

Response to reviewer 1, review 1 from 03.12.2023

We would like to thank the reviewer for their helpful recommendations. We very much appreciate the time and effort they took to carefully read the manuscript and to formulate detailed and constructive suggestions on how to improve it.

In the following, you find our responses to the individual recommendations (in bold).

0. The paper is in the direction of a community paper just published that they should cite in their manuscript: Laignel, B., Vignudelli, S., Almar, R., Becker, M., Bentamy, A., Benveniste, J., ... & Verpoorter, C. (2023). Observation of the coastal areas, estuaries and deltas from space. Surveys in Geophysics, 1-48

Response

The suggested reference provides a helpful and relevant summary of observations of the coastal zone, especially remote sensing observations. We therefore added it in the introduction as a motivation for this study. The passage reads now as follows (with old parts in grey):

Nowadays, there are several decades of remote sensing data available for coastal monitoring Laignel et al. (2023). Here we suggest to use observations for sea level and vertical land motion in combination with estimates of shoreline changes to quantify the geometrical relation between sea level and shoreline changes.

1. The authors use satellite radar altimetry to measure sea level. We know that this technique needs some specialized processing to retrieve data near coast. The authors use a coastal product that implements a retracker called ALES during two decades since 2002. However, the investigation of climate-related signals (e.g., trends) requires careful attention in ensuring homogeneity of processing between the various mission, removal of possible drifting in corrections, etc. For this reason ESA launched the Sea Level Climate change initiative to produce a validated gridded product that now is available through Copernicus (<https://cds.climate.copernicus.eu/portfolio/dataset/satellite-sea-level-global>). Moreover virtual stations are also provided here <https://climate.esa.int/en/projects/sea-level/data/> . The authors should use these products as benchmark to assess that their processing chain is consistent.

Response

In this study we used the OpenADB ALES-retracked along track product. In this product, the homogeneity of observations between different missions is achieved with cross-calibration with a global multi-mission crossover analysis (MMXO) (see Oelsmann et al. (2021), section 2.1). According to Oelsmann et al. (2021), possible drifts between the missions are almost completely removed by applying a radial correction to each single observation. A summary of the validation of this product is also provided in these slides: <https://mediatum.ub.tum.de/doc/1446336/document.pdf>

We considered using the ESA CCI virtual stations dataset at the beginning of the experiments, but it does not cover the study area.

We agree that a comparison of the along-track product with a gridded product, such as the one by ESA CCI might give further insights into uncertainties of coastal altimetry observations. However, as we found the area with the largest tide-gauge correlations is mainly positioned off-shore (> 50 km) on the open sea (see fig 4a), we don't expect to find very large differences between different altimetry products. We acknowledge that this could be different for other case studies. Nevertheless, in order to provide some guidance on this issue to the reader, we added the following paragraph in the new 'Limitations' section in the discussion:

For deriving an altimetry timeseries, we restricted ourselves to the use of one single dataset. This is an along-track product retracked with ALES, an algorithm specifically designed for coastal areas, provided by the OpenADB (see section 2.1). This OpenADB ALES product has been used successfully before in studies combining altimetry and tide gauges (e.g. Mangini et al., 2022; Oelsmann et al., 2021). Our comparison to the local tide gauges and the use for computing the Jarkus shorelines in comparison to the other solutions showed that offshore altimetry can be used to study shoreline changes. However, in order to get the full picture of uncertainties in altimetry datasets, it could be useful to additionally include other products, such as the ESA CCI gridded product (Copernicus Climate Change Service, 2018).

2. The authors mentions various papers related to the synergy of altimetry, tide gauge and GNSS data. I suggest to integrate with recent papers that provide updated inverse methods aiming at a better characterization of the errors in estimating the sea level trends

De Biasio F., Vignudelli S., Sea Level Change in the Mediterranean Sea from Satellite Altimetry and Tide Gauge. In Proceedings of Oceans from Space Conference (Editors: V. Barale, J.F.R. Gower, L. Alberotanza), 24-28 October 2022, Venice, Italy, 152-153, doi:10.57648/OceansFromSpaceV-2022-PROCEEDINGS.

De Biasio F., Vignudelli S., Baldin G.: Revisiting Vertical Land Motion and Sea Level Trends in the Northeastern Adriatic Sea Using Satellite Altimetry and Tide Gauge Data, Journal of Marine Science and Technology, 8(11), 949, doi:10.3390/jmse8110949, 2020.

Response

As both suggested references cover the same topic, we decided to include the latter one (De Biasio et al., 2020) in the introduction under point 1.1 Sea surface heights.

3. The GNSS time series contains discontinuities from antenna and receiver changes. It should be recalled that GNSS is a point, sometime not co-located with tide gauge. Estimation of the VLM depends on how the station is managed and how logs are updated. We have seen differences between the various services around the world. Our feeling is that only local people can assess well the significance of VLM trends and errors. Sometime using InSAR can help, but I don't want ask authors to add these data if they are not expert with this technique. I just like authors inform readers about caveats when using GNSS stations. In Table 1, please add error to your VLM estimation (versions 1,2,3)

Response

We agree that discontinuities in the GNSS timeseries and localised effects can be an issue when deriving rates of VLM, and that using only GNSS comes with some caveats. In this case study, the VLM rates are with magnitudes about 0.5 ± 0.4 mm/year relatively small compared to the rate of sea level rise with about 4.7 - 4.9 mm/year and therefore do not have a large effect on our main conclusions. To provide context for the reader, we added the following sentences to Section 2.3 Vertical land motion from GNSS (with old parts in grey):

Here, we decide to manually remove one, two or three of the bigger offsets (with 9 mm, 4.5 mm and 3.8 mm respectively) in order to get a time series clean of artificial jumps but still containing the signal of VLM. The resulting VLM rates are summarised in Table 1, together with estimates from other publications for the same GNSS station. These estimates cover slightly different time periods, but when assuming that VLM rates are stable over approximately four years, we see a rather wide spread between -0.18 ± 0.11 mm yr⁻¹ (Gravelle et al., 2023, ULR7A) and -0.63 ± 0.43 mm yr⁻¹ (Shirzaei et al., 2021). The differences in these outcomes of VLM rates indicate an uncertainty that approaches the magnitude of the signal. Another issue is that GNSS can only measure the component of VLM that takes place above the base of the GNSS station. Nevertheless, rates of GNSS height observations are currently the most accessible and up-to-date estimates of VLM. We therefore continue to work with the GNSS timeseries that results from removing the two largest offsets (version 2), as its VLM rate of -0.50 mm yr⁻¹ fits best in the range of estimates from earlier publications.

Furthermore, although we agree that the integration of more datasets could improve the understanding of all ongoing VLM processes at the study site we feel that it falls outside the scope of our paper. In order to provide some context on the shortcomings of using only GNSS as a source for VLM rates, we added the following paragraph in the new 'Limitations' section in the discussion:

Another correction applied to the tide gauges for the comparison with altimetry was the vertical land motion (VLM). Here we used only data from a GNSS station as a proxy for VLM. However, this approach may neglect other ongoing processes such as sediment compaction below the base of the GNSS station (Karegar et al., 2020). Additionally, we showed that identifying significant discontinuities in the GNSS timeseries due to antenna changes is not a straightforward task, leading to a relatively wide range of possible VLM rates between -0.18 mm yr⁻¹ and 1.15 mm yr⁻¹ (section 2.3). The picture of all VLM processes ongoing at Terschelling could be further improved by including InSAR (Interferometric SAR) data and GIA (Glacial Isostatic Adjustment) models.

Regarding the error estimates, the standard deviations computed in the least squares estimate are with magnitudes of 10^{-4} unrealistically small as no error correlations are known and considered in the input data. We therefore computed the standard deviation based on the actual residuals, and are reporting these in the table.

4. Detection of shoreline. Usage of state-of-the-art products is fine. However, Sentinel- 2 would provide more revisiting and better resolution. Landsat, like most other imaging satellites, is multispectral. The bands are Blue, Green, Red, near IR, and short wave IR, all with 30 m resolution There are one or two (depending on the

satellite) thermal IR bands with 60 m resolution There is a panchromatic image with 15 m resolution. In principle all spectral bands can contribute towards land-water discrimination, but in practice only a few bands provide robust and substantial leverage on classification land-water. Blue and green have the least contrast due to a combination of low and variable land albedo and possible strong and variable reflection from below water substrate. IR bands are generally better for water discrimination because there is a sharp increase in land albedo and increased absorption in water, leading to greater land-water contrast. There are two methods for improving the resolution: panchromatic sharpening and spectral un-mixing. The latest can improve detection from 30 m to 5 meters (see <http://meetingorganizer.copernicus.org/EGU2013/EGU2013-9681.pdf>). I suggest authors to discuss a bit the various methods and highlights pros and cons about using a customized processing or using global products.

Response

We appreciate the context provided by the reviewer and added some background and references for interested readers on optical satellite sensors and common methods to extract shorelines in the introduction subsection 1.2 Shoreline positions:

Shoreline positions extracted from optical satellite images are in the preceding literature usually referred to as satellite-derived shorelines. When working with images from optical satellite missions, there is usually a trade-off between spatial resolution and revisit period. The group of sensors with moderate resolution (about 250 m - 1000 m pixel size), such as Terra/Aqua MODIS, Envisat MERIS or Sentinel3 OLCI, have high revisit periods (about 0.5 - 3 days), but images are usually too coarse to extract shoreline geometries with sufficient accuracy relative to the width of the beach. The most commonly used optical sensors for shoreline extraction are high resolution sensors (ca 5 to 30 m pixel size). Since 1999, these satellites often carry additional panchromatic sensors that generate black and white images with a finer resolution, and can be used to downscale the multispectral images. Examples are the long-term Landsat missions (30 m resolution of multispectral channels, with a 15 m panchromatic band) with a revisit period of 16 days, Sentinel-2 MSI (10-20 m resolution) with a revisit period of 10 days (single satellite) or 5 days (two satellites in tandem) and long-term SPOT (5-20 m) with a revisit period of 26 days. Of these missions, SPOT is the only one whose data is not freely available. Finally, there are several commercial satellite missions with very high resolution (< 5 m) and short revisit periods (about 1-5 days) such as IKONOS, QuickBird, WorldView, or the cube satellite constellations by PlanetScope/Maxar. A more detailed review of optical satellite missions is given in Huang et al. (2018).

The process of extracting the shoreline from optical images starts usually by separating between land and water pixels. The easiest way to achieve this is to use a single band, which would preferably be one of the infrared bands where the differences in reflectance between water and land are the highest. The image of this band can be converted into a binary image by applying a threshold (e.g. Frazier and Page, 2000; Pardo-Pascual et al., 2012). This threshold can be chosen by a try-and-error procedure, or by computing it for example by using Otsu's method. Next to thresholding, the use of water indices (the ratios

of differences between bands) is very common to separate between land and water surfaces. There are several indices in use, where the choice depends on the type of the surroundings. For example, the Modified Normalised Difference Water Index (MNDWI) (Xu, 2006) was developed with the intent to better distinguish between water and buildings than the Normalised Difference Water Index (NDWI) (McFeeters, 1996). Usually the computation of an index is followed by the application of a threshold (e.g. Luijendijk et al., 2018; Dai et al., 2019; Almeida et al., 2021; Palomar-Vázquez et al., 2023), possibly also in combination with a classification (e.g. Vos et al., 2019b). Finally, there are advanced procedures to extract the shoreline at sub-pixel resolution, for example by using a marching squares algorithm to derive the shoreline contour (e.g. Bishop-Taylor et al., 2019a; Vos et al., 2019b) or by modelling the gradient of reflectances with polynomials and extracting the coordinates with the maximum gradient (e.g. Pardo-Pascual et al., 2012; Almonacid-Caballer et al., 2016; Sánchez-García et al., 2020).

Regarding the choice of Landsat or Sentinel-2, as pointed out in the manuscript, Almeida et al. (2021) state in their paper describing CASSIE that they use surface reflectances for Landsat, but Top-of-Atmosphere reflectances when using Sentinel-2. As explained in the manuscript, we were not sure which consequences we can expect when mixing up results from surface and TOA reflectances in one timeseries, and therefore decided to use only one of the sensors. Sentinel-2 has higher temporal and spatial resolution, but Landsat provides the longer timeseries that is required when studying long-term changes in response to climate change.

5. Comparison of altimetry with TG to estimate accuracy. I don't understand well how the tow measuring systems are made homogeneous. Comparison should be instantaneous. DAC and tides (if relevant) need to be removed as the two systems do not measure the same place. Some earth tides are seen partially by the TG. The recipe needs to be reported in appendix of the paper

Response

Observations from altimetry and from the two tide gauges were made comparable in terms of signal content by applying corrections for tides, atmospheric pressure and vertical land motion as described in the data section. We've additionally added a flowchart in the appendix (figure A2) to clarify the procedure.

6. Table 3 [should probably be table 4] : errors in trends need to be provided. Also significance of the trend should be checked (e.g. using the Mann–Kendall test).

Response

We added the error margins for the trends of cross-shore changes derived from the intersection of JARKUS profiles with a plane at sea level. Similar to the error margins for vertical land motion computed under point 3), the standard deviations derived from error propagation were unrealistically small as no error correlations are known and considered in the input data, so we computed the standard deviation based on the distance of a single cross-shore estimate to the linear model. We're very thankful for the reviewer to point this out, as it led to the discovery of one single transect that caused error margins for the trend of 322 m/year. When looking at the timeseries in detail, we found that it exhibited unrealistic jumps over 2000 m. As the problem is confined to this single transect and we

cannot identify the cause for these jumps, we decided to exclude this transect from all further computations. We've added the following sentence to section 3.2 Cross-shore changes from the intersection of land elevation data (JARKUS) with sea level to justify the exclusion of this transect.

Additionally, we also excluded one transect that exhibited unrealistic jumps larger than 2000 m from all computations.

The exclusion of the faulty transect led to a small change in the numbers for the trends averaged over the entire coastline up to 0.5 m/year, where the largest change was in the results for intersections with altimetry and the tide gauge timeseries reduced to the altimetry time period. Consequentially, some of the statistics when comparing cross-shore timeseries from JARKUS and from CASSIE change as well. In detail, the average bias increased from -80.6 m to -82.8 m, and the trend differences increased from -2.1 m/year to -2.3 m/year. On the other side, their standard deviations decreased by 13 % and 2 %, respectively.

Regarding the significance of the trends, we added the averaged results of the Mann-Kendall test (with 1 meaning the trend is significant within the 5-95 % confidence interval and 0 meaning that no significant trend was detected) as suggested by the reviewer to table 4.

7. A key point I would like mentioning is the replication of the approach to other sites and hopefully globally, following the promising validation in the study site. It is important to understand if a full remote sensing global application is feasible and if not the authors should explain how to fill the gaps. The authors highlight the need of land elevation data in high spatial and temporal resolution with high accuracy. Can SAR Interferometry fill the gap to measure land changes ? more and more SAR small satellites are going to be launched.

Response

We agree that the lack of global land elevation data with high horizontal resolution and high vertical accuracy is a key limitation when transferring the methodology to other sites. We included a short summary of currently known approaches to estimate the different parts of the topobathymetry from satellite remote sensing in the discussion under "Transferability to other sites" (old parts in grey).

The main limitation to transferability is the availability of land elevation data in high spatial and temporal resolution with high accuracy. While such data are available locally (e.g., Aquitaine in France (Nicolae Lerma et al., 2022), Narrabeen beach in Australia (Turner et al., 2016), Duck in USA (Larson and Kraus, 1994)), global datasets that cover also countries with less financial means are scarce. An alternative to land elevation data from in-situ and airborne LiDAR observations could be to estimate the topobathymetry from satellite remote sensing (e.g. Salameh et al., 2019; Gao, 2009). The topography can for example be derived from altimetry (e.g. Salameh et al., 2018), InSAR (e.g. Choi and Kim (2018)), stereo imagery (e.g. Almeida et al., 2019) or from a combination of sources (e.g. Pronk et al., 2024). For the bathymetry, there are different techniques that exploit the reflectance values from optical satellite imagery (e.g. Stumpf et al., 2003), that identify wave characteristics in optical or in SAR images (e.g. Bergsma et al., 2019), or that use a

combination of radiometry and wave kinematics (e.g. Najar et al., 2022). For intertidal zones, different studies exploited the corresponding tidal variability of shorelines and sea level (e.g. Bishop-Taylor et al., 2019b; Chen et al., 2023), for example by assigning sea surface heights to instantaneous shorelines ("waterline method", e.g. Mason et al. (1995)).

References

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