

1 **A large role of missing volatile organic compounds reactivity**
2 **from anthropogenic emissions in ozone pollution regulation**

3 Wenjie Wang^{1,2*}, Bin Yuan^{1*}, Hang Su², Yafang Cheng², Jipeng Qi¹, Sihang Wang¹,
4 Wei Song³, Xinming Wang³, Chaoyang Xue², Chaoqun Ma², Fengxia Bao², Hongli
5 Wang⁴, Shengrong Lou⁴, Min Shao¹

6 ¹ Institute for Environmental and Climate Research, Jinan University,
7 Guangzhou 511443, China

8 ² Multiphase Chemistry Department, Max Planck Institute for Chemistry, Mainz
9 55128, Germany

10 ³ State Key Laboratory of Organic Geochemistry, Guangzhou Institute of
11 Geochemistry, Chinese Academy of Sciences, Guangzhou 510640, China

12 ⁴ State Environmental Protection Key Laboratory of Formation and Prevention of
13 Urban Air Pollution Complex, Shanghai Academy of Environmental Sciences,
14 Shanghai 200233, China

15

16

17 *Correspondence to: Bin Yuan (byuan@jnu.edu.cn);
18 Wenjie Wang (Wenjie.Wang@mpic.de)

19

20 **Abstract:** There are thousands of VOC species in ambient air, while existing techniques
21 can only detect a small part of them (~ several hundred). The large number of
22 unmeasured VOCs prevents us from understanding the photochemistry of ozone and
23 aerosols in the atmosphere. The major sources and photochemical effects of these
24 unmeasured VOCs in urban areas remain unclear. The missing VOC reactivity, which
25 is defined as the total OH reactivity of the unmeasured VOCs, is a good indicator to
26 constrain the photochemical effect of unmeasured VOCs. Here, we identified the
27 dominant role of anthropogenic emission sources in the missing VOC reactivity
28 (accounting for up to 70%) by measuring missing VOC reactivity and tracer-based
29 source analysis in a typical megacity in China. Omitting the missing VOC reactivity
30 from anthropogenic emissions in model simulations will remarkably affect the
31 diagnosis of sensitivity regimes for ozone formation, overestimating the degree of
32 VOC-limited regime by up to 46%. Therefore, a thorough quantification of missing
33 VOC reactivity from various anthropogenic emission sources is urgently needed for
34 constraints of air quality models and the development of effective ozone control
35 strategies.

36
37

38 **1 Introduction**

39 Volatile organic compounds (VOCs) are key precursors of major photochemical
40 pollutants, including ozone (O₃) and secondary organic aerosols(Atkinson,
41 2000;Atkinson and Arey, 2003). Severe O₃ and particle pollution are frequently related
42 to high emissions of VOCs (Atkinson and Arey, 2003;Monks et al., 2015). There exist
43 thousands of VOC species in ambient air that are emitted from either natural processes
44 or anthropogenic activities (Goldstein and Galbally, 2007). No one instrument can
45 capture all VOCs out there and even when they can be measured there is information
46 missing on identification and properties (Yuan et al., 2017;Wang et al., 2014). Gas
47 chromatograph–mass spectrometer/flame ionization detector (GC–MS/FID) can
48 measure C2-C12 non-methane hydrocarbons (NMHCs) and C2-C6 oxygenated VOCs
49 (OVOCs) while cannot measure NMHCs and OVOCs with larger carbon number
50 (Wang et al., 2014). Proton-transfer-reaction time-of-flight mass spectrometer (PTR-
51 ToF-MS) is able to measure a huge number of OVOCs and aromatics and several
52 alkanes, but cannot measure most alkanes and alkenes, and cannot distinguish isomers
53 (Yuan et al., 2017). The 2,4-dinitrophenylhydrazine (DNPH)/high performance liquid
54 chromatography (HPLC) method can measure several carbonyls but cannot measure
55 non-polar organic species (Wang et al., 2009). The two-dimensional GC is able to
56 measure some intermediate-volatile and semi-volatile non-polar organics (Song et al.,
57 2022). A lack of standard gases prevents these technologies from accurate
58 quantification even if these technologies can identify more VOC species. In general,
59 many branched alkenes, OVOCs with complex functional groups, intermediate-volatile
60 and semi-volatile organics and complex biogenic VOCs cannot currently be well
61 quantified even if they can be identified by instruments. As a result, the total amount of
62 VOCs in ambient air has generally been underestimated. Currently, emission
63 inventories used in air quality models such as the Community Emissions Data System
64 (CEDS) emission inventory and the multi-resolution Emission Inventory for China
65 (MEIC) only include the VOC species that can be measured such as some C1-C9

删除了: By now

删除了: of VOCs

68 hydrocarbons and simple-structure OVOCs with small carbon number (<C6). This will
69 lead to an underestimation of the photochemical effect of total VOCs and thus causes
70 uncertainties in predicting secondary pollution. The quantification of the unmeasured
71 VOCs is crucial to assess secondary pollution precisely.

删除了: which

72 The total OH reactivity (R_{OH}), which can be directly measured, is an index for
73 evaluating the amount of reductive pollutants in terms of ambient OH loss. The total
74 OH reactivity is defined as:

删除了: The measurement of total OH reactivity (R_{OH}) provides an effective approach to quantify the total amount of reactive gases in terms of reacting with OH radicals.

$$75 \quad R_{OH} = \sum_i k_{OH+X_i} [X_i], \quad (1)$$

76 where X represents a reactive species including carbon monoxide (CO), nitrogen oxides
77 (NO_x) and VOCs etc., and k_{OH+X_i} is the reaction rate constant for the oxidation of
78 species X by OH. The measured R_{OH} is higher than that calculated based solely on the
79 measured reactive species, and the difference between them is mostly from unmeasured
80 VOCs (Yang et al., 2017). Missing VOC reactivity (missing VOC_R), defined as VOC
81 reactivity (VOC_R) of all unmeasured VOCs, can be obtained by subtracting the
82 calculated R_{OH} from the measured R_{OH} .

$$83 \quad \text{missing } VOC_R = \text{measured } R_{OH} - \text{calculated } R_{OH} \quad (2)$$

$$84 \quad \text{calculated } R_{OH} = \sum_i k_{OH+reactive\ species_i} [reactive\ species_i] \quad (3)$$

85 where reactive species represents measured VOCs and reactive inorganic species
86 including carbon monoxide (CO), nitric oxide (NO), nitrogen dioxide (NO_2), O_3 , sulfur
87 dioxide (SO_2), nitrous acid (HONO), and so on. The missing VOC_R provides a
88 constraint for evaluating the photochemical roles of unmeasured VOCs in the
89 atmosphere (Sadanaga et al., 2005; Yang et al., 2016b). The inclusion of the missing
90 VOC_R can help to improve the performance of box model and air quality models in
91 simulating photochemistry processes. Relatively high missing VOC_R have been found
92 in forests (Di Carlo et al., 2004; Hansen et al., 2014; Nakashima et al., 2014; Nölscher et
93 al., 2016; Praplan et al., 2019), urban areas (Shirley et al., 2006; Yoshino et al.,
94 2006; Dolgorouky et al., 2012; Yang et al., 2017) and suburban areas (Kovacs et al.,
95 2003; Yang et al., 2017; Fuchs et al., 2017; Lou et al., 2010), accounting for 10-75% of
96 total R_{OH} . Given that total VOC_R is one part of total R_{OH} , missing VOC_R would account

删除了: performance

102 for a larger percentage of total VOC_R (>10%-75%).

103 The potential sources of missing VOC_R include anthropogenic emissions, biogenic
104 emissions, soil emissions, and photochemical production, etc (Yang et al., 2016b).
105 Previous studies have reported that the missing VOC_R in forest areas was mainly from
106 either direct emissions or photochemical oxidation of biogenic VOCs (Di Carlo et al.,
107 2004;Hansen et al., 2014;Nakashima et al., 2014;Nölscher et al., 2016;Praplan et al.,
108 2019). Nevertheless, the dominant source of the missing VOC_R in urban and suburban
109 areas remains unclear or under debate.

110 Surface O₃ pollution has become a major public health concern in cities worldwide
111 (Paoletti et al., 2014;Lefohn et al., 2018). A critical issue in determining an emission
112 control strategy for ozone pollution is to understand the relative benefits of NO_x and
113 VOC emission controls. This is generally framed in terms of ozone precursor sensitivity,
114 i.e., whether ozone production is NO_x-limited or VOC-limited (Kleinman,
115 1994;Sillman et al., 1990). Nevertheless, the effect of missing VOCs on ozone
116 precursor sensitivity has not been well understood yet. Given that the missing VOC_R
117 could potentially account for a large part of total VOC_R, clearly clarifying the role of
118 missing VOC_R in determining ozone precursor sensitivity is an urgent need for the
119 diagnosis of ozone sensitivity regimes and formulation of an effective emission
120 reduction roadmap.

121 China has become a global hot spot of ground-level ozone pollution in recent years
122 (Lu et al., 2018;Wang et al., 2022). Pearl River Delta (PRD) remains one of the most
123 O₃-polluted regions in China (Li et al., 2022), although many control measures have
124 been attempted. Here, we measured RO_H in Guangzhou, a megacity in PRD and
125 quantified the missing VOC_R. The dominant source of the missing VOC_R and its impact
126 on ozone precursor sensitivity were comprehensively investigated.

删除了: s

2 Method

2.1 Overview of the measurement

The field campaign was conducted from 25 September to 30 October 2018 continuously at an urban site in downtown Guangzhou (113.2°E, 23°N). The sampling site is located on the ninth floor of a building on the campus of Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, 25 m above the ground level. This site is primarily influenced by industrial and vehicular emissions. ROH, VOCs, NOX, O3, HONO, SO2, CO, photolysis frequencies, and meteorological factors were simultaneously measured during the measurement period.

删除了:

2.2 ROH measurement

删除了: 1

Total ROH was measured by the comparative reactivity method (CRM) (Sinha et al., 2008). The CRM system consists of three major components, namely an inlet and calibration system, a reactor, and a measuring system. Here, pyrrole (C₄H₅N) was used as the reference substance in CRM and its concentration was quantified by a quadrupole proton-transfer-reaction mass spectrometer (PTR-MS) (Ionicon Analytik GmbH, Innsbruck, Austria). The CRM system was calibrated by propane, propene, toluene standards and 16 VOC mixed standard (acetaldehyde, methanol, ethanol, isoprene, acetone, acetonitrile, methyl vinyl ketone, methyl ethyl ketone, benzene, toluene, o-xylene, α -pinene, 1,2,4-trimethylbenzene, phenol, m-cresol and naphthalene). Measured and calculated ROH agreed well within 15% for all calibrations. The ROH measurement by the CRM method is interfered from ambient nitric oxide (NO), which produces additional OH radicals via the reaction of HO₂ radicals with NO (Sinha et al., 2008). To correct this interference, a series of experiments were conducted by introducing different levels of NO (0–160 ppb) and given amounts of VOC into the CRM reactor. A correction curve was acquired from these NO interference experiments, which can be used to correct the ROH thanks to the simultaneous measurement of ambient NO concentrations (Supplementary information S1; Fig. S1). The detection limits of the CRM method were around 2.5 s⁻¹, and the total uncertainty was estimated

删除了: The field campaign was conducted from 25 September to 30 October 2018 at an urban site in downtown Guangzhou (113.2°E, 23°N). This site is primarily influenced by industrial and vehicular emissions.

删除了: 100

163 to be about 15%. The CRM method has been successfully applied to measure OH
164 reactivity in urban areas with high NO_x levels in previous studies (Dolgorouky et al.,
165 2012; Yang et al., 2017; Hansen et al., 2015). The intercomparison between the CRM
166 method and pump-probe technique indicates that the CRM method can be used under
167 high-NO_x conditions (NO_x>10 ppb) if a NO_x-dependent correction is applied (Hansen
168 et al., 2015).

删除了: carefully

169 **2.3 VOCs measurements**

删除了: 2

170 Nonmethane hydrocarbons (NMHCs) were measured using a gas chromatograph-
171 mass spectrometer/flame ionization detector (GC-MS/FID) system coupled with a
172 cryogen-free preconcentration device (Wang et al., 2014). The system contains two-
173 channel sampling and GC column separation, which is able to measure C₂-C₅
174 hydrocarbons with the FID in one channel and measure C₅-C₁₂ hydrocarbons using
175 MS detector in the other channel. After removal of water vapor, VOCs were trapped at
176 -155 °C in a deactivated quartz capillary column (15 cm×0.53 mm ID) and a Porous
177 Layer Open Tubular (PLOT) capillary column (15 cm×0.53 mm ID) for the MS channel
178 and the FID channel, respectively. The system was calibrated weekly by TO-15 (Air
179 Environmental Inc., USA) and PAMS gas standards (Spectra Gases Inc., USA).
180 Detection limits for various compounds were in the range of 0.002-0.070 ppbv. A total
181 of 56 NMHCs species were measured (Table S1). The time resolution of the
182 measurement was 1 h. The uncertainties of VOC measurements by GC-MS/FID are in
183 the range of 15 %-20 %. More details of this method can be found in previous studies
184 (Wang et al., 2014; Yuan et al., 2012).

185 An online proton-transfer-reaction time-of-flight mass spectrometer (PTR-ToF-
186 MS) (Ionicon Analytic GmbH, Innsbruck, Austria) with H₃O⁺ and NO⁺ ion sources was
187 also used to measure VOCs. During the campaign, the PTR-ToF-MS automatically
188 switched between H₃O⁺ and NO⁺ chemistry every 10-20 min. The H₃O⁺ mode was
189 used to measure OVOCs and aromatics while the NO⁺ model was used to measure
190 alkanes with more carbons (C₈-C₂₀). When running in the H₃O⁺ ionization mode, the

193 drift tube was at a temperature of 50 °C, a pressure of 3.8 mbar, and a voltage of 920
194 V, leading to an operating E/N (E is the electric field, and N is the number density of
195 the gas in the drift tube) ratio of 120 Td. When running in the NO⁺ ionization mode, the
196 drift tube was at a temperature of 50 °C, a pressure of 3.8 mbar, and a voltage of 470
197 V, leading to an operating E/N ratio of 60 Td. PTR-ToF-MS technique is capable of
198 measuring oxygenated VOCs (OVOCs) and higher alkanes that GC-MS/FID cannot
199 measure (Wu et al., 2020; Wang et al., 2020a). The time resolution of PTR-ToF-MS
200 measurements was 10 s. A total of 31 VOCs were calibrated using either gas or liquid
201 standards (Table S2). For other measured VOCs, we used the method proposed by
202 Sekimoto et al. (2017) to determine the relationship between VOC sensitivity and
203 kinetic rate constants for proton transfer reactions of H₃O⁺ with VOCs. The fitted line
204 was used to determine the concentrations of those uncalibrated species. The
205 uncertainties of the concentrations for uncalibrated species were about 50 % (Sekimoto
206 et al., 2017). By this method, PTR-ToF-MS can additionally measure 128 VOCs which
207 were included in the analysis of this study. The detailed information for this method can
208 be found in Wu et al. (2020) and all VOC species measured by PTR-ToF-MS were
209 provided in table S4 of that article. The PTR-ToF-MS is capable of measuring
210 additional VOC species that GC-MS/FID cannot measure including alkanes with more
211 carbons (C12-C20) and OVOCs including aldehydes, ketones, carboxylic acids,
212 alcohols, and nitrophenols. Formaldehyde (HCHO) was measured by a custom-built
213 instrument based on the Hantzsch reaction and absorption photometry (Xu et al., 2022).

删除了:

删除了: NMHCs

214 **2.4 Other measurements**

删除了: 3

215 Nitrous acid (HONO) was measured by a custom-built LOPAP (Long Path
216 Absorption Photometer) based on wet chemical sampling and photometric detection
217 (Yu et al., 2022). The uncertainty of the measurement was 8 %. NO_x, O₃, SO₂, and CO
218 were measured by NO_x analyzer (Thermo Scientific, Model 42i), O₃ analyzer (Thermo
219 Scientific, Model 49i), SO₂ analyzer (Thermo Scientific, Model 43i), and CO analyzer
220 (Thermo Scientific, Model 48i), respectively. The meteorological data, including
221 temperature (T), relative humidity (RH) and wind speed and direction (WS, WD) were

删除了: sulfur dioxide (

删除了:)

227 recorded by Vantage Pro2 Weather Station (Davis Instruments Inc., Vantage Pro2) with
228 a time resolution of 1 min. Photolysis frequencies of O₃, NO₂, HONO, H₂O₂, HCHO₂
229 and NO₃ were measured by a spectrometer (Focused Photonics Inc., PFS-100) (Shetter
230 and Müller, 1999; Wang et al., 2019).

231 **2.5 Multiple linear regression**

232 The Multiple Linear Regression (MLR) has been successfully applied to quantify
233 the sources of air pollutants (Li et al., 2019; Yang et al., 2016a). In this study, a tracer-
234 based MLR analysis was used to decouple the individual contributions of
235 anthropogenic emissions, secondary production, biogenic emissions and background
236 level to missing VOC_R, as shown in Eq. (4).

$$237 \text{Missing VOC}_R = a\Delta\text{CO} + b[\text{O}_X] + c[\text{isoprene}_{\text{initial}}] + C_{\text{background}} \quad (4)$$

238 where O_X is defined as O₃+NO₂. ΔCO, [O_X] and [isoprene_{initial}] are concentrations
239 of tracers for anthropogenic emissions, secondary production and biogenic emissions,
240 respectively. ΔCO is the relative change between ambient CO and background CO of
241 150 ppb (Wang et al., 2020a). [isoprene_{initial}] represents the initial concentration of
242 isoprene from biogenic emissions that has not undergone any photochemical reactions,
243 which is calculated from observed isoprene and its photochemical products methyl
244 vinyl ketone (MVK) and methacrolein (MACR) (Xie et al., 2008). C_{background}
245 indicates the background level of missing VOC_R. a, b, c and C_{background} are fitted
246 coefficients by the multiple linear regression.

247 **2.6 Observation-based box model**

248 A zero-dimensional box model coupled with the Master Chemical Mechanism
249 (MCM) v3.3.1 chemical mechanism (Jenkin et al., 2003) was used to simulate the
250 photochemical production of RO_X (RO_X=OH+HO₂+RO₂) radicals and O₃ during the
251 field campaign. The model was constrained by the observations of meteorological
252 parameters, photolysis frequencies, VOCs, NO, NO₂, O₃, CO, SO₂ and HONO. The
253 model runs were performed in a time-dependent mode with a time resolution of 1 hour

删除了: 4

删除了: multiple linear regression

删除了: have

删除了: 5

258 and a spin-up of four days. A 24-h lifetime was introduced for all simulated species,
259 including secondary species and radicals, to approximately simulate dry deposition and
260 other losses of these species (Lu et al., 2013; Wang et al., 2020b). Sensitivity tests show
261 that this assumed physical loss lifetime has a relatively small influence on RO_x radicals
262 and ozone production rates.

263 Measured OVOCs such as HCHO, acetaldehyde and acetone were constrained in
264 the model and unmeasured OVOCs were simulated according to the photochemical
265 oxidation of NMHCs by OH radicals. RO₂, HO₂ and OH radicals were simulated by the
266 box model to calculate the net O₃ production rate (P(O₃)) and O₃ loss rate (L(O₃)) as
267 shown in Equations (5) and (6) as derived by Mihelcic et al. (2003)

$$268 P(O_3) = k_{HO_2+NO}[HO_2][NO] + \sum_i(k_{RO_2+NO}^i[RO_2^i][NO]) - k_{OH+NO_2}[OH][NO_2] - L(O_3)$$

269 (5)

$$270 L(O_3) = (\theta j(O^1D) + k_{OH+O_3}[OH] + k_{HO_2+O_3}[HO_2] + \sum_j(k_{alkene+O_3}^j[alkene^j]))[O_3]$$

271 (6)

272 where θ is the fraction of O¹D from ozone photolysis that reacts with water vapor, and
273 i and j represent the number of species of RO₂ and alkenes, respectively.

274 The box model was used to evaluate the impact of missing VOC_R on the O₃
275 production rate. In the base scenario, the box model was constrained by all measured
276 inorganic and organic gases but the missing VOC_R was not considered. To consider the
277 missing VOC_R in the box model, we additionally increased the concentration of
278 NMHCs to exactly compensate for the missing VOC_R by multiplying a factor, on the
279 basis of measured NMHC concentrations. We simulated four scenarios by increasing
280 the concentration of: (1) n-pentane, (2) ethylene, (3) toluene, (4) all measured 56
281 NMHCs. For the scenario of increasing all 56 NMHCs, concentrations of 56 NMHC
282 species were increased by multiplying the same factor. Given that the VOC_R of
283 unconstrained secondary products increases with the increase in the concentration of
284 NMHCs, several attempts of different values are needed to determine the increasing
285 factor.

286 **3 Results and discussion**

287 **3.1 Quantification of missing VOC_R during the campaign**

288 **Figure 1** shows the time series of measured RO_H, calculated RO_H according to all
289 measured reactive gases, and missing VOC_R (the gap between measured and calculated
290 RO_H) in Guangzhou. By using GC-MS/FID, we measured 56 NMHCs. By using PTR-
291 ToF-MS, we measured 159 VOCs and 128 of them were difficult to be measured before.
292 Besides the alkanes, with carbons less than 12, PTR-ToF-MS can also measure alkanes
293 with more carbons (C12–C20). With regard to OVOCs, not only common OVOC
294 species including formaldehyde and C2-C4 carbonyls but also carbonyls with more
295 carbons (C5–C10) and some N-containing OVOC species such as nitrophenol and
296 methyl nitrophenol, were measured by PTR-ToF-MS. Thanks to these additional
297 measured VOCs, the measured RO_H was close to the calculated RO_H within 20% in most
298 periods. In some periods the missing VOC_R was negative, which is probably due to the
299 uncertainty in the measurements of RO_H and reactive gases. The negative missing VOC_R
300 primarily occurred in the afternoon (12:00–17:00) when the photochemistry was most
301 active. Nevertheless, there were still some days exhibiting remarkable missing VOC_R.
302 The days with missing VOC_R of more than 25% of total RO_H, namely high missing-
303 VOC_R days, are indicated by yellow background in **Fig. 1a**. The largest missing VOC_R
304 occurred on October 15th, 16th, 25th and 26th, with average values of 16 s⁻¹. During the
305 period of October 24th to 26th, the total RO_H was highest and the missing VOC_R was also
306 relatively high among all days. **Figure 1b** shows the contribution of different species
307 classifications to total RO_H during high missing-VOC_R days. Inorganic species, NMHCs
308 and OVOCs account for 34%, 13% and 14% of total RO_H, respectively, with missing
309 VOC_R accounting for 39%. The fraction of missing VOC_R (39%) during the high
310 missing-VOC_R days is comparable to measurements in Los Angeles 2010 (Griffith et
311 al., 2016) and in Seoul 2016 (Sanchez et al., 2021).

312 We evaluated the uncertainty of the missing VOC_R. The uncertainty of the RO_H

删除了: many VOC species that

删除了: NMHCs

删除了: species

删除了: higher NMHCs

删除了: and several organic nitrates

删除了: -

319 measurement was 15%. In addition, according to reports of Jet Propulsion Laboratory
320 (Burkholder et al., 2020), reaction rate constants used for the calculation of R_{OH} in Eq
321 (3) have uncertainties of 5%–30%, depending on different species. We took the
322 uncertainties in the reaction rate constants and the measurements of all reactive gases
323 into account when calculating R_{OH} , according to error propagation. As a result, the
324 uncertainties in the missing VOC_R are 3.8 s^{-1} and 5.2 s^{-1} for the whole measurement
325 period and the high missing- VOC_R days, respectively. The average missing VOC_R
326 during the high missing- VOC_R days is 13 s^{-1} , which is significantly higher than the
327 uncertainty of 5.2 s^{-1} , suggesting that the missing VOC_R really exists during the high
328 missing- VOC_R days.

删除了: se

删除了: the

330 3.2 The sources of missing VOC_R

331 To explore the sources of missing VOC_R during the whole measurement period,
332 we investigated the correlation between missing VOC_R and tracers characterizing
333 primary emissions (CO , NO_X and $NMHCs$) and secondary production ($O_X=O_3+NO_2$
334 and formic acid). The correlation of missing VOC_R with CO , reactivity of $NMHCs$
335 ($NMHC_R$) and NO_X is moderate, with correlation coefficient (R) in the range of 0.47–
336 0.56 (Fig. 2a and b, and Fig. S2) while there is no significant correlation of missing
337 VOC_R with O_X and formic acid (Fig. 2c and Fig. S2). Furthermore, there is no
338 significant correlation between missing VOC_R and acetonitrile which is a tracer of
339 biomass burning (de Gouw et al., 2003; Wang et al., 2007) (Fig. S2), indicating that
340 biomass burning was not a major contributor to missing VOC_R during this campaign.
341 In terms of the diurnal variation, the missing VOC_R was higher in the morning (7:00–
342 10:00) and evening (18:00–22:00) when the anthropogenic emissions, especially
343 vehicle exhaust were intensive, and was lower in the afternoon when the
344 photochemistry was most active (Fig. 2d). The diurnal profile of missing VOC_R was
345 similar to those of CO , NO_X and $NMHC_R$. In contrast, the diurnal profiles of secondary
346 species including O_X , formic acid and acetic acid, which peaked in the afternoon,

删除了: S1

删除了: S1

删除了: S1

删除了: .

353 evidently differ from the diurnal profile of missing VOC_R (Fig. S3). Further, we
354 investigated the influence of air mass aging on missing VOC_R. The ratio of ethylbenzene
355 to m,p-xylene was used to characterize the degree of air mass aging (De Gouw et al.,
356 2005; Yuan et al., 2013). A higher ratio of ethylbenzene to m,p-xylene corresponds to a
357 higher degree of air mass aging as the m,p-xylene has a larger reaction rate constant
358 ($18.9 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$) than ethylbenzene ($7.0 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$) when
359 reacting with the major oxidant - OH radicals. As shown in Fig. 2e, missing VOC_R
360 decreases with the ratio of ethylbenzene to m,p-xylene. Given that secondary
361 production generally increased with air mass aging, this result further demonstrates that
362 missing VOC_R was not caused by enhanced secondary production.

363 During the high missing- VOC_R days, the correlation coefficient for missing VOC_R
364 versus CO is 0.76 (Fig. 3a), which is higher than that in the whole measurement period
365 (0.56) shown in Fig. 2a. We then quantify the sources of missing VOC_R during the high
366 missing- VOC_R days by applying MLR. The fitted coefficient a is $0.031 \text{ s}^{-1} \text{ ppb}^{-1}$, b is
367 $0.012 \text{ s}^{-1} \text{ ppb}^{-1}$, c is $1.8 \text{ s}^{-1} \text{ ppb}^{-1}$ and $C_{\text{background}}$ is 1.3 s^{-1} . The coefficient of determination
368 (R^2) for the MLR is 0.68. As shown in Fig. 3b, anthropogenic emissions were the largest
369 contributor to missing VOC_R, accounting for 70% of missing VOC_R. Secondary
370 production, biogenic emissions and background contribution played a minor role in
371 missing VOC_R (13%, 7%, 10%, respectively). The parametric relationship between
372 missing VOC_R and relevant tracers established by MLR provides a valid approach to
373 estimate the missing VOC_R according to readily available gases including CO, O_x and
374 isoprene.

375 Although anthropogenic emissions are identified to be the major source of missing
376 VOC_R, which species dominantly contribute to the missing VOC_R remains unclear. A
377 potential source is the unmeasured branched alkenes for their high reactivity, previously
378 observed from vehicle exhaust (Nakashima et al., 2010) and gasoline evaporation
379 emissions (Wu et al., 2015). Another possible source is emitted OVOCs with a more
380 complex functional group that cannot be accurately measured. In addition, directly
381 emitted semi-volatile and intermediate volatility organic compounds are also possible

删除了: S2

删除了: aging

删除了: es

删除了: aging degree of air masses

删除了: Given the larger missing VOC_R level during the high missing- VOC_R days, we focus on the high missing- VOC_R days in the following analysis.

删除了: In addition, the correlation between missing VOC_R and O_x is weak with $R=-0.25$ during the high missing-VOC_R days (Fig. 3b).

删除了: 3c

393 sources of missing VOC_R (Stewart et al., 2021).

394 3.3 The impact of missing VOC_R on O₃ sensitivity regimes

395 The reaction of OH with VOCs is key to the propagation and amplification of OH
396 radicals, thus determining the ozone production rate (Tonnesen and Dennis, 2000). The
397 box model was used to evaluate the impact of missing VOC_R on the O₃ production rate
398 during high missing-VOC_R days. The setting of model simulations for different
399 scenarios are depicted in Section 2.6. Under the base scenario, on average the measured
400 VOC_R of n-pentane, ethylene, toluene and all 56 NMHCs are 0.14 s⁻¹, 0.53 s⁻¹, 0.60 s⁻¹
401 and 4.6 s⁻¹ respectively. To consider the missing VOC_R (on average of 13 s⁻¹) in the
402 model, four scenarios were simulated by additionally increasing n-pentane, ethylene,
403 toluene and 56 NMHCs by a factor of 70, 16, 13.3 and 1.9, respectively. These
404 increasing factors led to an additional increase in VOC_R of both NMHCs and
405 unconstrained secondary products, which exactly compensated for the missing VOC_R.
406 **Figure 4** shows the simulated P(O₃) for the base scenario and the scenarios considering
407 missing VOC_R. The daytime average P(O₃) under the scenarios considering missing
408 VOC_R is a factor of 1.5-4.5 for the results under the base scenario. The difference in
409 added species has a large effect on P(O₃). Adding toluene causes a larger increase in
410 P(O₃) than adding n-pentane or ethene, as toluene has a stronger ability to amplify the
411 production of radicals.

412 O₃ precursor sensitivity depends on the dominant loss pathways of RO_X radicals
413 (RO_X=OH+HO₂+RO₂). O₃ production is NO_X-limited if the self-reaction of peroxy
414 radicals (HO₂ and RO₂) dominates the RO_X sink, and VOC-limited if the reaction of
415 NO₂ with OH dominates (Kleinman et al., 1997; Kleinman et al., 2001). Accordingly,
416 the ratio of RO_X sink induced by OH+NO₂ reaction to the total rate of the two RO_X
417 sinks, i.e., L_N/Q, is used to identify O₃ sensitivity regimes. O₃ production is NO_X-
418 limited if L_N/Q is lower than 0.5, otherwise, it is VOC-limited (Kleinman et al., 1997).

$$419 L_N/Q = \frac{k_{OH+NO_2}[OH][NO_2]}{k_{HO_2+RO_2}[HO_2][RO_2] + k_{HO_2+HO_2}[HO_2][HO_2] + k_{OH+HO_2}[OH][HO_2] + k_{OH+NO_2}[OH][NO_2]}$$

(7)

420

删除了:-

删除了: In the base scenario, the box model was constrained by all measured inorganic and organic gases but the missing VOC_R was not considered. To consider the missing VOC_R in the box model, we increased all measured NMHC species by a factor that can compensates for the missing VOC_R. In addition, we also try adding a single VOC species to represent the missing VOC_R. Three typical VOC species were added respectively, including n-pentane, ethylene and toluene.

删除了: one

删除了:-

删除了: The uncertainty in missing VOC_R leads to 13-17% uncertainties in the threshold of NO_X for scenarios considering missing VOC_R.

436 As shown in **Fig. 5a**, under the base scenario, L_N/Q remained at a stable and high
437 level (>0.9) during the daytime when photochemical production of ozone occurs,
438 indicating O_3 production was VOC-limited. Under the scenarios considering missing
439 VOC_R , L_N/Q decreased significantly regardless of which VOC species was added,
440 compared to the base scenario. Adding toluene caused the largest decrease in L_N/Q ,
441 followed by adding all measured NMHC species, adding the alkane and adding the
442 alkene. It is worth noting that adding toluene and all measured NMHC species caused
443 the L_N/Q to be close to 0.5 in the afternoon, indicating that the O_3 production shifted to
444 transitional or NO_X -limited regimes in these scenarios. **Fig. 5b** shows the changes in
445 radical sinks before and after considering missing VOC_R . All radical sinks including
446 self-reactions of peroxy radicals and $OH+NO_2$ reaction increased after considering
447 missing VOC_R . Nevertheless, the increased proportion of the self-reactions of peroxy
448 radicals was larger than that of $OH+NO_2$ reaction, leading to a decrease in L_N/Q and
449 thus a shift toward NO_X -limited regime.

450 **Figure 5c** shows the dependence of daily peak O_3 concentrations on NO_X
451 concentrations, which was calculated by the box model for the base scenario and the
452 scenario considering missing VOC_R . The NO_X concentration level corresponding to the
453 maximum of O_3 concentrations was determined. This NO_X concentration level reflects
454 the threshold to distinguish between VOC-limited and NO_X -limited regimes. The larger
455 threshold of NO_X represents a higher possibility of ozone production in NO_X limited
456 regime. The threshold of NO_X for the scenario considering missing VOC_R is 46% higher
457 than for the base scenario. Note that the uncertainty in missing VOC_R leads to 17%
458 uncertainty in the threshold of NO_X for the scenario considering missing VOC_R . Overall,
459 **Fig. 5** suggests that omitting the missing VOC_R will overestimate the degree of the
460 VOC-limited regime and thus overestimate the effect of VOCs abatement in reducing
461 ozone pollution, which in turn may mislead ozone control strategy.

462 3.4 Atmospheric implications

463 Although many previous studies have reported that photochemical production

464 processes and biogenic emissions are important sources of missing VOC_R (Lou et al.,
465 2010;Dolgorouky et al., 2012;Yang et al., 2017;Sanchez et al., 2021;Di Carlo et al.,
466 2004), we find that anthropogenic emissions may dominate the missing VOC_R in urban
467 regions. In zero-dimensional box models and three-dimensional chemistry-transport
468 models, the input of VOCs emission information mainly contains well-studied simple-
469 structure alkanes, alkenes and aromatics, while those unmeasured/unknown VOC
470 species have been neglected. This will lead to biases in quantifying ozone production
471 and diagnosing ozone sensitivity regimes. Our study demonstrates that the ambient
472 measurement of R_{OH} at urban sites can provide quantification of missing VOC_R, which
473 can be used in models to account for the missing VOC_R from anthropogenic emissions.
474 In addition, the parametric equation of missing VOC_R derived from MLR method (Eq
475 (4)) here can be used to estimate missing VOC_R according to measurements of CO, O_x
476 and isoprene. Further study should try to parse the specific sources of the missing VOC_R,
477 e.g., whether the missing VOC_R is from intermediate-volatility and semivolatile organic
478 compounds emitted from vehicles or whether it is from some other sources.
479 Furthermore, future studies can focus on direct measurements of missing VOC_R for
480 various emission sources to develop a comprehensive emission inventory of missing
481 VOC_R, which will help to improve O₃ pollution mitigation strategies.

删除了: versus CO developed

删除了: CO

删除了: Besides CO, other specific classes of hydrocarbons are also expected to be used as tracers for the development of the parametric equation.

483 Acknowledgement

484 This work was supported by the National Natural Science Foundation of China (grant
485 No. , 42121004, 42275103, 42230701, 42175135). This work was also supported by
486 Special Fund Project for Science and Technology Innovation Strategy of Guangdong
487 Province (Grant No.2019B121205004).

491 References

492 Atkinson, R.: Atmospheric chemistry of VOCs and NO_x, Atmos. Environ., 34, 2063-2101, 2000.

498 Atkinson, R., and Arey, J.: Atmospheric degradation of volatile organic compounds, *Chemical*
499 *reviews*, 103, 4605-4638, 2003.

500 Burkholder, J., Sander, S., Abbatt, J., Barker, J., Cappa, C., Crounse, J., Dibble, T., Huie, R., Kolb,
501 C., and Kurylo, M.: Chemical kinetics and photochemical data for use in atmospheric studies;
502 evaluation number 19, Pasadena, CA: Jet Propulsion Laboratory, National Aeronautics and
503 Space ..., 2020.

504 De Gouw, J., Middlebrook, A., Warneke, C., Goldan, P., Kuster, W., Roberts, J., Fehsenfeld, F.,
505 Worsnop, D., Canagaratna, M., and Pszenny, A.: Budget of organic carbon in a polluted
506 atmosphere: Results from the New England Air Quality Study in 2002, *J. Geophys. Res.-Atmos.*,
507 110, 2005.

508 de Gouw, J. A., Warneke, C., Parrish, D. D., Holloway, J. S., Trainer, M., and Fehsenfeld, F. C.:
509 Emission sources and ocean uptake of acetonitrile (CH₃CN) in the atmosphere, *J. Geophys.*
510 *Res.-Atmos.*, 108, <https://doi.org/10.1029/2002JD002897>, 2003.

511 Di Carlo, P., Brune, W. H., Martinez, M., Harder, H., Leshner, R., Ren, X. R., Thornberry, T., Carroll,
512 M. A., Young, V., Shepson, P. B., Riemer, D., Apel, E., and Campbell, C.: Missing OH
513 reactivity in a forest: Evidence for unknown reactive biogenic VOCs, *Science*, 304, 722-725,
514 10.1126/science.1094392, 2004.

515 Dolgorouky, C., Gros, V., Sarda-Estève, R., Sinha, V., Williams, J., Marchand, N., Sauvage, S.,
516 Poulain, L., Sciare, J., and Bonsang, B.: Total OH reactivity measurements in Paris during the
517 2010 MEGAPOLI winter campaign, *Atmos. Chem. Phys.*, 12, 9593-9612, 10.5194/acp-12-
518 9593-2012, 2012.

519 Fuchs, H., Tan, Z., Lu, K., Bohn, B., Broch, S., Brown, S. S., Dong, H., Gomm, S., Häsel, R., He,
520 L., Hofzumahaus, A., Holland, F., Li, X., Liu, Y., Lu, S., Min, K. E., Rohrer, F., Shao, M.,
521 Wang, B., Wang, M., Wu, Y., Zeng, L., Zhang, Y., Wahner, A., and Zhang, Y.: OH reactivity
522 at a rural site (Wangdu) in the North China Plain: contributions from OH reactants and
523 experimental OH budget, *Atmos. Chem. Phys.*, 17, 645-661, 10.5194/acp-17-645-2017, 2017.

524 Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the earth's
525 atmosphere, *Environ. Sci. Technol.*, 41, 1514-1521, 2007.

526 Griffith, S. M., Hansen, R., Dusanter, S., Michoud, V., Gilman, J., Kuster, W., Veres, P., Graus, M.,
527 de Gouw, J., and Roberts, J.: Measurements of hydroxyl and hydroperoxy radicals during
528 CalNex-LA: Model comparisons and radical budgets, *J. Geophys. Res.-Atmos.*, 121, 4211-4232,
529 2016.

530 Hansen, R. F., Griffith, S. M., Dusanter, S., Rickly, P. S., Stevens, P. S., Bertman, S. B., Carroll, M.
531 A., Erickson, M. H., Flynn, J. H., Grossberg, N., Jobson, B. T., Lefter, B. L., and Wallace, H.
532 W.: Measurements of total hydroxyl radical reactivity during CABINEX 2009 – Part 1:
533 field measurements, *Atmos. Chem. Phys.*, 14, 2923-2937, 10.5194/acp-14-2923-2014, 2014.

534 Hansen, R. F., Blocquet, M., Schoemaeker, C., Léonardis, T., Locoge, N., Fittschen, C., Hanoune,
535 B., Stevens, P. S., Sinha, V., and Dusanter, S.: Intercomparison of the comparative reactivity
536 method (CRM) and pump-probe technique for measuring total OH reactivity in an urban
537 environment, *Atmos. Meas. Tech.*, 8, 4243-4264, 10.5194/amt-8-4243-2015, 2015.

538 Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the
539 Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile
540 organic compounds, *Atmos. Chem. Phys.*, 3, 181-193, 10.5194/acp-3-181-2003, 2003.

541 Kleinman, L. I.: Low and high NO_x tropospheric photochemistry, *J. Geophys. Res.-Atmos.*, 99,
542 16831-16838, 1994.

543 Kleinman, L. I., Daum, P. H., Lee, J. H., Lee, Y. N., Nunnermacker, L. J., Springston, S. R.,
544 Newman, L., Weinstein-Lloyd, J., and Sillman, S.: Dependence of ozone production on NO and
545 hydrocarbons in the troposphere, *Geophys. Res. Lett.*, 24, 2299-2302, 1997.

546 Kleinman, L. I., Daum, P. H., Lee, Y. N., Nunnermacker, L. J., Springston, S. R., Weinstein-Lloyd,
547 J., and Rudolph, J.: Sensitivity of ozone production rate to ozone precursors, *Geophys. Res.*
548 *Lett.*, 28, 2903-2906, 2001.

549 Kovacs, T., Brune, W., Harder, H., Martinez, M., Simpasa, J., Frost, G., Williams, E., Jobson, T.,
550 Stroud, C., and Young, V.: Direct measurements of urban OH reactivity during Nashville SOS
551 in summer 1999, *Journal of Environmental Monitoring*, 5, 68-74, 2003.

552 Lefohn, A. S., Malley, C. S., Smith, L., Wells, B., Hazucha, M., Simon, H., Naik, V., Mills, G.,
553 Schultz, M. G., and Paoletti, E.: Tropospheric ozone assessment report: Global ozone metrics
554 for climate change, human health, and crop/ecosystem research, *Elem. Sci. Anth.*, 6, 2018.

555 Li, K., Jacob, D. J., Liao, H., Shen, L., Zhang, Q., and Bates, K. H.: Anthropogenic drivers of 2013–
556 2017 trends in summer surface ozone in China, *Proc. National Acad. Sci.*, 116, 422-427, 2019.

557 Li, X.-B., Yuan, B., Parrish, D. D., Chen, D., Song, Y., Yang, S., Liu, Z., and Shao, M.: Long-term
558 trend of ozone in southern China reveals future mitigation strategy for air pollution, *Atmos.*
559 *Environ.*, 269, 118869, 2022.

560 Lou, S., Holland, F., Rohrer, F., Lu, K., Bohn, B., Brauers, T., Chang, C., Fuchs, H., Häsel, R.,
561 and Kita, K.: Atmospheric OH reactivities in the Pearl River Delta–China in summer 2006:
562 measurement and model results, *Atmos. Chem. Phys.*, 10, 11243-11260, 2010.

563 Lu, K. D., Hofzumahaus, A., Holland, F., Bohn, B., Brauers, T., Fuchs, H., Hu, M., Haseler, R.,
564 Kita, K., Kondo, Y., Li, X., Lou, S. R., Oebel, A., Shao, M., Zeng, L. M., Wahner, A., Zhu, T.,
565 Zhang, Y. H., and Rohrer, F.: Missing OH source in a suburban environment near Beijing:
566 observed and modelled OH and HO₂ concentrations in summer 2006, *Atmospheric Chemistry*
567 *and Physics*, 13, 1057-1080, 10.5194/acp-13-1057-2013, 2013.

568 Lu, X., Hong, J. Y., Zhang, L., Cooper, O. R., Schultz, M. G., Xu, X. B., Wang, T., Gao, M., Zhao,
569 Y. H., and Zhang, Y. H.: Severe Surface Ozone Pollution in China: A Global Perspective,
570 *Environ. Sci. Technol. Lett.*, 5, 487-494, 10.1021/acs.estlett.8b00366, 2018.

571 Mihelcic, D., Holland, F., Hofzumahaus, A., Hoppe, L., Konrad, S., Müsgen, P., Pätz, H. W.,
572 Schäfer, H. J., Schmitz, T., and Volz-Thomas, A.: Peroxy radicals during BERLIOZ at
573 Pabstthum: Measurements, radical budgets and ozone production, *J. Geophys. Res.-Atmos.*, 108,
574 2003.

575 Monks, P. S., Archibald, A. T., Colette, A., Cooper, O., Coyle, M., Derwent, R., Fowler, D., Granier,
576 C., Law, K. S., Mills, G. E., Stevenson, D. S., Tarasova, O., Thouret, V., von Schneidemesser,
577 E., Sommariva, R., Wild, O., and Williams, M. L.: Tropospheric ozone and its precursors from
578 the urban to the global scale from air quality to short-lived climate forcer, *Atmos. Chem. Phys.*,
579 15, 8889-8973, 10.5194/acp-15-8889-2015, 2015.

580 Nakashima, Y., Kamei, N., Kobayashi, S., and Kajii, Y.: Total OH reactivity and VOC analyses for
581 gasoline vehicular exhaust with a chassis dynamometer, *Atmos. Environ.*, 44, 468-475,
582 <https://doi.org/10.1016/j.atmosenv.2009.11.006>, 2010.

583 Nakashima, Y., Kato, S., Greenberg, J., Harley, P., Karl, T., Turnipseed, A., Apel, E., Guenther, A.,
584 Smith, J., and Kajii, Y.: Total OH reactivity measurements in ambient air in a southern Rocky

585 mountain ponderosa pine forest during BEACHON-SRM08 summer campaign, *Atmos.*
586 *Environ.*, 85, 1-8, <https://doi.org/10.1016/j.atmosenv.2013.11.042>, 2014.

587 Nölscher, A. C., Yañez-Serrano, A. M., Wolff, S., de Araujo, A. C., Lavrič, J. V., Kesselmeier, J.,
588 and Williams, J.: Unexpected seasonality in quantity and composition of Amazon rainforest air
589 reactivity, *Nature Communications*, 7, 10383, 10.1038/ncomms10383, 2016.

590 Paoletti, E., De Marco, A., Beddows, D. C., Harrison, R. M., and Manning, W. J.: Ozone levels in
591 European and USA cities are increasing more than at rural sites, while peak values are
592 decreasing, *Environmental Pollution*, 192, 295-299, 2014.

593 Praplan, A. P., Tykka, T., Chen, D., Boy, M., Taipale, D., Vakkari, V., Zhou, P. T., Petaja, T., and
594 Hellen, H.: Long-term total OH reactivity measurements in a boreal forest, *Atmos. Chem. Phys.*,
595 19, 14431-14453, 10.5194/acp-19-14431-2019, 2019.

596 Sadanaga, Y., Yoshino, A., Kato, S., and Kajii, Y.: Measurements of OH reactivity and
597 photochemical ozone production in the urban atmosphere, *Environ. Sci. Technol.*, 39, 8847-
598 8852, 2005.

599 Sanchez, D., Seco, R., Gu, D., Guenther, A., Mak, J., Lee, Y., Kim, D., Ahn, J., Blake, D., Herndon,
600 S., Jeong, D., Sullivan, J. T., McGee, T., Park, R., and Kim, S.: Contributions to OH reactivity
601 from unexplored volatile organic compounds measured by PTR-ToF-MS – a case study in a
602 suburban forest of the Seoul metropolitan area during the Korea–United States Air Quality
603 Study (KORUS-AQ) 2016, *Atmos. Chem. Phys.*, 21, 6331-6345, 10.5194/acp-21-6331-2021,
604 2021.

605 Sekimoto, K., Li, S.-M., Yuan, B., Koss, A., Coggon, M., Warneke, C., and de Gouw, J.: Calculation
606 of the sensitivity of proton-transfer-reaction mass spectrometry (PTR-MS) for organic trace
607 gases using molecular properties, *International Journal of Mass Spectrometry*, 421, 71-94,
608 10.1016/j.ijms.2017.04.006, 2017.

609 Shetter, R. E., and Müller, M.: Photolysis frequency measurements using actinic flux
610 spectroradiometry during the PEM-Tropics mission: Instrumentation description and some
611 results, *J. Geophys. Res.-Atmos.*, 104, 5647-5661, <https://doi.org/10.1029/98JD01381>, 1999.

612 Shirley, T. R., Brune, W. H., Ren, X., Mao, J., Leshner, R., Cardenas, B., Volkamer, R., Molina, L.
613 T., Molina, M. J., Lamb, B., Velasco, E., Jobson, T., and Alexander, M.: Atmospheric oxidation
614 in the Mexico City Metropolitan Area (MCMA) during April 2003, *Atmos. Chem. Phys.*, 6,
615 2753-2765, 10.5194/acp-6-2753-2006, 2006.

616 Sillman, S., Logan, J. A., and Wofsy, S. C.: The sensitivity of ozone to nitrogen oxides and
617 hydrocarbons in regional ozone episodes, *J. Geophys. Res.-Atmos.*, 95, 1837-1851, 1990.

618 Sinha, V., Williams, J., Crowley, J., and Lelieveld, J.: The Comparative Reactivity Method—a new
619 tool to measure total OH Reactivity in ambient air, *Atmos. Chem. Phys.*, 8, 2213-2227, 2008.

620 Song, K., Gong, Y., Guo, S., Lv, D., Wang, H., Wan, Z., Yu, Y., Tang, R., Li, T., Tan, R., Zhu, W.,
621 Shen, R., and Lu, S.: Investigation of partition coefficients and fingerprints of atmospheric gas-
622 and particle-phase intermediate volatility and semi-volatile organic compounds using pixel-
623 based approaches, *Journal of Chromatography A*, 1665, 462808,
624 <https://doi.org/10.1016/j.chroma.2022.462808>, 2022.

625 Stewart, G. J., Nelson, B. S., Acton, W. J. F., Vaughan, A. R., Farren, N. J., Hopkins, J. R., Ward,
626 M. W., Swift, S. J., Arya, R., Mondal, A., Jangir, R., Ahlawat, S., Yadav, L., Sharma, S. K.,
627 Yunus, S. S. M., Hewitt, C. N., Nemitz, E., Mullinger, N., Gadi, R., Sahu, L. K., Tripathi, N.,
628 Rickard, A. R., Lee, J. D., Mandal, T. K., and Hamilton, J. F.: Emissions of intermediate-

629 volatility and semi-volatile organic compounds from domestic fuels used in Delhi, India, *Atmos.*
630 *Chem. Phys.*, 21, 2407-2426, 10.5194/acp-21-2407-2021, 2021.

631 Tonnesen, G. S., and Dennis, R. L.: Analysis of radical propagation efficiency to assess ozone
632 sensitivity to hydrocarbons and NO_x: 1. Local indicators of instantaneous odd oxygen
633 production sensitivity, *J. Geophys. Res.-Atmos.*, 105, 9213-9225, 2000.

634 Wang, C., Yuan, B., Wu, C., Wang, S., Qi, J., Wang, B., Wang, Z., Hu, W., Chen, W., Ye, C., Wang,
635 W., Sun, Y., Wang, C., Huang, S., Song, W., Wang, X., Yang, S., Zhang, S., Xu, W., Ma, N.,
636 Zhang, Z., Jiang, B., Su, H., Cheng, Y., Wang, X., and Shao, M.: Measurements of higher
637 alkanes using NO⁺ chemical ionization in PTR-ToF-MS: important contributions of higher
638 alkanes to secondary organic aerosols in China, *Atmospheric Chemistry and Physics*, 20,
639 14123-14138, 10.5194/acp-20-14123-2020, 2020a.

640 Wang, H., Zhang, X., and Chen, Z.: Development of DNPH/HPLC method for the measurement of
641 carbonyl compounds in the aqueous phase: applications to laboratory simulation and field
642 measurement, *Environmental Chemistry*, 6, 389-397, <https://doi.org/10.1071/EN09057>, 2009.

643 Wang, M., Zeng, L., Lu, S., Shao, M., Liu, X., Yu, X., Chen, W., Yuan, B., Zhang, Q., and Hu, M.:
644 Development and validation of a cryogen-free automatic gas chromatograph system (GC-
645 MS/FID) for online measurements of volatile organic compounds, *Anal. Methods*, 6, 9424-9434,
646 2014.

647 Wang, Q. Q., Shao, M., Liu, Y., William, K., Paul, G., Li, X. H., Liu, Y. A., and Lu, S. H.: Impact
648 of biomass burning on urban air quality estimated by organic tracers: Guangzhou and Beijing
649 as cases, *Atmos. Environ.*, 41, 8380-8390, 10.1016/j.atmosenv.2007.06.048, 2007.

650 Wang, W., Li, X., Shao, M., Hu, M., Zeng, L., Wu, Y., and Tan, T.: The impact of aerosols on
651 photolysis frequencies and ozone production in Beijing during the 4-year period 2012–2015,
652 *Atmos. Chem. Phys.*, 19, 9413-9429, 10.5194/acp-19-9413-2019, 2019.

653 Wang, W., Parrish, D. D., Li, X., Shao, M., Liu, Y., Mo, Z., Lu, S., Hu, M., Fang, X., and Wu, Y.:
654 Exploring the drivers of the increased ozone production in Beijing in summertime during 2005–
655 2016, *Atmos. Chem. Phys.*, 20, 15617-15633, 2020b.

656 Wang, W., Parrish, D. D., Wang, S., Bao, F., Ni, R., Li, X., Yang, S., Wang, H., Cheng, Y., and Su,
657 H.: Long-term trend of ozone pollution in China during 2014–2020: distinct seasonal and spatial
658 characteristics and ozone sensitivity, *Atmos. Chem. Phys.*, 22, 8935-8949, 10.5194/acp-22-
659 8935-2022, 2022.

660 Wu, C., Wang, C., Wang, S., Wang, W., Yuan, B., Qi, J., Wang, B., Wang, H., Wang, C., and Song,
661 W.: Measurement report: Important contributions of oxygenated compounds to emissions and
662 chemistry of volatile organic compounds in urban air, *Atmos. Chem. Phys.*, 20, 14769-14785,
663 2020.

664 Wu, Y., Yang, Y. D., Shao, M., and Lu, S. H.: Missing in total OH reactivity of VOCs from gasoline
665 evaporation, *Chinese Chemical Letters*, 26, 1246-1248, 10.1016/j.ccl.2015.05.047, 2015.

666 Xie, X., Shao, M., Liu, Y., Lu, S., Chang, C.-C., and Chen, Z.-M.: Estimate of initial isoprene
667 contribution to ozone formation potential in Beijing, China, *Atmos. Environ.*, 42, 6000-6010,
668 2008.

669 Xu, R., Li, X., Dong, H., Lv, D., Kim, N., Yang, S., Wang, W., Chen, J., Shao, M., and Lu, S.: Field
670 observations and quantifications of atmospheric formaldehyde partitioning in gaseous and
671 particulate phases, *Sci. Total Environ.*, 808, 152122, 2022.

672 Yang, Y., Liao, H., and Lou, S.: Increase in winter haze over eastern China in recent decades: Roles
673 of variations in meteorological parameters and anthropogenic emissions, *J. Geophys. Res.-*
674 *Atmos.*, 121, 13,050-013,065, <https://doi.org/10.1002/2016JD025136>, 2016a.

675 Yang, Y., Shao, M., Wang, X., Nölscher, A. C., Kessel, S., Guenther, A., and Williams, J.: Towards
676 a quantitative understanding of total OH reactivity: A review, *Atmos. Environ.*, 134, 147-161,
677 2016b.

678 Yang, Y., Shao, M., Keßel, S., Li, Y., Lu, K., Lu, S., Williams, J., Zhang, Y., Zeng, L., Nölscher,
679 A. C., Wu, Y., Wang, X., and Zheng, J.: How the OH reactivity affects the ozone production
680 efficiency: case studies in Beijing and Heshan, China, *Atmos. Chem. Phys.*, 17, 7127-7142,
681 10.5194/acp-17-7127-2017, 2017.

682 Yoshino, A., Sadanaga, Y., Watanabe, K., Kato, S., Miyakawa, Y., Matsumoto, J., and Kajii, Y.:
683 Measurement of total OH reactivity by laser-induced pump and probe technique—
684 comprehensive observations in the urban atmosphere of Tokyo, *Atmos. Environ.*, 40, 7869-
685 7881, <https://doi.org/10.1016/j.atmosenv.2006.07.023>, 2006.

686 Yu, Y., Cheng, P., Li, H., Yang, W., Han, B., Song, W., Hu, W., Wang, X., Yuan, B., Shao, M.,
687 Huang, Z., Li, Z., Zheng, J., Wang, H., and Yu, X.: Budget of nitrous acid (HONO) at an urban
688 site in the fall season of Guangzhou, China, *Atmos. Chem. Phys.*, 22, 8951-8971, 10.5194/acp-
689 22-8951-2022, 2022.

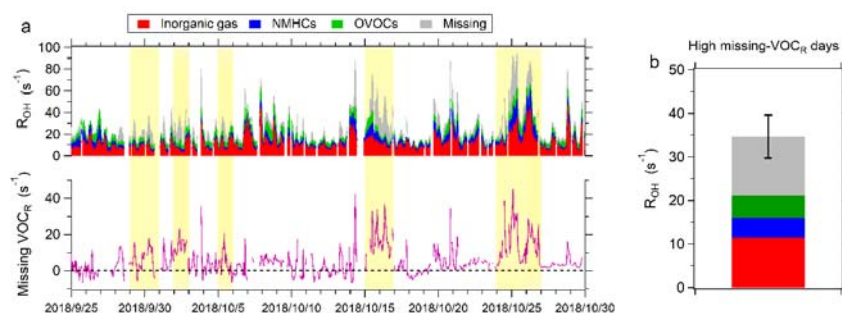
690 Yuan, B., Chen, W., Shao, M., Wang, M., Lu, S., Wang, B., Liu, Y., Chang, C.-C., and Wang, B.:
691 Measurements of ambient hydrocarbons and carbonyls in the Pearl River Delta (PRD), China,
692 *Atmos. Res.*, 116, 93-104, 2012.

693 Yuan, B., Hu, W., Shao, M., Wang, M., Chen, W., Lu, S., Zeng, L., and Hu, M.: VOC emissions,
694 evolutions and contributions to SOA formation at a receptor site in eastern China, *Atmos. Chem.*
695 *Phys.*, 13, 8815-8832, 2013.

696 Yuan, B., Koss, A. R., Warneke, C., Coggon, M., Sekimoto, K., and de Gouw, J. A.: Proton-
697 Transfer-Reaction Mass Spectrometry: Applications in Atmospheric Sciences, *Chemical*
698 *Reviews*, 117, 13187-13229, 10.1021/acs.chemrev.7b00325, 2017.

699

700



701

702 **Figure 1. The level of missing VOC_R during the measurements in Guangzhou. (a)**

703 Time series of measured R_{OH} and calculated R_{OH} from all measured reactive gases in

704 Guangzhou. Yellow background represents the high missing- VOC_R days with missing

705 VOC_R accounting for more than 30% of total R_{OH} . (b) Contributions of different

706 compositions to R_{OH} in high missing- VOC_R days. The error bar represents standard

707 deviation of missing VOC_R .

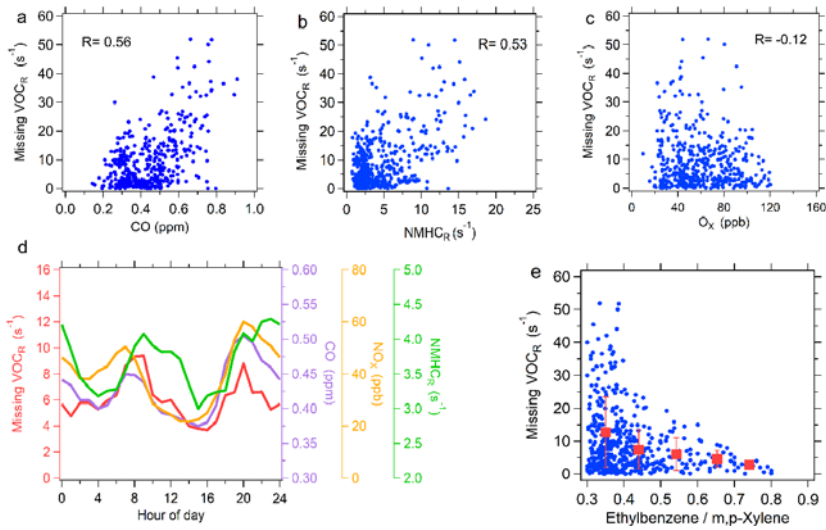
708

709

710

711

712



713

714 **Figure 2. Correlation of missing VOC_R with major tracers during the whole**

715 **measurement period.** (a-c) Correlation of missing VOC_R with CO, OH reactivity of

716 NMHCs (NMHC_R) and O_x. Each point represents hourly data. (d) Diurnal variations

717 in missing VOC_R, CO, NO_x and NMHCs. (e) The dependence of missing VOC_R on

718 ethylbenzene to m, p-xylene ratio. The red squares indicate the mean values of missing

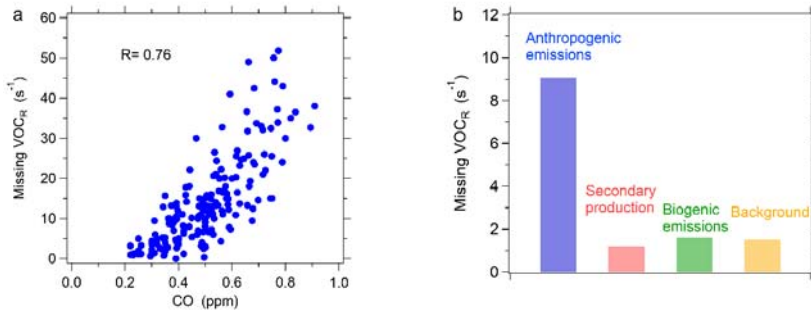
719 VOCR in different ranges of ethylbenzene/m,p-xylene with classification width of 0.1,

720 and the error bars represent standard deviation.

721

722

723



724

725 **Figure 3. The source apportionment of missing VOC_R in high missing-VOC_R days.**

726 (a) Correlation of missing VOC_R with CO. Each point represents hourly data. (b)

727 Contributions of different sources to missing VOC_R according to the MLR.

728

删除了: (b) Correlation of missing VOC_R with O_x. In (a) and (b), each

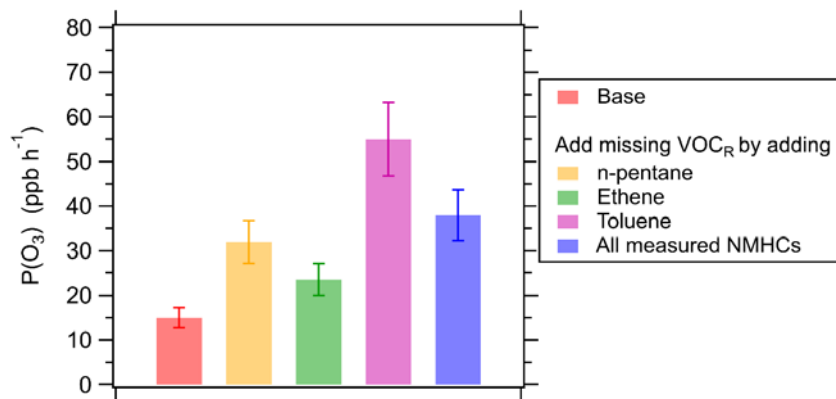


Figure 4. Simulated daytime mean $P(O_3)$ for the base scenario (without missing VOC_R) and the scenario considering missing VOC_R , respectively, in high-missing VOC_R days. The missing VOC_R is considered by adding individual species (n-pentane, ethene or toluene) or increasing all measured NMHCs to compensate for the missing VOC_R . The error bar represents standard deviation of $P(O_3)$ induced by the uncertainty of missing VOC_R .

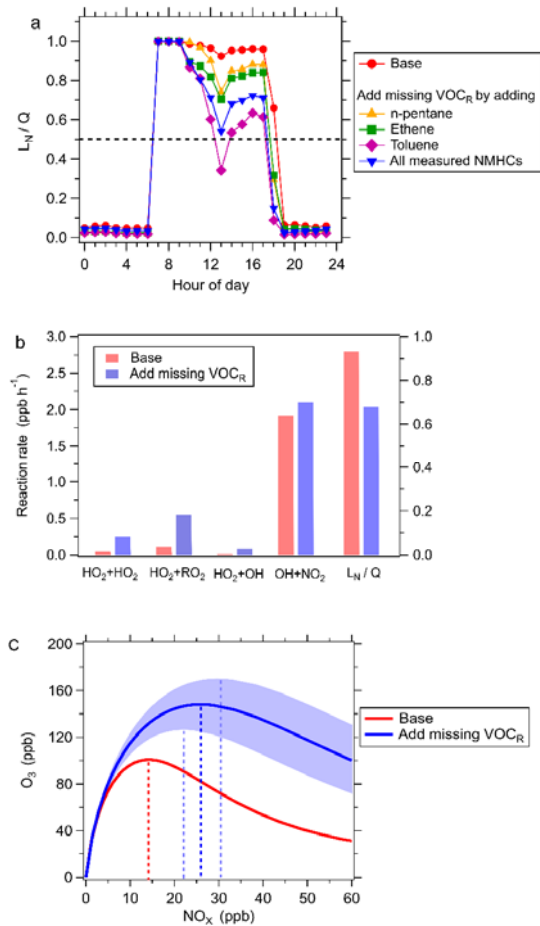


Figure 5. The impact of missing VOC_R on O₃ sensitivity for the high-missing VOC_R days. (a) Diurnal variations in L_N/Q for the base scenario and the scenarios considering missing VOC_R. The missing VOC_R is considered by adding individual species (n-pentane, ethene or toluene) or increasing all measured NMHCs to fill the missing VOC_R. The dashed line represents the threshold value of L_N/Q that distinguishes VOC-limited and NO_x-limited regimes. (b) The averages of radical sinks in the afternoon (12:00-18:00) for the base scenario (red bar) and the scenario considering missing VOC_R (blue bar) by increasing all measured NMHCs to fill the missing VOC_R. (c) Model-simulated dependence of daily peak O₃ concentrations on daily mean NO_x concentrations for the

删除了: (blue bar)

删除了: Model

base scenario (red curve) and the scenario considering missing VOC_R (blue curve) by increasing all measured NMHCs to fill the missing VOC_R . The dashed lines parallel to Y-axis represent the threshold of NO_X levels to distinguish between VOC-limited and NO_X -limited regimes. The shaded area represents standard deviation induced by the uncertainty in missing VOC_R .