Gobal and regional chemical influence of sprites: Reconciling modeling results and measurements

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Abstract. Mesospheric electrical discharges, known as sprites, formed by fast-propagating streamers, have been shown to create localized enhancements of atmospheric constituents such as N, O, NO_x, N_2O, and HO_x, as indicated by both, modeling results and space-based measurements. In this study, we incorporate the occurrence rate of sprites into a chemistry-climate model using meteorological parameters as a proxy. Additionally, we introduce the injection of chemical species by sprites into the model, based on electrodynamical modeling of individual sprite streamers and observations from space.

Our modeling results show a good agreement between the simulated sprite distribution and observed data on a global scale. While the global influence of sprites on the atmospheric chemistry is found to be negligible, our findings reveal their measurable chemical influence at regional scale, particularly for the concentration of HNO_3 and HNO_4 within the mesosphere. The simulations also suggest that sprites could be responsible for the observed NO_2 anomalies at an altitude of 52 km above thunderstorms, as reported by MIPAS. Finally, a projected simulation reveals that the occurrence rate of sprites could increase at a rate of 14% per 1 K rise in the global temperature.

1 Introduction

In 1925, Wilson (1925) predicted the existence of electrical discharges (nowadays called sprites) above thunderstorms, which were later confirmed by Franz et al. (1990). Sprites are a category of Transient Luminous Events (TLEs) that take place at altitudes ranging from 40 km to 90 km. They are initiated by the ionization resulting from the quasi-electrostatic field component of lightning discharges, as described by studies such as, for instance, Pasko et al. (1997); Stenbaek-Nielsen et al. (2000); Pasko et al. (2012). The quasi-electrostatic field primarily generates electromagnetic radiation within the Extremely Low Frequency (ELF) and the ~~Very Low Frequency (VLF~~Ultra low frequency (ULF) electromagnetic spectra, which result from continuing currents during discharges lasting several tens or hundreds of milliseconds (Brook et al., 1962). Consequently, sprites are commonly observed simultaneously with a discernible ELF and/or ~~VLF~~ ULF signal, emitted by lightning flashes characterized by continuing currents (Greenberg et al., 2009; Inan et al., 2010; Lu et al., 2017).

Sprites consist of fast-propagating streamers followed by long-standing luminous structures known as beads and glows. These events typically last between 1 and 100 milliseconds (Liu, 2010; Luque and Gordillo-Vázquez, 2010, 2011; Luque et al.,

2016; Malagón-Romero et al., 2020). The primary sources of optical emissions in sprite streamers, beads, and glows originate from various molecular components, including the first and second positive systems of molecular neutral nitrogen, the first negative system of molecular nitrogen ions, the Meinel band of molecular nitrogen ions, and the Lyman-Birge-Hopfield (LBH) band of molecular neutral nitrogen (Armstrong et al., 1998; Chen et al., 2003; Kuo et al., 2005; Stenbaek-Nielsen et al., 2007; Kanmae et al., 2007; Šimek, 2014; Sato et al., 2015; Hoder et al., 2016; Ihaddadene and Celestin, 2017; Gordillo-Vázquez et al., 2018; Pérez-Invernón et al., 2018b).

Observing sprites is challenging due to their short duration. However, ground-, space- and aircraft-based instruments have been successful in detecting sprites, providing valuable information about their occurrence (Armstrong et al., 1998; Stenbaek-Nielsen et al., 2007; Gordillo-Vázquez et al., 2018; Arnone et al., 2020). To overcome the limitations of direct observations, some researchers have proposed using ELF and ~~VLF~~ ULF lightning measurements from flashes with continuing currents as a proxy indicator for the occurrence of sprites (Füllekrug and Constable, 2000; Sato and Fukunishi, 2003; Andrey et al., 2022). Following this approach, Ignaccolo et al. (2006) estimated a global occurrence rate of 2.8 sprites per minute with an accuracy factor of ∼2-3. Chen et al. (2008) used satellite-based optical observations from the Imager of Sprites and Upper Atmospheric Lightning (ISUAL) experiment aboard the FORMOSAT-2 satellite to report a global occurrence rate of 0.5 sprites per minute, while also providing information about the polarity of the lightning parents and the distribution of sprites over land and ocean (Zhang et al., 2022). More recently, Andrey et al. (2022) estimated a global occurrence rate of about 0.6 sprites per minute based on global measurements of the energy radiated by cloud-to-ground (CG) lightning reported by the World Wide Lightning Location Network (WWLLN).

Electrodynamical and chemical models of sprites suggest a significant local production of NO_x ($NO+NO_2$), N_2O and HO_x ($H+OH+HO_2$) in the mesosphere (from about 40 km upwards) and the lower-ionosphere (Sentman et al., 2008; Gordillo-Vázquez, 2008, 2010; Gordillo-Vázquez et al., 2012; Evtushenko et al., 2013; Winkler and Nothold, 2014; Parra-Rojas et al., 2015; Pérez-Invernón et al., 2020; Winkler et al., 2021). According to Sentman et al. (2008), they calculated a production of 5×10^{19} molecules of NO per single streamer in sprites between altitudes of 65 km and 75 km by using a chemical model of sprites. Enell et al. (2008) reported a production of 3×10^{22} to 3×10^{23} NO molecules per complete sprite. Pérez-Invernón et al. (2020) calculated a production of N_2O and NO molecules per sprite between 68 and 75 km altitude of 2×10^{19} and 10^{21} molecules, respectively. Winkler et al. (2021) modeled the production of HO_2 by sprite-streamers and found that they could produce about 10^{20} molecules of HO_2 between 70 and 80 km altitudes. Finally, Malagón-Romero et al. (2023) extended the electrodynamical model of sprite-streamers of Pérez-Invernón et al. (2020) to estimate a production of 7.5×10^{18} molecules NO, 2.6×10^{18} molecules of NO_2, 2.6×10^{18} molecules of N_2O and a removal of 3.1×10^{22} molecules of O_3 by sprite-streamers between 49.75 km and 50 km.

The possibility of sprites producing NO_x and HO_x in the mesosphere has motivated several attempts of measuring the chemical production of sprites to determine their chemical role in the atmosphere. Arnone et al. (2008) combined NO_2 measurements obtained from the Michelson Interferometer for Passive Atmospheric Sounding (MIPAS) with lightning data sourced from the World Wide Lightning Location Network (WWLLN) for the period August to December 2003. They conducted a search for anomalies in nighttime measurements of NO_2 mixing ratios (at about 22.00 local time) at altitudes of 47 km, 52 km, and 60 km

above thunderstorms. This search was limited to latitudes within the range of 30°S to 20°N and over a instantaneous field of view with a footprint of 30 km in longitude × 1200 km in latitude at 52 km. To examine the relationship between lightning activity and NO_2 mixing ratios, they generated five sets of NO_2 measurements based on the accumulation of lightning events prior to the MIPAS overpass. The first set consisted of measurements taken in the absence of lightning flashes up to 60 minutes before the MIPAS overpass, while the remaining sets were similar, but included data from 10, 20, and 30 minutes prior to the overpass. They reported a maximum positive anomaly of the NO_2 mixing ratio of + 1 ppbV 20 minutes after the occurrence of lightning at 52 km. Subsequently, Arnone and Dinelli (2016) extended their investigation up to April 2004, corroborating the presence of an elevated mixing ratio of NO_2 above thunderstorms. However, when the analysis was further expanded to encompass the entire MIPAS2D dataset (Dinelli et al., 2010), no significant augmentation in the NO_2 mixing ratio was discernible at an altitude of 52 km above thunderstorms. These results collectively suggest that the chemical disturbance induced by sprites in the mesosphere resides on the cusp of current detection capabilities. Following these measurements, Arnone et al. (2014) introduced a parameterization scheme into the Whole Atmosphere Community Climate Model (WACCM) to explore how sprites influence the chemistry of the mesosphere. They incorporated the injection of sprite-generated NO_x based on the latest findings from sprite-streamer modeling, simulating a global rate of 2-3 sprites per minute. Their study encompassed both, summer and winter conditions, involving simulations covering 40 days each. Their results revealed an elevation of 0.015-0.15 ppbv of the NO_x mixing ratio at 70 km altitude in tropical regions, although this effect became insignificant on a global scale. Furthermore, they identified a potential localized increase of up to tens of percent of the NO_x mixing ratio within the altitude range of 60 to 85 km. This increase, while potentially detectable by current instruments like MIPAS, remains a localized phenomenon.

Sentman et al. (2003) reported that sprites produce no distinctive OH emissions at the 2% background brightness level, indicating an upper estimate in the perturbation of OH by sprites. Recently, Yamada et al. (2020) documented a notable increase of the HO_2 mixing ratio in three regions following the incidence of sprites. They used limb spectral measurements reported by the Submillimeter-Wave Limb-Emission Sounder (SMILES) and estimated that a single sprite could produce up to 10^{25} molecules of HO_2 between 75 km and 80 km altitude, which is considerably larger than the production of 10^{20} molecules of HO_2 estimated by Winkler et al. (2021). An injection of 10^{25} molecules of HO_2 per sprite implies that sprites could represent up to the 1% of the global source of nighttime background HO_2 in the upper mesosphere. Nevertheless, there remains uncertainty regarding whether measurements of sprite chemical activity (Yamada et al., 2020) might be influenced, either partially or entirely, by the chemical production of lightning-induced electron precipitation in the mesosphere (Xu et al., 2021).

In this study, we present the first parameterization of sprites based on ~~proxy meteorological parameters~~ the proxy meteorological parameter vertical velocity at the 450 hPa level for sprite activity. We implement this parameterization in the Modular Earth Submodel System (MESSy) for usage within the European Center HAMburg general circulation model (ECHAM) / MESSy Atmospheric Chemistry (EMAC) (Jöckel et al., 2010, 2016). The parameterization of sprites is based on the parameterization of the occurrence rate of Long-Continuing-Current (LCC) lightning developed by Pérez-Invernón et al. (2022), enabling us to investigate the global seasonal variability in the occurrence of sprites as well as their sensitivity to climate change. In addition, we introduce in the parameterization the injection of NO_x, N_2O and HO_2 by sprites as well as the direct depletion

of O_3 between 45 and 80 km altitude by using the modeling results of Pérez-Invernón et al. (2020); Winkler et al. (2021) and Malagón-Romero et al. (2023). In turn, we compare the simulated NO_2 mixing ratio resulting from model simulations of sprites with the nighttime positive anomalies in the NO_2 mixing ratio reported by MIPAS above thunderstorms (Arnone et al., 2014; Arnone and Dinelli, 2016) to assess the potential influence of sprites on these measurements.

2 Model

2.1 The EMAC model

The numerical chemistry-climate model EMAC couples ECHAM5 (Roeckner et al., 2006) with the MESSy framework to connect various multi-institutional computer codes, referred to as MESSy submodels (Jöckel et al., 2010, 2016). The submodels are employed to depict processes within the troposphere and middle atmosphere, as well as their interactions with oceans, land, and external factors originating from anthropogenic emissions.

The model is operated ~~at T42L90MA resolution, i.e.~~ with a triangular ~~trunctaion~~ truncation of the spectral resolution at wave number 42, corresponding to a quadratic Gaussian grid with a resolution of 2.8° in both latitude and longitude. It comprises 90 vertical levels, extending up to the 0.01 hPa pressure level, and employs a time step length of 720 seconds, as described by Jöckel et al. (2016) for the T42L90MA resolution. Additionally, the Tiedtke/Nordeng convection scheme (Tiedtke, 1989; Nordeng, 1994) implemented within the CONVECT submodel is utilized.

The LNOX submodel of MESSy is used to calculate the occurrence rate of sprite-triggering LCC lightning flashes. The LNOX submodel calculates the total lightning flash frequency, the LCC lightning flash frequency, and the production of NO_x by lightning from several lightning parameterizations selected by the user (Tost et al., 2007) and by fixing a scaling factor that results in a lightning occurrence rate of ~45 flashes per second globally (Christian et al., 2003; Cecil et al., 2014). For the present study, we used the parameterization of lightning flashes producing the best comparison between the simulated and the observed LCC lightning flash density (Bitzer, 2017), i. e., the lightning parameterization based on the Cloud Top Height (CTH) by Price and Rind (1992) for land, combined with a parameterization of lightning that used the updraft strength at 440 hPa pressure level (Allen and Pickering, 2002) for ocean, with a global scaling factor of 1.13 (Pérez-Invernón et al., 2022). The submodel LNOX calculates the occurrence rate of LCC flashes with a continuing current longer than 18 ms (LCC(>18 ms) lightning) from the updraught mass flux by employing the parameterization for the ratio of LCC to total lightning, as developed by Pérez-Invernón et al. (2022).

2.2 Parameterization of sprites

A new submodel named SPRITES is developed to include the parameterization of sprites in MESSy v2.54 and will be implemented in the submodel LNOX in future versions of MESSy. The submodel SPRITES calculates the sprite ~~flash~~ density and the production of NO, NO_2, N_2O and HO_2 and, in turn, the depletion of O_3 by sprites between 45 km and 75 km altitude.

2.2.1 Sprite occurrence

The occurrence rate of sprites in the submodel SPRITES is implemented by using the calculation of lightning density from the LNOX submodel. Sprites are generated by the charge moment change resulting from lightning, along with the duration over which the charge moment is attained (Cummer and Füllekrug, 2001). Lightning flashes with continuing currents, such as LCC(>18 ms) lightning, can ~~produce a significant~~ induce a substantial charge moment change and a high quasi-electrostatic field in the mesosphere (Gamerota et al., 2011)~~and, in turn, trigger the inception of sprites~~. Consequently, this process can trigger the initiation of sprites e.g., ~~(Pasko et al., 1997; Stenbaek-Nielsen et al., 2000; Pasko et al., 2012)~~Pasko et al. (1997); Stenbaek-Niel]. Therefore, we use the LCC(>18 ms) lightning density, computed by the LNOX submodel of MESSy (Tost et al., 2007; Pérez-Invernón et al., 2022) ~~is used~~ based on the vertical velocity at the 450 hPa pressure level, as a proxy for the occurrence of sprites. In addition, we imposed that nighttime sprites can only be produced after sunset, ~~as the~~ when the Sun is below the horizon. The absence of solar radiation during nighttime contributes to reduce the ionization of the lower ionosphere and, in turn, favors the electric breakdown of the air that triggers the inception of sprites (Pérez-Invernón et al., 2016). Finally, we imposed that only 20% of the nighttime LCC(>18 ms) lightning flashes have the potential to trigger sprites, following the ELF measurements of lightning with continuing current reported by Füllekrug and Constable (2000). Arnone et al. (2014) estimated that about 1/1,000 flashes could produce a sprite, while Pérez-Invernón et al. (2022) found that 7.4/1,000 flashes reported by LIS during one year have a continuing current lasting more than 18 ms. Therefore, the approximation of 1 sprites per 20% nighttime LCC(>18 ms) lightning flash is of the same order as the 1/1,000 sprite-to-flash estimate by Arnone et al. (2014). The assumed 20% is an upper limit, as Füllekrug and Constable (2000) reported that between 5 and 20% of the measured lightning flashes had the potential to produce air electric breakdown at sprite altitude. However, our approach lacks consideration for sprites triggered by lightning without continuing currents, which may lead to an underestimation of sprite occurrences (Inan et al., 2010). ~~Finally, it~~ Sprites are generated by the quasi-static removal of a relatively large lateral charge distribution where lightning continuing current is perhaps the most prominent indicator, but not necessarily the only mechanism that contributes. It is worth noting that Greenberg et al. (2009) found that approximately 33% of the 15 sprites analyzed in Europe were produced by lightning strikes unaccompanied by associated ELF transients. Similarly, Lu et al. (2017) reported that 15% of the 247 recorded sprites in North America were the result of negative cloud-to-ground flashes without detectable continuing currents.

2.2.2 Chemical influence of sprites

The submodel SPRITES introduces the chemical influence of sprites in the mesosphere by multiplying the calculated sprite density and the production/destruction of chemical species by single sprites. The HO_2 molecules produced by sprites are homogeneously distributed between 70 km and 75 km altitude. The submodel's namelist allows the user to choose a total injection of 10^{20} or 10^{25} molecules of HO_2 per sprite, based on ~~electrodynamic~~ modeling results (Winkler et al., 2021) and measurements (Yamada et al., 2020), respectively.

The injection of NO, NO_2 and N_2O molecules, as well as the direct depletion of O_3 molecules implemented in the SPRITE submodel are based on modeling results of single sprite streamers in the mesosphere-lower thermosphere (67-75 km) and

in the lower mesosphere (49.75-50 km) by Pérez-Invernón et al. (2020) and Malagón-Romero et al. (2023), respectively. The production of chemical species by sprites between the ranges of altitudes from 45 km to 49.75 km and from to 50 km and 67 km is estimated by following the approach developed by Malagón-Romero et al. (2023), i. e., by interpolating and extrapolating from the results by Pérez-Invernón et al. (2020) and Malagón-Romero et al. (2023). In particular, we estimate the production per meter of NO, NO_2, and N_2O molecules, as well as the depletion per meter of O_3 molecules, in the altitude range from 49.75 km to 50 km and from 67 km to 75 km. Subsequently, we extrapolate the production or removal of molecules from 45 km to 49.75 km and interpolate the production or removal of molecules from 50 km to 75 km. Following this approach, Malagón-Romero et al. (2023) obtained 6.2×10^{20} NO molecules, 2.6×10^{19} NO_2 molecules and 1.7×10^{20} N_2O molecules injected by a single sprite streamer in the range of altitudes between 45 km and 75 km, while they reported a removal of 3.1×10^{23} O_3 molecules. In addition, we apply a conversion factor between the chemical injection by a single sprite streamer and by a complete sprite in EMAC. While ELF radio measurements suggest the presence of over 1,000 streamers per individual sprite (Qin et al., 2012), it is important to recognize that the characteristics and production of chemical species within streamers can be heterogeneous (Stenbaek-Nielsen et al., 2013). Consequently, multiplying the injection of chemical species per streamer by the total number of streamers may lead to inaccuracies. To address this, we conduct a comparison between the simulated and observed total number of photons, allowing us to estimate the production of chemical species by observed sprites based on simulation results. Pérez-Invernón et al. (2020) reported a scaling factor ranging between 18 and ~~50.~~ 50 based on observed sprites. We have updated the estimation of this scaling factor by using recent detections of sprites by the Atmosphere-Space Interactions Monitor (ASIM). Gomez Kuri (2021) reported the detection of a sprite on July 10, 2019 by combining optical measurements of ASIM and ELF measurements from ground-based sensors. We have integrated the optical signal detected by ASIM in the range of wavelengths between 180 nm and 230 nm (ASIM photometer 2) during the 0.85 ms after the onset of the first and the second peaks associated with the sprite event at times 21:53:17.554 and 21:53:17.563 (Gomez Kuri, 2021, Figure 4.16), obtaining observed photometric fluxes of 10^{-10} J m^{-2} and 1.1×10^{-10} J m^{-2}, respectively. The synthetic flux of the streamer simulated by Pérez-Invernón et al. (2020) that would be observed by ASIM during 0.85 ms in the range of wavelengths between 180 nm and 230 nm is 10^{-12} J m^{-2}. Therefore, we can assume that a complete sprite emits about 100 times more photons than the simulated single sprite streamer. Therefore, we used a factor of 100 to convert the simulated injection of molecules by a single sprite streamer into the injection of molecules by a complete sprite.

2.3 Simulation set-up

Table 1 shows the overview of the performed simulations. Firstly, a purely dynamical simulation (SPRI) covering the present day climatic state is performed during the period 2000-2009 by nudging the model towards ERA-Interim reanalysis meteorological fields (Dee et al., 2011) to evaluate the sprite frequency parameterization. A projection simulation (RCP6.0) covering the years 2090-2095 is performed under Representative Concentration Pathway 6.0 (RCP6.0) in order to evaluate the sensitivity of sprites under climate change. We consider the years 2000 and 2090 as the spin-up phases. The RCP6.0 simulation is established following the simulation RC2-base-04 of Jöckel et al. (2016); Pérez-Invernón et al. (2023). The sea surface temperatures (SSTs) and the sea-ice concentrations (SICs) are prescribed from simulations with the Hadley Centre Global Environment

Table 1. Overview of the performed simulations.

Simulation	Mode	Years	Sprites	Sprites-chemistry
SPRI	Dynamical. Nudged towards ERA-Interim reanalysis	2000-2009	Yes	No
RCP6.0	Active chemistry. Projection RCP6.0	2090-2095	Yes	No
BASE	Active chemistry. Nudged towards ERA-Interim reanalysis	2000-2001	No	No
CTRL	Active chemistry. QCTM from BASE	2000-2001	No	No
SPRI-M	Active chemistry. QCTM from BASE	2000-2001	Yes	Yes (HO_x by Winkler et al. (2021))
SPRI-SMI	Active chemistry. QCTM from BASE	2000-2001	Yes	Yes (HO_x by Yamada et al. (2020))

Model version 2 - Earth System (HadGEM2-ES) Model (Collins et al., 2011; Bellouin et al., 2011). Projected greenhouse gases and SF_6 mixing ratios are included from Eyring et al. (2013). Anthropogenic emissions are taken from monthly values provided by Fujino et al. (2006) for the RCP6.0 scenario. The chemical influence of sprites in the atmosphere has been deactivated in this simulation in order to avoid unexpected perturbations in the chemistry. We refer to Jöckel et al. (2016) for more details about the simulations set-up. Pérez-Invernón et al. (2023) obtained a temperature increase of 4 K between present day and 2091-2095 by using the same setup.

In turn, a set of simulations with active chemistry are performed to evaluate the chemical role of sprites in the atmosphere in the Quasi Chemistry-Transport Model (QCTM) mode proposed by Deckert et al. (2011) to ensure that small chemical perturbations do not alter the simulated meteorology by introducing noise. Firstly, a two years simulation nudged towards ERA-Interim reanalysis meteorological fields (Dee et al., 2011; ECMWF, 2011) and without sprites is performed (BASE). The simulation is set-up following the simulation with interactive chemistry RC1SD-base-07 of Jöckel et al. (2016). The sea surface temperatures (SSTs) and the sea-ice concentrations (SICs) are used from ERA-Interim reanalysis data (Dee et al., 2011). The chemical kinetics is simulated by using the submodel MECCA (Module Efficiently Calculating the Chemistry of the Atmosphere) by Sander et al. (2019). Next, we conduct a present-time simulation in the QCTM mode, featuring active chemistry but excluding sprites (CTRL). This simulation is generated utilizing inputs for radiation calculations and methane oxidation from the BASE simulation. This approach allows us to separate dynamics from chemistry while maintaining consistent meteorological conditions, effectively suppressing meteorological variability. Lastly, we perform two additional present-day simulations, both in the QCTM mode and incorporating sprites. These simulations are denoted as SPRI-M and SPRI-SMI, respectively. In the SPRI-M simulation we have used modeling results of single sprite streamers (Pérez-Invernón et al., 2020; Malagón-Romero et al., 2023) to set the injection chemical species, while in the simulation SPRI-SMI we have modified the injection of HO_x accordingly to measurements from SMILES (Yamada et al., 2020) (see Section 2.2.2). Comparison of simulations CTRL, SPRI-M and SPRI-SMI allow us to determine the chemical role of sprites in the mesosphere. In this set of simulations, we consider the year 2000 as the spin-up phase.

3 Results and discussion

3.1 Geographical distribution of sprites

Details on the simulated global frequency of sprites and lightning are summarized in Table 2, while the simulated annually averaged sprite ~~flash~~ densities for present day and 2091-2095 are shown in Figure 1. We obtained a global sprite occurrence rate of 1.66 sprites per minute in 2000-2009, which is above the value reported by Chen et al. (2008) (0.5 sprites per minute) and below the value reported by Ignaccolo et al. (2006), ranging between 2 and 3 sprites per minute. In turn, we obtained a global occurrence rate of 2.55 sprites per minute in 2091-2095, which represents an increase of 54% (14% increase per 1 K increase in the global temperature between the period 2091-2095 and 2000-2009). The simulated increase of the global occurrence rate of sprites between present day and 2091-2095 is approximately similar to the simulated increase of the global occurrence rate of total lightning (47%).

The simulated latitudinal distribution of sprites has a pronounced peak between 10°N and 20°N and another less pronounced peak between 30°N and 40°N (see Figure 2(a)), in agreement with the climatology reported by (Zhang et al., 2022, Figure 5) and (Lu et al., 2022, Figure 2(b)) from ISUAL measurements. However, the simulated latitudinal distribution of sprites presents a main peak between 10°S and 20°S that was not detected by ISUAL. The disagreement between the observed and the simulated latitudinal distribution of sprites is influenced by the overestimation of simulated sprites over land between 10°S and 20°S (Figure 2(b)), i.e. in South America and Southeastern Asia. In turn, we found a good agreement between the simulated and the observed (Zhang et al., 2022; Lu et al., 2022) latitudinal distribution of sprites over ocean (Figure 2(c)), except for the observed peak near 40°N that is less pronounced in the simulations. We found a sprite land-ocean contrast of 0.9 : 1.1 and 1 : 1 for present day and 2091-2095, respectively. Chen et al. (2008) reported a sprite land-coast-ocean contrast of 4.7 : 3.2 : 1. The horizontal resolution of our simulations (2.8° × 2.8° quadratic Gaussian grid in latitude and longitude) does not allow us to distinguish between coast and land and coast and ocean. The parameterization of Andrey et al. (2022) produced 41.4% sprites over land and 58.6% over ocean, while we obtained 45% and 55%, respectively.

The simulated global density of sprites over land in present day simulations is in fairly good in agreement with the observations reported by Chen et al. (2008); Zhang et al. (2022), showing hotspots in Middle Africa, South America, Eastern North America, the Tornado Alley of North America, Western Europe and Southeastern Asia. The simulation produced an overestimation of the sprite density in Brazil, Southern Africa and China that can be explained by the low accumulative observation time of ISUAL in these regions (Chen et al., 2008). The high occurrence of sprites in the Mediterranean Sea and Western Europe is in agreement with the European climatology of sprites reported by Yair et al. (2015) and Arnone et al. (2020). The obtained sprite occurrence in Russia is in agreement with the sprite density derived from WWLLN data by Evtushenko et al. (2024), with the highest occurrence of sprites in the south.

Figure 3 shows the averaged sprite ~~flash~~ density during the seasons December-January-February (DJF), March-April-May (MAM), June-July-August (JJA) and September-October-November (SON) during 2000-2009. The maximum occurrence of sprites is reached during summer (DJF in the Southern Hemisphere, JJA in the Northern Hemisphere). During MAM the global density of sprites is shifted towards the Equatorial region and the Southern Hemisphere, while it is equally distributed

Table 2. Occurrence rate of lightning and sprites from the simulations SPRI and RCP6.0.

	2001-2009	2091-2095
Global occurrence rate of lightning flashes	45 flashes s^{-1}	66 flashes s^{-1}
Global occurrence rate of sprites	1.66 sprites min^{-1}	2.55 sprites min^{-1}
Lightning land-ocean contrast	1 : 1	1 : 1
Sprites land-ocean contrast	0.9 : 1.1	1 : 1

between both hemispheres in SON. ~~Sato and Fukunishi (2003)~~ (Sato and Fukunishi, 2003, Fig. 4) reported the global seasonal distribution of sprites based on ELF measurements for DJF, JJA and SON. There is a good agreement between the simulated

250 (Figure 3) and the reported ~~sprite distribution~~ (Sato and Fukunishi, 2003, Fig. 4(a)) spatial maximum of sprite density in JJA near the Equator and in Southeastern Asia, while the present parameterization underestimates the production of sprites in North and South America. In DJF, the simulation is in agreement with the climatology of sprites based on ELF measurements (Sato and Fukunishi, 2003) in all regions except for North America. Finally, we have found a good agreement between the simulations and the ELF based climatology during SON.

255 ### 3.2 Global chemical influence of sprite-streamers in the mesosphere-lower thermosphere

Figure 4 shows the simulated influence of sprites in the annually and globally averaged vertical profiles of NO_x=NO+NO$_2$, N$_2$O, HO_x=~~H+~~OH+HO$_2$, O$_3$, HNO$_3$ and HNO$_4$ by assuming that single sprites inject 10^{20} HO$_2$ molecules (Winkler et al., 2021). The obtained small variations between simulations with and without sprites clearly show that the influence of sprites is negligible on the global scale. The maximum effect of sprites in the vertical profiles of NO_x and N$_2$O are located in the

260 upper mesosphere, where the background abundance of these species is low. The found contribution of sprites to the global amount of N$_2$O in the upper mesosphere is about ~~0.002~~0.008%, in agreement with previous estimates by Pérez-Invernón et al. (2020) of 0.003% from electrodynamical simulations of streamers. ~~The background HO~~We obtained a marginal increase of approximately 0.007% in the background concentration of NO_x ~~mixing ratio in the upper mesosphere is slightly enhanced in the presence of sprites and it is dominated by the conversion of HO$_2$ into OH and H~~at an altitude of 70 km. This increment

265 is notably lower than the perturbation estimated by Arnone et al. (2014) due to sprites, which falls within the range of 2% to 20%. The variance in results can be attributed to the disparity in assumptions made by Arnone et al. (2014), who considered an injection of NO_x molecules ranging from 1.5×10^{23} to 1.5×10^{24}. In contrast, our study assumes a more conservative injection of 6.46×10^{22} NOx molecules. In addition, the sprite-NOx perturbation profile in this study is linear between the altitudes 45 km and 80 km, while the profile adopted by Arnone et al. (2014) peaks at about 65 km altitude. The amount of HO_x

270 is reduced in the ~~middle and lower~~ mesosphere as a consequence of the conversion of HO$_2$ and OH into the nitrogen reactive compounds HNO$_3$ and HNO$_4$ produced by the injection of NO_x. As a consequence, the mixing ratio of HNO$_3$ and HNO$_4$ increased in the upper mesosphere. In turn, the injection of NO and HO$_2$ by sprites produces an enhancement in the upper-
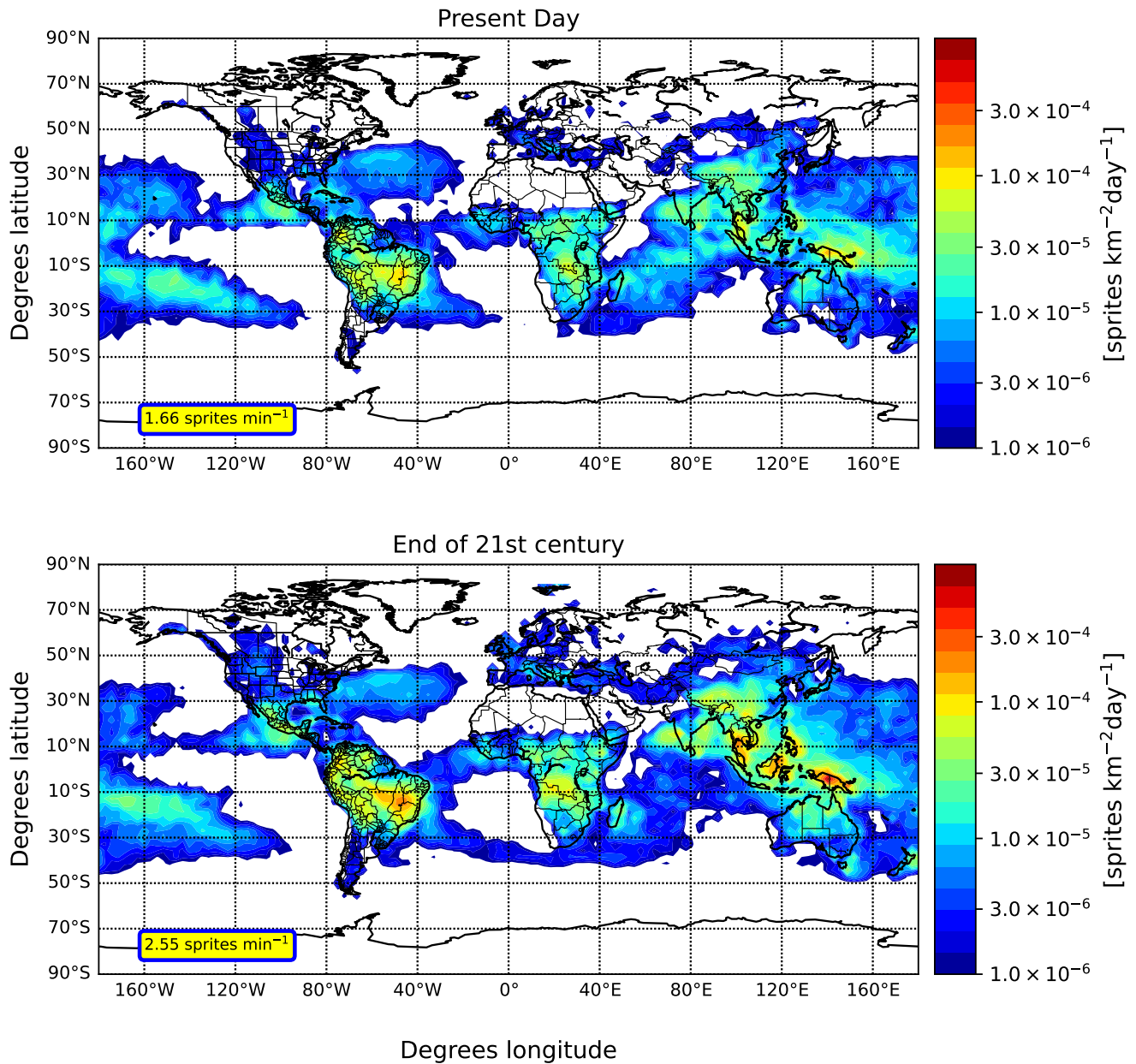
Figure 1. Simulated annually averaged sprite ~~flash~~ density in sprites per squared kilometer and day during 2001-2009 from the SPRI simulation (a) and during 2091-2095 from the RCP6.0 simulation (b). We annotate in boxes the annually averaged occurrence rate of sprites per minute.
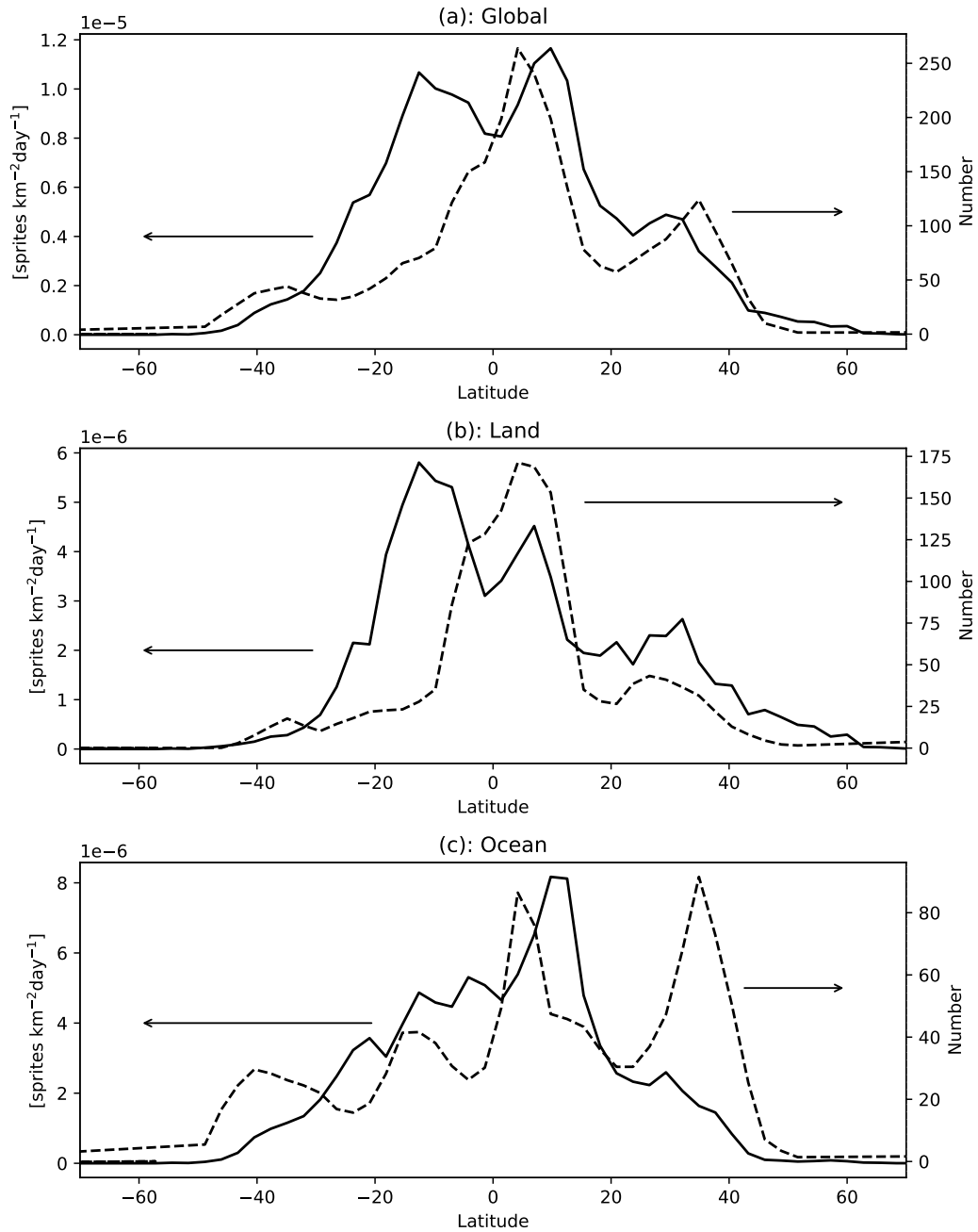
Figure 2. The solid lines depict the simulated annual average latitudinal sprite ~~flash~~ density, measured in sprites per square kilometer per day, spanning the years 2001 to 2009 (SPRI simulation), encompassing both, land and ocean regions. The dashed lines represent the total number of pure sprites and sprites accompanied by halos observed by ISUAL, as illustrated by (Lu et al., 2022, Figure 2(b)). The observations have been interpolated to the corresponding latitudes of the simulation results.
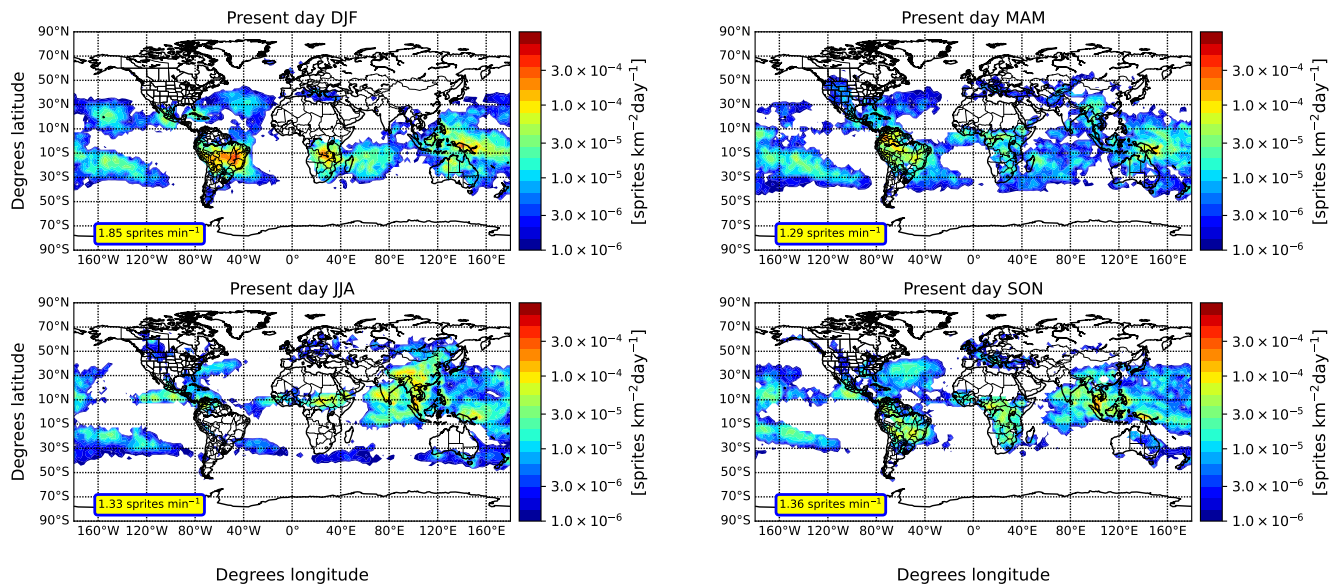
Figure 3. Simulated seasonally averaged sprite ~~flash~~ density in sprites per squared kilometer and day during 2001-2009 from the SPRI simulation. We annotate in boxes the seasonally averaged occurrence rate of sprites per minute.

mesospheric background O_3 mixing ratio, while the net contribution of sprites to O_3 in the middle and the lower mesosphere is negative.

275   We show in Figure 5 the influence of sprites on the annually and globally averaged vertical profiles of NO_x=NO+NO_2, N_2O, HO_x=~~H~~+OH+HO_2 and O_3 by assuming that single sprites inject 10^{25} HO_2 molecules (Yamada et al., 2020). There are some relevant differences between the enhancements of NO_x when introducing 10^{25} HO_2 molecules instead of 10^{20} HO_2 molecules per sprite. The injected HO_2 produces a modification of the contribution of NO and NO_2 to the total NO_x. The HO_2 reacts with NO, producing a slight decrease in the concentration of NO in the upper mesosphere and an increase of about ~~0.06~~0.07%

280   in NO_2. The injection of 10^{25} HO_2 molecules per sprite produces a ~~large HO_x mixing ratio~~ 0.01% increment in the mixing ratio of HO_x between 60 km and 80 km altitude, still too low to be considered a significant source of HO_x at a global scale. However, the conversion of OH and HO_2 into the reactive nitrogen compounds HNO_3 and HNO_4 in combination with NO_x led to a 0.3% and a 0.04% enhancement in the mixing ratio of HNO_4 and HNO_3 in the upper mesosphere, respectively. In turn, the influence of sprites on O_3 is different, when introducing 10^{25} HO_2 molecules per sprite. The injected HO_2 contributes to a

285   decrease of the background mixing ratio of O_3 in the upper mesosphere (about -0.002% at a global scale). The depletion of O_3 by HO_x can be due to the enhancement of NO_2, which contributes to deplete the mixing ratio of O_3.

We further analyze the geographical influence of sprites in the chemistry of the mesosphere. We show in Figure 6 the annual global difference of the mixing ratios of NO_x, N_2O, HO_x and O_3 at 72 km and 50 km altitude between two simulations with and without sprites. In this case, we have assumed that a single sprite injects 10^{20} HO_2 molecules (Winkler et al., 2021). The
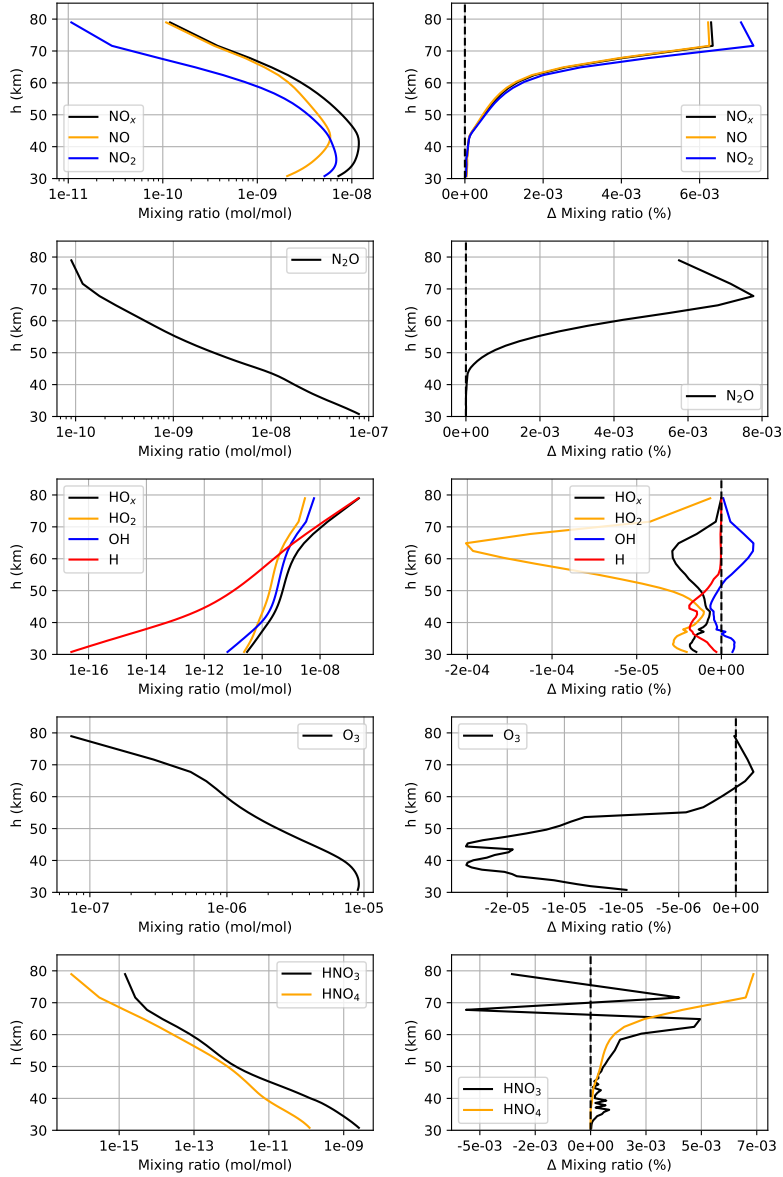
Figure 4. First column: Annually (2001) and globally averaged vertical profiles of the mixing ratio of NO_x, N_2O, HO_x, O_3, HNO_3 and HO_4 for a simulation without sprites (CTRL). Second column: Differences (in %) between the annually and globally averaged mixing ratio of the chemical species between the simulation with sprites (SPRI-M) and without sprites (CTRL). In the simulation with sprites, we have assumed that a single sprite can inject 10^{20} HO_2 molecules (Winkler et al., 2021).
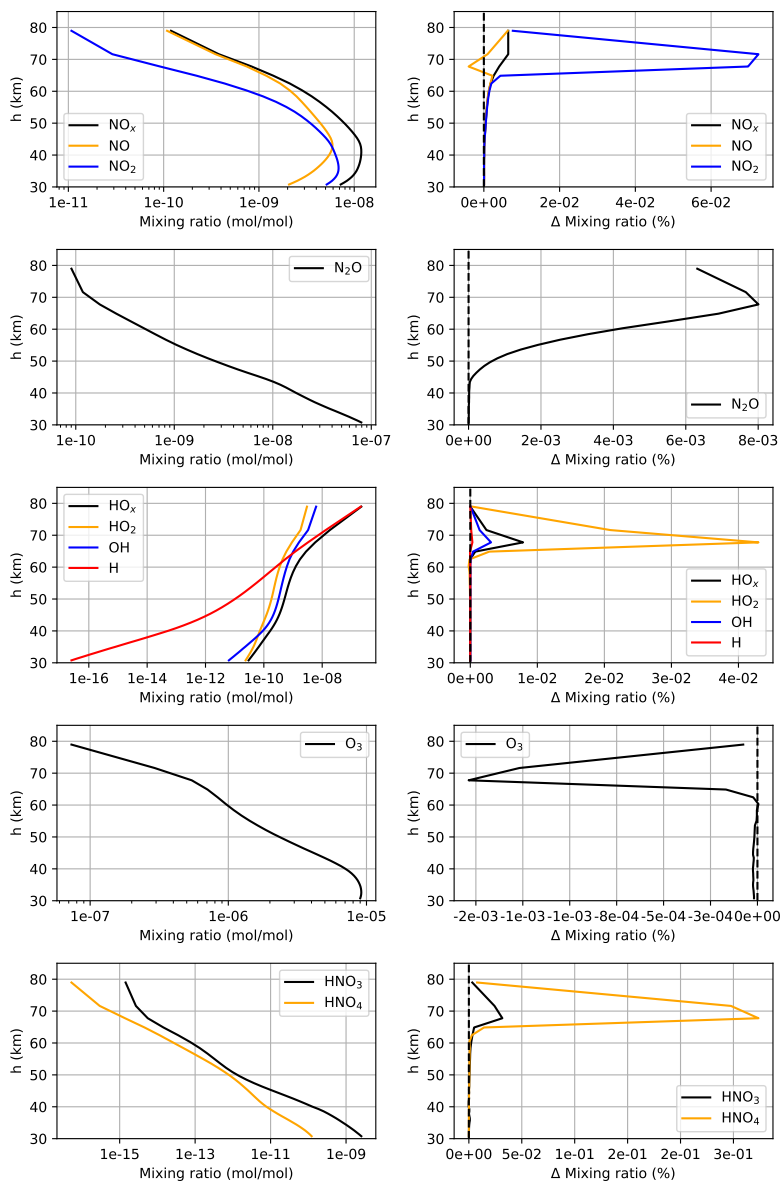
Figure 5. Same as Figure 4 but assuming that a single sprite injects 10^{25} HO_2 molecules (SPRI-SMI simulation) as reported by Yamada et al. (2020).

290   maximum increases ~~of the~~ in NO_x and N_2O mixing ratios at both altitudes are ~~located~~ observed in the tropical and middle latitudes~~, where the area with~~. This region coincides with the area that experiences the largest annual occurrence of sprites.. The chemical influence of sprites in the geographical distributions of HO_x and O_x are more complex. The mixing ratios of HO_x and O_3 decrease in the areas with a high occurrence rate of sprites at 72 km and 50 km altitude. The HO_x is depleted by the injected molecules of NO_x, while O_3 is directly depleted by sprites as prescribed by the developed parameterization.

295 The mixing ratio of O_3 increases near the poles at 72 km altitude, ~~while its tendency at polar latitudes is not homogeneous at 50 km. At 72 km altitude,~~ where the implemented parameterization of sprites imposes an injection of NO without NO_2. The NO injected at 72 km altitude at tropical and middle latitudes is transported polewards and produces O_3 in the presence of N. ~~The excess of O_3 at 72 km altitude at polar latitudes contributes to a depletion of HO_x molecules. The same relationship between O_3 and HO_x can be seen at 50 km altitude, where the inhomogeneous variations of the mixing ratios of HO_x and O_3~~

300 ~~are spatially correlated.~~

    Figure 7 shows the annual global difference between the mixing ratios of the analyzed chemical species assuming that a single sprite injects 10^{25} HO_2 molecules (Yamada et al., 2020) instead of 10^{20} HO_2 molecules (Winkler et al., 2021). The larger injection of HO_2 molecules does not produce any significant difference in the variation of NO_x and N_2O between the simulations with and without sprites. However, clear differences of the impact of sprites for the background mixing ratios of

305 HO_x and O_3 can be seen (see values and geographical distribution). There is a significant enhancement of the HO_x mixing ratio at 72 km in the regions with the largest occurrence of sprites, due to the direct injection of HO_2. The increase of the mixing ratio of HO_x produces a decrease of the O_3 mixing ratio, as O_3 molecules are depleted in the conversion between H, OH and HO_2 molecules, such as $OH + O_3 \rightarrow HO_2 + O_2$ and $HO_2 + O_3 \rightarrow OH + 2O_2$ (Jaeglé et al., 2001; Sander et al., 2011). At 50 km altitude, far from the vertical level where the HO_2 is injected (above 70 km), the variations of the HO_x and O_3 mixing

310 ratios are nearly similar as in the previous case (Figure 6).

    Finally, we show in Figure 8 the annual global difference of the mixing ratio of HNO_3 and HNO_4 at 72 km altitude between a simulation with sprites (assuming that a single sprite injects 10^{25} HO_2 molecules (Yamada et al., 2020)) and a simulation without sprites. There is a clear enhancement of the HNO_4 mixing ratio in regions with a large occurrence of sprites produced by the reaction between the injected NO_x and HO_x molecules, mainly $NO_2 + OH \rightarrow HNO_3$, $NO_2 + OH + M \rightarrow HNO_3 +$

315 M and $NO_2 + HO_2 \rightarrow HNO_4$ (Sander et al., 2011). According to Figure 8, the relative increase of the global HNO_4 mixing ratio (0.3%, see Figure 5) is significantly concentrated at tropical latitudes. In particular, sprites potentially represent a non-negligible source of upper-mesospheric HNO_4 at regional scale in South America and Southeastern Asia. In specific regions, the decrease of the HNO_3 mixing ratio is likely attributed to the elevated mixing ratio of OH, which interacts with HNO_3, leading to its depletion (Sander et al., 2011).

320 ### 3.3   Regional chemical influence of sprite-streamers: Comparison with NO_2 measurements by MIPAS

We now compare the simulated and the observed anomalies of the NO_2 mixing ratio at nighttime at the same local hour and time scale as reported by ~~Arnone et al. (2014)~~Arnone et al. (2008). They reported NO_2 anomalies over a instantaneous field of view footprint of 30 km $\times$ 500 km, an horizontal area that is about 6 times smaller than the area covered by each cell domain
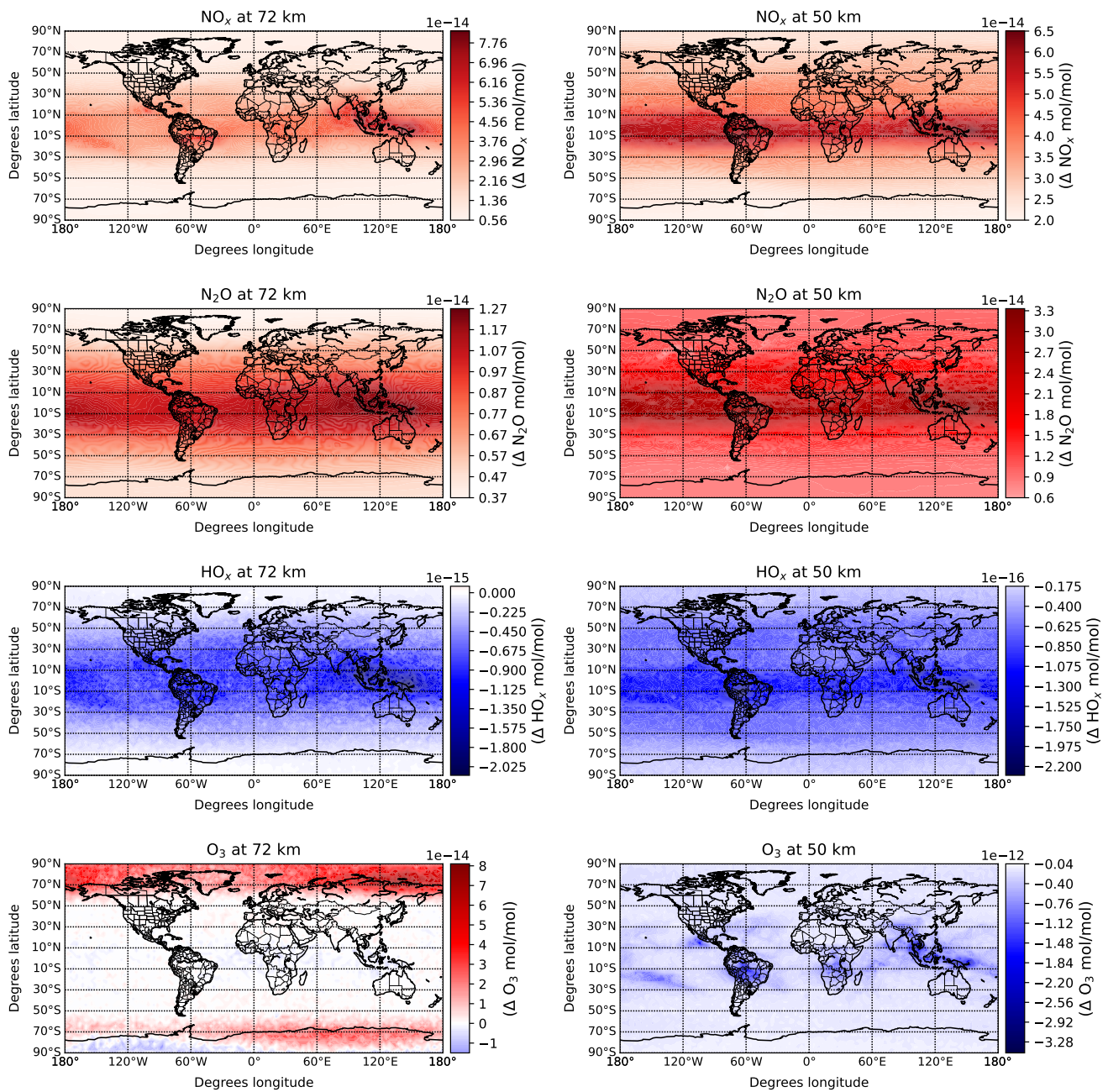
Figure 6. Annually (2001) ~~and globally~~ averaged differences of the NO_x, N_2O, HO_x and O_3 mixing ratios between a simulation with sprites (SPRI-M) and without sprites (CTRL) at 72 km (left) and 50 km (right) altitude. In the simulation with sprites, we have assumed that a single sprite can inject 10^{20} HO_2 molecules (Winkler et al., 2021).
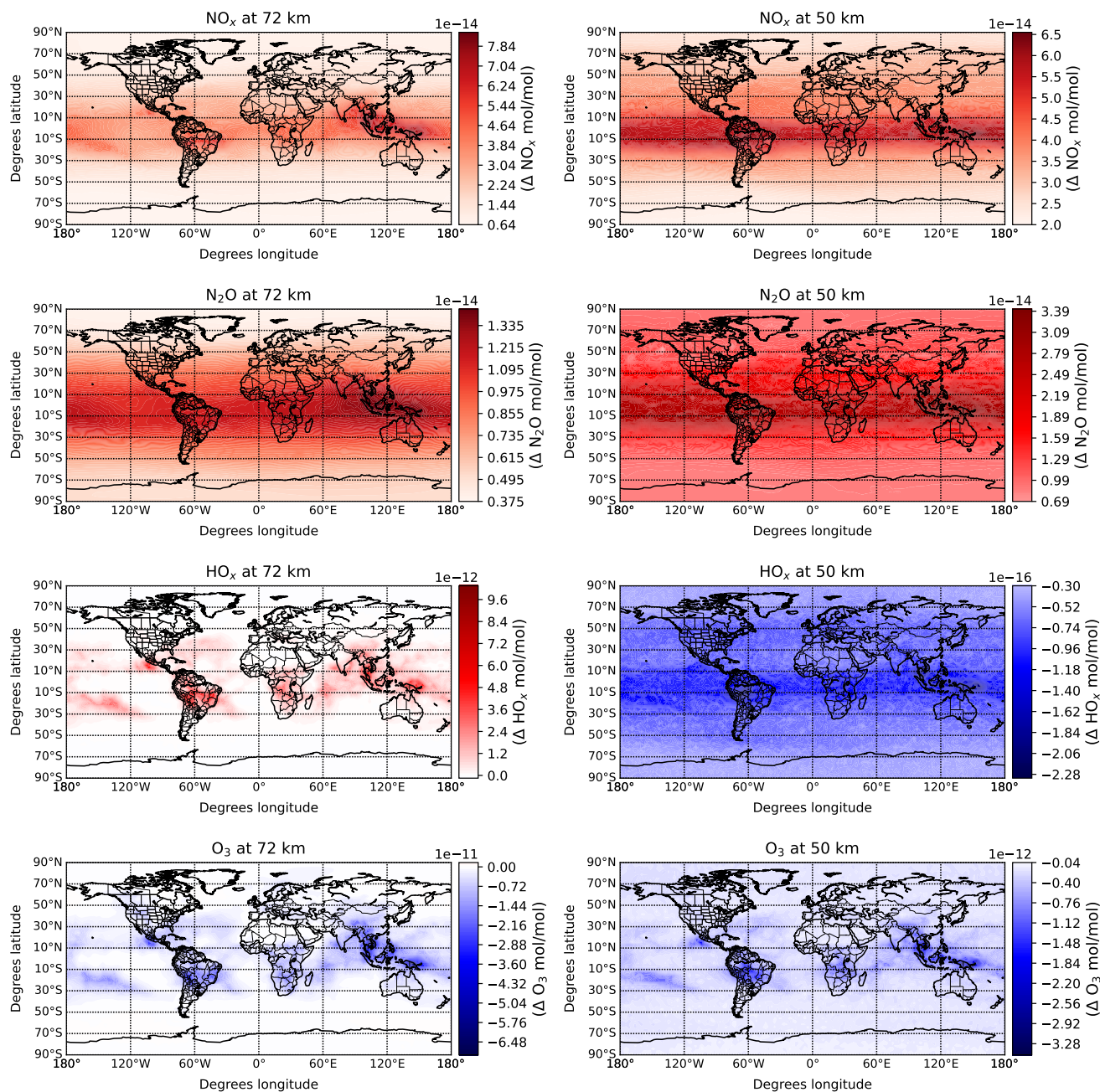
Figure 7. Same as Figure 6 but assuming that a single sprite injects 10^{25} HO$_2$ molecules (SPRI-SMI simulation) as reported by Yamada et al. (2020).
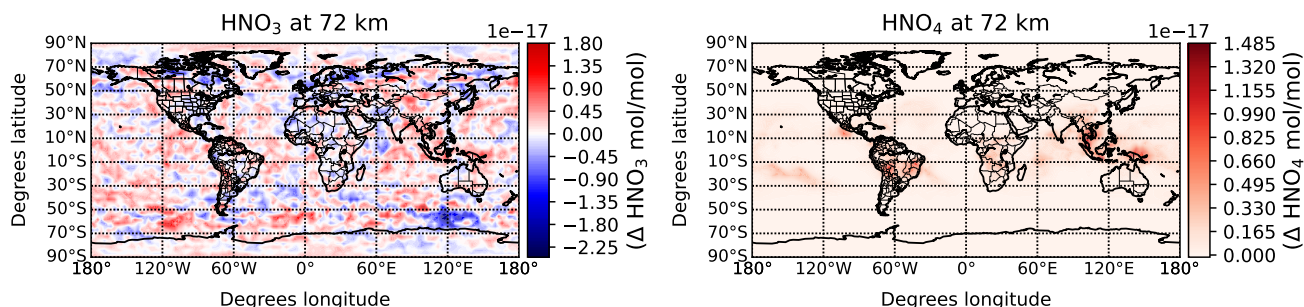
Figure 8. Same as Figure 7 but for the reactive nitrogen compounds HNO₃ and HNO₄.

of our simulations. We examine the ~~simulations CTRL and~~ simulation SPRI-M to generate vertical profiles of NO₂ and sprite frequency rates for each time step (720 s) between 22:00 and 23:00 local time ~~during August~~ in the period August to December 2001 within the latitude band 30°S to 20°N. Figure 9(a) shows the time series of the simulated NO₂ mixing ratio time series with and without sprites taking place up to 24 minutes before each given time according to the implemented parameterization of sprites. In total, Figure 9(a) comprises ~~366,358~~ 1,808,154 data points for NO₂ mixing ratios (black dots), from which ~~3,408~~ 11,850 correspond to post-sprite occurrences (red dots). Following ~~Arnone et al. (2014)~~Arnone et al. (2008), we have removed the NO₂ trend in Figure 9(b). The trend is calculated as the average of all the NO₂ mixing ratio data plotted in Figure 9(a). Comparison between the values shown in Figure 9(a,b) and ~~(Arnone et al., 2014, Fig. 2)~~ (Arnone et al., 2008, Fig. 2) confirms that the simulation is producing NO₂ mixing ratios at 52 km that are similar to the MIPAS measurements. The anomalies are analyzed in Figure 9(c), showing a positive NO₂ mixing ratio anomaly at 52 km altitude, and 24 minutes after the occurrence of sprites of +~~0.60~~0.56 ppbV, which is slightly lower than the +1 ppbV NO₂ anomaly reported by Arnone et al. (2014) above thunderstorms observed by MIPAS. We ~~have performed a t-test on the equality of means of distributions shown in Figure 9(c), obtaining a $p-value = 1.2 \times 10^{-185}$, which suggests that the difference is statistically significant .~~ computed the 95% confidence intervals for the mean and median of the distributions using a bootstrap method (Efron and Tibshirani, 1994) with 5000 resamples. The calculated confidence intervals are (-0.0022 ppbV, 0.0023 ppbV) and (0.5334 ppbV, 0.5843 ppbV) for the means of the distribution of simulated NO₂ mixing ratio anomalies at 52 km altitude and the distribution of NO₂ anomalies for cases in which sprites occurred less than 24 minutes before, respectively. Notably, these intervals do not overlap, suggesting significant differences between the distributions.

In their study, ~~Arnone et al. (2014)~~ Arnone et al. (2008) reported NO₂ anomalies over an instantaneous field of view with a footprint of 30 km × 1200 km, a size roughly three times smaller than the area covered by each cell domain in our simulations. As a result, it can be expected that the simulated anomalies are lower than those reported by ~~Arnone et al. (2014)~~ Arnone et al. (2008).
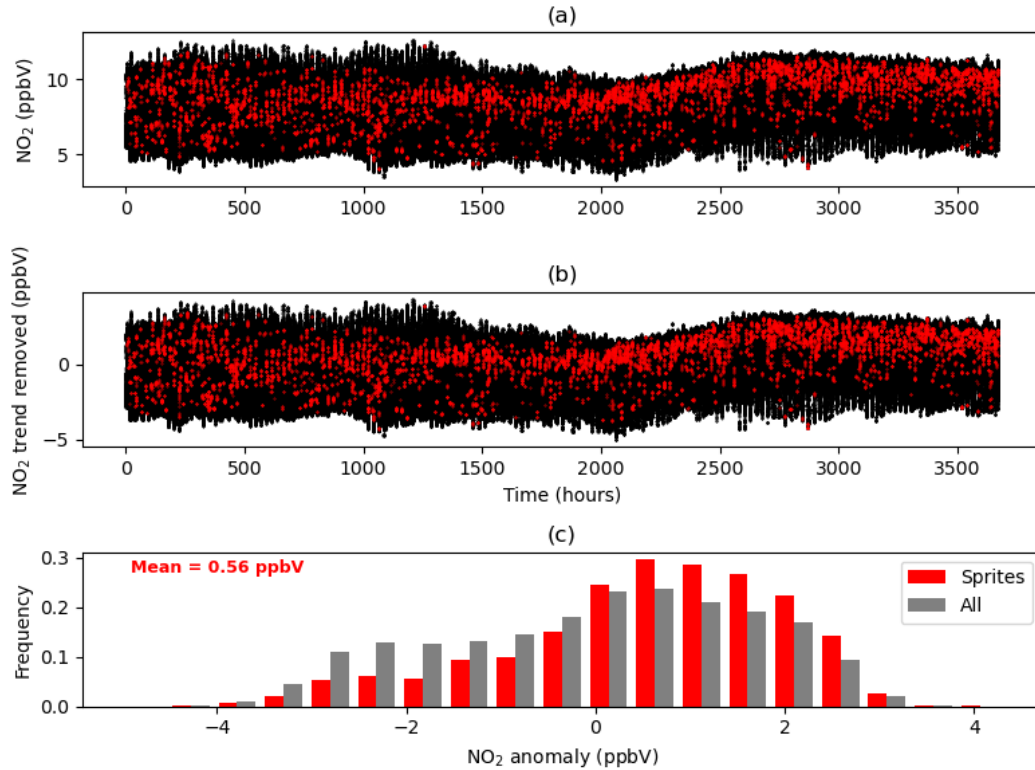
Figure 9. (a): Simulated NO$_2$ mixing ratio time series within the latitude band 30°S to 20°N between 22:00 and 23:00 local time every 720 s. The horizontal axis represents the time elapsed since August 1, 2001, at 00:00:00 UTC. Black dots denote NO$_2$ mixing ratios when no sprites occurred within the 24-minute window, whereas red dots signify NO$_2$ mixing ratios during the presence of ~~sprites~~ at least one sprite within the same 24-minute period. (b): Same as (a) but after NO$_2$ trend removal. (c): The ~~black~~ grey distribution shows the simulated NO$_2$ mixing ratio anomalies at 52 km altitude within the latitude band 30°S to 20°N between 22:00 and 23:00 local time every 720 s. The red distribution shows the NO$_2$ anomalies for cases in which sprite took place less than 24 minutes before.

4 Conclusions

We have developed and implemented in EMAC the first parameterization of sprites based on meteorological variables used as proxy. This parameterization has enabled us to simulate the global annual and seasonal global distributions of sprites and to estimate their sensitivity to climate change. In particular, we have obtained a future increase of the occurrence rate of sprites of 69% in 2091, which is larger than the predicted increase in lightning activity (about 47%). Recent modeling results and space-based measurements have been used to introduce the injection of chemical species by sprites in the model. We have found that the chemical influence of sprites in the mesosphere is not important at a global scale. However, our results indicate that sprites could be a non-negligible (measurable) source of HNO_4 at a regional scale, especially in the upper-mesosphere in South America and Southeastern Asia.

The analysis of simulated NO_2 mixing ratios above thunderstorms after the occurrence of sprites has confirmed that the anomalies in the nighttime NO_2 measurements reported by MIPAS after the occurrence of lightning (Arnone et al., 2014) can be due to sprites. In particular, our simulations indicate an enhancement of +~~0.6~~0.56 ppbV of the NO_2 mixing ratio above thunderstorms, at 52 km altitude within a 24 minutes window, while the increase reported by Arnone et al. (2014) was +1 ppbV.

The main conclusions of this study are:

1. The developed parameterization of sprites produces a good agreement between the simulated and the observed global distribution of sprites.

2. Implementation of sprites in EMAC ~~produces variations between +0.002% and~~ (see Section 3.2) produces a variation of +~~0.003~~0.008% of the mixing ratio of N_2O, ~~between +0.002% and~~ +~~0.06~~0.006% of NO_x, between -1 × 10^{-5}% and ~~0.0~~ +0.02% of HO_x, between -0.002% and -10^{-5}% of O_3, between +~~0.01~~ 0.005 and +0.05% of HNO_3 and between +~~0.0025~~0.007% and + 0.3% of HNO_4 between 60 km and 80 km altitude in the mesosphere.

3. The ~~globalscale~~ influence of sprites on the chemistry of the atmosphere at a global scale is negligible.

4. Our results confirm that NO_2 mixing ratio anomalies reported by MIPAS at 52 km altitude after the occurrence of lightning can be due to sprites.

5. The projected simulation with sprites (RCP6.0) indicates a 54% increase (14% per K) in their occurrence rate at the end of the 21st century, approximately similar to the expected increase of lightning activity.

Code and data availability. The Modular Earth Submodel System (MESSy) is continuously developed and applied by a consortium of institutions. The usage of MESSy and access to the source code are licensed to all affiliates of institutions which are members of the MESSy Consortium. Institutions can become a member of the MESSy Consortium by signing the MESSy Memorandum of Understanding. More information can be found on the MESSy Consortium website (http://www.messy-interface.org, last access: 29 September 2023). As the MESSy code is only available under license, the code cannot be made publicly available. The parameterization of sprites has been developed

based on MESSy version 2.54 and will be included in version 2.55. The data of the simulations generated in this study have been deposited in the Zenodo repository Pérez-Invernón et al. (2023)

Author contributions. F.J.P.I.: Conceptualization, methodology, validation, formal analysis, investigation, data curation, writing—original draft. F.J.G.V, A.M.R. and P. J.: Conceptualization, methodology, validation, formal analysis, investigation, data curation, writing—review and editing.

Competing interests. At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

Acknowledgements. The project that gave rise to these results received the support of a fellowship from "La Caixa" Foundation (ID 100010434). The fellowship code is LCF/BQ/PI22/11910026 (FJPI). Additionally, this work was supported by grants PID2019-109269RB-C43 (FJPI and FJGV) and PID2022-136348NB-C31 (F.J.P.I. and F.J.G.V) funded by MCIN/AEI/ 10.13039/501100011033 and "ERDF A way of making Europe". AMR acknowledges the financial support from the grant BEVP34A6840 funded by "Ramón Areces Foundation". FJPI and FJGV acknowledge financial support from the grant CEX2021-001131-S funded by MCIN/AEI/ 10.13039/501100011033. PJ acknowledges funding from the Initiative and Networking Fund of the Helmholtz Association through the project "Advanced Earth System Modelling Capacity (ESM)" and from the Helmholtz Association project "Joint Lab Exascale Earth System Modelling (JL-ExaESM)". The content of the paper is the sole responsibility of the author(s) and it does not represent the opinion of the Helmholtz Association, and the Helmholtz Association is not responsible for any use that might be made of the information contained. The high performance computing simulations (HPC) have been carried out on the DRAGO supercomputer of CSIC.

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