



1 **Increasing seasonal variation in the extent of rivers and lakes** 2 **from 1984 to 2022**

3 Björn Nyberg^{1,2,3*}, Roger Sayre⁴, Elco Luijendijk¹

4 Department of Earth Sciences, University of Bergen, Allegaten 41, 5020, Bergen, Norway.¹

5 Bjerknes Centre for Climate Research, Allegaten 70, 5020, Bergen, Norway.²

6 7Analytics, Innovation District Solheimsviken 7c, 5054, Bergen, Norway³

7 U.S. Geological Survey, 516 National Center, Reston, VA, 20192, USA.⁴

8

9

10 **Abstract**

11

12 Knowledge of the spatial and temporal distribution of surface water is important for water resource management,
13 flood risk assessment, monitoring ecosystem health, constraining estimates of biogeochemical cycles and
14 understanding our climate. While global scale spatial-temporal change detection of surface water has significantly
15 improved in recent years with planetary scale remote sensing and computing, it has remained challenging to
16 distinguish the changing characteristics of rivers and lakes. Here we analyze the spatial extent of permanent and
17 seasonal rivers and lakes globally over the past 38-years based on new data of river system extents and surface
18 water trends. Results show that while the total permanent surface area of both rivers and lakes has remained
19 relatively constant, the area with intermittent seasonal coverage has increased by 12% and 27% for rivers and
20 lakes, respectively. The increase is statistically significant in over 84% of global water catchments based on
21 Spearman rank correlations above 0.05 and p values less than 0.05. The results of our analysis are shared as the
22 Surface Area of Rivers and Lakes (SARL) database, which contributes to improved understanding of the
23 hydrological cycle and management of water resources.

24

25

26 **1. Introduction**

27

28 Climate change and population growth have placed considerable stress on our natural freshwater resources. Water
29 demand has increased nearly 8-fold over the past century with an estimated 70% of the total used to meet irrigation
30 needs (Siebert et al., 2010; Wada et al., 2016). To meet the increasing water demand, an estimated 16.7 million
31 reservoirs have been built holding a predicted 8070 km³ of freshwater (Lehner et al., 2016). Current water demand
32 represents only 10% of our approximate annual renewable freshwater resources (Oki and Kanae, 2006).
33 Nonetheless, water scarcity remains a significant problem around the world due to the variability of water in time
34 and space (Oki and Kanae, 2006; Mekonnen et al., 2016). The hydrological cycle is also crucial to the health of
35 our ecosystems and biodiversity that depend on the recurrence and seasonality of water to support life (Gleeson et
36 al., 2020). Furthermore, inland waters are also an important component in biogeochemical cycles of CO₂ and
37 methane that, by size, disproportionately contribute a significant portion to our total greenhouse emissions annually
38 (Bastviken et al., 2004; Allen and Pavelsky, 2018; Matthews et al., 2020).



39

40 While the surface area of water is only one part of the hydrological cycle, it is the most accessible portion
41 influencing human and ecosystem behavior and an important component in groundwater recharge (Oki and Kanae,
42 2006; Sibert et al., 2010; Gleeson et al., 2020). Knowledge of the type of waterbody, i.e. whether it is permanent,
43 intermittent, or seasonal and whether it is part of a river system or a lake, is important to understand the role of
44 water bodies in different hydrological processes, ecosystem support and biogeochemical cycles. The changing
45 physical environment and its waterbody type due to droughts, floods or direct human alteration also alters
46 migration patterns of humans, ecosystems, and their biodiversity (Neumann et al., 2015; Van Loon et al., 2016).
47 In addition, the perennial and seasonal state of both rivers and lakes has important implications for ecosystem
48 health (Messenger et al., 2021) and carbon cycles (Keller et al., 2020). The type, extent and seasonality of
49 waterbodies at a global basin scale is needed for improved water resource management sustained delivery of
50 ecosystem services (Sheffield et al., 2018).

51

52 Planetary scale computing and analysis of remotely sensed imagery have led to a number of studies revealing the
53 unprecedented impact of human resource management and climate change stress on the extent of water (Van Dijk
54 et al., 2011; Wada et al., 2016; Pekel et al., 2016, Donchyts, et al., 2016). In particular, significant advancements
55 have been made in identifying and quantifying the historical change in global reservoirs at a 0.01 to 100 km²
56 resolution (Donchyts et al., 2022). In addition, the temporal analysis of lakes extents has been analyzed up to 50
57 degrees North at a 0.1 km² resolution (Khandelwal et al., 2022). The identification of natural lakes in high latitude
58 regions has not been analyzed within a global water surface change context.

59

60 Far fewer studies have analyzed the change of river extents. Allen and Pavelsky (2018) mapped the observed
61 global surface area of rivers but only for a specific year and at mean annual water discharge. More recently, Feng
62 et al., (2022) quantified the temporal variability in global river widths over the past 37 years based on 30 m Landsat
63 imagery. However, this study does not map the changing surface area of rivers nor provides measurements at the
64 confluence and divergence of rivers common in anabranching and braided systems that comprise an estimated
65 52% of global rivers (Nyberg et al., 2023). As a result, there remains a significant knowledge gap in the temporal
66 variability of river surface area and its interaction with lakes and reservoirs. The aim of this paper is to compile
67 existing information to map the historical change in water surface area for rivers and lakes over the past 38-years,
68 and to examine the implications for water resource management, ecosystem health and biogeochemical cycles.

69

70

71 **2. Materials and Methods**

72

73 **2.1 Water Surface Area Classification**

74

75 To classify the permanent and seasonal extent of lakes and rivers, we utilize existing datasets describing the surface
76 area of water combined with a new and improved definition of river extents on a global scale. Our analysis is based
77 on the Global Surface Water (GSW) model by Pekel et al. (2016) that describes the permanent and seasonal extent
78 of open water from 1984 to 2023 based on 30 m Landsat imagery. Permanent surface area of water is defined as

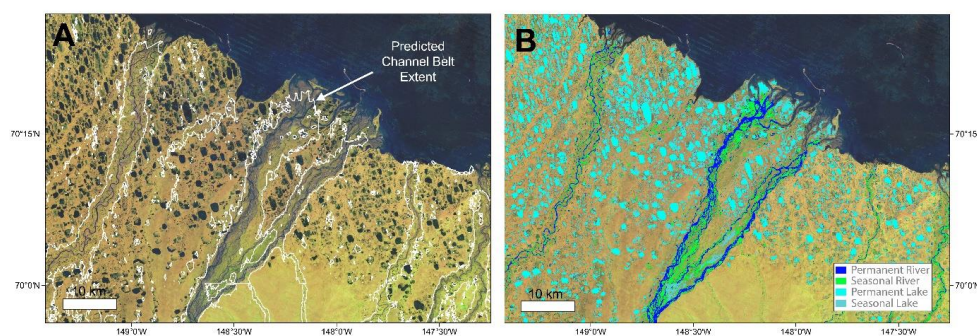


79 locations where open water is detected for all twelve months of a given year (or 100% of valid pixels). Seasonal
80 surface water was defined as any pixel location with at least one month during which water was detected. The
81 authors report less than 1% false positive open water classifications and less than 5% missed open water
82 classifications based on 40,000 randomly selected points.

83

84 To define the spatial extent of rivers is challenging given the dynamic nature of rivers, varying morphology, and
85 perennial versus non-perennial character of rivers. Nyberg et al. (2023) quantified the global extent of river channel
86 belts (GCB) based on 30 m resolution Landsat imagery using a machine learning method that spatially delineates
87 channel belt areas based on geomorphological properties. The river channel belt extents show the riverine
88 landforms of the river channel and its associated levees, bars and overbank deposits that therefore capture the
89 evolution of the riverine environment over time (Figure 1). The model reports a confidence value ranging from 0
90 to 100%, where a 0% confidence value indicates a non-riverine environment whereas a 100% confidence indicates
91 a riverine environment for a given pixel location. An evaluation of this model reported a 94% accuracy to the
92 validation dataset for channel belts wider than a 1 km. By combining the GCB and GSW datasets, we produce a
93 new global dataset mapping the historical change of the seasonal and permanent water surface area of lakes and
94 rivers (SARL) from 1984 to 2023. The seasonal extent of water within the channel belt shows rivers at bankfull or
95 larger flood events with inundation persisting for at least one month. Waterbodies outside the channel belt are
96 defined as a lakes or wetland regions that represent a body of water outside the delineated river channel belt extents
97 (Figure 1).

98



99 **Figure 1: Example permanent and seasonal water extent in rivers and lakes - A) Example Landsat 8 imagery for 2020**
100 **with overlain delineations of the maximum channel belt extent in white based on the GCB model (Nyberg et al., 2023).**
101 **Any pixel outside the channel belt is defined as lacustrine/wetland or floodplain. B) Change in permanent water extent**
102 **from 1984 to 2020 based on the GSW model (Pekel et al., 2016). Combined the two datasets provide a definition of the**
103 **seasonal and permanent change of water in riverine and lacustrine environments. Landsat imagery courtesy of the**
104 **USGS. See interactive map for further detail.**

105

106 To improve delineations of river channel belt extent we refine the classification of Nyberg et al., (2023) by utilizing
107 a 10% confidence on the GCB prediction below 60 degrees North and a 50% confidence above 60 degrees North.
108 This step was chosen to constrain rivers without distinct channel belts in high latitude regions (e.g., Canadian
109 Shield or Siberian Plateau) given most sedimentary basins with clearly defined channel belts occur around mid to



110 low latitude regions (Nyberg and Howell, 2016). In addition, previous databases of large lakes and reservoirs (~ >
111 10 km²) defined by the HydroLakes (Messenger et al., 2016) and OpenStreetMap (2022) datasets were included to
112 further improve delineations. The inclusion of the lacustrine databases reduced the global channel belt extent by
113 1.86% or 13.4 x 10⁴ km². These steps were processed on the Google Earth Engine platform (Gorelick et al., 2017)
114 resulting in a global database of the seasonal and permanent surface area of rivers and lakes from 1984 to 2022 at
115 a 30 m Landsat resolution.

116

117

118 **2.2 Temporal Water Surface Area Analysis**

119

120 Following the mapping of the SARL database, we analyze the data by aggregating results on drainage catchments
121 derived from the HydroSHEDS level 5 catchment dataset (Lehner et al., 2008). Considering that satellite imagery
122 is not available for certain years due to non-acquisition or excessive cloud cover, it is important to consider missing
123 values in the time series of surface water observations. Pekel et al., 2016 define the location of missing values for
124 each year in the GSW database but do not identify the waterbody type or seasonality of those missing values. To
125 rectify this omission, we take the averaged ratio between seasonal to permanent waterbody extent between 2015
126 to 2017 for each catchment as a baseline given there are no reported missing values during those years.
127 Subsequently, for each catchment the permanent and seasonal water extent for each time period without satellite
128 data is calculated proportional to number of missing values and known ratio of seasonal to permanent extent in
129 equation 1.

130

$$131 \quad pArea = pO + (nD * (1 - k)) \quad \text{if } nD * \frac{1-k}{p} < 0.05 \quad (1a)$$

$$132 \quad sArea = sO + (nD * k) \quad \text{if } nD * k < 0.05 \quad (1b)$$

133

134

135 where pArea is the permanent surface area, sArea is the seasonal surface area, pO is the observed permanent
136 surface area, sO is the observed seasonal surface area, nD is the number of no data values for the entire catchment,
137 k is the seasonal:permanent surface area ratio and p is the total number of water pixels (e.g. observed + missing
138 values).

139

140 Equation 1 is processed separately for rivers and lakes and only assigned to catchments where less than 5% of the
141 data is missing for any given year, waterbody type and seasonality. For years with more than 5% missing values,
142 the first year with valid observations for any given catchment is used. This creates an accompanying dataset
143 showing the absolute change, percentage change and annual percentage change in surface area of waterbodies for
144 each catchment from 1984 to 2022 (see Data Statement section for interactive map and supplementary figure S1).
145 While this approach does not account for the changing seasonal to permanent water ratios, it does provide an
146 approximation that allows for the analysis of water trends in situations where a few missing values are reported in
147 a much larger catchment region.

148



149 To statistically analyze trends in the surface area of rivers and lakes by catchment, we perform Spearman rank
150 correlations (Spearman, 1987). This correlation measures the monotonicity of the relationship between two
151 parameters, in this case the surface area of the waterbody extent versus time. This measure is ideal considering
152 waterbody extent is often not a linear relationship with time given the interannual variability. To perform this
153 analysis, we limit our analyses to catchments with more than 10 years of results, at least 1km² of detected water
154 surface area, and 95% data coverage for each year. Statistically significant trends are defined by p value less than
155 0.05 or 5%.

156

157 **2.3 Validation**

158

159 To validate and assess the accuracy of the SARL dataset over the 38-years of available water surface change data,
160 we compare the results to manually interpreted extent of river channel belt and lake environments for the year
161 2022. In total, 50 locations (see supplementary material Figure S2) measuring 50 km² were randomly selected
162 (excluding Greenland and Antarctica) to manually map the river channel belt extent using 30 m Landsat 8 imagery.
163 The Landsat imagery corresponds to the same spatial resolution and year of acquisition as the GCB dataset defined
164 by Nyberg et al. (2023) which is used in the current model to define the global channel belt extent (section 2.1).
165 The manual delineation of the channel belt is defined as the encompassing region of the active river and its
166 associated bars, over bank deposits and abandoned channels (Nyberg et al., 2023) to show the maximum
167 geomorphologically-observed extent of the river system through time.

168

169 Following the proposed automated method to classify the current SARL database (e.g., Figure 1), waterbodies
170 defined within the manually defined channel belt extent are classified as riverine whereas all other surface water
171 bodies are defined as lacustrine. Seasonal and permanent waterbody extent for both rivers and lakes are then
172 extracted based on the GSW model (Pekel et al., 2016). The manual river channel belt delineation is subsequently
173 used to compare the accuracy of the automated delineation in capturing the permanent and seasonal water extent
174 of the GSW dataset from 1984 to 2023 (see supplementary Table S2). While the GCB delineation is currently only
175 available for the year 2022, the method provides a baseline assessment for the accuracy of the automated method
176 in capturing the natural variability in seasonal and permanent water extent through time.

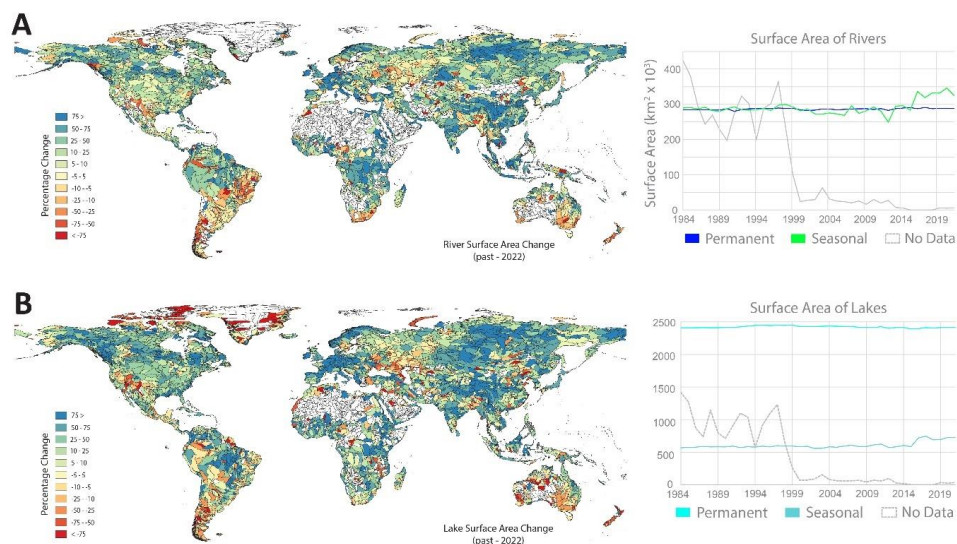
177

178 **3. Results**

179

180 **3.1 Total Water Surface Extent**

181 The permanent surface area of rivers has remained relatively steady over the past 38 years, increasing slightly by
182 1.1% to a total area of 2.9 x 10⁵ km² (Figure 2A). In contrast, the observed seasonal extent of rivers has increased
183 more significantly by 12% with a total area of 3.2 x 10⁵ km² by 2022. The yearly percentage of seasonal to
184 permanent river water extent ranges from 88 to 119%, increasing towards the end of 2022. Similarly, the spatial
185 extent of permanent lake surface area has increased by less than 1% since 1984 to a total area of 24.2 x 10⁵ km²
186 (Figure 2B). The seasonal extent of lakes has however increased significantly, by as much as 27% since 1984 to a
187 total area of 7.2 x 10⁵ km². The ratio of seasonal to permanent water change in lakes is considerably lower than
188 that of rivers, and ranges between 23-31% over the same period.



189

190

191 **Figure 2: Global Water Surface Area Change - The global surface water area of rivers (A) and lakes (B) based on the**
 192 **difference between the first recorded observation to 2022. Graphs show the temporal variability in permanent and**
 193 **seasonal water extent of rivers and lakes by year. See supplementary Figure S3 for percentage change by permanent**
 194 **and seasonal river and lake levels.**

195

196

197 **3.2 Annual Water Surface Area Trends**

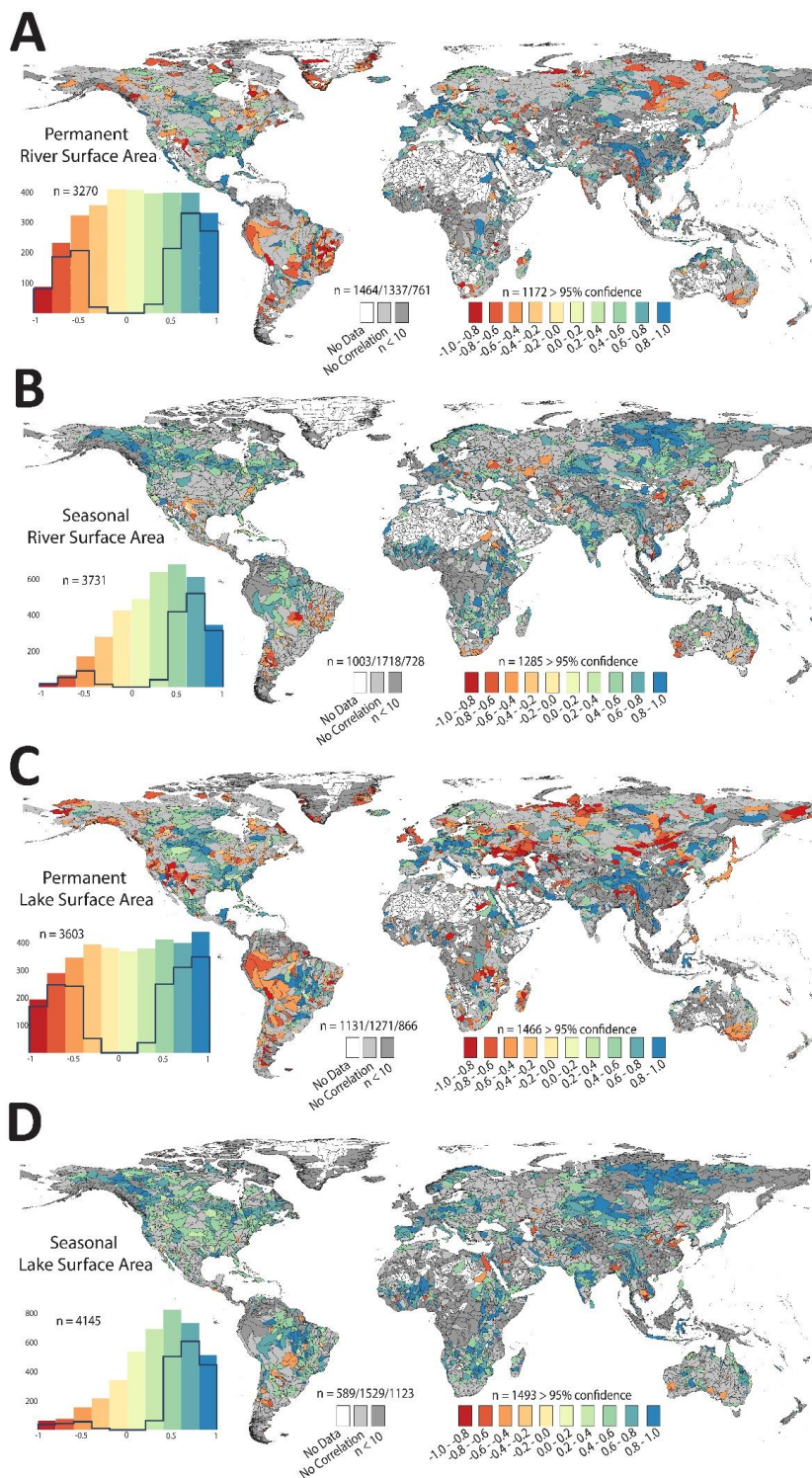
198

199 Figure 3 show trends in permanent versus seasonal water surface area of rivers and lakes from 1984 to 2022 based
 200 on the Spearman rank correlations. Rivers and lakes show a relative normal distribution in spearman correlations
 201 indicating permanent extent of rivers and lakes have experienced both decrease and increase in surface water area.
 202 In total, 47% and 54% of catchments show the permanent surface area of rivers (Figure 3A) and lakes (Figure 3C)
 203 have statistically changed. Out of all catchments with statistically significant changes, 60% and 54% of catchments
 204 have a positive increase in permanent water surface area for rivers and lakes, respectively. In comparison, the
 205 seasonal extent of rivers and lakes is strongly positive indicating an overall increased seasonality with 42% and
 206 49% of catchments showing a significant change for rivers and lakes, respectively (Figure 3B and 3D). Out of all
 207 catchments with statistically significant changes, 84% and 90% of catchments have a positive increase in seasonal
 208 water surface area extent for rivers and lakes, respectively.

209

210

211





213 **Figure 3: Spearman Rank Correlations: Catchments with statistically significant change in permanent (A,C) and**
214 **seasonal (B,D) water surface area since 1984 to 2020 for rivers and lakes. Spearman rank correlations are shown**
215 **ranging from -1 (red) to 1 (blue) where no correlation (light gray) indicate no statistically significant change. The**
216 **histogram for each map shows the spearman correlations with the line indicating the distribution of the statistically**
217 **significant samples (n > 9). See supplementary Figure S4 for entire spearman correlation dataset.**

218

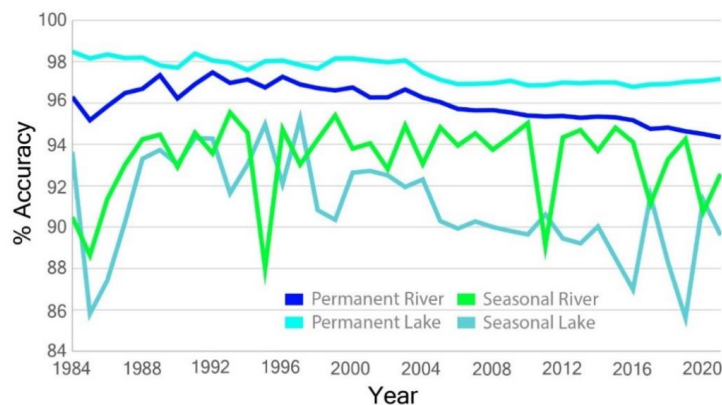
219

220 3.3 Accuracy

221

222 Overall, the SARL dataset showed a 93.8% accuracy to independent estimates of permanent and seasonal lake
223 coverage over the 38-year period as summarized in Figure 4. The permanent extent of lakes is the most consistent
224 and has the highest accuracy ranging between 96 and 98 percent. The accuracy of the permanent river extent is the
225 lowest and ranges from 84 to 91%. Finally, both the seasonal river and seasonal lakes waterbodies show on average
226 an accuracy over 90% but are also the most variable reflecting the variability in yearly water extent.

227



228

229 **Figure 4: Temporal Accuracy of the SARL Database - Overall accuracy of the SARL database for seasonal and**
230 **permanent water extent of lakes and rivers.**

231

232

233

234 4. Discussion

235

236 4.1 Ecosystem Health

237 Understanding temporal and spatial changes in rivers and lakes is key to the study of animal migrations and
238 community dynamics in and around lotic and lentic environments (Ngor et al., 2018). The availability of spawning
239 pathways from ocean to permanent mainstem rivers and lakes to permanent or seasonal tributary streams and lakes
240 is obviously a function of water presence and quality (Briggs et al., 2018). Site selection by aquatic organisms of
241 preferred habitat for feeding and nutrients will also depend on whether or not water is present at candidate locations



242 (Power et al., 2008). In general, riverine and lacustrine animals move in response to changing water extents, so the
243 ability to determine or predict the occurrence of water in global rivers and lakes is a powerful capability.

244 The ability to couple the area of river and stream surface water with adjacent, in-channel terrestrial area supports
245 increased understanding of freshwater microhabitats and freshwater biotic interactions. For example, while fish
246 are obviously restricted to in-water lotic microhabitats like riffles, pools, and runs (whose locations also change
247 spatially and temporally with changing surface water extent), amphibians and certain terrestrial invertebrates
248 regularly move between water and adjacent in-channel dry land (Lowe, 2009). Conceptualizing and delineating
249 freshwater ecosystems as only containing water is therefore under-representative of the area of occupancy and use
250 by many freshwater aquatic organisms.

251 Certain areas in the aqueous stream channel are utilized as flow refugia (Lancaster and Hildrew, 1993; Sakai et
252 al., 2021) or refugia from adverse stream acidification episodes (Baker et al., 1996) during discharge events. The
253 availability of flow refugia is a function of riverine surface water seasonality (Lancaster and Hildrew, 1993).
254 Clearly, the ability to bound the riverine environment as not just the water but the larger area within the channel
255 belt, and then to be able to distinguish between water and adjacent dryland spatially and temporally within that
256 riverine ecosystem, will advance understanding of the distribution and behavior of aquatic organisms in seasonally
257 changing environments. Moreover, current biodiversity models often overlook the importance of terrestrial and
258 aquatic ecosystems in non-perennial systems (Messenger et al., 2021). The current datasets may help to narrow this
259 knowledge gap.

260 The delineation of channel belt areas as geomorphologically-derived riverine environments also has potential for
261 addressing the ‘linearity’ challenge when delineating global freshwater ecosystems and habitats. For certain
262 applications like ecosystem conservation status reporting (e.g. such as is required by the UN Convention on
263 Biological Diversity) and ecosystem accounting (such as is characterized in the guidance from the UN’s System
264 for Environmental and Economic Accounting), area-based measures of ecosystem extent are needed. Except for
265 very large rivers, however, river features are nearly always represented spatially as vector networks where the
266 spatial entity for a river reach is a line segment. The segments may have an attribute for river width, but regardless,
267 the spatial entities representing river reaches are generally not area-based, and including freshwater ecosystems in
268 area-based assessments is challenging. As such, there may be utility in using the Global Channel Belt resource for
269 the spatial delineation of global freshwater ecosystems which would permit area-based assessments of their
270 condition.

271

272 **4.2 Biogeochemical cycles**

273 Existing observations of river surface area at a 30 m global resolution suggest an area of 4.6×10^5 km² at mean
274 annual water discharge (Allen and Pavelsky, 2018). However, our current study at the same 30 m resolution
275 suggests that the permanent extent of rivers is considerably lower at 2.9×10^5 km² or 37% less (Figure 2). On the
276 other hand, the seasonal extent of rivers may contribute another 3.2×10^5 km² for a total area of 6.1×10^5 km² or
277 an area approximately 32% larger than the previous observed estimate. Given rivers are known as a significant
278 source of carbon emissions through water-atmosphere controls, our observations suggest a strong seasonal
279 influence on biogeochemical cycles. Indeed, current estimates based on modelled monthly river surface area



280 suggest that rivers emit $2.0 \pm 0.2 \text{ Pg C y}^{-1}$ and that it is strongly based on seasonal river extents, particularly in
281 temperate and arctic rivers (Liu et al., 2022). The current study furthermore suggests that rivers have, over the past
282 38 years, increased in seasonality by as much as 12%. This entails that carbon emissions from rivers have
283 significantly increased in recent years, especially from high latitude Arctic rivers and high mountainous regions,
284 of for instance, the Tibetan Plateau that experience longer summer months.

285

286 The type of river system is also important in the carbon cycle with high discharge braided rivers recognized to
287 actively erode carbon rich floodplain material to reduce carbon oxidation into the atmosphere (Repasch et al.,
288 2021). The current SARL database show locations of laterally active river systems versus human-controlled river
289 systems. It is also important to recognize the perennial versus non-perennial nature of rivers and lakes which
290 suggests that CO_2 emissions from dry inland waters is overlooked in global calculations contributing an additional
291 6% ($\sim 0.12 \text{ Pg C y}^{-1}$; Keller et al., 2020; Messenger et al., 2021). The decreasing permanent extent of many lakes
292 and rivers (Figure 2) suggest that previously buried sediments with a disproportionately high amount of organic
293 carbon will be increasingly released to the atmosphere (Keller et al., 2020; Hao et al., 2021). Lastly, the damming
294 of rivers and its impact on flow and sedimentation patterns have been shown to eliminate another $48 \pm 11 \text{ Tg C y}^{-1}$
295 (Maavara et al., 2017), although there remain significant knowledge gaps at the local to regional level. The current
296 study and recent reservoir maps (Donchyts et al., 2022; Khandelwal et al., 2022) highlight that rivers and reservoirs
297 have undergone significant historical changes, which may help to further constrain these estimates.

298

299 **4.3 Water Resource Management**

300 Global observations of surface water extent are important to regulate and manage water at a basin scale for a range
301 of different sectors including agriculture, ecosystems, forestry, energy, water supply, environment protection,
302 flood and drought control, irrigation, wastewater treatment, governance and policy to name a few (Garcia et al.,
303 2016; Sheffield et al., 2018). In particular, the data is important to understand the impacts of hydrological extremes
304 of droughts and floods on the short and long-term trends on water allocation and management. The distinction of
305 rivers and lakes is important to understand the storage and transfer of water that impact different sectors and to
306 predict the future impacts of human and climate change stressors.

307 The permanent versus seasonal extent of water is also important to quantify pressures on water resources. In many
308 circumstances, the permanent extent of rivers and lakes has decreased over the past ~4 decades of observations
309 and replaced by an increased seasonality (Figure 3, Pekel et al., 2016; Donchyts, et al., 2016). A change from
310 perennial to seasonal surface water of river directly impacts groundwater recharge and the resulting water table
311 level (Gleeson et al., 2020). For instance, regions of central Australia, Caspian Sea, Aral Sea, western United
312 States, and southern Africa have seen a statistical decrease in permanent waters of both rivers and lakes (Figure 2)
313 that are related to well documented droughts and anthropogenic stresses (Micklin et al., 2007, Van Loon et al.,
314 2016; Mekonnen et al., 2016). There are also rivers that while have experienced relatively stable permanent river
315 water levels, also show decreased annual seasonal flooding, of for example, the Ob River in Siberia (Zemtsov,
316 2019).

317



318 On the other extreme, the Brahmaputra River in Bangladesh and India show an increase in the seasonal extent
319 (Figure 2) that is attributed to increased intensity of river discharge during monsoonal seasons due to increasing
320 Southern Oscillation Index extremes (Mirza, 2011). In other natural river systems, like the lower Amazon basin,
321 we see a slight, statistically significant increase in water surface area of rivers over past three decades (Figure 3).
322 The Amazon water catchment as one of the least human-modified river systems globally (Grill et al., 2019), has
323 increased water levels due to a strengthening Walker circulation (Barichivich et al., 2023). An increase in
324 permanent water surface area for both rivers and lakes is particularly noticeable in central Asia and the Tibetan
325 Plateau as well which has been associated with the acceleration of glacial melting and precipitation (Bo Huang,
326 2014). Furthermore, seasonal water expansion in Siberia can be related to increased thawing of permafrost lakes
327 in summer months (Matthews et al., 2020). Lastly, reservoir expansion, particularly around the Indian
328 subcontinent, eastern Brazil and China, attribute to the increased water surface area of lakes (Donchyts et al.,
329 2022).

330

331

332 **5. Conclusions**

333

334 Here we have presented a new classification on the long-term change in permanent and seasonal extents of both
335 rivers and lakes from 1984 to 2022. Our results show that while the global area of permanent rivers and lakes has
336 remained relatively steady over the past 4 decades (~1% change), the regional variability is considerably higher.
337 The global extent of seasonal rivers and lakes present a different trend, increasing in surface area by 12% and 27%,
338 respectively. For catchments with a statistically significant change, 84% are positively skewed showing an
339 increased seasonality in surface water coverage over the same period. The decreasing perennial extent of many
340 rivers and lakes is often reflected in an increased seasonality of those same water bodies. The results of our analysis
341 are shared as the SARL database, which includes waterbody type, seasonality and spatio-temporal change for
342 global rivers and lakes. This database is a valuable resource and framework for water resource monitoring and
343 assessment of ecosystem health and conservation in different waterbody types.

344

345 **Data Availability**

346 The SARL database developed in this study have been deposited in the Zenodo database under accession
347 code <https://doi.org/10.5281/zenodo.6895820>. An interactive map is available at

348 <https://bjornburnmyberg.users.earthengine.app/view/waterchange>

349

350 **Author Contribution**

351 BN conceived the original idea, designed the methodology and created the database. BN and EL analyzed the
352 original data and created the figures. BN, RS and EL wrote the manuscript.

353 **Competing interests**

354 The authors declare that they have no conflict of interest.

355



356 **Acknowledgements**

357 Any use of trade, product, or firm names is for descriptive purposes only and does not imply endorsement by the
358 U.S. Government. The authors are grateful for the journal-provided reviews and for helpful comments from John
359 W. Jones of the U.S. Geological Survey. Nyberg was funded by the Architectural Element Characterization of
360 Fluvial Systems project by AkerBP ASA.

361

362

363 **References**

364

365 Allen, G., & Pavelsky, T. (2018) Global extent of rivers and streams. *Science* 361, 585-588.

366 DOI:10.1126/science.aat0636

367 Barichivich, J., Gloor, E., Peylin, P., Brienen, R. J. W., Schöngart, J., Espinoza, J. C., & Pattayak, K. C. (2023).

368 Recent intensification of Amazon flooding extremes driven by strengthened Walker circulation. *Science*

369 *Advances*, 4(9), eaat8785. <https://doi.org/10.1126/sciadv.aat8785>

370 Bastviken, D., Cole, J., Pace, M., & Tranvik, L. (2004). Methane emissions from lakes: Dependence of lake
371 characteristics, two regional assessments, and a global estimate. *Global Biogeochemical Cycles*, 18(4).

372 <https://doi.org/https://doi.org/10.1029/2004GB002238>

373 Dottori, F., Szewczyk, W., Ciscar, J.-C., Zhao, F., Alfieri, L., Hirabayashi, Y., Bianchi, A., Mongelli, I., Frieler,

374 K., Betts, R. A., & Feyen, L. (2018). Increased human and economic losses from river flooding with
375 anthropogenic warming. *Nature Climate Change*, 8(9), 781–786. [https://doi.org/10.1038/s41558-018-0257-](https://doi.org/10.1038/s41558-018-0257-z)

376 [z](https://doi.org/10.1038/s41558-018-0257-z)

377 Donchyts, G., Winsemius, H., Baart, F., Dahm, R., Schellekens, J., Gorelick, N., Iceland, C., Schmeier, S., 2022.

378 High-resolution surface water dynamics in Earth's small and medium-sized reservoirs. *Sci. Rep.* 12,

379 13776. <https://doi.org/10.1038/s41598-022-17074-6>

380 Gleeson, T., Wang-Erlandsson, L., Porkka, M., Zipper, S. C., Jaramillo, F., Gerten, D., et al (2020). Illuminating

381 water cycle modifications and Earth system resilience in the Anthropocene. *Water Resources Research*,

382 56, e2019WR024957. <https://doi.org/10.1029/2019WR024957>

383 Gorelick, N., Hancher, M., Dixon, M., Ilyushchenko, S., Thau, D., & Moore, R. (2017). Google Earth Engine:

384 Planetary-scale geospatial analysis for everyone. *Remote Sensing of Environment*.

385 <https://doi.org/10.1016/j.rse.2017.06.031>

386 Grill, G., Lehner, B., Thieme, M., Geenen, B., Tickner, D., Antonelli, F., Babu, S., Borrelli, P., Cheng, L.,

387 Crochetiere, H., Ehalt Macedo, H., Filgueiras, R., Goichot, M., Higgins, J., Hogan, Z., Lip, B., McClain,

388 M.E., Meng, J., Mulligan, M., Nilsson, C., Olden, J.D., Opperman, J.J., Petry, P., Reidy Liermann, C., Sáenz,

389 L., Salinas-Rodríguez, S., Schelle, P., Schmitt, R.J.P., Snider, J., Tan, F., Tockner, K., Valdujo, P.H., van

390 Soesbergen, A., Zarfl, C., 2019. Mapping the world's free-flowing rivers. *Nature* 569, 215–221.

391 <https://doi.org/10.1038/s41586-019-1111-9>

392 Keller, P.S., Catalán, N., von Schiller, D., Grossart, H.-P., Koschorreck, M., Obrador, B., Frassl, M.A., Karakaya,

393 N., Barros, N., Howitt, J.A., Mendoza-Lera, C., Pastor, A., Flaim, G., Aben, R., Riis, T., Arce, M.I., Onandia,

394 G., Paranaíba, J.R., Linkhorst, A., del Campo, R., Amado, A.M., Cauvy-Fraunié, S., Brothers, S., Condon,

395 J., Mendonça, R.F., Reverey, F., Rõdm, E.-I., Datry, T., Roland, F., Laas, A., Obertegger, U., Park, J.-H.,



- 396 Wang, H., Kosten, S., Gómez, R., Feijoó, C., Elozegi, A., Sánchez-Montoya, M.M., Finlayson, C.M., Melita,
397 M., Oliveira Junior, E.S., Muniz, C.C., Gómez-Gener, L., Leigh, C., Zhang, Q., Marcé, R., 2020. Global
398 CO₂ emissions from dry inland waters share common drivers across ecosystems. *Nat. Commun.* 11, 2126.
399 <https://doi.org/10.1038/s41467-020-15929-y>
- 400 Lehner, B., Liermann, C. R., Revenga, C., Vörösmarty, C., Fekete, B., Crouzet, P., Döll, P., Endejan, M., Frenken,
401 K., Magome, J., Nilsson, C., Robertson, J. C., Rödel, R., Sindorf, N., & Wisser, D. (2011). High-resolution
402 mapping of the world's reservoirs and dams for sustainable river-flow management. *Frontiers in Ecology
403 and the Environment*, 9(9), 494–502. <https://doi.org/10.1890/100125>
- 404 Lehner, B., Verdin, K., & Jarvis, A. (2008). New Global Hydrography Derived From Spaceborne Elevation Data.
405 *Eos, Transactions American Geophysical Union*, 89(10), 93–94.
406 <https://doi.org/https://doi.org/10.1029/2008EO100001>
- 407 Matthews, E., Johnson, M. S., Genovese, V., Du, J., & Bastviken, D. (2020). Methane emission from high latitude
408 lakes: methane-centric lake classification and satellite-driven annual cycle of emissions. *Scientific Reports*,
409 10(1), 12465. <https://doi.org/10.1038/s41598-020-68246-1>
- 410 Mekonnen, M. M., & Hoekstra, A. Y. (2016). Four billion people facing severe water scarcity. *Science Advances*,
411 2(2), e1500323.
- 412 Messenger, M.L., Lehner, B., Cockburn, C., Lamouroux, N., Pella, H., Snelder, T., Tockner, K., Trautmann, T.,
413 Watt, C., Datry, T., 2021. Global prevalence of non-perennial rivers and streams. *Nature* 594, 391–397.
414 <https://doi.org/10.1038/s41586-021-03565-5>
- 415 Messenger, M.L., Lehner, B., Grill, G., Nedeva, I., Schmitt, O., 2016. Estimating the volume and age of water
416 stored in global lakes using a geo-statistical approach. *Nat. Commun.* 7, 13603.
417 <https://doi.org/10.1038/ncomms13603>
- 418 Micklin, P. (2007). The Aral Sea Disaster. *Annual Review of Earth and Planetary Sciences*, 35(1), 47–72.
419 <https://doi.org/10.1146/annurev.earth.35.031306.140120>
- 420 Mirza, M. M. Q. (2011). Climate change, flooding in South Asia and implications. *Regional Environmental
421 Change*, 11(1), 95–107. <https://doi.org/10.1007/s10113-010-0184-7>
- 422 Neumann, K., Sietz, D., Hilderink, H., Janssen, P., Kok, M., & van Dijk, H. (2015). Environmental drivers of
423 human migration in drylands – A spatial picture. *Applied Geography*, 56, 116–126.
424 <https://doi.org/https://doi.org/10.1016/j.apgeog.2014.11.021>
- 425 Nyberg, B., Howell, J. (2015). Is the present the key to the past? A global characterization of modern sedimentary
426 basins. *Geology* 43 (7): 643–646.
- 427 Nyberg, B., Henstra, G., Gawthorpe, R.L. Rodmar, R., Ahokas, J. (2023) Global scale analysis on the extent of
428 river channel belts. *Nat Commun* 14, 2163.
- 429 Oki, T., & Kanae, S. (2006). Global Hydrological Cycles and World Water Resources. *Science*, 313(5790), 1068
430 LP – 1072. <https://doi.org/10.1126/science.1128845>
- 431 Pekel, J.-F., Cottam, A., Gorelick, N., & Belward, A. S. (2016). High-resolution mapping of global surface water
432 and its long-term changes. *Nature*, 540(7633), 418–422. <https://doi.org/10.1038/nature20584>
- 433 Sheffield, J., Wood, E.F., Pan, M., Beck, H., Coccia, G., Serrat-Capdevila, A., Verbist, K., 2018. Satellite Remote
434 Sensing for Water Resources Management: Potential for Supporting Sustainable Development in Data-Poor
435 Regions. *Water Resour. Res.* 54, 9724–9758. <https://doi.org/https://doi.org/10.1029/2017WR022437>



- 436 Siebert, S., Burke, J., Faures, J. M., Frenken, K., Hoogeveen, J., Döll, P., & Portmann, F. T. (2010). Groundwater
437 use for irrigation – a global inventory. *Hydrol. Earth Syst. Sci.*, *14*(10), 1863–1880.
438 <https://doi.org/10.5194/hess-14-1863-2010>
- 439 Song, C., Huang, B., Richards, K., Ke, L., & Hien Phan, V. (2014). Accelerated lake expansion on the Tibetan
440 Plateau in the 2000s: Induced by glacial melting or other processes? *Water Resources Research*, *50*(4),
441 3170–3186. <https://doi.org/https://doi.org/10.1002/2013WR014724>
- 442 van Dijk, A. I. J. M., & Renzullo, L. J. (2011). Water resource monitoring systems and the role of satellite
443 observations. *Hydrol. Earth Syst. Sci.*, *15*(1), 39–55. <https://doi.org/10.5194/hess-15-39-2011>
- 444 Van Loon, A. F., Gleeson, T., Clark, J., Van Dijk, A. I. J. M., Stahl, K., Hannaford, J., Di Baldassarre, G., Teuling,
445 A. J., Tallaksen, L. M., Uijlenhoet, R., Hannah, D. M., Sheffield, J., Svoboda, M., Verbeiren, B., Wagener,
446 T., Rangelcroft, S., Wanders, N., & Van Lanen, H. A. J. (2016). Drought in the Anthropocene. *Nature*
447 *Geoscience*, *9*(2), 89–91. <https://doi.org/10.1038/ngeo2646>
- 448 Wada, Y., de Graaf, I. E. M., & van Beek, L. P. H. (2016). High-resolution modeling of human and climate impacts
449 on global water resources. *Journal of Advances in Modeling Earth Systems*, *8*(2), 735–763.
450 <https://doi.org/https://doi.org/10.1002/2015MS000618>
- 451 Zemtsov, V. A., Vershinin, D. A., Khromykh, V. V., & Khromykh, O. V. (2019). Long-term dynamics of maximum
452 flood water levels in the middle course of the Ob River. *IOP Conference Series: Earth and Environmental*
453 *Science*, *400*(1), 12004. <https://doi.org/10.1088/1755-1315/400/1/012004>