

# Evaluating an Earth system model from a water manager perspective

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**Abstract.** The large spatial scale of global Earth system models (ESM) is often cited as an obstacle to using the output by water resource managers in localized decisions. Recent advances in computing have improved the fidelity of hydrological responses in ESMs through increased connectivity between model components. However, the models are seldom evaluated for their ability to reproduce metrics that are important for and resonate with practitioners, or that allow practitioners to situate higher-resolution model outputs within a cascade of uncertainty stemming from different models and scenarios. We draw on the combined experience of the author team and water manager workshop participants to identify salient water management metrics and evaluate whether they are credibly reproduced over the conterminous U.S. by the Community Earth System Model v2 Large Ensemble (CESM2). We find that while the exact values may not match observations, aspects such as interannual variability can be reproduced by CESM2 for the mean wet day precipitation and length of dry spells. CESM2 also captures the proportion of annual total precipitation that derives from the heaviest rain days in watersheds that are not snow-dominated. Aggregating the 7-day mean daily runoff to HUC2 watersheds also shows rain-dominated regions capture the timing and interannual variability in annual maximum and minimum flows. We conclude there is potential for far greater use of large ensemble ESMs, such as CESM2, in long-range water management decisions to supplement high resolution regional projections.

## 1 Introduction

Water availability and water quality for human consumption, ecosystems, and agriculture are fundamental requirements, making pertinent assessments of future change crucial for adaptation planning (IPCC, 2022). Climate related changes in the hydrologic cycle will affect substantial portions of the world population, most directly through changes in water availability at or near the surface (Mankin et al., 2020; Sedláček and Knutti, 2014). The information required by water managers for decision making is not readily available in a relevant format, or at sufficient spatial or temporal resolutions from global Earth system models (ESM; e.g., Ekström et

36 al., 2018). We explore how the Community Earth System Model (CESM) represents the climatology of water  
37 availability, focusing on metrics that are familiar to decision makers in planning investment-scale decisions.  
38 The inability of ESMs to explicitly resolve sub-grid scale (~100 km) processes is often cited as the limitation  
39 preventing direct model use in decision making. Literature from large organizations making infrastructure  
40 decisions (e.g., Brekke, 2011; Brekke et al., 2009; Reclamation, 2016, 2014) emphasize downscaling climate  
41 model data closer to the scale of the watersheds they manage. These additional modeling steps add complexity  
42 and may increase statistical errors (Clark et al., 2015; Ekström et al., 2018). Extracting useful and robust  
43 information directly from ESMs would reduce such errors if metrics most important to decision makers, such as  
44 the timing of peak flow, were known to be robustly represented.

45

46 There are many comprehensive examples of metrics used to evaluate climate and hydrological models (e.g.,  
47 Ekström et al., 2018; Mizukami et al., 2019; Wagener et al., 2022), and communicate the impacts of climate  
48 change (e.g., Reed et al., 2022), or to identify decision-relevant metrics (e.g., Bremer et al., 2020; Mach et al.,  
49 2020; Underwood et al., 2018; Vano et al., 2014). However, very few have examined whether user defined metrics  
50 can be reliably reproduced by ESMs (Mankin et al., 2020), and if further model development and scale reduction  
51 is warranted instead of improved communication (Pacchetti et al., 2021). Better communication may also reduce  
52 the temptation of some users to calculate “standard hydroclimate metrics” that are not supported by the climate  
53 model data (Ekström et al., 2018).

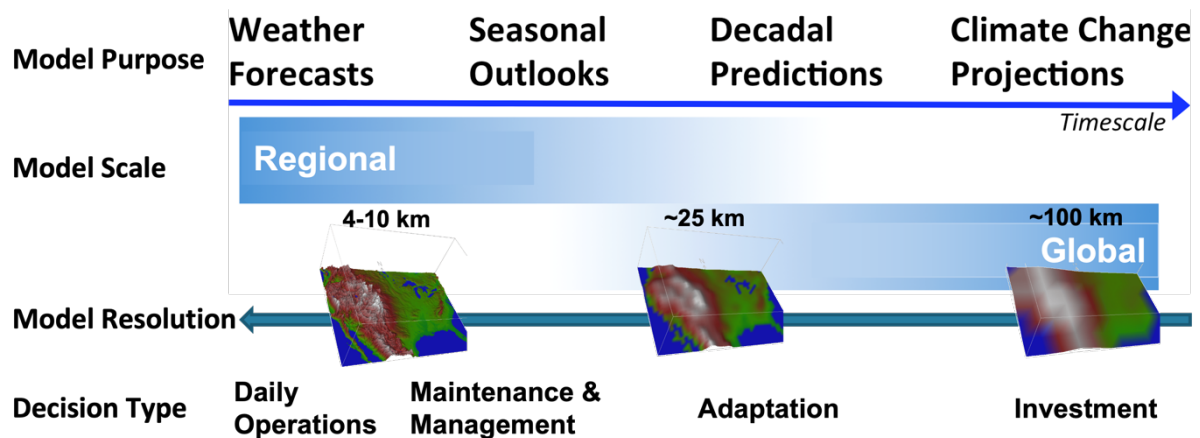
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55 In contrast, climate model output can be rejected unnecessarily when simulated annual minima from freely  
56 running simulations do not “match” the sequence of observed low flows (Ekström et al., 2018; Moise et al., 2015).  
57 Similarly, the benefits of a range of projected outcomes from different climate models are not widely appreciated  
58 beyond the climate model community (Tebaldi and Knutti, 2007). Large ensembles from a single climate model  
59 initialized with a range of atmospheric and ocean conditions, such as the CESM2 Large Ensemble (LENS2;  
60 Rodgers et al., 2021), help to bound the uncertainty that derives from a naturally chaotic system. Averaged over  
61 the full ensemble, they give a better estimate of the model’s response to internal and external forcing (Deser et al.,  
62 2012) and enable assessments of the rarity of projected extremes. The additional analysis to identify structural  
63 (i.e. model formulation) and internal variability within regional climate models means that there are fewer large  
64 ensembles at a high resolution (Deser et al., 2020).

65

66 Since different decision makers have different priorities and time-scales of interest, Shepherd et al. (2018)  
67 recommended the development of climate storylines to communicate with those using climate data to make  
68 decisions. Informed by prior surveys of water managers (e.g., Brekke, 2011; Brekke et al., 2009; Cantor et al.,  
69 2018; Raff et al., 2013; Wood et al., 2021), Fig. 1 aims to map the different types of water decisions (e.g., Raff et  
70 al., 2013 Fig. 3) to the different scales of model resolution (Meehl et al., 2009 Fig. 2). Water managers make daily  
71 operational decisions (e.g., to control instantaneous river flow) with the aid of fine-scale weather and flood models  
72 (<4 km) that reliably represent convective and local weather scale processes even though their predictability is  
73 relatively short lived (Yuan et al., 2019; far left side of Fig. 1). Larger watershed operations (such as reservoir  
74 management or groundwater recharge; e.g., Regional Water Authority, 2019) depend on seasonal outlooks

75 (middle left of Fig. 1). Smaller adaptation and mitigation projects take place at the typical policy or decadal  
 76 prediction scale (i.e., 4-10 years; middle right of Fig. 1). Finally, major public investments and inter-basin  
 77 agreements occur at the same time scales as climate projections (30-100 years; far right of Fig. 1) where persistent  
 78 and relatively predictable synoptic and planetary scale processes are well represented in lower resolution (~100  
 79 km) climate models (Phillips et al., 2020). While forecasts (seasonal or decadal) are re-initialized from specific  
 80 atmosphere, ocean or land states at regular time intervals, climate projections are run freely from a variety of  
 81 atmospheric and oceanic conditions that take several decades to converge to a mean climatology. In considering  
 82 the utility and useability of information directly from ESMs we focus on decisions made over decadal to climate  
 83 scales at larger spatial scales.  
 84



85  
 86 **Figure 1: Mapping the temporal and spatial scales of models to the timeframes for water management decisions.**

87 Given that ESMs have advanced immeasurably in the recent decade, it is time to re-evaluate whether their direct  
 88 output can support decision makers. Such an evaluation needs to focus on how well the models can reproduce  
 89 metrics used by decision makers, and whether the results are credible (Briley et al., 2020; Jagannathan et al.,  
 90 2021). Here we evaluate the credibility of one ESM in generating metrics known to be salient for water  
 91 management decisions; specifically, decisions for water management infrastructure project investments.  
 92

93 The motivation for this paper is to identify:

- 94 ● a set of water availability metrics that resonates with decision makers and supports their investment-scale  
 95 decisions;
- 96 ● how well CESM2 represents the climatology and recent observed behaviors of those metrics; and
- 97 ● the range of CESM2 structural uncertainty and internal variability for these metrics.

98  
 99 This paper builds off a decade of collaboration between scientists at the National Center for Atmospheric Research  
 100 (NCAR) and US water agencies that led to a virtual workshop (Tye, 2023), and presents a test case for improved  
 101 communication with water management decision makers. The focus is on the Conterminous United States  
 102 (CONUS) to match the interest of workshop participants.

103 **2 Climate Information Needs from Prior Research**

104 Information needs vary greatly, from 5-minute rainfall totals at a point (ASCE, 2006), to basin-wide measures of  
105 annual minimum and maximum total runoff. Water management decision metrics can be grouped into similar  
106 types such as timing, frequency, magnitude, extreme values, variability, and duration of events (Ekström et al.,  
107 2018). While some aspects of timing, magnitude, or variability can be reliably reproduced by ESMs (e.g., Deser  
108 et al., 2020; Tebaldi and Knutti, 2007), others such as short duration extremes are less reliable.

109

110 Methods of evaluation and data use also differ. For instance, Clifford et al. (2020) reported that predicting general  
111 changes in the frequency of extreme precipitation events is more useful for future planning than the precise  
112 prediction of mean values evaluated by model developers. Lehner et al. (2019) emphasized that models need to  
113 be evaluated for their ability to reproduce sensitivities (e.g., streamflow changes in response to temperature and  
114 precipitation changes) in addition to mean states. However, metrics that are meaningful for evaluating a model's  
115 capabilities (e.g., the ratio of precipitation to runoff) are less valuable for management decisions (Lehner et al.,  
116 2019; McMillan, 2021; Mizukami et al., 2019). When reporting results, water managers are more familiar with  
117 the 'water year', rather than the calendar year, to capture the full annual hydrological cycle (Ekström et al., 2018).  
118 While the use of water years is a nuance that does not add substantial value to climate model assessments,  
119 communication with decision makers is improved by presenting data in a familiar format (Briley et al., 2020).

120

121 There is a need for information at the local scale that is unlikely to be met directly by raw outputs from the current  
122 generation of ESM. But better communication of the variability in future daily precipitation and associated runoff  
123 can add value to the detailed models by bringing in the added statistical context and perspective of the large  
124 ensembles. Thus, we believe that ESMs can produce useful information about hydro-meteorological extremes  
125 when presented at different spatial or temporal scales, and offer the benefits of large climate model ensembles to  
126 constrain future impact uncertainty.

127

128 Appendix A summarizes potential hydrological metrics used in water management decisions (Jagannathan et al.,  
129 2021) or statistical assessments of extremes (Zhang et al., 2011), and model evaluations (Phillips et al., 2020).  
130 Metrics in bold are presented in this paper. We only considered a simplistic measure of meteorological drought  
131 (absence of rain) in the current work, as drought is sensitive to the definition (Bachmair et al., 2016) and local  
132 conditions (Mukherjee et al., 2018), and so not suited to a generalized assessment. Similarly, snow measures are  
133 not included in this assessment in part due to limited availability of high-quality, long-duration, quality-controlled,  
134 observational data (McCrary et al., 2017); and partly due to the biases in snow distribution arising from the  
135 smoothed topography in GCMs (McCrary et al., 2022).

136 **3 Data and Methods**

137 **3.1 Climate Model Data**

138 CESM2 (Danabasoglu et al., 2020) is a fully coupled global model that simulates the Earth's climate system  
139 through interactive models for atmosphere, ocean, land, sea-ice, river runoff, and land-ice. Variables considered

140 in this project are taken from the Community Atmosphere Model version 6 (CAM6) and the Community Land  
141 Model version 5.0 (CLM5; Lawrence et al., 2019) and are part of the default model outputs. A schematic of the  
142 model components is included in Appendix B. This project uses daily values scaled up to annual (e.g., annual  
143 maximum daily precipitation) on a ~1 degree resolution grid. Data were extracted over the CONUS from 10  
144 ensemble members of LENS2 (Rodgers et al., 2021) for model validation in the current era (1981-2010).

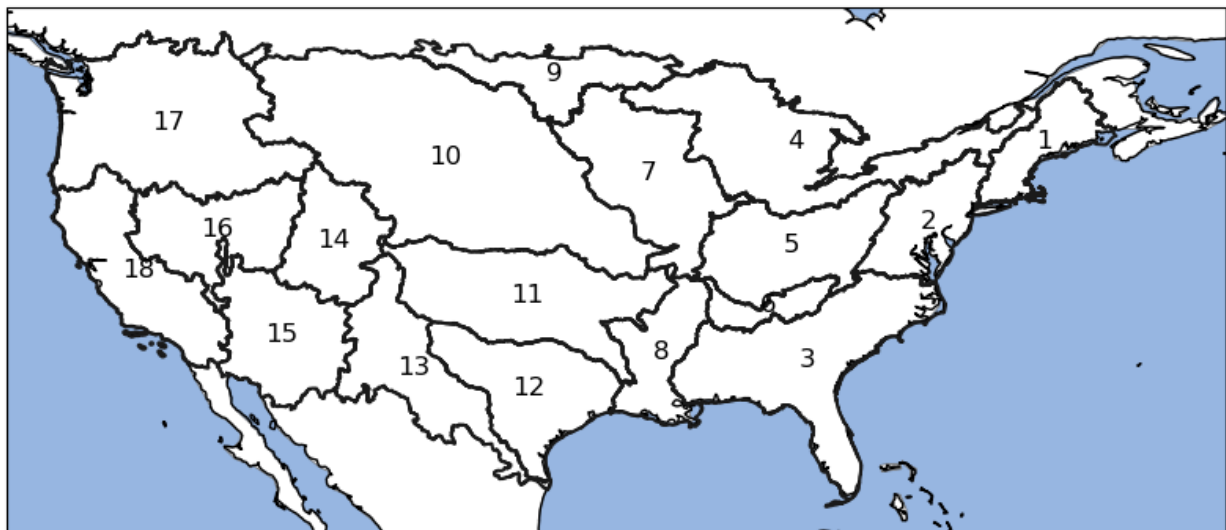
### 145 3.2 Observations

146 Gridded daily observations of precipitation at 1/16° horizontal resolution (~6 km) were obtained from the Livneh  
147 et al. (2013) dataset covering CONUS and southern Canada for the control period (1981-2010), hereafter referred  
148 to as “Livneh”. Pierce et al. (2021) provided an update to the Livneh data set to address time adjustments that  
149 result in an underestimation of the most extreme daily precipitation totals and resultant runoff and flood potential  
150 (Pierce et al. 2021). However, as we are also interested other measures of precipitation and in runoff minima, we  
151 did not employ the updated gridded observations.

152

153 Livneh daily temperature maxima and minima, and precipitation were used to force the Variable Infiltration  
154 Capacity Model (VIC; Liang et al., 1994) version 4.1.2 to obtain runoff estimates for years 1980-2005 as evaluated  
155 in Livneh et al. (2013). Hereafter referred to as “Livneh-VIC”.

156



157

158 **Figure 2: HUC 2 regions used in data validation and analysis. Regions defined by USGS (2013): Region 01 New England**  
159 **(NE); Region 02 Mid-Atlantic (MA); Region 03 South Atlantic-Gulf (SA); Region 04 Great Lakes (GL); Region 05 Ohio**  
160 **(OH); Region 06 Tennessee (TN); Region 07 Upper Mississippi (UM); Region 08 Lower Mississippi (LM); Region 09**  
161 **Souris-Red-Rainy (RR); Region 10 Missouri (MR); Region 11 Arkansas-White-Red (ARK); Region 12 Texas-Gulf**  
162 **(GUL); Region 13 Rio Grande (RIO); Region 14 Upper Colorado (UC); Region 15 Lower Colorado (LCO); Region 16**  
163 **Great Basin (GB); Region 17 Pacific Northwest (PN);**  
164 **Region 18 California (CA)**

### 165 3.3 Methods

166 All analyses were carried out using the North American water year (1 October to 30 September) to facilitate later  
167 communication.

### 168 **3.3.1 Remapping**

169 For ease of comparison, model output were re-gridded using a conservative second-order remapping (Jones, 1999)  
170 to place both datasets on the same scale grid and assess anomalies. Data were also calculated as areal averages or  
171 totals over the 2-digit Hydrological Unit Code (HUC2) regions (Seaber et al., 1987). HUC2 basins represent 18  
172 watersheds, covering areas ranging from 41,000 mi<sup>2</sup> (~105,000 km<sup>2</sup>; Tennessee) to 520,960mi<sup>2</sup> (1,350,000 km<sup>2</sup>;  
173 Missouri), shown in Fig. 2. While the scale of HUC2 regions may be large for some local decision-makers, it is  
174 also a more appropriate and conservative scale to compare to ESMs as demonstrated by Lehner et al. (2019).

### 175 **3.3.2 Percentile-based thresholds**

176 The threshold for very heavy rain days (Q95) was calculated at each individual grid cell using only days with  $\geq 1$   
177 mm rain (“wet days”). Thresholds were derived empirically for each model ensemble member, with the ensemble  
178 mean threshold (Q95) used to identify the days per year exceeding the threshold (N95) and total annual rainfall  
179 from those days (P95).

180 Runoff was aggregated over each HUC2 watershed and multiplied by the respective area to generate total volume  
181 per day. Volume per day was then converted to measurements more familiar to users, such as acre feet per day or  
182 cubic meters per second. Daily time series of total volumetric runoff had a 7-day running mean smoother applied,  
183 then annual maximum, minimum and mean values were extracted. The highest and lowest 7-day average runoff  
184 expected once per decade (7Q90, 7Q10) were estimated empirically from the 25 ranked values of annual maxima  
185 and minima per watershed. Stationarity was assumed over the climatological period for the purposes of these  
186 analyses, acknowledging that changes may have already occurred in the frequency of these events.

## 187 **4 Model Evaluation**

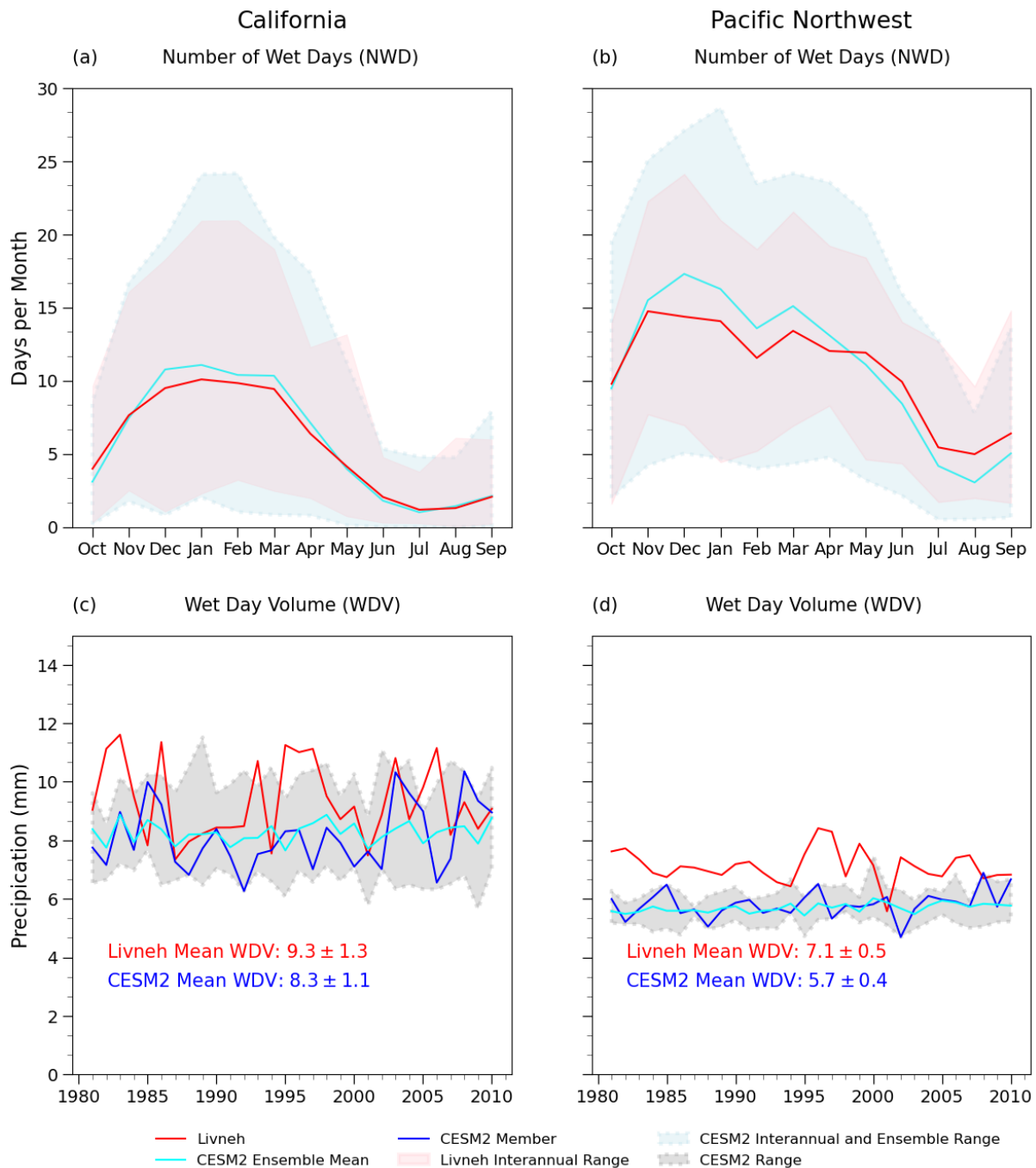
188 The metrics used to evaluate CESM2’s ability to reproduce large scale features and physical behaviors (e.g.,  
189 Danabasoglu and Lamarque, 2021 and the associated Special Issue) are not necessarily those employed by  
190 decision makers. ESMs are designed to represent large-scale atmospheric processes and fluxes not specific local  
191 responses (Gettelman and Rood, 2016), but this design assumption may not be sufficiently well communicated to  
192 decision makers. The purpose of our evaluation is to establish whether CESM2 output is also fit for local decision  
193 purposes, or if the breadth of information from ESM ensembles remains unsuitable for immediate use in targeted  
194 water management decisions.

### 195 **4.1 Rainfall metrics**

196 While broad spatial patterns of seasonal mean daily rainfall are reproduced well (Danabasoglu et al., 2020; Feng  
197 et al., 2020; Simpson et al., 2022), CESM2 fails to capture details over high topography, and overestimates  
198 summer precipitation where convective extremes dominate summer rainfall (Appendix C). The seasonal mean  
199 precipitation also fails to capture some important watershed-level processes, such as the seasonal variability in the  
200 number of days with precipitation and the associated intensity.

201

202 Estimates of mean annual rainfall on wet days, or wet day volume, are in broad agreement between Livneh and  
203 CESM2 output. Figure 3 shows an example of the mean number of wet days per month (NWD), and mean wet  
204 day volume (WDV) averaged over the California and Pacific Northwest. While CESM2 represents the NWD  
205 annual cycle very well in regions such as California (Fig. 3a, 3c) and the Pacific Northwest (Fig. 3b, 3d), it does  
206 not capture NWD in many central and snow dominated regions (Fig. S1 and Fig. S2). This is likely due to the  
207 smoother topography of CESM2 missing the influence of orographic uplift, and large spatial scale missing sub-  
208 grid scale convective systems (e.g., over the Central Plains). The figures also highlight the scale of model  
209 (structural and internal variability) uncertainty present in the ensemble. As noted in previous sections, water  
210 management decision-makers are aware of the potential scale of uncertainty and expressed a desire for the full  
211 ensemble range to be presented to them instead of ensemble means.



213

214 **Figure 3: Average number of wet days per month (a, b) and interannual variability in mean annual precipitation on**  
 215 **wet days for Livneh climatological mean (red) with interannual spread (pink) and CESM2 mean (cyan) with**  
 216 **interannual and ensemble spread (gray); and (c,d) between 1981-2010 for observations derived from Livneh (red) and**  
 217 **an individual CESM2 ensemble member (blue) and Ensemble spread (gray) in (a,c) Region 18 California (CA);**  
 218 **and (b,d) Region 17 Pacific Northwest (PN).**

219 The annual variability in WDV, both year-to-year variations as well as the overall range of minima and maxima,  
 220 is well captured by each of the model members for the different HUC2 regions, even if the absolute values do not  
 221 match (Fig. 3 c,d). As expected, the specifics of which years have high or low values of WDV are not the same  
 222 for each ensemble member (i.e. demonstrating internal variability). As a result, the ensemble mean value of WDV  
 223 (cyan) does not reflect the same year-to-year variability as the observations or individual ensemble members



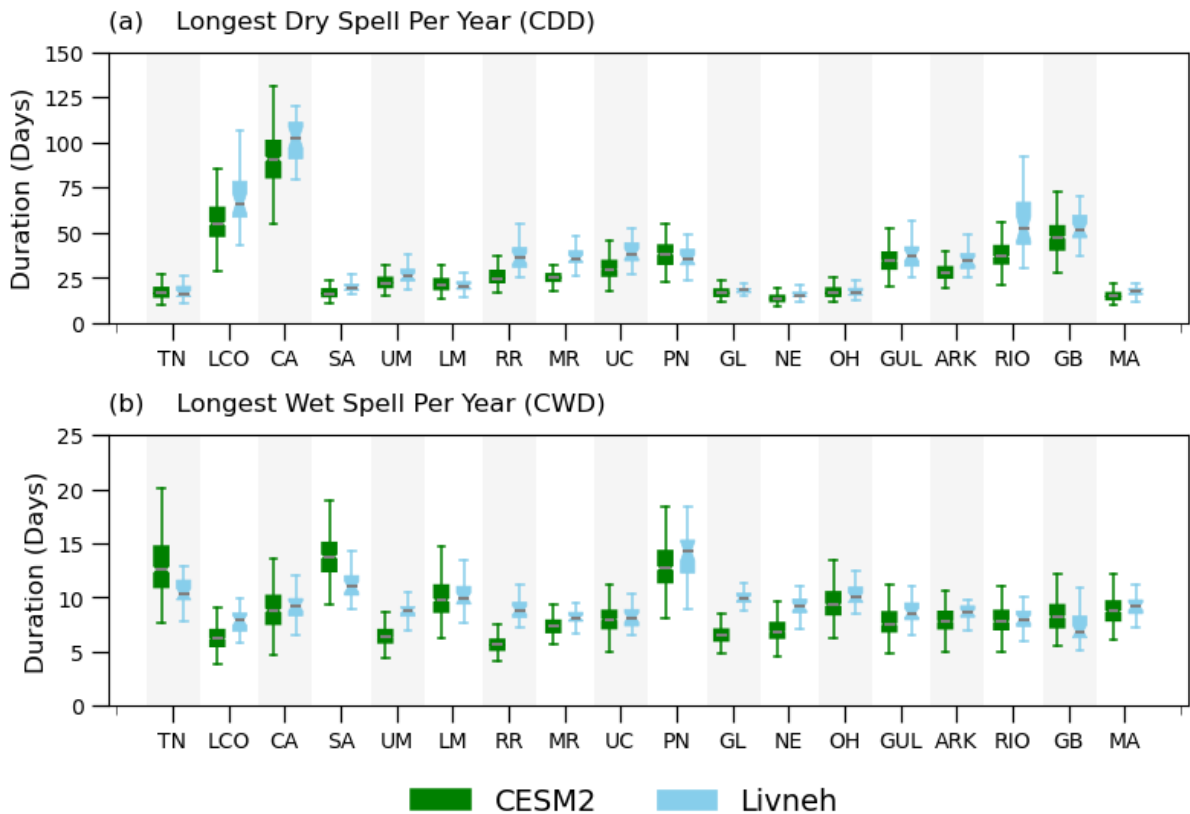
224 (blue). Decision makers expressed that the interannual variability demonstrated by each model member is more  
225 valuable to demonstrate the credibility of the data than the ensemble mean (Tye, 2023). We recommend that the  
226 full range of values of each metric (i.e. after computation for each ensemble member individually) are  
227 communicated in addition to the climatological means to help bound uncertainty around decisions (Wilby et al.,  
228 2021).

229

230 The magnitude of interannual variability in WDV (i.e., the absolute differences between the maximum and  
231 minimum values in each member time series) is typically within 10% of observations in all regions as illustrated  
232 for two regions in Fig. 3. Exceptions are the Lower Colorado, South Atlantic-Gulf and Upper Mississippi where  
233 the simulated distributions are too narrow. Many different sources of error may contribute to this discrepancy,  
234 such as the inability to resolve convective precipitation (Chen et al., 2021) in addition to elevation changes not  
235 captured by the coarse model resolution, or the “drizzle effect” that is common in GCMs (Chen et al., 1996; Dai,  
236 2006).

237

238 CESM2 captures the longest spells of consecutive dry days per year (CDD; Fig. 4a) and consecutive wet days per  
239 year (CWD; Fig. 4b), and their variability. Many regions capture both the interannual variability and the  
240 climatological mean duration of CWD, particularly those regions that are subject to large-scale synoptic systems  
241 (e.g., Pacific Northwest, Mid Atlantic-Gulf, California). Several regions either overestimate (South Atlantic-Gulf)  
242 or underestimate (Great Lakes, Souris-Red-Rainy) the absolute durations of the longest wet spells, but do reflect  
243 the magnitude of interannual variability. The exception is Tennessee, where both interannual variability and mean  
244 CWD are overestimated. At the grid scale, broad spatial patterns of CWD are correct but the finer atmospheric  
245 processes arising from topographic features are incorrect, as expected from the coarse model resolution. A similar  
246 pattern is present in CDD, except that some drier regions with CDD >30 days do not capture the full range of  
247 interannual variability (Souris-Red Rainy, Missouri, Rio Grande). As GCMs have a tendency to produce drizzle,  
248 adjusting for a higher wet day threshold (e.g., 2 mm) might improve dry spell representation in those regions. It  
249 is also important to communicate such model sensitivities to users more effectively.

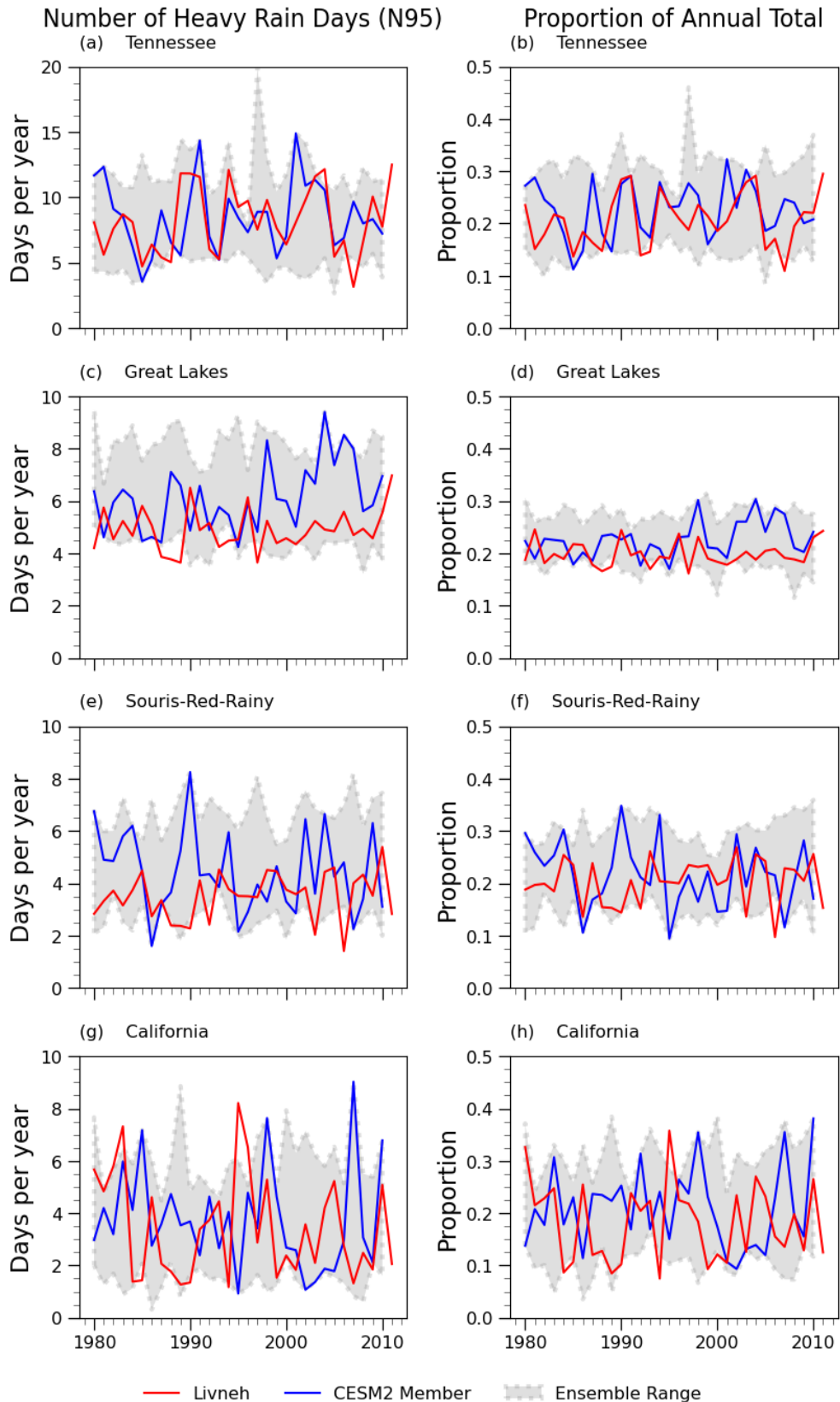


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251 **Figure 4: a) Longest duration per year of consecutive days <1 mm rain (longest dry spell) for Livneh over all years**  
 252 **(green) and CESM2 ensemble range over all years (blue) for all HUC2 regions; and b) Longest duration per year of**  
 253 **consecutive days with  $\geq 1$  mm rain (longest wet spell). Regional Acronyms defined in Fig. 2.**

254 The thresholds for heavy and very heavy rain days (P95, P99) are defined with respect to the wet days, and  
 255 calculated individually and compared for Livneh and CESM2 both to understand whether the intensity of more  
 256 extreme rainfall is captured, and to evaluate model behavior. A comparison of the thresholds reflects the  
 257 considerable improvements in modeling capabilities in recent years (Gettelman et al., 2022). For instance, earlier  
 258 versions of CESM underestimated extreme precipitation intensity by 10-30 mm/day east of the Rockies, and  
 259 overestimated intensity by 5-10 mm/day to the west (Gervais et al., 2014). We found CESM2 still underestimates  
 260 the most extreme rainfall, but that errors have approximately halved. As these differences are still inadequate for  
 261 many engineering and major infrastructure decisions (Wright et al., 2019), we focus on CESM2’s ability to capture  
 262 the relative contributions of P95 and P99 to the annual total and the interannual variability in their frequency. A  
 263 result with considerable useability is the proportion of annual total precipitation derived from the heaviest rain  
 264 days, or “Proportional Contribution of Extreme Days” (P95Tot). This proportion and its interannual variability is  
 265 well represented by CESM2 at the HUC2 scale and has shown to be skillful in other models (Tebaldi et al., 2021).  
 266

267 The interannual variability in the frequency (N95) and intensity of extreme rainfall, as represented by P95Tot, are  
 268 illustrated in Fig. 5 and 6. In several HUC2 regions the simulations report more frequent events, and proportionally  
 269 higher totals (e.g., Rio-Grande, Missouri, Upper Colorado and Lower Colorado; Fig. S3). Overall, there is good  
 270 subjective agreement between Livneh and CESM2, identifying an opportunity to inform local decisions from large  
 271 scale ESMs.



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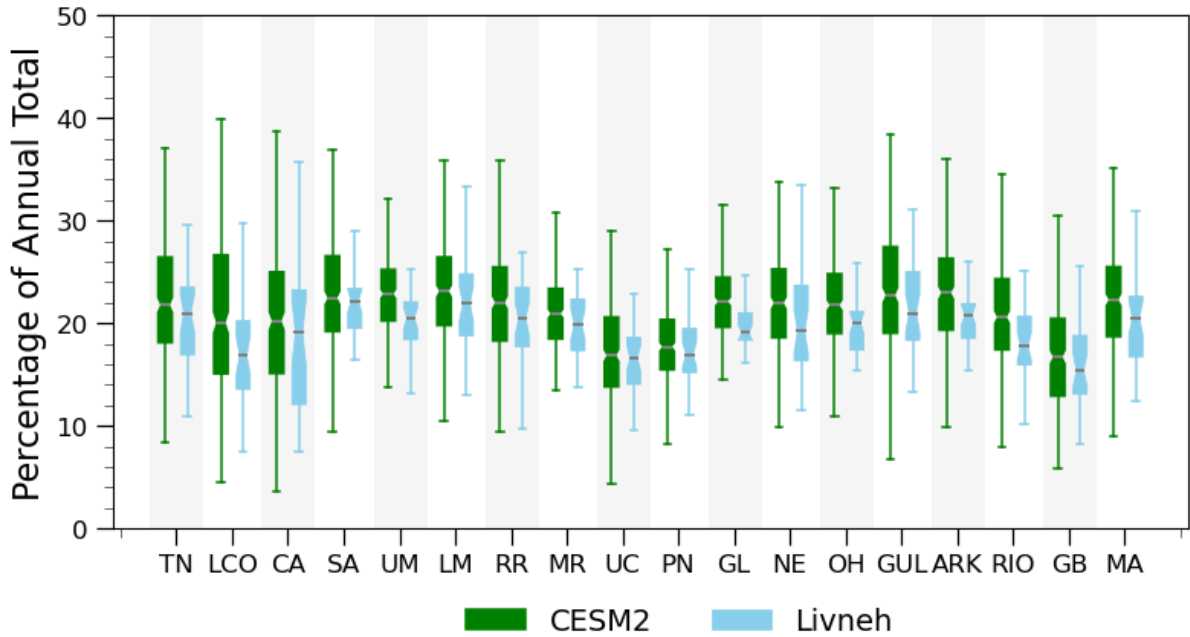
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Figure 5: a, c, e, g) Number of very heavy rain days per year; and b, d, f, h) total rain from very heavy rain days as a proportion of the annual total for a, b) Tennessee (TN); c,d) Great Lakes (GL); e,f) Souris-Red-Rainy (RR); and g,h) California (CA) HUC2 regions. Observations in red; CESM2 ensemble spread in gray, single randomly selected ensemble member in blue.



277  
 278 **Figure 6: Box plots of the interannual range of contributions to annual total rainfall from very heavy days (P95Tot)**  
 279 **shown as percentages for: Observations (light blue), and ensemble range for CESM2 (green) for all HUC2 regions.**  
 280 **Boxes are bound by the interquartile range, black lines indicate the median, notches indicate the degree of spread from**  
 281 **the median and bars extend to the full data range.**

#### 282 4.2 Runoff metrics

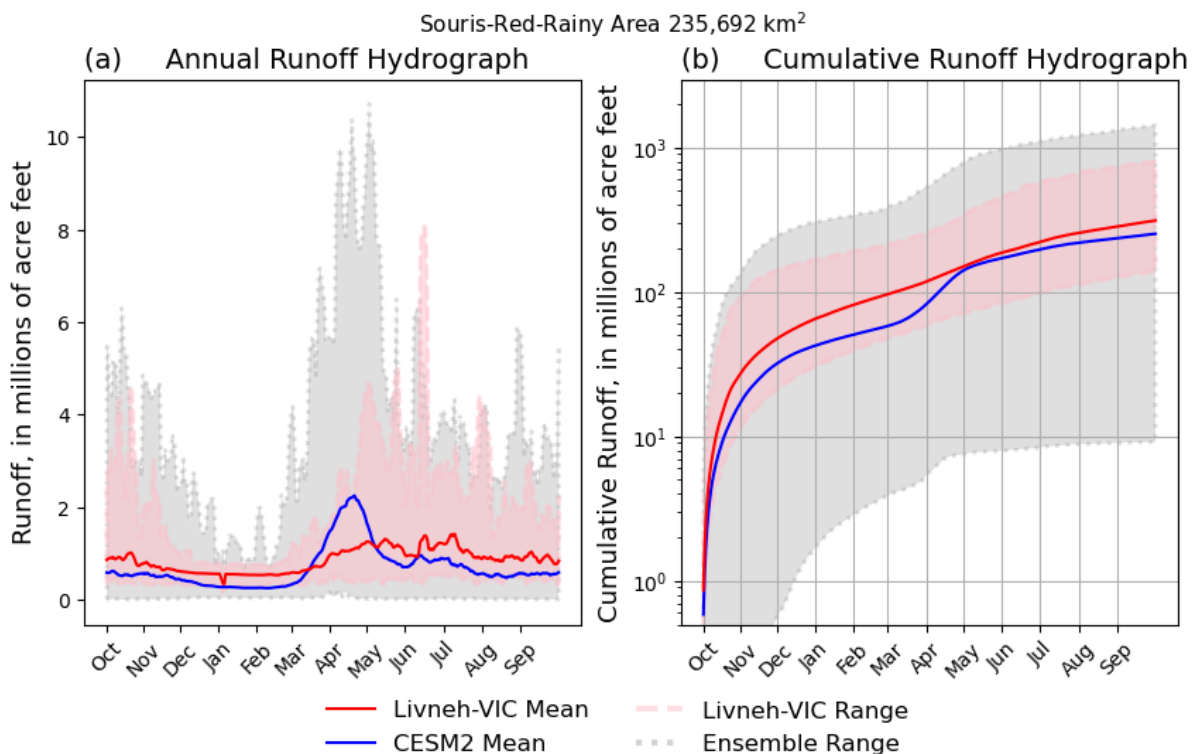
283 Runoff estimates are taken from the individual components of surface and subsurface runoff generated within  
 284 CLM5 (Lawrence et al., 2019) and compared to the Livneh forced VIC runoff (“Livneh-VIC”).

285  
 286 Assessing the skill of runoff in large-scale models is complicated by many factors, including the mismatch of  
 287 scales between in-channel flow ( $\sim 1\text{-}10^2$  m) and the grid scale ( $\sim 10^5$  m). Thus, metrics of climate model runoff  
 288 should be selected carefully and the runoff should be aggregated or combined with other metrics, rather than used  
 289 directly (Lehner et al., 2019). Appendix D demonstrates the discrepancies between the grid-scale representation  
 290 of runoff from Livneh-VIC and CESM2. The large discrepancies arise from different processes that are not  
 291 captured adequately, such as groundwater, topography, and associated snow ablation and melt, in addition to  
 292 meteorological biases.

293  
 294 However, water management decisions are made over watersheds in units such as acre-feet<sup>1</sup> or cubic meters, while  
 295 model data are output as a depth of runoff over each grid cell (e.g., mm/day per km<sup>2</sup>). We aggregated the 7-day  
 296 running mean daily runoff (Q7) within each HUC2 region to generate Q7 time series in each basin. Fig. 7a  
 297 illustrates the 25-year mean seasonal cycle for Livneh-VIC in red and CESM2 in blue, and the full range of values  
 298 over all years and ensemble members for the Souris-Red-Rainy basin (HUC Region 9), additional basins are  
 299 included in Fig. S4. Data are presented in millions of acre feet, to align with decision maker needs. The minimum  
 300 simulated Q7 in any year considerably underestimates the lowest flows in this region compared to Livneh-VIC.

<sup>1</sup> 1 Acre-foot is the volume of water it would take to cover 1 acre of land to a depth of 1 foot. Equal to 325,852 gallons or 1,233 m<sup>3</sup> (USGS Water Science).

301 In contrast, the largest total runoff volume is overestimated and peaks too early in the water year. Figure 7b plots  
 302 the same information as the cumulative runoff volume from the start of the water year, highlighting that the lowest  
 303 runoff volume is underestimated by a factor of ten for this region, other regions are illustrated in Fig. S5. Low  
 304 runoff volumes were typically underestimated in smaller regions (e.g., NE, TN). High runoff volumes were only  
 305 underestimated in three regions (LM, ARK, GUL) and considerably overestimated in seven regions. Snow-  
 306 dominated regions perform particularly poorly for both QMax and QMin as snowpack and the timing of associated  
 307 runoff are not well simulated. Transitional regions that straddle both snow- and rain-dominated hydrology also  
 308 fail to capture QMax, but better estimate Qmin (not shown). Only the South Atlantic region reproduces both  
 309 QMax and QMin.



310  
 311 **Figure 7: Interannual variability in runoff in Souris Red Rainy Region for a) the mean seasonal cycle; and b) the**  
 312 **cumulative watershed runoff over the water year. Livneh-VIC climatological mean in red, range of all years in pink;**  
 313 **CESM2 ensemble mean in blue and ensemble range in gray. Figure highlights the underestimation of the lowest runoff**  
 314 **volume by CESM2 by a factor of ten.**

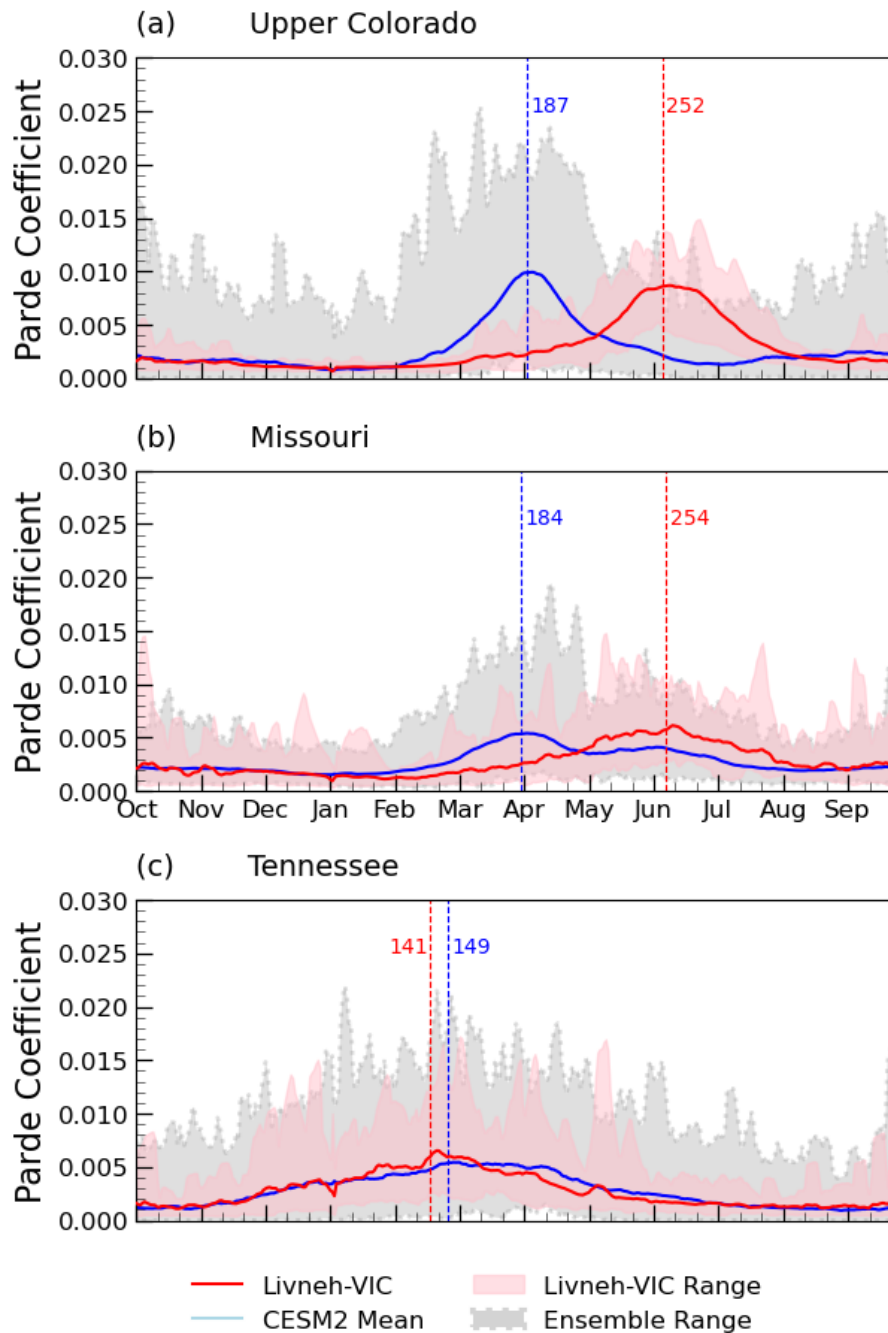
315  
 316 We explored the relationship between the highest and total annual runoff (QMax/QTot), and lowest and total  
 317 annual runoff (QMin/QTot). Some regions performed well for QMax/QTot, others performed better for  
 318 QMin/QTot but there was no consistent relationship that could be utilized by decision makers.

319  
 320 Participants at the NSF NCAR workshop (Tye, 2023) emphasized that the exact numbers produced by climate  
 321 models are not very important for future decisions. Others have also emphasized the importance of well-  
 322 represented processes in the model (Reed et al., 2022) and correlations with known experiences (Mach et al.,  
 323 2020; Shepherd et al., 2018). Focusing on fidelity to the historical climate exaggerates the importance of model  
 324 performance instead of robustness to different conditions without ensuring that model predictions are useful or

325 reliable (Brunner et al., 2021; Wagener et al., 2022). Runoff estimates in transitional catchments may be  
326 inadequate in the current climate but plausible in the future, if the model reproduces rain-dominated hydrological  
327 processes (McMillan, 2021).

328

329 Climatological mean runoff cycles are estimated from Pardé coefficients — calculated as  $Q7/Q_{Tot}$  on each  
330 calendar day — a dimensionless value that enables comparison across regions. Figure 8 depicts the mean seasonal  
331 cycle for representative snow-dominated (Upper Colorado), transitional (Missouri) and rain-dominated  
332 (Tennessee) regions, demonstrating how an imperfect representation of snow in the Upper Colorado results in  
333 CESM2 peak runoff occurring two months earlier than Livneh-VIC (Fig. 8a). The runoff regimes display very  
334 different seasonal characteristics, with CESM2 having a “mid late spring” runoff regime rather than Livneh-VIC’s  
335 “extreme early summer” regime (Fig. 8a; Haines et al., 1988). Peak runoff is also too early in transitional regions,  
336 but closer to Livneh-VIC than in snow-dominated regions (Fig. 8b). Rain-dominated regions capture both the  
337 timing of  $Q_{Max}$  and overall seasonal hydrograph shape (Fig. 8c).



338

339 **Figure 8 : Seasonal patterns of runoff for HUC2 regions a) Upper Colorado (UC); b) Tennessee (TN);**  
 340 **and c) Missouri (MR). Constructed from normalized series of the ratio of 7-day mean runoff to the mean annual total.**  
 341 **Livneh-VIC runoff climatological mean (red), climatological range (pink), CESM2 ensemble mean (blue) and ensemble**  
 342 **range (gray with dashed border). Vertical lines indicate the mean date of peak runoff with number of days since the**  
 343 **start of the water year.**

344 7Q10 and 7Q90 are estimated empirically from annual minima and maxima as occurring once per decade.  
 345 Projected changes in the frequency of very low or very high runoff volumes are deemed credible where CESM2  
 346 replicates the standard deviation of annual minima and maxima according to a  $\chi^2$  test at the 5% significance level.  
 347 Table 1 reports CESM2 and Livneh-VIC regional estimates of 7Q10 and 7Q90 and standard deviations of the  
 348 annual maxima and minima; values in bold indicate where estimates are statistically similar. It should be noted  
 349 that the values in Table 1 have  $\leq 10\%$  of occurring in any year, and so represent the tails of the runoff distribution.

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**Table 1 : Very low (7Q10) and very high (7Q90) regional runoff, and standard deviation in regional annual minima ( $\sigma$  QMin) and annual maxima ( $\sigma$  QMax) for Livneh and CESM2. Values in bold indicate where CESM2 and Livneh-VIC regional runoff are statistically similar according to a  $\chi^2$  test.**

Region		Livneh-VIC				CESM2			
		7Q10	7Q90	$\sigma$ QMin	$\sigma$ QMax	7Q10	7Q90	$\sigma$ QMin	$\sigma$ QMax
NE	1	4.1	132.4	1.3	25.5	8.6	215.1	4.7	39.9
MA	2	<b>6.9</b>	103.5	<b>2.5</b>	25.7	<b>7.4</b>	220.7	<b>3.6</b>	47.9
SA	3	<b>21.1</b>	<b>240.4</b>	<b>8.4</b>	<b>50.7</b>	<b>20.5</b>	<b>258.6</b>	<b>11.9</b>	<b>45.8</b>
GL	4	<b>6.9</b>	122.5	<b>2.2</b>	23.8	<b>7.8</b>	331.0	<b>4.3</b>	58.0
OH	5	<b>7.8</b>	187.6	<b>2.3</b>	53.0	<b>9.4</b>	260.9	<b>4.5</b>	56.4
TN	6	2.1	<b>90.5</b>	0.8	<b>23.1</b>	0	<b>98.7</b>	0.3	<b>21.7</b>
UM	7	2.1	78.2	1.7	16.9	7.9	122.3	4.7	31.5
LM	8	3.9	212.2	1.1	36.1	8.0	81.0	5.1	14.7
RR	9	1.0	<b>24.3</b>	0.5	<b>7.1</b>	0	<b>33.0</b>	0.1	<b>8.4</b>
MR	10	2.3	103.0	1.6	28.1	5.2	147.4	4.2	30.4
ARK	11	2.2	130.5	0.7	36.2	3.2	93.9	4.5	18.1
GUL	12	1.5	99.1	0.5	35.5	1.3	70.7	2.8	16.7
RIO	13	<b>0.5</b>	<b>22.5</b>	<b>0.2</b>	<b>5.8</b>	<b>0.4</b>	<b>29.5</b>	<b>1.3</b>	<b>7.3</b>
UC	14	0.6	27.3	0.2	7.2	0	74.7	0.2	15.3
LCO	15	0.5	19.4	0.2	7.5	0.3	46.7	0.7	11.6
GB	16	0.7	33.3	0.3	10.3	1.8	71.5	1.3	21.1
PN	17	20.6	266.5	7.9	50.2	4.4	449.6	2.6	87.3
CA	18	1.6	323.2	0.4	101.9	1.3	233.4	1.1	61.3

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Grid-scale estimates such as mean daily runoff readily highlight why decision makers have low confidence in CESM2 output: the metrics are not salient and appear to have no skill. After aggregating the 7-day mean daily runoff to watershed scales, some skill emerges in the annual minima and maxima, and seasonal cycles. Snow-dominated watersheds perform poorly with regard to peak runoff volume and timing of the peaks and lows, as expected (McCrary et al., 2022). Rain-dominated watersheds capture the inter-annual variability and magnitudes of peak and low flows, and the seasonal hydrographs. While CESM2 at this coarse scale does not represent the local topography and cannot represent finer scale snow, our analysis indicates the land surface model correctly



362 simulates the overall bulk water budget for most watersheds as illustrated in Figures 7 and 8. However, the tail  
363 behavior of highest and lowest total watershed runoff is only captured by a few basins and so caution needs to be  
364 exercised in the interpretation and use of model results, as biases may propagate into the future. This is premised  
365 on the understanding of *why* the model can produce accurate results, and whether the accuracy can be reliably  
366 reproduced for the future climate (Wagener et al., 2022).

367

368 While participants at the NSF NCAR workshop stated that precise estimates are not necessary, they also  
369 emphasized their desire for high confidence in the projected scale and direction of any changes. We note that  
370 “confidence” is derived from a combination of 1) credible process representation; 2) agreement with historical  
371 trends, given internal variability; 3) agreement across multiple models. It is worth noting that trends in extremes  
372 may be important without being statistically significant, as a limited sample of points (e.g. one per year) from a  
373 stochastic series is inherently noisy. However, some of these trends may emerge from the noise in the distribution  
374 and so are important to monitor.

## 375 **6 Discussion**

376 As decision makers have become more immersed in developing water management adaptation plans, the role of  
377 “climate services” in developing salient climate information has increased (Briley et al., 2020; Brugger et al.,  
378 2016; Dilling et al., 2019). We tested our hypothesis that recent improvements in ESMs can allow decision-  
379 relevant metrics to be produced directly, by leveraging the combined experience of the author team, results from  
380 the NCAR workshop, and the wealth of literature on actionable knowledge (Bremer et al., 2020; Jagannathan et  
381 al., 2021; Mach et al., 2020; Vano et al., 2014). Given that no model can perfectly address all decision needs, we  
382 identified and evaluated multiple metrics that can frame specific water management decisions within the known  
383 constraints of the data (Lempert, 2021), or within the decision makers’ experiences (Austin, 2023; Clifford et al.,  
384 2020; Reed et al., 2022; Shepherd et al., 2018).

385

386 It is important to communicate the original purpose of the model and associated weaknesses, so that decision  
387 makers fully understand which information is appropriate to use in other applications (Fisher and Koven, 2020;  
388 Gettelman and Rood, 2016; Wagener et al., 2022). Given the balance between model fidelity and model  
389 complexity (Clark et al., 2015) and the absence of detailed global scale observation data (e.g., Gleason and Smith,  
390 2014; Reba et al., 2011) CESM2 provides a plausible representation of Earth system processes and moisture  
391 fluxes, but may not capture basin-scale specifics (Ek, 2018; Lehner et al., 2019). That said, there are continued  
392 efforts to improve the simulation of land surface processes and analyses such as those presented in this article can  
393 flag weaknesses for future improvement (Lawrence et al., 2019).

394

395 Establishing model fidelity also requires distinguishing an accurate representation of the climate processes from  
396 serendipitous correlation with observations. Whether the model has good process representation overall, or  
397 exactitude in one simulation can be established through internal variability analyses using large ensembles (e.g.,  
398 Deser et al., 2020; Tebaldi et al., 2021). Repeating the analyses with several different ESMs to establish the degree

399 of agreement (Mankin et al., 2020) would further strengthen the usability of metrics presented in this article. It is  
400 also worth noting that the analysis presented here only used one reference dataset. As different reanalysis and  
401 observational datasets can have large discrepancies, a thorough model evaluation would also benefit from  
402 comparison to several products (Kim et al., 2020; Newman et al., 2015), including an assessment of how removing  
403 temporal adjustments in observations affects the statistics of extremes (Pierce et al., 2021).

404

405 While the precise details of precipitation and runoff may not be well simulated by CESM2, we found some aspects  
406 are credible. The frequency of wet days highlighted regions where current seasonal behavior is well captured, and  
407 may support planning around flood and drought control or wildfire risk when used in combination with other  
408 models or data sources (Austin, 2023; Clifford et al., 2020; Jagannathan et al., 2021; Reclamation, 2016).

## 409 7 Conclusions

410 This paper presented an assessment of whether a standard resolution (~100 km grid) Earth system model is capable  
411 of producing information that water users typically employ in their decisions. Our motivation was to explore  
412 whether it is possible to reduce the need for intermediate downscaling, and to broaden the use of large model  
413 ensembles to quantify the influence of internal variability on localized decisions. We drew on the combined  
414 experience of the project team and workshop participants to identify potential metrics and familiar modes of  
415 visualization. This project used only CESM2 over the conterminous United States to develop example metrics  
416 that may be explored within other models and over other regions. CESM2 is unable to reproduce some metrics  
417 given the lack of topographical detail. A companion paper by Rugg et al. (2023) examines potential improvements  
418 to the subgrid-scale simulation of land processes to improve the representation of the hydrological cycle in  
419 mountainous regions.

420

421 We encourage others working in the decision space between climate data producers and users to be forthcoming  
422 about specific regions and reasons where model data are not credible, or where the model has particular  
423 weaknesses (such as the drizzle effect) that may be overcome with a different analysis approach.

424 For future model assessors, the following metrics were found to be salient for water users and were skillfully  
425 reproduced in many regions.

426

427 Rainfall:

- 428 ■ Number of wet days ( $\geq 1$ mm of rain) per year/season
- 429 ■ Mean precipitation on wet days
- 430 ■ Duration of the longest wet and dry spells per year
- 431 ■ Number of days with rain  $>$  95th percentile of current climate wet day totals
- 432 ■ Proportion of the annual total derived from days  $>$  95th percentile of wet day totals

433

434 Runoff (aggregated up to basin level, as a volume for 3- and 7-day averages):

- 435 ■ Annual maxima and minima

- 436 ■ Frequency of very high or very low flows (< 10% annual chance of occurring in the current climate)
- 437 ■ Proportion of averaged daily runoff to annual total

438

439 The work presented in this paper is a small step toward establishing greater usability of climate model output by  
 440 decision makers. The present evaluation is also only the first step in evaluating ESM performance. Additional  
 441 research is needed to support water managers placing these results and their uncertainty in the context of additional  
 442 observational data (such as remote sensing) that may already be available to them. Continued collaboration is  
 443 essential to improve the transfer of knowledge (e.g., data requirements, model assumptions, decision constraints)  
 444 between communities.

445

446 **Appendix A**

447

448 **Table A1: Hydro-meteorological responses used in water management decisions, and the specific metrics that have**  
 449 **potential for representation in ESMs. Metrics in bold are presented in this article.**

<b>Hydro-meteorological Responses</b>	<b>Typical Water Management Decision</b>	<b>Metric</b>	<b>Description</b>
Annual rainfall	Water supply and drought monitoring	<b>Total Precipitation (PRCPTOT)</b>	<b>Total annual precipitation measured as rainfall or snow water equivalent.</b>
Seasonal rainfall cycle	Seasonal water supply, reservoir operations management	<b>Number of Wet Days (NWD), Mean Wet Day Volume (WDV)</b>	<b>Frequency of days with <math>\geq 1</math>mm precipitation (NWD)</b> per month, season or year, <b>Mean precipitation on wet days</b> calculated from PRCPTOT/NWD
Rainfall extreme	Flood and stormwater management	95th percentile (Q95) <b>Number of very heavy rain days (N95)</b> Very heavy rain volume (P95) <b>Proportional contribution of very heavy rain (P95tot)</b>	Rainfall percentile threshold that is exceeded by 5% rain events per year on average, and calculated from wet days only <b>Frequency of days with rainfall exceeding Q95</b> Total rain falling on days exceeding Q95 <b>Proportional of annual total derived from very heavy rain</b> , calculated as P95/PRCPTOT
Rainfall extreme (dry)	Water supply planning and drought monitoring/planning including water rights and restrictions.	<b>Consecutive dry days (CDD)</b>	<b>Maximum duration of spell with consecutive days measuring &lt; 1 mm precipitation.</b>

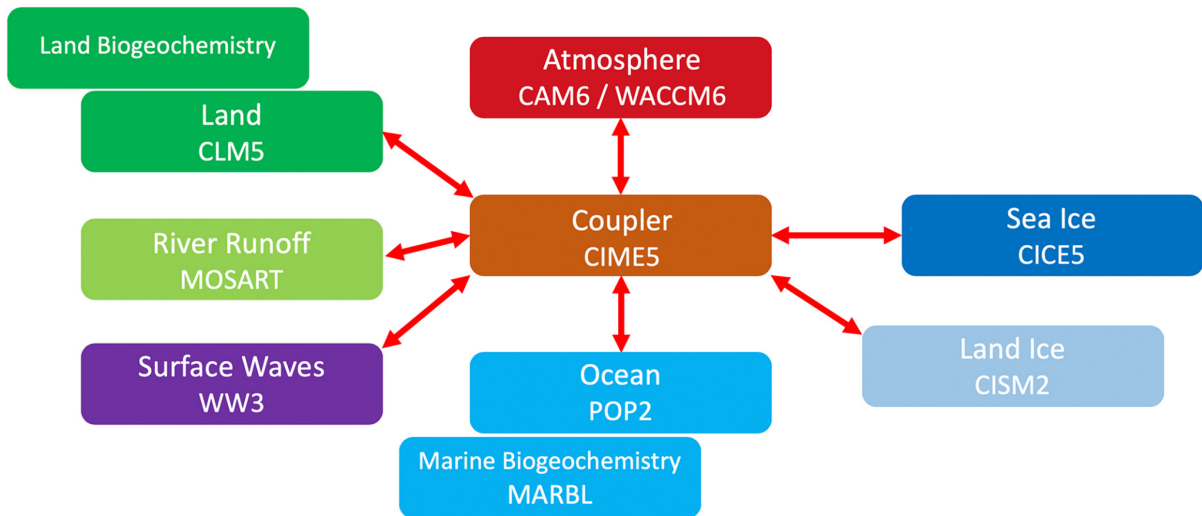
<b>Hydro-meteorological Responses</b>	<b>Typical Water Management Decision</b>	<b>Metric</b>	<b>Description</b>
Rainfall extreme (wet)	Stormwater management, water supply planning	<b>Consecutive wet days (CWD)</b>	<b>Maximum duration of spell with consecutive days measuring <math>\geq 1</math> mm precipitation.</b>
High streamflow	Reservoir management and flood control, water quality management and water supply management, including use of supplemental water supplies	<b>Annual maximum runoff (QMax)</b> Description (JMaxF) Description (HFD)	<b>Annual maximum daily volume of basin-wide runoff</b> Julian day of QMax/ day of the water year Duration of high flows
Low streamflow	Water supply management, assessment of water shortages with respect to seasonal demands	<b>Annual minimum runoff (QMin)</b> Description (JMinF) Description (LFD)	<b>Annual minimum daily volume of basin-wide runoff</b> Julian day of QMin/ day of the water year Duration of low flows
Streamflow	Water supply planning, water quality management, reservoir operations management, planning future investment needs	<b>7-day mean runoff (Q7)</b>	<b>Daily volume of basin-wide runoff averaged over 7 days.</b> Often presented as percentage of annual total volume of runoff or Pardé coefficient (Pardé, 1933)
Very low streamflow	Water quality management for discharge permits, conservation management, drought planning	<b>7-day “10-year” low runoff (7Q10)</b>	<b>7-day averaged basin-wide lowest volume of runoff with &lt;10% annual probability of occurring.</b> Estimated from Qmin series.
Very high flow	Flood management and planning, reservoir operations	<b>7-day “10-year” high runoff (7Q90)</b>	<b>7-day averaged basin-wide highest volume of runoff with &lt;10% annual probability of occurring.</b> Estimated from Qmax series.
Streamflow	Water supply planning, reservoir operations management	Central Tendency (CT) Description (Q <sub>25</sub> , Q <sub>50</sub> , Q <sub>75</sub> )	Day of the water year when the cumulative annual runoff exceeds 50% of the total annual runoff Annual quartiles of cumulative annual runoff estimated from daily streamflow.
Snowpack	Reservoir operations and flood management, water supply planning	Snow Water Equivalent (SWE) Maximum (SWEMax) SWEMax Date	Volume of peak snow water equivalent Day of the water year when peak SWE occurs

Hydro-meteorological Responses	Typical Water Management Decision	Metric	Description
		SWE Duration	Total length of snow accumulation and ablation
Snowmelt	Flood management and reservoir operations	Snowmelt onset	Day of water year of snowmelt onset

450

451 **Appendix B**

452 Schematic of the Community Earth System Model version 2 (CESM2) model components, reproduced from  
 453 Danabasoglu et al. (2020) Figure 1.

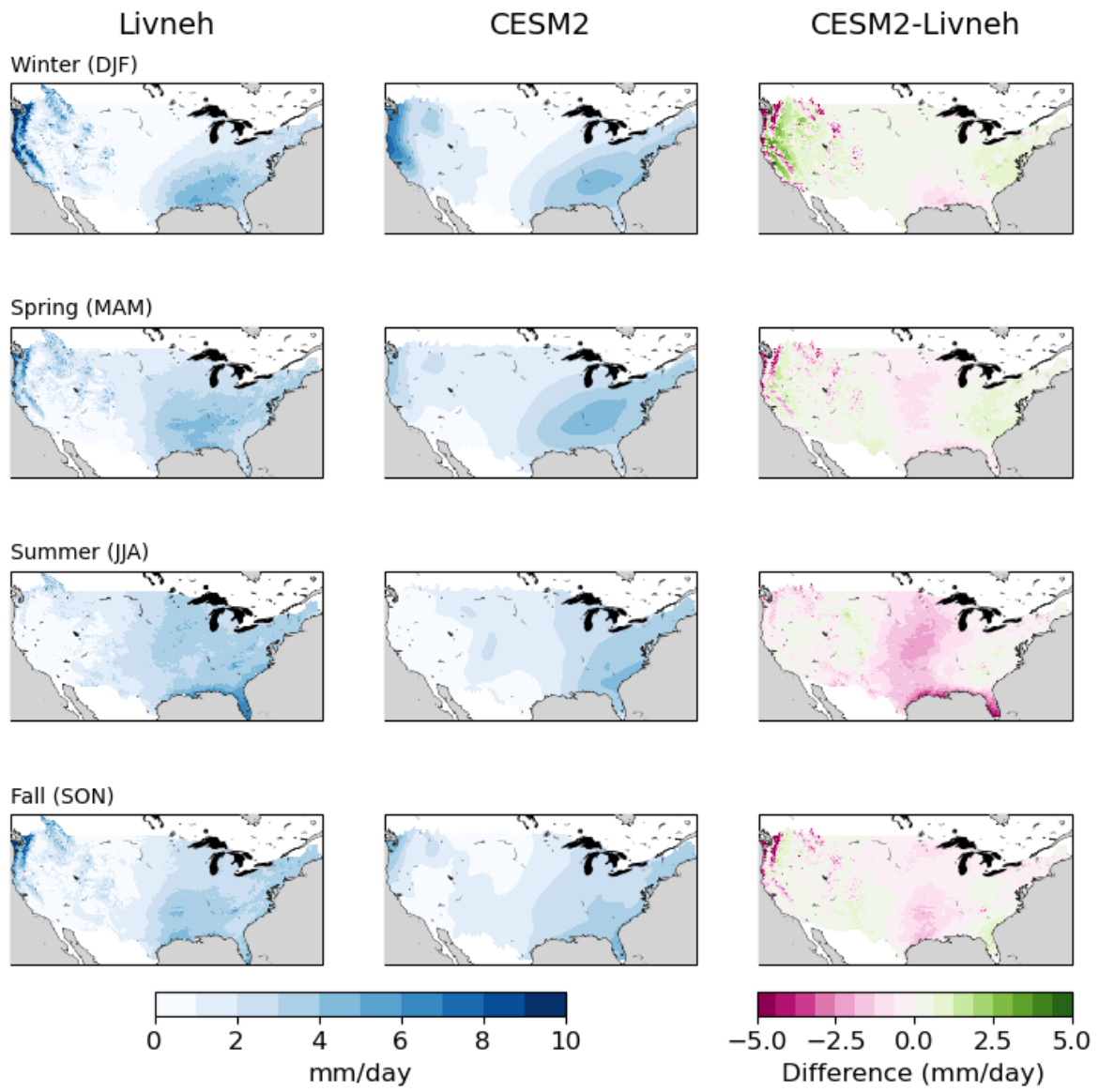


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456 **Appendix C**

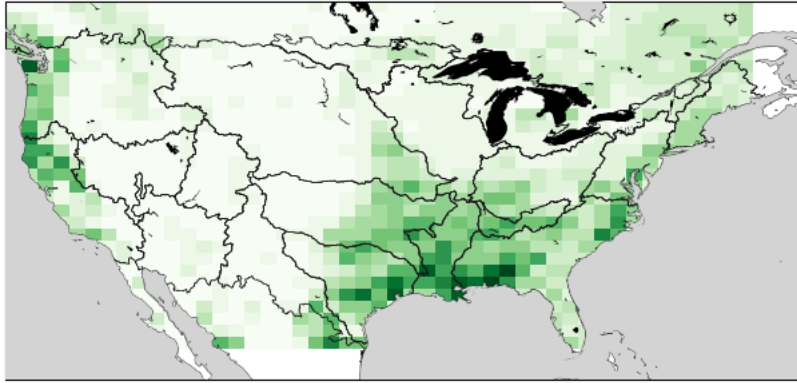
457 **Seasonal Mean Precipitation for Winter (top row), Spring (row 2), Summer (row 3) and Fall (bottom row) as shown in**  
458 **Livneh (left column) and CESM2 (middle column), and difference CESM2-Livneh (right column)**



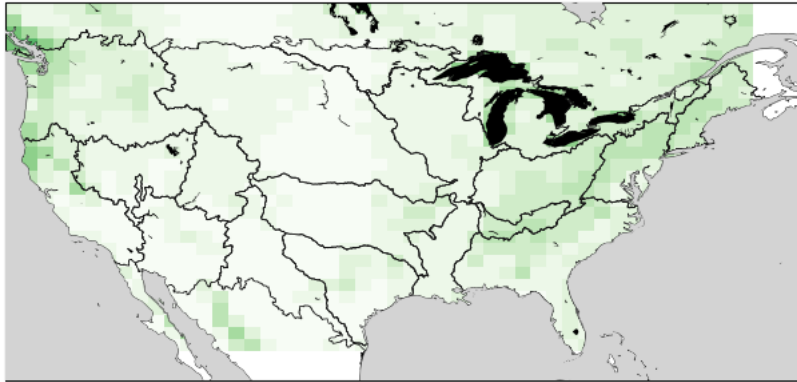
459

**Maximum Daily Runoff (1981-2005)**

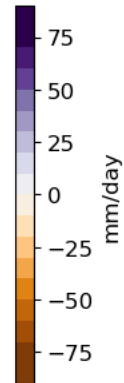
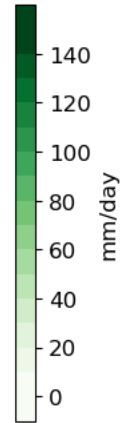
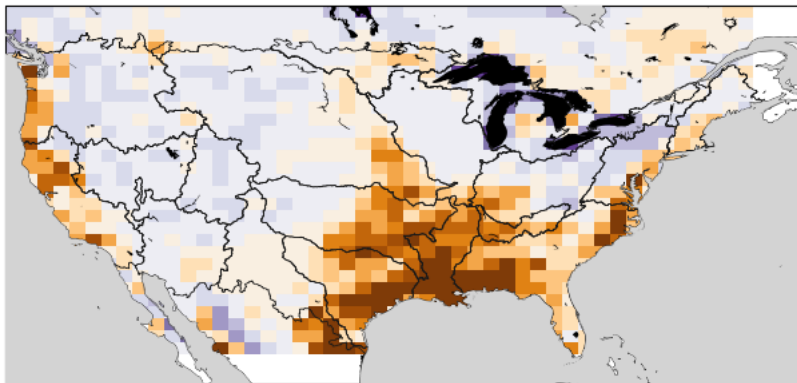
(a) Observations (Livneh-VIC)



(b) CESM2



(c) Difference (CESM2 - Livneh-VIC)



465 **Data availability**

466 All data generated for this study (e.g., CESM2 and Livneh-VIC calculated indices) along with Jupyter notebooks  
467 to recreate tables and figures are available in the repository [https://github.com/maritye/PSIF\\_water\\_avail](https://github.com/maritye/PSIF_water_avail)

468 **Author Contribution**

469 Conceptualization, M.T., J.R., E.G., A.N., A.W. and R.M.; Methodology, M.T., J.R., E.G.; Investigation, M.G.,  
470 M.T.; Data Curation, M.G., M.T.; Writing - original draft, M.T., A.R., and R.M.; Writing – reviewing and editing,  
471 M.T., J.R., E.G., A.N., A.W., R.M., A.R., F.L., C.B., and S.H.; Visualization, C.B., M.G. and M.T.; Supervision,  
472 J.R., E.G., A.N., F.L. and A.W.; Funding Acquisition, J.R., E.G., A.N., A.W., F.L., C.B., S.H. and M.T.; Project  
473 Administration J.R.

474 **Competing Interests**

475 The authors declare that they have no conflict of interest.

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