



1 **Aggravated surface O₃ pollution primarily driven by meteorological variation**
2 **in China during the early COVID-19 pandemic lockdown period**

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28 **Abstract**

29 Due to the lockdown during the COVID-19 pandemic in China from late January to early April in
30 2020, a significant reduction of primary air pollutants has been identified by satellite and ground
31 observations. However, this reduction is in contrast with the increase of surface O₃ concentration
32 in many parts of China during the same period. The reasons for this contrast are studied here from
33 two perspectives: emission changes and inter-annual meteorological variations. Based on top-
34 down constraints of NO_x emissions from TROPOMI measurements and GEOS-Chem model
35 simulations, our analysis reveals that NO_x and volatile organic compound (VOC) emission
36 reductions as well as meteorological variations lead to 8%, -3%, and 1% changes in O₃ over North
37 China, respectively. In South China, however, we find that meteorological variations cause ~30%
38 increases in O₃, which is much larger than -1% and 2% changes due to VOC and NO_x emission
39 reductions, respectively, and the overall O₃ increase is consistent with the surface observations.
40 The higher temperature is the main reason that leads to the surface O₃ increase in South China.
41 Overall, inter-annual meteorological variations have a larger impact than emission reductions on
42 the aggravated surface O₃ pollution in China during the early lockdown period of COVID-19
43 pandemic.

44

45 **1. Introduction**

46 Surface ozone (O₃), an important air pollutant that is harmful to human health (Jerrett et
47 al., 2009) and stomatal conductance of green vegetations (Gong et al., 2020), is produced by
48 photochemical reactions of nitrogen oxides (NO_x) and volatile organic compounds (VOC) (Liu et
49 al., 1987; Sillman et al., 1990). In addition to emissions, meteorological conditions, such as
50 temperature, solar radiation and relative humidity, also have large impacts on surface O₃ formation
51 (Lu et al., 2019).

52 Ground observations show that surface O₃ increased dramatically during the COVID-19
53 lockdown period in China by around 40% on average (Tong et al., 2023) and even larger than 100%
54 (Shi and Brasseur, 2020; Liu et al., 2021) depending on the time period and region. The reduction
55 of economic activities during the lockdown period led to a significant decrease of several primary
56 air pollutants emissions. The NO₂ vertical column density (VCD) from satellite measurements and
57 surface NO₂ concentration from ground measurements were reduced by 40% - 60% in China
58 during the lockdown period (Bauwens et al., 2020; Shi and Brasseur, 2020; Liu et al., 2020a; Zhang



59 et al., 2020). A lower but discernible reduction of sulfur dioxide (SO₂), carbon monoxide (CO),
60 and formaldehyde (HCHO) have also been identified by satellite or ground-based observations in
61 China (Shi and Brasseur, 2020; Levelt et al., 2022; Ghahremanloo et al., 2021). However, during
62 this period surface O₃ concentrations increased, and the respective roles of meteorological factor
63 and emission reduction for the aggravated surface O₃ pollution during the lockdown in China need
64 to be further quantified.

65 This study provides a quantitative analysis of the causes for the unexpectedly aggravated
66 surface O₃ pollution in China during the early lockdown period of the pandemic from two
67 perspectives using GEOS-Chem model. One is anthropogenic emission reduction of NO_x and
68 VOC in response to the lockdown possibly under a VOC-limiting chemical regime of surface O₃
69 production (Guo et al., 2023), while the other is the impact of natural variability of meteorological
70 conditions. Previous studies have reported the enhanced surface O₃ due to NO_x emission decline
71 during the lockdown period in North China using chemical transport model (CTM) simulations
72 without controlling for the impacts of meteorological variability (Zhang et al., 2021; Huang et al.,
73 2020; Miyazaki et al., 2020). Other studies quantified or excluded the meteorological impacts on
74 the surface O₃ using statistical analysis instead of CTM that account for the physical and chemical
75 processes (Venter et al., 2020; Bi et al., 2022; Tong et al., 2023). Although a few studies have
76 investigated the contributions from both emission reduction and meteorological variability to
77 surface O₃ increase using CTMs, most of their results have uncertainties due to the limitations of
78 their analysis. For example, some of them keep the emissions unchanged (Zhao et al., 2020) or
79 assume an arbitrarily uniform emission reduction instead of constraining the emission based on
80 observations (Le et al., 2020; Liu et al., 2021). In cases where the emissions were constrained by
81 the observations, the focus was limited to several cities in China (Liu et al., 2020b). Furthermore,
82 in the past studies, the surface O₃ increase during the lockdown period of 2-4 weeks is quantified
83 in reference to the time period right before the lockdown instead of the same period in previous
84 years; such comparisons by design cannot exclude the possibility that the seasonal variation of
85 meteorology from early January to early April may have dominated the cause for the surface O₃
86 increase. A comprehensive analysis of the contributions from emission reductions and
87 meteorological variations to the surface O₃ increase during the first round of the lockdown period
88 with respect to the same time period in previous years in China is therefore overdue.



89 Here, we apply a top-down method to update NO_x and VOC emission in February and
90 March in 2020 based on the TROPOMI NO₂ and formaldehyde (HCHO) product. A set of GEOS-
91 Chem model simulations with NO_x and VOC emissions and meteorological fields in different time
92 periods are then conducted. Based on the difference in surface O₃ concentration in different
93 modeling sensitivity experiments, we quantitatively assess the respective roles of emission and
94 meteorology in regulating surface O₃ concentration in continental China. The ground observations
95 of surface O₃ and NO₂ concentration are compared with the model simulations to verify our
96 analysis. Section 2 introduces the satellite and ground-based measurements, NO_x emission update
97 scheme, and the configurations of GEOS-Chem simulation experiments. Section 3 provides an
98 evaluation of the constrained NO_x emission and surface O₃ simulations. The analysis of the
99 mechanism of the aggravated surface O₃ pollution is presented in Section 4, followed by the
100 summary and conclusions in Section 5.

101

102 **2. Datasets and Methods**

103 **2.1 TROPOMI NO₂ and HCHO product**

104 We used tropospheric NO₂ and HCHO level 2 VCD product provided by the Tropospheric
105 Monitoring Instrument (TROPOMI) onboard the Sentinel-5 Precursor (S5P) satellite (Veefkind et
106 al., 2012). S5P is a sun-synchronous polar orbit satellite launched on 13 October 2017, which
107 covers the near-global domain in a single day. TROPOMI provides NO₂ and HCHO retrievals at
108 an approximately 7 km x 3.5 km spatial resolution (5.5km x 3.5 km since 6 August 2019) from the
109 ascending orbit with an equatorial crossing time of ~13:30 local time (Van Geffen et al., 2020; De
110 Smedt et al., 2018). The datasets were obtained from the NASA Goddard Earth Sciences Data and
111 Information Services Center (<https://daac.gsfc.nasa.gov>). A quality control procedure similar to
112 Bauwens et al. (2020) but with slightly stricter criteria is adopted for TROPOMI NO₂ and HCHO
113 data. The TROPOMI retrievals under one or more than one of the following conditions are
114 screened out for data quality control. (1) Quality assurance value is no larger than 0.5; (2) cloud
115 radiance fraction within NO₂ or HCHO retrieval window is larger than 0.3; (3) solar zenith angle
116 is larger than 70°; and (4) viewing zenith angle is larger than 70°.

117

118 **2.2 Ground O₃ and NO₂ measurements**



119 Surface measurements of O₃ and NO₂ were collected from ~1600 operational air quality
120 monitoring stations over the mainland China managed by the China National Environmental
121 Monitoring Center (<http://www.cnemc.cn/en/>). We calculated daily maximum 8-hour average
122 (MDA8) O₃ concentration from hourly in situ measurements. Surface O₃ are measured by
123 ultraviolet photometric method and Indigo disulfonate spectrophotometry, following the national
124 environmental standards of HJ 590-2010 and HJ 504-2009. Surface NO₂ concentrations are
125 measured by the chemiluminescence method (Zhang and Cao, 2015), which can cause a positive
126 bias in the NO₂ measurements (Steinbacher et al., 2007). The true NO₂ concentrations only account
127 for 43%-76% and 70%-83% of measured values for rural and urban sites (Steinbacher et al., 2007).
128 Following Wang et al. (2020b), we also applied a correction factor but with a lower value of 0.75
129 to the measured NO₂, considering that we included both rural and urban sites. The sampling ports
130 are placed at 3 to 15 meters above the ground following the national environmental monitoring
131 method standard HJ 664-2013. The measured data are reported in the unit of $\mu\text{g m}^{-3}$ under standard
132 temperature (273.15 K) and pressure (101.325 kPa) according to national environmental standards
133 GB 3095-2012.

134

135 **2.3 GEOS-Chem model and its adjoint**

136 The global 3-D chemical transport model GEOS-Chem (Bey et al., 2001) version 12.7.2 is
137 used here. We apply the nested-grid version of GEOS-Chem (Chen et al., 2009; Wang et al., 2004)
138 with the horizontal resolution of $0.25^\circ \times 0.3125^\circ$ and 47 vertical hybrid-sigma levels over East Asia
139 (70°E - 140°E , 15°N - 55°N). The boundary conditions are obtained from the $2^\circ \times 2.5^\circ$ global
140 simulation. The model is driven by the GEOS-FP meteorological field provided by NASA Global
141 Modeling and Assimilation Office (GMAO). A detailed O₃-NO_x-hydrocarbon chemistry (Mao et
142 al., 2010; Mao et al., 2013; Travis et al., 2016) is included in the GEOS-Chem model. The altitude
143 of the surface O₃ output from GEOS-Chem is specified at 9 meters above the ground to match the
144 in-situ measurements (Travis et al., 2017; Zhang et al., 2012). Through our sensitivity test using
145 GEOS-Chem, the variation of surface O₃ from 3 to 9 meters above the surface is generally less
146 than 0.723 ppb (75th percentile), and the median bias is 0.283 ppb. Travis et al. (2017) reported
147 from 60 m to 10 m above the ground, the MDA8 O₃ could decrease by ~3 ppb. Therefore, when
148 comparing GEOS-Chem surface O₃ with in-situ measurements, the differences caused by
149 inconsistent reported altitudes (9 m versus 3-15 m) can be ignored.



150 The global anthropogenic emission used in GEOS-Chem model is the Community
151 Emissions Data System (CEDS) inventory (Hoesly et al., 2018), which is replaced by the MIX
152 inventory (Li et al., 2017) over the Asian region. Biogenic emissions for VOCs follows the Model
153 of Emissions of Gases and Aerosols from Nature (MEGAN) inventory (Guenther et al., 2012).
154 Natural NO_x emissions includes biomass burning from GFED4 inventory (Van Der Werf et al.,
155 2017), soil NO_x emissions (Hudman et al., 2012) and lightning sources (Murray et al., 2012; Ott
156 et al., 2010).

157 The adjoint of the GEOS-Chem model (Henze et al., 2007; Henze et al., 2009) is a
158 component of the 4D-Var inversion method that can efficiently optimize spatially disaggregated
159 aerosol and gas emissions. This is done through iterative minimization of a cost function using the
160 model adjoint to calculate the gradient of the cost function with respect to a large number of model
161 parameters (such as anthropogenic NO_x emissions in each grid box) simultaneously. The cost
162 function is the sum of the error weighted difference between forward model outputs and
163 observations and the divergence of posterior model parameters from the prior estimate (Section
164 2.4). We developed and validated the observation operator for TROPOMI NO₂ in the GEOS-Chem
165 adjoint model version 35n similar to Wang et al. (2020a) and used it to optimize the anthropogenic
166 NO_x emission during the lockdown period in China. The NO_x emission optimization is
167 implemented using the 4D-Var method with GEOS-Chem adjoint at the nested grid with the
168 resolution of 0.25°×0.3125°. The prior anthropogenic NO_x emission used in the GEOS-Chem
169 adjoint is HTAP version 2 (Janssens-Maenhout et al., 2015), which is equivalent to the MIX
170 inventory in East Asia (Li et al., 2017).

171

172 **2.4 NO_x and VOC emission updates**

173 Two approaches are used to update the emissions during the lockdown period in 2020. The
174 first is a simple mass balance approach (Leue et al., 2001; Martin et al., 2003; Vinken et al., 2014)
175 for updating the NO_x emission by assuming a constant NO_x lifetime and NO_x/NO₂ ratio. In the
176 period from 2010 to 2019, the anthropogenic NO_x emissions have declined significantly as a result
177 of the clean air actions of Chinese government (Zheng et al., 2018). We scale the anthropogenic
178 NO_x emission from year 2010 to 2019 using the spatially gridded ratio of mean TROPOMI
179 tropospheric NO₂ VCD in Feb.-Mar. 2019 to GEOS-Chem simulated NO₂ column with default
180 MIX 2010 emission (Appendix A), to obtain the baseline anthropogenic NO_x emission in 2019,



181 which is denoted as MIX 2019. To derive anthropogenic NO_x emissions in 2020 in China during
182 the COVID-19 lockdown (MIX 2020), the spatially gridded ratio of mean TROPOMI tropospheric
183 NO₂ VCD in 2020 Feb.-Mar. to that in 2019 Feb.-Mar. is taken as a scaling factor for the updated
184 baseline anthropogenic NO_x emission in 2019 (MIX 2019). The two-month mean of TROPOMI
185 NO₂ VCD in 2019 and 2020 are calculated with the physical oversampling procedure (Sun et al.,
186 2018). Scaling factors in regions where mean TROPOMI tropospheric NO₂ VCD in 2019 Feb.-
187 Mar. is less than 0.1 Dobson unit (DU) are set to 1 for emission updates in both 2020 and 2019,
188 assuming that the lockdown only affects the populated areas (that have high NO₂ in 2019).

189 The second method for updating NO_x emission is 4D-Var via the GEOS-Chem adjoint
190 model. The anthropogenic NO_x emissions in 2020 lockdown period derived from the GEOS-Chem
191 adjoint is denoted as 2020 Adjoint. Following Wang et al. (2020a), the cost function J for
192 optimizing the NO_x emission is defined as

$$193 \quad J = \frac{1}{2} \sum_{\mathbf{c} \in \Omega} [H(\mathbf{c}) - \mathbf{s}]^T \mathbf{S}_{\text{obs}}^{-1} [H(\mathbf{c}) - \mathbf{s}] + \frac{1}{2} \gamma (\boldsymbol{\sigma} - \boldsymbol{\sigma}_a)^T \mathbf{S}_a^{-1} (\boldsymbol{\sigma} - \boldsymbol{\sigma}_a) \quad (1)$$

194 where \mathbf{s} is the tropospheric slant column density of TROPOMI NO₂, which is the product of
195 TROPOMI NO₂ VCD and air mass factor. H is the TROPOMI NO₂ observation operator that maps
196 the modeled NO₂ concentrations \mathbf{c} to the observations in time and space and calculates the
197 corresponding slant column density to make an apple-to-apple comparison of the model to
198 TROPOMI. Ω is the spatial and temporal domain where both model simulations and observations
199 are available. $\boldsymbol{\sigma}$ is the scaling factor of anthropogenic NO_x emissions to be optimized, and $\boldsymbol{\sigma}_a$ is
200 the prior emission scaling factors, which equals 1. \mathbf{S}_{obs} and \mathbf{S}_a are observational and prior error
201 covariance matrices, respectively. γ is the regularization factor that balances the weights of the
202 observational term and prior term. We assumed \mathbf{S}_{obs} to be diagonal following Wang et al. (2020a)
203 with the diagonal values calculated as the square of the standard error of tropospheric NO₂ slant
204 column density from the TROPOMI product. The prior error of the NO_x emissions is assumed to
205 be 100%. The spatial correlation of NO_x emissions is considered in this study, and the off-diagonal
206 elements of \mathbf{S}_a are computed by assuming an exponentially decaying error correlation with a fixed
207 decaying distance of 150 km following Qu et al. (2017). The γ value was determined as 500 via
208 the total error minimization and L-curve test (Henze et al., 2009; Qu et al., 2017).

209 We developed the observation operator for TROPOMI NO₂ product in the GEOS-Chem
210 adjoint model with GEOS-Chem NO₂ vertical profiles and TROPOMI NO₂ averaging kernel



211 applied to minimize the discrepancies between the assumptions in TROPOMI NO₂ retrieval and
212 GEOS-Chem model simulation. See Appendix B for additional details. The observation operator
213 has been validated using the finite difference method (Appendix C).

214 For anthropogenic VOC emissions update, we only applied the mass balance method based
215 on the TROPOMI HCHO data. The default anthropogenic VOC emissions used in the GEOS-
216 Chem is also MIX 2010 (Li et al., 2017). We ignore the change of anthropogenic VOC emissions
217 from 2010 to 2019. The baseline VOC emission in 2019 (MIX 2019) is identical to that of MIX
218 2010. The updated anthropogenic VOC emissions during the lockdown period is denoted as MIX
219 2020. HCHO is one species of VOC and may not be able to represent other VOC species. Different
220 from NO_x, biogenic sources, meteorological impacts, and large retrieval uncertainty of HCHO due
221 to its low optical depth prevent accurately quantifying the emission decline due to lockdown from
222 satellite retrievals (Levelt et al., 2022). To optimize the signal, we spatially aggregate the ratio of
223 TROPOMI HCHO during the lockdown period to that before the lockdown to the resolution of
224 0.5°, which are used as the scaling factors for updating the anthropogenic VOC emissions during
225 the lockdown period. The aggregation is based on the oversampling of TROPOMI HCHO at 0.01°
226 resolution, and the ratio is computed as the mean of the lowest 25th percentile of all ratios at 0.01°
227 resolution in each 0.5°×0.5° grid box, which ensures that only statistically significant changes are
228 considered. We assumed the change of anthropogenic VOC emissions over sparsely populated
229 areas (TROPOMI NO₂ in 2019 Feb.-Mar. less than 0.1 DU) is insignificant and assigned the ratio
230 values as one. To further evaluate the uncertainties associated with this approach, we also
231 conducted sensitivity study by using different threshold in the aggregation.

232 We assess the results from model experiments (as described in Section 2.5) adopting the
233 updated NO_x emission by comparing mean tropospheric NO₂ VCD from GEOS-Chem and from
234 TROPOMI observations in Feb.-Mar. of 2019 and 2020. The averaging kernel of TROPOMI NO₂
235 is applied to modeled NO₂ column for this comparison, following Sha et al. (2021). Further
236 quantitative evaluation of the model results also used the surface observation of O₃ and NO₂.

237

238 **2.5 GEOS-Chem model experiments**

239 A series of sensitivity experiments is conducted over China with different NO_x and VOC
240 emissions and GEOS-FP meteorological fields in different years using GEOS-Chem (v12.7.2)
241 model. All simulations are conducted from Jan. 15 to Mar. 31. The 17 days before Feb. 1 are used



242 for spin up, and the model output for Feb. and Mar. are used for the analysis. The configurations
 243 of different simulations are listed in Table 1.

244

245 **Table 1.** Configurations of model sensitivity experiments.

Experiments	Abbreviation	Meteorology	NOx Emission	VOC Emission
Baseline (2019)	2019B	GEOS-FP 2019	MIX 2019	MIX 2019
2020 Default	2020D	GEOS-FP 2020	MIX 2019	MIX 2019
2020 NOx	2020N	GEOS-FP 2020	MIX 2020	MIX 2019
2020 VOC	2020V	GEOS-FP 2020	MIX 2019	MIX 2020
2020 Lockdown	2020L	GEOS-FP 2020	MIX 2020	MIX 2020
2020 Adjoint	2020A	GEOS-FP 2020	Adjoint 2020	MIX 2020

246

247 We use the following equations to quantify the contributions from NOx and VOC emission
 248 reduction due to COVID-19 and meteorological variation to the increase of surface O₃.

249
$$\Delta O_3^{\text{NOx}} = \frac{O_3^{2020A} - O_3^{2020V}}{O_3^{2019B}} \times 100\% \quad (2)$$

250
$$\Delta O_3^{\text{VOC}} = \frac{O_3^{2020L} - O_3^{2020N}}{O_3^{2019B}} \times 100\% \quad (3)$$

251
$$\Delta O_3^{\text{ems}} = \frac{O_3^{2020A} - O_3^{2020D}}{O_3^{2019B}} \times 100\% \quad (4)$$

252
$$\Delta O_3^{\text{met}} = \frac{O_3^{2020D} - O_3^{2019B}}{O_3^{2019B}} \times 100\% \quad (5)$$

253 Where ΔO_3^{NOx} , ΔO_3^{VOC} and ΔO_3^{ems} are the relative differences in surface O₃ concentration caused
 254 by emission decline of NOx, VOC, and both NOx and VOC resulting from COVID-19. ΔO_3^{met}
 255 represents the relative contribution to the surface O₃ change from the meteorological variation
 256 between 2 years. O_3^{2019B} , O_3^{2020D} , O_3^{2020N} , O_3^{2020V} , O_3^{2020L} and O_3^{2020A} are mean MDA8 surface
 257 O₃ concentration simulated by modeling experiments Baseline (2019), 2020 Default, 2020 NOx,
 258 2020 VOC, 2020 Lockdown and 2020 Adjoint, respectively (Table 1).

259 The difference in simulated surface O₃ between 2020 and 2019, is the result of both
 260 emission reductions and meteorological variations and is denoted as ΔO_3^{all} . It is calculated as



261 follows and is evaluated against the observed relative difference of mean MDA8 O₃ in Feb. to Mar.
262 between 2019 and 2020 at all ground sites:

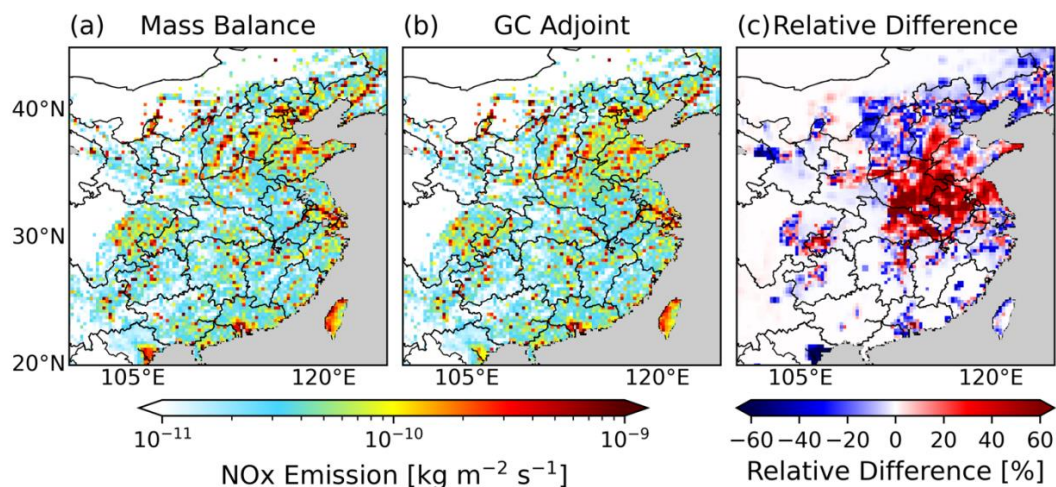
$$263 \quad \Delta O_3^{\text{all}} = \frac{O_3^{2020A} - O_3^{2019B}}{O_3^{2019B}} \times 100\% \quad (6)$$

264

265 **3. Results of model development, emissions, and validation**

266 **3.1 Changes of NO_x and VOC emissions during COVID**

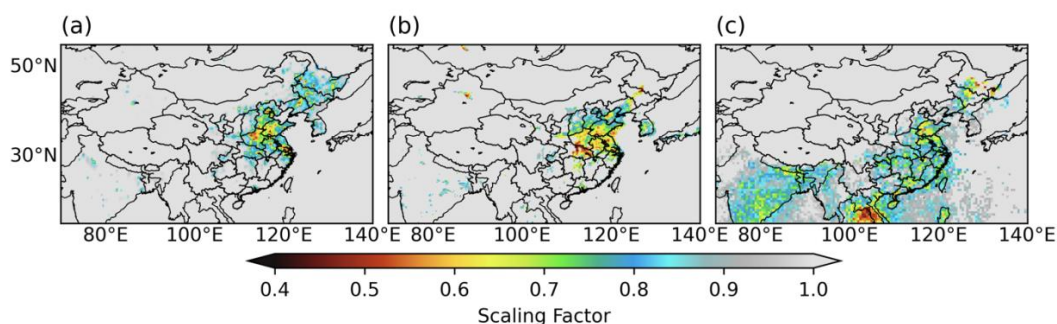
267 We updated the anthropogenic NO_x emissions during the COVID lockdown using both
268 4D-Var and mass balance methods (Fig. 1 and 2). The NO_x emissions from the 4D-Var inversion
269 share a similar spatial pattern and magnitude with those found using the mass balance method (Fig.
270 1). However, the NO_x emissions from the 4D-Var inversion are lower overall than those from the
271 mass balance method over North China by ~10% and larger over central China by ~40%. Fig. 2(a-
272 b) shows that the 4D-Var NO_x emission reduction is more severe over urban regions and displays
273 a smoother spatial pattern than that from the mass balance approach, which is caused by the
274 arbitrary cut off with 0.1 DU of NO₂ VCD in the latter. Furthermore, the 4D-Var inversion
275 captured the NO_x emission decline in Northeast China where the mass balance approach did not
276 because of the low NO₂ VCD. During Feb.-Mar. 2020, the anthropogenic NO_x emissions in East
277 China decreased by ~30% compared to those in the same period in 2019. We also scale the
278 anthropogenic VOC emissions based on the TROPOMI HCHO data (Fig. 2(c)). The VOC
279 emissions decrease by ~20%-30% in East and South Asia. The anthropogenic VOC emission
280 changes in sparsely populated areas over Northwest China are neglected.



281

282 **Figure 1.** Updated anthropogenic NO_x emission during Feb.-Mar. 2020 from (a) mass balance
283 method, (b) 4D-Var method and (c) their relative difference.

284



285

286 **Figure 2.** Scaling factors for anthropogenic NO_x emission in Feb.-Mar. from 2019 to 2020 as
287 derived from (a) 4D-Var, (b) mass balance. Scaling factors for anthropogenic VOC emissions from
288 the mass balance are in (c).

289

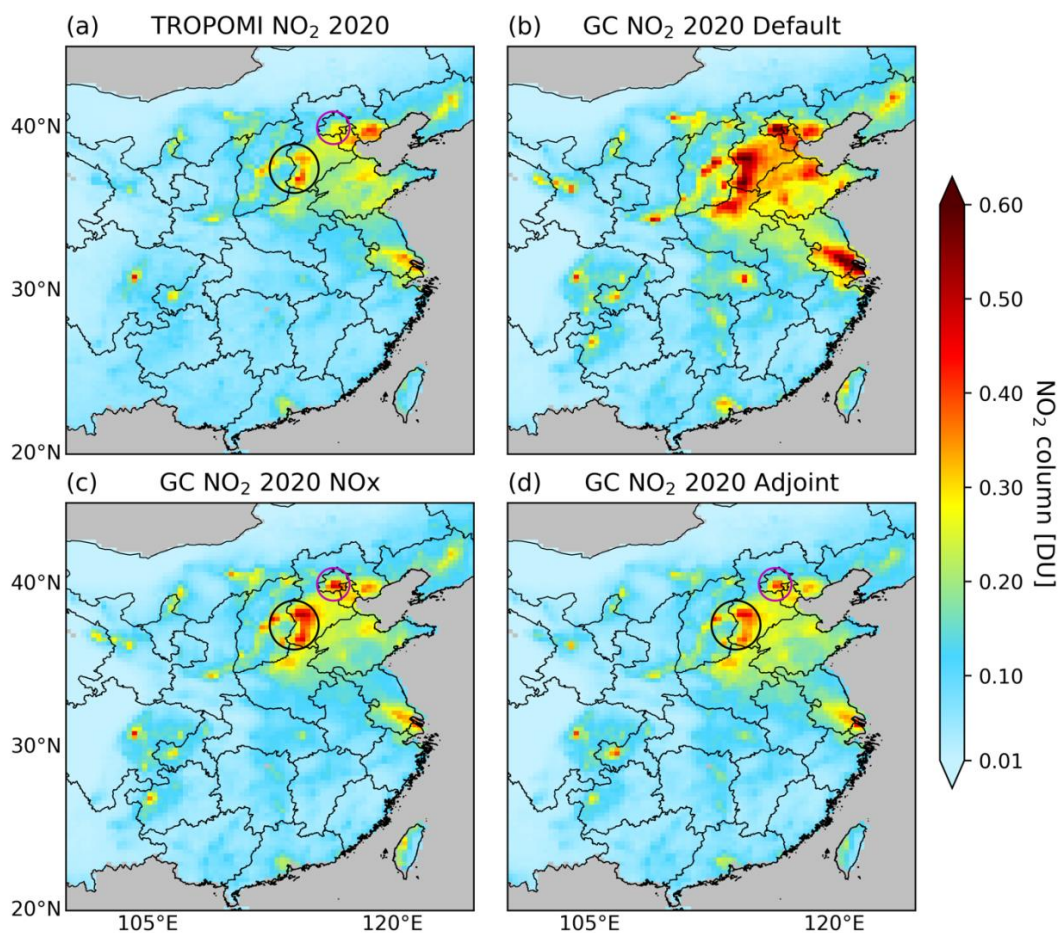
290 3.2 Validation of NO₂ simulations

291 We further assess our updated anthropogenic NO_x emissions by comparing the NO₂ VCD
292 from TROPOMI with that from GEOS-Chem with the anthropogenic NO_x emissions before and
293 after the scaling (Fig. 3 and 4). Before updating the NO_x emissions, the 2020 Default (Fig. 3(b))
294 simulation significantly overestimates the NO₂ VCD compared to the TROPOMI NO₂
295 observations (Fig. 3(a)). With the NO_x emissions updated, the simulations 2020 NO_x (Fig. 3(c))



296 and 2020 Adjoint (Fig. 3(d)) exhibit a much better agreement with TROPOMI NO₂ observation
297 than 2020 Default. However, Fig. 3(c) shows the GEOS-Chem simulation with the NO_x emissions
298 from mass balance approach overestimated the NO₂ VCD over Beijing and southwest of Hebei
299 Province (pink and black circles in Fig.3) compared with TROPOMI data. The reason is that
300 scaling factors are applied only to anthropogenic NO_x emissions, not total NO_x emissions, so it is
301 expected that the model may still overestimate the NO₂ column after scaling part of the total NO_x
302 emission. With the anthropogenic NO_x emissions optimized by the 4D-Var method, the
303 overestimation of NO₂ VCD over Beijing and southwest of Hebei Province (pink and black circles
304 in Fig. 3) is mitigated compared with the NO_x emissions from mass balance approach.

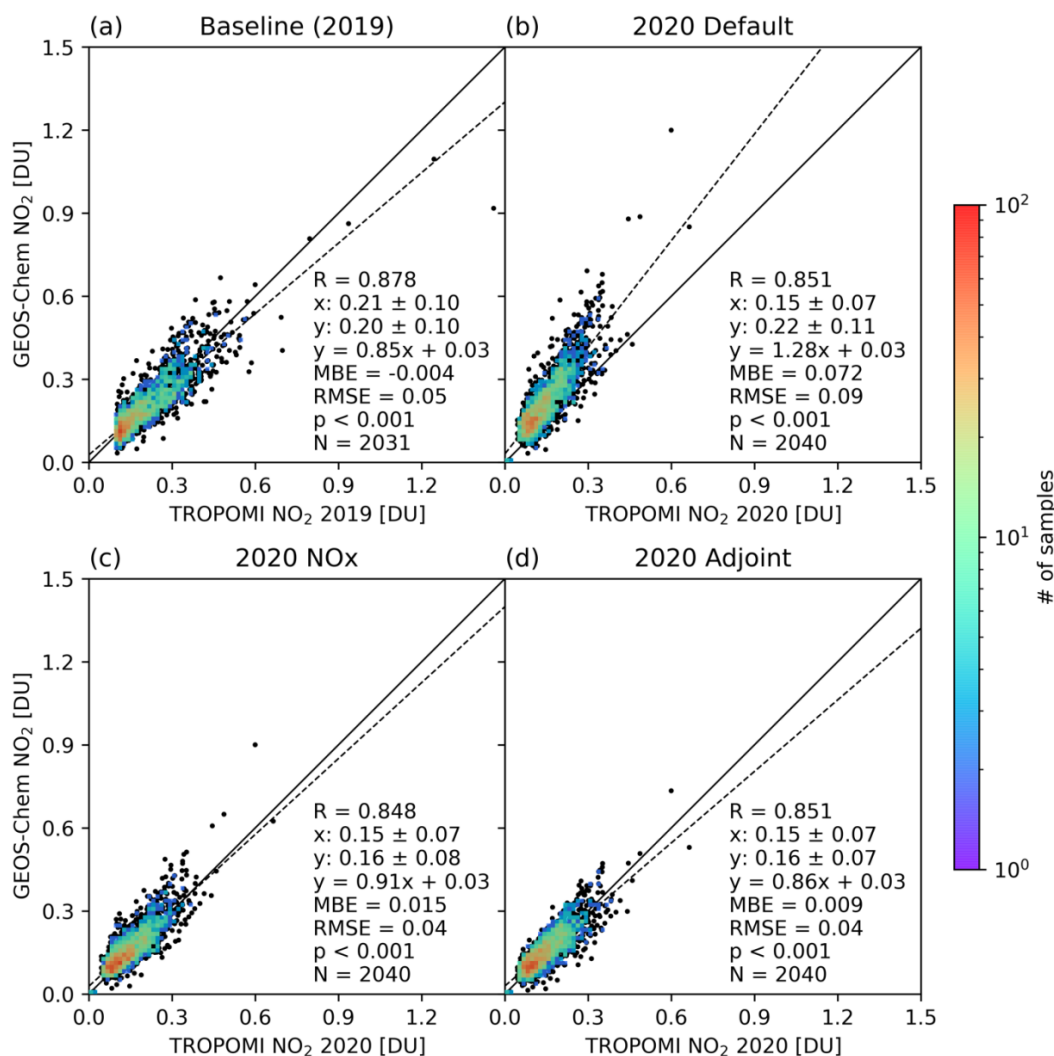
305 Fig. 4 further displays the statistics for the comparison between the TROPOMI NO₂ and
306 GEOS-Chem simulations via the scatterplot. The Baseline (2019) simulation captures the
307 magnitude of NO₂ VCD observations in 2019 well (Fig. 4(a)). The root-mean-square-error (RMSE)
308 and mean bias error (MBE) for the simulation with 2020 NO_x emission derived from mass balance
309 method (Fig. 4(b)) decreased by 0.050 DU and 0.057 DU as compared to the 2020 Default (Fig.
310 4(c)). Compared with the result from GEOS-Chem simulation 2020 NO_x, emissions from 2020
311 Adjoint (Fig. 4(d)) further led to the reduction of the MBE of the NO₂ VCD by 0.006 DU and
312 improve the correlation coefficient by 0.003. The significant overestimation of several pixels with
313 TROPOMI NO₂ VCD larger than 0.4 DU by the simulation 2020 NO_x is also mitigated by 2020
314 Adjoint. The MBE between GEOS-Chem and TROPOMI for Baseline (2019), 2020 NO_x and
315 2020 Adjoint are -0.004, 0.015 and 0.009 DU, respectively. The corresponding relative bias are
316 1.9%, 10% and 6.0%, which are all less than the relative uncertainty of ~30% for TROPOMI
317 tropospheric NO₂ VCD over East China (Van Geffen et al., 2022). The improved agreement
318 between the simulation with updated NO_x emission and TROPOMI NO₂ provides a basis for
319 further analyzing the mechanism of aggravated surface O₃ pollution.



320

321 **Figure 3.** Comparison of tropospheric NO₂ VCD from (a) TROPOMI product in 2020 Feb.-Mar
322 with that from GEOS-Chem simulations (b) 2020 Default, (c) 2020 NO_x and (d) 2020 Adjoint.
323 The pink and black circles mark the areas where NO_x emissions from 4D-Var mitigated the NO₂
324 overestimation by mass balance method. The emissions and meteorology configurations for
325 GEOS-Chem simulations 2020 Default, 2020 NO_x and 2020 Adjoint are listed in Table 1.

326



327

328 **Figure 4.** Scatter plot of TROPOMI NO₂ VCD versus the GEOS-Chem simulations for (a)
329 Baseline (2019), (b) 2020 Default, (c) 2020 NO_x and (d) 2020 Adjoint, respectively. TROPOMI
330 data in Feb.-Mar. of 2019 was used in (a), and that of 2020 was used in (b-d). The emissions and
331 meteorology configurations for GEOS-Chem simulations are listed in Table 1. Only pixels with
332 TROPOMI NO₂ VCD in 2019 Feb.-Mar. larger than 0.1 DU are included in all comparisons.

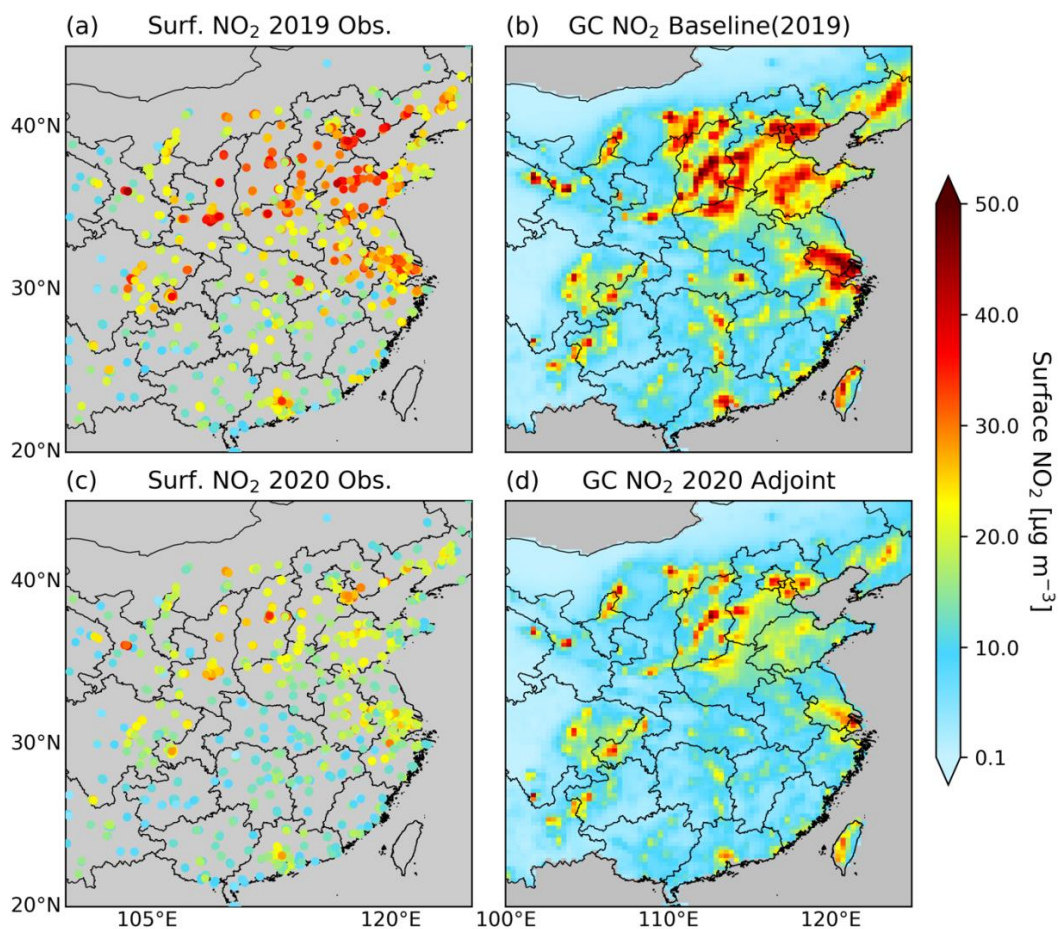
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334



335 Fig. 5 and Fig. 6 show the comparison of surface NO₂ between ground measurements and
336 GEOS-Chem simulations. The GEOS-Chem simulations Baseline (2019) (Fig. 5(b)) and 2020
337 Adjoint (Fig. 5(d)) both capture the spatial pattern and magnitude of surface NO₂ measurements
338 in Feb.-Mar. of 2019 (Fig. 5(a)) and 2020 (Fig. 5(c)) well, respectively. Fig. 6 further displays the
339 good agreements of surface NO₂ from Baseline (2019) (Fig. 6(a)) and 2020 Adjoint (Fig. 6(b)) to
340 the in-situ measurements via scatter plots. Table 2 displays the evaluation statistics, including the
341 correlation coefficient (R), MBE, RMSE and the slope and intercept of the linear regression, for
342 the simulated surface NO₂ from various simulation experiments compared with the in-situ
343 measurements. The correlation coefficient, MBE and RMSE between the simulation Baseline
344 (2019) and ground measurements in 2019 Feb.-Mar. are 0.724, 1.572 μg m⁻³ and 8.49 μg m⁻³,
345 respectively. Without updating the NO_x emissions in 2020, the simulation 2020 Default
346 overestimate the ground measurements of surface NO₂ in 2020 Feb.-Mar (Table 2). The slope for
347 the linear regression is 1.19, and the MBE and RMSE are 6.021 μg m⁻³ and 10.43 μg m⁻³,
348 respectively (Table 2). After updating the NO_x emissions, the GEOS-Chem simulations 2020 NO_x
349 and 2020 Adjoint have good agreements with the in-situ measurements in 2020 Feb.-Mar. The
350 correlation coefficient between the simulation 2020 Adjoint versus the in-situ measurements is
351 0.651, higher than that of 0.608 for the simulation 2020 NO_x versus the ground measurements
352 (Table 2). The MBE and RMSE of 2020 Adjoint (0.683 μg m⁻³ and 6.68 μg m⁻³) are lower than
353 those of 2020 NO_x (1.726 μg m⁻³ and 7.74 μg m⁻³) (Table 2). This result further indicates the
354 superiority of 4D-Var for optimizing NO_x emissions compared with the mass balance method
355 (Cooper et al., 2017; Streets et al., 2013).

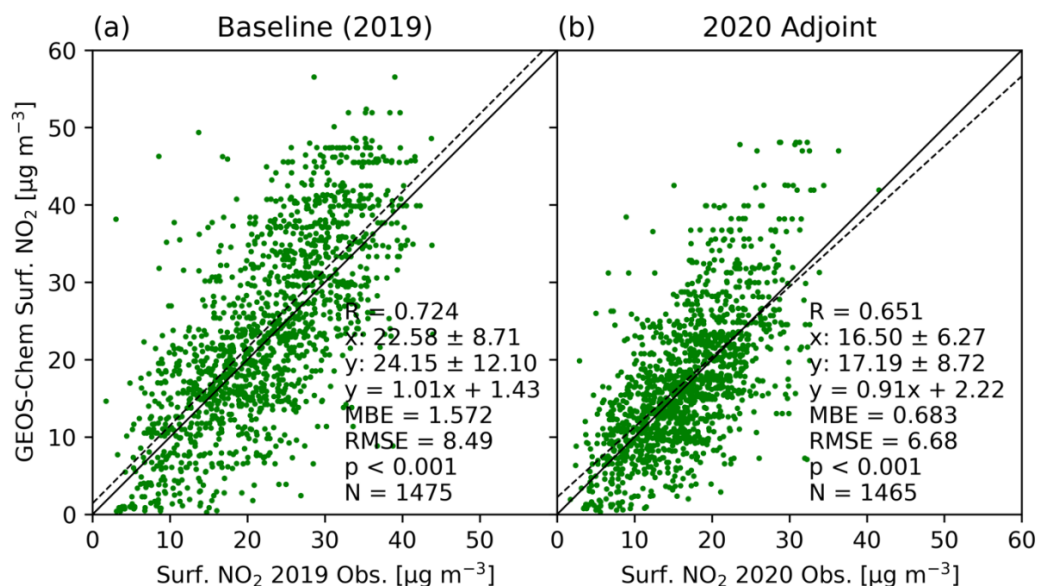
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357

358 **Figure 5.** Comparison of surface NO₂ concentrations from ground measurements for (a) 2019
359 Feb.-Mar. and (c) 2020 Feb.-Mar. versus those from GEOS-Chem simulations (b) Baseline (2019)
360 and (d) 2020 Adjoint. Grey color means no data is presented.

361



362

363 **Figure 6.** Scatter plots for comparing the surface NO₂ concentrations from GEOS-Chem
 364 simulations and ground measurements. (a) GEOS-Chem simulation Baseline (2019) versus ground
 365 measurements in 2019 Feb.-Mar. (b) GEOS-Chem simulation 2020 Adjoint versus ground
 366 measurements in 2020 Feb.-Mar. Note: the number of ground sites differ in these two years.

367

368

369 **Table 2.** Evaluation statistics for modeled surface NO₂ compared with the in-situ measurements*.

Experiments	R	MBE ($\mu\text{g m}^{-3}$)	RMSE ($\mu\text{g m}^{-3}$)	Slope	Intercept ($\mu\text{g m}^{-3}$)
Baseline (2019)	0.724	1.572	8.49	1.01	1.43
2020 Default	0.661	6.021	10.43	1.19	2.95
2020 NO_x	0.608	1.726	7.74	0.92	3.03
2020 Adjoint	0.651	0.683	6.68	0.91	2.22

370

* The simulation experiment Baseline (2019) is compared with the ground measurements in 2019
 371 Feb.-Mar. Other three experiments are compared with the ground measurements in 2020 Feb.-Mar.

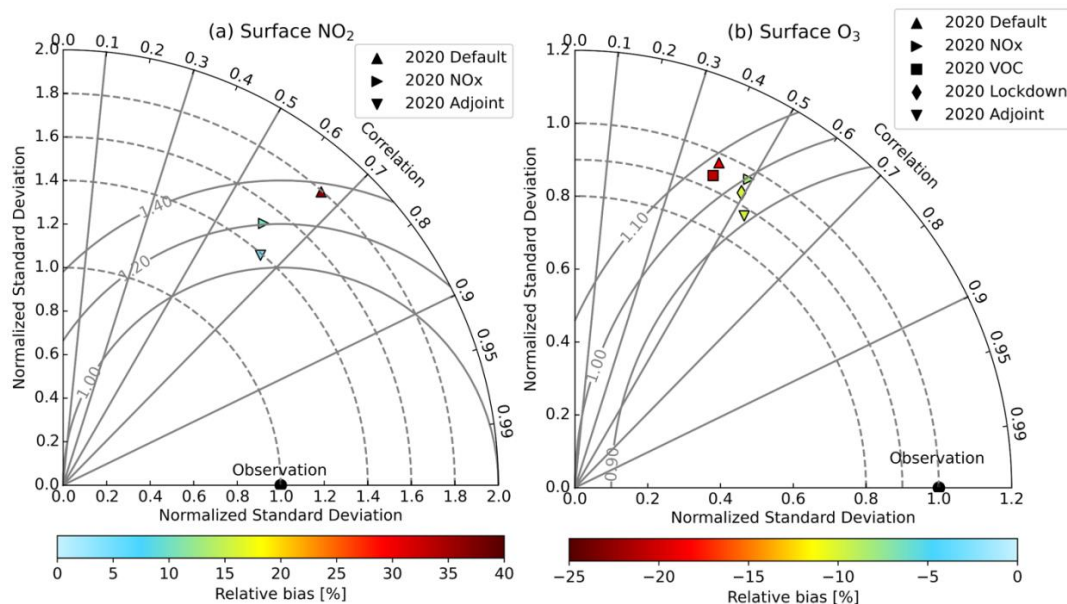
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373

Fig. 7 (a) is the Taylor diagram for evaluating the GEOS-Chem simulations of surface NO₂
 374 concentrations from 2020 Default, 2020 NO_x and 2020 Adjoint using the in-situ measurements.



375 The simulation 2020 Adjoint (inverted triangle in Fig. 7(a)) has the best performance among these
376 three simulations with the lowest relative bias and lowest normalized centered RMSE. Without
377 updating the NO_x emission, 2020 Default features a relative bias of ~37%. After updating the NO_x
378 emissions, 2020 NO_x reduces the relative bias, normalized centered RMSE and normalized
379 standard deviation from around 37%, 1.38 and 1.80 to around 10%, 1.20 and 1.51 compared with
380 2020 Default, but the correlation coefficient also decreases. By using 4D-Var method, 2020
381 Adjoint further reduces the relative bias, normalized centered RMSE and normalized standard
382 deviation and increases the correlation coefficient compared with 2020 NO_x.



383

384 **Figure 7.** Taylor diagram for evaluating the GEOS-Chem simulations of (a) surface NO₂ and (b)
385 surface O₃ during lockdown period (2020 Feb.-Mar.) using ground observations for different
386 simulation experiments listed in Table 1. The evaluation of surface O₃ only includes the areas
387 where the NO_x emissions optimized by 4D-Var reduced by more than 10%.

388

389

390 3.3 Evaluation of surface O₃ simulations

391

392

393

We evaluated the GEOS-Chem simulations of MDA8 surface O₃ from different simulation
experiments listed in Table 1 using ground measurements. Fig. 7(b) is the Taylor diagram for
comparing the surface O₃ concentrations during 2020 Feb.-Mar. from ground measurements and

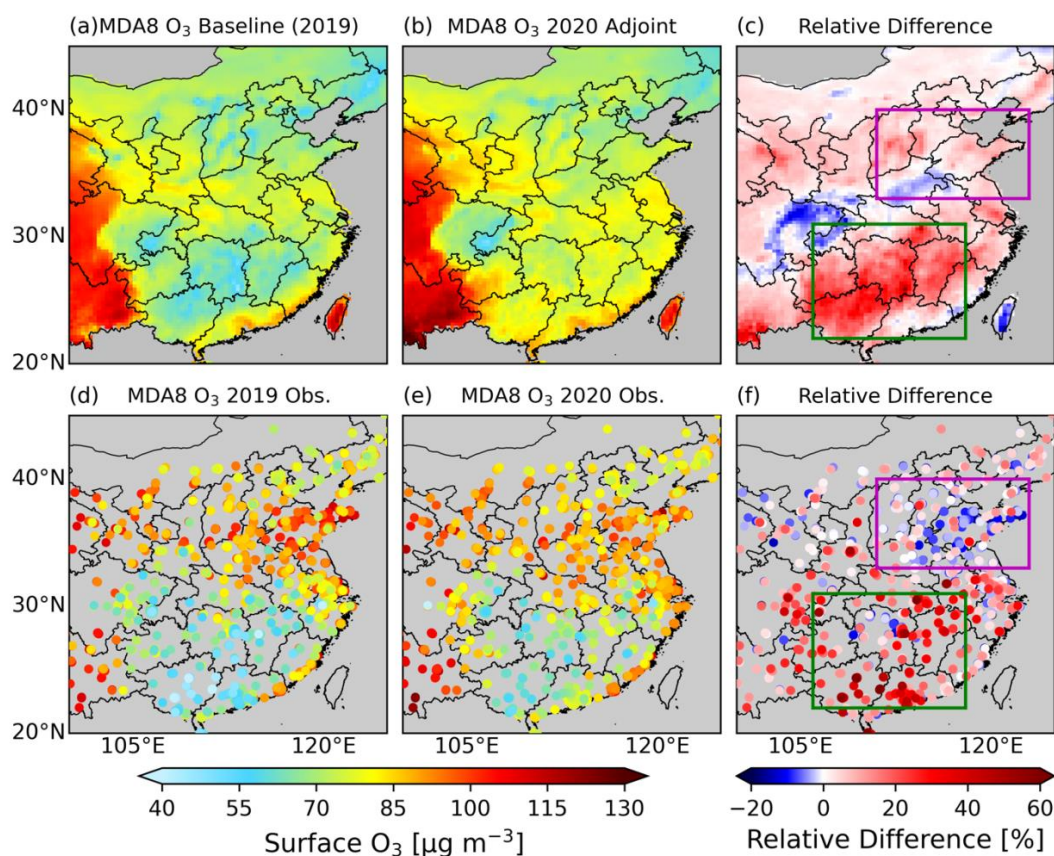


394 GEOS-Chem simulations. We focused on areas with significant NO_x emissions reduction to better
395 assess the role of updated NO_x emissions in improving surface O₃ simulations. The ground sites
396 where the NO_x emissions from 4D-Var decline by less than 10% are excluded. The correlation
397 coefficient between the simulation Baseline (2019) and ground observations is ~0.53 and the
398 relative bias is around -25%. By applying 2020 meteorological fields and scaling the VOC
399 emissions, the correlation coefficients decreased to ~0.40 for model simulations 2020 Default and
400 2020 VOC, with little reduction in the relative bias. By updating the NO_x emissions, the relative
401 bias reduced to around -10% while the correlation coefficients remain at ~0.50 for model
402 simulations 2020 NO_x, 2020 Lockdown and 2020 Adjoint. This indicates the NO_x emission
403 updates significantly improve the surface O₃ simulations. Comparing the simulations 2020 Default
404 and 2020 VOC, or 2020 NO_x and 2020 Lockdown, the results show that scaling VOC emissions
405 does not improve the surface O₃ simulations significantly over the continental China, but over
406 South China, VOC emissions update reduces the relative bias by 3%. Among all simulations, 2020
407 Adjoint exhibits the best performance with the lowest normalized centered RMSE, largest
408 correlation coefficient and a low relative bias of ~10%. This result further confirms the superiority
409 of the 4D-Var with respect to the mass balance method for optimizing NO_x emissions. Therefore,
410 we used the 2020 Adjoint to evaluate the impacts of NO_x emission on surface O₃ in the following
411 analysis.

412 Fig. 8 compares the modeled surface O₃ in Feb.-Mar. of 2019 (Fig. 8(a)) and 2020 (Fig.
413 8(b)) and the relative difference (Fig. 8(c)) computed from Equation (6) with the in-situ
414 measurements (Fig. 8(d-f)). The ground observations show that the highest level of surface O₃
415 pollution occurs in North China and southwest of China. The average MDA8 O₃ in two months
416 can reach up to ~110 μg m⁻³ at STP (~51.4 ppbv), which is higher than the China National Ambient
417 Air Quality Standard daily maximum 8-hour Grade I standard of 100 μg m⁻³. GEOS-Chem model
418 underestimates the surface O₃ over North China for both years compared with ground observations,
419 which could be a result of the out-of-date VOC emissions, but it captures the magnitude and spatial
420 distribution of surface O₃ and the increasing trend in South China well. In South China, the
421 measured surface O₃ in 2020 Feb.-Mar. increases by 30-50%, while over North China, it increases
422 generally by less than 20% even decreases in some regions. The relative differences of simulated
423 surface O₃ between two years is comparable to the ground observations over South China (green
424 box in Fig. 8(c, f)). Over North China (pink box in Fig. 8(c, f)), the average relative difference



425 between two years from the model and observation are 4.27% and -3.01%, respectively, both of
426 which are much smaller than their counterparts in South China. While the relative difference from
427 model simulations has different signs as compared to that of observations on average, both the
428 change of O_3 is indeed small and the model is able to capture the part of O_3 decrease in the
429 southwest part of the North China domain (Fig. 8(c)). We note that some previous studies showed
430 large increase of O_3 in North China, but such increase is in comparison with the O_3 in the month
431 right before the lockdown (not the same time in 2019; (Shi & Brasseur, 2020; Y. M. Liu et al.,
432 2021).
433



434
435 **Figure 8.** Comparison of MDA8 surface O_3 in 2019 and 2020 Feb.-Mar. and the relative difference
436 between two years from GEOS-Chem model (a-c) versus ground observations (d-f). GEOS-Chem
437 mean MDA8 O_3 at 9 m above the surface under standard temperature and pressure (STP; 273.15



438 K, 101.325 kPa) from (a) Baseline (2019) and (b) 2020 Adjoint simulation (Table 1) together with
439 (c) their relative difference. Ground observed mean MDA8 surface O₃ under STP in (d) 2019 Feb.-
440 Mar.; (e) 2020 Feb.-Mar. and (f) their relative difference. The pink and green boxes in (c) and (f)
441 define the North China and South China domain.

442

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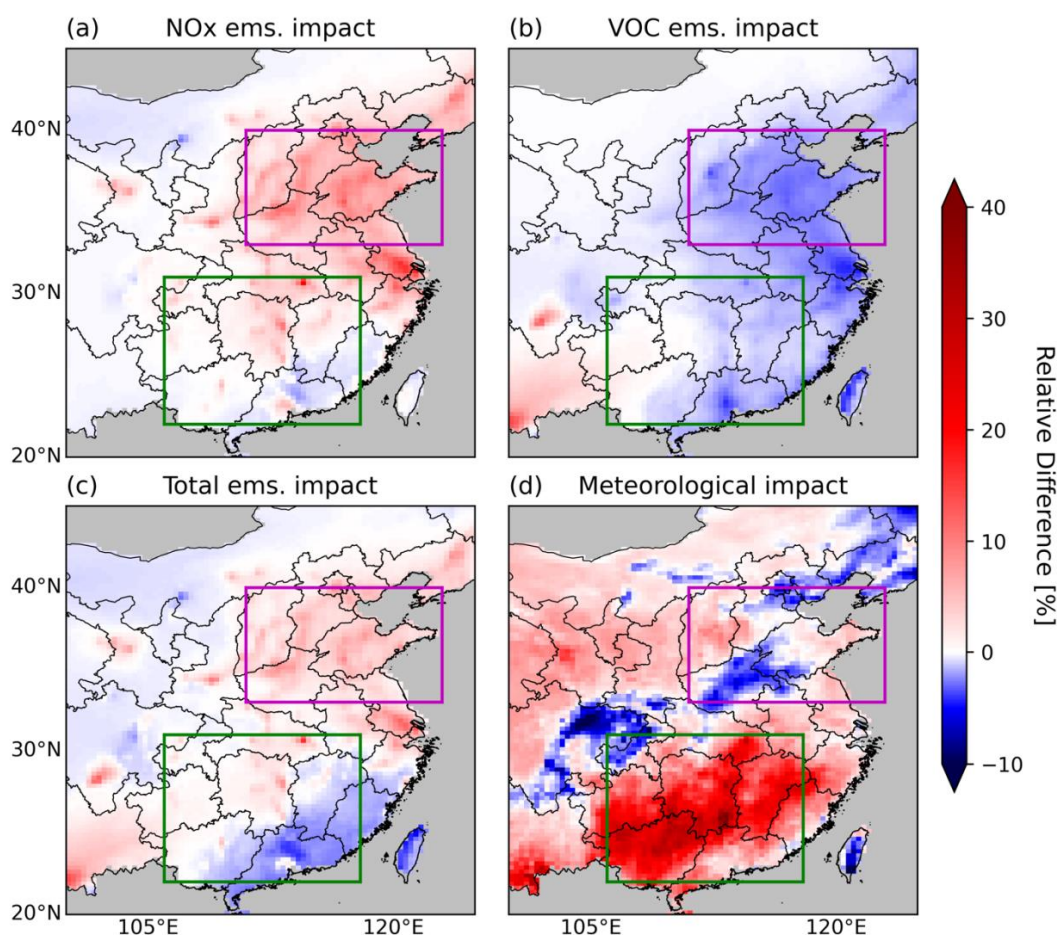
444 **4. Mechanism of aggravated surface O₃ pollution**

445 **4.1 Relative contribution from declining emissions and meteorological variations**

446 From equations (2)-(5) we can analyze the mechanism of surface O₃ increase in China
447 during the COVID-19 pandemic (Fig. 9). NO_x emission reduction as a result of COVID-19
448 lockdown leads to a ~8% increase in the mean MDA8 surface O₃ over North China (pink boxes in
449 Fig. 9) between 2019 and 2020 Feb.-Mar. (Fig. 9(a)), while the VOC emission decline causes ~3%
450 of O₃ decrease (Fig. 9(b)). The average contribution of the meteorological variations to the surface
451 O₃ change is less than 1% in North China (Fig. 9(d)). However, in South China, the inter-annual
452 meteorological variations dominate the surface O₃ increases, causing a ~30% increase (Fig. 9(d)),
453 while the NO_x and VOC emission reduction has little impacts. The overall magnitude of emissions
454 contribution to the surface O₃ change over North China is ~5%, similar to that of the
455 meteorological effects, but meteorological variations lead to both O₃ increases and decreases in
456 different regions. Over South China, the meteorological effect is much larger than the net effects
457 of declining emissions. Overall, the impact of inter-annual meteorological variations between 2019
458 and 2020 is almost 30 times larger than the overall emissions impacts on the aggregated surface
459 O₃ pollution in China. Our results are consistent with the conclusion from Zhao et al. (2020) that
460 meteorological variation has larger impacts than emissions reduction on surface O₃ in the southern
461 city of Guangzhou, but in Beijing, emission reduction has a larger impact during 23-29 January.
462 Liu et al. (2020b) reported that the surface O₃ increase in the major cities of the Yangtze River
463 Delta region were driven by both emission reduction and meteorological variations to a similar
464 degree from pre-lockdown period (Jan. 1-22, 2020) to lockdown period (Jan. 23-Feb. 29, 2020).
465 However, Zhao et al. (2020) and Liu et al. (2020b) only focused on the lockdown period of one
466 week in reference to the time period right before the lockdown instead of the same period in
467 previous years, which cannot exclude the effects of seasonal variation of meteorology and did not
468 provide a comprehensive analysis over the whole lockdown period. Moreover, Liu et al. (2020b)



469 only analyzed four representative cities instead of showing the analysis at a national scale. Further,
470 Zhao et al. (2020) did not update the anthropogenic emissions during the lockdown period, which
471 brings significant uncertainties to their analysis.
472
473



474
475 **Figure 9.** Relative difference in simulated surface O₃ caused by (a) NO_x emission reduction, (b)
476 VOC emission reduction, (c) overall emission reduction and (d) meteorological variations due to
477 COVID-19 lockdown. The pink and green boxes in each panel define the North China and South
478 China domain.

479
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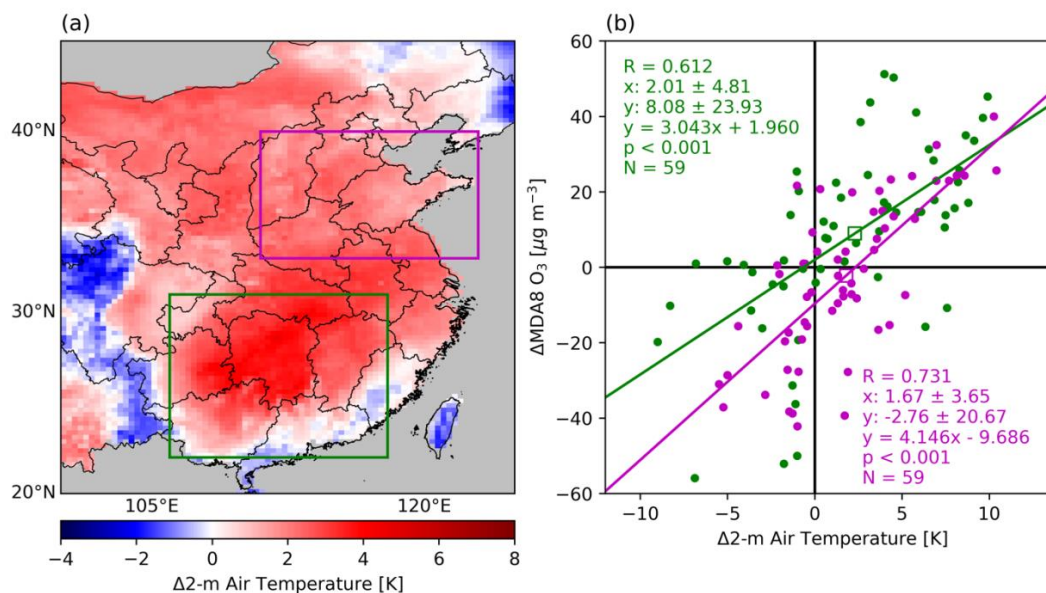


481 **4.2 Higher temperature leading to aggravated surface O₃ pollution in South China**

482 The positive correlation between the surface O₃ and temperature is widely observed and
483 reported in the literature (Pusede et al., 2015). Higher temperature leads to higher concentrations
484 of surface O₃ because it improves the O₃ production rate by affecting the organic reactivity,
485 production of HOx radicals, formation and decomposition of peroxy nitrates and alkyl nitrates
486 (Pusede et al., 2015). We calculated the daily difference in Feb.-Mar. between 2020 and 2019
487 (excluding Feb. 29, 2020) for the daily mean of MDA8 O₃ from ground measurements and 2-meter
488 air temperature from GEOS-FP data used in our GEOS-Chem simulations for the South China
489 (106°E - 118°E, 22°N - 31°N, green box in Fig. 10(a)) and North China (111°E - 123°E, 33°N -
490 40°N, pink box in Fig. 10(a)). Fig. 10 displays the difference of 2-month mean 2-meter air
491 temperature in Feb.-Mar. between 2020 and 2019 (Fig. 10(a)) and the scatter plot between the
492 daily difference of measured surface O₃ concentration and 2-meter air temperature over both South
493 China (green dots in Fig. 10(b)) and North China (pink dots in Fig. 10(b)). We found the 2-meter
494 air temperature increased by ~2.3°C in South China, and the daily difference of surface O₃
495 concentration and 2-meter air temperature are well correlated with a positive correlation
496 coefficient of 0.612. Therefore, the significant aggravated surface O₃ pollution in South China
497 could be attributed to the temperature increase. The reason for the temperature increase is the lower
498 cloud fraction. Via analyzing the GEOS-FP data, we found the cloud fraction decreases by ~5%,
499 and the downward visible direct flux at surface increased by 5 W m⁻² over South China. The lower
500 cloud fraction increases the downward solar radiation at the surface during the lockdown period,
501 leading to higher surface air temperature. The enhanced solar radiation at the surface could also
502 promote the production of O₃ via photochemical reactions. In North China, 2-meter air temperature
503 also increased by 1.8°C, but the measured surface MDA8 O₃ decreased by 3% (Fig. 8(f)). Fig.
504 10(b) shows the daily difference of MDA8 O₃ and 2-meter air temperature over North China also
505 has a high correlation coefficient of 0.731. However, the intercept of the linear regression line is
506 negative, so that the surface O₃ could decrease even though the temperature increases. The
507 predicted average change of surface MDA8 O₃ in South China and North China are marked by the
508 green and pink open squares respectively in Fig. 10(b) based on the linear regression. Because of
509 the different intercepts, the predicted MDA8 O₃ in South China increases by ~9.0 μg m⁻³, while it
510 decrease by 2.2 μg m⁻³ in North China, although the average temperature increased in both South
511 and North China.



512



513

514 **Figure 10.** (a) The change of 2-meter air temperature from 2019 Feb.-Mar. to 2020 Feb.-Mar. (b)
515 The scatter plots between the daily difference of surface O_3 measurements and 2-meter air
516 temperature in Feb.-Mar. between 2020 and 2019 over South China (green dots) and North China
517 (pink dots). The green and pink open squares mark the predicted average change of surface MDA8
518 O_3 in South China (green open square) and North China (pink open square), respectively, based
519 on the linear regression against temperature change.

520

521 **5. Summary**

522 A significant reduction in primary air pollutants has been identified by surface and satellite
523 observations during the COVID-19 pandemic in China (Bauwens et al., 2020; Miyazaki et al.,
524 2020), which is in contrast to the increase of surface O_3 . In this study, we analyzed the reasons for
525 the enhanced surface O_3 pollution from two perspectives: anthropogenic emissions reduction and
526 inter-annual meteorological variations. We constrain the NO_x emissions based on the TROPOMI
527 NO_2 product using both the mass balance and 4D-Var methods. The VOC emissionw were also
528 updated based on the TROPOMI HCHO product via the mass balance approach. We analyzed the
529 contributions from emissions reduction and meteorological variations to surface O_3 increases
530 through a series of sensitivity simulations using the GEOS-Chem model.



531 The updated NO_x emissions from the 4D-Var and mass balance approaches share a similar
532 spatial pattern. However, the NO_x emissions from 4D-Var are lower than those from the mass
533 balance method over North China by ~10% but larger over central China by ~40%. The evaluation
534 of the simulations with the updated emissions against the TROPOMI NO₂, in-situ measurements
535 of surface NO₂ and O₃ indicate that the NO_x emissions from the 4D-Var inversion leads to better
536 model performance than that from the mass balance approach.

537 The anthropogenic NO_x emission decreased by ~30% over East China during 2020 Feb.-
538 Mar. compared to the same period in 2019. Over North China, NO_x emission reduction leads to a
539 ~8% increase in the mean MDA8 surface O₃, while the VOC emissions decline causes O₃ to
540 decrease by ~3%. The average contribution of meteorological variations to the surface O₃ change
541 is less than 1% in North China. However, in South China, the inter-annual meteorological variation
542 dominates the surface O₃ increase, causing a ~30% increase, while the NO_x and VOC emission
543 reduction has nearly no impacts on O₃. Overall, the impact of inter-annual meteorological
544 variations between 2019 and 2020 is almost 30 times larger than the impact of emissions on the
545 enhanced surface O₃ pollution in China.

546 The significant increase of surface O₃ in South China could be attributed to the higher
547 temperature during the lockdown period, which is caused by the lower cloud fraction. The lower
548 cloud fraction increases the downward solar radiation at the surface during the lockdown period,
549 leading to higher surface air temperature. The enhanced solar radiation at the surface could also
550 promote the production of O₃ via photochemical reactions.

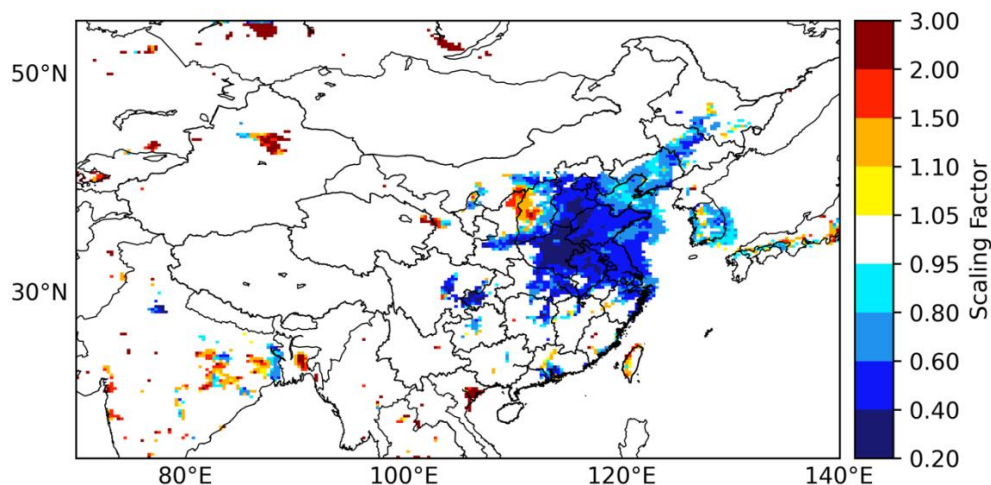
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553 **Appendix A: NO_x emission reduction in China from 2010 to 2019**

554 The default anthropogenic NO_x emission over East Asia in GEOS-Chem is MIX 2010 (Li
555 et al., 2017). To generate the anthropogenic NO_x emission in 2019, we calculated the ratio of mean
556 TROPOMI tropospheric NO₂ VCD in Feb.-Mar. 2019 to GEOS-Chem simulated NO₂ VCD with
557 the default MIX 2010 emission as the scaling factor (Fig. A1). The scaling factors in regions where
558 mean TROPOMI tropospheric NO₂ VCD in 2019 Feb.-Mar. less than 0.1 DU are set to 1. From
559 2010 to 2019, the anthropogenic NO_x emission has declined significantly as a result of the clean
560 air actions of Chinese government (Zheng et al., 2018).

561



562

563 **Figure A1.** The scaling factor of anthropogenic NO_x emission from year 2010 to 2019.

564

565

566 **Appendix B: Applying the TROPOMI NO₂ averaging kernel in the observation**
 567 **operator**

568 To optimize the NO_x emissions and minimize the cost function (Equation (1)) with the 4D-
 569 Var method, GEOS-Chem adjoint needs to compute the derivative of the cost function with respect
 570 to the model parameters to be optimized, which is the scaling factors of the anthropogenic NO_x
 571 emissions in this study. An essential step is to calculate the adjoint forcing \mathbf{F} , which is the
 572 derivative of the cost function with respect to the modeled NO₂ concentration shown as Equation
 573 (B1).

574
$$\mathbf{F} = \frac{\partial J}{\partial \mathbf{c}} = \mathbf{S}_{\text{obs}}^{-1} [H(\mathbf{c}) - \mathbf{s}] \frac{\partial H(\mathbf{c})}{\partial \mathbf{c}} \quad (\text{B1})$$

575 For each single TROPOMI NO₂ observation, the adjoint forcing component f and cost function
 576 component j are computed as Equation (B2) and Equation (B3).

577
$$f = \frac{M_{\text{gc}} v_{\text{gc}} - M_{\text{obs}} v_{\text{obs}}}{e_{\text{obs}} M_{\text{obs}}} M_{\text{gc}} \quad (\text{B2})$$

578
$$j = \frac{0.5 f (M_{\text{gc}} v_{\text{gc}} - M_{\text{obs}} v_{\text{obs}})}{M_{\text{gc}}} \quad (\text{B3})$$



579 Here M_{gc} is GEOS-Chem air mass factor applying the GEOS-Chem NO₂ vertical profiles and
580 TROPOMI NO₂ averaging kernel. M_{obs} is TROPOMI air mass factor. v_{gc} and v_{obs} are the
581 tropospheric NO₂ VCD from GEOS-Chem model and TROPOMI observation, respectively. The
582 product of air mass factor and NO₂ VCD is NO₂ slant column density. e_{obs} is the standard error of
583 TROPOMI tropospheric NO₂ VCD.

584 We calculated the GEOS-Chem air mass factor M_{gc} as Equation (B4) following Qu et al.
585 (2019).

$$586 \quad M_{gc} = \frac{\sum_{i \in \text{trop.}} c_i^{gc} \Delta p_i^{gc} w_i^{gc}}{\sum_{i \in \text{trop.}} c_i^{gc} \Delta p_i^{gc}} \quad (\text{B4})$$

587 Here c_i^{gc} is GEOS-Chem NO₂ mixing ratio at vertical layer i , Δp_i^{gc} is the pressure difference
588 between the GEOS-Chem vertical layer i and $i+1$. w_i^{gc} is the scattering weight at the GEOS-Chem
589 vertical layer i , which is calculated by the linear interpolation of the scattering weights at the
590 vertical coordinate of the model TM5 used for TROPOMI NO₂ retrieval. The scattering weight at
591 the TM5 vertical layer l (w_l^{TM5}) is computed as the product of TROPOMI air mass factor and the
592 TROPOMI averaging kernel at the TM5 vertical layer l (A_l^{TM5}) using Equation (B5) (Eskes and
593 Boersma, 2003).

$$594 \quad w_l^{\text{TM5}} = M_{obs} A_l^{\text{TM5}} \quad (\text{B5})$$

595

596 **Appendix C: Validation of the TROPOMI NO₂ observation operator**

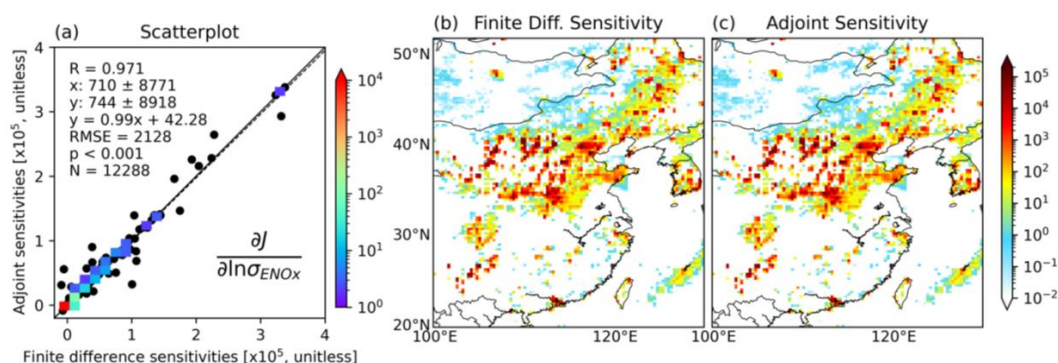
597 We validated the observation operator by comparing the sensitivity of the cost function
598 with respect to the emission scaling factor from GEOS-Chem adjoint and a finite difference
599 estimation as shown in Equation (C1). We shut down the transport and exclude a priori term from
600 the cost function for the validation, so that the gradient of cost function component in each grid
601 cell to the local emission scaling factor equals to the gradient of total cost function to the emission
602 scaling factor in the same grid cell.

$$603 \quad \frac{\partial J(\ln \sigma)}{\partial \ln \sigma} \approx \frac{J(\ln(\sigma + 0.05)) - J(\ln(\sigma - 0.05))}{\ln(\sigma + 0.05) - \ln(\sigma - 0.05)} \quad (\text{C1})$$

604 Fig. C1 compared the cost function sensitivities calculated from GEOS-Chem adjoint and
605 the finite difference method for the nested grids with the spatial resolution of 0.25°×0.3125°. The
606 spatial pattern and magnitude of the cost function sensitivities from the two methods match with



607 each other with a correlation coefficient of 0.97. The statistics show that the agreement of the
608 adjoint sensitivities and finite difference sensitivities in this study is comparable to that in Wang
609 et al. (2020a) although we constrain the NO_x emission at a much finer resolution of 0.25°×0.3125°
610 than in their study (2°×2.5°).
611



612
613 **Figure C1.** Comparison of adjoint sensitivities and finite difference sensitivities. (a) Scatter plot
614 of the adjoint sensitivity of the cost function with respect to the logarithm of NO_x emission scaling
615 factor versus the finite difference sensitivities. The color scheme for panel (a) encodes the number
616 of samples (the legend on the right of panel (a)). (b) Map of finite difference sensitivity. (c) Map
617 of adjoint sensitivity.

618
619

620 Data availability

621 The TROPOMI NO₂ and HCHO product are available at the NASA Goddard Earth
622 Sciences Data and Information Services Center (<https://daac.gsfc.nasa.gov>). The ground O₃ and
623 NO₂ measurements are available at the China National Environmental Monitoring Center
624 (<http://www.cnemc.cn/en/>).

625

626 Author contribution

627 ZL and JW designed the research, and ZL conducted the research. YW, DKH and XC
628 contributed to the research design. ZL and JW wrote the manuscript, and XC and DKH contributed
629 to the writing. YW and TS developed the codes for comparing the tropospheric NO₂ VCD from
630 model and TROPOMI data. KS developed the codes for oversampling.



631

632 **Competing interests**

633 The authors declare that they have no conflict of interest.

634

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639

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