodisperse particles at the outlet of a Differential 148 Mobility Analyser (DMA, Model: 3081, TSI Inc., USA). Particles of different mobility sizes were selected by applying varying voltage to the DMA, maintaining a specific sheath flow rate (Knutson and Whitby, 1975). 149 Instead of employing the standard TSI Aerosol Instrument Manager control software, the DMA was governed by 150 151 custom-made software, enabling the selection of specific diameters and the control of residence time at those 152 diameters (Moore et al., 2010; Deng et al., 2011). A total of 29 mobility diameters ranging from 29 nm to 612 nm, at logarithmic intervals, were chosen, and each diameter was maintained for 30 seconds. This time was selected 153 154 in consideration of the time lag (estimated to be less than 20s) between the instruments following the DMA (Fig. 155 1a). A complete size scan took approximately 15 minutes. The DMA's sheath and sample flow rates were 156 calibrated both before and after the experiment, and the mobility size selected by the DMA was validated using PSLs. 157

158 The monodisperse number concentration of all condensation nuclei ([CN]), [CCN], and both [BCc] and [BCf] was

159 measured by a Condensation Particles Counter (CPC, Model 3772, TSI Inc., USA) with a flowrate of 0.30 L/min,

and a Cloud Condensation Nucleus counter (CCNc, DMT Inc., USA) with a flowrate of 0.50 L/min, and the other

161 Single Particle Soot Photometer (SP2, DMT Inc., USA) with a flowrate of 0.12 L/min, respectively (Fig. 1a).





Three SSs at 0.1%, 0.2%, 0.3% were set for the CCNc, and with an interval of 30 minutes, allowing two size scans for each SS. The sample/sheath flowrate and also the SS of the CCNc were calibrated before and after the experiment (Rose et al., 2008). Moreover, the monodisperse [*CCN*], [*CN*], and [*BCc*] were corrected for multiple charge effects, which performed from large particles (only have one negative charge) to smaller particles according to charging probabilities for different mobility size (Wiedensohler, 1988).

167 The CCN activity of all-particles was determined based on the size resolved activation ratio (F_{act}) ([CCN]/[CN])

- 168 follow the method by Moore et al. (2010). This size resolved ratio, [CCN]/[CN], reflect how many parts of CN
- 169 could serve as CCN can be directly measured by the CCNc and CPC with size scanning by the DMA. [CCN]/[CN]
- 170 increase with the particle size, the ratio equals 1 at larger particle sizes (e.g., greater than 300 nm), whereas it
- 171 reduces to 0 at smaller sizes (e.g., 30 nm). The specific size at which the [CCN]/[CN] ratio reaches 0.5 is
- 172 identified as the activation diameter (D50). This critical value is generally derived through sigmoid fitting of the
- 173 size-resolved [CCN]/[CN] (Rose et al., 2008).

174 The CCN activated number concentration of BCc $[BC_{act}]$ is discriminated from BCf and directly measured in this 175 study follow the method by Hu et al. (2021a), which compare the difference between the un-activated particle 176 number concentration ($[CN_{unact}]$) and [BCc] to derived how many BCc are activate into CCN. $[CN_{unact}]$ is derived 177 by subtracting the [CCN] from the total [CN]. $[CN_{unact}]$ is zero at larger particle sizes, indicating that all particles, 178 including BCc, are activated. Conversely, when $[CN_{unact}]$ equals the total [CN], it signifies that none of the 179 particles are activated. It is well recognized activation of BCc into CCN is more challenging compare to that of 180 BCf at the same diameter, due to BCc's lower κ (Dusek et al., 2006; Bond et al., 2013). Only when the $[CN_{unact}]$ is

smaller than [BCc], it means some BCc have been activated into CCN, while $[CN_{unacc}]$ is zero, it means all BCc

are activated. So, the $[BC_{act}]$ and activation ratios $(F_{act,BC})$ of BCc are calculated as:

183
$$[BC_{act}] = [BCc] - ([CN] - [CCN])$$
(4)

184
$$F_{act,BC} = ([BC_{act}]/[BCc])$$
(5)

185 $[BC_{act}]$ is the difference between the [BCc] and $[CN_{unact}]$, and the $F_{a,BC}$ is calculated by $[BC_{act}]$ over the [BCc]. 186 When $[CN_{unact}]$ is larger than [BCc] means none of BC could activate, the $[BC_{act}]$ and $F_{act,BC}$ is 0.

187 **2.4 Determination of the Photochemical Age**

188 The PTR-TOF-AMS mass concentration measurement was calibrated by the CH3COCH4+ (m/z 59.0490) and 189 H3O+ (m/z 21.0226) following the procedures described by Bruns et al. (2016). The equivalent photochemical 190 age (t_{age}), reflected the OH exposure dose of the air masses, was calculated as the change of toluene over benzene 191 concentration ratio due to their different reaction coefficient with OH, and expressed as:

192
$$t_{age} = \frac{1}{[OH](k_{toluene} - k_{benzene})} \times \left[\ln \left(\frac{[toluene]}{[benzene]} \right)_0 - \ln \left(\frac{[toluene]}{[benzene]} \right) \right]$$
(6)

where [OH] is the mean ambient OH concentration at 2.5×10^6 molecule cm⁻³ during summer daytime in this region (Yuan et al., 2012). $k_{toluene}$ and $k_{benzene}$ is the reaction rate with OH of toluene (5.63×10^{-12} cm⁻³ molecule⁻¹ s⁻¹) and benzene (1.22×10^{-12} cm⁻³ molecule⁻¹ s⁻¹), respectively. ($\frac{[toluene]}{[benzene]}$)₀ is the ratio of freshly emitted toluene and benzene which is assumed to be 3.3, and $\frac{[toluene]}{[benzene]}$ is the ratio of measured toluene and here are also a structure of the ratio of formed for the destine to measure toluene and

197 benzene in the atmosphere. The calculation of t_{age} was only performed for the daytime to warrant it indicates the





198 photochemical process. The dataset of t_{age} in this study was already test successfully to obtain the reaction or 199 production rates of secondary OA and VOCs (Wu et al., 2022).

200 2.5. Backward Trajectory Analysis

The potential source contribution to the target site was modelled by using Hybrid Single-Particle Lagrangian Integrated Trajectory Model (HYSPLIT) ensemble calculation with 3-hourly, 1°×1° GDAS horizontal and vertical reanalysis wind field to retrieve backward 36-hour air mass to the target site, and air mass clusters came from four different directions were classified during the entire observation periods (Fig. 1b). Cluster 1 (C1) was from the southern polluted area, while C2-C4 were from the northern clean areas (Fig. 1b).

206 3. RESULTS AND DISCUSSION

207 3.1 Overview of the Aerosol Physical and Chemical Properties

208 Figure 2a-2d shows the time series of aerosol physical and chemical properties with the colour labels at the top represent different air masses influence periods, and the statistical comparison of aerosol properties between 209 210 different air masses were show in Fig. 2e-2i. Consistent with previous studies, the aerosol physic-chemical 211 properties were highly dependent on air masses (Tian et al., 2020; Zhao et al., 2020). Higher pollutant 212 concentrations including PM₁ Sulphate, Nitrate, OA, and BC were observed during C1 period with air from 213 southern part of North China Plain (NCP) (average PM₁ value of $33.2 \pm 16.4 \,\mu g \, m^{-3}$) (Fig. 2e and Fig. 2f), where contain high pollutant concentration due to the intensive emissions (Fig. 1b). When the air come from northern 214 215 clean part of NCP during C3 and C4 period (Fig. 1b), lower pollutant concentrations were observed (average PM₁ 216 value of $10.4 \pm 5.8 \ \mu g \ m^{-3}$) (Fig. 2a and Fig. 2e), which suggest pollutants were diluted and dispersed during the 217 fast transport process. Thus, C1 was estimated as polluted condition, while C3 and C4 was classified as clean 218 environment. However, C2 came from the same direction as C4 but with less distances of transportation, which 219 means smaller wind speeds to dilute the emissions, so the PM₁ mass concentration $(15.4 \pm 10.6 \,\mu g \,m^{-3})$ of C2 was higher than clean environment, but still one half lower than that of C1, and classified as light pollution condition 220 221 (Fig. 2e).

222 During the transition from clean to pollution condition, the mass concentrations of all chemical components 223 experienced substantial increase, as shown in Fig. 2a. Notably, the increase in inorganic aerosols (including 224 NH₄HSO₄, (NH₄)₂SO₄, and NH₄NO₃) was more pronounced than that of OA (Fig. 2b), which lead to lower/higher 225 contribution of OA to all-particles during polluted/clean period (Fig. 2b and Fig. 2e). As κ_{all} is determined based 226 on the volume fraction of OA and inorganic aerosol according to the ZSR internal mixing rule, this result distinct 227 κ_{all} characteristic for different airmasses influence periods (Fig. 2d and Fig. 2h).

228 **3.2 The Hygroscopicity of BCc and All-particles**

229 The κ_{all} (particles average hygroscopicity) varied from 0.13 to 0.45 with an average value of 0.29 during 230 observation (Fig. 2d), the κ_{all} during C1 polluted period (average value of 0.38) was 1.5 times larger than that 231 during C3/C4 clean period (average value 0.25) (Fig. 2h), this was attributed to the increased contribution of 232 inorganic aerosol, known for their high hygroscopicity, under polluted condition.

233 The κ_{BC} is influenced by a combination of κ_{all} and the coating thickness (reflected by V_{coating}/V_{BC}) (Eq. 3). The

234 $V_{\text{coating}}/V_{BC}$ of BCc during C1 polluted period was 4.9±4.2, which were about 3 - 5 times larger than that during

235 C3/C4 clean period (Fig. 3c and Fig. 3g), implying thick coating of BCc under polluted condition. Due to both

236 increase of $V_{\text{coating}}/V_{\text{BC}}$ and κ_{all} , the κ_{BC} show similar variation characteristic as κ_{all} also with higher value under

polluted condition (average value of 0.23), while lower value under clean condition (average value of 0.11) (Fig.

238 2d and Fig. 2h).





239 $V_{\text{coating}}/V_{\text{BC}}$ exhibits a larger variation, ranging from 1.3 to 20, compared to κ_{all} , indicating that $V_{\text{coating}}/V_{\text{BC}}$ has a 240 substantial influence on κ_{BC} (Fig. 2d). This can be evidenced by the maximum value of κ_{BC} correlates with that of 241 $V_{\text{coating}}/V_{BC}$ (Fig. 2c and Fig. 2d), and the divergence of the maximum values of κ_{all} and κ_{BC} (Fig. 2d). The coating of BCc is formed through by the photo-chemistry of VOCs precursors during the atmosphere aging process, 242 243 which is determined by VOCs precursors and also aging time. Comparing to the clean clusters C3 and C4, the air 244 mass movement of C1 was much slower, combined with passing the highest pollution level region, the slowly 245 transported air mass from the polluted southern region was characterized with a longer transport time than other 246 air mass clusters (Fig. 1b). As expected, the observation of PTR-ToF-MS shows higher VOCs values and also the 247 longer t_{age} during C1 polluted condition (Fig. S2). Figure 2i show that the t_{age} during C1 polluted period (13h) was about one times larger than that during C3/C4 clean period (9h), this reflected the air during polluted period was 248 249 more aged.

Figure 3 shows diurnal variation of PM₁, BC mass concentration, OA fraction, $V_{\text{coating}}/V_{BC}$, κ_{all} and κ_{BC} for the 250 251 different air masses respectively. More pronounced diurnal patterns were found during the C3/C4 clean period, as 252 the fast-moving clean air masses dilute the pollutant and were mainly influenced by local emissions, however, regional transport and secondary aerosol formation influence the diurnal pattern during C1 and C2 period (Fig. 253 254 2a). For chemical inert BC particles, clear BC mass peaking at morning and afternoon rush-hour (marked as grey 255 bar) (Fig. 3b) was observed which is typical time period of heavy traffic with nascent BC emitted, but C1 period showed a continuous transport and increased BC mass loading throughout the midday, even when developed 256 257 planetary boundary layer (PBL) dilute the pollutants.

Fresher BCc showed lowest $V_{coating}/V_{BC}$ during the morning rush-hours, and within a few hours the $V_{coating}/V_{BC}$ quickly increase and peak at noon when solar radiation was strongest for all-time period (Fig. 3d). The increase during midday was in consistent with the OA fraction (Fig. 3c), which suggested that the acquire coating of BC particles might be secondary OA through the photochemical oxidation, inconsistent with recently studies (Cui et al., 2022; Wu et al., 2019). As enhanced OA fraction at noon due to photochemical SOA formation (Fig. 3c), result a lower value of κ_{all} during noontime (Fig. 3e). κ_{BC} is more influenced by the $V_{coating}/V_{BC}$, the κ_{BC} was found increase after morning rush-hours due to the increase of coating on BC.

265 3.3. The CCN Activation Properties of All-particles and BCc

266 **3.3.1 Determining the D50 of BCc**

267 The measurement of the activation properties of all-particles and BCc into CCN was discussed in Sec. 2.3.2. 268 Figure 4 illustrates a representative example of size-resolved activation of all-particle and BCc at SS = 0.1% and 269 0.3% respectively. Activation commences for all-particles when [CCN] is greater than 0 (depicted by blue lines), 270 while BCc (indicated by black lines) begins to activate when the [CNunact] (represent by green lines) falls below that of [BCc] (Fig. 4). $[BC_{acl}]$ (denoted by the grey shadow) can be derived using Equation 4, and the size-271 272 resolved $F_{act,BC}$ can be calculated in accordance with Equation 5. It can be seen that with the increase of particle size, BCc started to get activated at 0.19µm (0.13µm) at SS=0.1% (0.3%), which is higher than all-particle by 44% 273 274 (46%), the higher D50 of BCc also indicates that the lower κ_{BC} than κ_{all} .

We would like to emphasize that the principal uncertainty in the method employed stems from the assumption regarding the lower κ for BC compared to other particles. The assumption was consistent with concurrent theoretical understanding, pure BC was hydrophobic ($\kappa_{BC}=0$), though the κ_{BC} increase by acquire hygroscopic coating through atmospheric aging (Grimonprez et al., 2018; Weingartner et al., 1997), the κ_{BC} remains lower than other components under the same atmospheric oxidation process based on ZSR volume mixing rule. Most importantly, even though some internally mixing BCc activate into CCN previous to other chemical components, the results here was underestimated the [BC_{act}] and $F_{act,BC}$, so the CCN activity of BCc derived through this

²⁸² method could represent the low bound of the atmospheric BCc.





283 3.3.2 D50 of BCc and All-particles

284 Figure 5 illustrate the temporal evolution of size distribution of BCc, CN and CCN, respectively. The [CCN] 285 increase during polluted condition due to the increase BCc and all-particles concentration, and the [CCN] is 286 mainly contributed by all-particles (Fig. 5c). The derived D50 for BCc and all-particles under water SS at 0.1%, 287 0.2%, and 0.3% is shown as lines in Fig. 5a-c for the temporal evolution, and mean value is shown as vertical 288 dash lines in Fig. 5d. The mean D50 of BCc particles (0.17µm, 0.13µm, and 0.12µm under SS= 0.1%, 0.2%, and 0.3%) was approximate 1.4 times larger than that of all-particles (0.11µm, 0.08µm, and 0.07µm under SS= 0.1%, 289 290 0.2%, and 0.3%) (Fig. 5), which was expected from κ -Köhler theory (Petters and Kreidenweis, 2007) that lower/higher D50 means the higher/lower κ . The higher D50 of BCc also reflected lower κ and BCc should be 291 activated in the last order. 292

293 Large variation of D50 of BCc and all-particles was also found for different air massed (Fig. 5), the D50 of all-294 particles $(0.13\mu m \pm 0.01\mu m)$ and BCc $(0.20\mu m \pm 0.01\mu m)$ during C1 polluted period was lower that $(0.15\mu m \pm 295 \ 0.02\mu m)$ and $(0.23\mu m \pm 0.02\mu m)$ during C3/C4 clean period. This is also inconsistent with the variation of κ_{all} 296 and κ_{BC} (discussed in 3.2), suggested the more inorganic fraction increase κ_{all} and the thick coating increase the 297 κ_{BC} during the polluted condition.

Figure 6a and 6b illustrate the diurnal variation of D50 of all-particles and BCc under SS=0.10%, the other two SSs at 0.20% and 0.30% are shown in Fig. S3 and Fig. S4. The diurnal variation was more pronounce under clean period (i.e. C3+C4), when local emission dominate the pollution sources. Higher D50 of all-particles and BCc was found during the rush-hours when more primary hydrophobic OA and fresh external mixed BCc emitted. The D50 of all-particles and BCc will decrease after photochemistry oxidation secondary OA formation and coating on BCc which was also in consistent with that of κ_{all} and κ_{BC} (Fig. 3).

304 3.3.3 Droplet Activation Fraction of BCc and All-particles

Figure 5d shows the averaged size distribution of BCc, CN and CCN under three SSs for different periods classified by air mass clusters (C1-C4). Notably BCc had consistently larger peak size than all-particles for all air masses. For example, BCc peak from 185 to 210nm, which was about 2-4 times larger than that of all-particles, which peaked from 58 to 115nm. The more polluted air masses (C1 and C2) had larger peak size of BCc and CN than cleaner air masses (C3 and C4), because of the condensation process during the polluted episode when gases heavily partitioned to aerosol phase and enlarged particle size.

Though the BC core size is smaller, after attaching on larger particles the overall coated size is even significantly larger than BCf. This means in the typical sub-urban environment, BC had been significantly associated with non-BC substance through condensation and coagulation process (Yu et al., 2022), and the bare BC without containing non-BC substance had taken a less proportion and most BCc had been enlarged during the aging process. This is consistent with previous studies that BC was associated with larger aerodynamic size than BCf (Pan et al., 2019; Wang et al., 2021). Previous studies also observed rapid aging of BC in a few hours in urban environment (Peng et al., 2016).

318 Figure 6d and 6e show the diurnal variations F_{act} (in number and mass) for all-particles and BCc under SS=0.1%.

319 The other two SSs at 0.2% and 0.3% are shown in Figure S2 and Figure S3. With SS increase from 0.1% to 0.2%,

- 320 the number $F_{act,BC}$ and $F_{act,all}$ increased from 42%±15% to 69%±15% and from 30%±15% to 51%±19%. Figure 6f
- 321 show the statistical F_{act} results compare for different air masses. the number $F_{act,BC}$ increased from 30% to 53%
- 322 under SS=0.1% (Fig. 6f) with the pollution increase from C3/C4 to C1. The mass $F_{act,BC}$ show similar variation
- 323 characteristic as the number $F_{act,BC}$, which increase from 58 % to 76 % under SS=0.1% from C3/C4 to C1 period
- 324 (Fig. 6f), and from 81% to 89% and from 87% to 91%, under SS at 0.2% and 0.3%, respectively. BCc acquires
- 325 thicker hygroscopic coating through longer atmospheric aging time during the long range regional transport C1
- 326 polluted period, the hygroscopic coating not only increase κ_{BC} and lead to a lower D50, also increasing the BCc





327 size. The decrease of D50 and increase of particle size both led to an increase of F_{act} . The $F_{act,BC}$ mainly depended 328 on the V_{coating}/V_{BC}, as the D50 and size of BCc was dependent on V_{coating}/V_{BC}. The diurnal variation also show 329 lowest $F_{act,BC}$ during rush hours due to lower V_{coating}/V_{BC}, and increased in the daytime with increased V_{coating}/V_{BC} 330 (Fig. 6e).

The number $F_{act,BC}$ was always higher than $F_{act,all}$ under all SSs (Fig. 6, Fig. S2 and Fig. S3), due to the size of BCc is larger than all-particles. The mass $F_{act,BC}$ was about 1.61 higher than in number, compared with about 2.88 difference of mass and number F_{act} for all-particles. This is also caused by the larger size distribution of BCc than all-particle. The large portion of all-particles in small size drives the number F_{act} to be lower than BCc. Although D50 of BC is higher, the larger size distribution of BCc than all-particles means BCc can have a higher droplet removing efficiency than BCf.

The large F_{act,BC} than BCf particles, and thickly coated BCc have preference in CCN activation was also observed 337 338 through cloud residues observation at a mountain site in south China (Fu et al., 2022). The measured F_{act} was 339 generally consistent with experiment conducted in a South China mega city using the same method (Hu et al., 340 2021a), but are higher than cloud residues observation in China (Zhang et al., 2017), which reported an average 341 value 33% of the mass $F_{act,BC}$ and number F_{act} of BCc (all-particle) ranged in 5%-45% (7%-60%). We attribute 342 the difference to unknow ambient SS for the cloud, which have great influence on the Fact, BC. According to 343 previous cloud residues observation at Jungfraujoch, the number $F_{act,BC}$ increase from 50% to 80% when SS 344 increase from 0.21% to 0.50% (Motos et al., 2019a; Hammer et al., 2014). The ambient SS show large variation can be range from 0.01% to 2.00% (Hammer et al., 2014), however, it is hardly to direct observe the SS from the 345 346 field observation. Through this method the number and mass $F_{act,BC}$ activated into droplets can be quantitatively 347 investigate under certain ambient SS condition, thus the results here provide a reference to investigate the lifetime 348 of BCc.

349 3.4. Time Scale for the Evolution of CCN Activity of BCc

350 The temporal evolution of parameters related to CCN activity including κ , D50 and F_{act} against photochemical 351 age (t_{age}) for BCc and all-particles are shown in Fig. 7. It shows that the aging time (t_{ag}) can well explain the 352 variation of all parameters for BCc and all-particles with a linear function (\mathbb{R}^{2} >0.4).

The volume ratio of coating over rBC ($V_{\text{coating}}/V_{BC}$) showed an increase rate of 0.14 h⁻¹ (Fig. S4), and can be fitted as:

355
$$V_{coatina}/V_{BC} = 0.55 + 0.14 \cdot t_{aae}$$
 (7)

This results in a linear increase rate for the hygroscopicity parameter, κ_{BC} showed a linear rate of 0.09 h⁻¹ ((Fig. 7a, 7b, and 7c)) at all SSs, can be fitted as:

358
$$\kappa_{BC} = 0.09 + 0.01 \cdot t_{age}$$
 (8)

359 In addition, κ_{BC} also follow a linear function with t_{age} as:

360
$$\kappa_{all} = 0.16 + 0.014 \cdot t_{age}$$
 (9)

361 This means the κ for the large proportion of non-BC is enhanced after aging due to enhanced formation of 362 inorganic salt and secondary organic aerosols (Wu et al., 2022). The results show it will take about 24 and 10 363 hours to reach the global mean κ of 0.27-0.30 (Pringle et al., 2010) for BCc and all particles, respectively.

364 The aging scale of a few hours observed here is generally consistent with previously measured BC aging 365 influenced by anthropogenic pollution, such as an aircraft measurement (Moteki et al., 2012) which observed the 366 conversion time scale of BC was a few hours under polluted environment. The results were also comparable with 367 previous chamber studies which investigated the κ evolution of BC (Tritscher et al., 2011; Lambe et al., 2015;





Peng et al., 2017), the increase rate of κ_{BC} (0.01h⁻¹) in this study was similar, but the initial κ_{BC} here (0.08-0.26) was higher than the fresh combustion soot as used in their results (0-0.12). These consistently showed a few hours tended to be a typical aging time scale to convert appreciable amount of initially hydrophobic BC to be hydrophilic, hereby CCN active. The exact time scale may according to pollution level, such as a higher concentration of gas precursors and favoured meteorological condition may promote the growth rate of hygroscopic substances on BC.

374 Correspondingly, the D50 of BCc particles and all-particles decreased with t_{age} , with a rate of -3 nm h⁻¹ and -3nm 375 h⁻¹(Fig. 7d), -1 nm h⁻¹ and -2nm h⁻¹ (Fig. 7e), and -0.5 nm h⁻¹ and -2 nm h⁻¹ (Fig. 7f) under SS = 0.10%, 0.20%, 376 and 0.30%, respectively. The results are summarized in Table 1.

377 The number activation fraction $F_{act,all}$ and $F_{act,BC}$ also show linear corelation with t_{age} . A notable higher CCN

activation fraction was observed for BCc than all particles at the same t_{age} , this was attributed to the larger coated

379 size of BCc than other non-BC particles (Fig. 5). This difference between BCc and other particles was reduced at

380 longer t_{age} because the sufficient aging time had allowed all particles to grow to similar large sizes.

381 After 8 hours of aging, $F_{act,BC}$ increased from 35% to 63% as SS rose from 0.1% to 0.2%, and then increase from

382 63% to 77% as SS increase from 0.2% to 0.3%. This illustrates a notable increase in $F_{act,BC}$ from SS=0.1% to

383 0.20%, but a less significantly changed $F_{act,BC}$ from SS=0.2% to 0.3%. Furthermore, the rate increase of $F_{act,BC}$

384 with t_{age} was 2% h⁻¹ at SS=0.10% (Fig.7g), which was twice as large as at SS=0.2% (1% h⁻¹) (Fig. 7h) and 0.3%

 $(1\% h^{-1})$ (Fig. 7i). This indicated that the activation of BCc was more sensitive at lower SS (<0.2%), but may

reach a plateau when SS>0.2%. This is because the activation diameter at this typical SS range was close to the

387 mode size of BCc number distribution (Fig. 5d), and a small increase of SS will cause a considerable change of

388 fraction which is larger than the activation diameter. The SS range from 0.1% to 0.2% is for the typical formation

389 of fog and stratiform thin clouds (Hammer et al., 2014).

390 4. Conclusion

In this study, we conduce directly measurements of the droplet activation of BC-containing particles under water 391 392 supersaturation conditions in a sub-urban environment. Both the hygroscopicity and CCN activation diameters 393 reveal that the hygroscopicity of all-particles and BCc concurrently increase with polluted levels, primarily 394 governed by the amounts of coating. As pollution intensifies, the volume ratio of coating over rBC increases from 395 1.8 to 4.7, causing the hygroscopicity of BCc to increases from 0.11 to 0.23. The increase of coating not only increase BCc's hygroscopicity but also shifts these particles to larger sizes, significantly enhance their CCN 396 397 activity. The number activated fraction of BCc was 42% under SS at 0.1%, higher than the 30% activation of all 398 particles. Despite the lower hygroscopicity, the elevated activation fraction of BCc can be attributed to their 399 larger size compared to all particles. It was also found that the number activated fraction of BCc was more sensitive in the SS range from 0.1% to 0.2%, a typical range for fog and stratiform thin clouds formation. This 400 401 means a small increase of SS even for the environment with relatively low moisture and weak updraft the BC can 402 be activated. Considering these clouds have a lower precipitation rate, BC in these clouds may have more 403 persistent indirect radiative impacts.

Furthermore, we successfully parameterized the evolution of hygroscopicity and CCN activation characteristics of BCc and all particles in relation to equivalent photochemical age. The hygroscopicity of BCc was found positive linearly correlated with aging time with a rate of 0.01 h^{-1} , and the activation diameter negative correlated with a rate of 1-3 nm h⁻¹. Within 10 hours, 65% of BCc activated under SS=0.2% in this typical sub-urban

408 environment, this is in line with a wide range of recent ambient studies that BC can be half activated under





- 409 SS=0.1-0.2%. This means these BCc may be efficiently incorporated into clouds by serving as CCN. This aging 410 process by acquiring more hygroscopic substances of BC may be particularly important under polluted 411 environment, where the condensation and coagulation processes intensively occurred. The results here thus 412 provide direct evidence of efficient droplet activation of BCc, and the aging scale associated with their CCN
- 413 activation can be used as a refere to validate the regional models regarding aerosol-cloud interactions of BC.
- 414 **Data availability.** All data in this paper are available from the authors upon request (tianping@bj.cma.gov.cn).
- 415 **Competing interests.** At least one of the (co-)authors is a member of the editorial board of Atmospheric 416 Chemistry and Physics.
- 417 Author contribution. D. D., and M. H. led and designed the study; P. T. and D. L. designed the study, set up the 418 experiment, analysed the data, and wrote the paper. K. H. and Y. W. conducted the field observation and 419 contributed the data analysis. H. H. and J. S performed the calibration of the instruments. C. Y. and D. H 420 contributed to scientific discussion.
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- 638 Table 1. A summary of linear fitting results for activation diameter (D50) and activation fraction (F_{act}) of all
- 639 particles and BCc-containing particles.

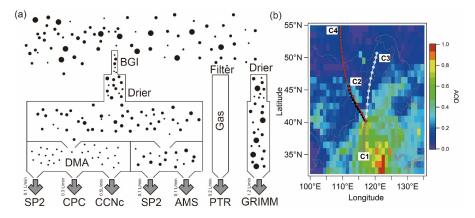
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Initial, Slope	Activation diameter (D50)		Activation fraction (F_{act})	
	All particles	BC-containing particles	All particles	BC-containing particles
SS=0.1%	170nm, -3nm h ⁻¹	240nm, -3nm h ⁻¹	4%, 2% h ⁻¹	19%, 2% h ⁻¹
SS=0.2%	120nm, -2nm h ⁻¹	170nm, -1nm h ⁻¹	18%, 0.3%h ⁻¹	55%, 0.1% h ⁻¹
SS=0.3%	100nm, -2nm h ⁻¹	150nm, -0.5nm h ⁻¹	26%, 0.3% h ¹	69%, 0.1% h ⁻¹





642 **Figures and captions**

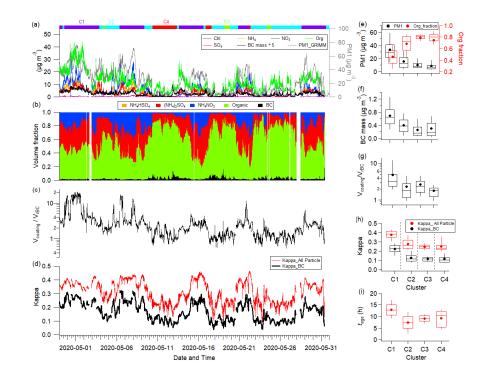


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Figure 1. Schematic of the experiment. a) Experimental setup showing aerosols are sampled through a $PM_{2.5}$ impactor following a drier. Mono-dispersed aerosol was measured downstream of a differential mobility analyser (DMA), including SP2, CPC and CCNc, and additional line measure the poly-dispersed aerosols with SP2 and HR-ToF-AMS. All sizes of dried aerosol are measured by a GRIMM, and the VOC concentration was measured by a PTR-TOF-AMS. b) Mean aerosol optical depth (AOD) during the experimental month, with three lines showing the clustered 36h backward trajectories initialized from the experimental site with each markers denoting 6h transport time.





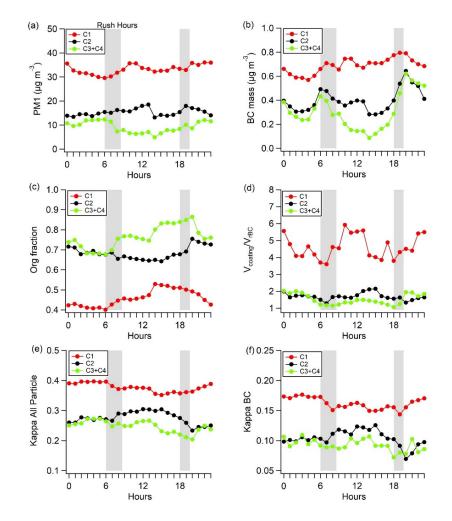


653 Figure 2. Time series of a) mass concentrations for chloride (Chl), ammonium (NH₄), nitrate (NO₃), organic 654 (Org), sulfate (SO²₄), BC mass and PM₁ measured by HR-TOF-AMS, SP2 and GRIMM. b) Volume fraction of 655 particle components for NH₄HSO₄, (NH₄)₂SO₄, NH₄NO₃, Organic and BC, c) $V_{\text{coating}}/V_{BC}$, d) κ for all particle and BC respectively. The flag on the top of the graph represents different clusters by different colours. Mass 656 657 concentrations of e) Organic fraction and PM1 by summing HR-ToF-AMS components and BC mass from SP2, f) BC mass, g) $V_{\text{coating}}/V_{BC}$, h) κ , i) t_{age} for all particle (red) and BC (black) for four clusters, with solid circle 658 indicating the average; in each box, the solid line and top and bottom boundaries represent the median, 75th and 659 660 25th percentiles, respectively.

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Figure 3. Diurnal variations of a) PM₁, b) BC mass, c) Organic fraction, d) $V_{\text{coating}}/V_{BC}$, e) κ for all particles and f) κ for BC for four clusters, respectively. Grey bars indicate the traffic rush hours, and the red, black, and green line represent the C1, C2 and C3/C4 period, respectively.





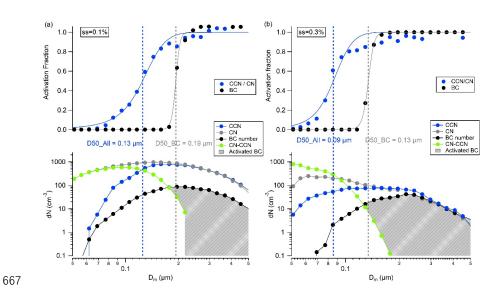
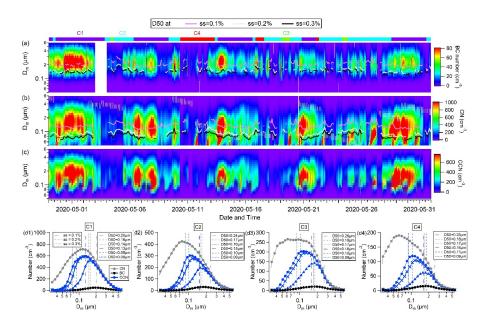


Figure 4. A typical example of size-resolved number concentrations of CN, CCN and BCc under supersaturation of a) 0.1% and b) 0.3%. The green lines show the un-activated number concentration and the grey shadows indicate the activated BCc number concentrations at each size. The blue and black dots in the upper panel show the activation fraction of all-particle and BCc. The blue and grey lines indicate the sigmoid fitting on the sizeresolved activation fraction of all-particle and BC respectively, with the vertical dash lines indicating the respective activation diameters.







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Figure 5. Temporal evolution of numbers of a) BCc, b) CN, c) CCN and activation diameter (D50) of BCc and CN particles under SS=0.1%, 0.2% and 0.3%. d) Mobility particles size distribution of CN, BCc and CCN numbers under SS=0.1%, 0.2% and 0.3% for different clusters. The blue and grey dash lines on the images denote the D50 for all-particles and BCc under different SS, respectively. Bottom panels show the mean particles size distributions of CN, CCN (under the three SSs) and BCc, for the classified air masses.





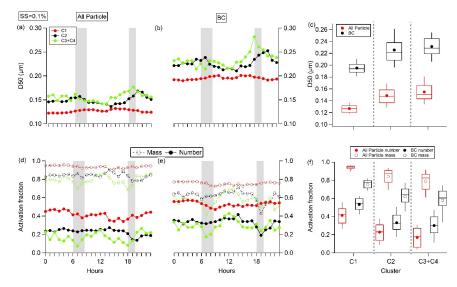


Figure 6. Diurnal variations of D50 of a) all-particles, and b) BCc particles under SS=0.1% for different air mass clusters, respectively, c) the statistical of D50 for different air mass cluster, with solid circle indicating the average; in each box, the solid line and top and bottom boundaries represent the median, 75th and 25th, respectively. c) diurnal variation of the activation fraction of all-particles, and e) BCc particles, with the solid and dash line represent the number and mass fraction, the mass (number) activation fraction was calculated by summing the mass (number) of particles with diameter larger than D50, and divided by the total mass (number). f) statistical of the activation fraction.

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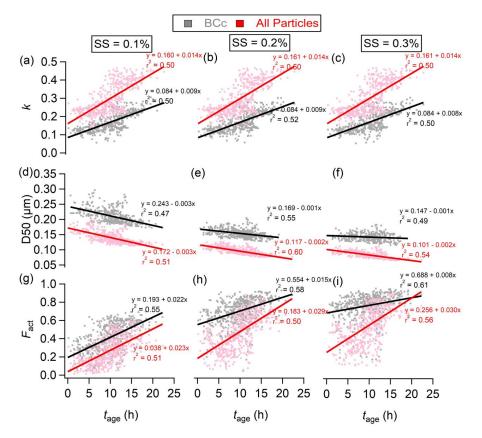


Figure 7. The evolution of hygroscopicity parameter (κ) (a-c), activation diameter (D50) (d-f), and number activation fraction (F_{act}) (g-i) with photochemical age (t_{age}) under SS at 0.1%, 0.2%, c) 0.3%. The black and red denote the BCc particles and all-particles, respectively. The lines show the least-square lineal regression.