

1 **Aggravated surface O₃ pollution primarily driven by meteorological variation**
2 **in China during the ~~e2020~~early COVID-19 pandemic lockdown period**

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30 Abstract

31 Due to the lockdown during the COVID-19 pandemic in China from late January to early April in
32 2020, a significant reduction of primary air pollutants, as compared to the same time period in
33 2019, has been identified by satellite and ground observations. However, this reduction is in
34 contrast with the increase of surface O₃ concentration in many parts of China during the same
35 period from 2019 to 2020. The reasons for this contrast are studied here from two perspectives:
36 emission changes and inter-annual meteorological variations. Based on top-down constraints of
37 NO_x emissions from TROPOMI measurements and GEOS-Chem model simulations, our analysis
38 reveals that NO_x and volatile organic compound (VOC) emission reductions as well as
39 meteorological variations lead to 8%, -3%, and 1% changes in O₃ over North China, respectively.
40 In South China, however, we find that meteorological variations cause ~30% increases in O₃,
41 which is much larger than -1% and 2% changes due to VOC and NO_x emission reductions,
42 respectively, and the overall O₃ increase in the simulations is consistent with the surface
43 observations. The higher temperature associated with increase of solar radiation and declined
44 relative humidity are the main reasons that leads to the surface O₃ increase in South China.
45 OverallCollectively, inter-annual meteorological variations have a larger impact than emission
46 reductions on the aggravated surface O₃ pollution in China during the ~~early~~ lockdown period of
47 COVID-19 pandemic.

48

49 1. Introduction

50 Surface ozone (O₃), an important air pollutant that is harmful to human health (Jerrett et
51 al., 2009) and stomatal conductance of green vegetations (Gong et al., 2020), is produced by
52 photochemical reactions of nitrogen oxides (NO_x) and volatile organic compounds (VOC) (Liu et
53 al., 1987; Sillman et al., 1990). In addition to emissions, meteorological conditions, such as
54 temperature, solar radiation and relative humidity, also have large impacts on surface O₃ formation
55 (Lu et al., 2019).

56 Ground observations show that surface O₃ increased dramatically during the COVID-19
57 lockdown period in China by around 40% on average (Tong et al., 2023) and even larger than 100%
58 (Shi and Brasseur, 2020; Liu et al., 2021) depending on the time period and region. The reduction
59 of economic activities during the lockdown period led to a significant decrease of several primary
60 air pollutants emissions. The NO₂ vertical column density (VCD) from satellite measurements and

61 surface NO₂ concentration from ground measurements were reduced by 40% - 60% in China
62 during the lockdown period (Bauwens et al., 2020; Shi and Brasseur, 2020; Liu et al., 2020a; Zhang
63 et al., 2020). A lower but discernible reduction of sulfur dioxide (SO₂), carbon monoxide (CO),
64 and formaldehyde (HCHO) have also been identified by satellite or ground-based observations in
65 China (Shi and Brasseur, 2020; Levelt et al., 2022; Ghahremanloo et al., 2021). However, during
66 this period surface O₃ concentrations increased, and the respective roles of meteorological factor
67 and emission reduction for the aggravated surface O₃ pollution during the lockdown in China need
68 to be further quantified.

69 This study provides a quantitative analysis of the causes for the unexpectedly aggravated
70 surface O₃ pollution in China during the ~~early~~ lockdown period of the pandemic from two
71 perspectives using GEOS-Chem model. One is anthropogenic emission reduction of NO_x and
72 VOC in response to the lockdown possibly under a VOC-limiting chemical regime of surface O₃
73 production (Guo et al., 2023), while the other is the impact of natural variability of meteorological
74 conditions. Previous studies have reported the enhanced surface O₃ due to NO_x emission decline
75 during the lockdown period in North China using chemical transport model (CTM) simulations
76 without controlling for the impacts of meteorological variability (Zhang et al., 2021; Huang et al.,
77 2020; Miyazaki et al., 2020). Other studies quantified or excluded the meteorological impacts on
78 the surface O₃ using statistical analysis instead of CTM that account for the physical and chemical
79 processes (Venter et al., 2020; Bi et al., 2022; Tong et al., 2023). Although a few studies have
80 investigated the contributions from both emission reduction and meteorological variability to
81 surface O₃ increase using CTMs, most of their results have uncertainties due to the limitations of
82 their analysis. For example, some of them keep the emissions unchanged (Zhao et al., 2020) or
83 assume an arbitrarily uniform emission reduction instead of constraining the emission based on
84 observations (Le et al., 2020; Liu et al., 2021). In cases where the emissions were constrained by
85 the observations, the focus was limited to several cities in China (Liu et al., 2020b). Furthermore,
86 in the past studies, the surface O₃ increase during the lockdown period of 2-4 weeks is quantified
87 in reference to the time period right before the lockdown instead of the same period in previous
88 years; such comparisons by design cannot exclude the possibility that the seasonal variation of
89 meteorology from early January to early April may have dominated the cause for the surface O₃
90 increase. A comprehensive analysis of the contributions from emission reductions and

91 meteorological variations to the surface O₃ increase during the first round of the lockdown period
92 with respect to the same time period in previous years in China is therefore overdue.

93 Here, we apply a top-down method to update NO_x and VOC emission in February and
94 March in 2020 based on the TROPOMI NO₂ and formaldehyde (HCHO) product. A set of GEOS-
95 Chem model simulations with NO_x and VOC emissions and meteorological fields in different time
96 periods are then conducted. Based on the difference in surface O₃ concentration in different
97 modeling sensitivity experiments, we quantitatively assess the respective roles of emission and
98 meteorology in regulating surface O₃ concentration in continental China. The ground observations
99 of surface O₃ and NO₂ concentration are compared with the model simulations to verify our
100 analysis. Section 2 introduces the satellite and ground-based measurements, NO_x emission update
101 scheme, and the configurations of GEOS-Chem simulation experiments. Section 3 provides an
102 evaluation of the constrained NO_x emission and surface O₃ simulations. The analysis of the
103 mechanism of the aggravated surface O₃ pollution is presented in Section 4, followed by the
104 summary and conclusions in Section 5.

105

106 **2. Datasets and Methods**

107 **2.1 TROPOMI NO₂ and HCHO product**

108 We used tropospheric NO₂ and HCHO level 2 VCD product provided by the Tropospheric
109 Monitoring Instrument (TROPOMI) onboard the Sentinel-5 Precursor (S5P) satellite (Veefkind et
110 al., 2012). S5P is a sun-synchronous polar orbit satellite launched on 13 October 2017, which
111 covers the near-global domain in a single day. TROPOMI provides NO₂ and HCHO retrievals at
112 an approximately 7 km x 3.5 km spatial resolution (5.5km x 3.5 km since 6 August 2019) from the
113 ascending orbit with an equatorial crossing time of ~13:30 local time (Van Geffen et al., 2020; De
114 Smedt et al., 2018). The datasets were obtained from the NASA Goddard Earth Sciences Data and
115 Information Services Center (<https://daac.gsfc.nasa.gov>). A quality control procedure similar to
116 Bauwens et al. (2020) but with slightly stricter criteria is adopted for TROPOMI NO₂ and HCHO
117 data. The TROPOMI retrievals under one or more than one of the following conditions are
118 screened out for data quality control. (1) Quality assurance value is no larger than 0.5; (2) cloud
119 radiance fraction within NO₂ or HCHO retrieval window is larger than 0.3; (3) solar zenith angle
120 is larger than 70°; and (4) viewing zenith angle is larger than 70°.

121

122 2.2 Ground O₃ and NO₂ measurements

123 Surface measurements of O₃ and NO₂ were collected from ~1600 operational air quality
124 monitoring stations over the mainland China managed by the China National Environmental
125 Monitoring Center (<http://www.cnemc.cn/en/>). We calculated daily maximum 8-hour average
126 (MDA8) O₃ concentration from hourly in situ measurements. Surface O₃ are measured by
127 ultraviolet photometric method and Indigo disulfonate spectrophotometry, following the national
128 environmental standards of HJ 590-2010 and HJ 504-2009. Surface NO₂ concentrations are
129 measured by the chemiluminescence method (Zhang and Cao, 2015) that quantifies the NO₂
130 concentrations by measuring the NO decomposed from NO₂, which can cause a positive bias in
131 the NO₂ measurements (Steinbacher et al., 2007) because NO_z (compounds produced from the
132 atmospheric oxidation of NO_x) can also be decomposed to NO. The true NO₂ concentrations only
133 account for 43%-76% and 70%-83% of measured values for rural and urban sites (Steinbacher et
134 al., 2007). Following Wang et al. (2020b), we also applied a correction factor but with a lower
135 value of 0.75 to the measured NO₂, considering that we included both rural and urban sites. The
136 sampling ports are placed at 3 to 15 meters above the ground following the national environmental
137 monitoring method standard HJ 664-2013. The measured data are reported in the unit of μg m⁻³
138 under standard temperature (273.15 K) and pressure (101.325 kPa) according to national
139 environmental standards GB 3095-2012.

140

141 2.3 GEOS-Chem model and its adjoint

142 The global 3-D chemical transport model GEOS-Chem (Bey et al., 2001) version 12.7.2 is
143 used here. We apply the nested-grid version of GEOS-Chem (Chen et al., 2009; Wang et al., 2004)
144 with the horizontal resolution of 0.25°×0.3125° and 47 vertical hybrid-sigma levels over East Asia
145 (70°E-140°E, 15°N-55°N). The boundary conditions are obtained from the 2°×2.5° global
146 simulation. The model is driven by the GEOS-FP meteorological field provided by NASA Global
147 Modeling and Assimilation Office (GMAO). A detailed O₃-NO_x-hydrocarbon chemistry (Mao et
148 al., 2010; Mao et al., 2013; Travis et al., 2016) is included in the GEOS-Chem model. The altitude
149 of the surface O₃ output from GEOS-Chem is specified at 9 meters above the ground to match the
150 in-situ measurements (Travis et al., 2017; Zhang et al., 2012). Through our sensitivity test using
151 GEOS-Chem, the variation of surface O₃ from 3 to 9 meters above the surface is generally less
152 than 0.723 ppb (75th percentile), and the median bias is 0.283 ppb. Travis et al. (2017) reported

153 from 60 m to 10 m above the ground, the MDA8 O₃ could decrease by ~3 ppb. Therefore, when
154 comparing GEOS-Chem surface O₃ with in-situ measurements, the differences caused by
155 inconsistent reported altitudes (9 m versus 3-15 m) can be ignored.

156 The global anthropogenic emission used in GEOS-Chem model is the Community
157 Emissions Data System (CEDS) inventory (Hoesly et al., 2018), which is replaced by the MIX
158 inventory (Li et al., 2017) over the Asian region. Biogenic emissions for VOCs follows the Model
159 of Emissions of Gases and Aerosols from Nature (MEGAN) inventory (Guenther et al., 2012).
160 Natural NO_x emissions includes biomass burning from GFED4 inventory (Van Der Werf et al.,
161 2017), soil NO_x emissions (Hudman et al., 2012) and lightning sources (Murray et al., 2012; Ott
162 et al., 2010).

163 The adjoint of the GEOS-Chem model (Henze et al., 2007; Henze et al., 2009) is a
164 component of the 4D-Var inversion method that can efficiently optimize spatially disaggregated
165 aerosol and gas emissions. This is done through iterative minimization of a cost function using the
166 model adjoint to calculate the gradient of the cost function with respect to a large number of model
167 parameters (such as anthropogenic NO_x emissions in each grid box) simultaneously. The cost
168 function is the sum of the error weighted difference between forward model outputs and
169 observations and the divergence of posterior model parameters from the prior estimate (Section
170 2.4). We developed and validated the observation operator for TROPOMI NO₂ in the GEOS-Chem
171 adjoint model version 35n similar to Wang et al. (2020a) and used it to optimize the anthropogenic
172 NO_x emission during the lockdown period in China. The [monthly](#) NO_x emission optimization is
173 implemented using the 4D-Var method with GEOS-Chem adjoint at the nested grid with the
174 resolution of 0.25°×0.3125° [via assimilating the daily TROPOMI NO₂ measurements](#). The prior
175 anthropogenic NO_x emission used in the GEOS-Chem adjoint is HTAP version 2 (Janssens-
176 Maenhout et al., 2015), which is equivalent to the MIX inventory in East Asia (Li et al., 2017).

177

178 **2.4 NO_x and VOC emission updates**

179 Two approaches are used to update the emissions during the lockdown period in 2020. The
180 first is a simple mass balance approach (Leue et al., 2001; Martin et al., 2003; Vinken et al., 2014)
181 for updating the NO_x emission by assuming a constant NO_x lifetime and NO_x/NO₂ ratio. In the
182 period from 2010 to 2019, the anthropogenic NO_x emissions have declined significantly as a result
183 of the clean air actions of Chinese government (Zheng et al., 2018). We scale the anthropogenic

184 NOx emission from year 2010 to 2019 using the spatially gridded ratio of mean TROPOMI
 185 tropospheric NO₂ VCD in Feb.-Mar. 2019 to GEOS-Chem simulated NO₂ column with default
 186 MIX 2010 emission (Appendix A), to obtain the baseline anthropogenic NOx emission in 2019,
 187 which is denoted as MIX 2019. To derive anthropogenic NOx emissions in 2020 in China during
 188 the COVID-19 lockdown (MIX 2020), the spatially gridded ratio of mean TROPOMI tropospheric
 189 NO₂ VCD in 2020 Feb.-Mar. to that in 2019 Feb.-Mar. is taken as a scaling factor for the updated
 190 baseline anthropogenic NOx emission in 2019 (MIX 2019). The two-month mean of TROPOMI
 191 NO₂ VCD in 2019 and 2020 are calculated with the physical oversampling procedure (Sun et al.,
 192 2018). Scaling factors in regions where mean TROPOMI tropospheric NO₂ VCD in 2019 Feb.-
 193 Mar. is less than 0.1 Dobson unit (DU) are set to 1 for emission updates in both 2020 and 2019,
 194 assuming that the lockdown only affects the populated areas (that have high NO₂ in 2019).

195 The second method for updating NOx emission is 4D-Var via the GEOS-Chem adjoint
 196 model. The anthropogenic NOx emissions in 2020 lockdown period derived from the GEOS-Chem
 197 adjoint is denoted as 2020 Adjoint. Following Wang et al. (2020a), the cost function J for
 198 optimizing the NOx emission is defined as

$$199 \quad J = \frac{1}{2} \sum_{\mathbf{c} \in \Omega} [H(\mathbf{c}) - \mathbf{s}]^T \mathbf{S}_{\text{obs}}^{-1} [H(\mathbf{c}) - \mathbf{s}] + \frac{1}{2} \gamma (\boldsymbol{\sigma} - \boldsymbol{\sigma}_a)^T \mathbf{S}_a^{-1} (\boldsymbol{\sigma} - \boldsymbol{\sigma}_a) \quad (1)$$

200 where \mathbf{s} is the tropospheric slant column density of TROPOMI NO₂, which is the product of
 201 TROPOMI NO₂ VCD and air mass factor. H is the TROPOMI NO₂ observation operator that maps
 202 the modeled NO₂ concentrations \mathbf{c} to the observations in time and space and calculates the
 203 corresponding slant column density to make an apple-to-apple comparison of the model to
 204 TROPOMI. Ω is the spatial and temporal domain where both model simulations and observations
 205 are available. $\boldsymbol{\sigma}$ is the scaling factor of anthropogenic NOx emissions to be optimized, and $\boldsymbol{\sigma}_a$ is
 206 the prior emission scaling factors, which equals 1. \mathbf{S}_{obs} and \mathbf{S}_a are observational and prior error
 207 covariance matrices, respectively. γ is the regularization factor that balances the weights of the
 208 observational term and prior term. We assumed \mathbf{S}_{obs} to be diagonal following Wang et al. (2020a)
 209 with the diagonal values calculated as the square of the standard error of tropospheric NO₂ slant
 210 column density from the TROPOMI product. The prior error of the NOx emissions is assumed to
 211 be 100%. The spatial correlation of NOx emissions is considered in this study, and the off-diagonal
 212 elements of \mathbf{S}_a are computed by assuming an exponentially decaying error correlation with a fixed

213 decaying distance of 150 km following Qu et al. (2017). The γ value was determined as 500 via
214 the total error minimization and L-curve test (Henze et al., 2009; Qu et al., 2017).

215 We developed the observation operator for TROPOMI NO₂ product in the GEOS-Chem
216 adjoint model with GEOS-Chem NO₂ vertical profiles and TROPOMI NO₂ averaging kernel
217 applied to minimize the discrepancies between the assumptions in TROPOMI NO₂ retrieval and
218 GEOS-Chem model simulation. See Appendix B for additional details. The observation operator
219 has been validated using the finite difference method (Appendix C).

220 For anthropogenic VOC emissions update, we only applied the mass balance method based
221 on the TROPOMI HCHO data. The default anthropogenic VOC emissions used in the GEOS-
222 Chem is also MIX 2010 (Li et al., 2017). We ignore the change of anthropogenic VOC emissions
223 from 2010 to 2019 ([Appendix D](#)). The baseline VOC emission in 2019 (MIX 2019) is identical to
224 that of MIX 2010. The updated anthropogenic VOC emissions during the lockdown period is
225 denoted as MIX 2020. HCHO is one species of VOC and may not be able to represent other VOC
226 species. Different from NO_x, biogenic sources, meteorological impacts, and large retrieval
227 uncertainty of HCHO due to its low optical depth prevent accurately quantifying the emission
228 decline due to lockdown from satellite retrievals (Levelt et al., 2022). Vigouroux et al. (2020)
229 [reported that TROPOMI HCHO tends to be overestimated by ~26% for HCHO column lower than](#)
230 [2.5×10¹⁵ molecules cm⁻² and underestimated by ~31% for HCHO column higher than 8.0×10¹⁵](#)
231 [molecules cm⁻²](#). To optimize the signal, we spatially aggregate the ratio of TROPOMI HCHO [in](#)
232 [2020 Feb.-Mar. during the lockdown period](#) to that [in 2019 Feb.-Mar. before the lockdown](#) to the
233 resolution of 0.5°, which are used as the scaling factors for updating the anthropogenic VOC
234 emissions during the lockdown period. The aggregation is based on the oversampling of
235 TROPOMI HCHO at 0.01° resolution, and the ratio is computed as the mean of the lowest 25th
236 percentile of all ratios at 0.01° resolution in each 0.5°×0.5° grid box, which ensures that only
237 statistically significant changes are considered. We assumed the change of anthropogenic VOC
238 emissions over sparsely populated areas (TROPOMI NO₂ in 2019 Feb.-Mar. less than 0.1 DU) is
239 insignificant and assigned the ratio values as one. To further evaluate the uncertainties associated
240 with this approach, we also conducted sensitivity study by using different threshold in the
241 aggregation.

242 We assess the results from model experiments (as described in Section 2.5) adopting the
243 updated NO_x emission by comparing mean tropospheric NO₂ VCD from GEOS-Chem and from

244 TROPOMI observations in Feb.-Mar. of 2019 and 2020. The averaging kernel of TROPOMI NO₂
 245 is applied to modeled NO₂ column for this comparison, following Sha et al. (2021). Further
 246 quantitative evaluation of the model results also used the [TROPOMI measurements of HCHO and](#)
 247 surface observation of O₃ and NO₂.

248

249 2.5 GEOS-Chem model experiments

250 A series of sensitivity experiments is conducted over China with different NO_x and VOC
 251 emissions and GEOS-FP meteorological fields in different years using GEOS-Chem (v12.7.2)
 252 model. All simulations are conducted from Jan. 15 to Mar. 31. The 17 days before Feb. 1 are used
 253 for spin up, and the model output for Feb. and Mar. are used for the analysis. The configurations
 254 of different simulations are listed in Table 1.

255

256 **Table 1.** Configurations of model sensitivity experiments.

Experiments	Abbreviation	Meteorology	NO _x Emission	VOC Emission
Baseline (2019)	2019B	GEOS-FP 2019	MIX 2019	MIX 2019
2020 Default	2020D	GEOS-FP 2020	MIX 2019	MIX 2019
2020 NO_x	2020N	GEOS-FP 2020	MIX 2020	MIX 2019
2020 VOC	2020V	GEOS-FP 2020	MIX 2019	MIX 2020
2020 Lockdown	2020L	GEOS-FP 2020	MIX 2020	MIX 2020
2020 Adjoint	2020A	GEOS-FP 2020	Adjoint 2020	MIX 2020

257

258 We use the following equations to quantify the contributions from NO_x and VOC emission
 259 reduction due to COVID-19 and meteorological variation to the increase of surface O₃.

$$260 \quad \Delta O_3^{\text{NO}_x} = \frac{O_3^{2020A} - O_3^{2020V}}{O_3^{2019B}} \times 100\% \quad (2)$$

$$261 \quad \Delta O_3^{\text{VOC}} = \frac{O_3^{2020L} - O_3^{2020N}}{O_3^{2019B}} \times 100\% \quad (3)$$

$$262 \quad \Delta O_3^{\text{ems}} = \frac{O_3^{2020A} - O_3^{2020D}}{O_3^{2019B}} \times 100\% \quad (4)$$

$$263 \quad \Delta O_3^{\text{met}} = \frac{O_3^{2020D} - O_3^{2019B}}{O_3^{2019B}} \times 100\% \quad (5)$$

264 Where ΔO_3^{NOx} , ΔO_3^{VOC} and ΔO_3^{ems} are the relative differences in surface O_3 concentration caused
 265 by emission decline of NOx, VOC, and both NOx and VOC resulting from COVID-19. ΔO_3^{met}
 266 represents the relative contribution to the surface O_3 change from the meteorological variation
 267 between 2 years. O_3^{2019B} , O_3^{2020D} , O_3^{2020N} , O_3^{2020V} , O_3^{2020L} and O_3^{2020A} are mean MDA8 surface
 268 O_3 concentration simulated by modeling experiments Baseline (2019), 2020 Default, 2020 NOx,
 269 2020 VOC, 2020 Lockdown and 2020 Adjoint, respectively (Table 1).

270 The difference in simulated surface O_3 between 2020 and 2019, is the result of both
 271 emission reductions and meteorological variations and is denoted as ΔO_3^{all} . It is calculated as
 272 follows and is evaluated against the observed relative difference of mean MDA8 O_3 in Feb. to Mar.
 273 between 2019 and 2020 at all ground sites:

$$\Delta O_3^{\text{all}} = \frac{O_3^{2020A} - O_3^{2019B}}{O_3^{2019B}} \times 100\% \quad (6)$$

275

276 **3. Results of model development, emissions, and validation**

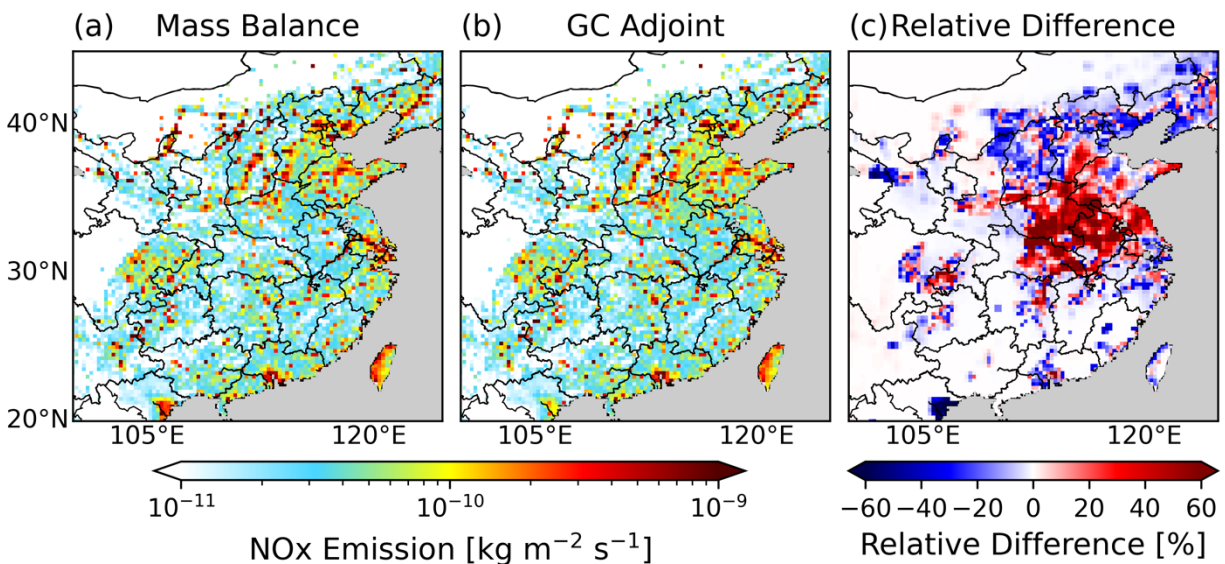
277 **3.1 Changes of NOx and VOC emissions during COVID**

278 We updated the anthropogenic NOx emissions during the COVID lockdown using both
 279 4D-Var and mass balance methods (Fig. 1 and 2). The NOx emissions from the 4D-Var inversion
 280 share a similar spatial pattern and magnitude with those found using the mass balance method (Fig.
 281 1). However, the NOx emissions from the 4D-Var inversion are lower overall than those from the
 282 mass balance method over North China by ~10% and larger over central China by ~40%. Fig. 2(a-
 283 b) shows that the 4D-Var NOx emission reduction is more severe over urban regions and displays
 284 a smoother spatial pattern than that from the mass balance approach, which is caused by the
 285 arbitrary cut off with 0.1 DU of NO₂ VCD in the latter. Furthermore, the 4D-Var inversion
 286 captured the NOx emission decline in Northeast China where the mass balance approach did not
 287 because of the low NO₂ VCD. During Feb.-Mar. 2020, the anthropogenic NOx emissions in East
 288 China decreased by ~30% compared to those in the same period in 2019.

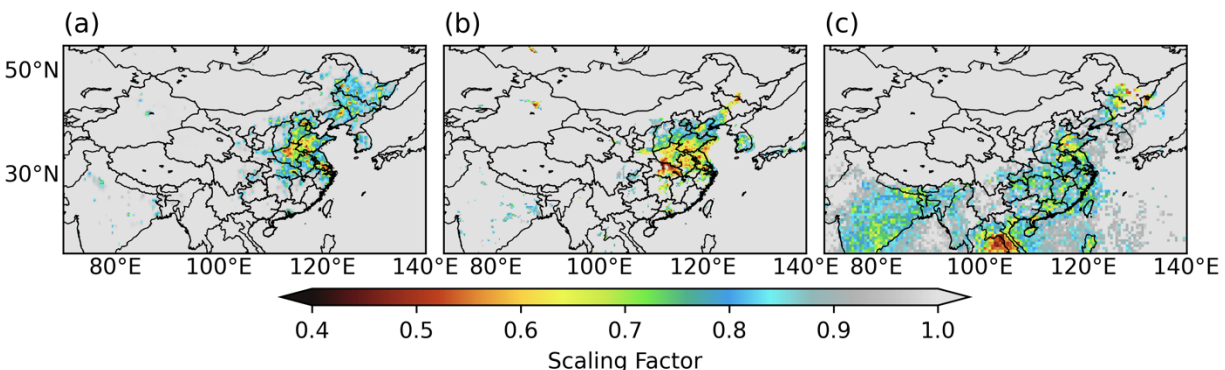
289 We also scale the anthropogenic VOC emissions based on the TROPOMI HCHO data (Fig.
 290 2(c)). The VOC emissions decrease by ~20%-30% in East and South Asia. The anthropogenic
 291 VOC emission changes in sparsely populated areas over Northwest China are neglected.

292 TROPOMI HCHO data cannot distinguish the anthropogenic emissions from biogenic and
 293 biomass burning sources for the Indochinese Peninsula in Southeast Asia because of the dense

294 vegetation in this region. However, this study investigated the O₃ pollution in China, the Southeast
 295 Asia with the dense vegetation is out of our study domain. The impact of VOC emission bias in
 296 Southeast Asia on the surface O₃ pollution in China is negligible considering the lifetime of
 297 biogenic VOC is generally short (Atkinson, 2000). For the populated urban regions in China,
 298 where the surface O₃ pollution exerts more significant health impacts, the anthropogenic source
 299 dominates the VOC emissions (Williams and Kopppmann, 2007).



300
 301 **Figure 1.** Updated anthropogenic NOx emission during Feb.-Mar. 2020 from (a) mass balance
 302 method, (b) 4D-Var method and (c) their relative difference. The relative difference is calculated
 303 as dividing the difference of (b) minus (a) by (a).



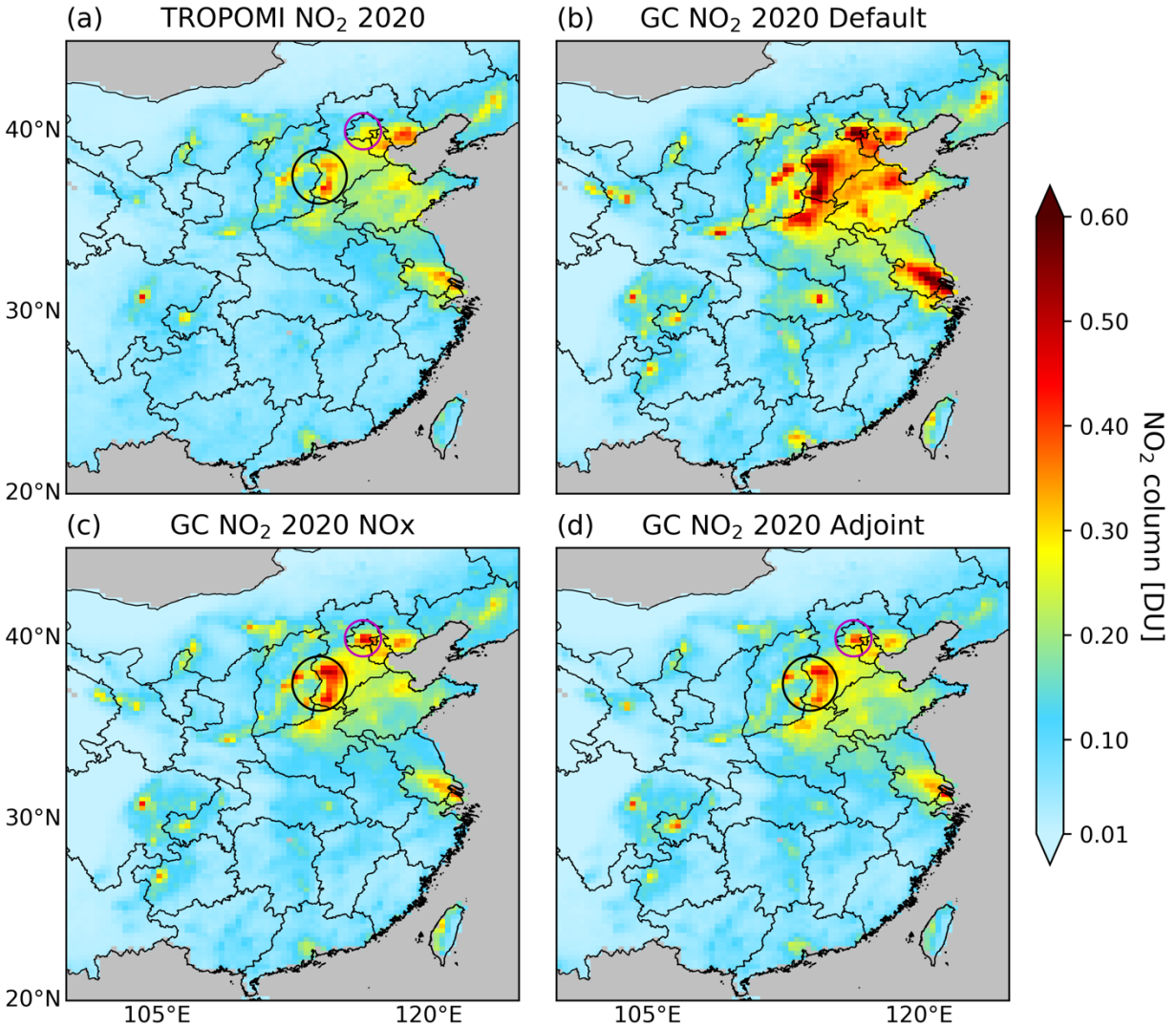
305
 306 **Figure 2.** Scaling factors for anthropogenic NOx emission in Feb.-Mar. from 2019 to 2020 as
 307 derived from (a) 4D-Var, (b) mass balance. Scaling factors for anthropogenic VOC emissions from
 308 the mass balance are in (c).

309

310 **3.2 Validation of NO₂ simulations**

311 We further assess our updated anthropogenic NO_x emissions by comparing the NO₂ VCD
312 from TROPOMI with that from GEOS-Chem with the anthropogenic NO_x emissions before and
313 after the scaling (Fig. 3 and 4). Before updating the NO_x emissions, the 2020 Default (Fig. 3(b))
314 simulation significantly overestimates the NO₂ VCD compared to the TROPOMI NO₂
315 observations (Fig. 3(a)). With the NO_x emissions updated, the simulations 2020 NO_x (Fig. 3(c))
316 and 2020 Adjoint (Fig. 3(d)) exhibit a much better agreement with TROPOMI NO₂ observation
317 than 2020 Default. However, Fig. 3(c) shows the GEOS-Chem simulation with the NO_x emissions
318 from mass balance approach overestimated the NO₂ VCD over Beijing and southwest of Hebei
319 Province (pink and black circles in Fig.3) compared with TROPOMI data. The reason is that
320 scaling factors are applied only to anthropogenic NO_x emissions, not total NO_x emissions, so it is
321 expected that the model may still overestimate the NO₂ column after scaling part of the total NO_x
322 emission. With the anthropogenic NO_x emissions optimized by the 4D-Var method, the
323 overestimation of NO₂ VCD over Beijing and southwest of Hebei Province (pink and black circles
324 in Fig. 3) is mitigated compared with the NO_x emissions from mass balance approach.

325 Fig. 4 further displays the statistics for the comparison between the TROPOMI NO₂ and
326 GEOS-Chem simulations via the scatterplot. The Baseline (2019) simulation captures the
327 magnitude of NO₂ VCD observations in 2019 well (Fig. 4(a)). The root-mean-square-error (RMSE)
328 and mean bias error (MBE) for the simulation with 2020 NO_x emission derived from mass balance
329 method (Fig. 4(b)) decreased by 0.050 DU and 0.057 DU as compared to the 2020 Default (Fig.
330 4(c)). Compared with the result from GEOS-Chem simulation 2020 NO_x, emissions from 2020
331 Adjoint (Fig. 4(d)) further led to the reduction of the MBE of the NO₂ VCD by 0.006 DU and
332 improve the correlation coefficient by 0.003. The significant overestimation of several pixels with
333 TROPOMI NO₂ VCD larger than 0.4 DU by the simulation 2020 NO_x is also mitigated by 2020
334 Adjoint. The MBE between GEOS-Chem and TROPOMI for Baseline (2019), 2020 NO_x and
335 2020 Adjoint are -0.004, 0.015 and 0.009 DU, respectively. The corresponding relative bias are
336 1.9%, 10% and 6.0%, which are all less than the relative uncertainty of ~30% for TROPOMI
337 tropospheric NO₂ VCD over East China (Van Geffen et al., 2022). The improved agreement
338 between the simulation with updated NO_x emission and TROPOMI NO₂ provides a basis for
339 further analyzing the mechanism of aggravated surface O₃ pollution.



340

341 **Figure 3.** Comparison of tropospheric NO₂ VCD from (a) TROPOMI product in 2020 Feb.-Mar

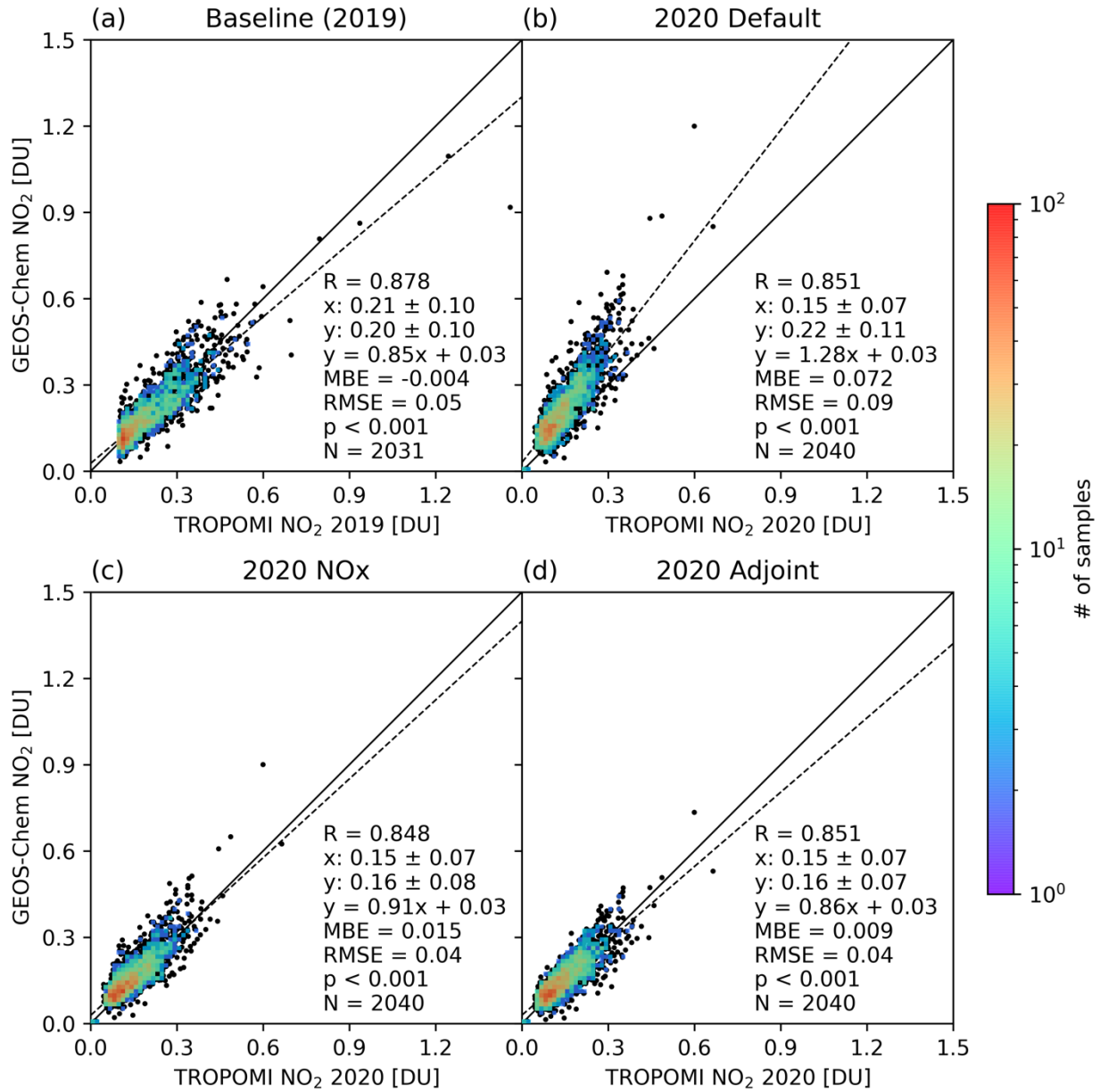
342 with that from GEOS-Chem simulations (b) 2020 Default, (c) 2020 NO_x and (d) 2020 Adjoint.

343 The pink and black circles mark the areas where NO_x emissions from 4D-Var mitigated the NO₂

344 overestimation by mass balance method. The emissions and meteorology configurations for

345 GEOS-Chem simulations 2020 Default, 2020 NO_x and 2020 Adjoint are listed in Table 1.

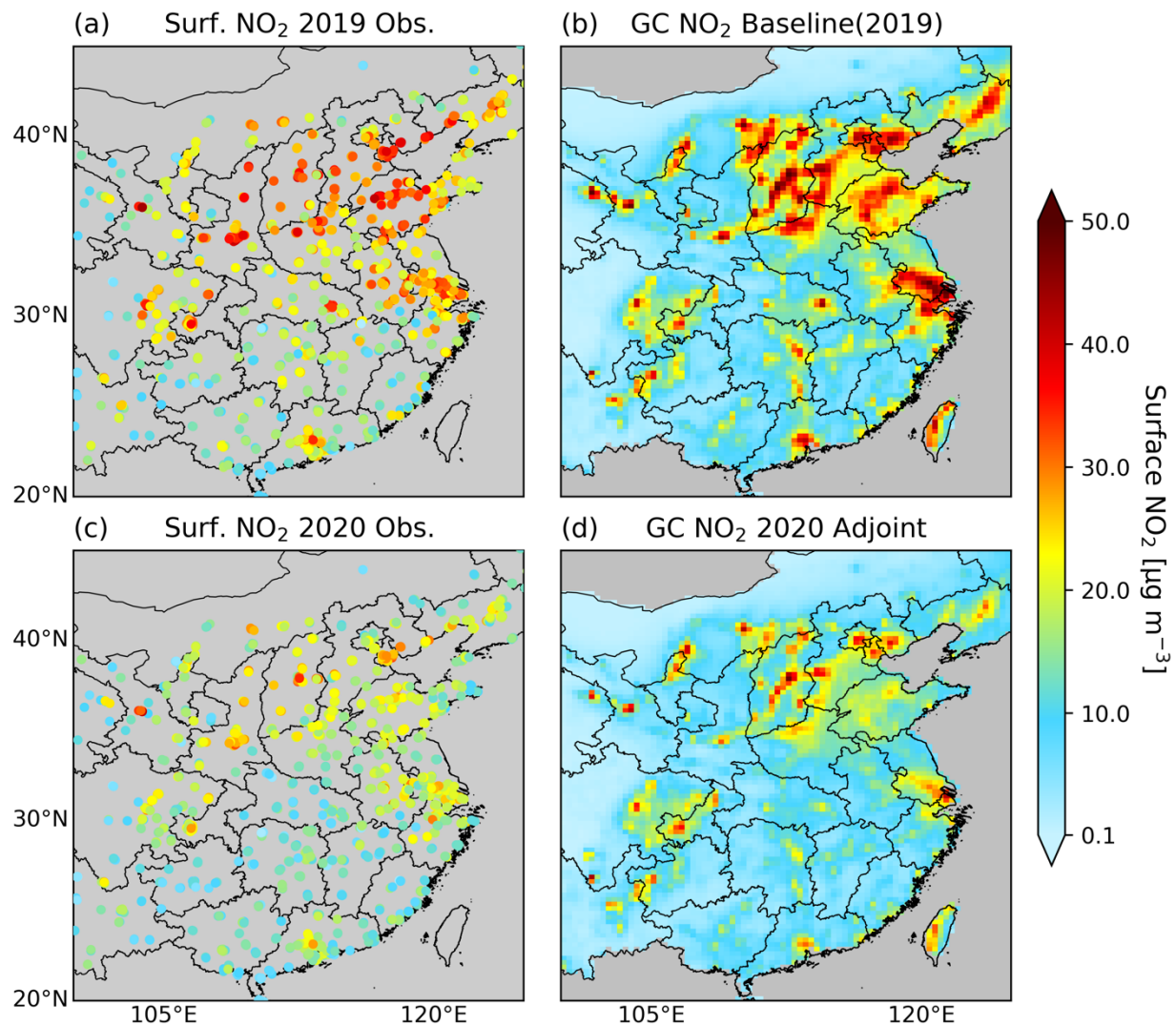
346



347
 348 **Figure 4.** Scatter plot of TROPOMI NO₂ VCD versus the GEOS-Chem simulations for (a)
 349 Baseline (2019), (b) 2020 Default, (c) 2020 NO_x and (d) 2020 Adjoint, respectively. TROPOMI
 350 data in Feb.-Mar. of 2019 was used in (a), and that of 2020 was used in (b-d). The emissions and
 351 meteorology configurations for GEOS-Chem simulations are listed in Table 1. Only pixels with
 352 TROPOMI NO₂ VCD in 2019 Feb.-Mar. larger than 0.1 DU are included in all comparisons.

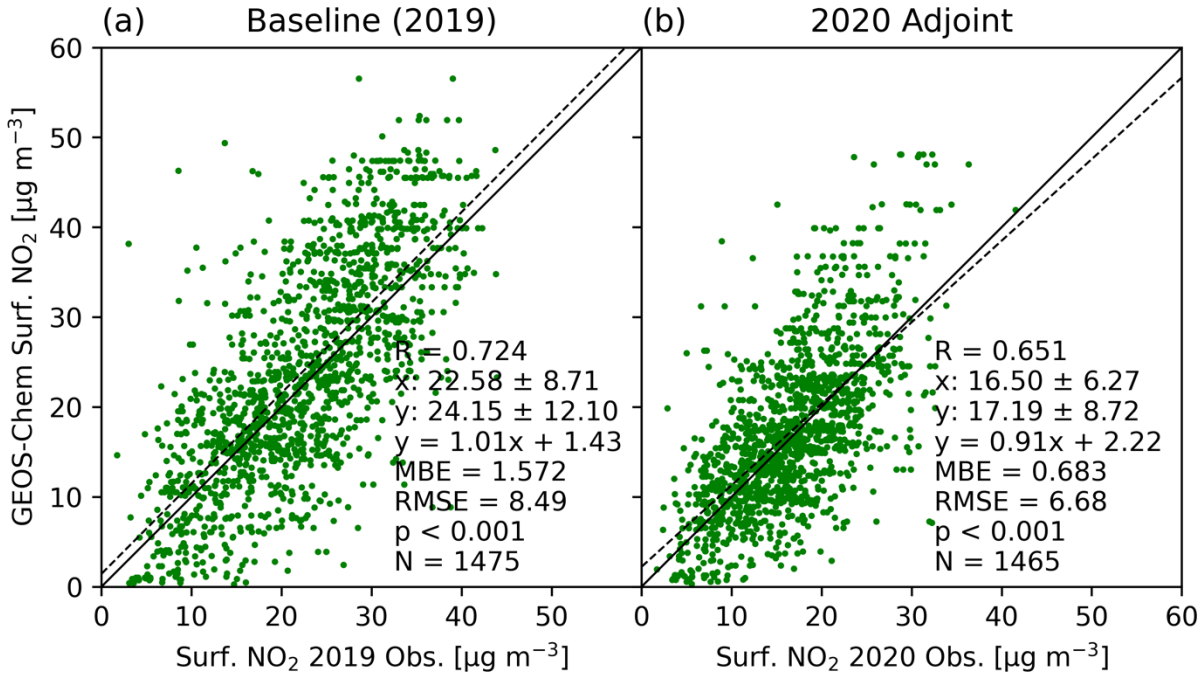
353
 354

355 Fig. 5 and Fig. 6 show the comparison of surface NO₂ between ground measurements and
356 GEOS-Chem simulations. The GEOS-Chem simulations Baseline (2019) (Fig. 5(b)) and 2020
357 Adjoint (Fig. 5(d)) both capture the spatial pattern and magnitude of surface NO₂ measurements
358 in Feb.-Mar. of 2019 (Fig. 5(a)) and 2020 (Fig. 5(c)) well, respectively. Fig. 6 further displays the
359 good agreements of surface NO₂ from Baseline (2019) (Fig. 6(a)) and 2020 Adjoint (Fig. 6(b)) to
360 the in-situ measurements via scatter plots. Table 2 displays the evaluation statistics, including the
361 correlation coefficient (R), MBE, RMSE and the slope and intercept of the linear regression, for
362 the simulated surface NO₂ from various simulation experiments compared with the in-situ
363 measurements. The correlation coefficient, MBE and RMSE between the simulation Baseline
364 (2019) and ground measurements in 2019 Feb.-Mar. are 0.724, 1.572 μg m⁻³ and 8.49 μg m⁻³,
365 respectively. Without updating the NO_x emissions in 2020, the simulation 2020 Default
366 overestimate the ground measurements of surface NO₂ in 2020 Feb.-Mar (Table 2). The slope for
367 the linear regression is 1.19, and the MBE and RMSE are 6.021 μg m⁻³ and 10.43 μg m⁻³,
368 respectively (Table 2). After updating the NO_x emissions, the GEOS-Chem simulations 2020 NO_x
369 and 2020 Adjoint have good agreements with the in-situ measurements in 2020 Feb.-Mar. The
370 correlation coefficient between the simulation 2020 Adjoint versus the in-situ measurements is
371 0.651, higher than that of 0.608 for the simulation 2020 NO_x versus the ground measurements
372 (Table 2). The MBE and RMSE of 2020 Adjoint (0.683 μg m⁻³ and 6.68 μg m⁻³) are lower than
373 those of 2020 NO_x (1.726 μg m⁻³ and 7.74 μg m⁻³) (Table 2). This result further indicates the
374 superiority of 4D-Var for optimizing NO_x emissions compared with the mass balance method
375 (Cooper et al., 2017; Streets et al., 2013).
376



377
 378 **Figure 5.** Comparison of surface NO₂ concentrations from ground measurements for (a) 2019
 379 Feb.-Mar. and (c) 2020 Feb.-Mar. versus those from GEOS-Chem simulations (b) Baseline (2019)
 380 and (d) 2020 Adjoint. Grey color means no data is presented.

381



382

383 **Figure 6.** Scatter plots for comparing the surface NO₂ concentrations from GEOS-Chem
 384 simulations and ground measurements. (a) GEOS-Chem simulation Baseline (2019) versus ground
 385 measurements in 2019 Feb.-Mar. (b) GEOS-Chem simulation 2020 Adjoint versus ground
 386 measurements in 2020 Feb.-Mar. Note: the number of ground sites differ in these two years.

387

388

389 **Table 2.** Evaluation statistics for modeled surface NO₂ compared with the in-situ measurements*.

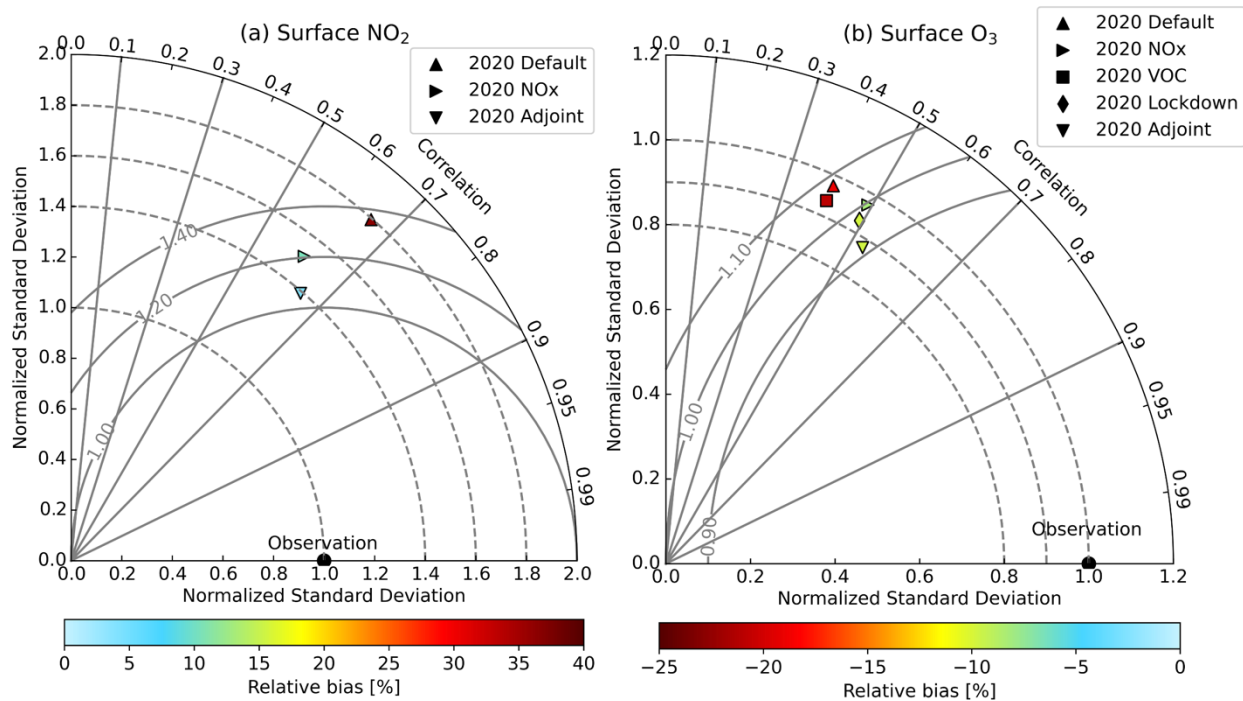
Experiments	R	MBE ($\mu\text{g m}^{-3}$)	RMSE ($\mu\text{g m}^{-3}$)	Slope	Intercept ($\mu\text{g m}^{-3}$)
Baseline (2019)	0.724	1.572	8.49	1.01	1.43
2020 Default	0.661	6.021	10.43	1.19	2.95
2020 NOx	0.608	1.726	7.74	0.92	3.03
2020 Adjoint	0.651	0.683	6.68	0.91	2.22

390 * The simulation experiment Baseline (2019) is compared with the ground measurements in 2019
 391 Feb.-Mar. Other three experiments are compared with the ground measurements in 2020 Feb.-Mar.

392

393 Fig. 7 (a) is the Taylor diagram for evaluating the GEOS-Chem simulations of surface NO₂
 394 concentrations from 2020 Default, 2020 NOx and 2020 Adjoint using the in-situ measurements.

395 The simulation 2020 Adjoint (inverted triangle in Fig. 7(a)) has the best performance among these
 396 three simulations with the lowest relative bias and lowest normalized centered RMSE. Without
 397 updating the NOx emission, 2020 Default features a relative bias of ~37%. After updating the NOx
 398 emissions, 2020 NOx reduces the relative bias, normalized centered RMSE and normalized
 399 standard deviation from around 37%, 1.38 and 1.80 to around 10%, 1.20 and 1.51 compared with
 400 2020 Default, but the correlation coefficient also decreases. By using 4D-Var method, 2020
 401 Adjoint further reduces the relative bias, normalized centered RMSE and normalized standard
 402 deviation and increases the correlation coefficient compared with 2020 NOx. We also validated
 403 the VOC emissions by comparing the simulated HCHO VCD with TROPOMI measurements
 404 (Appendix D).



406
 407 **Figure 7.** Taylor diagram for evaluating the GEOS-Chem simulations of (a) surface NO₂ and (b)
 408 surface O₃ during lockdown period (2020 Feb.-Mar.) using ground observations for different
 409 simulation experiments listed in Table 1. The evaluation of surface O₃ only includes the areas
 410 where the NOx emissions optimized by 4D-Var reduced by more than 10%.

411

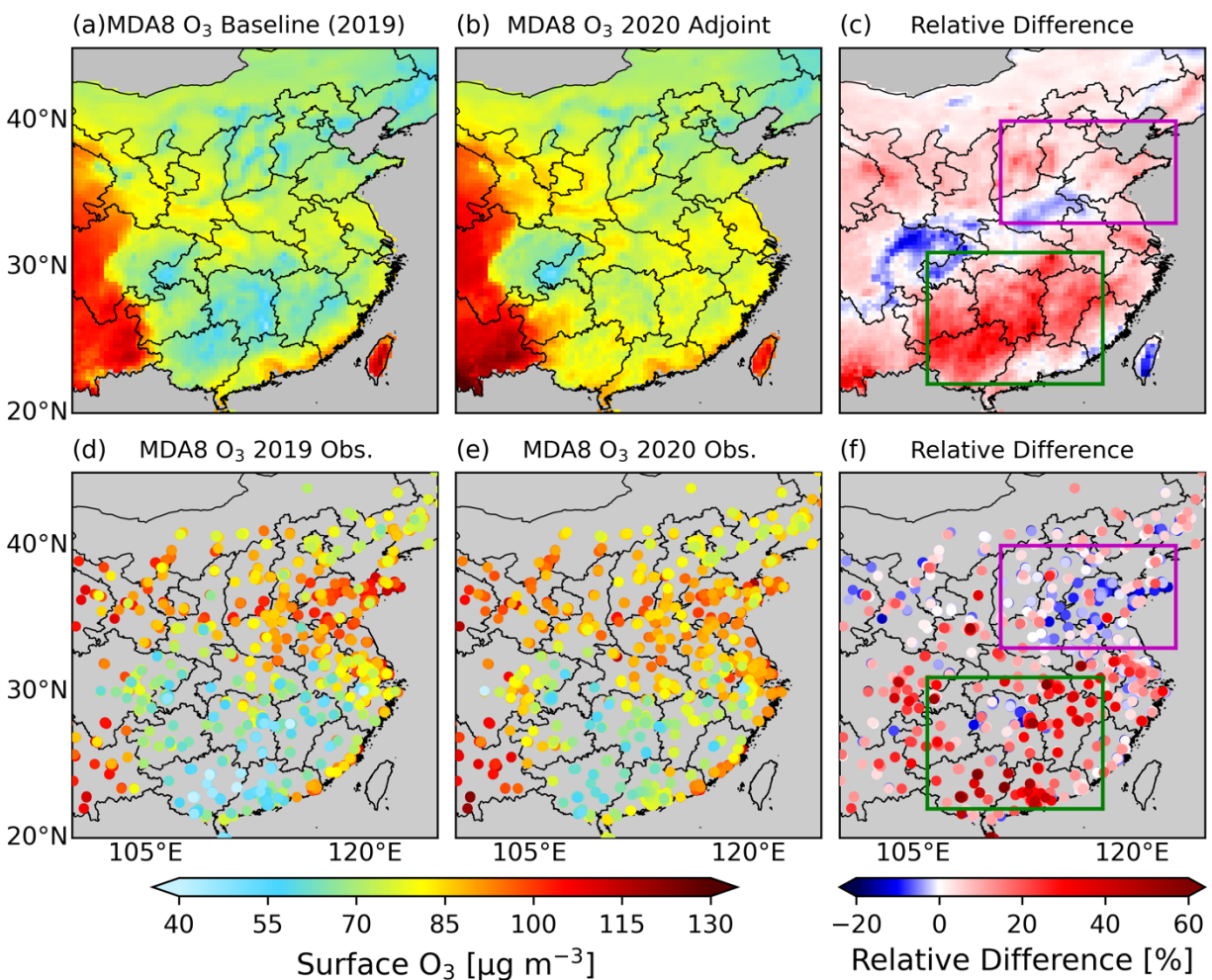
412

413 3.3 Evaluation of surface O₃ simulations

414 We evaluated the GEOS-Chem simulations of MDA8 surface O₃ from different simulation
415 experiments listed in Table 1 using ground measurements. Fig. 7(b) is the Taylor diagram for
416 comparing the surface O₃ concentrations during 2020 Feb.-Mar. from ground measurements and
417 GEOS-Chem simulations. We focused on areas with significant NO_x emissions reduction to better
418 assess the role of updated NO_x emissions in improving surface O₃ simulations. The ground sites
419 where the NO_x emissions from 4D-Var decline by less than 10% are excluded. The correlation
420 coefficient between the simulation Baseline (2019) and ground observations is ~0.53 and the
421 relative bias is around -25%. By applying 2020 meteorological fields and scaling the VOC
422 emissions, the correlation coefficients decreased to ~0.40 for model simulations 2020 Default and
423 2020 VOC, with little reduction in the relative bias. By updating the NO_x emissions, the relative
424 bias reduced to around -10% while the correlation coefficients remain at ~0.50 for model
425 simulations 2020 NO_x, 2020 Lockdown and 2020 Adjoint. This indicates the NO_x emission
426 updates significantly improve the surface O₃ simulations. Comparing the simulations 2020 Default
427 and 2020 VOC, or 2020 NO_x and 2020 Lockdown, the results show that scaling VOC emissions
428 does not improve the surface O₃ simulations significantly over the continental China, but over
429 South China, VOC emissions update reduces the relative bias by 3%. Among all simulations, 2020
430 Adjoint exhibits the best performance with the lowest normalized centered RMSE, largest
431 correlation coefficient and a low relative bias of ~10%. This result further confirms the superiority
432 of the 4D-Var with respect to the mass balance method for optimizing NO_x emissions. Therefore,
433 we used the 2020 Adjoint to evaluate the impacts of NO_x emission on surface O₃ in the following
434 analysis.

435 Fig. 8 compares the modeled surface O₃ in Feb.-Mar. of 2019 (Fig. 8(a)) and 2020 (Fig.
436 8(b)) and the relative difference (Fig. 8(c)) computed from Equation (6) with the in-situ
437 measurements (Fig. 8(d-f)). The ground observations show that the highest level of surface O₃
438 pollution occurs in North China and southwest of China. The average MDA8 O₃ in two months
439 can reach up to ~110 μg m⁻³ at STP (~51.4 ppbv), which is higher than the China National Ambient
440 Air Quality Standard daily maximum 8-hour Grade I standard of 100 μg m⁻³. GEOS-Chem model
441 underestimates the surface O₃ over North China for both years compared with ground observations,
442 which could be a result of the underestimation of biogenic out-of-date VOC emissions (Appendix
443 D). The underestimation of the simulated O₃ over North China will not significantly affect our
444 study purpose since this study focuses on revealing the impacts of emissions and meteorology

445 change on the surface O₃ change by each region. The bias is predominantly systematic and will be
446 substantially cancelled when we compute the relative difference of the surface O₃. ~~The model~~
447 ~~it~~ captures the magnitude and spatial distribution of surface O₃ and the increasing trend in South
448 China well. In South China, the measured surface O₃ in 2020 Feb.-Mar. increases by 30-50%,
449 while over North China, it increases generally by less than 20% even decreases in some regions.
450 The relative differences of simulated surface O₃ between two years is comparable to the ground
451 observations over South China (green box in Fig. 8(c, f)). In Sichuan Basin, the trend of the surface
452 O₃ change from the model is opposite to the measurements, which is probably caused by the
453 inaccurate simulation of the meteorological effects (See Section 4.1) due to the complex terrain
454 features in this region. The localized improvement of the model simulation is further needed for
455 the regional study. Over North China (pink box in Fig. 8(c, f)), the average relative difference
456 between two years from the model and observation are 4.27% and -3.01%, respectively, both of
457 which are much smaller than their counterparts in South China. While the relative difference from
458 model simulations has different signs as compared to that of observations on average, both the
459 change of O₃ is indeed small and the model is able to capture the part of O₃ decrease in the
460 southwest part of the North China domain (Fig. 8(c)). We note that some previous studies showed
461 large increase of O₃ in North China, but such increase is in comparison with the O₃ in the month
462 right before the lockdown (not the same time in 2019; (Shi & Brasseur, 2020; Y. M. Liu et al.,
463 2021).
464



465
 466 **Figure 8.** Comparison of MDA8 surface O₃ in 2019 and 2020 Feb.-Mar. and the relative difference
 467 between two years from GEOS-Chem model (a-c) versus ground observations (d-f). GEOS-Chem
 468 mean MDA8 O₃ at 9 m above the surface under standard temperature and pressure (STP; 273.15
 469 K, 101.325 kPa) from (a) Baseline (2019) and (b) 2020 Adjoint simulation (Table 1) together with
 470 (c) their relative difference. Ground observed mean MDA8 surface O₃ under STP in (d) 2019 Feb.-
 471 Mar.; (e) 2020 Feb.-Mar. and (f) their relative difference. The pink and green boxes in (c) and (f)
 472 define the North China and South China domain.

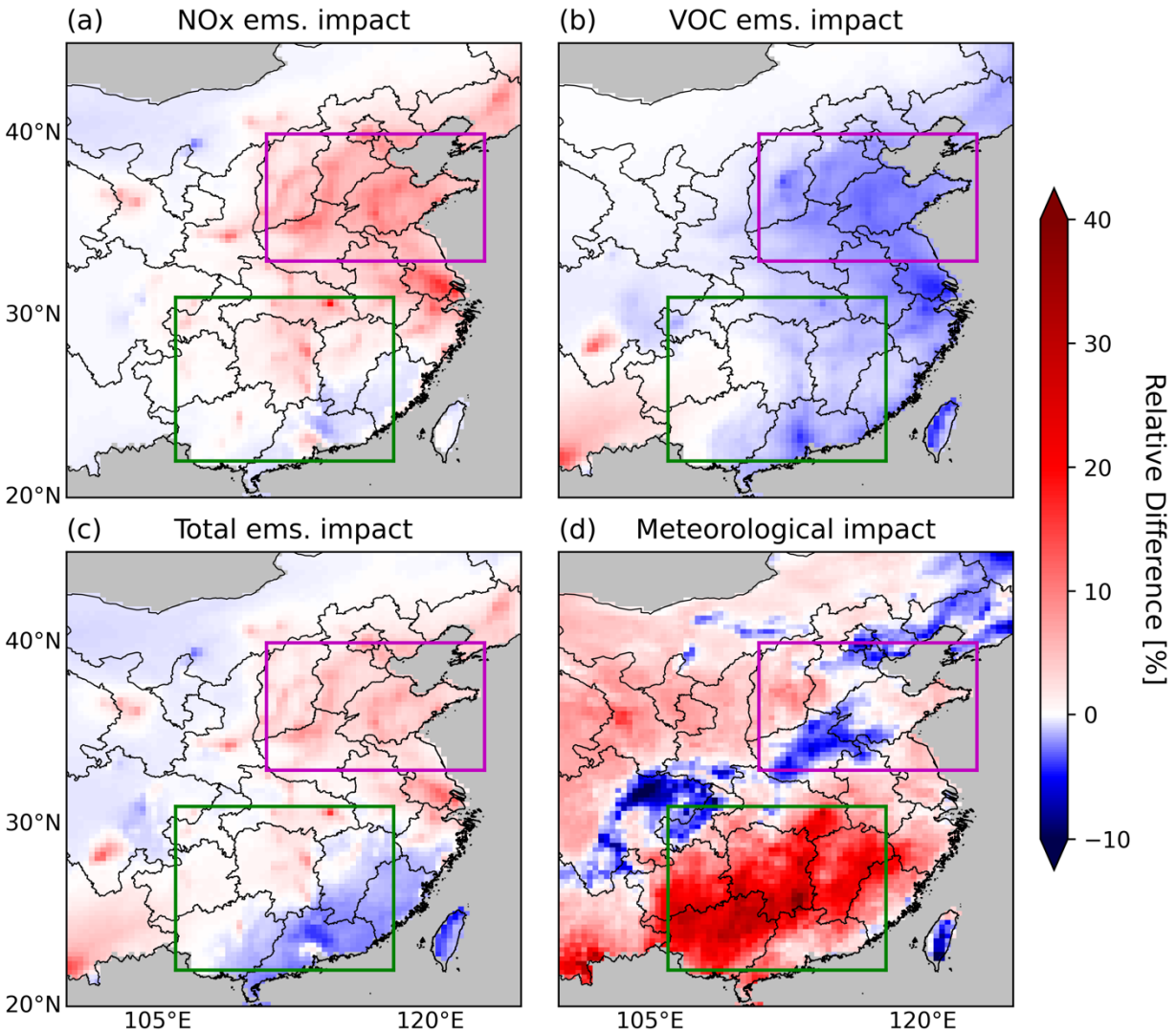
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 474

475 **4. Mechanism of aggravated surface O₃ pollution**

476 **4.1 Relative contribution from declining emissions and meteorological variations**

477 From equations (2)-(5) we can analyze the mechanism of surface O₃ increase in China
478 during the COVID-19 pandemic (Fig. 9). NO_x emission reduction as a result of COVID-19
479 lockdown leads to a ~8% increase in the mean MDA8 surface O₃ over North China (pink boxes in
480 Fig. 9) between 2019 and 2020 Feb.-Mar. (Fig. 9(a)), while the VOC emission decline causes ~3%
481 of O₃ decrease (Fig. 9(b)). The average contribution of the meteorological variations to the surface
482 O₃ change is less than 1% in North China (Fig. 9(d)). However, in South China, the inter-annual
483 meteorological variations dominate the surface O₃ increases, causing a ~30% increase (Fig. 9(d)),
484 while the NO_x and VOC emission reduction has little impacts. The overall magnitude of emissions
485 contribution to the surface O₃ change over North China is ~5%, similar to that of the
486 meteorological effects, but meteorological variations lead to both O₃ increases and decreases in
487 different regions. Over South China, the meteorological effect is much larger than the net effects
488 of declining emissions. Overall, the impact of inter-annual meteorological variations between 2019
489 and 2020 is almost 30 times larger than the overall emissions impacts on the aggregated surface
490 O₃ pollution in China. Our results are consistent with the conclusion from Zhao et al. (2020) that
491 meteorological variation has larger impacts than emissions reduction on surface O₃ in the southern
492 city of Guangzhou, but in Beijing, emission reduction has a larger impact during 23-29 January.
493 Liu et al. (2020b) reported that the surface O₃ increase in the major cities of the Yangtze River
494 Delta region were driven by both emission reduction and meteorological variations to a similar
495 degree from pre-lockdown period (Jan. 1-22, 2020) to lockdown period (Jan. 23-Feb. 29, 2020).
496 However, Zhao et al. (2020) and Liu et al. (2020b) only focused on the lockdown period of one
497 week in reference to the time period right before the lockdown instead of the same period in
498 previous years, which cannot exclude the effects of seasonal variation of meteorology and did not
499 provide a comprehensive analysis over the whole lockdown period. Moreover, Liu et al. (2020b)
500 only analyzed four representative cities instead of showing the analysis at a national scale. Further,
501 Zhao et al. (2020) did not update the anthropogenic emissions during the lockdown period, which
502 brings significant uncertainties to their analysis. Previous studies found that TROPOMI NO₂
503 product has a negative bias of -7% to -20% (Verhoelst et al., 2021; Judd et al., 2020; Li et al.,
504 2021). The sensitivity simulations indicate this low bias does not significantly affect the model
505 evaluation and our main conclusions (figures not shown).

506
507



508
 509 **Figure 9.** Relative difference in simulated surface O₃ caused by (a) NO_x emission reduction, (b)
 510 VOC emission reduction, (c) overall emission reduction and (d) meteorological variations due to
 511 COVID-19 lockdown. The pink and green boxes in each panel define the North China and South
 512 China domain.

513
 514

515 **4.2 Critical meteorological variablesHigher temperature causing leading to aggravated**
 516 **surface O₃ pollution in South China**

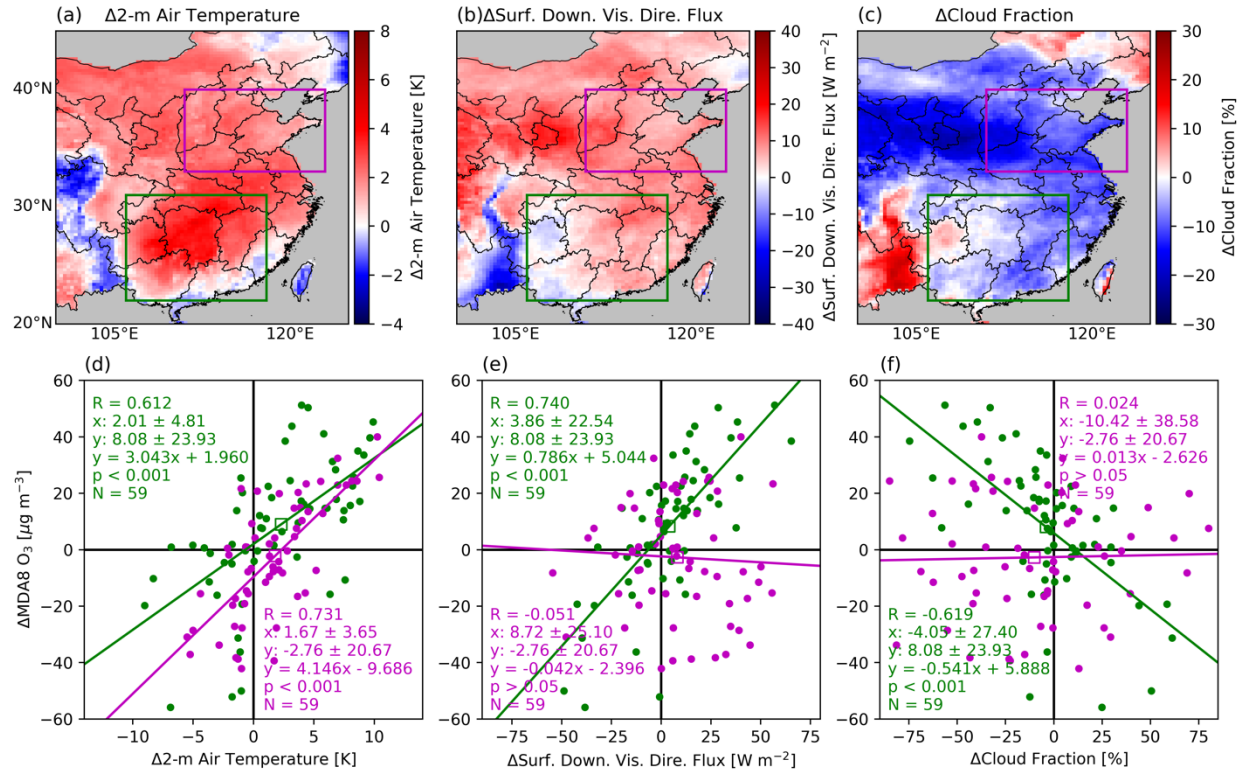
517 To identify the critical meteorological variables that lead to the aggravated surface O₃
 518 pollution in South China, we investigated the correlations between the surface O₃ concentrations
 519 and 2-m air temperature, downward visible direct flux at surface, cloud fraction, relative humidity,

520 and wind speed. The positive correlation between the surface O₃ and temperature is widely
521 observed and reported in the literature (Pusede et al., 2015). Higher temperature leads to higher
522 concentrations of surface O₃ because it improves the O₃ production rate by affecting the organic
523 reactivity, production of HOx radicals, formation and decomposition of peroxy nitrates and alkyl
524 nitrates (Pusede et al., 2015). We calculated the daily difference in Feb.-Mar. between 2020 and
525 2019 (excluding Feb. 29, 2020) for the daily mean of MDA8 O₃ from ground measurements and
526 2-meter air temperature, downward visible direct flux at surface and cloud fraction from GEOS-
527 FP data used in our GEOS-Chem simulations for the South China (106°E - 118°E, 22°N - 31°N,
528 green box in Fig. 10(a-c)) and North China (111°E - 123°E, 33°N - 40°N, pink box in Fig. 10(a-
529 c)). Fig. 10 displays the difference of 2-month mean 2-meter air temperature (Fig. 10(a)),
530 downward visible direct flux at surface (Fig. 10(b)) and cloud fraction (Fig. 10(c)) in Feb.-Mar.
531 between 2020 and 2019 (Fig. 10(a)) and the scatter plot between the daily difference of measured
532 surface O₃ concentration and 2-meter air temperature (Fig. 10(d)), downward visible direct flux
533 at surface (Fig. 10(e)) and cloud fraction (Fig. 10(f)) over both South China (green dots in Fig.
534 10(d-fb)) and North China (pink dots in Fig. 10(d-fb)). We found the 2-meter air temperature
535 increased by ~2.3°C in South China, and the daily difference of surface O₃ concentration and 2-
536 meter air temperature are well correlated with a positive correlation coefficient of 0.612. Therefore,
537 the temperature increase contributed to the significant aggravated surface O₃ pollution in South
538 China ~~could be attributed to the temperature increase.~~ The enhanced solar radiation at the surface
539 could also promote the production of O₃ via photochemical reactions. The correlation coefficient
540 between the daily difference of surface O₃ concentration and downward visible direct flux at
541 surface is as high as 0.740 in South China (Fig. 10(e)). The reason for the increase of temperature
542 ~~increase and solar radiation at surface~~ is the lower cloud fraction. Via analyzing the GEOS-FP data,
543 we found the cloud fraction decreases by ~5% (Fig. 10(c)), and the downward visible direct flux
544 at surface increased by 5 W m⁻² (Fig. 10(b)) over South China. The lower cloud fraction increases
545 the downward solar radiation at the surface during the lockdown period, leading to higher surface
546 air temperature. The change of cloud fraction is negatively correlated with the change of surface
547 O₃ in South China with the correlation coefficient of -0.619 (Fig. 10(f)).

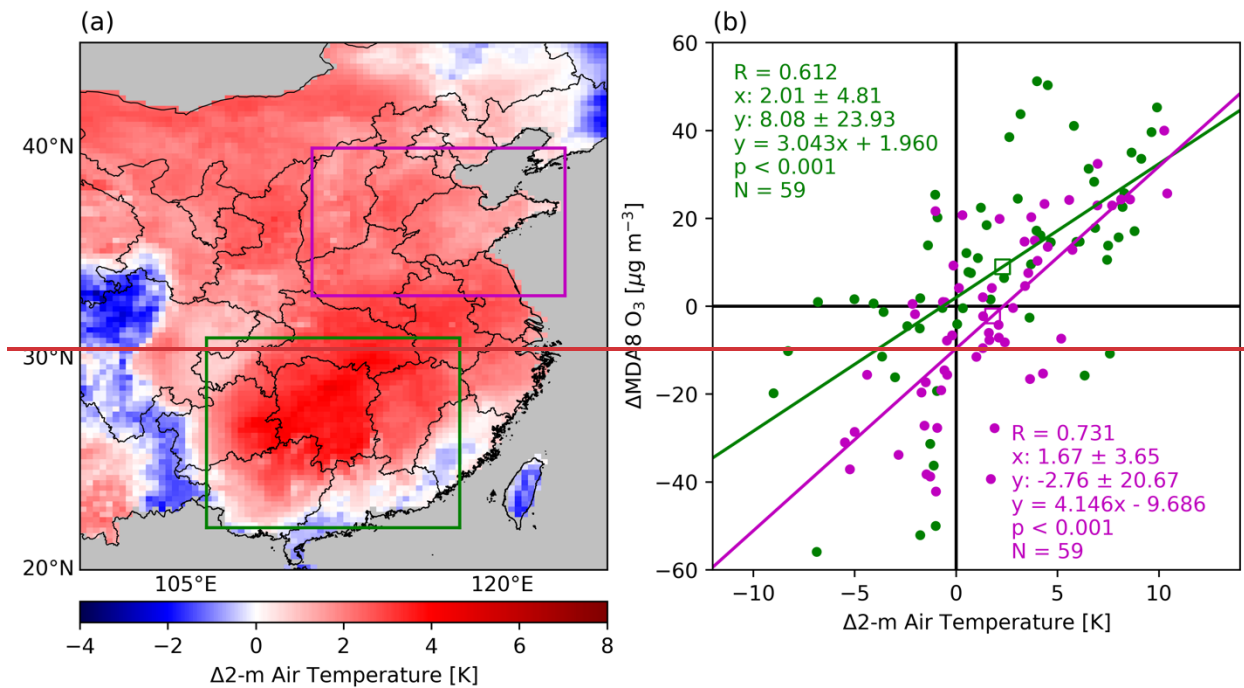
548 ~~The enhanced solar radiation at the surface could also promote the production of O₃ via~~
549 ~~photochemical reactions.~~ In North China, 2-meter air temperature also increased by 1.8°C, but the
550 measured surface MDA8 O₃ decreased by 3% (Fig. 8(f)). Fig. 10(d**b**) shows the daily difference

551 of MDA8 O₃ and 2-meter air temperature over North China also has a high correlation coefficient
552 of 0.731. However, the intercept of the linear regression line is negative, so that the surface O₃
553 could decrease even though the temperature increases. This negative intercept is caused by the net
554 effects of factors other than temperature, including chemistry, emissions, and other meteorological
555 factors. It is a challenge to quantify the contributions of each individual factors, because these
556 factors are thermodynamically or dynamically related. The predicted average change of surface
557 MDA8 O₃ in South China and North China are marked by the green and pink open squares
558 respectively in Fig. 10(d**b**) based on the linear regression. Because of the different intercepts, the
559 predicted MDA8 O₃ in South China increases by ~9.0 μg m⁻³, while it decreases by 2.2 μg m⁻³ in
560 North China, although the average temperature increased in both South and North China. On the
561 contrary to South China, the change of solar radiation at surface and cloud fraction are poorly
562 correlated with the change of surface O₃ concentrations in North China (Fig. 10(e, f)), mainly due
563 to a more important role of emission change in regulating the surface O₃ in North China.

564



565



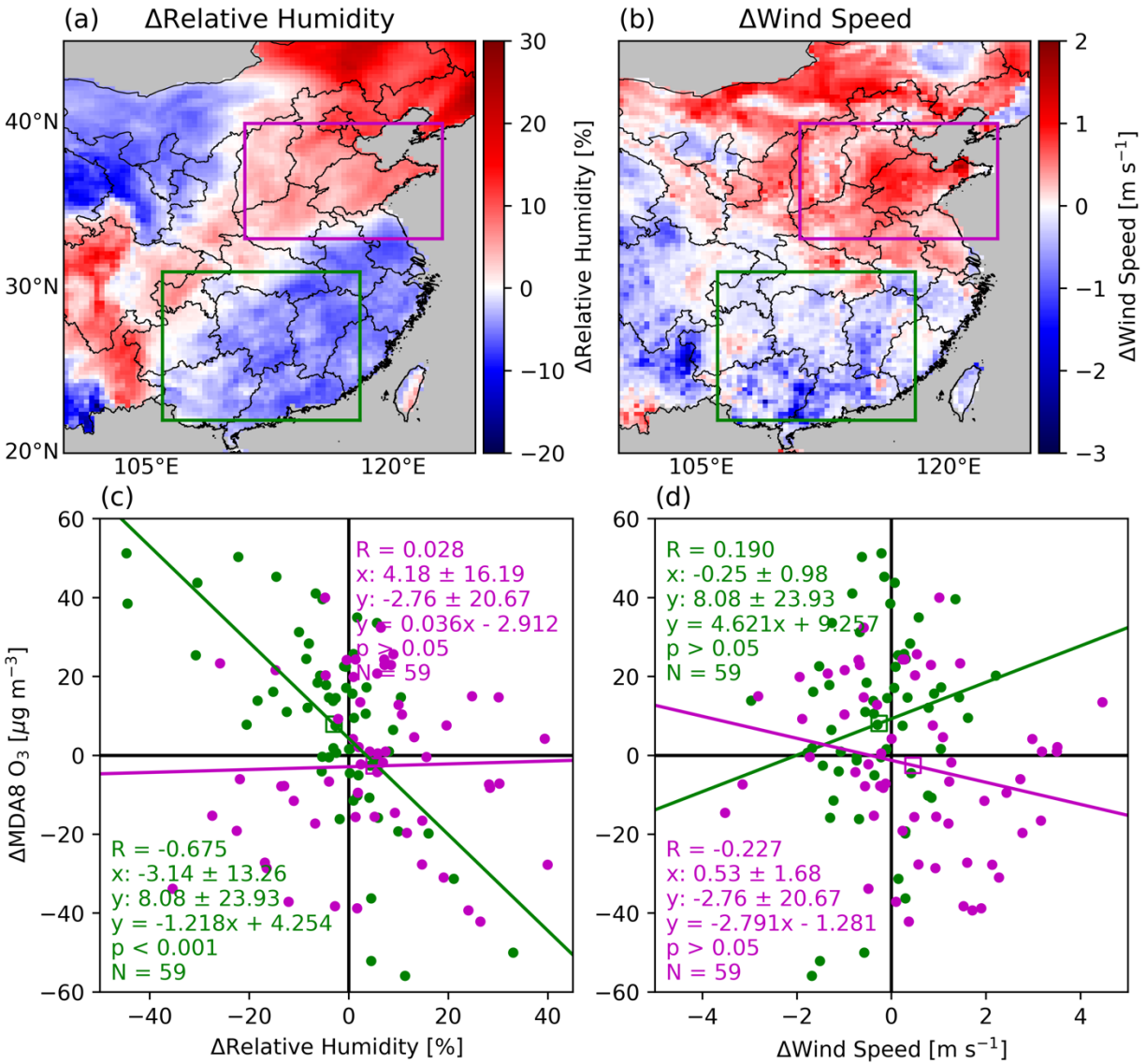
566

567 **Figure 10.** (a) The changes of (a) 2-meter air temperature, (b) downward visible direct flux at
 568 surface and (c) cloud fraction from 2019 Feb.-Mar. to 2020 Feb.-Mar. (b) and (d) The scatter plots
 569 between the daily difference of surface O₃ measurements and (d) 2-meter air temperature, (e)

570 downward visible direct flux at surface and (f) cloud fraction in Feb.-Mar. between 2020 and 2019
571 over South China (green dots) and North China (pink dots). The green and pink open squares in
572 (d-f) mark the predicted average change of surface MDA8 O₃ in South China ~~(green open square)~~
573 and North China ~~(pink open square)~~, respectively, based on the linear regression against
574 ~~temperature~~ the change of meteorological variables ~~change~~.

575

576 The impacts of relative humidity and wind speed on the surface O₃ change are also
577 investigated (Fig. 11). The relative humidity increased by ~5.1% in North China and decreased by
578 ~3.0% in South China on average. The strong correlation (R = -0.675) between the change of
579 relative humidity and surface O₃ in South China indicates the decrease of relative humidity also
580 contributes to the increase of surface O₃ pollution in South China, but the correlation between them
581 in North China is very low. The wind speed also changed oppositely in South China and North
582 China, but we cannot identify any significant impact of wind speed on the surface O₃ pollution
583 since the correlation coefficients are low in both South China and North China. In summary, the
584 significant increase of surface O₃ pollution during the lockdown period in South China could be
585 primarily attributed to the higher temperature, enhanced solar radiation at surface and declined
586 relative humidity.



587

588 **Figure 11.** Same as Fig. 10 but for (a, c) relative humidity and (b, d) wind speed.

589

590 5. Summary

591 A significant reduction in primary air pollutants has been identified by surface and satellite
 592 observations during the COVID-19 pandemic in China (Bauwens et al., 2020; Miyazaki et al.,
 593 2020), which is in contrast to the increase of surface O_3 . In this study, we analyzed the reasons for
 594 the enhanced surface O_3 pollution from two perspectives: anthropogenic emissions reduction and
 595 inter-annual meteorological variations. We constrain the NO_x emissions based on the TROPOMI
 596 NO_2 product using both the mass balance and 4D-Var methods. The VOC emissions_{sw} were also

597 updated based on the TROPOMI HCHO product via the mass balance approach. We analyzed the
598 contributions from emissions reduction and meteorological variations to surface O₃ increases
599 through a series of sensitivity simulations using the GEOS-Chem model.

600 The updated NO_x emissions from the 4D-Var and mass balance approaches share a similar
601 spatial pattern. However, the NO_x emissions from 4D-Var are lower than those from the mass
602 balance method over North China by ~10% but larger over central China by ~40%. The evaluation
603 of the simulations with the updated emissions against the TROPOMI NO₂, in-situ measurements
604 of surface NO₂ and O₃ indicate that the NO_x emissions from the 4D-Var inversion leads to better
605 model performance than that from the mass balance approach.

606 The anthropogenic NO_x emission decreased by ~30% over East China during 2020 Feb.-
607 Mar. compared to the same period in 2019. Over North China, NO_x emission reduction leads to a
608 ~8% increase in the mean MDA8 surface O₃, while the VOC emissions decline causes O₃ to
609 decrease by ~3%. The average contribution of meteorological variations to the surface O₃ change
610 is less than 1% in North China. However, in South China, the inter-annual meteorological variation
611 dominates the surface O₃ increase, causing a ~30% increase, while the NO_x and VOC emission
612 reduction has nearly no impacts on O₃. Overall, the impact of inter-annual meteorological
613 variations between 2019 and 2020 is almost 30 times larger than the impact of emissions -on the
614 enhanced surface O₃ pollution in China.

615 The significant increase of surface O₃ in South China could be attributed to the higher
616 temperature, enhanced solar radiation at surface and declined relative humidity during the
617 lockdown period, which is caused by the lower cloud fraction. The lower cloud fraction increases
618 the downward solar radiation at the surface during the lockdown period, leading to higher surface
619 air temperature. We cannot identify any significant impact of wind speed on the surface O₃
620 pollution. The enhanced solar radiation at the surface could also promote the production of O₃ via
621 photochemical reactions.

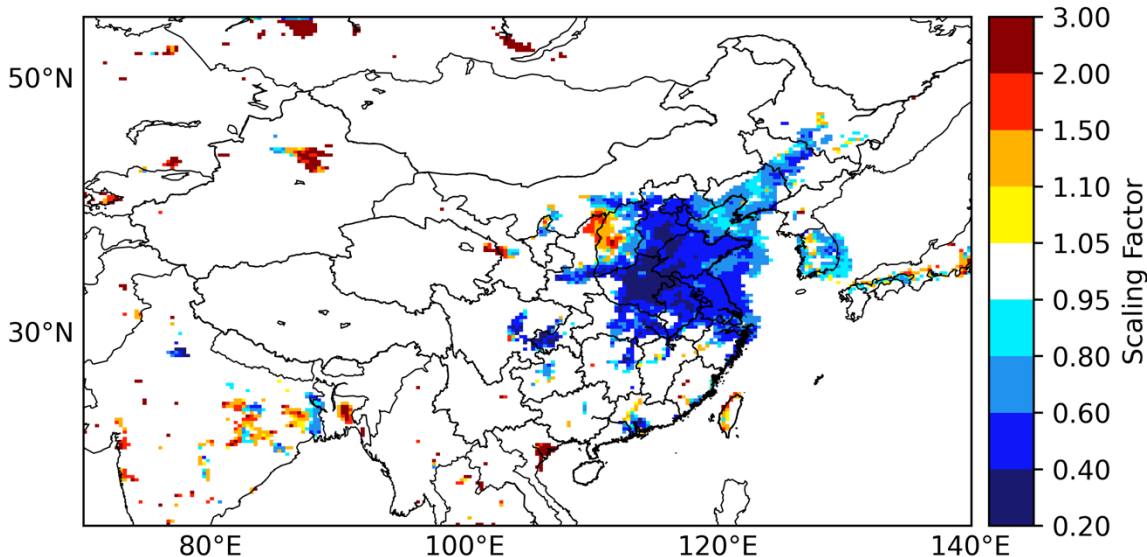
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623

624 **Appendix A: NO_x emission reduction in China from 2010 to 2019**

625 The default anthropogenic NO_x emission over East Asia in GEOS-Chem is MIX 2010 (Li
626 et al., 2017). To generate the anthropogenic NO_x emission in 2019, we calculated the ratio of mean
627 TROPOMI tropospheric NO₂ VCD in Feb.-Mar. 2019 to GEOS-Chem simulated NO₂ VCD with

628 the default MIX 2010 emission as the scaling factor (Fig. A1). The scaling factors in regions where
 629 mean TROPOMI tropospheric NO₂ VCD in 2019 Feb.-Mar. less than 0.1 DU are set to 1. From
 630 2010 to 2019, the anthropogenic NO_x emission has declined significantly as a result of the clean
 631 air actions of Chinese government (Zheng et al., 2018).
 632



633
 634 **Figure A1.** The scaling factor of anthropogenic NO_x emission from year 2010 to 2019.
 635

636
 637 **Appendix B: Applying the TROPOMI NO₂ averaging kernel in the observation**
 638 **operator**

639 To optimize the NO_x emissions and minimize the cost function (Equation (1)) with the 4D-
 640 Var method, GEOS-Chem adjoint needs to compute the derivative of the cost function with respect
 641 to the model parameters to be optimized, which is the scaling factors of the anthropogenic NO_x
 642 emissions in this study. An essential step is to calculate the adjoint forcing \mathbf{F} , which is the
 643 derivative of the cost function with respect to the modeled NO₂ concentration shown as Equation
 644 (B1).

645
$$\mathbf{F} = \frac{\partial J}{\partial \mathbf{c}} = \mathbf{S}_{\text{obs}}^{-1} [H(\mathbf{c}) - \mathbf{s}] \frac{\partial H(\mathbf{c})}{\partial \mathbf{c}} \quad (\text{B1})$$

646 For each single TROPOMI NO₂ observation, the adjoint forcing component f and cost function
 647 component j are computed as Equation (B2) and Equation (B3).

648
$$f = \frac{M_{gc}v_{gc} - M_{obs}v_{obs}}{e_{obs}M_{obs}}M_{gc} \quad (B2)$$

649
$$j = \frac{0.5f(M_{gc}v_{gc} - M_{obs}v_{obs})}{M_{gc}} \quad (B3)$$

650 Here M_{gc} is GEOS-Chem air mass factor applying the GEOS-Chem NO₂ vertical profiles and
 651 TROPOMI NO₂ averaging kernel. M_{obs} is TROPOMI air mass factor. v_{gc} and v_{obs} are the
 652 tropospheric NO₂ VCD from GEOS-Chem model and TROPOMI observation, respectively. The
 653 product of air mass factor and NO₂ VCD is NO₂ slant column density. e_{obs} is the standard error of
 654 TROPOMI tropospheric NO₂ VCD.

655 We calculated the GEOS-Chem air mass factor M_{gc} as Equation (B4) following Qu et al.
 656 (2019).

657
$$M_{gc} = \frac{\sum_{i \in \text{trop.}} c_i^{gc} \Delta p_i^{gc} w_i^{gc}}{\sum_{i \in \text{trop.}} c_i^{gc} \Delta p_i^{gc}} \quad (B4)$$

658 Here c_i^{gc} is GEOS-Chem NO₂ mixing ratio at vertical layer i , Δp_i^{gc} is the pressure difference
 659 between the GEOS-Chem vertical layer i and $i+1$. w_i^{gc} is the scattering weight at the GEOS-Chem
 660 vertical layer i , which is calculated by the linear interpolation of the scattering weights at the
 661 vertical coordinate of the model TM5 used for TROPOMI NO₂ retrieval. The scattering weight at
 662 the TM5 vertical layer l (w_l^{TM5}) is computed as the product of TROPOMI air mass factor and the
 663 TROPOMI averaging kernel at the TM5 vertical layer l (A_l^{TM5}) using Equation (B5-B6) (Eskes
 664 and Boersma, 2003; Palmer et al., 2001).

665
$$w_l^{TM5} = \frac{M_{obs}A_l^{TM5}}{M_{geo}} \frac{M_{obs}A_l^{TM5}}{M_{obs}A_l^{TM5}} \quad (B5)$$

666
$$M_{geo} = \sec \theta_0 + \sec \theta \quad (B6)$$

667 where M_{geo} is geometric air mass factor, θ_0 and θ are solar zenith angle and viewing zenith angle,
 668 respectively.

669

670 **Appendix C: Validation of the TROPOMI NO₂ observation operator**

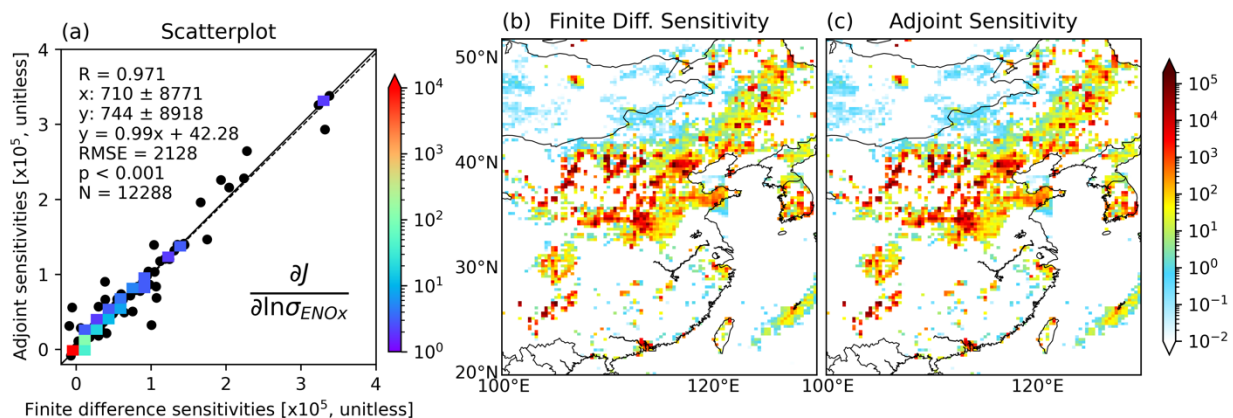
671 We validated the observation operator by comparing the sensitivity of the cost function
 672 with respect to the emission scaling factor from GEOS-Chem adjoint and a finite difference
 673 estimation as shown in Equation (C1). We shut down the transport and exclude a priori term from

674 the cost function for the validation, so that the gradient of cost function component in each grid
 675 cell to the local emission scaling factor equals to the gradient of total cost function to the emission
 676 scaling factor in the same grid cell.

$$677 \quad \frac{\partial J(\ln \sigma)}{\partial \ln \sigma} \approx \frac{J(\ln(\sigma + 0.05)) - J(\ln(\sigma - 0.05))}{\ln(\sigma + 0.05) - \ln(\sigma - 0.05)} \quad (C1)$$

678 Fig. C1 compared the cost function sensitivities calculated from GEOS-Chem adjoint and
 679 the finite difference method for the nested grids with the spatial resolution of $0.25^\circ \times 0.3125^\circ$. The
 680 spatial pattern and magnitude of the cost function sensitivities from the two methods match with
 681 each other with a correlation coefficient of 0.97. The statistics show that the agreement of the
 682 adjoint sensitivities and finite difference sensitivities in this study is comparable to that in Wang
 683 et al. (2020a) although we constrain the NO_x emission at a much finer resolution of $0.25^\circ \times 0.3125^\circ$
 684 than in their study ($2^\circ \times 2.5^\circ$).

685



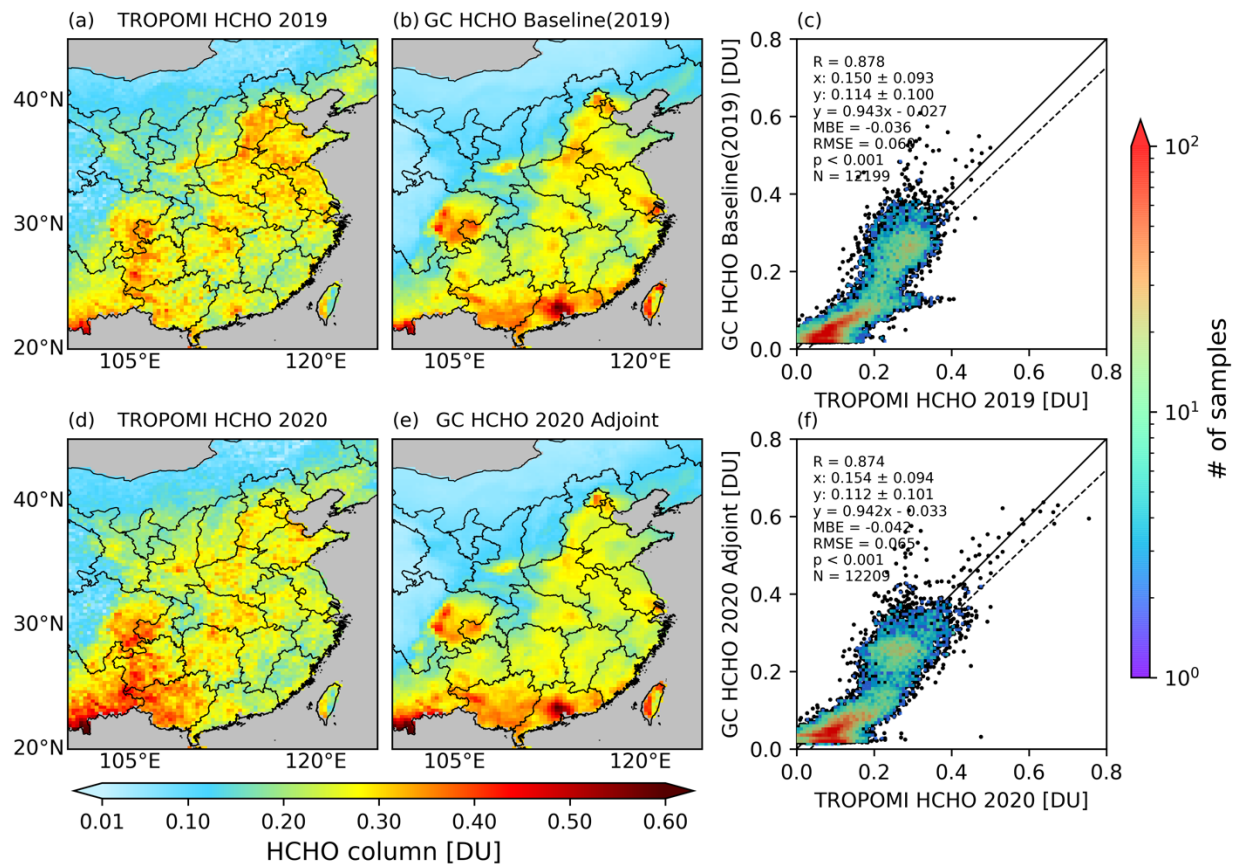
686
 687 **Figure C1.** Comparison of adjoint sensitivities and finite difference sensitivities. (a) Scatter plot
 688 of the adjoint sensitivity of the cost function with respect to the logarithm of NO_x emission scaling
 689 factor versus the finite difference sensitivities. The color scheme for panel (a) encodes the number
 690 of samples (the legend on the right of panel (a)). (b) Map of finite difference sensitivity. (c) Map
 691 of adjoint sensitivity.

692

693 Appendix D: Validation of HCHO simulations

694 It is not practical to validate the VOC emissions species by species in this study. We
 695 compare the GEOS-Chem simulation of HCHO with TROPOMI data for a rough validation of the
 696 VOC emissions. Overall, the GEOS-Chem simulations Baseline (2019) and 2020 Adjoint agree

697 well with TROPOMI HCHO in Feb.-Mar. of 2019 and 2020, respectively (Fig. D1). The
 698 correlation coefficients for the two comparisons are 0.878 and 0.874, and the MBE are -0.036 DU
 699 and -0.042 DU, respectively. This good agreement also supports our assumption of ignoring the
 700 change of VOC emissions from 2010 to 2019, since the anthropogenic VOC emission used for
 701 Baseline (2019) is equivalent to MIX 2010. However, Fig. D1 shows GEOS-Chem tends to
 702 overestimate the HCHO VCD over urban regions and underestimate it over rural regions, which
 703 indicates the partition of total emissions may overestimate the anthropogenic emissions and
 704 underestimate the biogenic emissions, at least for HCHO. A more comprehensive evaluation for
 705 various species and emission sectors would be helpful for further improvement of the VOC
 706 emissions used in the model.



708
 709 **Figure D1.** Comparison of tropospheric HCHO VCD from TROPOMI and GEOS-Chem. (a):
 710 TROPOMI HCHO in 2019 Feb.-Mar. (b): GEOS-Chem HCHO from Baseline (2019). (c): Scatter
 711 plot between (a) and (b). (d): TROPOMI HCHO in 2020 Feb. -Mar. (e): GEOS-Chem HCHO from

712 [2020 Adjoint. \(f\): Scatter plot between \(d\) and \(e\). The emissions and meteorology configurations](#)
713 [for GEOS-Chem simulations Baseline \(2019\) and 2020 Adjoint are listed in Table 1.](#)

717 **Data availability**

718 The TROPOMI NO₂ and HCHO product are available at the NASA Goddard Earth
719 Sciences Data and Information Services Center (<https://daac.gsfc.nasa.gov>). The ground O₃ and
720 NO₂ measurements are available at the China National Environmental Monitoring Center
721 (<http://www.cnemc.cn/en/>).

723 **Author contribution**

724 ZL and JW designed the research, and ZL conducted the research. YW, DKH and XC
725 contributed to the research design. ZL and JW wrote the manuscript, and XC and DKH contributed
726 to the writing. YW and TS developed the codes for comparing the tropospheric NO₂ VCD from
727 model and TROPOMI data. KS developed the codes for oversampling.

729 **Competing interests**

730 The authors declare that they have no conflict of interest.

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