

## **$\delta\text{O}_2/\text{N}_2$ datasets and gas loss effect**

### **a) Data validation**

Table S1 and Table S2 present information about the unpublished and published datasets respectively. Information on the datasets used in Figure 3 can be found in the main body of text in Table 3. Currently no accumulation rate and temperature data for FP and SIR and therefore not included in Figure 3.

**Table S1** Information on unpublished datasets used in this study. Core history and measurement information are presented to the best of our knowledge, along with the depth range for each dataset, and the initial form of the gases in the ice: bubble, clathrate, brittle zone (or bubble clathrate transition zone (BCTZ), or mixed. The laboratory at which the measurements were performed, and the method used are stated in the 'Lab' and 'Method' columns, where M-RF stands for melt refreeze method and ME for melt extraction method.

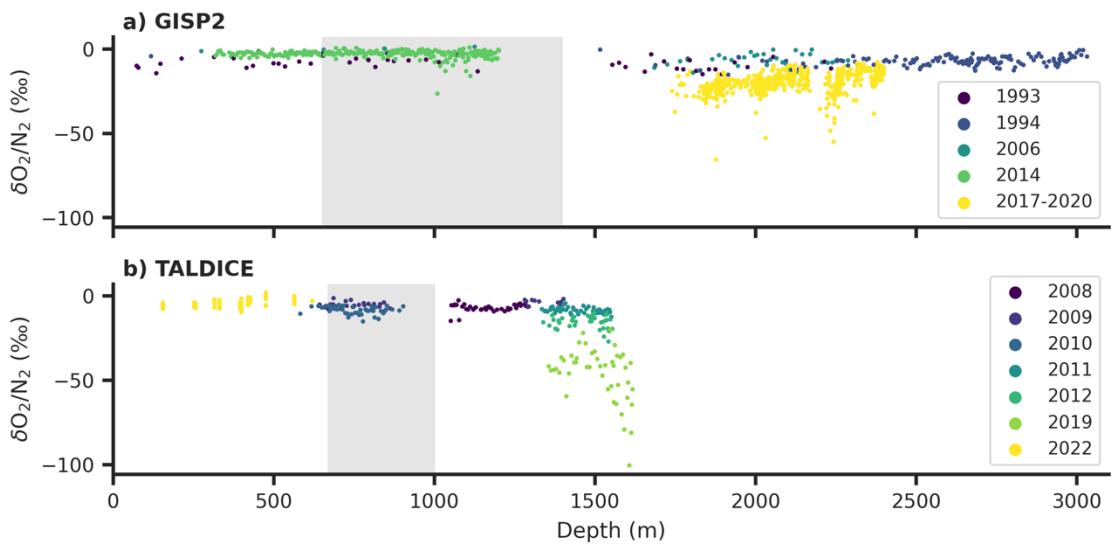
Site	Cored	Measured	Storage T	Lab	Method	Depth	Bubbly or clathrate
	(yr)	(yr)	(°C)			(m)	
<b>BI</b>	2003-2005	2010-2011	-22	LSCE	M-RF	608.85-693.55	BCTZ
<b>EDML</b>	2000-2006	2007	-25	LSCE	ME	596-860	Bubbly
<b>FP</b>	2011-2012	2015	-22	LSCE	ME	289.3-387.75	Bubbly
<b>GISP2</b>	1989-1993	2014	-35	Scripps	M-RF	318.67-1201.22	All
<b>GISP2</b>	1989-1993	2017-2020	-35	Scripps	M-RF	1740.27-2399.45	All
<b>JRI</b>	2008	2011	-22	LSCE	ME	52.25-363.55	Bubbly
<b>NEEM</b>	2008-2012	2011	-20	LSCE	M-RF	1757.25-2524.5	Clathrate
<b>SIR</b>	2018-2019	2021	-22	LSCE	ME	299.75-436.15	Bubbly
<b>TALDICE</b>	2005-2007	2008	-20	LSCE	ME	1003-1291	Clathrate
<b>TALDICE</b>	2005-2007	2009	-20	LSCE	ME	617.48-1401.92	All
<b>TALDICE</b>	2005-2007	2010	-20	LSCE	ME	582.48-902.93	Bubbly and BCTZ
<b>TALDICE</b>	2005-2007	2011	-20	LSCE	ME	1334.92-1550.92	Clathrate
<b>TALDICE</b>	2005-2007	2012	-20	LSCE	ME	1338.95-1553.95	Clathrate
<b>TALDICE</b>	2005-2007	2019	-20	LSCE	ME	1356-1617	Clathrate
<b>TALDICE</b>	2005-2007	2022	-20	LSCE	ME	155-620	Bubbly

**Table S2** Information on published datasets used in this study. Core storage history are presented to the best of our knowledge, along with the depth range for each dataset, the initial state of the gases in the ice, and the reference.

Site	Cored	Measured	Storage T	Depth	Bubbly or clathrate	Reference
	(yr)	(yr)	(°C)	(m)		
<b>DF</b>	1993-1996	2006	-50	1204.05-2500.36	All	Kawamura et al., 2007
<b>DF</b>	1993-1996	2017	-50	112.99-2001.06	All	Oyabu et al., 2022
<b>EDC</b>	1999-2004	2005-2008	-50	2483-3188 1300 – 1903, 2595	Clathrate	Landias et al., 2012
<b>EDC</b>	1999-2004	2012	-50	- 3260	Clathrate	Bazin et al., 2016
<b>EDC</b>	1999-2004	2016	-50	1904-2562 1489.95 - 1832.6, 1995.95 - 2350.15, 2555.85 - 2633.4, 2744.5 - 2797.85,	Clathrate	Extier et al., 2018
<b>EDC</b>	1999-2004	2021-2022	-50	2873.75 - 2910.6	Clathrate	Bouchet et al., 2023
<b>GISP2</b>	1989-1993	1993	-35	72.85-2235.05	All	Bender et al., 1994
<b>GISP2</b>	1989-1993	1994	-35	117.7-3032	All	Smith, 1998
<b>GISP2</b>	1989-1993	2006	-35	274.24-2287.88	All	Suwa and Bender, 2008
<b>DE08</b>	2008	-	-	175.15-218.2	Bubbly	Buizert et al., 2020
<b>DSSW20k</b>	1997	-	-	60.88-63.08	Bubbly	Buizert et al., 2020
<b>NGRIP</b>	1996-2004	-	-20	2464.09-2591.69	Clathrate	Bazin et al., 2016
<b>RICE</b>	2011-2013	2015	-36	116-146	Bubbly	Buizert et al., 2020
<b>RICE</b>	2011-2013	2017	-36	60.2-761	Bubbly	Lee et al., 2020
<b>SD</b>	1997-1999	-	-25	69.22-69.66	Bubbly	Buizert et al., 2020
<b>SD</b>	1997-1999	-	-25	74.95-972.98	All	Severinghaus et al., 2009
<b>SP</b>	2014-2016	-	-35	158-188.35	Bubbly	Buizert et al., 2020
<b>SP</b>	2014-2016	-	-35	125.3-1751.01	All	Severinghaus et al., 2019
<b>TALDICE</b>	2005-2007	2020	-20	1350-1600	All	Crotti et al., 2022
<b>VK</b>	1990-1998	2006	-35	174.5-3348	All	Suwa and Bender, 2008
<b>WAISD</b>	2006-2011	2011	-35	80.29-3396.61	All	Severinghaus et al., 2015
<b>WAISD</b>	2006-2011	2015	-35	119.35-190	Bubbly	Buizert et al., 2020

### b) GISP2 and TALDICE $\delta\text{O}_2/\text{N}_2$ data

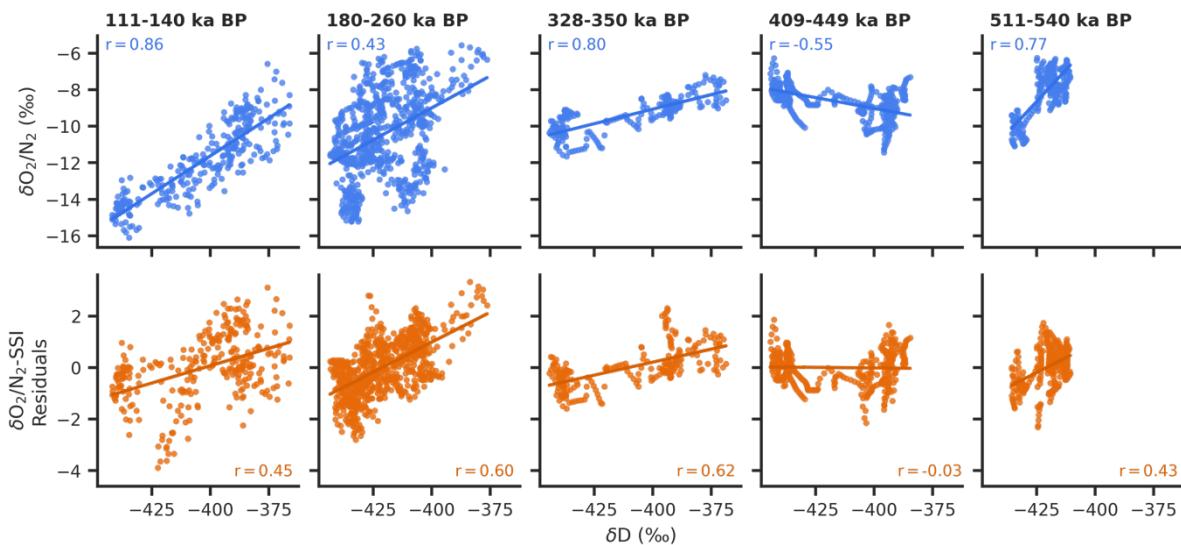
Successive measurements on clathrate ice shows  $\delta\text{O}_2/\text{N}_2$  depletion through time both from the GISP2 and TALDICE ice cores. A similar effect has been documented at EDC by Bouchet et al., 2023. Interestingly, the 2006 measurements on the GISP2 core clathrate ice (Suwa and Bender, 2008) are comparable to those made more than 10 years earlier (Bender et al., 1994; Smith, 1998). However, there is a depletion of up to 20 ‰ in measurements from between 2017 and 2020 (this study). Data from bubbly ice and within the reported brittle zone shows no gas loss effect.



**Figure S1**  $\delta\text{O}_2/\text{N}_2$  series from a) the GISP2 ice core and b) the TALDICE ice core after different storage conditions. The years of measurements are indicated and correspond to the different colours of the series. Grey shaded regions represented the reported brittle zones (Neff, 2014).

For the TALDICE ice core, all measurements have been performed on the melt extraction line and storage was only possible at -20°C. Figure S2 is displaying the series of  $\delta\text{O}_2/\text{N}_2$  measured after different periods of storage at -20°C. There is a clear decrease of the mean value of  $\delta\text{O}_2/\text{N}_2$  with storage length at -20°C from 2 years (measurements performed in 2008) to 13 years (measurements performed in 2019) in the clathrate ice. However, we find that the mean value of  $\delta\text{O}_2/\text{N}_2$  in the bubbly ice is still very high even after 16 years of storage at -20°C (mean value comparable to the mean  $\delta\text{O}_2/\text{N}_2$  value for samples measured in 2008) which suggests less or no gas loss for bubbly ice in the TALDICE ice core.

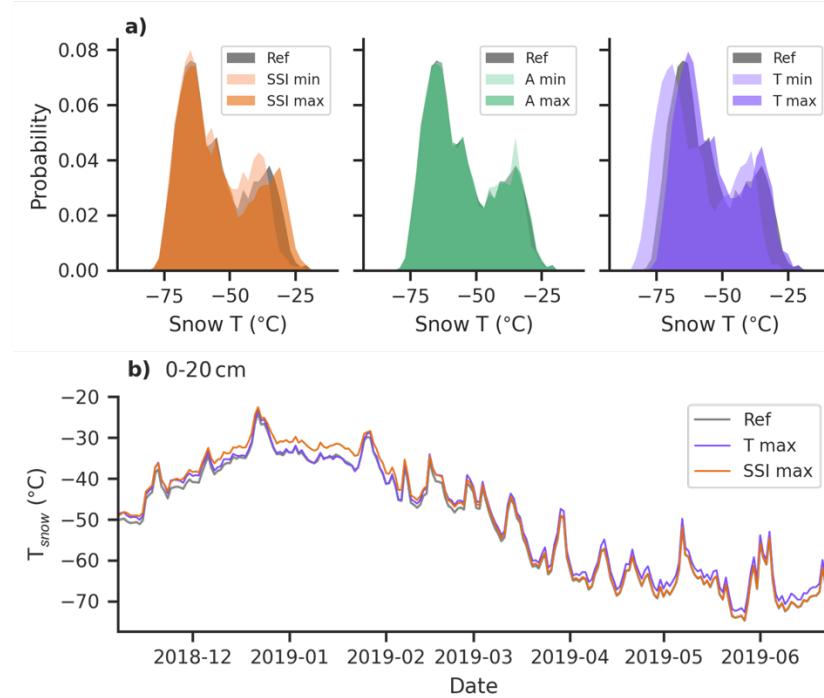
#### Temporal variability over different time slices from EDC



**Figure S3** Scatter plots showing each high-resolution section from the EDC core. The upper panel shows the linear regressions between  $\delta\text{O}_2/\text{N}_2$  and  $\delta\text{D}$ , and the lower shows the same for  $\delta\text{O}_2/\text{N}_2$  -SSI residuals and  $\delta\text{D}$ .

The following section shows that the strong correlation between  $\delta O_2/N_2$ -SSI residuals and  $\delta D$  during period between 190-260 ka BP are representative the entire core. We run the same analysis as presented in Section 3.1.2 but for the four other sections of high-resolution  $\delta O_2/N_2$  measurements from EDC (Bouchet et al., 2023). We observe significant positive correlations between  $\delta O_2/N_2$ -SSI residuals and  $\delta D$  (and  $\delta O_2/N_2$  and  $\delta D$ ) for all time periods, except for the period between 409-449 ka BP.

### Crocus model temperature evaluation



**Figure S4** Comparison of simulated snow temperatures in the upper 20 cm at Dome C. Panel a) shows the distributions for each set of simulation; SSI (orange), accumulation rate (green) and temperature (purple) the reference simulation (grey); compared to the reference simulation (grey). A comparison of mean 20 cm snow temperature from the reference (grey), increased temperature (T max), and increase SSI (SSI max) simulations are shown in panel b) between November 2018 and June 2019.

Figure S4 shows the simulated snow temperature distributions over the top 20 cm for the different forcing scenarios. Summer snow temperatures are most sensitive to perturbations in shortwave radiation forcings (SSI), but there is negligible effect in winter due to polar night conditions remaining constant between forcings. Perturbations in accumulation rate appear to have little influence on snow temperature. In contrast, snow temperature is strongly modified by perturbations in air temperature, specifically during winter and in the decreased temperature scenario (T min). Interestingly, the increased temperature scenario (T max) does not substantially influence summer snow temperatures. Figure S4 panel b), showing the daily mean snow temperatures over the top 20 cm, further highlights the non-linear effects of temperature increase where T max results in negligible change in summer temperatures and higher winter temperatures than the reference simulation. As discussed in the main body of text, this is expected to be due to the increased sensitivity of snow temperatures to insolation.

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