

# 1 Evaluating an Earth system model from a water manager 2 perspective

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## 14 Abstract

15 The large spatial scale of global Earth system models (ESM) is often cited as an obstacle to using the output by  
16 water resource managers in localized decisions. Recent advances in computing have improved the fidelity of  
17 hydrological responses in ESMs through increased connectivity between model components. However, the  
18 models are seldom evaluated for their ability to reproduce metrics that are important for practitioners, or present  
19 the results in a manner that resonates with the users. We draw on the combined experience of the author team  
20 and stakeholder workshop participants to identify salient water resource metrics and evaluate whether they are  
21 credibly reproduced over the conterminous U.S. by the Community Earth System Model v2 Large Ensemble  
22 (CESM2). We find that while the exact values may not match observations, aspects such as interannual  
23 variability can be reproduced by CESM2 for the mean wet day precipitation and length of dry spells. CESM2  
24 also captures the proportion of annual total precipitation that derives from the heaviest rain days in watersheds  
25 that are not snow-dominated. Aggregating the 7-day mean daily runoff to HUC2 watersheds also shows  
26 rain-dominated regions capture the timing and interannual variability in annual maximum and minimum flows.  
27 We conclude there is potential for far greater use of large ensemble ESMs, such as CESM2, in long-range water  
28 resource decisions to supplement high resolution regional projections.

## 29 1 Introduction

30 Water availability and water quality for human consumption, ecosystems, and agriculture are fundamental  
31 requirements, making pertinent assessments of future change crucial for adaptation planning (IPCC, 2022).  
32 Climate related changes in the hydrologic cycle will affect substantial portions of the world population, most  
33 directly through changes in water availability at or near the surface (Mankin et al., 2020; Sedláček and Knutti,  
34 2014). The information required by water resource managers for decision making is not readily available in a

35 relevant format, or at sufficient spatial or temporal resolutions from global Earth system models (ESM; e.g.,  
36 Ekström et al., 2018). We explore how the Community Earth System Model (CESM) represents the climatology  
37 of water availability, focussing on metrics that are familiar to decision makers in planning investment-scale  
38 decisions.

39 The inability of ESMs to explicitly resolve sub-grid scale (~100 km) processes is often cited as the limitation  
40 preventing direct model use in decision making. Literature from large organizations making infrastructure  
41 decisions (e.g., Brekke, 2011; Brekke et al., 2009; Reclamation, 2016, 2014) emphasize downscaling climate  
42 model data closer to the scale of the watersheds they manage. These additional modeling steps add complexity  
43 and may increase statistical errors (Clark et al., 2015; Ekström et al., 2018). Extracting useful and robust  
44 information directly from ESMs would reduce such errors if metrics most important to decision makers, such as  
45 the timing of peak flow, were known to be robustly represented.

46

47 There are many comprehensive examples of metrics used to evaluate climate and hydrological models (e.g.,  
48 Ekström et al., 2018; Mizukami et al., 2019; Wagener et al., 2022), and communicate the impacts of climate  
49 change (e.g., Reed et al., 2022), or to identify decision-relevant metrics (e.g., Bremer et al., 2020; Mach et al.,  
50 2020; Underwood et al., 2018; Vano et al., 2014). However, very few have examined whether user defined  
51 metrics can be reliably reproduced by ESMs (Mankin et al., 2020), and if further model development and scale  
52 reduction is warranted instead of improved communication (Pacchetti et al., 2021). Better communication may  
53 also reduce the temptation of some users to calculate “standard hydroclimate metrics” that are not supported by  
54 the climate model data (Ekström et al., 2018).

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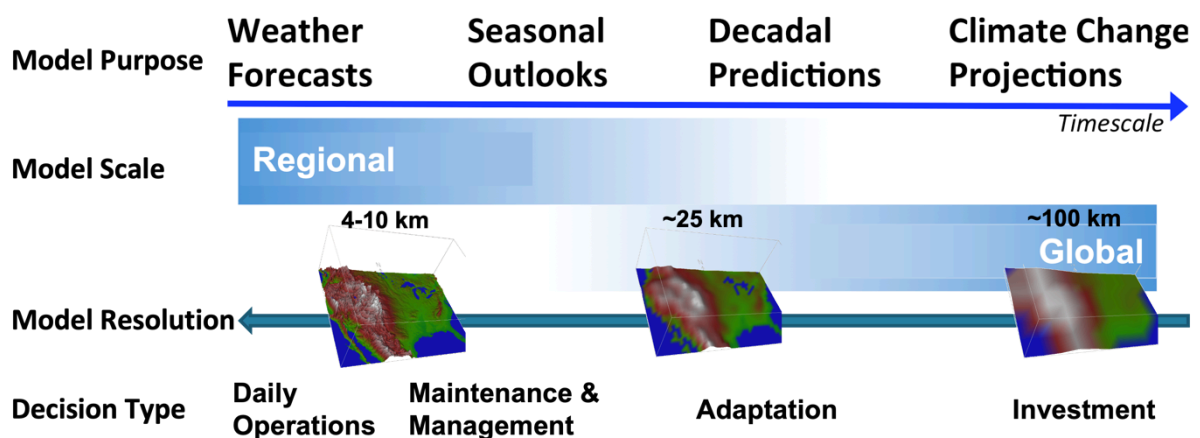
56 In contrast, climate model output can be rejected unnecessarily when simulated annual minima from freely  
57 running simulations do not “match” the sequence of observed low flows (Ekström et al., 2018; Moise et al.,  
58 2015). Similarly, the benefits of a range of projected outcomes from different climate models are not widely  
59 appreciated beyond the climate model community (Tebaldi and Knutti, 2007). Large ensembles from a single  
60 climate model initialized with a range of atmospheric and ocean conditions, such as the CESM2 Large  
61 Ensemble (LENS2; Rodgers et al., 2021), help to bound the uncertainty that derives from a naturally chaotic  
62 system. Averaged over the full ensemble, they give a better estimate of the model’s response to internal and  
63 external forcing (Deser et al., 2012) and enable assessments of the rarity of projected extremes. The additional  
64 analysis to identify structural (i.e. model formulation) and internal variability within regional climate models  
65 means that there are fewer large ensembles at a high resolution (Deser et al., 2020).

66

67 Since different decision makers have different priorities and time-scales of interest, Shepherd et al. (2018)  
68 recommended the development of climate storylines to communicate with those using climate data to make  
69 decisions. Informed by prior surveys of water managers (e.g., Brekke, 2011; Brekke et al., 2009; Cantor et al.,  
70 2018; Raff et al., 2013; Wood et al., 2021), Fig. 1 aims to map the different types of water decisions (e.g., Raff  
71 et al., 2013 Fig. 3) to the different scales of model resolution (Meehl et al., 2009 Fig. 2). Water managers make  
72 daily operational decisions (e.g., to control instantaneous river flow) with the aid of fine-scale weather and flood  
73 models (<4 km) that reliably represent convective and local weather scale processes even though their

74 predictability is relatively short lived (Yuan et al., 2019; far left side of Fig. 1). Larger watershed operations  
 75 (such as reservoir management or groundwater recharge; e.g., Regional Water Authority, 2019) depend on  
 76 seasonal outlooks (middle left of Fig. 1). Smaller adaptation and mitigation projects take place at the typical  
 77 policy or decadal prediction scale (i.e., 4-10 years; middle right of Fig. 1). Finally, major public investments and  
 78 inter-basin agreements occur at the same time scales as climate projections (30-100 years; far right of Fig. 1)  
 79 where persistent and relatively predictable synoptic and planetary scale processes are well represented in lower  
 80 resolution (~100 km) climate models (Phillips et al., 2020). While forecasts (seasonal or decadal) are  
 81 re-initialized from specific atmosphere, ocean or land states at regular time intervals, climate projections are run  
 82 freely from a variety of atmospheric and oceanic conditions that take several decades to converge to a mean  
 83 climatology. In considering the utility and useability of information directly from ESMs we focus on decisions  
 84 made over decadal to climate scales at larger spatial scales.

85



86

87 **Figure 1: Mapping the temporal and spatial scales of models to the timeframes for water management decisions.**

88 Given that ESMs have advanced immeasurably in the recent decade, it is time to re-evaluate whether their direct  
 89 output can support decision makers. Such an evaluation needs to focus on how well the models can reproduce  
 90 metrics used by decision makers, and whether the results are credible (Briley et al., 2020; Jagannathan et al.,  
 91 2021). Here we evaluate the credibility of one ESM in generating metrics known to be salient for water  
 92 management decisions; specifically, decisions for water management infrastructure project investments.

93

94 The motivation for this paper is to identify:

- 95 • a set of water availability metrics that resonates with decision makers and supports their
- 96 investment-scale decisions;
- 97 • how well CESM2 represents the climatology and recent observed behaviors of those metrics; and
- 98 • how such metrics are projected to change.

99

100 This paper builds off a decade of collaboration between scientists at the National Center for Atmospheric  
 101 Research (NCAR) and US water agencies that led to a virtual workshop (Tye, 2023), and presents a test case for  
 102 improved communication with water resources decision makers. The focus is on the Conterminous United  
 103 States (CONUS) to match the interest of workshop participants.

## 104 2 Climate Information Needs from Prior Research

105 Information needs vary greatly, from 5-minute rainfall totals at a point (ASCE, 2006), to basin-wide measures of  
106 annual minimum and maximum total runoff. Water management decision metrics can be grouped into similar  
107 types such as timing, frequency, magnitude, extreme values, variability, and duration of events (Ekström et al.,  
108 2018). While some aspects of timing, magnitude, or variability can be reliably reproduced by ESMs (e.g., Deser  
109 et al., 2020; Tebaldi and Knutti, 2007), others such as short duration extremes are less reliable.

110

111 Methods of evaluation and data use also differ. For instance, Clifford et al. (2020) reported that predicting  
112 general changes in the frequency of extreme precipitation events is more useful for future planning than the  
113 precise prediction of mean values evaluated by model developers. Lehner et al. (2019) emphasized that models  
114 need to be evaluated for their ability to reproduce sensitivities (e.g., streamflow changes in response to  
115 temperature and precipitation changes) in addition to mean states. However, metrics that are meaningful for  
116 evaluating a model's capabilities (e.g., the ratio of precipitation to runoff) are less valuable for management  
117 decisions (Lehner et al., 2019; McMillan, 2021; Mizukami et al., 2019). When reporting results, water managers  
118 are more familiar with the 'water year', rather than the calendar year, to capture the full annual hydrological  
119 cycle (Ekström et al., 2018). While the use of water years is a nuance that does not add substantial value to  
120 climate model assessments, communication with decision makers is improved by presenting data in a familiar  
121 format (Briley et al., 2020).

122

123 There is a need for information at the local scale that is unlikely to be met directly by raw outputs from the  
124 current generation of ESM. But better communication of the variability in future daily precipitation and  
125 associated runoff can add value to the detailed models by bringing in the added statistical context and  
126 perspective of the large ensembles. Thus, we believe that ESMs can produce useful information about  
127 hydro-meteorological extremes when presented at different spatial or temporal scales, and offer the benefits of  
128 large climate model ensembles to constrain future impact uncertainty.

129

130 Appendix A summarizes potential hydrological metrics used in water management decisions (Jagannathan et al.,  
131 2021) or statistical assessments of extremes (Zhang et al., 2011), and model evaluations (Phillips et al., 2020).  
132 Metrics in bold are presented in this paper. We only considered a simplistic measure of meteorological drought  
133 (absence of rain) in the current work, as drought is sensitive to the definition (Bachmair et al., 2016) and local  
134 conditions (Mukherjee et al., 2018), and so not suited to a generalized assessment. Similarly, snow measures are  
135 not included in this assessment in part due to limited availability of high-quality, long-duration,  
136 quality-controlled, observational data (McCrary et al., 2017); and partly due to the biases in snow distribution  
137 arising from the smoothed topography in GCMs (McCrary et al., 2022).

## 138 3 Data and Methods

### 139 3.1 Climate Model Data

140 CESM2 (Danabasoglu et al., 2020) is a fully coupled global model that simulates the Earth's climate system  
141 through interactive models for atmosphere, ocean, land, sea-ice, river runoff, and land-ice. Variables considered  
142 in this project are taken from the Community Atmosphere Model version 6 (CAM6) and the Community Land  
143 Model version 5.0 (CLM5; Lawrence et al., 2019) and are part of the default model outputs. A schematic of the  
144 model components is included in Appendix B. This project uses daily values scaled up to annual (e.g., annual  
145 maximum daily precipitation) on a ~1 degree resolution grid. Data were extracted over the CONUS from 10  
146 ensemble members of LENS2 (Rodgers et al., 2021) for model validation in the current era (1981-2010).

### 147 3.2 Observations

148 Gridded daily observations of precipitation at 1/16° horizontal resolution (~6 km) were obtained from the  
149 Livneh et al. (2013) dataset covering CONUS and southern Canada for the control period (1981-2010), hereafter  
150 referred to as "Livneh". While the time adjustment in the Livneh dataset results in an underestimation of the  
151 most extreme daily precipitation totals and resultant runoff and flood potential (Pierce et al. 2021), we are also  
152 interested in precipitation and runoff minima. As a result we did not employ the updated gridded observations  
153 (Pierce et al. 2021).

154

155 Livneh daily temperature maxima and minima, and precipitation were used to force the Variable Infiltration  
156 Capacity Model (VIC; Liang et al., 1994) version 4.1.2 to obtain runoff estimates for years 1980-2005 as  
157 evaluated in Livneh et al. (2013). Hereafter referred to as "Livneh-VIC".

158



159

160 Figure 2: HUC 2 regions used in data validation and analysis. Regions defined by USGS (2013): Region 01 New  
161 England (NE); Region 02 Mid-Atlantic (MA); Region 03 South Atlantic-Gulf (SA); Region 04 Great Lakes (GL);  
162 Region 05 Ohio (OH); Region 06 Tennessee (TN); Region 07 Upper Mississippi (UM); Region 08 Lower Mississippi  
163 (LM); Region 09 Souris-Red-Rainy (RR); Region 10 Missouri (MR); Region 11 Arkansas-White-Red (ARK); Region  
164 12 Texas-Gulf (GUL); Region 13 Rio Grande (RIO); Region 14 Upper Colorado (UC); Region 15 Lower Colorado

165 (LCO); Region 16 Great Basin (GB); Region 17 Pacific Northwest (PN);  
166 Region 18 California (CA)

### 167 3.3 Methods

168 All analyses were carried out using the North American water year (1 October to 30 September) to facilitate  
169 later communication.

#### 170 3.3.1 Remapping

171 For ease of comparison, model output were re-gridded using a conservative second-order remapping (Jones,  
172 1999) to place both datasets on the same scale grid and assess anomalies. Data were also calculated as areal  
173 averages or totals over the 2-digit Hydrological Unit Code (HUC2) regions (Seaber et al., 1987). HUC2 basins  
174 represent 18 watersheds, covering areas ranging from 41,000 mi<sup>2</sup> (~105,000 km<sup>2</sup>; Tennessee) to 520,960mi<sup>2</sup>  
175 (1,350,000 km<sup>2</sup>; Missouri), shown in Fig. 2. While the scale of HUC2 regions may be large for some local  
176 decision-makers, it is also a more appropriate and conservative scale to compare to ESMs as demonstrated by  
177 Lehner et al. (2019).

#### 178 3.3.2 Percentile-based thresholds

179 The threshold for very heavy rain days (Q95) was calculated at each individual grid cell using only days with  $\geq$   
180 1 mm rain (“wet days”). Thresholds were derived empirically for each model ensemble member, with the  
181 ensemble mean threshold (Q95) used to identify the days per year exceeding the threshold (N95) and total  
182 annual rainfall from those days (P95).

183 Runoff was aggregated over each HUC2 watershed and multiplied by the respective area to generate total  
184 volume per day. Volume per day was then converted to measurements more familiar to users, such as acre feet  
185 per day or cubic meters per second. Daily time series of total volumetric runoff had a 7-day running mean  
186 smoother applied, then annual maximum, minimum and mean values were extracted. The highest and lowest  
187 7-day average runoff expected once per decade (7Q90, 7Q10) were estimated empirically from the 25 ranked  
188 values of of annual maxima and minima per watershed.

### 189 4 Model Evaluation

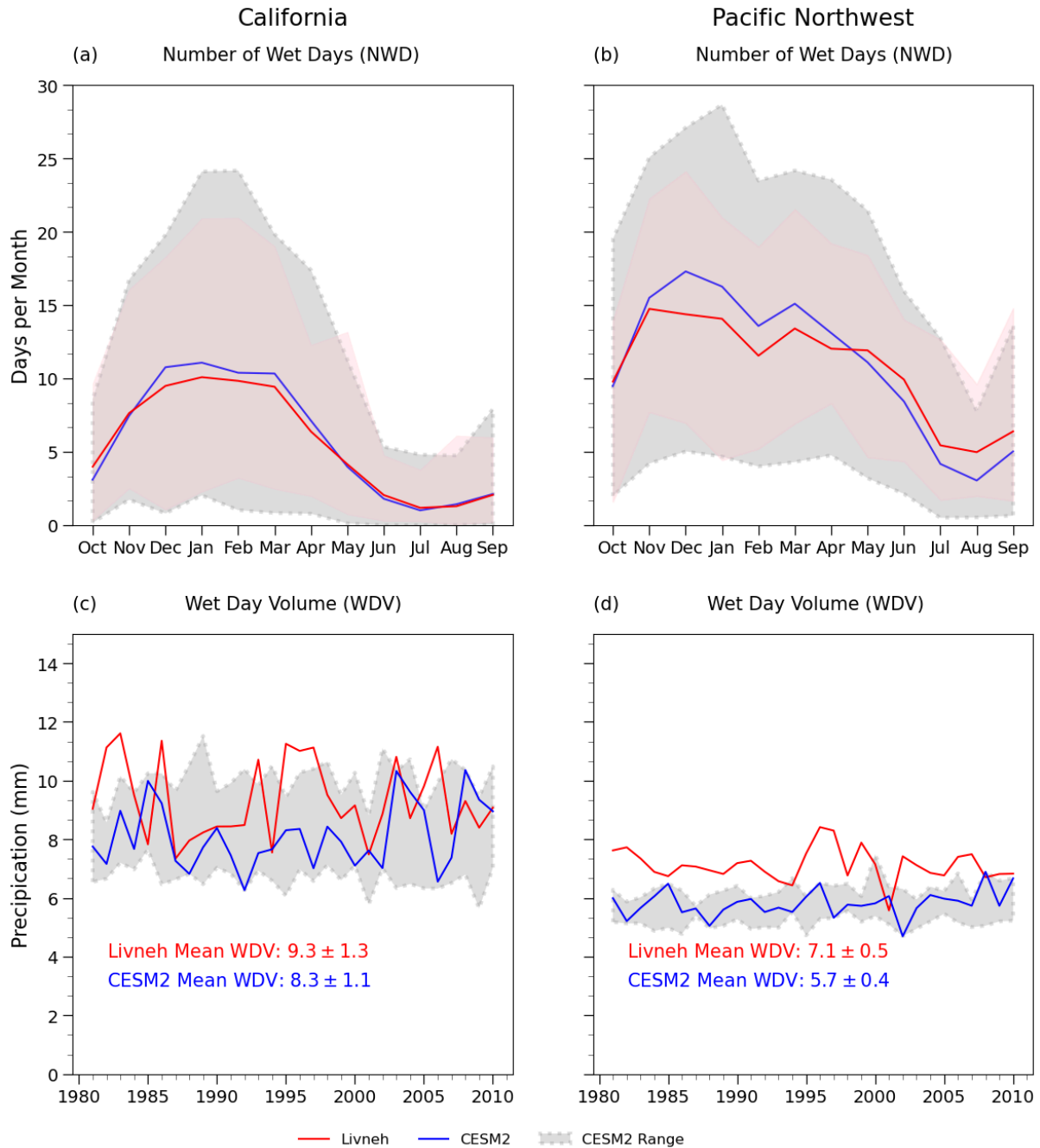
190 The metrics used to evaluate CESM2’s ability to reproduce large scale features and physical behaviors (e.g.,  
191 Danabasoglu and Lamarque, 2021 and the associated Special Issue) are not necessarily those employed by  
192 decision makers. ESMs are designed to represent large-scale atmospheric processes and fluxes not specific local  
193 responses (Gettelman and Rood, 2016), but this design assumption may not be sufficiently well communicated  
194 to decision makers. The purpose of our evaluation is to establish whether CESM2 output is also fit for local  
195 decision purposes, or if the breadth of information from ESM ensembles remains unsuitable for immediate use  
196 in targeted water management decisions.

#### 197 4.1 Rainfall metrics

198 While broad spatial patterns of seasonal mean daily rainfall are reproduced well (Danabasoglu et al., 2020; Feng  
199 et al., 2020; Simpson et al., 2022), CESM2 fails to capture details over high topography, and overestimates  
200 summer precipitation where convective extremes dominate summer rainfall (Appendix B). The seasonal mean  
201 precipitation also fails to capture some important watershed-level processes, such as the seasonal variability in  
202 the number of days with precipitation and the associated intensity.

203

204 Estimates of mean annual rainfall on wet days, or wet day volume, are in broad agreement between Livneh and  
205 CESM2 output. Figure 3 shows an example of the mean number of wet days per month (NWD), and mean wet  
206 day volume (WDV) averaged over the Mid Atlantic and Pacific Northwest. While CESM2 represents the NWD  
207 annual cycle very well in regions such as California (Fig. 3a, 3c) and the Pacific Northwest (Fig. 3b, 3d), it does  
208 not capture NWD in many central and snow dominated regions. This is likely due to the smoother topography of  
209 CESM2 missing the influence of orographic uplift, and large spatial scale missing sub-grid scale convective  
210 systems (e.g., over the Central Plains).



211

212 **Figure 3: Average number of wet days per month (a, b) and interannual variability in mean annual precipitation on**  
 213 **wet days for Livneh climatological mean (red) with interannual spread (pink) and CESM2 mean (blue) with**  
 214 **interannual and ensemble spread (gray); and (c,d) between 1981-2010 for observations derived from Livneh (red)**  
 215 **and CESM2 ensemble mean (blue) and spread (gray) in (a,c) Region 18 California (CA);**  
 216 **and (b,d) Region 17 Pacific Northwest (PN).**

217 The annual variability in WDV, both year-to-year variations as well as the overall range of minima and maxima,  
 218 is well captured by each of the model members for the different HUC2 regions, even if the absolute values do  
 219 not match (Fig. 3 c,d). As expected, the specifics of which years have high or low values of WDV are not the  
 220 same for each ensemble member (i.e. demonstrating internal variability). As a result, the ensemble mean value  
 221 of WDV (blue) does not reflect the same year-to-year variability as the observations. Decision makers expressed  
 222 that the interannual variability demonstrated by each model member is more valuable to demonstrate the



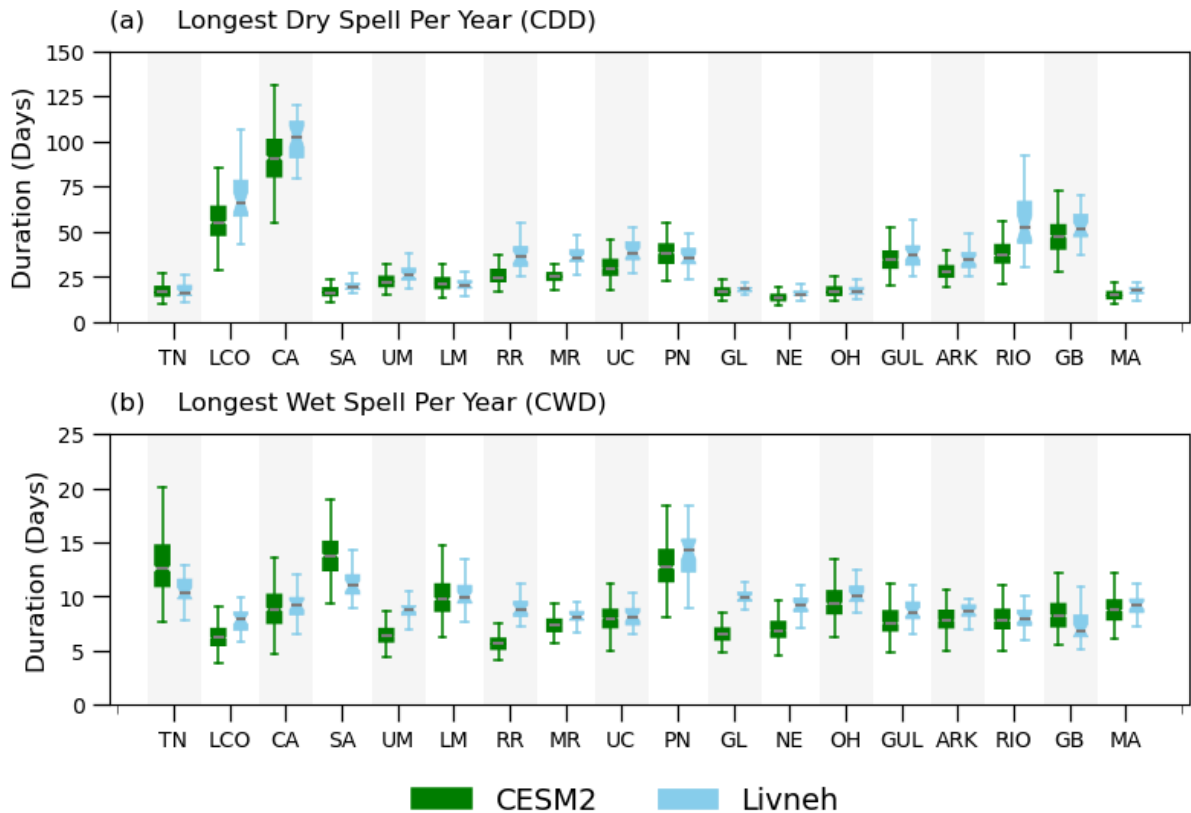
223 credibility of the data than the ensemble mean (Tye, 2023). We recommend that the full range of values of each  
224 metric (i.e. after computation for each ensemble member individually) are communicated in addition to the  
225 climatological means to help bound uncertainty around decisions (Wilby et al., 2021).

226

227 The magnitude of interannual variability in WDV (i.e., the absolute differences between the maximum and  
228 minimum values in each member time series) is typically within 10% of observations in all regions as illustrated  
229 for two regions in Fig. 3. Exceptions are the Lower Colorado, South Atlantic-Gulf and Upper Mississippi where  
230 the simulated distributions are too narrow. Many different sources of error may contribute to this discrepancy,  
231 such as the inability to resolve convective precipitation (Chen et al., 2021) in addition to elevation changes not  
232 captured by the coarse model resolution, or the “drizzle effect” that is common in GCMs (Chen et al., 1996; Dai,  
233 2006).

234

235 CESM2 captures the longest spells of consecutive dry days per year (CDD; Fig. 4a) and consecutive wet days  
236 per year (CWD; Fig. 4b), and their variability. Many regions capture both the interannual variability and the  
237 climatological mean duration of CWD, particularly those regions that are subject to large-scale synoptic systems  
238 (e.g., Pacific Northwest, Mid Atlantic-Gulf, California). Several regions either overestimate (South  
239 Atlantic-Gulf) or underestimate (Great Lakes, Souris-Red-Rainy) the absolute durations of the longest wet  
240 spells, but do reflect the magnitude of interannual variability. The exception is Tennessee, where both  
241 interannual variability and mean CWD are overestimated. At the grid scale, broad spatial patterns of CWD are  
242 correct but the finer atmospheric processes arising from topographic features are incorrect, as expected from the  
243 coarse model resolution. A similar pattern is present in CDD, except that some drier regions with CDD >30 days  
244 do not capture the full range of interannual variability (Souris-Red Rainy, Missouri, Rio Grande). As GCMs  
245 have a tendency to produce drizzle, adjusting for a higher wet day threshold (e.g., 2 mm) might improve dry  
246 spell representation in those regions. It is also important to communicate such model sensitivities to users more  
247 effectively.



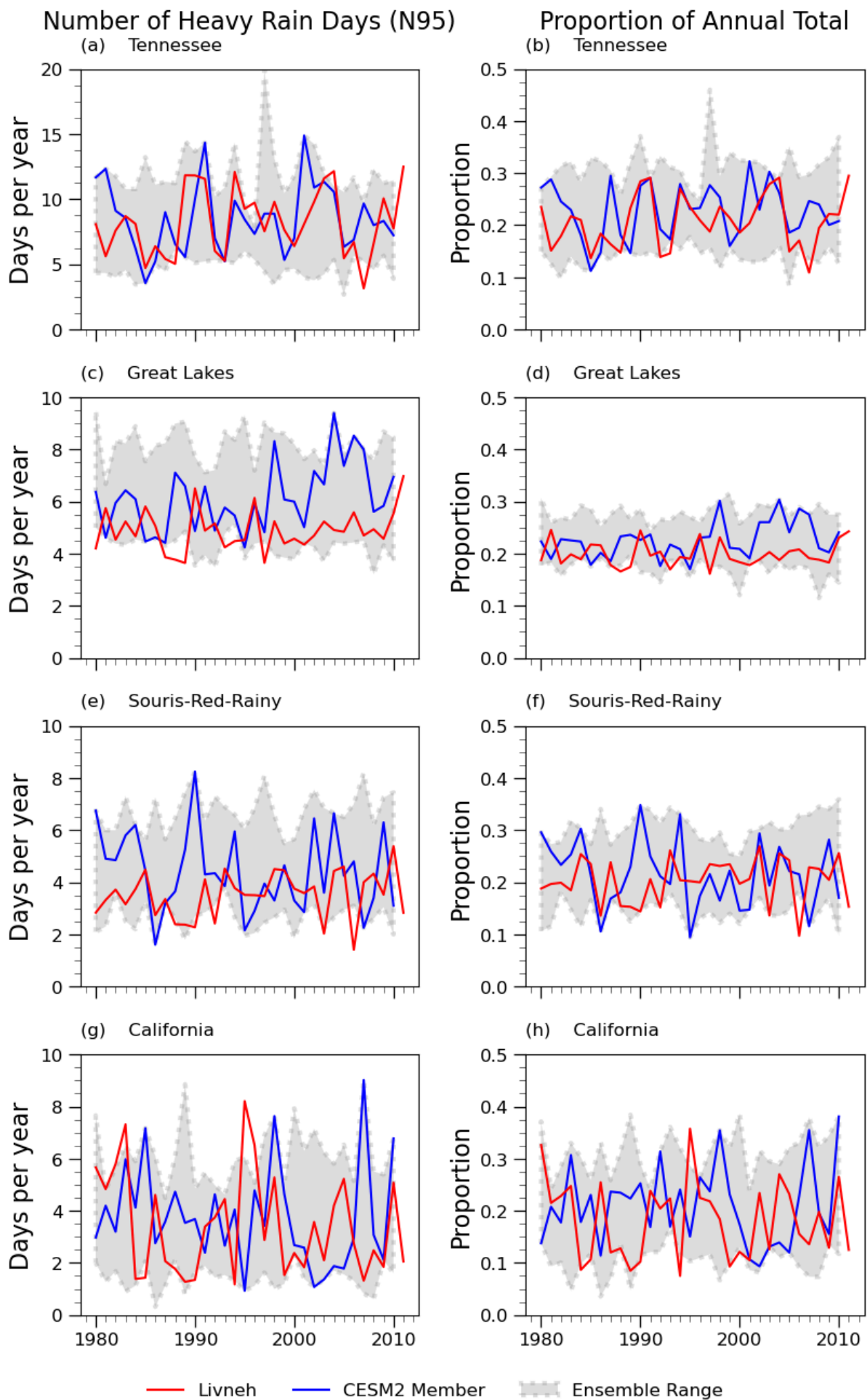
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249 **Figure 4: a) Longest duration per year of consecutive days <1 mm rain (longest dry spell) for Livneh over all years**  
 250 **(green) and CESM2 ensemble range over all years (blue) for all HUC2 regions; and b) Longest duration per year of**  
 251 **consecutive days with  $\geq 1$  mm rain (longest wet spell). Regional Acronyms defined in Fig. 2.**

252 The thresholds for heavy and very heavy rain days (P95, P99) are defined individually for Livneh and CESM2  
 253 both to understand whether the intensity of more extreme rainfall is captured, and to evaluate model behavior. A  
 254 comparison of the thresholds reflects the considerable improvements in modeling capabilities in recent years  
 255 (Gettelman et al., 2022). For instance, earlier versions of CESM underestimated extreme precipitation intensity  
 256 by 10-30 mm/day east of the Rockies, and overestimated intensity by 5-10 mm/day to the west (Gervais et al.,  
 257 2014). We found CESM2 still underestimates the most extreme rainfall, but that errors have approximately  
 258 halved. As these differences are still inadequate for many engineering and major infrastructure decisions  
 259 (Wright et al., 2019), we focus on CESM2's ability to capture the relative contributions of P95 and P99 to the  
 260 annual total and the interannual variability in their frequency. A result with considerable useability is the  
 261 proportion of annual total precipitation derived from the heaviest rain days, or "Proportional Contribution of  
 262 Extreme Days" (P95Tot). This proportion and its interannual variability is well represented by CESM2 at the  
 263 HUC2 scale and has shown to be skillful in other models (Tebaldi et al., 2021).

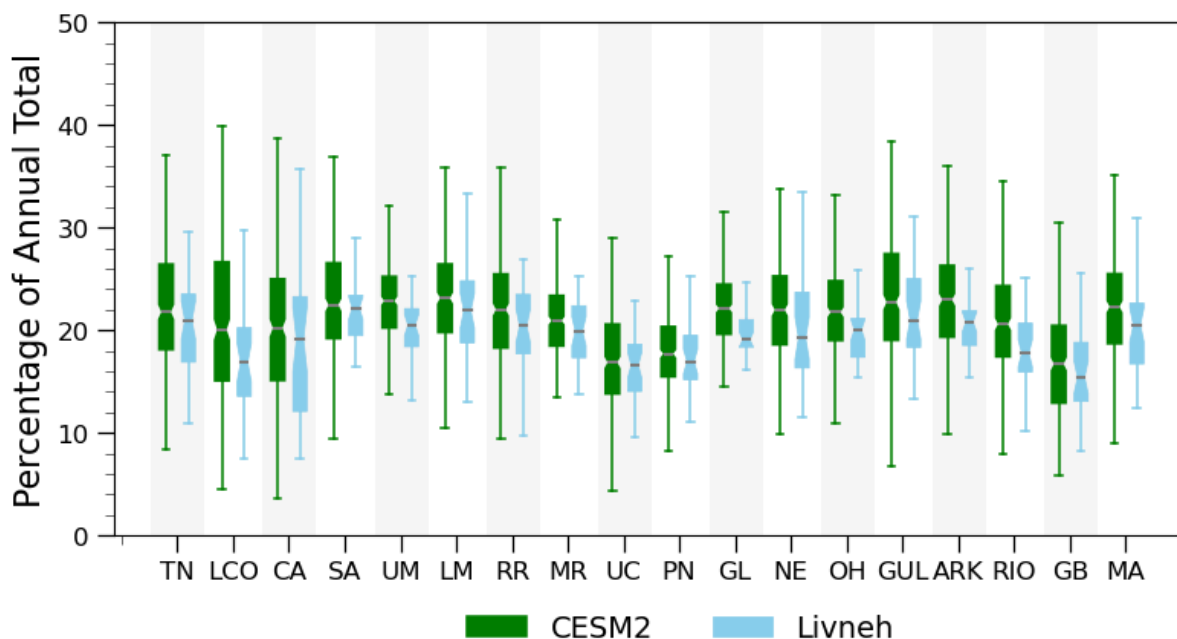
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265 The interannual variability in the frequency (N95) and intensity of extreme rainfall, as represented by P95Tot,  
 266 are illustrated in Fig. 5 and 6. In several HUC2 regions the simulations report more frequent events, and  
 267 proportionally higher totals (e.g., Great Lakes, Rio-Grande, Missouri, Upper Colorado and Lower Colorado).  
 268 Overall, there is good agreement between Livneh and CESM2, identifying an opportunity to inform local  
 269 decisions from large scale ESMs.



— Livneh    — CESM2 Member    Ensemble Range

271 Figure 5: a, c, e, g) Number of very heavy rain days per year; and b, d, f, h) total rain from very heavy rain days as a  
 272 proportion of the annual total for a, b) Tennessee (TN); c,d) Great Lakes (GL); e,f) Souris-Red-Rainy (RR); and g,h)  
 273 California (CA) HUC2 regions. Observations in red; CESM2 ensemble spread in gray, single ensemble member in  
 274 blue.



275  
 276 Figure 6: Box plots of the interannual range of contributions to annual total rainfall from very heavy days (P95Tot)  
 277 shown as percentages for: Observations (light blue), and ensemble range for CESM2 (green) for all HUC2 regions.  
 278 Boxes are bound by the interquartile range, black lines indicate the median, notches indicate the degree of spread  
 279 from the median and bars extend to the full data range.

## 280 4.2 Runoff metrics

281 Runoff estimates are taken from the individual components of surface and subsurface runoff generated within  
 282 CLM5 (Lawrence et al., 2019) and compared to the Livneh forced VIC runoff (“Livneh-VIC”).

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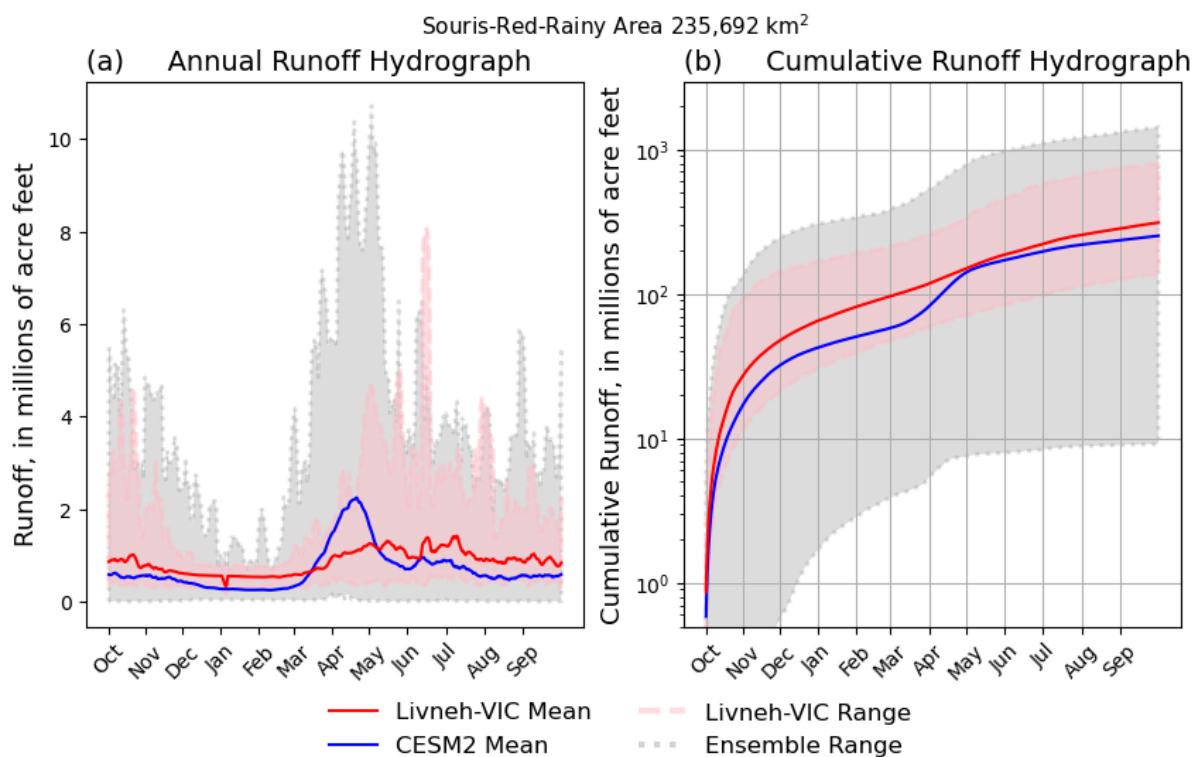
284 Assessing the skill of runoff in large-scale models is complicated by many factors, including the mismatch of  
 285 scales between in-channel flow ( $\sim 1\text{-}10^2$  m) and the grid scale ( $\sim 10^5$  m). Thus, metrics of climate model runoff  
 286 should be selected carefully and the runoff should be aggregated or combined with other metrics, rather than  
 287 used directly (Lehner et al., 2019). Appendix C demonstrates the discrepancies between the grid-scale  
 288 representation of runoff from Livneh-VIC and CESM2. The large discrepancies arise from different processes  
 289 that are not captured adequately, such as groundwater, topography, and associated snow ablation and melt, in  
 290 addition to meteorological biases.

291

292 However, water management decisions are made over watersheds in units such as acre-feet<sup>1</sup> or cubic meters,  
 293 while model data are output as a depth of runoff over each grid cell (e.g., mm/day per km<sup>2</sup>). We aggregated the  
 294 7-day running mean daily runoff (Q7) within each HUC2 region to generate Q7 time series in each basin. Fig.  
 295 7a illustrates the 25-year mean seasonal cycle for Livneh-VIC in red and CESM2 in blue, and the full range of

296 <sup>1</sup> 1 Acre-foot is the volume of water it would take to cover 1 acre of land to a depth of 1 foot. Equal to 325,852  
 297 gallons or 1,233 m<sup>3</sup> (USGS Water Science).

298 values over all years and ensemble members for the Souris-Red-Rainy basin (HUC Region 9). Data are  
 299 presented in millions of acre feet, to align with decision maker needs. The minimum simulated Q7 in any year  
 300 considerably underestimates the lowest flows in this region compared to Livneh-VIC. In contrast, the largest  
 301 total runoff volume is overestimated and peaks too early in the water year. Figure 7b plots the same information  
 302 as the cumulative runoff volume from the start of the water year, highlighting that the lowest runoff volume is  
 303 underestimated by a factor of ten. Low runoff volumes were typically underestimated in smaller regions (e.g.,  
 304 NE, TN). High runoff volumes were only underestimated in three regions (LM, ARK, GUL) and considerably  
 305 overestimated in seven regions. Snow-dominated regions perform particularly poorly for both QMax and QMin  
 306 as snowpack and the timing of associated runoff are not well simulated. Transitional regions that straddle both  
 307 snow- and rain-dominated hydrology also fail to capture QMax, but better estimate Qmin (not shown). Only the  
 308 South Atlantic region reproduces both QMax and QMin.



309

310 **Figure 7: Interannual variability in runoff in Souris Red Rainy Region for a) the mean seasonal cycle; and b) the**  
 311 **cumulative watershed runoff over the water year. Livneh-VIC climatological mean in red, range of all years in pink;**  
 312 **Cesium2 ensemble mean in blue and ensemble range in gray. Figure highlights the underestimation of the lowest**  
 313 **runoff volume by Cesium2 by a factor of ten.**

314

315 We explored the relationship between the highest and total annual runoff (QMax/QTot), and lowest and total  
 316 annual runoff (QMin/QTot). Some regions performed well for QMax/QTot, others performed better for  
 317 QMin/QTot but there was no consistent relationship that could be utilized by decision makers.

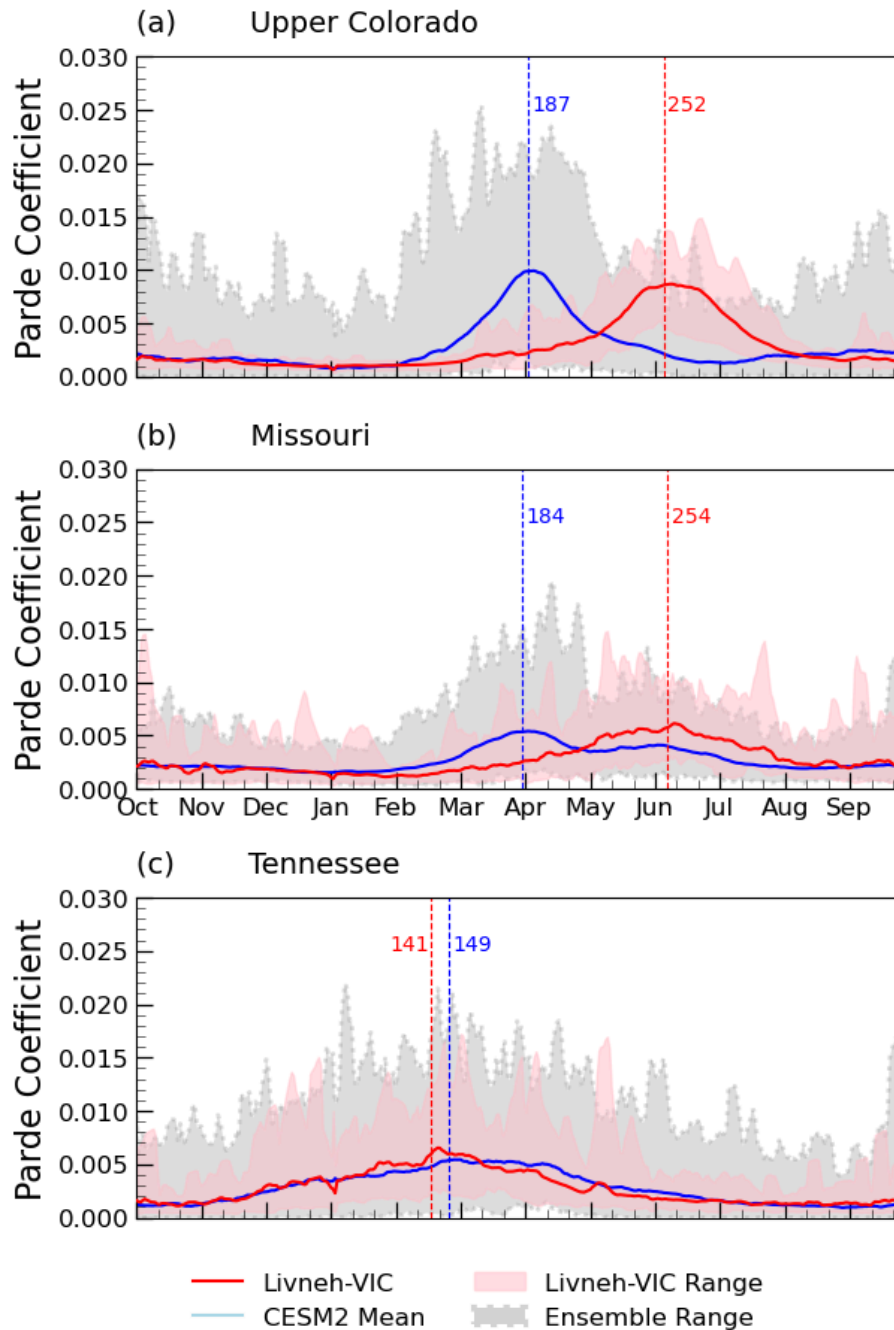
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319 Participants at the NCAR workshop (Tye, 2023) emphasized that the exact numbers produced by climate models  
 320 are not very important for future decisions. Others have also emphasized the importance of well-represented  
 321 processes in the model (Reed et al., 2022) and correlations with known experiences (Mach et al., 2020;

322 Shepherd et al., 2018). Focussing on fidelity to the historical climate exaggerates the importance of model  
323 performance instead of robustness to different conditions without ensuring that model predictions are useful or  
324 reliable (Brunner et al., 2021; Wagener et al., 2022). Runoff estimates in transitional catchments may be  
325 inadequate in the current climate but plausible in the future, if the model reproduces rain-dominated  
326 hydrological processes (McMillan, 2021).

327

328 Climatological mean runoff cycles are estimated from Pardé coefficients — calculated as  $Q7/Q_{Tot}$  on each  
329 calendar day — a dimensionless value that enables comparison across regions. Figure 8 depicts the mean  
330 seasonal cycle for representative snow-dominated (Upper Colorado), transitional (Missouri) and rain-dominated  
331 (Tennessee) regions, demonstrating how an imperfect representation of snow in the Upper Colorado results in  
332 CESM2 peak runoff occurring two months earlier than Livneh-VIC (Fig. 8a). The runoff regimes display very  
333 different seasonal characteristics, with CESM2 having a “mid late spring” runoff regime rather than  
334 Livneh-VIC’s “extreme early summer” regime (Fig. 8a; Haines et al., 1988). Peak runoff is also too early in  
335 transitional regions, but closer to Livneh-VIC than in snow-dominated regions (Fig. 8b). Rain-dominated  
336 regions capture both the timing of  $Q_{Max}$  and overall seasonal hydrograph shape (Fig. 8c).



337

338 Figure 8 : Seasonal patterns of runoff for HUC2 regions a) Upper Colorado (UC); b) Tennessee (TN);  
 339 and c) Missouri (MR). Constructed from normalized series of the ratio of 7-day mean runoff to the mean annual  
 340 total. Livneh-VIC runoff climatological mean (red), climatological range (pink), CISM2 ensemble mean (blue) and  
 341 ensemble range (gray with dashed border). Vertical lines indicate the mean date of peak runoff with number of days  
 342 since the start of the water year.

343 7Q10 and 7Q90 are estimated empirically from annual minima and maxima as occurring once per decade.  
 344 Projected changes in the frequency of very low or very high runoff volumes are deemed credible where CISM2  
 345 replicates the standard deviation of annual minima and maxima according to a  $\chi^2$  test at the 5% significance  
 346 level. Table 1 reports CISM2 and Livneh-VIC regional estimates of 7Q10 and 7Q90 and standard deviations of  
 347 the annual maxima and minima; values in bold indicate where estimates are statistically similar. It should be

348 noted that the values in Table 1 have  $\leq 10\%$  of occurring in any year, and so represent the tails of the runoff  
 349 distribution.

350

351 **Table 1 : Very low (7Q10) and very high (7Q90) regional runoff, and standard deviation in regional annual minima ( $\sigma$**   
 352 **QMin) and annual maxima ( $\sigma$  QMax) for Livneh and CESM2. Values in bold indicate where CESM2 and**  
 353 **Livneh-VIC regional runoff are statistically similar according to a  $\chi^2$  test.**

Region		Livneh-VIC				CESM2			
		7Q10	7Q90	$\sigma$ QMin	$\sigma$ QMax	7Q10	7Q90	$\sigma$ QMin	$\sigma$ QMax
NE	1	4.1	132.4	1.3	25.5	8.6	215.1	4.7	39.9
MA	2	<b>6.9</b>	103.5	<b>2.5</b>	25.7	<b>7.4</b>	220.7	<b>3.6</b>	47.9
SA	3	<b>21.1</b>	<b>240.4</b>	<b>8.4</b>	<b>50.7</b>	<b>20.5</b>	<b>258.6</b>	<b>11.9</b>	<b>45.8</b>
GL	4	<b>6.9</b>	122.5	<b>2.2</b>	23.8	<b>7.8</b>	331.0	<b>4.3</b>	58.0
OH	5	<b>7.8</b>	187.6	<b>2.3</b>	53.0	<b>9.4</b>	260.9	<b>4.5</b>	56.4
TN	6	2.1	<b>90.5</b>	0.8	<b>23.1</b>	0	<b>98.7</b>	0.3	<b>21.7</b>
UM	7	2.1	78.2	1.7	16.9	7.9	122.3	4.7	31.5
LM	8	3.9	212.2	1.1	36.1	8.0	81.0	5.1	14.7
RR	9	1.0	<b>24.3</b>	0.5	<b>7.1</b>	0	<b>33.0</b>	0.1	<b>8.4</b>
MR	10	2.3	103.0	1.6	28.1	5.2	147.4	4.2	30.4
ARK	11	2.2	130.5	0.7	36.2	3.2	93.9	4.5	18.1
GUL	12	1.5	99.1	0.5	35.5	1.3	70.7	2.8	16.7
RIO	13	<b>0.5</b>	<b>22.5</b>	<b>0.2</b>	<b>5.8</b>	<b>0.4</b>	<b>29.5</b>	<b>1.3</b>	<b>7.3</b>
UC	14	0.6	27.3	0.2	7.2	0	74.7	0.2	15.3
LCO	15	0.5	19.4	0.2	7.5	0.3	46.7	0.7	11.6
GB	16	0.7	33.3	0.3	10.3	1.8	71.5	1.3	21.1
PN	17	20.6	266.5	7.9	50.2	4.4	449.6	2.6	87.3
CA	18	1.6	323.2	0.4	101.9	1.3	233.4	1.1	61.3

354

355 Grid-scale estimates such as mean daily runoff readily highlight why decision makers have low confidence in  
 356 CESM2 output: the metrics are not salient and appear to have no skill. After aggregating the 7-day mean daily  
 357 runoff to watershed scales, some skill emerges in the annual minima and maxima, and seasonal cycles.  
 358 Snow-dominated watersheds perform poorly with regard to peak runoff volume and timing of the peaks and  
 359 lows, as expected (McCrary et al., 2022). Rain-dominated watersheds capture the inter-annual variability and  
 360 magnitudes of peak and low flows, and the seasonal hydrographs. While CESM2 at this coarse scale does not



361 represent the local topography and cannot represent finer scale snow, our analysis indicates the land surface  
362 model correctly simulates the overall bulk water budget for most watersheds as illustrated in Figures 7 and 8.  
363 However, the tail behavior of highest and lowest total watershed runoff is only captured by a few basins and so  
364 caution needs to be exercised in the interpretation and use of model results, as biases may propagate into the  
365 future. This is premised on the understanding of *why* the model can produce accurate results, and whether the  
366 accuracy can be reliably reproduced for the future climate (Wagener et al., 2022).

367 While participants at the NCAR workshop stated that precise estimates are not necessary, they also emphasized  
368 their desire for high confidence in the projected scale and direction of any changes. We note that “confidence” is  
369 derived from a combination of 1) credible process representation; 2) agreement with historical trends, given  
370 internal variability; 3) agreement across multiple models. It is worth noting that trends in extremes may be  
371 important without being statistically significant, as a limited sample of points (e.g. one per year) from a  
372 stochastic series is inherently noisy. However, some of these trends may emerge from the noise in the  
373 distribution and so are important to monitor.

374

375

376

377 CESM2-LENS projections could helpfully augment RCM output in rain-dominated regions such as Tennessee,  
378 Ohio, and California, where CESM2 most closely reproduces Livneh-VIC, by providing supplementary  
379 information on the relative uncertainty in the models. This is also true for transitional basins such as the Rio  
380 Grande, Northeast, and Lower Colorado, where seasonal snowpack may become more ephemeral and change  
381 the seasonal hydrological responses.

382

## 383 **6 Discussion**

384 As decision makers have become more immersed in developing water resource management adaptation plans,  
385 the role of “climate services” in developing salient climate information has increased (Briley et al., 2020;  
386 Brugger et al., 2016; Dilling et al., 2019). We tested our hypothesis that recent improvements in ESMs can allow  
387 decision-relevant metrics to be produced directly, by leveraging the combined experience of the author team,  
388 results from the NCAR workshop, and the wealth of literature on actionable knowledge (Bremer et al., 2020;  
389 Jagannathan et al., 2021; Mach et al., 2020; Vano et al., 2014). Given that no model can perfectly address all  
390 decision needs, we identified and evaluated multiple metrics that can frame specific water management  
391 decisions within the known constraints of the data (Lempert, 2021), or within the decision makers’ experiences  
392 (Austin, 2023; Clifford et al., 2020; Reed et al., 2022; Shepherd et al., 2018).

393

394 It is important to communicate the original purpose of the model and associated weaknesses, so that decision  
395 makers fully understand which information is appropriate to use in other applications (Fisher and Koven, 2020;  
396 Gettelman and Rood, 2016; Wagener et al., 2022). Given the balance between model fidelity and model  
397 complexity (Clark et al., 2015) and the absence of detailed global scale observation data (e.g., Gleason and  
398 Smith, 2014; Reba et al., 2011) CESM2 provides a plausible representation of Earth system processes and

399 moisture fluxes, but may not capture basin-scale specifics (Ek, 2018; Lehner et al., 2019). That said, there are  
400 continued efforts to improve the simulation of land surface processes and analyses such as those presented in  
401 this article can flag weaknesses for future improvement (Lawrence et al., 2019).

402

403 Establishing model fidelity also requires distinguishing an accurate representation of the climate processes from  
404 serendipitous correlation with observations. Whether the model has good process representation overall, or  
405 exactitude in one simulation can be established through internal variability analyses using large ensembles (e.g.,  
406 Deser et al., 2020; Tebaldi et al., 2021). Repeating the analyses with several different ESMs to establish the  
407 degree of agreement (Mankin et al., 2020) would further strengthen the usability of metrics presented in this  
408 article. It is also worth noting that the analysis presented here only used one reference dataset. As different  
409 reanalysis and observational datasets can have large discrepancies, a thorough model evaluation would also  
410 benefit from comparison to several products (Kim et al., 2020; Newman et al., 2015), including an assessment  
411 of how removing temporal adjustments in observations affects the statistics of extremes (Pierce et al., 2021).

412

413 While the precise details of precipitation and runoff may not be well simulated by CESM2, we found some  
414 aspects are credible. The frequency of wet days highlighted regions where current seasonal behavior is well  
415 captured, and may support planning around flood and drought control or wildfire risk when used in combination  
416 with other models or data sources (Austin, 2023; Clifford et al., 2020; Jagannathan et al., 2021; Reclamation,  
417 2016).

## 418 7 Conclusions

419 This paper presented an assessment of whether a standard resolution (~100 km grid) Earth system model is  
420 capable of producing information that water users typically employ in their decisions. Our motivation was to  
421 explore whether it is possible to reduce the need for intermediate downscaling, and to broaden the use of large  
422 model ensembles to quantify the influence of internal variability on localized decisions. We drew on the  
423 combined experience of the project team and workshop participants to identify potential metrics and familiar  
424 modes of visualization. This project used only CESM2 over the conterminous United States to develop example  
425 metrics that may be explored within other models and over other regions. CESM2 is unable to reproduce some  
426 metrics given the lack of topographical detail. A companion paper by Rugg et al. (2023) examines potential  
427 improvements to the subgrid-scale simulation of land processes to improve the representation of the  
428 hydrological cycle in mountainous regions.

429

430 We encourage others working in the decision space between climate data producers and users to be forthcoming  
431 about specific regions and reasons where model data are not credible, or where the model has particular  
432 weaknesses (such as the drizzle effect) that may be overcome with a different analysis approach.

433 For future model assessors, the following metrics were found to be salient for water users and were skillfully  
434 reproduced in many regions.

435

436 Rainfall:

- 437     ▪   Number of wet days ( $\geq 1$ mm of rain) per year/season
- 438     ▪   Mean precipitation on wet days
- 439     ▪   Duration of the longest wet and dry spells per year
- 440     ▪   Number of days with rain  $> 95$ th percentile of current climate wet day totals
- 441     ▪   Proportion of the annual total derived from days  $> 95$ th percentile of wet day totals

442 Runoff (aggregated up to basin level, as a volume for 3- and 7-day averages):

- 443     ▪   Annual maxima and minima
- 444     ▪   Frequency of very high or very low flows ( $< 10\%$  annual chance of occurring in the current climate)
- 445     ▪   Proportion of averaged daily runoff to annual total

446

447 The work presented in this paper is a small step toward establishing greater usability of climate model output by  
 448 decision makers. Continued collaboration is essential to improve the transfer of knowledge (e.g., data  
 449 requirements, model assumptions, decision constraints) between communities.

450

451 **Appendix A**

452

453 **Table A1: Hydro-meteorological responses used in water management decisions, and the specific metrics that have**  
 454 **potential for representation in ESMs. Metrics in bold are presented in this article.**

<b>Hydro-meteorological Responses</b>	<b>Typical Water Management Decision</b>	<b>Metric</b>	<b>Description</b>
Annual rainfall	Water supply and drought monitoring	Total Precipitation (PRCPTOT)	Total annual precipitation measured as rainfall or snow water equivalent.
Seasonal rainfall cycle	Seasonal water supply, reservoir operations management	Number of Wet Days (NWD), Mean Wet Day Volume (WDV)	Frequency of days with $\geq 1$ mm precipitation (NWD) per month, season or year, Mean precipitation on wet days calculated from PRCPTOT/NWD
Rainfall extreme	Flood and stormwater management	95th percentile (Q95) Number of very heavy rain days (N95) Very heavy rain volume (P95)	Rainfall percentile threshold that is exceeded by 5% rain events per year on

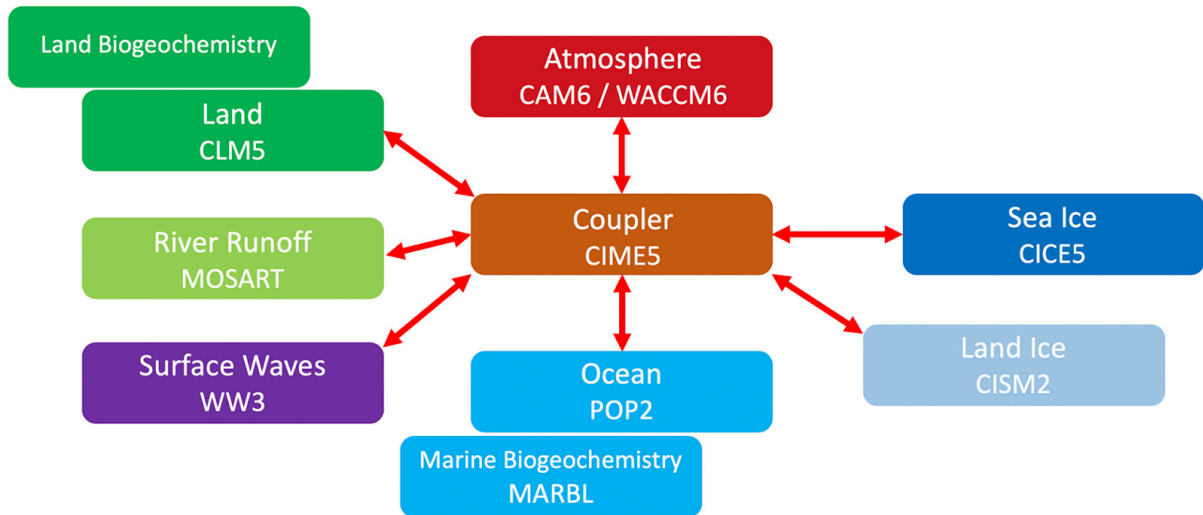
<b>Hydro-meteorological Responses</b>	<b>Typical Water Management Decision</b>	<b>Metric</b>	<b>Description</b>
		Proportional contribution of very heavy rain (P95tot)	average, and calculated from wet days only Frequency of days with rainfall exceeding Q95 Total rain falling on days exceeding Q95 Proportional of annual total derived from very heavy rain, calculated as P95/PRCPTOT
Rainfall extreme (dry)	Water supply planning and drought monitoring/planning including water rights and restrictions.	Consecutive dry days (CDD)	Maximum duration of spell with consecutive days measuring < 1 mm precipitation.
Rainfall extreme (wet)	Stormwater management, water supply planning	Consecutive wet days (CWD)	Maximum duration of spell with consecutive days measuring ≥ 1 mm precipitation.
High streamflow	Reservoir management and flood control, water quality management and water supply management, including use of supplemental water supplies	Annual maximum runoff (QMax) Description (JMaxF) Description (HFD)	Annual maximum daily volume of basin-wide runoff Julian day of QMax/ day of the water year Duration of high flows
Low streamflow	Water supply management, assessment of water shortages with respect to seasonal demands	Annual minimum runoff (QMin) Description (JMinF) Description (LFD)	Annual minimum daily volume of basin-wide runoff Julian day of QMin/ day of the water year Duration of low flows
Streamflow	Water supply planning, water quality management, reservoir operations management, planning future investment needs	7-day mean runoff (Q7)	Daily volume of basin-wide runoff averaged over 7 days. Often presented as percentage of annual total volume of runoff or Pardé coefficient (Pardé, 1933)
Very low streamflow	Water quality management for discharge permits,	7-day "10-year" low runoff (7Q10)	7-day averaged basin-wide lowest volume of runoff with

<b>Hydro-meteorological Responses</b>	<b>Typical Water Management Decision</b>	<b>Metric</b>	<b>Description</b>
	conservation management, drought planning		<10% annual probability of occurring. Estimated from Qmin series.
Very high flow	Flood management and planning, reservoir operations	7-day "10-year" high runoff (7Q90)	7-day averaged basin-wide highest volume of runoff with <10% annual probability of occurring. Estimated from Qmax series.
Streamflow	Water supply planning, reservoir operations management	Central Tendency (CT) Description (Q <sub>25</sub> , Q <sub>50</sub> , Q <sub>75</sub> )	Day of the water year when the cumulative annual runoff exceeds 50% of the total annual runoff Annual quartiles of cumulative annual runoff estimated from daily streamflow.
Snowpack	Reservoir operations and flood management, water supply planning	Snow Water Equivalent (SWE) Maximum (SWEMax) SWEMax Date SWE Duration	Volume of peak snow water equivalent Day of the water year when peak SWE occurs Total length of snow accumulation and ablation
Snowmelt	Flood management and reservoir operations	Snowmelt onset	Day of water year of snowmelt onset

455

456 **Appendix B**

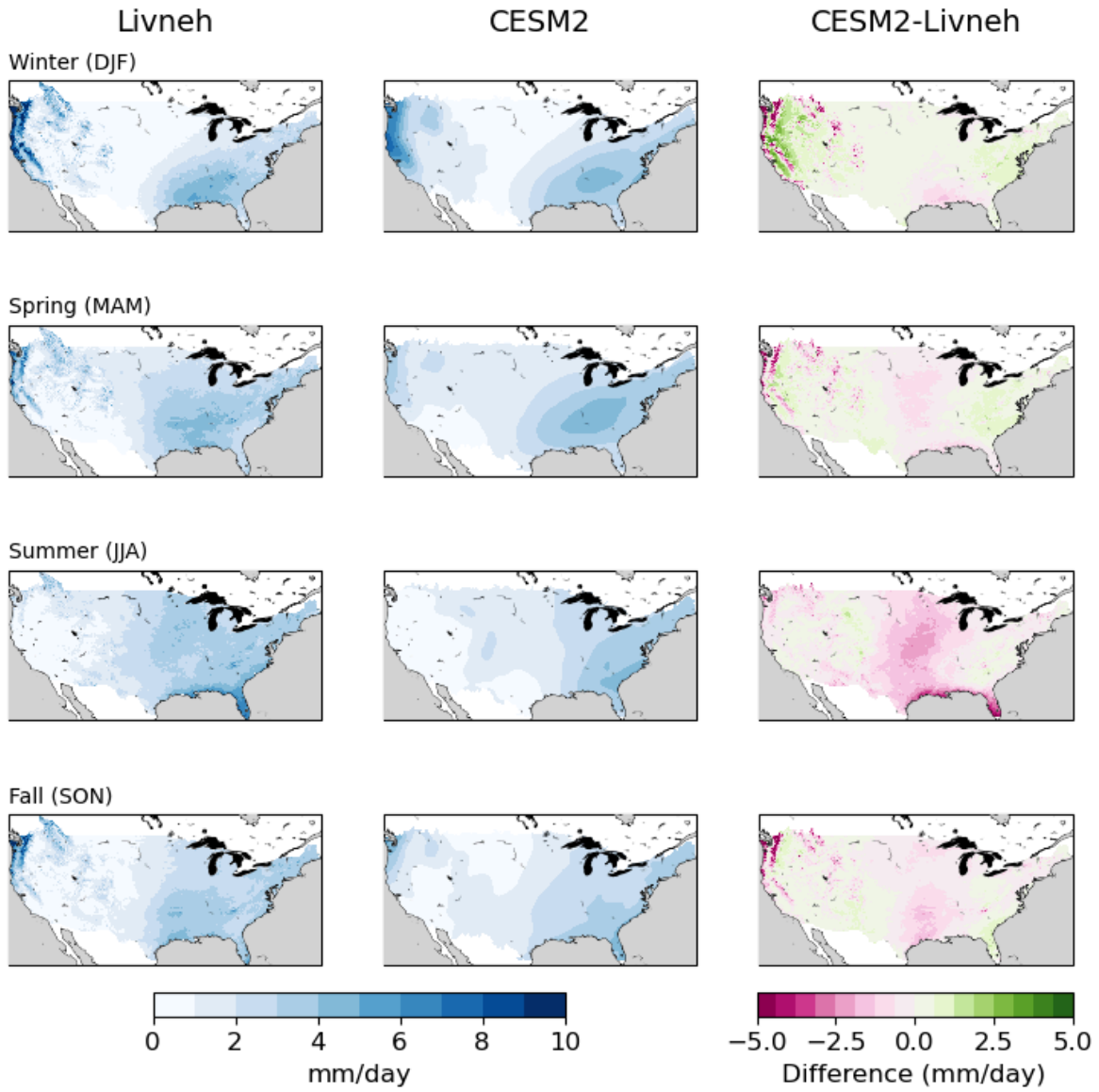
457 Schematic of the Community Earth System Model version 2 (CESM2) model components, reproduced from  
458 Danabasoglu et al. (2020) Figure 1.



459

460 **Appendix C**

461 Seasonal Mean Precipitation for Winter (top row), Spring (row 2), Summer (row 3) and Fall (bottom row) as shown  
462 in Livneh (left column) and CESM2 (middle column), and difference CESM2-Livneh (right column)

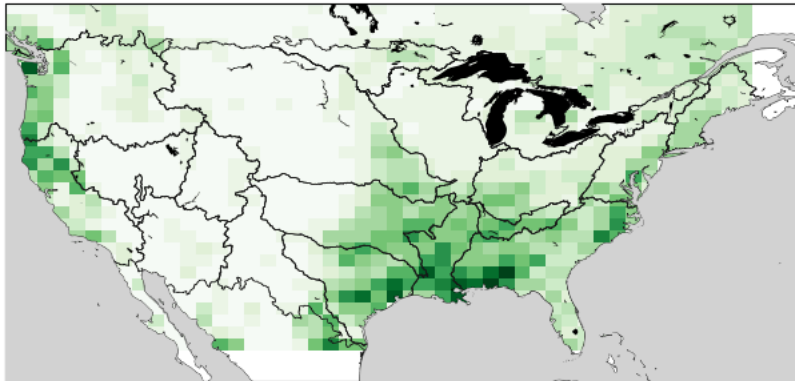


463

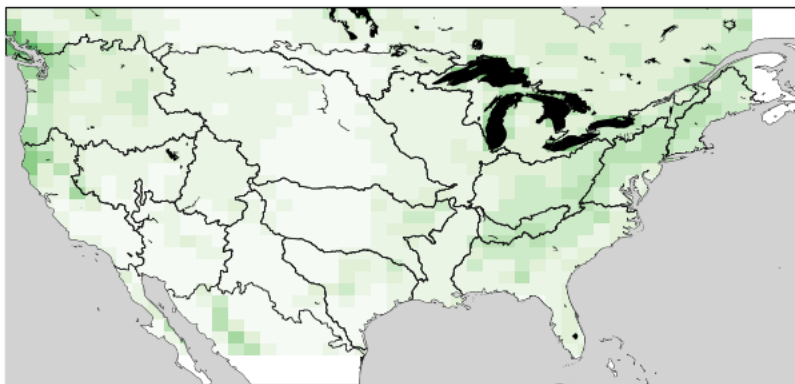
464

### Maximum Daily Runoff (1981-2005)

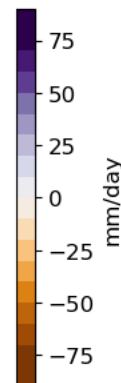
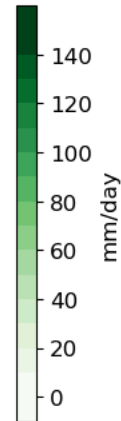
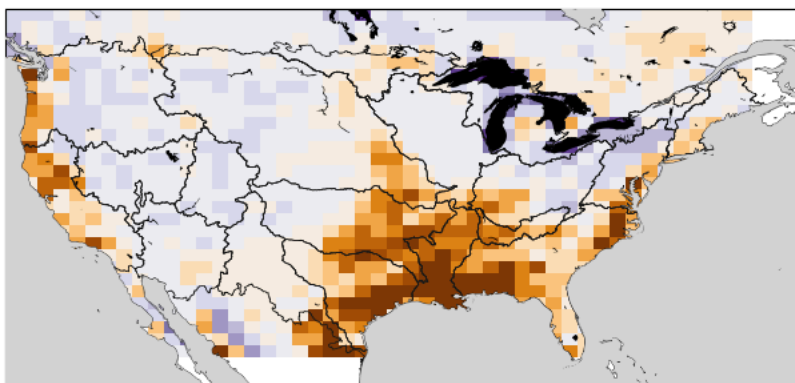
(a) Observations (Livneh-VIC)



(b) CESM2



(c) Difference (CESM2 - Livneh-VIC)





#### 470 Data availability

471 All data generated for this study (e.g., CESM2 and Livneh-VIC calculated indices) along with Jupyter  
472 notebooks to recreate tables and figures are available in the repository  
473 [https://github.com/maritye/PSIF\\_water\\_avail](https://github.com/maritye/PSIF_water_avail)

#### 474 Author Contribution

475 Conceptualization, M.T., J.R., E.G., A.N., A.W. and R.M.; Methodology, M.T., J.R., E.G.; Investigation, M.G.,  
476 M.T.; Data Curation, M.G., M.T.; Writing - original draft, M.T., A.R., and R.M.; Writing – reviewing and  
477 editing, M.T., J.R., E.G., A.N., A.W., R.M., A.R., F.L., C.B., and S.H.; Visualization, C.B., M.G. and M.T.;  
478 Supervision, J.R., E.G., A.N., F.L. and A.W.; Funding Acquisition, J.R., E.G., A.N., A.W., F.L., C.B., S.H. and  
479 M.T.; Project Administration J.R.

#### 480 Competing Interests

481 The authors declare that they have no conflict of interest.

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