

Authors:

We thank the referees for carefully reading and reviewing our manuscript. Your critical point of view and your constructive suggestions have certainly helped us to improve our manuscript. Below, the reviewer's comments are shown and our responses are added in blue *"italic"* text.

Response to reviewer 1:

General Comments:

In this manuscript, the contamination of long-range transported volcanic aerosol particles was captured in the local Planetary Boundary Layer (PBL) above the Ocean Science Center at Mindelo, Cabo Verde using ground-based lidar observations. More specifically, the authors use two case studies, one before (typical local PBL) and one during the 2021 eruption of the Cumbre Vieja volcano at La Palma island which is located at a distance of about 1500 km from the measurement site to showcase changes in the lidar optical properties and therefore demonstrate the importance of volcanic aerosol detection for health related purposes even further away from the volcanic activity. More specifically, the authors support on a single case study and they report the particle extinction coefficient, the particle extinction-to-backscatter ratio and the particle linear particle depolarization ratio at three wavelengths (355, 532 and 1064nm). The authors use auxiliary information to support that the changes inside the PBL are due to the volcanic activity alone using, for example, AERONET data, backward trajectories, fire presence using FIRMS (although not shown) and FLEXPART simulations. Given the need for accurate detection of volcanic particles in the atmosphere and the scarce lidar observations during volcanic eruptions the study has potential but in the current version it lacks scientific interest and clarity. Therefore, the manuscript may be considered for publication after major revisions.

Thanks for the critical comment. We have now changed the focus of the manuscript a bit, so that the focus is more on the novel and unique optical properties rather than on the pollution event itself while preserving all main information. We also have changed the title so that it becomes clear that we focus on case studies. Furthermore, we completely revised the abstract so that the content of the manuscript should be much more obvious. With all these changes, we believe that the manuscript has significantly improved and also its scientific interest and clarity has become much clearer.

Specific comments:

The title is misleading. In the manuscript, two case studies using a PollyXT lidar are shown. One before the eruption at La Palma island and a second one during the volcanic activity. Then, the authors use the two case studies to discuss the changes in the lidar optical properties over the measurement site. Therefore, I suggest choosing a more suitable title including the word case study. Please note and correct throughout the manuscript: The name of the island is La Palma and should not be confused with Las Palmas which is the capital of Gan Canaria.

Thank you for making us realize that the title was misleading. We have now changed our title to "Tropospheric sulfate from Cumbre Vieja (La Palma) observed over Cabo Verde contrasted to background conditions – lidar case study of aerosol extinction, backscatter, depolarization"

and lidar ratio profiles at 355, 532 and 1064 nm”, which reflects better the contents of the manuscript. Additionally, we have also corrected the spelling of “La Palma” throughout the manuscript. Many thanks for this hint!

The abstract should be short, clear, and summarize the findings of the study. As written, the abstract is misleading and lacks scientific importance. It is misleading because, the authors mention the full duration of the volcanic activity, but they do not mention that the findings rise from one individual case study. It also lacks scientific importance because although it is mentioned that a special version of a PollyXT system was used which allowed the calculation of the particle extinction coefficient, the particle extinction-to-backscatter ratio and the particle linear particle depolarization ratio at three wavelengths (355, 532 and 1064nm), there is no mentioning of the 1064nm wavelength which is the added value compared to the standard high-power lidars which operate at 355 and/or 532nm. In fact, there is no comprehensive summary of the findings per wavelength per aerosol optical property. More specifically, the lidar ratio and linear particle depolarization ratio which are of great importance in aerosol classification are absent from the abstract.

Thank you for your critical but constructive comment. We have restructured and shortened the abstract to point out that we performed a case study and not a long-term analysis and highlighted the availability of the measurements at 1064 nm. Furthermore, we added the results, especially for the lidar ratio and the depolarization ratio, at the particular wavelengths.

The authors support the changes in the PBL optical properties during the volcanic eruption to be caused by the presence of sulphate aerosols and they exclude the presence of other aerosol sources using FLEXPART and backward trajectories together with fire location from FIRMS. It would be beneficial to include a FIRMS figure and I was also wondering whether there are in situ observations at Mindelo site to check the presence of black/organic carbon. This will solidify your conclusions regarding the higher lidar ratio observed during the volcanic activity and its origin.

Thank you for your recommendations! We added a FIRMS figure in combination with backward trajectories to the appendix B of the manuscript. We also agree that in-situ observations would be helpful for the argumentation and checked the availability of in-situ data (e.g., in the Global Atmosphere Watch World data center <https://ebas-data.nilu.no/>) but, unfortunately, no measurement data of black or organic carbon at Mindelo is publicly available for that time period.

Then, I really missed a long-term report of the lidar optical properties including the full period of the volcanic activity. The authors advertise in quite a few points in the manuscript that the lidar has captured the full volcanic activity. So, why did you choose to focus on one case study only and not include the full dataset? Furthermore, it will be of great importance to go a step further and estimate the mass concentration of the long-range transported sulphate aerosols.

Thank you for this comment! In this manuscript, we focus on the case study analysis and the respective optical properties. Given the revised title and abstract it should become clearer now. But we are aware that a long-term study of the full period of volcanic activity would be of large interest. Continuous lidar measurements were performed during that time, but a complete quality control data set including successful and reliable cloud screening as a

prerequisite for a high-quality aerosol optical data set is not yet available. That is why we, for now, analyzed single days within this period. However, a long-term study based on automatically retrieved lidar optical properties may be subject of a future publication as we added in the outlook. For now, we reworked the misleading paragraphs in the manuscript. Furthermore, we now added a profile of the sulfate mass concentration to Fig. 5 and a corresponding paragraph in lines 347–353 of the revised manuscript to highlight the potential of such lidar measurements for studies of air quality.

Overall, in its current form the manuscript is a report of a single case study in which the optical properties are the result of marine and sulphate aerosol mixture of unknown contribution with limited added value to the scientific community and at places highly speculative. The presentation and usage of English should be also substantially improved.

With the new title and the updated abstract, the focus of the manuscript should be clearer. Several new aspects are considered with the case study: It is the first time, optical properties at 1064 nm for sulfate aerosol are measured. We can clearly disentangle the effect of the sulfate from marine by comparing to background conditions and, thus, show the impact of the volcanic eruption at Mindelo on this specific day. We could show that no ash was transported from the volcano towards Mindelo. Furthermore, we reworked the manuscript linguistically. With all these changes, we believe that the manuscript is of high value for the scientific community.

L122-124: Can the authors comment on the stability of the calibration for this wavelength? What is the expected error in the optical products? The references in this sentence point to another lidar system and not PollyXT. What is the error estimation for this system? Furthermore, the 1064nm depolarization capability is also a new feature. What is the uncertainty of the particle depolarization ratio for this system?

The calculation of the extinction coefficient at 1064 nm via the rotational Raman method follows the methodology described in Haarig et al., 2016. The spectral cross-talk calibration methodology uses a liquid cloud as it was introduced in Haarig et al., 2022 and is referenced in the manuscript but briefly described here for your convenience: the strong backscatter signal at the cloud base is used to iteratively determine the spectral cross-talk correction factor. The elastic signal (1064 nm) multiplied with the spectral cross-talk correction factor is subtracted from the rotational Raman signal (1058 nm) so that the particle induced strong backscatter signal at cloud base (elastic backscattering process) is not contaminating the rotational Raman signal any more. The spectral cross-talk correction factor changes only when the neutral density filters are changed in one of the two channels (either 1064 or 1058 nm). As there was no change in neutral density filters between 24 September and 4 October 2021, the spectral cross-talk correction factor of $6.7e-4 \pm 0.3e-4$ is valid for the whole period.

The calibration of the depolarization ratio at 1064 nm and the estimation of its uncertainties followed the same approach as the calibration at 355 and 532 nm (Engelmann et al., 2016). The Delta 90° calibration (Freudenthaler et al., 2009) with a linear polarizer after the pinhole was applied.

We also included a dedicated section describing the uncertainty estimation for all optical products now in the manuscript – see lines 141–150 of the revised manuscript. Following this section, the error for the particle depolarization ratio is about 0.01 at 1064 nm.

L178-179: Is there a reference to support the statement that the measurement on the 16th of September represent the typical situation over the location?

The lidar-derived optical properties have been already studied, quite intensively, in the framework of ASKOS and L2A+ (ESA funded projects). A comprehensive overview of the PollyXT lidar measurements conducted during the ASKOS intensive measurement periods is shown in Fig. S1. The attenuated backscatter coefficient at 1064 nm (Fig. 1a, 1c, 1e) in combination with the volume depolarization ratio at 532 nm (Fig. 1b, 1d, 1f) reveal the typical aerosol conditions above Mindelo, which are a clean marine boundary layer (MBL; non-depolarizing spherical particles), with a dust aerosol layer (depolarizing non-spherical particles) on top of that.

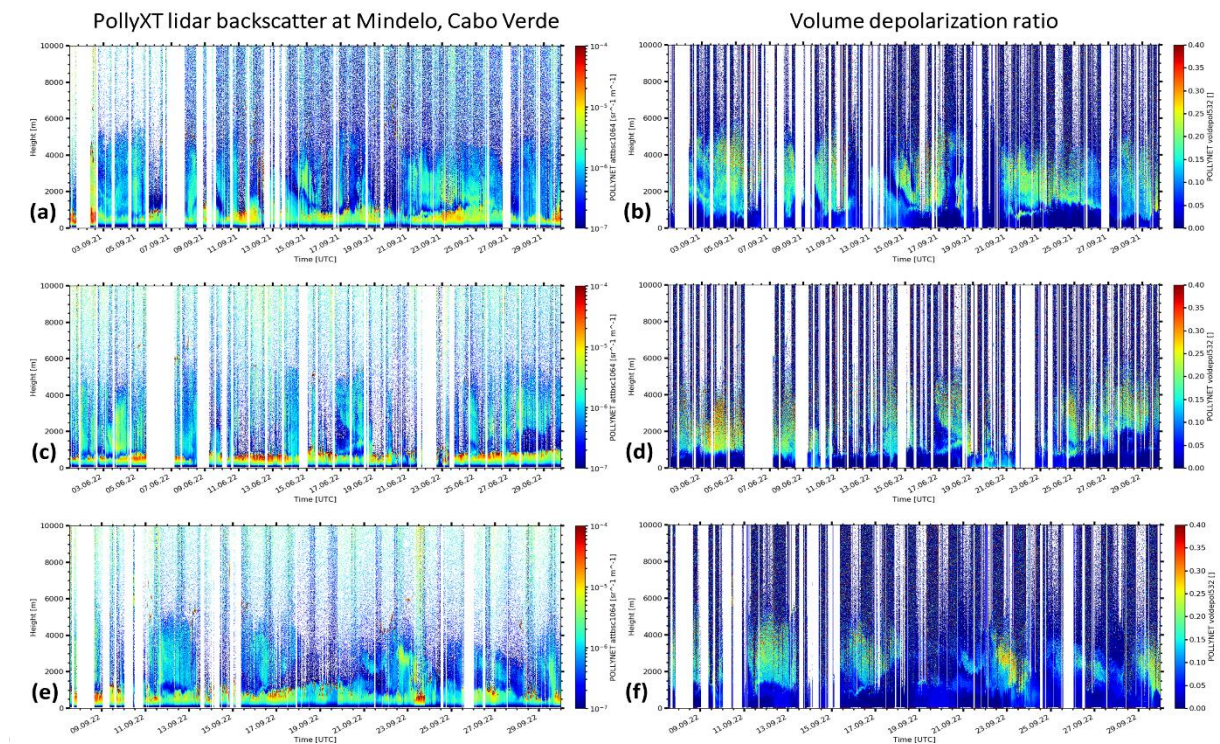


Figure S1: Overview of the lidar attenuated backscatter coefficient at 1064 nm (left column) and volume depolarization ratio at 532 nm (right column) as retrieved from the PollyXT lidar during the ASKOS operations in September 2021 (a, b), June 2022 (c, d), and September 2022 (e, f).

These results have been published (for now) in project deliverables (e.g., <https://l2a.space.noa.gr/backend/assets/e7a1b125-e2b2-4489-a205-720bd4f8077a?download>, last access: 11 January 2024). Additionally, the data are available in native format via the ESA Validation Data Center (EVDC) under DOI: 10.60621/jatac.campaign.2021.2022.caboverde (2023). We have now included the relevant citations in lines 208–211 of the revised manuscript.

Figures 3-5: What the error bars refer to? Do they include the systematic errors or just the variability caused by the time averaging?

The error bars show the statistical error in case of the particle extinction coefficient and a relative error (minimized systematic error + statistical error) of 15 % for the particle backscatter coefficient. The errors of the lidar ratio and the Ångström exponent were calculated using the error propagation. For the particle linear depolarization ratio, constant absolute errors of 0.02 at 355 nm and 0.01 at 532 and 1064 nm are considered as described now in the manuscript. Thus, minimized systematic errors are included as well as statistical

errors. We added a detailed explanation of the uncertainties in Sect. 2.1 (lines 141–150) and clarified the meaning of the error bars in the figure caption.

Technical corrections:

Lines 67-69: Please give a reference.

Done, please see line 70 in the revised manuscript.

Lines 75-86: Note this publication about Cumbre Vieja volcanic eruption:

Bedoya-Velásquez, A.E.; Hoyos-Restrepo, M.; Barreto, A.; García, R.D.; Romero-Campos, P.M.; García, O.; Ramos, R.; Roininen, R.; Toledano, C.; Sicard, M.; et al. Estimation of the Mass Concentration of Volcanic Ash Using Ceilometers: Study of Fresh and Transported Plumes from La Palma Volcano. *Remote Sens.* 2022, 14, 5680.

<https://doi.org/10.3390/rs14225680>

Thank you for pointing to this interesting publication that was missing in our review! We included it now!

Line 104: Please check that the coordinates of the measurement site are correct. I think it should be W, not E.

We corrected the wrong coordinate of the measurement site as well as of the volcano. Thank you for notifying!

Figures 2-5: Use the same thickness for all lines.

The different line thickness was chosen to satisfy the requirements for the color blindness tests but thanks to your comment we noticed that this solution may be confusing. Thus, we standardized the line thicknesses now.

Figure 3: Be consistent on the heights at which the error bars appear. For example, Figs. 5a, b and c all have the error bar in different locations.

We chose the different heights on purpose to ensure a better visibility of the error bars but we also acknowledge your comment as, of course, choosing consistent heights ensures a better comparability. Thus, we changed it according to your recommendation.

Figure 4d: ref. lines are missing from the legend. For example, the dotted orange line which most probably refers to the bae532/1064 (RR) on the 16th of September.

We added the missing symbols and labels in the legend.

Figure C1: Include the 16th of September 2021 as it is one of the case studies. The figure could also be a bit more zoomed to the region of interest.

Thanks for this suggestion. We have included it now accordingly.

Response to reviewer 2:

General comments:

The manuscript describes the measurements of a volcanic aerosol cloud from the Cumbre Vieja volcano eruption in Sept. 2021 that was transported 1500 km over the Atlantic Ocean from the island La Palma to the measurement site in Mindelo on the Cape Verde Islands. Measurements were performed with a multi-wavelength Raman and depolarisation lidar and a sun- and moon-photometer. Additionally, trajectory calculations are used to verify the origin of the measured air masses in the boundary layer and in the free troposphere. The comparison of the measured values of the optical properties with the ones from a clean reference period a few days earlier and from lidar measurements of other campaigns indicate that the measured aerosol in the boundary layer is sulfate aerosol from the volcanic eruption.

As there are only few high quality multi-wavelength lidar measurements of aerosol stemming from volcanic eruption - especially spanning the wavelength range from 355 nm to 1064 nm, the retrieved aerosol optical properties add valuable data to the global database that can be used for aerosol characterization. The manuscript is very well organized and written.

We greatly appreciate the review and detailed comments provided. As many comments were made with respect to the uncertainties of the lidar-derived optical products, we now have included a dedicated paragraph on this issue in the manuscript. You can find it in lines 141–150. Our responses to the specific comments are as follows.

Specific comments:

Las Palmas is a municipality and city on Gran Canaria. The Cumbre Vieja volcano is located on the island called La Palma.

Thank you for the correction! This point was also raised by RC1. We corrected the spelling of the island throughout the whole manuscript.

The dual-field-of-view channels are mentioned in chapter 2.1 but unfortunately not used. Why? And if not used, why mentioning it?

Thank you for your question. The dual-field-of-view (FOV) channels are mentioned in the manuscript as they are now part of the standard PollyXT setup. Since the dual-FOV polarization lidar technique was implemented only recently (Jimenez et al., 2020a, 2020b) it is not included in the instrument's overview paper (Engelmann et al., 2016). The technique is powerful, allowing the determination of microphysical liquid-water properties, which combined with the lidar-derived aerosol properties can be used for studying aerosol-cloud interactions. Since the scope of the paper is not aerosol-cloud-interactions we did not use the dual-FOV channels, nevertheless, their availability is worth mentioning. We have also included a relevant statement in lines 118–121 of the revised manuscript.

The absolute scale of the "attenuated backscatter coefficient" shown in the fig. 2 need reference values at reference heights for each plot. This is even more necessary as measurements of two different days are compared (figs. 2c and 2d).

It is worth to mention that here the calibrated attenuated backscatter coefficient is shown. Thus, it is a quantitative property which can be compared to each other. However, it is clear that different attenuation in lower altitudes leads to different values at higher altitudes even though aerosols properties there are similar. Thus, values are not comparable in terms of backscatter intensity at a given altitude. The aim of providing the plot is to show the vertical structure on the days of interest. We made this now clearer in lines 187–189 of the text.

Figs. 3e and 4e: the x scales should be the same for an easier comparison of the values in the two plots.

Thanks for the hint! We changed the x scales of Fig. 4e accordingly to the one of Fig. 3e. For the particle backscatter and extinction coefficient, we explicitly chose different x scales and prefer to leave it like that to provide as much detail as possible for the very different atmospheric conditions on these two days. Instead, we inserted the profiles from the Fig. 3 as reference in grey lines.

Fig. 5b and 5c: the red 1064 nm (RR) line does not match the description in the figure caption. If an increased smoothing of 742.5m is used for the particle extinction coefficient and the lidar ratio at 1064nm, then there can be only one value in the considered height range between 0.25 km and 1.0 km, but there is a line of values between 0.75 km and 1.0 km.

*The vertical smoothing is performed by using a moving average filter, i.e., a vertical distance of 7.5 m between the data points is preserved. However, for an arbitrary smoothing length “s”, each data point contains information of the height range from 0.5*s below this point to 0.5*s above it. Thus, with increasing s, the height, where the profile starts, increases, namely it starts always at 0.5*s, but then it continues in steps of 7.5 m. For the 1064 RR products, with considering a smoothing length of 742.5 m, the data points of the extinction coefficient and lidar ratio in Figs. 5b and 5c (reaching now from 650m to 800m) contain information of the height range 278.5–1171.5 m. This limitation to this height range is done to exclude effects of the incomplete overlap (lower end) and noise (upper end). Please note the new height range of the 1064 extinctions products (650–800 m instead of 750–1000 m). We found a mistake in our plotting and had to adapt the range accordingly. This correction also led to minor changes with regard to the layer mean value, which we updated as well. Furthermore, we added an explanation concerning the smoothing to Sect. 2.1 (lines 141–150).*

The error bars in figs. 3 to 5 are not explained:

1. What are the error bars showing? Should be mentioned in each figure caption.

The error bars show the statistical error in case of the particle extinction coefficient and a relative error (minimized systematic error + statistical error) of 15 % for the particle backscatter coefficient. The errors of the lidar ratio and the Ångström exponent were calculated using the error propagation. For the particle linear depolarization ratio, constant absolute errors of 0.02 at 355 nm and 0.01 at 532 and 1064 nm are considered as described now in the manuscript in a dedicated section. We also clarified it now in the caption of Fig. 3 and refer to the new paragraph concerning the errors. As we avoided to repeat redundant information, which are the same in all three figures, we prefer to add the explanation of the error bars only to the caption of Fig. 3.

2. Why does the error not change over large height ranges for e.g. in figs. 3c, 3d, and 3e, etc?

This point should be clearer now with the better explanation of the errors in the dedicated section in lines 141–150. As the errors provided are relative ones, and the quantities itself in Fig. c,d,e are only slightly changing with height, it appears that the error bars are nearly constant.

3. Figs. 3e and 4e should have the same x-scale for a better comparison.

Done.

4. The error bars in fig. 5b of the 355 nm and 532 nm are unrealistically small.

We have added now a dedicated section (see lines 141–150) concerning the error calculation in the methodology part. Thus, the way we calculated the uncertainties should be clear. For the extinction, we obtain the uncertainty from the error of the linear fit made to get the deviation in the extinction formula. The resulting error is in the order of 10 Mm^{-1} for the plotted height ranges and, thus, seems to be small compared to the large values we measured.

5. In fig. 5c the error bar of the 1064 nm RR curve is much larger than the layer variability. What does it show?

The error was calculated using the error propagation and includes the errors of the particle backscatter and extinction coefficient. With the added explanation of the errors in Sect. 2.1 (lines 141–150) it should be clearer now.

The header of table 1 is a bit confusing:

We rephrased the header now.

1. In fig. 5b the values at 1 km height are already less than half of the mean values below. What does then the "edge effect" mean?

"Edge effect" meant that we did not want to include the transition from the PBL to the lofted layer into the mean values for the PBL (and vice versa). Thus, we reflected that choosing 1 km as the upper boundary for the PBL was not appropriate and we changed it to 800 m. But we concluded, that "edge effect" is a misleading term and rephrased it accordingly.

2. It is unclear what "standard deviation" means and how it is derived. Is it the uncertainty due to signal noise or the parameter variability over height?

In this context, "standard deviation" means the statistical error due to the averaging over the layer, i.e., it is the parameter variability over height. As this quantity strongly depends on the vertical smoothing, we now decided to replace these values for the intensive optical properties with the layer mean of the given errors, which were used for the error bars in Figs. 3–5. We clarified this now also in the header of table 1.

3. Is it possible that table 1 does not show any uncertainties?

In the table of the preprint, the +- values did not include the uncertainties, which were used for the error bars in Figs. 3–5, but only the standard deviation (see answer above). We changed it now for the intensive optical properties, which do not show a significant vertical variability within a homogenous aerosol layer. For these parameters, the table now shows the layer mean error based on the uncertainties of the error bars. We clarified it in the table and, together with the new paragraph in Sect. 2.1 concerning errors/uncertainties, it should be clear, what is meant.

4. Why is there no 1064 nm AOD?

Thanks for this hint. We now added it.

5. What do the +- values of the AOD mean? Standard deviation?

The uncertainty values of the AOD describe the parameter variability. The AOD was calculated from the layer mean extinction coefficient while its uncertainty values were derived using the Gaussian error propagation with the standard deviation of the mean extinction as input. With the improved presentation of table 1, it now is clearer.

The error bars / standard deviation values in the plots and table need a better discrimination and clearer description. This is especially necessary if these values should be used in other studies for aerosol typing.

Thank you very much for the detailed feedback concerning the uncertainties! Of course, they are of particular importance! Thus, we added a new paragraph to Sect. 2.1 of the revised version of the manuscript, where all errors are explained in detail. Furthermore, we included a description of the used errors to the captions of the respective figures and the table.

Furthermore, are systematic uncertainties considered?

We calibrate our system according to ACTRIS/EARLINET standards. Thus, if we aware of any systematic error, we correct for it (e.g., depolarization calibration, polarization effects in the receiver unit). Remaining systematic errors are considered as described in the new paragraph concerning the uncertainties.

Considering the uncertainties and variability, is the increase of the lidar ratio mentioned in lines 237ff really significant?

You are right, it is not significant. We rephrased it accordingly.

In line 293f it is stated about the AE values:

During the measurement period, the values decreased to 0.7 ± 0.1 and 1.7 ± 0.3 , respectively, due to hygroscopic growth of the sulfate particles.

=> It is not clear, why the AE decreases. What is the difference between the aerosol parcels at the start and at the end of the "measurement period" - with respect to particle growth? Why do the latter grow more than the former?

You are right, it is not clear, why the AE decreases. It was stated like this in the article of Navas-Guzman et al., 2013, but not further explained, how the hygroscopic growth may have

changed during the night. To answer your question, we can only speculate. One reason could be that the relative humidity increased during the night or that particles arriving at the end of the measurement period followed a different or longer trajectory so that they were exposed longer to high relative humidity. Also, a changing portion of sulfate particles and sulfuric acid droplets is possible. Either, more sulfate particles, which are the larger ones, are present at the end of the night, or there are more sulfuric acid droplets, which have a stronger hygroscopic change [Navas-Guzman et al., 2013]. Of course, coating and mixing with locally-produced aerosol cannot be excluded, leading to a change in the effective radius, too. As the aim of mentioning these findings was just to give reference values for aerosol optical properties of volcanic sulfate it is not meaningful to extend this discussion in the text. Thus, we clarified in the text the origin of this statement.

Some correction proposals:

Line 28: diameter lower than => diameter smaller than

Done.

Line 63: ...which is in the range of $30 \pm 5\%$ for pure dust ... => at which wavelength?

Done.

Line 65: ...lower particle linear depolarization ratio ... => typical values?

Done.

Line 66: ; John et al., 2011). => (John et al., 2011).

Done.

Line 145: In addition, the add(horizontal) distribution of the volcanic plume was monitored.

Done.

Line 150: A time series of the AOD => A time series of the CIMEL AOD at Mindelo

We clarified that the columnar AOD from the sun photometer is meant and not the lidar-derived one.

Line 151: the hourly mean AOD was about 0.4. => at which wavelength?

Done.

Line 154: AODs, with values close to 1.0, => at which wavelength?

Done.

Line 193: than the once measured => ones

Done.

Line 195: The lidar ratio at 1064nm is in line with dust observations at Leipzig, Germany => are there no changes due to long range transport to be expected?

The lidar ratio can be linked to the mineralogical composition of mineral dust, which depends mainly on the source region and does not change much during the transport. At least, previous observations of the lidar ratio at 355 and 532 nm pointed to this dependence (Veselovskii et al, 2020). The dust source region (Mauritania, Algeria or in more general terms Western Sahara) was similar for the observations at Leipzig in 2021 and the ones presented in the current manuscript. Therefore, we do expect a similar behavior for the dust lidar ratio. A comment about the similar source regions was added to the manuscript in lines 223–225.

Line 207: In addition, vertical smoothing was reduced, which improves the accuracy of the near-field profiles => in which sense does the accuracy improve? The resolution is improved. But for which purpose?

The resolution of the profiles in Fig. 5 was increased to reduce the overlap effect and to display the profiles down to lower altitudes than in Fig. 3 and 4, where a larger vertical smoothing was required to avoid too much noise in altitudes above 1 km. Furthermore, an improved vertical resolution allows us to better illustrate the vertical variability within the PBL. We clarified this now in the text. But you are right, the words accuracy, precision etc. should be used with more caution.

Line 278: the measured quantities => Do you mean aerosol "quantities" or aerosol "optical properties"?

You are right. "Aerosol optical properties" seems to be the more appropriate phrasing. We changed it now, also at equivalent positions in the text.

Line 321: The intense pollution caused an unusually high AOD of more than 1.0 at different spectral bands. => change "different spectral bands" to "at the smaller wavelengths" - or so.

Done.

Line 351ff: " Having measurements at all three wavelengths allows us to get new insights in lidar-based aerosol typing and to enlarge our data sets. The findings of this study provide useful insights on the lidar-derived optical properties of volcanic aerosol."

=> This are diffuse statements. What are the new and useful insights?

Thank you for this comment. You are right, these sentences were not clearly formulated and are redundant in the current state. We rephrased the paragraph to better point out the benefit of our study. The new and useful aspects of this study are that we provide aerosol optical properties measured with a multiwavelength-Raman-polarization lidar even at 1064 nm, which is a novel feature and rarely done so far. It was the first time that volcanic sulfate could be studied applying this technique. Our findings enlarge the existing data sets and can, thus, help to improve the lidar-based aerosol typing. Besides this aspect, they will also help to improve the quantification and modeling of the radiative effects of tropospheric volcanic aerosol.

Further changes:

During the work on the revision, we found some further points not raised by the reviewers, which we wanted to improve on our own accord. Especially the intense discussions on the error bars made us realizing that it is more reasonable to provide the height-constant absolute error for the particle linear depolarization ratio and to show the layer mean error for the intensive optical properties, while for the extensive optical properties we kept the standard deviation. To include a better description of the errors in the captions of the figures and the table, we decided to shorten the original captions. We also added the uncertainties to the running text. Furthermore, we found and corrected an error in the plotting routine leading to minor changes in the values summarized in the table. As the FLEXPART simulations were performed within the TRACE model, we added a paragraph concerning TRACE in lines 160–164.

Tropospheric sulfate from Cumbre Vieja (La Palma) observed over Cabo Verde contrasted to background conditions - lidar case study of aerosol extinction, backscatter, depolarization and lidar ratio profiles at 355, 532 and 1064 nm

~~Tropospheric sulfate from Cumbre Vieja volcano at Las Palmas, transported towards Cabo Verde - lidar measurements of aerosol extinction, backscatter and depolarization at 355, 532 and 1064 nm~~

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Abstract. In September 2021, volcanic aerosol (mainly freshly formed sulfate plumes) originating from the eruption of Cumbre Vieja on La Palma, Canary Islands, Spain, crossed Cabo Verde at altitudes below 2 km. On 24 September 2021, an extraordinarily large aerosol optical depth (AOD) close to 1.0 (daily mean at 500 nm) was observed at Mindelo, Cabo Verde. This event provided favorable conditions to obtain lidar-derived profiles of extinction and backscatter coefficients, lidar ratio and depolarization ratio at 355, 532 and 1064 nm in the sulfate aerosol plume. A novel feature of the lidar system operated at Mindelo is the availability of extinction, lidar ratio and depolarization measurements at 1064 nm in addition to the standard wavelengths of 355 and 532 nm. Having measurements of these parameters at all three wavelengths is a major advantage for the aerosol characterization and in aerosol typing efforts as the lidar ratio and the particle linear depolarization ratio are key parameters for this purpose. In this article, we present the key results of the lidar observations obtained on one specific day, namely on 24 September 2021, 04:38–05:57 UTC, including the first ever measurements of the particle extinction coefficient, the lidar ratio and the depolarization ratio at 1064 nm for volcanic sulfate, and discuss the findings in terms of aerosol optical properties and mass concentrations by comparison to a reference observation (16 September 2021) representing the typical background conditions before the start of the eruptions. We found an unusual high particle extinction coefficient of 721 ± 51 , 549 ± 38 and $178 \pm 13 \text{ Mm}^{-1}$ and an enhanced lidar ratio of 66.9 ± 10.1 , 60.2 ± 9.2 and $30.8 \pm 8.7 \text{ sr}$ at 355, 532 and 1064 nm, respectively, in the sulfate-dominated planetary boundary layer (PBL). The particle linear depolarization ratio was $\leq 0.9\%$ at all respective wavelengths. It is the first time that lidar-derived intensive aerosol optical properties could be derived for volcanic sulfate at all three wavelengths and, thus, it is a highly valuable data set for global aerosol characterization. The lidar analysis also revealed

a sulfate-related AOD of about 0.35 ± 0.03 at 532 nm of the total PBL-related AOD of 0.43. The rest of the AOD contribution was caused by a lofted Saharan dust layer extending from 1.4 to 5 km and leading to a total AOD of 0.79 at 532 nm. Volcanic ash contribution to the observed aerosol plumes could be mostly excluded based on trajectory analysis and the observed optical properties. Peak mass concentration was $178.5 \pm 44.6 \mu\text{g m}^{-3}$ in the volcanic influenced – sulfate dominated polluted PBL showing the hazardous potential of such sulfate plumes to significantly worsen local air quality at even remote locations. From 19 September to 13 December 2021, volcanic eruptions took place at the Cumbre Vieja ridge, Las Palmas, Canary Islands. Thereby, fine ash and volatiles, like sulfur dioxide (SO_2), were emitted and transported over hundreds to thousands of kilometers away from the island. Continuous lidar observations with the multiwavelength-Raman-polarization-lidar Polly^{XT} were performed at the Ocean Science Center at Mindelo, Cabo Verde, in the framework of the Joint Aeolus-Tropical Atlantic Campaign (JATAC) 2021/2022 enabling the characterization of the atmospheric state above Mindelo during the eruption period. A special feature of the system operated at Mindelo is, that measurements of the particle extinction coefficient, the particle extinction-to-backscatter ratio (lidar ratio) and the particle linear depolarization ratio are available at all three wavelengths (355, 532 and 1064 nm). The typical aerosol conditions over Mindelo are a clean marine planetary boundary layer (PBL) up to approx. 1 km and above a Saharan dust layer (SAL, up to 6 km) during northern hemispheric summer and fall. A particle extinction coefficient smaller than 200 Mm^{-1} , a lidar ratio smaller than 30 sr and a particle linear depolarization ratio close to 0% have been typically observed within the planetary boundary layer, while a lidar ratio between 40 and 60 sr and a linear depolarization ratio between 20 and 30% are characteristic for the SAL above. In contrast, during the time of the volcanic eruptions, a strongly polluted PBL was observed on specific days beginning on the 23 September 2021, whereby the particle extinction coefficient and the lidar ratio increased up to 800 Mm^{-1} and 80 sr (at 355 nm), respectively. On 24 September, the aerosol optical depth, determined by an AERONET (Aerosol Robotic Network) sun-photometer, was as high as 0.9 and 1.1 (daily averages at 500 and 340 nm). HYSPLIT (Hybrid Single-Particle Lagrangian Integrated Trajectory) trajectories indicate air mass transport from Canary Islands to Mindelo at heights below 2 km. The observed pollution in the PBL over Mindelo is attributed to sulfate aerosol from the volcanic eruption at Las Palmas as the particle linear depolarization ratio was low ($\leq 3\%$) and, thus, does not indicate non-spherical particles, such as Saharan dust or volcanic ash. We thus conclude that sulfate aerosol formed from gaseous precursors during the transport (2–3 days for a distance of 1500 km) from Las Palmas towards Cabo Verde. No indications of volcanic ash over Mindelo were found in the SAL. This finding is supported by the HYSPLIT trajectories, which show that air masses in higher altitudes originate from the African continent and not from the Canary Islands.

1 Introduction

Volcanic eruptions are of large importance for the Earth's climate (Hansen et al., 1997; Robock, 2000) because the emitted particles and gases can be transported several hundreds of kilometers away from the source and influence the global radiation budget (Solomon et al., 2011; Groß et al., 2012; Martin et al., 2014). Typically emitted products of volcanic activity are ash particles with a diameter smaller than 2 mm during explosive phases, as well as volatiles such as sulfur dioxide (SO_2);

McGonigle et al., 2004; Aiuppa et al., 2008; Carracedo et al., 2022). SO₂ is the most abundant gas emitted by volcanoes (Kampouri et al., 2021) of which 10–20Mt are released into the troposphere each year (Martin et al., 2014). While in the stratosphere this gas has a lifetime of multiple weeks, it persists in the troposphere for around 1–3 days (Navas-Guzmán et al., 2013; Pattantyus et al., 2018). In a chemical reaction with water and further atmospheric components (hydroxyl radical (OH) in clear air conditions or hydrogen peroxide (H₂O₂) in cloudy air), it is quickly converted to sulfate aerosol (SO₄²⁻ bearing substances and sulfuric acid droplets; Ansmann et al., 2011b; Martin et al., 2014; Pattantyus et al., 2018). The efficiency of the conversion of SO₂ to sulfate aerosol is influenced by multiple factors and increases with temperature and relative humidity (Eatough et al., 1994; Yang et al., 2018). The lifetime of sulfate aerosol in the troposphere of 1 to 3 weeks is much longer than the one of SO₂ or volcanic ash so that it can be transported over long distances (Pappalardo et al., 2004; Filonchyk et al., 2022). If it reaches the higher troposphere/lower stratosphere it can remain even for several years (Jäger, 2005; Deshler, 2008; Martin et al., 2014).

Sulfate aerosol particles ~~are impacting the climate in several ways~~~~have several climate impacts~~ since they reflect solar radiation (Pappalardo et al., 2004) and scatter light even more efficiently with increasing relative humidity due to hygroscopic growth (Miffre et al., 2012). Furthermore, they act as cloud condensation nuclei (CCN) and ice nucleation particles (INPs) and, thus, influence the precipitation cycle (Pappalardo et al., 2004). Especially in cities, sulfate aerosol is of large importance with regard to air quality. It is one of the major components of urban PM_{2.5} (Zhang et al., 2004; Yang et al., 2018). Although anthropogenic SO₂ emissions are with 110Mt per year 5–10 times higher than volcanic emissions, volcanic eruptions are one of the greatest natural sources for sulfur emissions (Martin et al., 2014). Furthermore, their emissions have a larger ~~impact on the climate~~~~climate impact~~ due to the release of SO₂ in higher altitudes, which provides a longer lifetime of the formed aerosol particles (Kampouri et al., 2021). In addition, sulfate particles can be emitted directly as well (Martin et al., 2014). Moreover, volcanic eruptions have not only climatological but also economical ~~consequences~~~~impact~~ with regard to aviation. For example, volcanic ash can cause engine damage at aircraft or ~~the~~ air traffic is even suspended, as happened during the eruption of the Iceland volcano Eyjafjallajökull in spring 2010 (Groß et al., 2012). To reduce the risks, ash-dispersion simulations are used in early warning systems. Assimilation of satellite products like Aeolus wind measurements improve the ash plume forecast, as shown in a recent study of Amiridis et al. (2023). Besides the climatological and economical consequences, volcanic gases and particles regionally lead to strong pollution events, too, ~~so that in fact~~ during volcanic eruptions villages in the proximity even have to be evacuated. Often, the visibility is reduced and extremely unfavorable air quality is caused (Pattantyus et al., 2018). As a dominant component of PM_{2.5}, sulfate aerosol has a negative impact on human health as it infiltrates deeply into the lung and can cause asthma, sinusitis or further respiratory disease (Businger et al., 2015).

Lidar observations have expanded our knowledge on volcanic aerosol in the troposphere. In the case of the eruption of Eyjafjallajökull in Iceland in 2010 (Ansmann et al., 2010, 2011b; Groß et al., 2012; Pappalardo et al., 2013) and Etna in Italy in 2002 (Pappalardo et al., 2004) and 2019 (Kampouri et al., 2021), pure volcanic ash was observed. Lidar ratios in the range of 30–60sr and a particle linear depolarization ratio of 35–37% were measured at 355 and 532nm. In general, it is challenging to distinguish volcanic ash from other depolarizing aerosol types, especially from desert dust, because of the very similar lidar ratios of both types. The main quantity for the distinction between volcanic ash and desert dust is the particle linear

depolarization ratio, which is in the range of $30 \pm 5\%$ (at 355, 532, 710 and 1064 nm) for pure dust (Ansmann et al., 2010) and, thus, smaller than the aforementioned values for ash. Sulfate aerosol instead can be distinguished more easily from volcanic ash due to a much lower particle linear depolarization ratio, which is close to zero, and the different size ranges of the aerosol particles. While volcanic ash is in the coarse mode (diameter $> 2 \mu\text{m}$), sulfate aerosol is in the fine mode (diameter $\leq 2 \mu\text{m}$) (John et al., 2011). ~~(fine mode (diameter $\leq 2 \mu\text{m}$) for sulfate aerosol and coarse mode (diameter $> 2 \mu\text{m}$) for volcanic ash; John et al., 2011)~~ A separation of volcanic sulfate and ash based on the particle linear depolarization ratio was successfully introduced by Ansmann et al. (2011b). Sulfate particles produce a larger lidar ratio (55 up to 80 sr) and a particle linear depolarization ratio close to zero (4–5%) as multiwavelength-Raman lidar observations at 355 and 532 nm during the eruptions of Eyjafjallajökull and Etna have shown (Pappalardo et al., 2004; Mona et al., 2012; Navas-Guzmán et al., 2013). In the case of Eyjafjallajökull, Navas-Guzmán et al. (2013) observed two distinct aerosol layers over Granada, Spain, consisting of ~~to~~ 82% of sulfate aerosol. Sulfate aerosol from Eyjafjallajökull mixed with continental aerosol was furthermore observed in the planetary boundary layer (PBL) ~~PBL~~ over Potenza, Italy, (Mona et al., 2012). One of the first multiwavelength-Raman lidar ~~measurements~~ observations of tropospheric volcanic aerosol, and especially sulfate particles (mixed with a low amount of soot), was also performed at Potenza by Pappalardo et al. (2004), capturing the eruption of Etna in 2002.

One of the most recent volcanic eruptions, which was highly present in the European media, took place at the Cumbre Vieja volcanic ridge (28.62°N , 17.88°W ~~28.34°N , 17.50°E~~ , 1949 m a.s.l. (Filonchik et al., 2022)) at La Palma ~~Las Palmas~~, Canary Islands. The event is described in detail by Carracedo et al. (2022). Further studies concerning its impact on air quality were performed by Filonchik et al. (2022) and Milford et al. (2023). ~~A ceilometer-based study of the mass concentration of volcanic ash at La Palma and its distribution to the south of France was performed by Bedoya-Velázquez et al. (2022).~~ Volcanic activity started on 19 September 2021. The last eruption was recorded on 13 December 2021. ~~The volcanic activity during this period was classified as a basaltic fissure eruption dominated by strombolian activity and with episodic phreatomagmatic pulses.~~ The eruptive column usually reached to an altitude of 3500 m a.s.l. and peaked at 8500 m a.s.l. on 13 December. During the whole time of volcanic activity, fine lapilli (diameter 2–64 mm) were constantly produced. In addition, ash ($< 2 \text{ mm}$) and more than 10 kt SO_2 per day were emitted (Filonchik et al., 2022) so that at different measurement sites at La Palma ~~locally~~ the European air quality hourly threshold of $350 \mu\text{g m}^{-3}$ was exceeded on multiple days (Milford et al., 2023). The SO_2 emissions were largest at the beginning of the period with a maximum of 125 kt on 23 September 2021 (Milford et al., 2023). According to SO_2 -dispersion forecasts (Carracedo et al., 2022), the emission products were transported over long distances reaching Central Europe and the Caribbean. ~~A mixture of fine and coarse mode aerosol originating from La Palma was detected at Toulouse, France, on 24–25 September 2021 (Bedoya-Velázquez et al., 2022).~~

Volcanic aerosol of this eruption was also transported towards Mindelo on the Cabo Verdean Islands, which are located 1500 km southwest of the Canary Islands. Since June 2021, the multiwavelength-Raman-polarization lidar Polly^{XT} (Engelmann et al., 2016; Baars et al., 2016) has been operated there and was able to ~~capture~~ observe the volcanic aerosol plume. On 24 September 2021 ~~On specific days~~, the volcanic particles caused a high aerosol optical depth (AOD) of around 1.0, as measured with the co-located sun photometer, and a strong pollution in the PBL with extinction coefficient values more than twice as much as the typical background conditions ~~leading to~~ and a highly reduced visibility. ~~Since Cabo Verde is located in the trade~~

~~wind zone with preferred direction of inflow from north-east, i.e., from the Canary Islands, w~~We will show that the volcanic aerosol reached the measurement site at a low altitude and, thus, had a significant relevance with regard to air quality and human health~~impacts~~. In this paper, we present a case study of lidar observations conducted on ~~the~~ 24 September 2021 (period of volcanic activity) at Mindelo, Cabo Verde ~~contrasted to a reference measurement from 16 September 2021, before the start~~ of the eruptions. ~~The lidar observations indicate a strong pollution by volcanic sulfate aerosol.~~ In the following section, the methodology is described, including information about the instruments and models, the measurement site and the ~~methodway~~ of data processing. In Sect. 3, the results for the case study are presented and discussed in Sect. 4, before a conclusion is drawn in Sect. 5.

2 Methodology

130 2.1 Measurement site and instrumentation

In the frame of the ground-based part (ASKOS) of the Joint Aeolus-Tropical Atlantic Campaign (JATAC) (Amiridis et al., 2022; Fehr et al., 2023; Marinou et al., 2023), a temporary ACTRIS (Aerosol, Clouds and Trace Gases Research Infrastructure) ~~remote sensing~~ station was set up at the Ocean Science Center Mindelo (OSCM) at Mindelo, Cabo Verde, (16.878°N, 24.995°W ~~E~~) in June 2021. The OSCM is located on the west coast of the island of São Vicente, ~~with lowso-that-the~~ anthropogenic influence ~~is low~~. The island itself is located 1500 km southwest of the Canary Islands and ~~La PalmaLas Palmas~~ and in the trade wind zone with usual advection of air masses from north-easterly direction. Typically, cumulus convection occurs at Mindelo during nighttime.

Amongst others, this station is equipped with a Polly^{XT} multiwavelength-Raman-polarization lidar. The lidar deployed at Mindelo has a few improvements, compared to previous instruments (Althausen et al., 2009; Engelmann et al., 2016). For instance, it uses a diode pumped Nd:YAG laser, which has a higher repetition rate (100Hz) than the typical ~~flashlamp pumped~~ Nd:YAG laser (20–30Hz) of Polly^{XT}. This feature offers the possibility to retrieve profiles of the optical properties with a lower temporal averaging down to 10 min. It is an important capability with regard to this study, since at Mindelo ~~often~~ small clouds ~~often~~ occur ~~atduring~~ night and cloud-free periods are quite short. Furthermore, the receiver consists of 15 channels and enables measurements of the elastic backscatter coefficient at 355, 532 and 1064 nm, the inelastic backscatter at 387, 607 and 145 1058 nm, the cross-polar signal at 355, 532 and 1064 nm and the inelastic signal from water vapour at 407 nm. Additionally ~~to the far-field (ff) measurements~~, near-field (nf) measurements are available for the 355 and 532 nm elastic channels and the 387 and 607 nm Raman channels. The instrument has also a dual-field-of-view depolarization channel (Jimenez et al., 2020a) (~~Jimenez et al., 2020b~~), which is a powerful technique, allowing the determination of microphysical liquid-water properties (Jimenez et al., 2020b). Thus, in combination with the lidar-derived aerosol optical properties, it can be used to study aerosol-150 cloud interactions, which is, however, not the scope of the study we are presenting here. With the described~~is~~ setup, several aerosol optical ~~properties~~quantities can be determined. These are the particle backscatter coefficient, the particle extinction coefficient, the lidar ratio (ratio of particle extinction-to-backscatter coefficient) and the particle linear depolarization ratio, all at 355, 532 and 1064 nm, as well as the backscatter-related Ångström exponent between the different wavelengths and

the extinction-related Ångström exponent. The availability of the extinction coefficient, the lidar ratio and the particle linear depolarization ratio at 1064 nm as well as the backscatter coefficient at this wavelength determined via the rotational Raman (RR) channel is a new feature of this device. The calculation of the extinction coefficient at 1064 nm via the rotational Raman method follows the methodology described in Haarig et al. (2016). The spectral cross-talk calibration using a liquid cloud with a constant cloud base height was introduced in Haarig et al. (2022). Here, a liquid cloud base on 4 October 2021 was used for calibration, which led to a spectral cross-talk correction factor of $6.7e-4 \pm 0.3e-4$. There was no change in neutral density filters between 24 September and 4 October 2021 and therefore the spectral cross-talk correction factor remained stable. The calibration of the depolarization ratio at 1064 nm and the estimation of its uncertainties followed the same approach as the calibration at 355 and 532 nm (Engelmann et al., 2016). The $\Delta 90^\circ$ calibration (Freudenthaler et al., 2009) with a linear polarizer after the pinhole was applied. ~~A liquid cloud base on 4 October 2021 was used to estimate the cross talk of the elastic signal in the rotational Raman signal at 1058 nm (Haarig et al., 2016, 2022).~~ Like every Polly^{XT} lidar system, the Polly^{XT} lidar that operates at Mindelo is part of Polly^{NET} (Baars et al., 2016) and vertical profiles of the lidar optical properties are automatically derived by the Polly^{NET} processing chain (Yin and Baars, 2021). However, for this study, the profiles were analyzed manually. Due to the frequent occurrence of boundary layer clouds, a more tailored data analysis was needed. ~~Otherwise, optical profiles were not derived by the automatic processing chain having a strict cloud screening and other quality control measures.~~ Furthermore, to reduce noise, the profiles were smoothed vertically by using a moving average filter. The result is again a continuous profile with the distance of 7.5 m between the single data points but starting at an altitude, which is half of the smoothing length (s). Thus, each data point contains information of the height range from $0.5s$ below to $0.5s$ above this point.

Concerning the uncertainties of the lidar-derived aerosol optical properties, it is worth to mention that systematic errors (e.g., polarization effects in the receiver unit, dead time effects, overlap effects) are generally corrected as the lidar system is calibrated according to ACTRIS/EARLINET standards (e.g., ACTRIS guidelines, 2024). To account for remaining unknown systematic errors (e.g. reference height and value) and the statistical uncertainties, a relative error of 15 % is considered for the particle backscatter coefficient determined with the Raman method (Althausen et al., 2009; Baars et al., 2012; Engelmann et al., 2016; Baars et al., 2016). For the particle extinction coefficient, the statistical error is calculated from the error of the linear fit of the derivative without considering systematic uncertainties. For the particle depolarization ratio, we consider a remaining constant absolute error of 0.02 at 355 nm and of 0.01 at 532 and 1064 nm retrieved after intense calibration approaches (ACTRIS guidelines, 2024). The uncertainties of the lidar ratio and the Ångström exponents were then calculated using the Gaussian error propagation.

In addition, a CIMEL Sun Sky Lunar photometer of type CE318-T was used for this study, which is operating in the Aerosol Robotic Network (AERONET; Holben et al., 1998). It measures solar irradiances at 8 different wavelengths (340, 380, 440, 500, 675, 870, 1020 and 1640 nm) from which the AOD (at the same wavelengths), the columnar Ångström exponent (for 6 wavelength pairs), the volume size distribution, the refractive index, the single scattering albedo, the absorption AOD, the extinction AOD, the asymmetry factor, and the phase function are derived. A new capability of the latest type CE318-T is that it measures during night as well using the moonlight to determine nighttime AODs.

2.2 Air mass source attribution

To describe the origin of the observed air masses, Hybrid Single-Particle Lagrangian Integrated Trajectories (Stein et al., 2015; Rolph et al., 2017; HYSPLIT, 2024) were used. Ensemble trajectories with 27 members were calculated for 5 days back in time, i.e., towards the day when volcanic activity at [La PalmaLas Palmas](#) started. The meteorological input data was taken from the Global Data and Assimilation Service (GDAS1, 2024). Furthermore, simulations with [the air mass source attribution tool TRACE](#) (Radenz et al., 2021; Radenz, 2021) were performed, which is a combination of the FLEXible PARTicle dispersion model FLEXPART (Pisso et al., 2019) and a simplified version of the MODIS land cover classification (Broxton et al., 2014) or custom defined geographical areas. For FLEXPART, the meteorological input data was taken from the Global Forecast System (GFS; National Centers for Environmental Prediction, National Weather Service, NOAA, U.S. Department of Commerce (2000)).~~the Lagrangian dispersion model FLEXPART (Radenz et al., 2021) were performed using meteorological data fields from the Global Forecast System (GFS; National Centers for Environmental Prediction, National Weather Service, NOAA, U.S. Department~~
In this case, 5-day backward simulations were calculated for 500 air parcels, which arrive at Mindelo at different altitudes from 0 to 10km in steps of 500m with a temporal resolution of 3h. Evaluating both backward simulation models allows us to ensure more certainty with respect to the origin of the air masses.

In addition, the [horizontal](#) distribution of the volcanic plume was monitored. Therefore, the transport of SO₂ and its advection towards Mindelo was tracked. For this purpose, the TROPOspheric Monitoring Instrument (Veefkind et al., 2012; TROPOMI, 2024) on board the polar orbiting Sentinel-5 Precursor satellite [was used](#), which offers daily global measurements of the amount of SO₂ molecules in a column per surface area. Its horizontal resolution is $3.5 \times 5.5 \text{ km}^2$.

3 Results

A time series of the [AERONET](#) AOD at different wavelengths and the columnar Ångström exponent between 440 and 870nm is shown in Fig. 1. Before the start of the eruption, the hourly mean AOD was [around about 0.4 at the shown wavelengths](#). Hourly mean Ångström exponent values of 0.2 were usually observed until 22 September. During the time of volcanic activity, a change in the behaviour of the Ångström exponent and the AOD could be seen since 22 September and, thus, 3 days after the eruption started. A strong increase of the Ångström exponent to values higher than 0.8 was measured. In that time, high AODs, ~~with~~ values close to 1.0 [at wavelengths \$\leq 500 \text{ nm}\$](#) were recorded, e.g., on 24 and 29 September 2021. On 24 September, the daily mean AOD was 1.1 at 340nm and 0.9 at 500nm. The vertically-resolved lidar optical properties are presented in a case study for 24 September (Sect. 3.2) contrasted to background conditions before the volcanic eruption (16 September, Sect. 3.1), representing a clean [PBLplanetary boundary layer](#) (marine influenced).

3.1 Reference case (16 September 2021)

To contrast the differences between the volcanic influenced aerosol conditions over Mindelo and the typical situation before the start of the eruption at [La PalmaLas Palmas](#), the 16th September 2021 was selected as reference observation. The corresponding

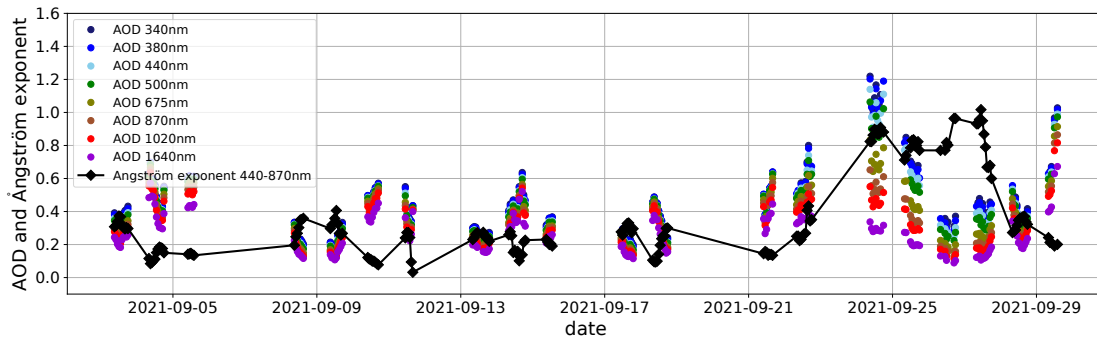


Figure 1. Time series of Level 2.0 hourly averages of the aerosol optical depth (AOD) at different wavelengths and the columnar Ångström exponent between 440 and 870nm measured with an AERONET sun photometer at Mindelo in September 2021. Data points for the 16 September were cut out, since on that day, a cirrus was present, which was not correctly screened out by the AERONET retrieval. The original data, including the cirrus, is shown added in the Appendix (Fig. A1). ~~The data was taken from AERONET (2024)~~

height-resolved temporal development of the calibrated attenuated backscatter coefficient at 1064 nm and the volume depolarization ratio at 532 nm is shown in Fig. 2 (left side). The figure is provided to illustrate the vertical structure of the aerosol layers for the days of interest. As it is not corrected for the atmospheric attenuation, it does not allow a quantitative comparison of the backscatter intensity at a given altitude. For this purpose, vertical profiles of the backscatter coefficient are presented later on. The vertical structure on the 16 September showed two different aerosol layers. During that day, two different aerosol layers were present. The PBL planetary boundary layer reached up to 0.8 km height. In that layer, no depolarization occurred (Fig. 2b). Above, a lofted layer, which was strongly depolarizing, was located between 1.2 and 6 km height. Small clouds were frequently present in the PBL as indicated by a very strong backscatter signal and complete attenuation (no signal) some few hundreds of meters above the cloud base. They can be identified in Fig. 2a as red spots above which stripes of missing signal are visible. Such a vertical structure This layering has been typically observed over Mindelo from June to October 2021 and is in agreement with previous studies on this archipelago (Ansmann et al., 2011a; Groß et al., 2011; Rittmeister et al., 2017). Additionally, on the reference day, a cirrus occurred, which was not correctly screened out by the AERONET algorithm (cf. Fig. A1 and Fig. A2). Thus, there are no usable sun photometer data for the 16 September.

Vertical profiles of the lidar optical properties (Fig. 3) were derived with the Raman method (Ansmann et al., 1992) for a 48-min interval in the evening (22:24–23:12 UTC, red rectangle in Fig. 2, left side), since this was the longest cloud-free period during nighttime. The corresponding mean values are summarized in Table 1. The uncertainties given in the text are always the standard deviation (parameter variability within the layer) or the layer mean error as described in the caption of Table 1. In the lofted layer, between 1.3 and 5.3 km height, the mean lidar ratio of 58.4 ± 8.8 and 47.3 ± 7.2 sr (at 355 and 532 nm) and the particle linear depolarization ratio of 24.5 ± 2.0 , 28.1 ± 1.0 and $24.1 \pm 1.0\%$ (at 355, 532 and 1064 nm) are in a typical range for desert dust (Haarig et al., 2017; Floutsi et al., 2023). In the PBL (up to 0.6 km height), relatively clean marine conditions (Bohmann et al., 2018) were observed, which were characterized by a mean lidar ratio of 17.3 ± 2.8 and 23.8 ± 4.2 sr 23.7 sr

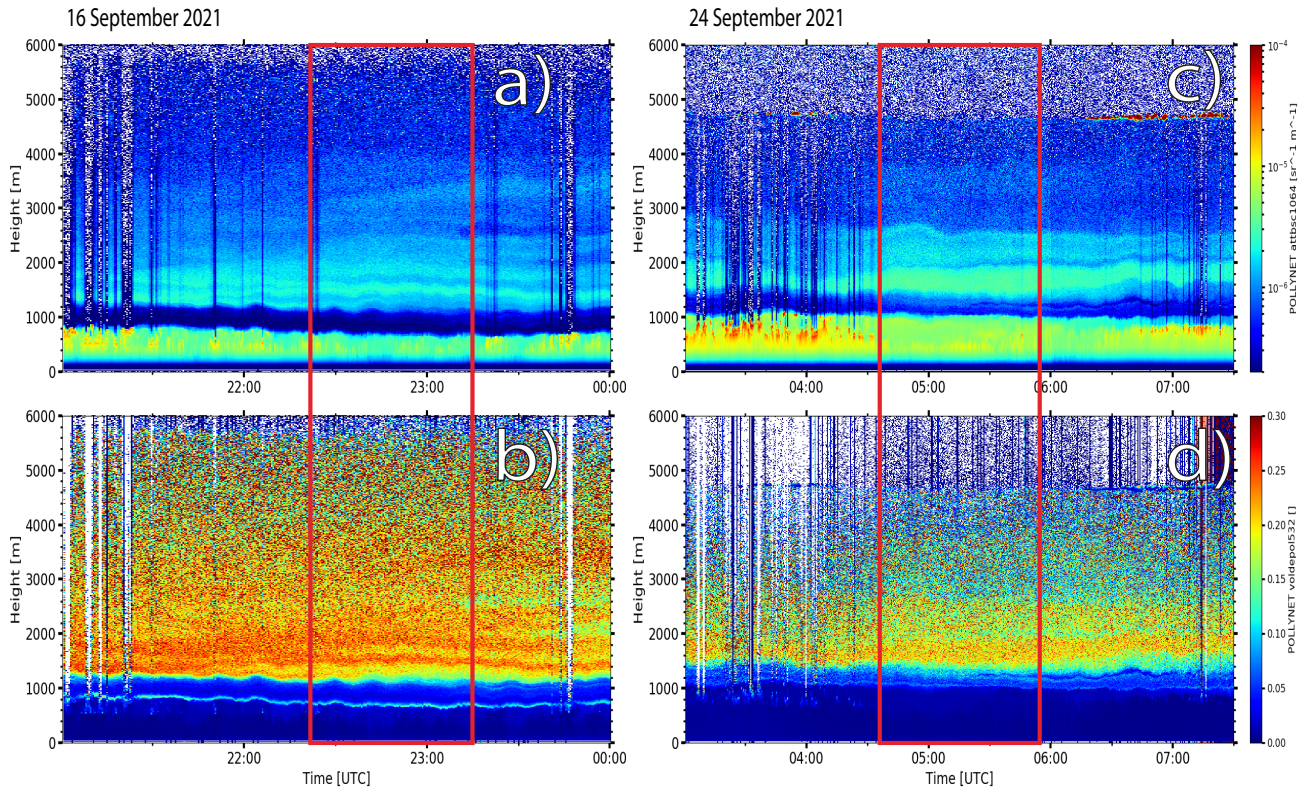


Figure 2. Temporal evolution of the height-resolved **calibrated** attenuated backscatter coefficient at 1064 nm (top) and the volume depolarization ratio at 532 nm (bottom) measured by Polly^{XT} at Mindelo, Cabo Verde, during 16 September 2021, 21:00–24:00 UTC, (left) and 24 September 2021, 03:00–07:30, (right).

240 (at 355 and 532 nm) **between 0.25 and 0.6 km height** and a mean particle linear depolarization ratio of $\leq 1.1\%$ **0.7, 1.1 and 1.0%**
 (at 355, 532 and 1064 nm) **between 0.06 and 0.6 km height**. The mean particle extinction coefficient was about 114 ± 20 and
 $130 \pm 24 \text{ Mm}^{-1}$ (at 355 and 532 nm) **between 0.25 and 0.6 km height**. Unfortunately, for that day the rotational Raman profiles
 at 1064 nm were not available since the analyzed time period of 48 min is too short to obtain reasonable results. However, the
 measurement from **the** 16 September represents the typical values, which we usually observed over Mindelo during that time
 245 of the year, as **the** lidar studies in the framework of ASKOS and L2A+ confirm (L2A+, 2024; EVDC, 2024). This statement is
 valid especially for the PBL.

3.2 Volcanic influence (24 September 2021)

For **the-the** 24 September 2021, the height-resolved temporal development of the attenuated backscatter coefficient at 1064 nm
 is shown in Fig. 2c and the volume depolarization ratio at 532 nm in Fig. 2d. Again, two distinct aerosol layers are visible – a

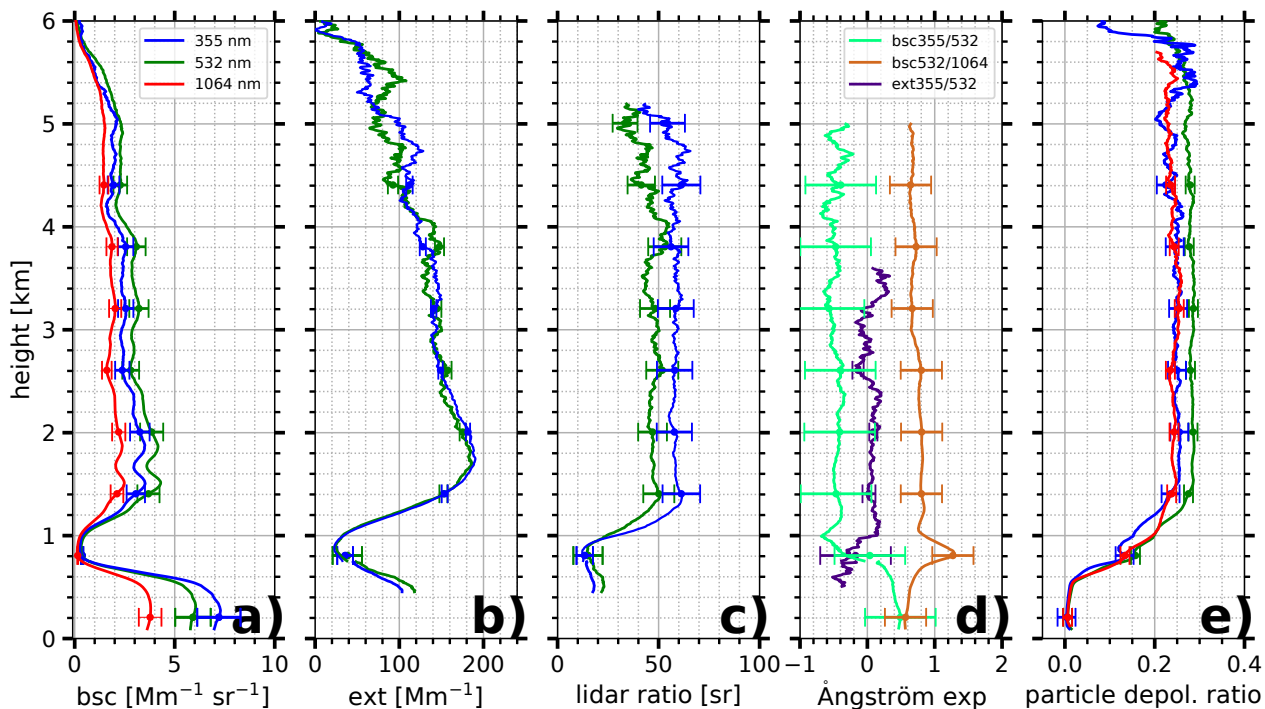


Figure 3. Measured with Polly^{XT} at Mindelo, Cabo Verde, on 16 September 2021, between 22:24 and 23:12 UTC: vertical profiles of (a) the particle backscatter coefficient, (b) the particle extinction coefficient, (c) the lidar ratio, (d) the Ångström exponent and (e) the particle linear depolarization ratio. Vertical smoothing: 187.5 m for (a), (e) and the backscatter-related Ångström exponent and 742.5 m for (b), (c) and the extinction-related Ångström. Near- and far-field measurements are merged at 750 m. The error bars show the uncertainties described in Sect. 2.1. Backscatter and depolarization profiles as well as the backscatter-related Ångström exponent are plotted with a vertical smoothing of 187.5 m. The vertical smoothing for the remaining quantities is 742.5 m. Near- and far-field measurements are merged at 750 m.

250 very low depolarizing PBL (Fig. 2d) up to about 1 km height and a strongly depolarizing lofted layer from 1.4 to 5 km height. As on 16 September, often small clouds occurred frequently in the PBL. Before first daylight appeared at 07:30 UTC, a longer cloud free period evolved. Thus, optical properties were retrieved with the Raman method for an 1:19 h interval (04:38–05:57 UTC, indicated by a red rectangle in Fig. 2, right side).

The corresponding vertical profiles are shown in Fig. 4. On that day, all lidar-derived optical quantities are available at all 255 three wavelengths. For the lofted layer, mean values, as depicted in Table 1, were retrieved based on the far-field measurements considering a vertical extent of 1.4–4.4 km and a vertical smoothing as in Fig. 4. The particle backscatter coefficient varied from $0.8 \text{ Mm}^{-1} \text{ sr}^{-1}$ (minimum at 355 nm) to $5.0 \text{ Mm}^{-1} \text{ sr}^{-1}$ (maximum at 532 nm) and the particle extinction coefficient was in the range of $114\text{--}168 \text{ Mm}^{-1}$ at 355, 532 and 1064 nm. from 16 Mm^{-1} (minimum at 532 nm) to 246 Mm^{-1}

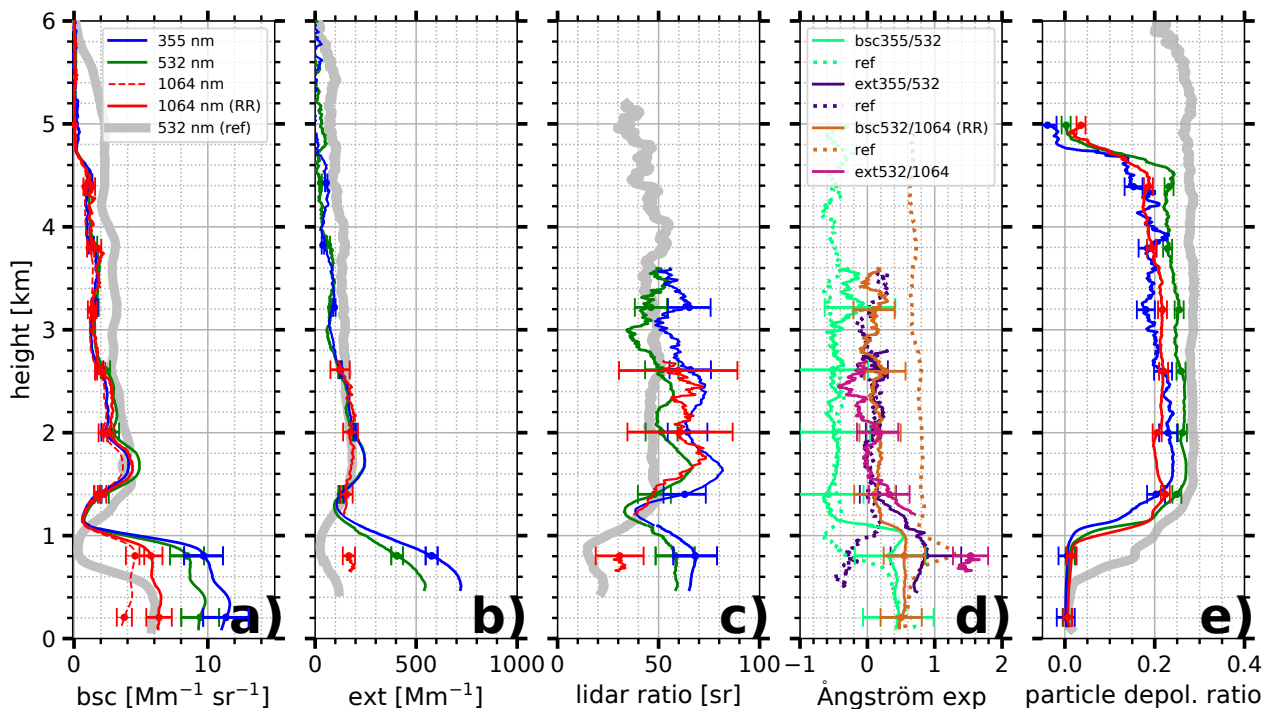


Figure 4. Same as in Fig. 3 but for the 24 September 2021, between 04:38 and 05:57 UTC (note the changed x-scale for (a) and (b)), including the particle extinction coefficient and lidar ratio at 1064 nm and the extinction-related Ångström exponent between 532 and 1064 nm (vertical smoothing: 397.5 m below 1.2 km and 1492.5 m above 1.2 km). Near- and far-field measurements are merged at 1100 m. Reference profiles at 532 nm from 16 September 2021 are shown as thick grey line and in (d) as dotted lines and labelled as "ref". For that day, profiles of the particle extinction coefficient and the lidar ratio are available at 1064 nm, too, as well as the extinction-related Ångström exponent between 532 and 1064 nm. For them, the vertical smoothing is 397.5 m (below 1.2 km) and 1492.5 m (above 1.2 km). Near- and far-field measurements are merged at 1100 m. Additionally, the reference profiles from 16 September 2021 at 532 nm (thick grey line and dotted lines for the Ångström exponent and labelled as "ref") are shown, representing the typical aerosol conditions over Mindelo with a clean marine PBL.

(maximum at 532 nm). Measurements of the lidar ratio led to layer mean values of 64.8 ± 10.2 , 50.9 ± 8.3 and 61.8 ± 8.6 sr
 260 65.3 , 50.7 and 68.7 sr (355, 532, and 1064 nm, respectively). These values are slightly larger than the ones measured
 on 16 September, which were similar compared to previous observations of pure Saharan dust (50–60 sr at 355 and 532 nm)
 (Ansmann et al., 2011a; Ansmann et al., 2011; Ansmann et al., 2023). The increase of the lidar ratio from 532 nm to 1064 nm
 by 21%–36% is in line with dust observations at Leipzig, Germany (increase by 24–38%; Haerig et al., 2022), because of similar
 source regions of the dust in the Western Sahara of both observations. The enhanced lidar ratio at 355 nm compared to 532 nm
 265 indicates specific source regions in the Sahara as reported from lidar observations in Senegal (Veselovskii et al., 2020). The

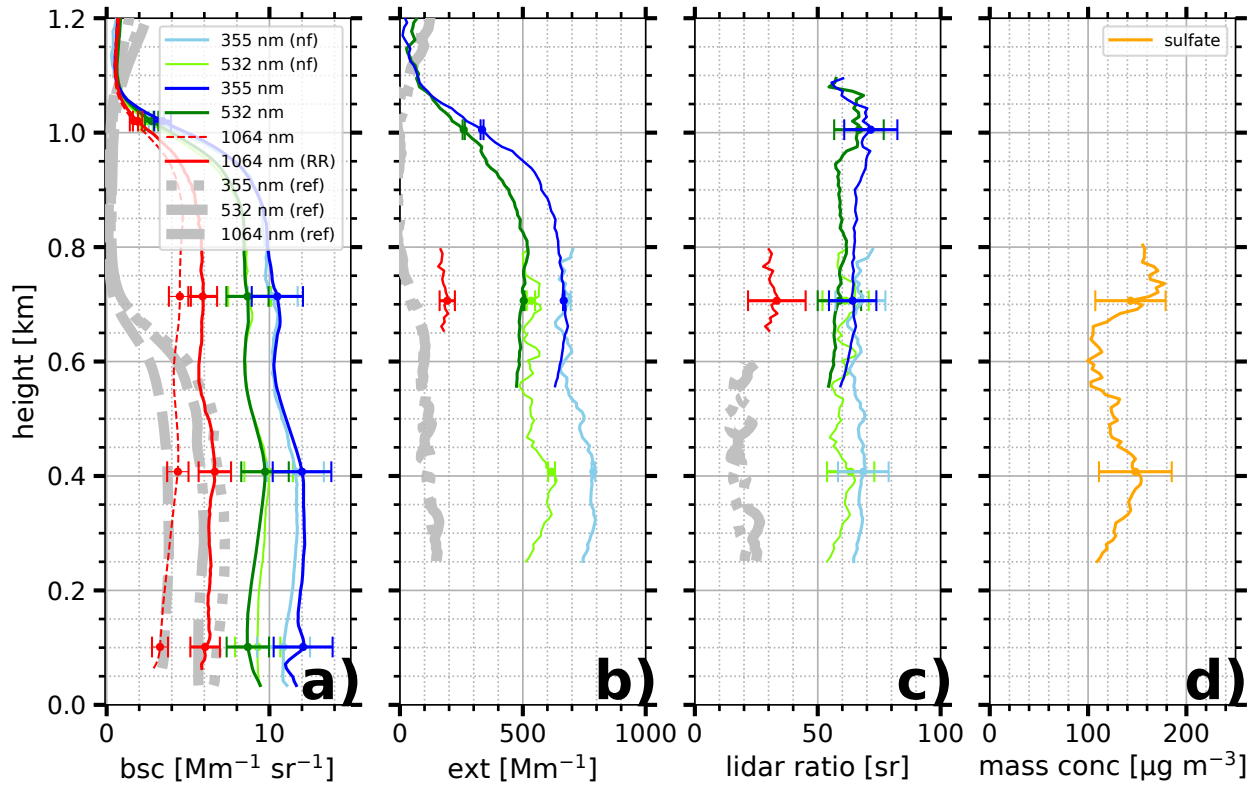


Figure 5. (a)–(c): same profiles as in Fig. 4 (a)–(c), but only up to a height of 1.2 km. Vertical smoothing: (a) 67.5 m, (b) and (c) 187.5 m at 355 and 532 nm and 742.5 m at 1064 nm. Near- and far-field measurements are shown separately. (d): sulfate mass concentration with height constant relative error of 25%. The same profiles as in Fig. 4 are shown here, but only for the PBL up to a height of 1.2 km. Furthermore, vertical smoothing was reduced to 67.5 m for the particle backscatter coefficient and to 187.5 m for the particle extinction coefficient and the lidar ratio at 355 and 532 nm and increased to 742.5 m for the particle extinction coefficient and the lidar ratio at 1064 nm. Near-field (nf) and far-field measurements are shown separately.

lidar ratio at 1064 nm is in line with dust observations at Leipzig, Germany (Haarig et al., 2022). From 532 nm to 1064 nm, the lidar ratio increased by 36%, which is in agreement with the previous observations at Leipzig (increase by 24–38%; Haarig et al. (2022)). The measured particle linear depolarization ratio of 20.6–25.0% for the three different wavelengths indicates the presence of non-spherical particles, i.e., desert dust, but is somewhat smaller than what was typically observed for pure dust (Freudenthaler et al., 2009; Floutsi et al., 2023), indicating the presence of some spherical non-dust particles. Consider-
 270 ing the wavelength dependence of the particle linear depolarization ratio, a decrease of 18% from 532 towards 1064 nm was observed. Similar findings were made at Leipzig, Germany, and Morocco during SAMUM (decrease by 13–31%; Freuden-

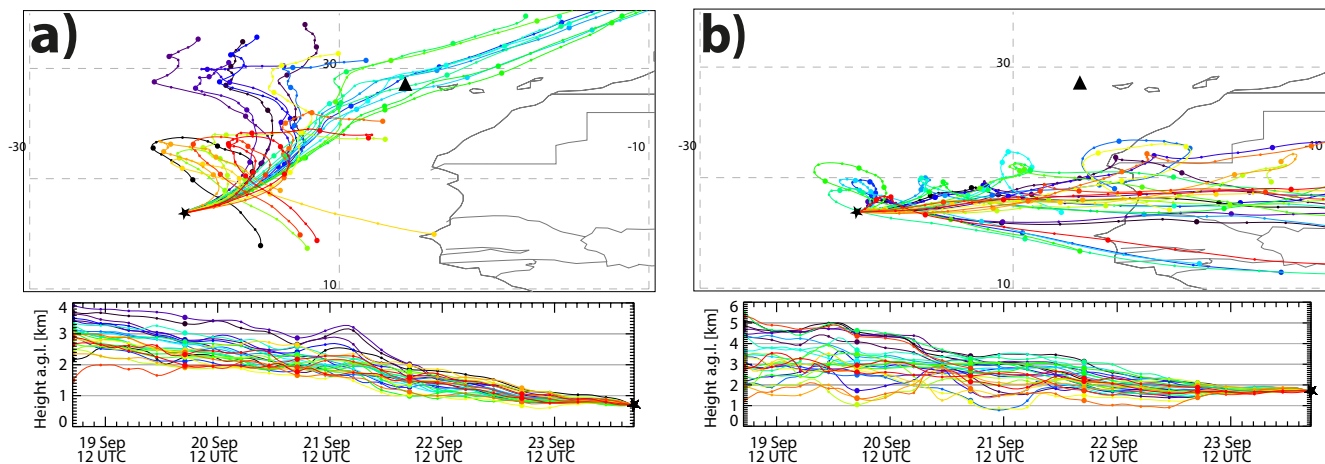


Figure 6. HYSPLIT ensemble trajectories for 120 hours back in time are shown. Backward trajectories of air masses arriving at Mindelo (black star) on 24 September 2021, 5 UTC at (a) 0.7 km and (b) 1.7 km were computed. In the lower altitude, they mainly originate from Las Palmas (black triangle), whereas in the higher altitude, they were advected from the Sahara.

thaler et al., 2009; Haarig et al., 2022). The backscatter-related Ångström exponent in the lofted layer is on average around 0.4 ± 0.31 for the wavelength pair 532/1064 nm, indicating large particles (i.e., desert dust). Considering the higher lidar ratio (especially at 355 nm) and the lower particle linear depolarization ratio on 24 September compared to the typical values of pure desert dust, we conclude that the dust on 24 September was slightly polluted.

In contrast to the almost typical aerosol conditions in the lofted layer, an unusual strong pollution was observed in the PBL. The findings are highlighted in Fig. 5 showing zoomed profiles. In addition, vertical smoothing was reduced, which improves the resolution, reduces the overlap effect especially for the near-field profiles and, thus, allows to include information from lower altitudes above the lidar accuracy of the near-field profiles. All mean values for the PBL are listed in Table 1 as well. On this day, extremely high values of the particle extinction coefficient were observed with layer mean values of 721 ± 51 , 549 ± 38 and $178 \pm 13 \text{ Mm}^{-1}$ (at 355, 532 and 1064 nm) in the PBL between 0.25 and 1 km. The maximum values were even higher with 794 ± 7 , 640 ± 13 and $198 \pm 26 \text{ Mm}^{-1}$ (at 355, 532 and 1064 nm, mentioned here with statistical errors). These values are 3–4 times higher than what was observed under clean marine conditions as shown for 16 September and indicated as grey lines in Fig. 5. Additionally, the particle extinction coefficient was strongly decreasing with increasing wavelength. A similar behavior was observed for the lidar ratio. Mean values of 66.9 ± 10.1 , 60.2 ± 9.2 and $30.8 \pm 8.7 \text{ sr}$ (at 355, 532 and 1064 nm) were found, showing a decrease by 48% from 532 towards 1064 nm. The mean values of the lidar ratio are notably quite high compared to the clean marine conditions and are typical for pollution or even smoke (Floutsi et al., 2023). However, the decrease of the lidar ratio at 1064 nm compared to the value at 532 nm points rather to pollution than to smoke. In the case of wildfire smoke an increase of the lidar ratio at 1064 nm was observed (Haarig et al., 2018). The high lidar ratio values point out the presence of particles, which are strongly attenuating the incoming solar radiation

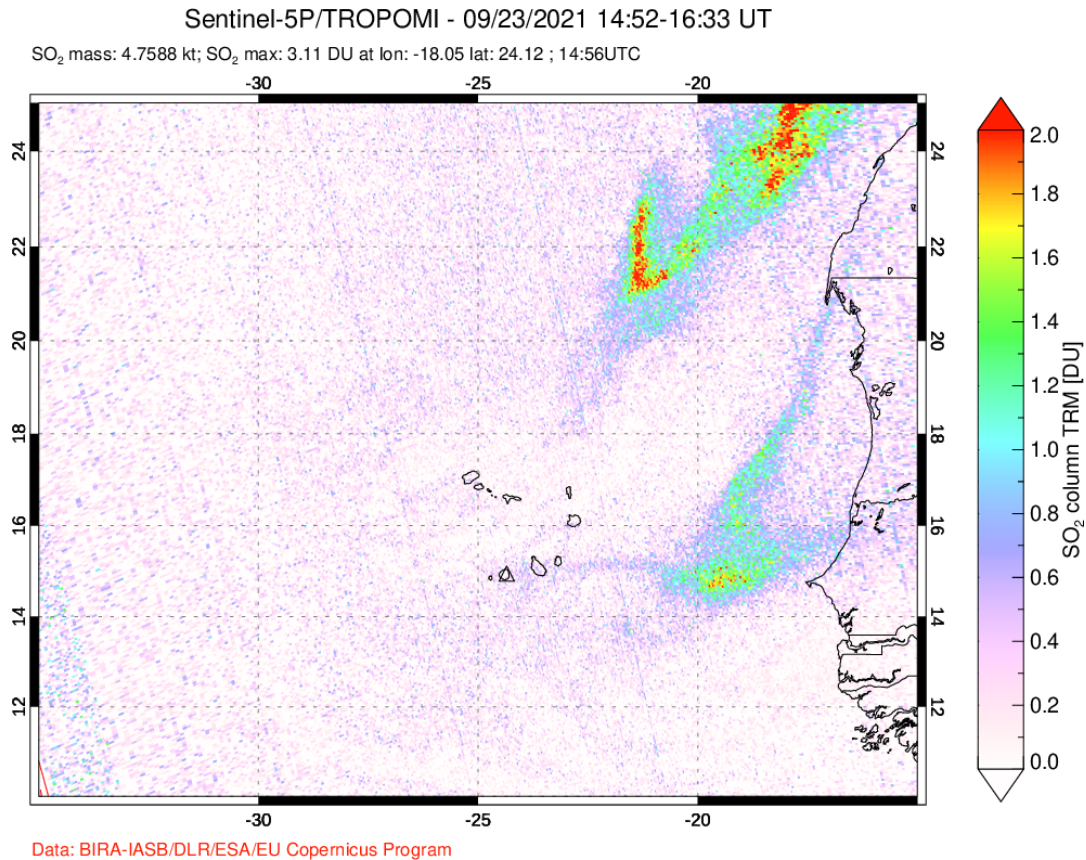


Figure 7. Satellite observations of TROPOMI on Sentinel-5P, which show the column integrated of SO₂ mass for the Cabo Verdean region on 23 September 2021 (TROPOMI, 2024).

by scattering and absorption (Wandinger et al., 2023). The large extinction close to the ground indicates a strong pollution and explains the unusual high daily mean AOD of 1.1, 0.9 and 0.51+and 0.9, which was measured with the sun photometer at 340, 500 and 1020 nm 340 nm and 500 nm on this day. The lidar-derived total AOD between 04:38 and 05:57 UTC was 0.96 ± 0.28, 0.79 ± 0.26 and 0.57 ± 0.17 at 355, 532 and 1064 nm, respectively, 1.07 (355 nm) and 0.86 (532 nm) and, thus, in agreement with the values of the sun photometer measured during daytime. The AOD for the boundary layer only (including the sulfate and marine contribution), as derived from the lidar measurements, was 0.58 ± 0.03, 0.43 ± 0.02 and 0.18 ± 0.01 at 355, 532 and 1064 nm 0.69 at 355 nm and 0.5 at 532 nm and, thus, covered 54–60% 58–64% of the total lidar-derived AOD in case of 355 and 532 nm and 32% in case of 1064 nm. Furthermore, visibility was strongly reduced on that day. Based on the maximum particle extinction coefficient at 532 nm and using the Koschmieder equation (Koschmieder, 1924), we calculated the visibility to be the visibility was around 6 km. The presence of relatively small particles is indicated further supported by the moderate wavelength dependence represented indicated by the backscatter-related Ångström exponent between 532 and 1064 nm (RR), which was

0.54 ± 0.31 and the mean extinction-related Ångström exponent of only 0.68 ± 0.07 . The low values of the particle linear depolarization ratio $\rho_{\parallel} \leq 0.9\%$ indicate that the observed particles are spherical.

305 4 Discussion

To summarize, both days – the 16th September 2021 (before the start of the eruption at [La Palma](#)) and the 24th September 2021 (during the volcanic episode) – had a similar aerosol layering structure with a PBL ≤ 1 km and a lofted layer of Saharan dust up to 6 km, typical for this time of the year at Cabo Verde. Both measurements were taken under similar meteorological conditions and, thus, well suitable to assess the influence of the volcanic activity.

310 Although for the lofted layer, the lidar-derived aerosol optical properties slightly vary concerning the extent and intensity of the lofted layer between the 16 and 24 September 2021 with a lower layer top height, a lower particle linear depolarization ratio and a higher lidar ratio on the 24 September, we can conclude that on both days the predominant aerosol type in the lofted layer was Saharan dust. Obviously, no volcanic ash was included in this layer on 24 September, because in that case we would have observed much higher values of the particle linear depolarization ratio (Groß et al., 2012). The higher lidar ratio
315 observed on 24 September, however, indicate the presence of stronger absorbing particles slightly contaminating the Saharan Air Layer (SAL) on that day. Considering the particle linear depolarization ratio and the lidar ratio together, a contamination with continental pollution or smoke is feasible. To corroborate the origin of the lofted aerosol layer on this day, 120 h HYSPLIT ensemble backward trajectories are shown in Fig. 6. Simulations of air mass arrival at Mindelo on 24 September, 5 UTC, at 0.7 km (Fig. 6a) and at 1.7 km (Fig. 6b) have been calculated. The trajectories for the higher altitude (Fig. 6b) show that the
320 lofted layer was influenced by an easterly flow so that air masses were advected directly from the Saharan desert, which makes the occurrence and predominance of Saharan dust evident. However, partial mixing with sulfate during the transport over the Atlantic Ocean cannot be ruled out as well as smoke and pollution contamination over the African continent. Fire spot analysis with the Fire Information for Resource Management System (FIRMS, 2024, Fig. B1) ([FIRMS, firms.modaps.eosdis.nasa.gov](#), not shown) revealed only little fire activity along the transport path within the 120 h. Fires were detected at the eastern border
325 of Algeria and close to the Mediterranean. Thus, smoke contamination may have led to the slight contamination of the SAL. However, we consider the presence of volcanic ash based on this analysis and the eruption mechanisms at Cumbre Vieja to be unlikely.

In contrast, the aerosol conditions in the PBL strongly differed between both analyzed measurement periods. While on 16 September, a clean marine PBL was present, a strong pollution was observed on 24 September with layer mean values of
330 the particle extinction coefficient up to almost 800 Mm^{-1} , compared to $\leq 130 \text{ Mm}^{-1}$ during the clean marine conditions. The lidar ratio on that day was strongly enhanced as well with values around 60 sr compared to values for pure marine conditions of around 20 sr. The aerosol load in the PBL was furthermore responsible for 54–60% of the total AOD at 355 and 532 nm and for 32% at 1064 nm in the analyzed time period. In contrast, for the 16 September, the contribution of the AOD in the PBL to the total AOD was only 12% (at 355 and 532 nm), i.e., more than 80% of the total AOD was caused by the SAL.
335 The observed pollution is associated to air masses coming from [La Palma](#) containing volcanic aerosol, which is

supported by the HYSPLIT backward trajectories depicted in Fig. 6a. They illustrate a distinct advection of air masses from Canary Islands and, thus, from the volcano on [La PalmaLas Palmas](#). Additionally, [TRACE backwardFLEXPART](#) simulations (Fig. C1) confirm these findings as they show air masses accumulating over the Atlantic Ocean on 18 September before they pass from northwest over [La PalmaLas Palmas](#) on 21 September and move further to Mindelo on 23 September. [TRACE backward simulation](#) were also performed for air parcels arriving at Mindelo at 500m altitude on 16 September 2021, 0 UTC, shown in Fig. C2. The pathway of the air parcels was similar to the one from 24 September. While on 16 September, they also partly crossed the African continent, they passed only over the Atlantic Ocean and the Canary Islands on 24 September. This finding emphasizes even more the impact of the volcanic eruption on the pollution observed in the PBL at Mindelo. As the [correspondingmeasured](#) particle linear depolarization ratio is low [in the PBL](#), the presence of ash particles can be excluded.

345 Instead, volcanic sulfate seemed to be the dominating aerosol type in the PBL. It becomes more evident if the large amount of sulfur dioxide released by the volcano is taken into account. The SO₂ emissions were greatest at the beginning of the active period, reaching a maximum of 125kt on 23 September 2021 (Milford et al., 2023). SO₂ was advected towards Mindelo as can be seen in the satellite measurements of TROPOMI onboard Sentinel-5P (Fig. 7), showing the amount of SO₂ around the Cabo Verdean region during afternoon of 23 September 2021. The presence of SO₂ offered the possibility for secondary aerosol formation to sulfate particles, which is assumed to be the source of the observed particles. SO₂ quickly oxidates to sulfate aerosol with a high efficiency at warm temperatures and high relative humidity (Eatough et al., 1994; Yang et al., 2018). [FavorableFavored](#) conditions seemed to be [prevalentprovided](#) since the air masses [wereare](#) transported only over the Atlantic Ocean in a tropic region. According to Pattanyus et al. (2018), conversion rates are large (3–50% s^{-1}) especially in cloudy air, which is given [as well](#) due to the frequently occurring small clouds in the PBL as observed over the Cabo Verdean region.

355 In addition, not only SO₂ could have been advected from [La PalmaLas Palmas](#) but also sulfate particles themselves. Filonchik et al. (2022) identified, based on the single scattering albedo and the dissection of the AOD into a coarse and fine mode component, that on 24 September 2021 coarse-mode particles were almost absent at [La PalmaLas Palmas](#). Instead the presence of non-absorbing but scattering fine-mode particles attributed to sulfate aerosol was shown, which could have been formed locally or were emitted directly by the volcano. As these observations are valid for a time period in which our case study was included, it strengthens our hypothesis that we measured volcanic sulfate at Mindelo originating from Cumbre Vieja.

The presence of sulfate aerosol from the eruption at [La PalmaLas Palmas](#) also becomes evident since the measured [aerosol optical propertiesquantities](#) are in agreement with previous lidar observations of volcanic sulfate (e.g., Pappalardo et al., 2004; [Mona et al., 2012; Navas-Guzmán et al., 2013](#)). Furthermore, in this study, it was the first time ever that tropospheric volcanic sulfate was measured with a lidar at 1064nm. During the eruption of Eyjafjallajökull in 2010, Navas-Guzmán et al. (2013) observed lofted aerosol layers between 1.5 and 3.5km consisting to 82% of fine-mode aerosol particles, i.e., sulfate particles, over Granada, Spain. The corresponding values of the lidar ratio were 55 and 75sr (355, 532nm). [Mona et al. \(2012\)](#) recorded values of the lidar ratio up to 80sr for volcanic sulfate from Eyjafjallajökull mixed with continental aerosol in the PBL over Potenza, Italy, while during the eruption of Mt. Etna in 2002 a lidar ratio of 55 ± 4 sr (355nm) was measured by Pappalardo et al. (2004) in a lofted aerosol layer of young sulfate particles mixed with a low amount of soot between 4 and 4.5km over

Potenza. With 66.9 ± 10.1 and 60.2 ± 9.2 sr (355, 532 nm), the observations over Mindelo on 24 September fit well into the range of values of the lidar ratio observed during the previous eruptions. Also, the observed wavelength dependence with decreasing lidar ratio by 48% from 532 towards 1064 nm confirms the assumptions in the CALIPSO aerosol typing which uses a lidar ratio of 50 sr at 532 nm and of 30 sr at 1064 nm for sulfate (Kim et al., 2018; Tackett et al., 2023). The particle linear depolarization ratio measured on 24 September 2021 was with $\leq 0.9\%$ even smaller than the values observed for sulfate from Eyjafjallajökull, which was 4–5% (Navas-Guzmán et al., 2013), indicating more clearly the presence of spherical (sulfate) particles. For the volcanic sulfate from Eyjafjallajökull, the backscatter-related Ångström exponent measured over Granada was 1.1 ± 0.2 for the wavelength pair 355/532 and 2.1 ± 0.1 for 532/1064. During the measurement period, the values decreased to 0.7 ± 0.1 and 1.7 ± 0.3 , respectively. Navas-Guzmán et al. (2013) stated that the temporal evolution in the Ångström exponent arose from an increasing particle size in the accumulation mode driven by hygroscopic growth and a potential change in the chemical composition rather than an increasing contribution of coarse-mode particles., due to hygroscopic growth of the sulfate particles. In a second layer of volcanic sulfate, values of 1.7 ± 0.1 (355/532) and 1.4 ± 0.2 (532/1064) were observed by Navas-Guzmán et al. (2013), also decreasing significantly during the measurement period. Compared to these studies, the backscatter-related Ångström exponent measured over Mindelo on 24 September 2021 was smaller having values of only 0.42 ± 0.52 and 0.54 ± 0.31 (355/532, 532/1064 (RR)). These low values can be explained by hygroscopic growth of the sulfate particles since they were exposed to high humidity during their transport over the Atlantic Ocean before they reached Mindelo. Furthermore, we expect that some marine particles were present in the PBL above Mindelo, which are also spherical at high relative humidity (accounting to the low depolarization ratio). As marine particles are larger than the sulfate particles, they reduced the backscatter-related Ångström exponent in contrast to the aforementioned observations where air masses were influenced by the European continent. The extinction-related Ångström exponent between 355 and 532 nm observed over Mindelo was 0.68 ± 0.07 and, thus, closer to the previous observations at Granada where values of 0.7 ± 0.1 and 0.8 ± 0.1 were observed.

As a further step, an attempt to estimate the mass concentration of the observed sulfate aerosol was done according to the method used in Ansmann et al. (2011b). A sulfate conversion factor of $0.2 \cdot 10^{-6}$ was obtained leading to the vertical profile of the mass concentration, which is shown in Fig. 5d, and a layer mean value for the PBL of about $133.1 \pm 20.3 \mu\text{g m}^{-3}$. Such a high sulfate concentration indicates extremely polluted conditions. For comparison, the aerosol pollution level should not exceed $50 \mu\text{g m}^{-3}$ in European cities to avoid unhealthy situations. For our estimation, the error of the sulfate mass concentration was assumed to be similar to the one of the dust mass concentration, which is about 20–30% (Ansmann et al., 2019). Thus, for our estimation, we used a height constant relative error of 25%.

In this paper, it is the first time ever that we can report the optical properties for the volcanic plume mixed in the marine PBL for all three (aerosol lidar) wavelengths by extending the observational capabilities towards 1064 nm. While the lidar ratios at 355 and 532 nm are in agreement with previous observations, the lidar ratio at 1064 nm of 30.8 ± 8.7 sr and the extinction-related Ångström exponent of 1.53 ± 0.26 between 532 and 1064 nm have never been reported so far. Thus, it is a milestone for the characterization of volcanic sulfates with remote sensing techniques.

405 5 Summary and conclusions

In the frame of ESA's JATAC campaign to validate the Aeolus satellite, the multiwavelength-Raman-polarization lidar Polly^{XT} was installed at Mindelo, Cabo Verde, in June 2021 together with further instruments, e.g., an AERONET sun photometer. During the [intensive observation period](#) of the campaign [in September 2021](#), a volcanic eruption at the Cumbre Vieja ridge at [La PalmaLas Palmas](#), Canary Islands, took place, starting on 19 September 2021. Volcanic activity was recorded until 13
410 December 2021. Due to the location of Mindelo in the trade wind zone, the preferred wind direction is northeast, i.e., from [the](#) Canary Islands. Thus, advected air masses contaminated with volcanic aerosol were observed within the local PBL while the SAL above seemed little affected. The occurrence of volcanic aerosol at Mindelo was indicated by an increase of the columnar Ångström exponent and the AOD as measured by the sun photometer after 22 September 2021. Volcanic aerosol was furthermore observed with the Polly^{XT} lidar, [which is shown in a case study of 24 September 2021. on multiple days during the](#)
415 [whole time of volcanic activity. Exemplarily, a case study was shown for the 24 September 2021.](#) On that day, a pronounced pollution was seen over Cabo Verde, strongly contrasting the conditions observed before the start of the eruption. The intense pollution caused an unusually high AOD of [around more than](#) 1.0 at [wavelengths \$\leq 500\$ nm \(AERONET daily mean\), different spectral bands.](#) For a more detailed view, the vertically-resolved optical properties derived from the lidar were analyzed. They were compared to the lidar measurements from 16 September, which was before the start of the eruption and represents the
420 typical aerosol conditions over Cabo Verde at this time of the year.

The lidar measurements for both days showed the presence of two distinct aerosol layers – the [PBLplanetary boundary layer](#) and a lofted layer of Saharan dust. For the 24 September, HYSPLIT trajectory calculations and [TRACEFLEXPART](#) simulations indicated a distinct advection of air masses from [La PalmaLas Palmas](#) in the PBL. Air masses of the lofted layer originated from the Saharan desert. With the lidar, a strong pollution in the PBL was revealed. It led to an unusual high particle extinction
425 coefficient of [721 \$\pm\$ 51, 549 \$\pm\$ 38 and 178 \$\pm\$ 13 Mm⁻¹718, 547 and 126 Mm⁻¹](#) and an enhanced lidar ratio of [66.9 \$\pm\$ 10.1, 60.2 \$\pm\$ 9.2 and 30.8 \$\pm\$ 8.7 sr66.7, 60.1 and 33.7 sr](#) (mean values at 355, 532 and 1064 nm) in contrast to ≤ 130 Mm⁻¹ and ≤ 23.8 sr ≤ 23.7 sr within the clean marine PBL on the 16 September. Thus, on 24 September, the attenuation in the [PBLplanetary boundary layer](#) was increased by a factor of 3–4 compared to the background conditions. The visibility significantly decreased during these days down to 6 km. According to the measured particle extinction coefficient, the AOD for the [PBLplanetary boundary](#)
430 [layer](#) was [0.58 \$\pm\$ 0.03, 0.43 \$\pm\$ 0.02 and 0.18 \$\pm\$ 0.01 at 355, 532 and 1064 nm, respectively.0.69 at 355 nm and 0.5 at 532 nm.](#) It accounts for [54-60%58-64%](#) of the total AOD [in the case of 355 and 532 nm and for 32% at 1064 nm.](#) Compared to the AOD in the PBL of 0.08 (355 and 532 nm) during the clean marine PBL on 16 September 2021, we can conclude that the pollution on 24 September accounted for [81-86%84-88%](#) of the AOD in the PBL [on that day](#) (AOD caused by pollution: [0.5 \$\pm\$ 0.04 0.64 and 0.35 \$\pm\$ 0.030.42](#) at 355 and 532 nm, respectively), i.e., only [14-19%12-16%](#) of the AOD in the PBL were caused
435 by marine aerosol. Since the particle linear depolarization ratio in the PBL was close to 0%, the presence of volcanic ash could be excluded. Instead, sulfate aerosol due to the volcanic eruption at [La PalmaLas Palmas](#) seemed to be the dominating particle type in the low altitudes. This finding was furthermore supported by satellite measurements of Sentinel-5P, showing an advection of SO₂ towards Mindelo on 23 September, which was transformed to sulfate aerosol reaching Mindelo the day after.

In contrast, no indication for pure volcanic aerosol in the lofted layer could be found. The lidar ratio of 64.8 ± 10.2 , 50.9 ± 8.3 and 61.8 ± 8.6 sr (mean values at 355, 532 and 1064 nm) were slightly higher compared to 58.4 ± 8.8 and 47.3 ± 7.2 sr (355, 532 nm) on 16 September 2021. Instead, the particle linear depolarization ratio of 20.8 ± 2.0 , 25.0 ± 1.0 and 20.6 ± 1.0 % were lower than 24.5 ± 2.0 , 28.1 ± 1.0 and 24.1 ± 1.0 % (355, 532, 1064 nm), which were observed on 16 September. However, Saharan dust as the major contributor can still be identified within this layer (SAL) but probably slightly contaminated with smoke, pollution and/or sulfate.

While observations of Saharan dust have already been captured during several campaigns (e.g., SAMUM; Ansmann et al., 2011a; Tesche et al., 2011), it was the first time that the optical properties of volcanic aerosol were observed at Cabo Verde with a multiwavelength-Raman-polarization lidar. Lidar observations of volcanic ash exist for distinct eruptions like Eyjafjallajökull (Ansmann et al., 2010; Groß et al., 2012) but lidar measurements of tropospheric volcanic sulfate aerosol are very rare, yet. Thus, it is important to enlarge the knowledge about the aerosol optical properties of volcanic sulfate, which is aimed by our study. As the observations were made in a usually clean marine PBL, the influence of mixing with other aerosol types is low. Besides this point, we show in our study that far-range transported volcanic aerosol can also effect air quality, indicated by sulfate mass concentrations of more than $100 \mu\text{g m}^{-3}$, and, thus, may have an impact on human health, even more than 1000 km away from the emission source. One additional benefit of this study is the first ever availability of measurements of the particle extinction coefficient and the lidar ratio at 1064 nm for volcanic sulfate aerosol. Having measurements at all three wavelengths is a major advantage with regard to lidar-based aerosol typing and enlarges our data sets. The findings of this study allows us to get new insights in lidar-based aerosol typing and to enlarge our data sets. The findings of this study provide useful insights on the lidar-derived optical properties of volcanic aerosol. They can in turn be used to further improve the aerosol typing by multiwavelength-Raman-polarization lidars, as well as space-borne lidar missions as NASA's CALIPSO CALIOP or ESA's Aeolus and EarthCARE, or assist in the development of new aerosol typing schemes. Besides this point, our findings will be helpful for studying the radiative effects of tropospheric volcanic aerosol, which are still not properly quantified and modeled. As the focus of the campaign at Cabo Verde was on the Aeolus validation activities, there is also the possibility for further research on the potential of Aeolus to capture the volcanic plume on its way to Cabo Verde, which is planned for future studies. Furthermore, a long-term study of the influence of the eruption of Cumbre Vieja on the atmosphere over Cabo Verde based on the ground-based lidar measurements at Mindelo is foreseen but first needs a more robust cloud screening in the automatically derived lidar products. Finally, the observation of this event highlights the necessity for ground-based lidar stations in remote areas. With respect to that, a permanent aerosol and cloud remote sensing station within the framework of ACTRIS has been set up in Mindelo.

Data availability. The Polly^{XT} lidar data will be made available via ACTRIS services, but for now it is available at <https://doi.org/10.5281/zenodo.10650879> (Gebauer et al., 2024). Near-real-time measurement quicklooks can be found at <https://polly.tropos.de/>. AERONET data (station name "Mindelo_OSCM") was downloaded from https://aeronet.gsfc.nasa.gov/cgi-bin/draw_map_display_aod_v3?long1=-180&long2=180&lat1=-89&lat2=90&multiplier=2&what_map=4&nachal=1&formatter=0&level=2&place_code=10 https://aeronet.gsfc.nasa.gov/cgi-bin/data_displa

(AERONET, 2024), last access: ~~14 February 2024~~~~9 March 2023~~. HYSPLIT trajectories were calculated using the online tool on <https://www.ready.noaa.gov/hypub-bin/trajtype.pl?runtype=archive> with meteorological input data from GDAS1 (<https://www.ready.noaa.gov/gdas1.php>), last access: ~~23 January 2024~~~~20 January 2023~~. The data for the FLEXPART analysis was taken from <https://doi.org/10.5065/D6M043C6>, last access: ~~29 January 2024~~~~24 September 2021~~. The TROPOMI SO₂ plot [wasis](https://so2.gsfc.nasa.gov/pix/daily/ixxxza/troploop5pca.php?yr=21&mo=09&dy=23&bn=cverde) taken from <https://so2.gsfc.nasa.gov/pix/daily/ixxxza/troploop5pca.php?yr=21&mo=09&dy=23&bn=cverde>. The underlying data was downloaded from the S5P-PAL data portal (<https://data-portal.s5p-pal.com/>) from July 2022 onward, and from BIRA-IASB distributions website (<https://distributions.aeronomie.be/>) for older data, last access: ~~6 February 2024~~~~13 March 2023~~.

Table 1. Layer-mean values of the lidar-derived aerosol optical and microphysical properties for the PBL and the lofted layer on 24 September 2021, 04:38–05:57 UTC, and on 16 September 2021, 22:24–23:12 UTC. The values are given along with the standard deviation (parameter variability within the layer) for the extensive aerosol properties and the sulfate mass concentration and with the layer-mean errors (errors as described in Sec. 2.1) for the intensive aerosol properties. Geometric information of the aerosol layers is also provided (note that for the calculation of the extinction-related properties in the PBL the layer bottom height is 0.25 km due to the overlap configurations). Layer mean values and standard deviation of the lidar-derived aerosol optical properties for the planetary boundary layer and the lofted layer on 24 September averaged from 04:38 to 05:57 UTC and on 16 September 2021 from 22:24 to 23:12 UTC. To avoid edge and overlap effects, the mean values were calculated for the following height ranges: lofted layer (ff measurements): 1.3–5.3 km (16 September with vertical smoothing as in Fig. 3) and 1.4–4.4 km (24 September with vertical smoothing as in Fig. 4); PBL (nf measurements): 0.06–0.6 km for backscatter-related properties and 0.25–0.6 km for extinction-related properties (16 September) and 0.06–1 km and 0.25–1 km, respectively (24 September), both with vertical smoothing as in Fig. 5.

Layer mean optical and microphysical properties					
		PBL (nf, smoothing as in Fig. 5)		lofted layer (ff, smoothing as in Figs. 3 and 4)	
		24 Sep	16 Sep	24 Sep	16 Sep
		0.06–0.8 km	0.06–0.6 km	1.3–5.3 km	1.4–4.4 km
		sulfate	marine	Saharan dust	
Extensive aerosol optical properties and microphysical properties					
Particle backscatter coefficient	355 nm	10.9 ± 0.6	6.8 ± 0.5	$2.01_9 \pm 0.9$	2.4 ± 0.5
	532 nm	9.2 ± 0.58	$5.65_7 \pm 0.4$	$2.32_4 \pm 1.1$	2.9 ± 0.6
	1064 nm	4.1 ± 0.44	3.6 ± 0.3	1.8 ± 0.8	1.8 ± 0.3
	1064 nm (RR)	6.1 ± 0.34	-	2.2 ± 0.92	-
Particle extinction coefficient	355 nm	721 ± 51	114 ± 20	120 ± 64	138 ± 31
	532 nm	549 ± 38	130 ± 24	$114_{113} \pm 65$	134 ± 32
	1064 nm	178 ± 13	-	168 ± 21	-
Aerosol optical depth	355 nm	0.58 ± 0.03	0.08 ± 0.01	0.38 ± 0.25	0.57 ± 0.21
	532 nm	0.43 ± 0.02	0.08 ± 0.01	0.36 ± 0.24	0.57 ± 0.20
	1064 nm	0.18 ± 0.01	-	0.39 ± 0.16	-
sulfate mass concentration ($\mu\text{g m}^{-3}$)		133.1 ± 20.3	-	-	-
Intensive aerosol optical properties					
Lidar ratio (sr)	355 nm	66.9 ± 10.1	17.3 ± 2.8	64.8 ± 10.2	58.4 ± 8.8
	532 nm	60.2 ± 9.2	23.8 ± 4.2	50.9 ± 8.3	47.3 ± 7.2
	1064 nm	30.8 ± 8.7	-	61.8 ± 8.6	-
Particle linear depolarization ratio (%)	355 nm	0.3 ± 2.0	0.7 ± 2.0	20.8 ± 2.0	24.5 ± 2.0
	532 nm	0.7 ± 1.0	1.1 ± 1.0	25.0 ± 1.0	28.1 ± 1.0
	1064 nm	0.9 ± 1.0	1.0 ± 1.0	20.6 ± 1.0	24.1 ± 1.0
Ångström exponent	ext 355/532 nm	0.68 ± 0.07	-0.32 ± 0.29	0.10 ± 0.14	0.06 ± 0.08
	ext 532/1064 nm	1.53 ± 0.26	-	-0.06 ± 0.53	-

Appendix A: AOD and lidar quicklooksprofiles

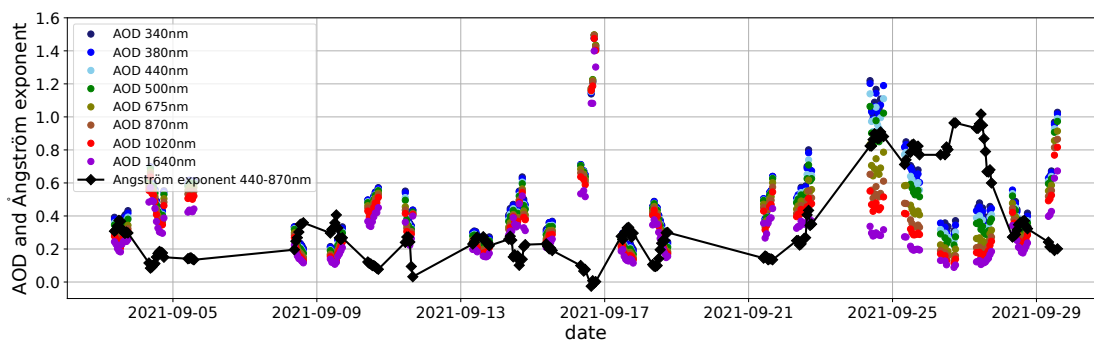


Figure A1. Same as Figure 1 but including the measurement of 16 September 2021, which was contaminated by a cirrus cloud.

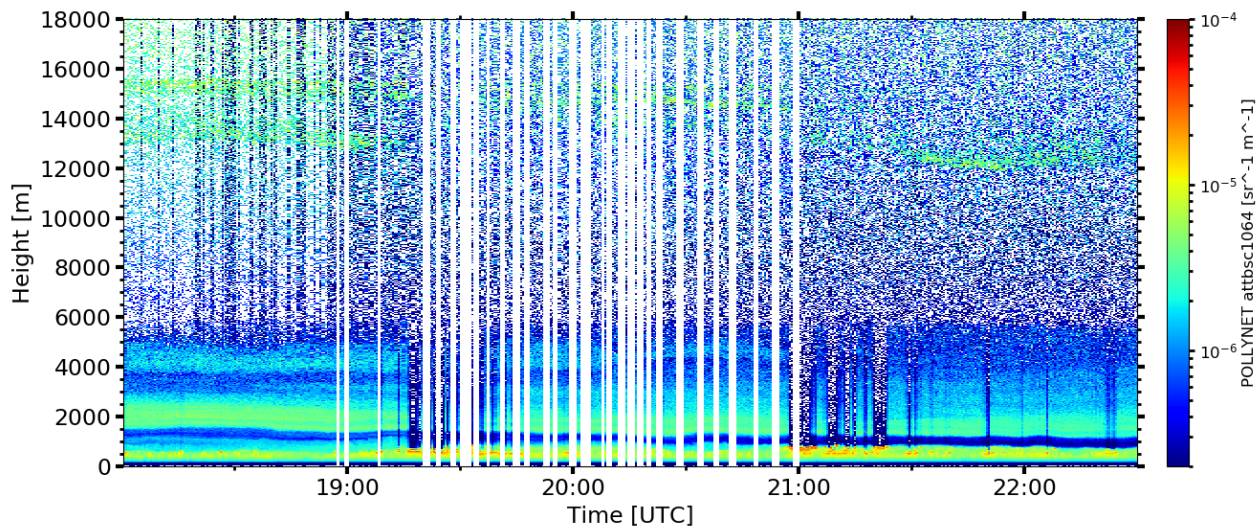


Figure A2. Temporal evolution of the calibrated attenuated backscatter coefficient at 1064 nm measured by Polly^{XT} at Mindelo, Cabo Verde, during 16 September 2021, 18–22:30 UTC. In an altitude between 12 and 16 km the Cirrus cloud was located, which was not correctly screened out by the AERONET algorithm.

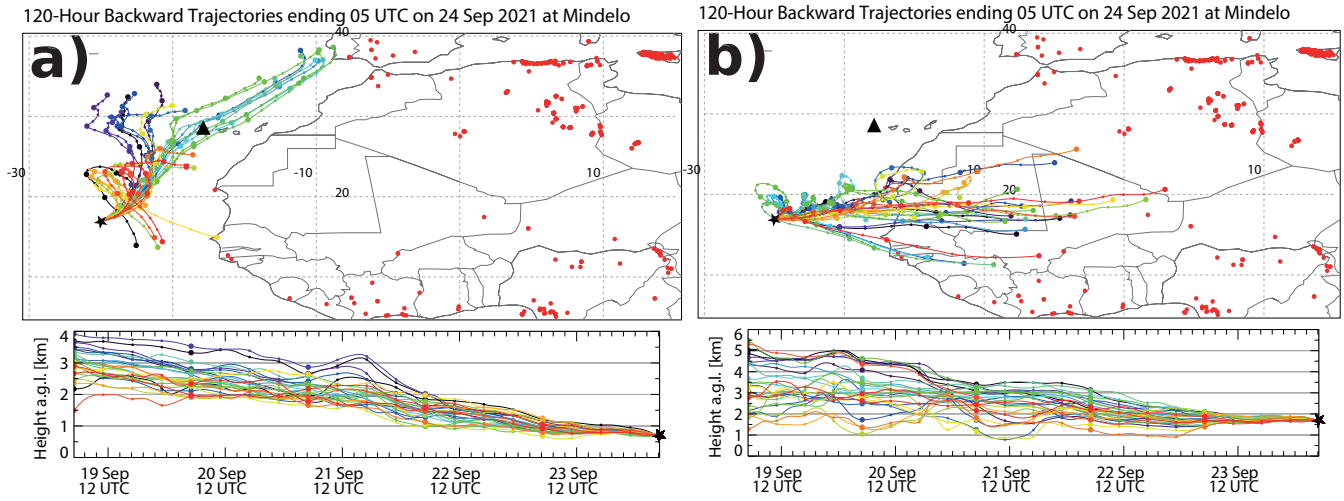


Figure B1. Same as Fig. 6 but with fire spot analysis from MODIS (FIRMS, firms.modaps.eosdis.nasa.gov) between 14 and 24 September 2021.

Appendix C: TRACEFLEXPART simulations

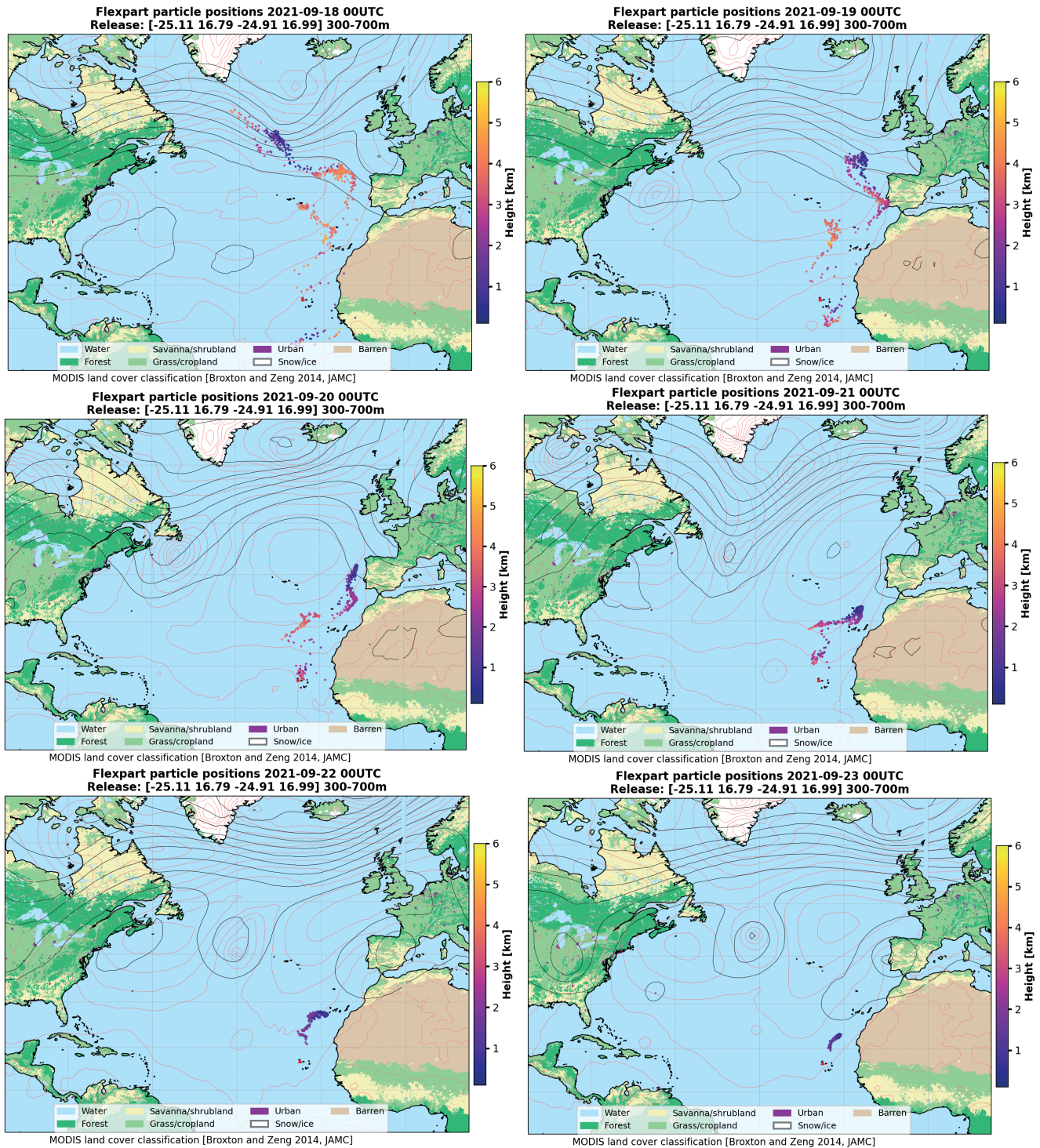


Figure C1. Selected **TRACEFlexpart** simulations show the location of single air parcels (labelled as "particle positions") between 18 and 23 September 2021, each at 0 UTC, before they arrive at Mindelo (red triangle) at 500m on 24 September 2021, 6 UTC. The colour of the dots indicates their height above ground. MODIS land cover classification according to Broxton et al. (2014).

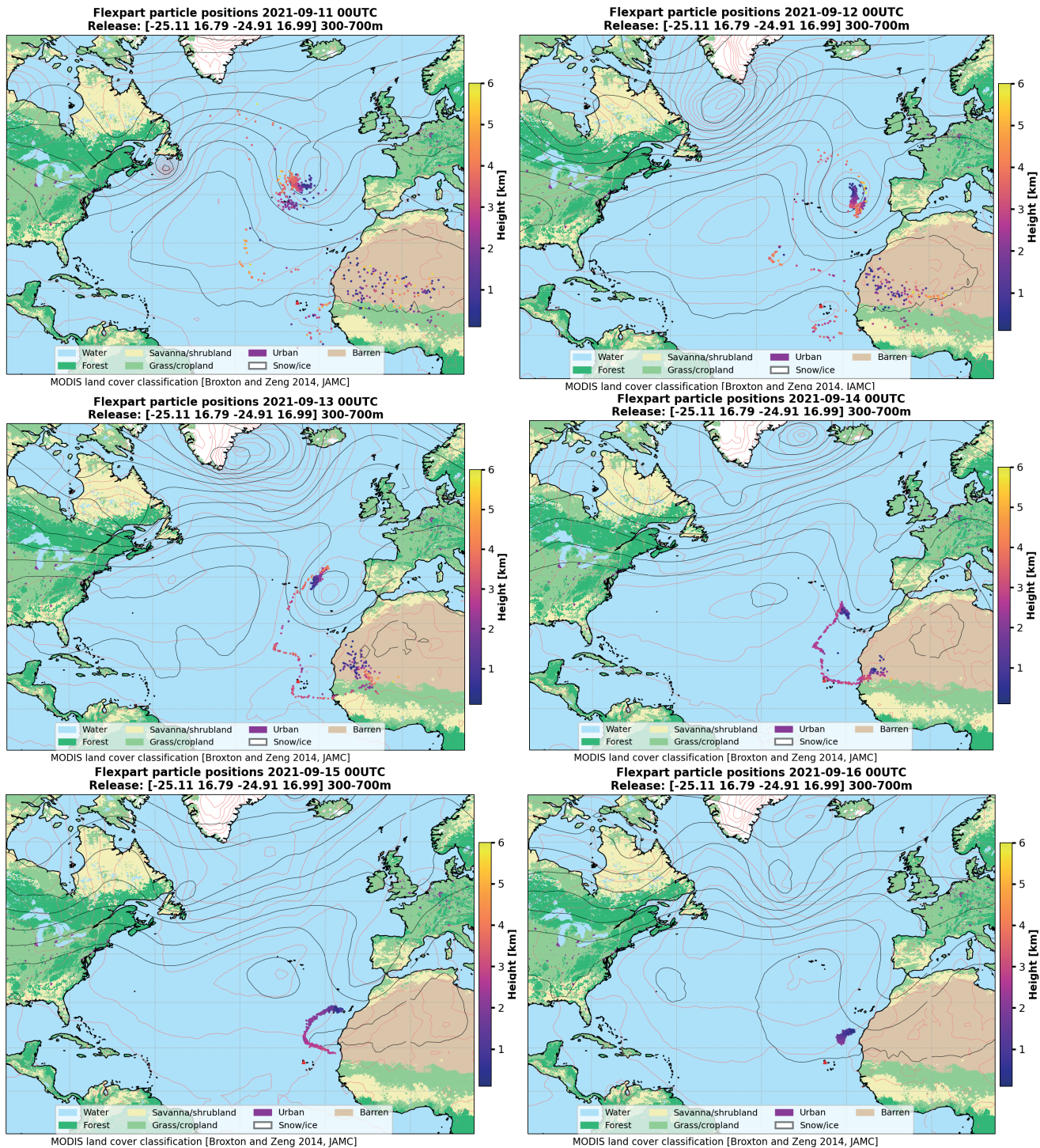


Figure C2. Same as Fig. C1 but for air parcels arriving on 16 September 2021, 0 UTC.

Author contributions. HG conceptualized the manuscript together with HB and AAF. MH provided the data of the lidar extinction measurements at 1064 nm. MR performed the [TRACEFLEXPART](#) simulations. AA contributed his expertise on lidar data analysis and volcanic aerosol. RE, DA, HB, and AS have been responsible for the deployment and operation of the ground-based instruments at Mindelo. CZ coordinates the scientific activities at OSCM, Cabo Verde. All coauthors were actively involved in the extended discussions and the elaboration of the final design of the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

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