Cotc 11:4-	A Train Calastics Criteria	E2GM COGD-2.0	Diagnostia spellestions
Satellite	A-Train Selection Criteria	E3SM-COSPv2.0	Diagnostic applications
Composite	December of a CLWIC 1 and	Selection Criteria	GI WIG 1 12 1
MODIS and CloudSat	Based on the SLWC detection scheme described in Suzuki et al. (2010), with updated Cloud Optical Thickness (COT) threshold for consistency with COSPv2.0 WRDs: • CloudSat reflectivity profiles (2B-GEOPROF R05) are matched to MODIS cloud profiles (MOD06-1KM-AUX R05). • Cloud tops and bottom are determined where reflectivity > -30 dBZ. • Single layer clouds are selected where the MODIS cloud layer flag ('Cloud_Multi_Layer_Flag') indicates one layer and COT > 0.3. • MODIS cloud top pressure > 500 hPa. • MODIS cloud top effective radius 5 ≤ Re ≤ 30 μm • To select warm liquid clouds, the ECMWF-AUX temperature profiles were matched to the Cloud Profiling Radar (CPR) footprint. • Profiles are selected where the ECMWF-AUX cloud top temperature and MODIS cloud top temperature ≥ 273 K. • Profiles selected where CPR cloud mask ('cpr_cmask') values are ≥ 30, indicating a good or strong echo with high-	Based on the WRDs originally implemented in COSPv2.0 (Michibata et al., 2019), with modifications described in main text Sect. 2.2. Subcolumns selected where: • MODIS liquid water path (LWP) > 0 g/kg • MODIS liquid COT > 0.3 • MODIS lice Water Path (IWP) ≤ 0 g/kg • MODIS lice COT < 0.3 • MODIS liquid cloud top effective radius 5 ≤ Re ≤ 30 μm • CloudSat reflectivity ≥ -30 dBZ for one or more contiguous layers • Temperature at cloud top (determined by CloudSat reflectivity threshold described above) ≥ 273 K	 SLWC cloud fraction maps, binned by CloudSat reflectivity CFODDs binned by MODIS cloud top R_e MODIS COT PDFs binned by MODIS cloud top R_e

Table S2. PI base cloud state for the 12 sensitivity experiments. Dash ("-") indicates the KK2000 coefficient value was unchanged from the default E3SMv2 parameterization (equal to the "CNTL" simulation value).

Name	A	α	β	accre	PI LWP (kg m ⁻²)	PI SLWC Cloud Fraction	PI SWCRE (W m ⁻²)
CNTL	3.05E+04	3.19	-1.4	1.75	0.107	0.052	-12.1
alpha01	-	4.22	-	-	0.180	0.049	-14.1
alpha02	-	3	-	-	0.080	0.052	-10.7
beta01	-	-	-1		0.087	0.050	-10.4
beta02	-	-	-1.79	-	0.124	0.052	-13.0
beta03	-	-	-3.01	-	0.161	0.051	-14.1
acoef0.05x	1.35E+03	=	-	-	0.150	0.052	-13.9
acoef5x	1.53E+05	=	-	-	0.079	0.050	-10.1
acoef10x	3.05E+05	-	-	-	0.066	0.047	-8.9
acoef50x	1.53E+06	-	-	-	0.039	0.034	-5.2
acoef100x	3.05E+06	-	-	-	0.030	0.026	-3.6
accre	=	-	-	5	0.077	0.049	-10.2

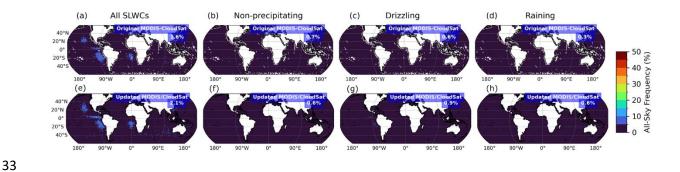


Figure S1. All-sky frequencies of total SLWCs June 2006 – Apr 2011, non-precipitating ($Z_{max} < -15 \ dBZ_e$), drizzling ($-15 \ dBZ_e \le Z_{max} < 0 \ dBZ_e$) and raining ($Z_{max} \ge 0 \ dBZ_e$) ocean-only SLWCs according to original reference analysis of MODIS and CloudSat observations (Michibata et al., 2019a, 2019b) (a-d), and updated reference MODIS and CloudSat analysis (as in Fig. 1), but increasing the lower MODIS COT threshold from 0.3 to 15.

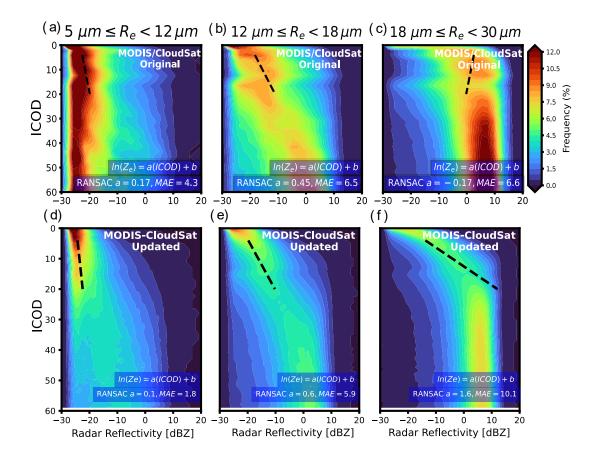


Figure S2. Contoured frequency by optical depth diagrams (CFODDs) for SLWCs June 2006 – April 2011 binned by MODIS cloud top effective radius (R_e) from original reference MODIS-CloudSat observations analysis (a-c) and updated reference MODIS-CloudSat observations analysis (d-f) as in Fig. 2, but increasing the lower MODIS COT threshold from 0.3 to 15. Random Sample Consensus (RANSAC) linear regressions were applied to the CFODD at $4 \le ICOD \le 20$ to estimate droplet collection efficiencies. RANSAC slopes and Median Absolute Error (MAE) values are shown in blue boxes.

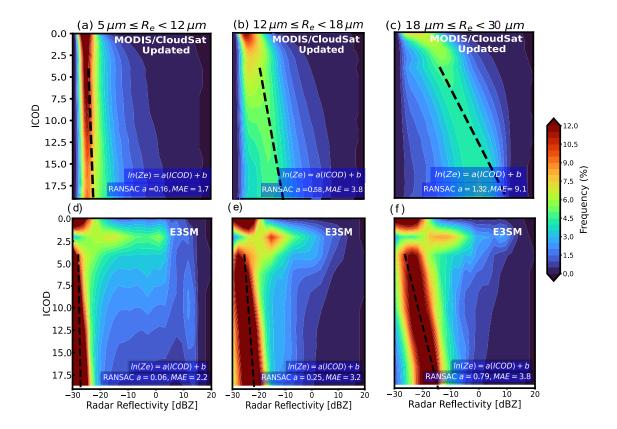


Figure S3. Contour frequency by optical depth diagrams (CFODDs) for subset of SLWCs with max CloudSat reflectivity < 20 dBZ and COT < 20, June 2006 – April 2011 binned by MODIS cloud top effective radius (Re) from updated reference MODIS-CloudSat observations analysis (a-c) and the E3SMv2 simulation (d-f). CloudSat reflectivities are binned by MODIS in-cloud optical depth (ICOD) to construct CFODDs. Random Sample Consensus (RANSAC) linear regressions were applied to the CFODD at $4 \le ICOD \le 20$ to estimate droplet collection efficiencies. RANSAC slopes and Median Absolute Error (MAE) values are shown in blue boxes. E3SM-COSP CFODDs shows discontinuity in CloudSat reflectivity frequencies near cloud top, and decreased droplet collection efficiencies compared to observations.

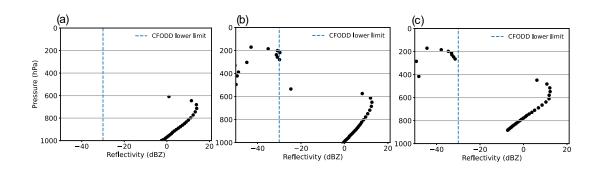


Figure S4. Example E3SMv2 SLWC reflectivity profiles from the CloudSat simulator output in COSPv2.0. E3SMv2 SLWCs exhibit reflectivity > 0 dBZ at cloud top with high frequency compared to MODIS-CloudSat observations (see Fig. 2, Sect. 3). A CloudSat ground-clutter mask that was implemented in the WRDs for improved comparison with observations is not shown here.

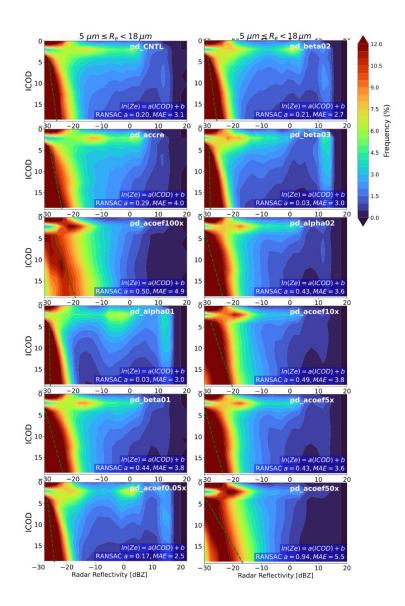


Figure S5. CFODDs for E3SMv2 PD simulations in 12 experiments featuring variations of the default E3SMv2 autoconversion and accretion parameterizations (Table 1), for SLWCs with MODIS R_e between 5 and 18 μm and COT between 4 and 20. RANSAC linear regressions were applied to the CFODDs at $4 \le ICOD \le 20$. RANSAC slopes and MAE values are shown in blue boxes.

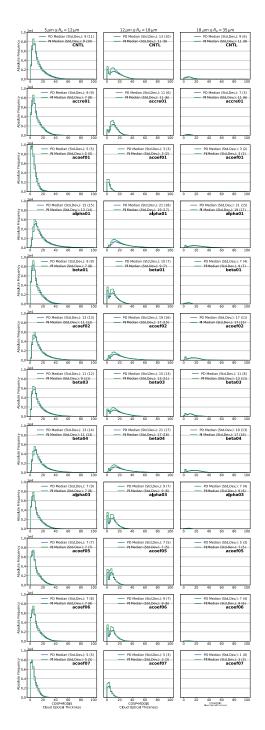


Figure S6. Absolute frequency of SLWCs in E3SMv2 in 12 warm rain process sensitivity experiments, binned by simulated MODIS R_e . Blue and green PDFs indicate the PD and PI simulation results, respectively.

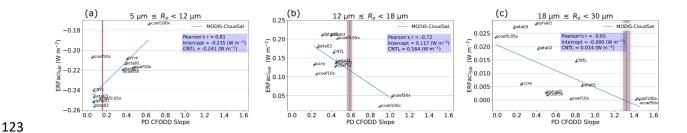


Figure S7. Linear regression between E3SMv2 ERFacisw and CFODD slopes in 12 PD autoconversion and accretion sensitivity experiments, binned by MODIS R_e . Results show that SLWCs in the small and medium R_e size bin contribute to ERFacisw in equal magnitude but opposite sign, and SLWCs with large R_e make a relatively small positive contribution to ERFacisw compared to the small or medium R_e populations. The positive correlation in the small R_e size bin indicates that increasing droplet collection efficiency weakens ERFacisw for this SLWC subset. The positive ERFacisw values that diminish with increasing CFODD slope in the medium and large R_e size bins indicate that increased aerosol yields decreased small and medium Re SLWC cloud fraction (see Figs. S12-S13), but that increased droplet collection efficiencies oppose the aerosol effect. Grey and pink shaded regions indicate the 68 and 98% confidence intervals for the MODIS-CloudSat CFODD slope, respectively. Labels indicate the sensitivity experiment names (Table 1).

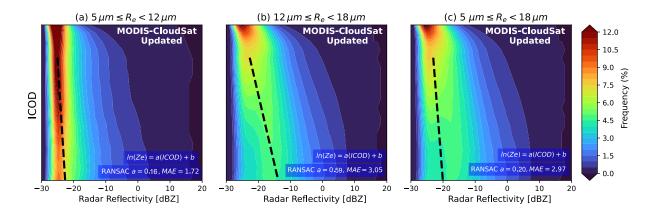


Figure S8. CFODDs as in Fig. S3 for subset of SLWCs with max CloudSat reflectivity < 20 dBZ and COT < 20, June 2006 – April 2011 binned by MODIS R_e from updated reference MODIS-CloudSat observations analysis (a-c), but with combined "small" and "medium" Re SLWCs in (c). CloudSat reflectivities are binned by MODIS ICOD to construct CFODDs. Random Sample Consensus (RANSAC) linear regressions were applied to the CFODD at $4 \le ICOD \le 20$ to estimate droplet collection efficiencies. RANSAC slopes and Median Absolute Error (MAE) values are shown in blue boxes.

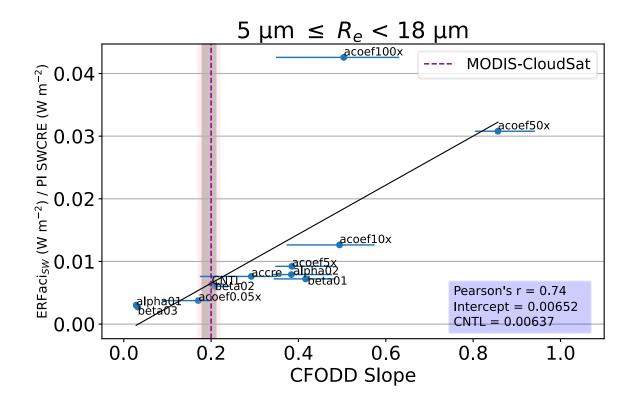


Figure S9. Linear regression between E3SMv2 ERFacisw normalized by SWCRE and CFODD slopes in 12 PD autoconversion and accretion sensitivity experiments, generated from SLWCs with MODIS R_e between 5 and 18 μ m. Error bars represent 1-sigma error estimated from RANSAC-fit bootstrapping (Sect. 2). Grey and pink shaded regions indicate the 68 and 98% confidence intervals for the MODIS-CloudSat CFODD slope, respectively. Labels indicate the sensitivity experiment names (Table 1).

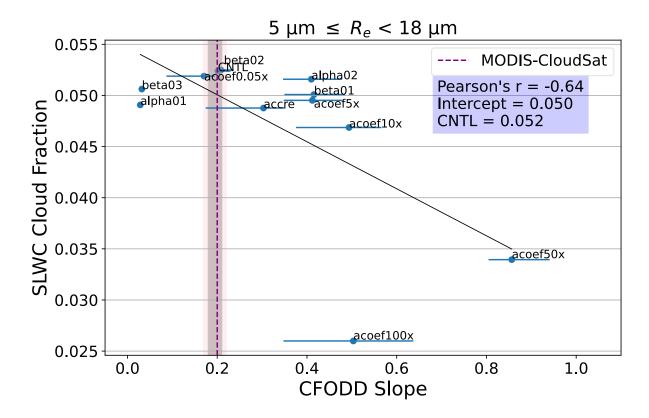


Figure S10. Linear regression between PI E3SMv2 SLWC cloud fraction and PD CFODD slopes in 12 autoconversion and accretion sensitivity experiments, generated from SLWCs with MODIS R_e between 5 and 18 μ m. Error bars represent 1-sigma error estimated from RANSAC-fit bootstrapping (Sect. 2). Grey and pink shaded regions indicate the 68 and 98% confidence intervals for the MODIS-CloudSat CFODD slope, respectively. Labels indicate the sensitivity experiment names (Table 1).

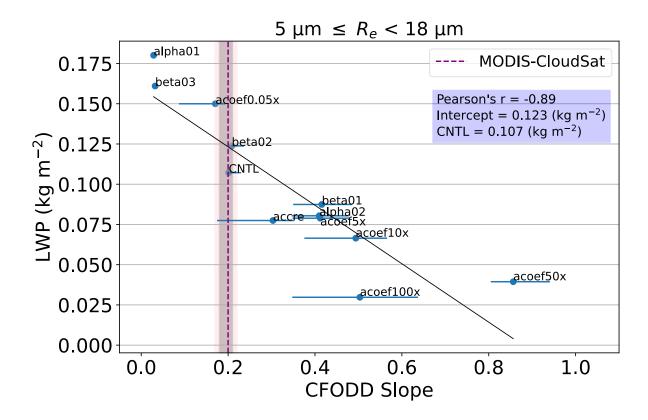


Figure S11. Linear regression between PI E3SMv2 SLWC LWP and PD CFODD slopes in 12 autoconversion and accretion sensitivity experiments, generated from SLWCs with MODIS R_e between 5 and 18 μ m. Error bars represent 1-sigma error estimated from RANSAC-fit bootstrapping (Sect. 2). Grey and pink shaded regions indicate the 68 and 98% confidence intervals for the MODIS-CloudSat CFODD slope, respectively. Labels indicate the sensitivity experiment names (Table 1).

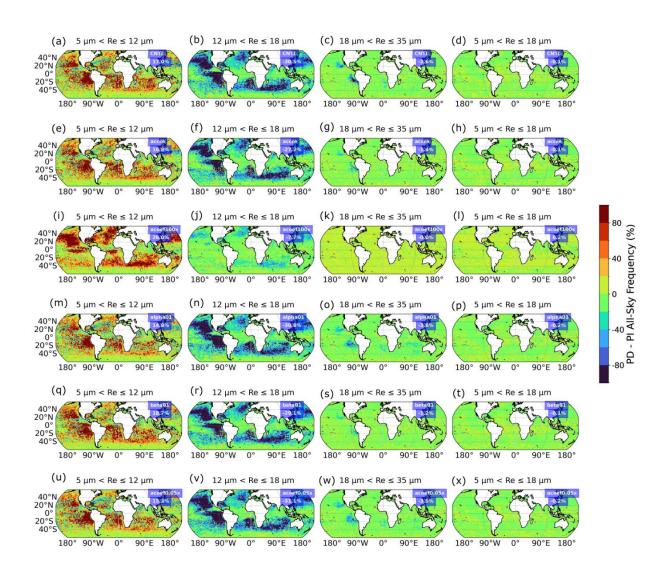


Figure S12. Difference between PD and PI all-sky SLWC cloud fraction in 6 of 12 warm rain process sensitivity experiments, binned by simulated MODIS R_e. Labels indicate experiment name (Table 1) and global mean cloud fraction difference.

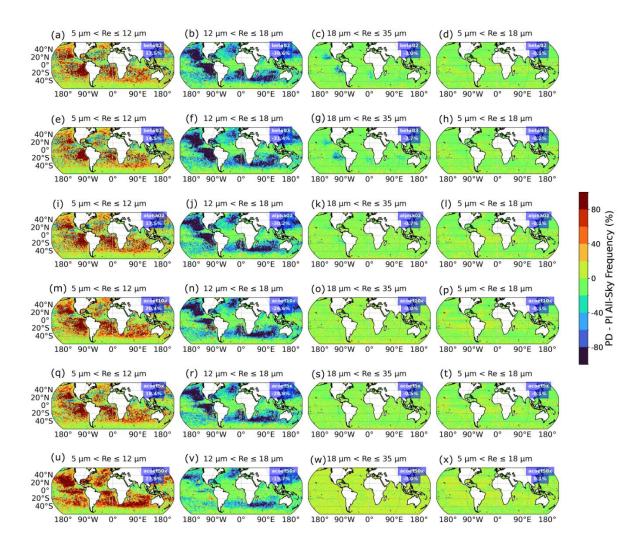


Figure S13. Difference between PD and PI all-sky SLWC cloud fraction in 6 of 12 warm rain process sensitivity experiments, binned by simulated MODIS R_e. Labels indicate experiment name (Table 1) and global mean cloud fraction difference.

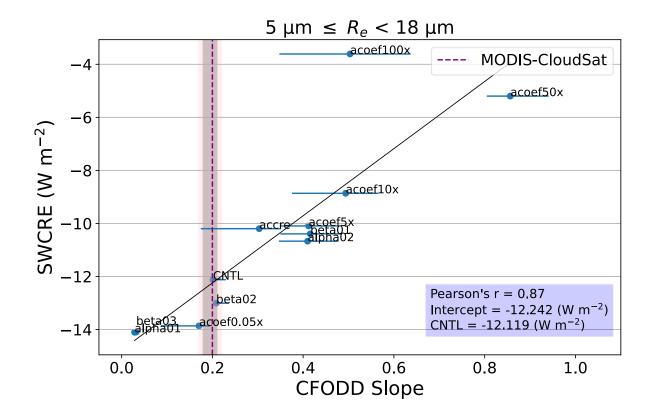


Figure S14. Linear regression between PI E3SMv2 SLWC SWCRE and PD CFODD slopes in 12 autoconversion and accretion sensitivity experiments, generated from SLWCs with MODIS R_e between 5 and 18 μ m. Error bars represent 1-sigma error estimated from RANSAC-fit bootstrapping (Sect. 2). Grey and pink shaded regions indicate the 68 and 98% confidence intervals for the MODIS-CloudSat CFODD slope, respectively. Labels indicate the sensitivity experiment names (Table 1).

235 References 236 Mace, G. G., & Zhang, Q. (2014). The CloudSat radar-lidar geometrical profile product (RL-GeoProf): Updates, 237 improvements, and selected results. Journal of Geophysical Research: Atmospheres, 119(15), 9441–9462. 238 https://doi.org/https://doi.org/10.1002/2013JD021374 239 Marchand, R., Mace, G. G., Ackerman, T., & Stephens, G. (2008). Hydrometeor Detection Using Cloudsat—An 240 Earth-Orbiting 94-GHz Cloud Radar. Journal of Atmospheric and Oceanic Technology, 25(4), 519-533. 241 https://doi.org/10.1175/2007JTECHA1006.1 242 Michibata, T., Suzuki, K., Ogura, T., & Jing, X. (2019a). Incorporation of inline warm rain diagnostics into the 243 COSP2 satellite simulator for process-oriented model evaluation. Geoscientific Model Development, 12(10), 4297-4307. https://doi.org/10.5194/gmd-12-4297-2019 244 245 Michibata, T., Suzuki, K., Ogura, T., & Jing, X. (2019b). Incorporation of inline warm rain diagnostics into the 246 COSP2 satellite simulator for process-oriented model evaluation. Geoscientific Model Development, 12(10), 247 4297–4307. https://doi.org/10.5194/gmd-12-4297-2019 248 Platnick, S., Meyer, K. G., King, M. D., Wind, G., Amarasinghe, N., Marchant, B., Arnold, G. T., Zhang, Z., 249 Hubanks, P. A., Holz, R. E., Yang, P., Ridgway, W. L., & Riedi, J. (2017). The MODIS Cloud Optical and 250 Microphysical Products: Collection 6 Updates and Examples From Terra and Aqua. IEEE Transactions on 251 Geoscience and Remote Sensing, 55(1), 502-525. https://doi.org/10.1109/TGRS.2016.2610522 252 Suzuki, K., Nakajima, T. Y., & Stephens, G. L. (2010). Particle Growth and Drop Collection Efficiency of Warm 253 Clouds as Inferred from Joint CloudSat and MODIS Observations. Journal of the Atmospheric Sciences, 254 67(9), 3019–3032. https://doi.org/10.1175/2010JAS3463.1 255