# Influence of Ocean Alkalinity Enhancement with Olivine or Steel Slag on a Coastal Plankton Community in Tasmania

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# 10

11 Abstract. Ocean alkalinity enhancement (OAE) aims to increase atmospheric CO<sub>2</sub> sequestration in the oceans through the 12 acceleration of chemical rock weathering. This could be achieved by grinding rocks containing alkaline minerals and 13 adding the rock powder to the surface ocean where it dissolves and chemically locks CO<sub>2</sub> in seawater as bicarbonate. 14 However, CO<sub>2</sub> sequestration during dissolution coincides with the release of potentially bio-active chemicals and may 15 induce side effects. Here, we used 53 L microcosms to test how coastal plankton communities from Tasmania respond to 16 OAE with olivine (mainly Mg<sub>2</sub>SiO<sub>4</sub>) or steel slag (mainly CaO and Ca(OH)<sub>2</sub>) as alkalinity sources. Three microcosms were 17 left unperturbed and served as a control, three were enriched with olivine powder (1.9 g L<sup>-1</sup>), and three with steel slag 18 powder (0.038 g L<sup>-1</sup>). Olivine and steel slag powders were of similar grain size. Olivine was added in a higher amount than 19 the steel slag since previous tests evidenced that it would have released less alkalinity over the 3-week experiment. Olivine 20 and steel slag powders were of similar grain size, but the amount of added olivine needed to be much higher than the steel 21 slag because less alkalinity is released by the olivine than the steel slag over the 3 week experiment. Phytoplankton and 22 zooplankton community responses as well as some biogeochemical parameters were monitored for 21 days. Olivine and 23 steel slag additions increased total alkalinity by 29 µmol kg<sup>-1</sup> and 361 µmol kg<sup>-1</sup> respectively, which corresponds to a 24 theoretical increase of 0.9 % and 14.8 % of the seawater storage capacity for atmospheric CO<sub>2</sub>. Olivine and steel slag 25 released silicate nutrients into the seawater column, but steel slag released considerably more and also significant amounts 26 of phosphate. After 21 days, no significant difference was found in dissolved iron concentrations (>100 nmol  $L^{-1}$ ) in the 27 treatments and the control. Both minerals released dissolved aluminium (>50 nmol L<sup>4</sup>). The slag addition increased 28 dissolved manganese concentrations (771 nmol L<sup>-1</sup>), while olivine increased dissolved nickel concentrations (37 nmol L<sup>-1</sup>) 29 1). Correspondingly, tThe slag treatment increased the total particulate manganese concentrations (22 nmol L<sup>+</sup>), while olivine increased the total particulate nickel (5 nmol L<sup>-1</sup>), which was consistent with the increase in the dissolved 30 31 concentrations of these trace metals in seawater. There was no significant difference in total chlorophyll a concentrations 32 between the treatments and the control, likely due to nitrogen limitation of the phytoplankton community. However, flow 33 cytometry results indicated an increase in the cellular abundance of several smaller (~<20 µm) phytoplankton groups in 34 the olivine treatment-compared to the slag treatment and the control. The abundance of larger phytoplankton (~>20 µm) 35 decreased much more in the control than in the mineral addition-treatments after day 10. Furthermore, the maximum 36 quantum yields of photosystem II ( $F_v/F_m$ ) were higher in slag and olivine treatments, suggesting that mineral additions 20

37 increased photosynthetic performance. The zooplankton community composition was also affected with the most notable

38 changes being observed in the dinoflagellate Noctiluca scintillans and the appendicularian Oikopleura sp. in the olivine

39 <u>treatment</u>. Overall, steel slag is much more efficient for CO<sub>2</sub> removal with OAE than olivine and appears to induce less

40 change in the plankton community when relating the CO<sub>2</sub> removal potential to the level of environmental impact that was
41 observed here.

42

## 43 1 Introduction

Keeping global warming below 2 °C requires immediate emissions reduction. Additionally, between 450-1100 Gigatonnes
of carbon dioxide (CO<sub>2</sub>) need to be removed from the atmosphere by 2100 (Smith et al., 2023). This could be achieved
with a portfolio of terrestrial and marine Carbon Dioxide Removal (CDR) methods. Ocean alkalinity enhancement (OAE)
is a marine CDR method that could theoretically contribute significantly to the global CDR portfolio (Ilyina et al., 2013;
Feng et al., 2017; Lenton et al., 2018).

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Alkalinity is generated naturally when rock weathers and it has control on the ocean's chemical capacity to store  $CO_2$ (Schuiling and Krijgsman, 2006). Natural rock weathering is currently responsible for about 0.5 Gt of atmospheric  $CO_2$ sequestration every year (Renforth and Henderson, 2017). The idea behind OAE is to accelerate natural rock weathering by extracting calcium- or magnesium-rich rocks, such as olivine, pulverizing them, and spreading them onto the sea surface to increase chemical weathering rates (Hartmann et al., 2013). The weathering (i.e., dissolution) of these alkaline minerals will consume protons (H<sup>+</sup>), which shifts the carbonate chemistry equilibrium in seawater from  $CO_2$  towards increasing bicarbonate (HCO<sub>3</sub><sup>-</sup>) and carbonate ion (CO<sub>3</sub><sup>2-</sup>) concentrations:

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58 
$$CO_2+H_2O \Rightarrow H_2CO_3 \Rightarrow HCO_3^- + H^+ \Rightarrow CO_3^{2-} + 2H^+$$

(1)

thereby making new space for atmospheric  $CO_2$  to be dissolved in seawater and permanently stored. Previous model studies have shown that OAE can mitigate climate change significantly by increasing the oceanic uptake of  $CO_2$  from the atmosphere (Kohler et al., 2010; Paquay and Zeebe, 2013; Keller et al., 2014; Lenton et al., 2018). For example, the study by Burt et al. (2021) suggested that the total global mean dissolved inorganic carbon (DIC) inventories would increase by 156 GtC after total alkalinity is enhanced at a rate of 0.25 Pmol year<sup>-1</sup> in 75-year simulations.

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There are a variety of alkaline minerals that could be used for OAE. A widely considered naturally occurring mineral is forsterite, a (Mg<sub>2</sub>SiO<sub>4</sub>)-rich olivine. This type of olivine is abundant in ultramafic rock such as dunite, constituting at least 88 % of the rock composition (Ackerman et al., 2009; Su et al., 2016). Olivine occurs in the Earth's crust but is more abundant in the upper mantle. There are at least several billion tons of olivine resources on Earth (Caserini et al., 2022). However, the extraction of olivine in 2017 was only around 8.4 Mt year<sup>-1</sup> (Reichl et al., 2018), which is about two orders of magnitude below the mass needed for climate-relevant OAE with olivine (Caserini et al., 2022). The net reaction for CO<sub>2</sub> sequestration with Mg<sub>2</sub>SiO<sub>4</sub> is:

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$$Mg_2SiO_4 + 4CO_2 + 4H_2O \rightarrow 2Mg^{2+} + 4HCO_3^- + H_4SiO_4$$
 (2)

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Another potential OAE source material is steel slag (Renforth, 2019), a by-product of steel manufacturing. During steel manufacturing, high-purity calcium oxide (CaO) is used to improve the quality of the steel through accumulation of unwanted materials such as sulphur and phosphorus. Steel slag mainly contains CaO, SiO<sub>2</sub>, Al<sub>2</sub>O<sub>3</sub>, Fe<sub>2</sub>O<sub>3</sub>, MgO, and MnO (Kourounis et al., 2007), and the chemical composition can vary depending on the manufacturing process (Wang et al., 2011). Due to the presence of CaO and potentially other alkaline components, steel slag can increase alkalinity when dissolved in seawater. The chemical reaction for CO<sub>2</sub> sequestration with CaO is:

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$$CaO + H_2O \rightarrow Ca(OH)_2$$
 and  $Ca(OH)_2 + 2CO_2 \rightarrow Ca^{2+} + 2HCO^-$  (3)

Some of the steel slag that is produced during steel manufacturing is further used (e.g., for road construction and civil
engineering) but in some countries like China, 70.5 % of steel slag is left unused and stored in dumps (Guo et al., 2018).
In 2016, more than 300 million tons of steel slag was not used effectively, thereby occupying the land and raising
environmental concerns (Guo et al., 2018). The effective alkaline composition, availability, and relatively low cost of the
raw materials make olivine and steel slag potential source materials for OAE.

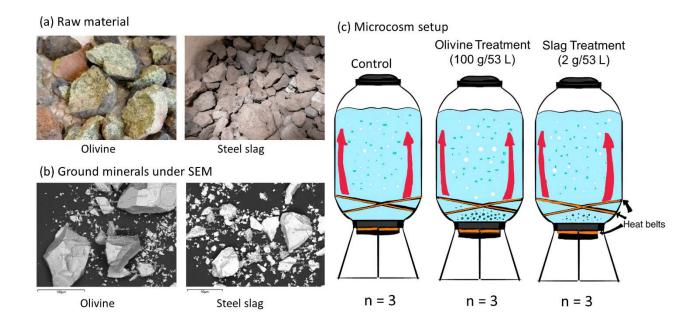
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91 To assess whether OAE is viable, it needs to be understood how its application may affect marine biota such as plankton 92 and the biogeochemical fluxes they drive. Some data on the effects of OAE with sodium hydroxide (NaOH) on plankton 93 communities have recently been published (Ferderer et al., 2022; Subhas et al., 2022), but to the best of our knowledge, no 94 such data are available for olivine- and/or slag-based OAE. Chemical perturbations via olivine and slag should be like 95 those by NaOH in that they increase seawater pH and shift the carbonate chemistry equilibrium (see Eq. 1). However, there 96 would be additional chemical perturbations because minerals contain a variety of potentially bioactive elements that are 97 released into the environment when they dissolve in seawater (Bach et al., 2019). One particular concern is that natural and anthropogenic minerals such as olivine and steel slag are rich in bioactive metals that are usually scarce in the ocean, such 98 99 as iron (Fe), copper (Cu), nickel (Ni), manganese (Mn), zinc (Zn), cadmium (Cd), and chromium (Cr). Many of these trace 100 metals are essential micronutrients for phytoplankton growth (Sunda, 2000; Sunda, 2012), such as being co-factors for 101 various metalloenzymes (summarized by Twining and Baines, 2013). It is possible that the addition of alkaline minerals 102 may benefit phytoplankton by providing trace metals currently limiting phytoplankton growth (Falkowski, 1994; Basu and 103 Mackey, 2018). For instance, the addition of Fe is well known to stimulate phytoplankton blooms in those vast ocean 104 regions where Fe levels limit growth (Boyd et al., 2007; Moore et al., 2013). However, some trace metals can also inhibit 105 phytoplankton growth, and different phytoplankton species have different requirements and tolerances for trace metals 106 (Sunda, 2012) so the addition of trace metals via OAE may change phytoplankton community composition.

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- Here, we describe a microcosm experiment with coastal Tasmanian plankton communities that was used to investigate: (1)
   how effectively OAE via the application of finely ground olivine and steel slag could sequester atmospheric CO<sub>2</sub>, and (2)
- 110 if /how olivine and steel slag additions affect various components of the plankton community.
- 111

#### 112 2 Methodology

#### 113 2.1 Microcosm setup



#### 114

Fig. 1. Experimental design and alkalinity sources. (a) Raw materials used as alkalinity sources: olivine (left) and steel slag (right).
 Olivine and steel slag were originally larger than 20 mm. (b) Ground minerals observed with a scanning electron microscope (SEM). (c)
 Microcosm setup: each microcosm enclosed ~ 53 L of surface seawater with natural plankton communities. Olivine and steel slag
 treatments and the control were kept in a temperature-controlled room and two heat belts were attached to the bottom of each microcosm
 to create convective circulation.

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121 We used nine 53 L transparent Kegland® Fermzilla conical unitank fermenters (polyethylene terephthalate) (Fig. 1) as 122 microcosms to incubate natural plankton communities. All microcosms were prewashed with hydrochloric acid (10 % v/v) 123 and rinsed five times with 18.2 MΩ Milli-Q water. Seawater with coastal plankton communities was collected at Battery 124 Point, Tasmania (42.892°S, 147.337°E) within 2 hours by lowering the microcosms into the ocean with a crane and filling 125 them in a manner similar to a Niskin bottle, as described in detail in Ferderer et al. (2022). A sieve with a mesh size of 2 126 mm was attached to the top and bottom of the microcosms during filling to avoid the entrapment of large and patchily distributed organisms in the microcosms. The enclosed seawater weight was initially between 52.35-54.70 kg. After 127 128 seawater collection, filled microcosms were immediately transported back to the Institute for Marine and Antarctic Studies 129 (University of Tasmania) on a truck and transferred within 75 min into a temperature-controlled room set to 7.5-8 °C. Two 130 heat belts were attached to the bottom of each microcosm to induce a convective mixing current (Ferderer et al., 2022). 131 Seawater temperature inside the microcosms was about 13.5 °C due to the heating effects of the heat belts and was the 132 same as the sampled region. LED light strips were used to provide an average light intensity of 236  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup> (ranging from 208 to 267  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup>) with a daily light-dark cycle of 10:14 hours. The light intensity was the 133 134 average light intensity in each microcosm measured with a LICOR light meter at 0.15 m depth within the microcosm. 135 Microcosms positioned in the temperature-controlled room were shuffled anti-clockwise every day to ensure similar light

- 136 intensity for each microcosm throughout the experiment. Treatments were established 24 hours after collecting the seawater.
- 137 The total alkalinity released per amount of mineral powder added was much higher for the slag powder than the olivine
- powder in our preliminary test trials. So, three microcosms were enriched with 100 g of olivine powder, three microcosms
- 139 with 2 g of steel slag powder, while the remaining three microcosms were left unperturbed and served as controls.
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#### 142 **2.2** Preparation of olivine and steel slag powder

143 The olivine rocks were provided by Moyne Shire Council who sourced the mineral from a quarry in Mortlake, Victoria, 144 Australia. The Basic Oxygen Slag (hereafter referred to as "slag") was provided by Bradley Mansell who sourced the 145 material from Liberty Primary Steel Whyalla Steelworks in Whyalla, South Australia, Australia. Upon delivery, the olivine 146 rocks were 40-80 mm in diameter, and slag aggregates were 20-50 mm in diameter. These were crushed to smaller than 10 147 mm pieces using a hydraulic crusher. The crushed material was further ground with a ring mill with a chrome milling pot. 148 Afterwards, finely-ground samples were sieved to get samples with  $150 \sim 250 \mu m$  grain size. The sieved olivine and slag grains were inspected for their appearance and elemental composition using a Hitachi SU-70 analytical field emission 149 150 scanning electron microscope (SEM), and energy dispersive spectrometers (Central Science Laboratory (CSL), University 151 of Tasmania). Grain size spectra were determined with a Sympatec QICPIC particle size analyser LIXCELL (CSL, 152 University of Tasmania).

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## 154 2.3 Seawater sampling

Seawater was transferred with a peristaltic pump from the microcosms at a depth of about 0.15 m into 1 L acid-washed sampling bottles (LDPE) using an acid-washed silicon tube. Seawater in these bottles was then subsampled for dissolved trace metal samples, filtrations, Fast Repetition Rate fluorometry (FRRf), and flow cytometry analysis. Samples for nutrients and total alkalinity (TA) were transferred using the same pump but through a silicone tube into 80 mL HDPE bottles. Total alkalinity and macronutrient samples were filtered during this process through a 0.2  $\mu$ m nylon filter attached to the silicone tube to remove all particles and organisms > 0.2  $\mu$ m.

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#### 162 2.4 Salinity, nutrients, carbonate chemistry, and trace metal analysis

163 Salinity was measured before and at the end of the experiment using a HACH HQ40d portable meter. The pH<sub>T</sub> (total scale) 164 and temperatures were measured daily (2-3 hours after the onset of the light period) using a pH meter (914 pH/Conductometer Metrohm). We recorded voltages and temperature from the pH meter and calibrated the  $pH_T$  at original 165 temperature at sampled time using the certified reference material (CRM) Tris buffer following the method described in 166 167 SOP6a by Dickson et al. (2007). Briefly, the standard buffer's pH and voltage at different temperature gradients were 168 recorded, and temperature vs. voltage polynomial regression data were generated for calculating calibrated pH values (pH<sub>T</sub>) 169 (refer to Eq. 3 in SOP6a of Dickson et al. (2007)). The regression could then be used to obtain a CRM pH value for each 170 temperature and to calibrate the pH measured in the microcosms to the total pH scale.

172 Total alkalinity was sampled every four days. It was measured in duplicate using a Metrohm 862 Compact Titrosampler 173 coupled with an Aquatrode Plus with PT1000 temperature sensor following the SOP3b open-cell titration protocol 174 described in Dickson et al. (2007). Filtered TA samples were stored at 8 °C for a maximum of 23 days before measurement. 175 Titration curves were evaluated using the "calkulate" script within PyCO2sys by Humphreys et al. (2022). The carbon 176 chemistry equilibrium was calculated with the R package "seacarb" Gattuso et al. (2023) from pH<sub>T</sub>, TA, phosphate, silicate, 177 temperature, and salinities using stoichiometric equilibrium constants from Lueker et al. (2000). Dissolved macronutrients 178 were measured every second day using standard spectrophotometric methods developed by Hansen and Koroleff (1999) 179 on the day the samples were taken from the microcosms.

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181 Dissolved trace metal concentrations were measured four times during the experiment: a few hours before olivine and slag 182 were added, a few hours after these minerals were added on day 2, near the middle of the experiment on day 13, and at the 183 end of the experiment on day 22. Sixty mL of seawater was collected using an acid-washed 60 mL syringe, and the seawater 184 was filtered through 25 mm diameter 0.2 µm pore size polycarbonate filters. Unfortunately, we did not notice that 0.2 µm 185 pore size nylon filters (acid washed) were used during sampling on days 1 and 2 so we refiltered these seawater samples 186 again using 0.2 µm pore size polycarbonate filters after one month. All seawater samples were diluted approximately 20-187 fold by weight using Milli-Q water (18.2 MQ·cm grade) and acidified using 1 % ultrapure HCl. These samples were 188 analysed using Sector Field Inductively Coupled Plasma Mass Spectrometry (SF-ICP-MS) employing multiple resolution 189 settings to overcome major spectral interferences. Due to the presence of abundant major metal ions in our samples, such 190 as Na and Mg, natural open-ocean seawater from the Southern Ocean with very low trace metal concentrations was diluted 191 20 times with Milli-Q water and used as a representative blank. The same Southern Ocean seawater was enriched with 192 different gradients of trace metal standards to calculate the samples' trace metal concentrations. Five of the total 36 samples 193 had abnormal trace metal concentrations, and 2 of them were from day 1. We considered values as outliers using the 194 interquartile range (IQR) criterion on pre-addition data, and if values are more than 10 times higher than replicates, they 195 are also considered as outliers. These samples containing outliers were excluded from the data analysis (Table S1.). The 196 major likely source of these metal contaminations is sampling in the temperature control room, where precautions were 197 insufficiently implemented.

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#### **199 2.5 Particulate matter and plankton community analysis**

200 Chlorophyll *a* was sampled every second day by filtering the seawater through glass fibre filters (GF/F, pore size =  $0.7 \mu m$ , 201 diameter =25 mm), and filters were stored in 15 mL polypropylene tubes wrapped with aluminium foil and stored at -80 °C 202 for 50-70 days before measurement. Each filter was immersed in 10 mL 100 % methanol for 18-20 h to extract chlorophyll 203 from phytoplankton and these samples were analysed on a Turner fluorometer (Model 10-AU) following the method 204 described by Evans et al. (1987).

205

Phytoplankton flow cytometry samples were fixed with 40 µL of a mixture of formaldehyde-hexamine (18 %:10 % v/w) added to 1400 µL of seawater sample. All bacteria samples (700 µL) were fixed with 14 µL glutaraldehyde (Electron-microscope grade, 25 %). After mixing samples with fixatives, samples were stored for 25 minutes at 10 °C, then flash-frozen in liquid nitrogen, and stored at -80 °C until measurement 83-86 days later. Directly before the measurement, samples were thawed at 37 °C. Bacteria samples were stained with SYBR green I (diluted in dimethylsulfoxide) at a final

211 ratio of 1:10000 (SYBR Green I: sample).

212

A Cytek Aurora flow cytometer (Cytek Biosciences) was used to quantify the abundance of fluorescing particles such as phytoplankton or stained bacteria. Phytoplankton groups were distinguished based on their fluorescence signal intensity of different laser excitation/emission wavelength combinations and forward scatter (FSC). The yellow-green laser (centre wavelength: 577 nm), in combination with FSC signal strength, was used to separate cyanobacteria and cryptophytes from other phytoplankton. The violet laser (centre wavelength: 664 nm) in combination with FSC was used to distinguish picoeukaryotes, nanoeukaryotes, and microphytoplankton. The blue laser (centre wavelength: 508 nm) in combination with FSC was used to distinguish bacteria from other living (i.e., DNA-containing) particles (Fig. S. 1).

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221 The biovolume of each classified flow cytometry phytoplankton type was calculated using the equation:

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223 
$$Biovolume = Cell number count \times \left(\frac{FSC}{10248}\right)^{2.14}$$
 (4)

224

where biovolume is the biovolume of the phytoplankton ( $\mu$ m<sup>3</sup>), cell number is the cell count per mL of sample, and the FSC is the forward scatter signal value from the flow cytometry. This equation is calculated based on the relationship between biovolume and FSC for different phytoplankton species (Selfe, 2022). The biovolume of each phytoplankton type was then divided by the total biovolume of all phytoplankton type to calculate the biovolume proportion of each phytoplankton type (Biovolume prop.). This derived value was used to estimate the phytoplankton composition in each microcosm.

231

Phytoplankton photosynthetic performance was estimated from the rapid light curves measured with an FRRf (FastOcean Sensor FRRf3, Chelsea Instruments Group) every second day following the protocol adapted from Schallenberg et al. (2020). Samples were kept in the dark for 20 minutes before the measurement and then added to the FRR fluorometry cuvette, which was temperature-controlled at 13.5 °C. Filtered natural seawater was used for blank correction. A channel with three light wavelengths (450, 530, and 624 nm) was used in each acquisition sequence. At least 10 acquisitions were measured for each sample. The maximum electron transport rate (ETR<sub>max</sub>), initial slope of the rapid light curve ( $\alpha$ ), and the light-saturation parameter (E<sub>k</sub>) were calculated using the equation described by Platt et al. (1980) without photoinhibition:

240 
$$ETR = ETR_{max} \left[ 1 - e^{-\frac{\alpha E}{ETR_{max}}} \right]$$
(5)

241

239

These parameters together with the maximum quantum yield of PSII ( $F_v/F_m$ ) were used to compare the photosynthetic performance of the phytoplankton communities in different microcosms.

244

Seawater was sampled before the treatment and at the end of the experiment for particulate trace metal concentrations. Samples of 100 mL were filtered through an acid-cleaned polycarbonate filter (25 mm diameter, 0.8  $\mu$ m pore size) and placed in an acid-cleaned polypropylene filter holder in a trace metal-clean laminar flow bench. The filters were washed with the EDTA-oxalate reagent (1.4 mL) twice (8 min total) and rinsed with chelexed NaCl solution (0.6 mol L<sup>-1</sup> with 2.38 mmol L<sup>-1</sup> of HCO<sub>3</sub><sup>-</sup>, pH=8.2) 10 times (1.5 mL aliquots) (Tovar-Sanchez et al., 2003; Tang and Morel, 2006). Filters were stored in acid-washed well plates at -20 °C before analysis. The digestion process followed the method reported by Bowie et al. (2010). Briefly, all samples and triplicate certified reference materials plankton standards (50 mg/vial) were digested in a mixture of strong ultrapure acids (750  $\mu$ L 12 mol L<sup>-1</sup> HCl, 250  $\mu$ L 40 % HF, 250  $\mu$ L 14 mol L<sup>-1</sup> HNO<sub>3</sub>) in 15 mL Teflon perfluoroalkoxy (PFA) vials on a 95 °C hot plate for 12 h in a fume hood. They were then dry evaporated for 4 h and resuspended in 10 % v-v ultrapure HNO<sub>3</sub>. All prepared solutions had indium as internal standard added to a final concentration of 10  $\mu$ g L<sup>-1</sup>. Three pre-mixed multi-element standard solutions (MISA) were prepared as external calibration standards.

257

Particulate organic carbon (POC) was sampled by filtering 100 mL of seawater from each microcosm. Glass fibre filters
(Whatman GF/F, pore size =0.7 μm, diameter =13 mm) were pre-combusted at 400 °C for 6 h. Filters were stored at -20 °C
before measurement. Samples were treated via fuming with 2N HCl to remove carbonates overnight and dried in the oven
for 4h. Finally, filters were folded into silver cups and stored in a desiccator until analysis. Samples were analysed for
carbon with a Thermo Finnigan EA 1112 Series Flash Elemental Analyser (CSL, University of Tasmania).

263

Biogenic silica (BSi) concentrations were analysed every 4 days by filtering 100 mL of seawater from each microcosm. Mixed Cellulose Ester (MCE) membrane filters (diameter = 25 mm, pore size =  $0.8 \mu$ m) were used for BSi samples. BSi filters were placed in a plastic petri dish and stored at -20 °C before measurement. Filters were processed using the hot NaOH digestion method of Nelson et al. (1989). The final solution was measured using the same process as the dissolved silicate (see section 2.4).

269

A self-made plastic zooplankton net (20 mm height and 15 mm width) with a 210  $\mu$ m mesh size was acid-washed first and then used to collect zooplankton from microcosms before mineral addition on day 2, near the middle (day 13), and at the end of the experiment (day 23). Samples were stored in 10 % formalin seawater solutions and kept at room temperature until measurements. Zooplankton were quantified and identified under a Leica M165C microscope fitted with a Canon 5D camera. The number of zooplankton from one mini-trawl in each collection was converted to the unit of individual L<sup>-1</sup> and used for data analysis. The diversity of zooplankton communities was estimated with the Shannon Diversity Index (H) calculated as:

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279

278  $H = -\sum (pi \times \ln (pi))$ (6)

where pi is the proportion of the entire zooplankton community made up of individual species abundance, and ln is the natural logarithm.

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## 284 2.6 Statistic analysis

R studio was used for data analyses. Generalized additive models (GAMs) from the package "mgcv" were fitted to the data
to predict the changes over time. The GAMs all shared the same equations:

287

 $288 \qquad Y = s(Day),$ 

289

(7)

in which Y presents the dependent variable and s(Day) is the smooth term of the day of the experiment. Another GAM wasused to detect significant differences between treatments and the control:

292

293 
$$Y = Treatment + s(Day) + s(Day, by = oTreatment)$$
(8)

294

In this equation, the variable "Treatment" includes three conditions: "Control", "Slag" and "Olivine"; while "oTreatment"
is the ordered factor of the variable "Treatment" which allowed us to compare the GAMs smooth terms from different
treatments and the control (Simpson, 2017).

298

When comparing GAMs, P-means represent the p-value obtained from comparing two GAMs, such as the control and the olivine treatment. If P-means is below 0.05, it indicates that the mean values of the two GAMs exhibit significant differences over the course of the experiment. Conversely, if P-means is equal to or greater than 0.05, it suggests that the two GAMs have similar mean values. In contrast, P-smooths represents the p-value derived from comparing the smooth terms of two GAMs. If P-smooths is below 0.05, it indicates that the two GAMs demonstrate significantly different trends in their change over time.

305

For the analysis of trace metal concentrations and zooplankton abundance, Generalized Linear Models (GLMs) from the
 'stats' package were fitted to the data to determine significant differences between treatments and the control. The selection
 of specific GLMs was based on the distribution of the raw data. One GLM equation is

310 
$$Y = Treatment + \frac{Day}{22} + (\frac{Day}{22})^2$$
 (9)

311

309

with family = Gamma, where Y represents the measured parameter (abundance of a zooplankton species and dissolved
trace metal concentrations); treatment is the conditions ("Control", "Slag" and "Olivine"); and Day represents the day of
the experiment. The other GLM equation,

315

$$316 \quad Y = Treatment + Day \tag{10}$$

317

with family = Gaussian, was employed for particulate trace metal data and the Shannon Diversity Index. To compare the
contribution of the three treatments on the measured parameters, Tukey's significant difference test was conducted on the
GLMs using the 'glht' function.

321

## 322 3. Results

## 323 **3.1** Elemental composition and grain size of the finely-ground minerals

324 SEM analysis revealed the approximate elemental composition of olivine and slag powder (Table 1). Based on this analysis 325 the olivine composition resembles the Mg-rich olivine mineral "forsterite" ( $Mg_2SiO_4$ ). The particle size spectrum of olivine 326 powder is shown in detail in Fig. S2. Roughly 69 % of the olivine particles, when measured by volume, fell within the diameter range of 35 - 300  $\mu$ m. Additionally, SEM analysis revealed high levels of Ca and O in the slag, indicative of the considerable Ca(OH)<sub>2</sub> and CaO content of the powder (Table 1; please note that H cannot be measured with the applied method). The particle size measurement (Fig. S2) showed that 78 % of the ground slag particles were between 35 - 300

330

μm.

331

**Table 1.** The weight percentage of elements from two minerals. Unit: wt %.

Element	0	Ca	Mn	Si	Mg	Fe	Al	Ti	Cr	Ni
Olivine	39.9	0.4		19.9	26.4	13.0	1.0			0.8
Steel slag	41.9	36.0	7.0	6.5	4.3	3.7	3.4	1.7	1.6	

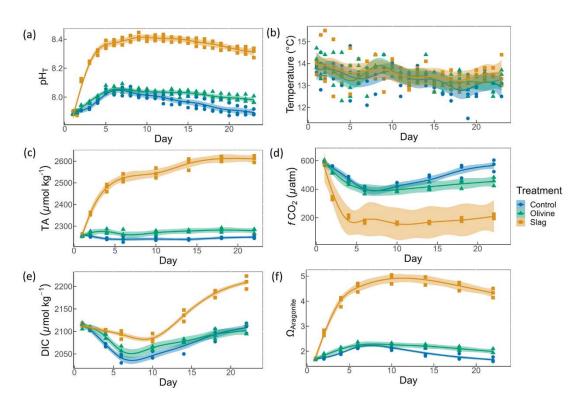
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334

#### 335 **3.2** Physical and chemical conditions over the course of the experiment.

On day 2 of the experiment, when olivine particles were introduced into the microcosms, the smallest fraction of the powder 336 337 remained suspended, causing the seawater to become highly turbid for several days. The resulting milky appearance of the 338 seawater eventually faded over a period of approximately five days, and by day 5, the turbidity had visually become like 339 the slag treatment and the control. This effect was not anticipated, and as a result, we decided to investigate its impact on 340 light intensity. To do so, a test was conducted after the main experiment in which olivine powder was added to a microcosm 341 identical to those used in the experiment, and light intensity was measured daily at a depth of 0.15 m. The results showed 342 that the addition of olivine caused an initial reduction in light intensity of 18.5 % at 15 mins after addition, which declined 343 to 7.4 %, 3.7 %, 3.7 % and 0 % after 1, 2, 3, and 4 days, respectively. These findings indicate that olivine additions can 344 significantly affect the light environment in the microcosms, whereas no such effect was observed in the slag treatment.

345



347 Fig. 2. Carbonate chemistry conditions. The temporal development of (a) pH<sub>T</sub>, (b) temperature, (c) total alkalinity (TA), (d) CO<sub>2</sub> fugacity

348 ( $fCO_2$ ) computed at *in situ* temperature and atmospheric pressure, (e) dissolved inorganic carbon (DIC), and (f) aragonite saturation state 349 ( $\Omega_{aragonite}$ ). The dots represent the raw data (n=3 for each treatment per sampling time), and the fitted curve is the generalized additive 350 model (GAM). The shading represents the 95 % confidence interval of the fitted GAM.

351

352 The pH<sub>T</sub> of all microcosms increased from day 1 to day 5 (Fig. 2a). This was due to photosynthetic CO<sub>2</sub> drawdown in the 353 control or photosynthetic CO<sub>2</sub> drawdown in combination with alkalinity release from minerals in the treatments. During 354 the peak of the bloom,  $pH_T$  was  $8.037 \pm 0.010$  in the control -(average values  $\pm$  standard error),  $8.054 \pm 0.014$  in the olivine 355 treatment and  $8.411 \pm 0.015$  in the slag treatment. The pH<sub>T</sub> was significantly higher in the slag than the olivine treatment 356 and the control throughout the experiment (control and olivine pH<sub>T</sub> were not significantly different). The final pH<sub>T</sub> on day 357 23 of the control, olivine, and slag treatments were  $7.893 \pm 0.012$ ,  $7.978 \pm 0.015$ , and  $8.309 \pm 0.019$ , respectively. The 358 temperature inside of the microcosms varied between replicates, which may have added noise in the biological response 359 data. However, on average there was no statistically significant difference between control/treatments during the 360 experiment.

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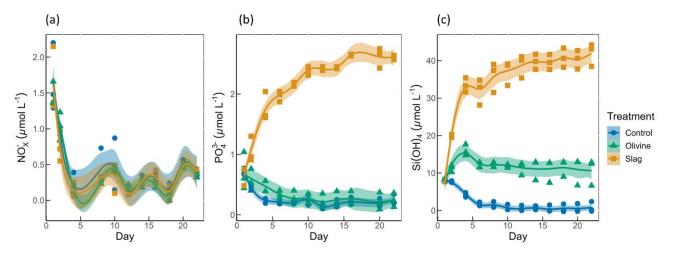
362 In our data analysis, all the fitted GAMs from the treatments and the control exhibited significant differences in  $pH_T$  from 363 each other, as evidenced by the p-values of both P-means and P-smooths being smaller than 0.001. For detailed results of 364 the GAM p-values, please refer to Table S2.

365

Total alkalinity increased marginally from  $2255 \pm 2$  to  $2262 \pm 13 \mu mol kg^{-1}$  within the first 6 days after olivine addition while it increased more substantially from  $2259 \pm 1$  to  $2522 \pm 11 \mu mol kg^{-1}$  in the same time span in the slag treatment (Fig. 2c). The TA in the control decreased from  $2261 \pm 2 \mu mol kg^{-1}$  to  $2240 \pm 7 \mu mol kg^{-1}$  from day 1 to day 6 but remained stable thereafter. The TA reached  $2279 \pm 6 \mu mol kg^{-1}$  in the olivine treatment and  $2611 \pm 9 \mu mol kg^{-1}$  in the slag treatment on day 22. The slag treatment reached a significantly higher TA than the olivine treatment and the control (P-smooths < 0.001). The mean TA from GAM in olivine treatment was higher than the control (P-means < 0.001).

372

The CO<sub>2</sub> fugacity (*f*CO<sub>2</sub>) computed at *in situ* temperature and atmospheric pressure decreased continuously in the first 6 days in all microcosms (Fig. 2d). Then it increased again in the control and olivine treatments while staying lower in the slag treatment (P-means and P-smooths  $\leq 0.001$  between either treatment or the control). Dissolved inorganic carbon (Fig. 2e) and the aragonite saturation state ( $\Omega_{aragonite}$ ; Fig. 2f) revealed a similar trend over the course of the experiment in the control and the olivine treatment. In contrast, the slag treatment had higher DIC and  $\Omega_{aragonite}$  values throughout the experiment (P-means < 0.001).



380 Fig. 3. Macronutrients concentrations over the course of the study. (a) Nitrate and nitrite concentrations. (b) Phosphate concentrations. 381 (c) Silicic acid concentrations. The dots represent the raw data (n=3 for each treatment per collection), and the fitted curve is the 382 generalized additive model.

384 Initial nitrate and nitrite (NO<sub>x</sub><sup>-</sup>), phosphate (PO<sub>4</sub><sup>3-</sup>), and silicic acid (Si(OH)<sub>4</sub>) concentrations were  $1.58 \pm 0.12, 0.69 \pm 0.59,$ and 8.04  $\pm$  0.10 µmol L<sup>-1</sup>, respectively (Fig. 3). NO<sub>x</sub>-declined rapidly in all microcosms once the experiment had 385 commenced to values below 0.5 µmol L<sup>-1</sup> and no significant difference was detected between treatments and control (P-386 smooths >0.05; Fig. 3a). In both the olivine treatment and the control, the PO<sub>4</sub><sup>3-</sup> concentration decreased in the first six 387 388 days (Fig. 3b). In the slag treatment,  $PO_4^{3-}$  increased to a maximum of  $2.65 \pm 0.01 \mu mol L^{-1}$ , which was significantly higher 389 than in the olivine treatment and the control (P-means <0.001). The Si(OH)<sub>4</sub> concentration increased to a maximum of 390  $15.99 \pm 0.87 \mu$ mol L<sup>-1</sup> in the olivine treatment, increased to a maximum of  $41.92 \pm 1.75 \mu$ mol L<sup>-1</sup> in the slag treatment, but 391 decreased below the detection limit in the control (Fig. 3c). Significant differences were observed in the development of 392 Si(OH)<sub>4</sub> between all treatments and the control (Table S2).



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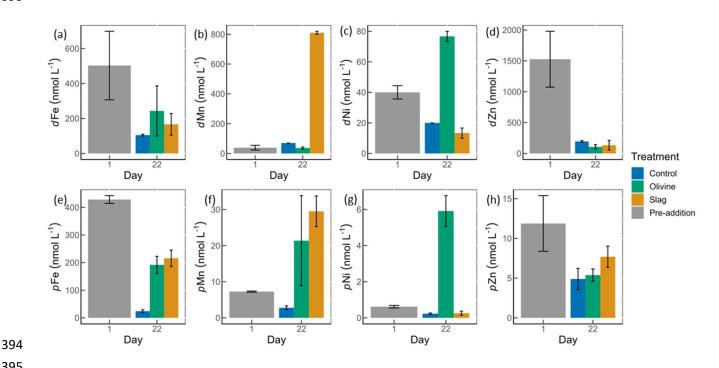


Fig. 4. Dissolved and particulate trace metal concentrations in microcosm seawater. (a)-(d) are dissolved trace metal concentrations, and (e)-(h) are total particulate trace metal concentrations. The error bars represent the standard error from measured samples. The preaddition data shown in (a)-(d) represent the average of 7 microcosms before addition of slag or olivine. The data for the control on day 22 in (a)-(d) and for the pre-addition on day 1 in (e)-(h) were based on two of three microcosm replicates. The remaining data were based on all three microcosm replicates.

401

402 The dissolved trace metal concentrations measured from microcosms are presented in Fig. S3. While the mass of olivine 403 added to the microcosms was 50-fold greater than in steel slag (100 g vs 2 g), it's noteworthy that the variation in dissolved 404 trace metal concentrations between the two treatments were much smaller than 50 folds. After 21 days of experiment, the 405 treatments showed an increase in dissolved Al concentrations from  $920 \pm 286$  to  $970 \pm 228$  nmol L<sup>-1</sup> in olivine treatment, 406 and from  $920 \pm 286$  to  $1093 \pm 77$  nmol L<sup>-1</sup> in slag treatment, while in the control dissolved Al decreased to  $230 \pm 10$  nmol 407  $L^{-1}$  (Fig. S3). The fitted GLMs were compared, and the p-value revealed how much influence a treatment had on the 408 dissolved metal concentrations (Table S3). The results indicate that the slag and olivine additions led to significantly higher 409 Al concentrations than in the control (p-values < 0.05), but no significant difference was found between the two treatments 410 (p-value = 0.189). The Cu concentration in the olivine on day 22 was significantly higher than the slag treatment and the 411 control (p-value <0.05) (Fig. S3). The addition of olivine and slag released some dissolved Fe, but overall, the concentration of Fe did not differ significantly between treatments (Fig. 4a, Table S3). The slag released a substantial amount of dissolved 412 413 Mn (maximum  $810 \pm 10$  nmol L<sup>-1</sup> on day 22) (Fig. 4b), leading to significantly higher concentrations than in the olivine 414 treatment and the control (p-values < 0.001). A significant amount of dissolved Ni (maximum 77 ± 3 nmol L<sup>-1</sup> on day 22) 415 was released from the olivine powder (p-values <0.001) (Fig. 4c). The initial concentration of dissolved Zn in seawater 416 was much higher than on day 22 in all microcosms, and no significant difference in Zn concentrations was found between 417 the treatments and the control.

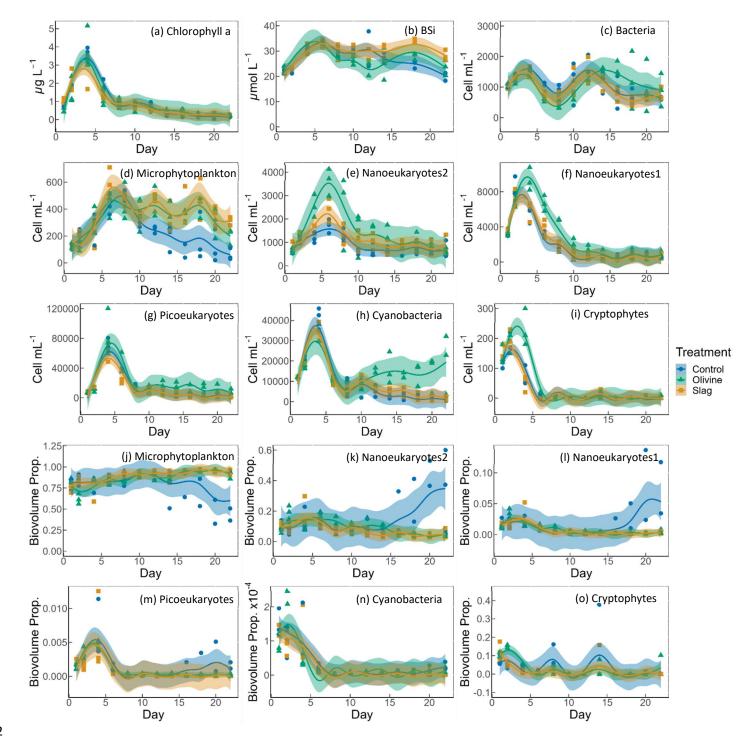
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Particulate concentrations of some trace metals also differed between treatments. The total particulate Fe decreased in all microcosms on day 22 comparing with the pre-addition level, but both mineral addition treatments had higher particulate Fe concentrations than the control (Fig. 4e). The addition of slag elevated particulate Mn concentrations to a level higher than the pre-addition and the control on day 22 (Fig. 4f), while the addition of olivine increased the particulate Ni concentrations to a level higher than the slag, the control, and the pre-addition (Fig. 4g). The particulate Zn concentrations in general decreased by the end of the experiment (Fig. 4h), and no significant differences were found between the treatments and the control.

426

427 The POC on day 1 and day 22 from all microcosms were very similar,  $10.99 \pm 0.58$  and  $11.03 \pm 0.41 \mu$ mol L<sup>-1</sup> respectively 428 (Fig. S4) so the metal:POC results were consistent with the particulate trace metal results (Fig. 4 e-h). In general, the non-429 surface metal:POC are positively correlated with the total metal:POC ratios (Fig. S5). The ratio of non-surface to total 430 particulate trace metal concentrations is summarized in Table S5. Both non-surface and total Fe concentrations decreased 431 in microcosms on day 22 compared with the pre-addition level. Iron:POC ratios were significantly higher in the treatments 432 than in the control on day 22 (p-values <0.05. Table S3), and there was no significant difference between mineral addition 433 treatments. The non-surface to total Fe:POC ratios were > 0.94 in all microcosms on both day 1 and day 22. The total and 434 non-surface Mn:POC ratio was the highest in the slag treatment. These ratios were higher than the pre-addition level and 435 the control at the end of the experiment. The total particulate Ni concentrations in the olivine treatment were significantly 436 higher than before olivine addition. The olivine treatment led to a >22-fold higher Ni:POC ratio compared to the other two

- treatments (p-value <0.001).
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- 439
- 440



# 441 3.3 Development and physiology of the plankton community



443 Fig. 5. Temporal development of chlorophyll a concentration (chl-a), BSi, and different eukaryotic and bacterial plankton groups as
444 determined with flow cytometry. (a) chlorophyll a; (b) BSi; cell concentrations of (c) heterotrophic bacteria, (d) microphytoplankton, (e)

nanoeukaryotes2, (f) nanoeukaryotes1 (g) picoeukaryotes, (h) cyanobacteria, and (i) cryptophytes; biovolume proportion of (j)
microphytoplankton, (k) nanoeukaryotes2, (l) nanoeukaryotes1 (m) picoeukaryotes, (n) cyanobacteria, and (o) cryptophytes. The figure
data points represent the raw data, and the fitted curve is the generalized additive model. The shaded area represents the 95 % confidence
interval.

449

The chl-a concentration in all microcosms increased from day 1 to day 4 from 1  $\mu$ g L<sup>-1</sup> to 3-4  $\mu$ g L<sup>-1</sup> (Fig. 5a). The chl-a concentration then decreased rapidly from day 4 to day 8, then continued to decrease, though more slowly, to <0.3  $\mu$ g L<sup>-1</sup> until the end of the experiment. The GAMs of chl-a did not show any difference between treatments and the control (both P-means and P-smooths >0.05, see Table S2).

454

455 The BSi concentration increased from day 1 to day 6 in all microcosms (Fig. 5b). In the olivine treatments, BSi 456 concentrations decreased slightly after the peak until day 12 but then increased again. In the slag treatment, BSi 457 concentrations remained relatively stable after the initial phytoplankton bloom. In contrast, BSi concentration decreased continuously in the control after the initial peak. Olivine particles suspended in seawater after the mineral addition (see 458 459 section 3.2) partially ended up on BSi filters during filtration. This led to extremely high BSi measurements on days 2 and 460 4 that were removed from Fig. 5b. Without these outliers, the mean of fitted BSi GAM in the olivine treatment was lower 461 than the control and the slag treatment (Table S2), and the slag treatment had the highest average BSi over the course of 462 the experiment. Overall, the BSi trends in the two treatments were similar (P-smooths = 0.269), and both were significantly 463 different from the control (P-smooths < 0.05).

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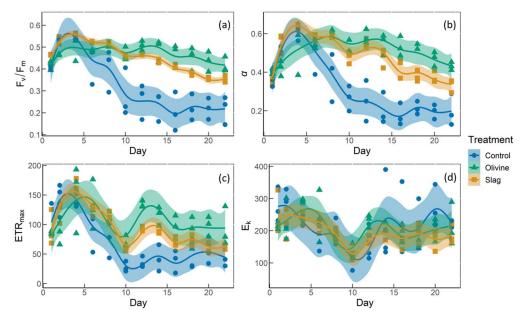
465 The development of the phytoplankton community composition showed significant differences between the treatments and 466 the control. In general, most phytoplankton groups exhibited similar patterns to chl-a, with peak cell numbers occurring on 467 day 4 (Fig. 5f-i) apart from microphytoplankton and nanoeukaryotes2 which had the peak delayed for 1-2 days (Fig. 5d-468 e). Please be aware that flow cytometers may not capture some large and chain-forming phytoplankton. After reaching peak values during the bloom, phytoplankton abundance generally decreased steadily. Microphytoplankton displayed 469 470 similar trends to the results for BSi. Before day 10, all microcosms had similar microphytoplankton abundances (Fig. 5d). 471 However, in the control, microphytoplankton abundance declined continuously and at a faster rate compared to the two treatments (P-smooths values <0.03). From day 2 to day 6, the abundance of nanoeukaryotes1, nanoeukaryotes2, 472 473 picoeukaryotes, and cryophytes was higher in the olivine treatment compared to the slag treatment and the control. After 474 day 8, their abundance in the olivine treatment decreased to a similar level as the slag treatment and the control. Notably, 475 there were few significant differences observed between the slag treatment and the control in terms of the abundances of 476 nanoeukaryotes1, nanoeukaryotes2, picoeukaryotes, cyanobacteria, and cryptophytes throughout the experiment. In the 477 olivine treatment, cyanobacteria experienced a second bloom after day 10, which was significantly different from the other 478 two groups (P-smooths <0.01). Heterotrophic bacteria exhibited an increase and decline pattern following the 479 phytoplankton bloom until day 8 (Fig. 5c). Subsequently, bacteria abundance increased again, reaching a second peak 480 during days 12-14, followed by a decline until the end of the experiment. The decline in bacteria abundance was slower in 481 the olivine treatment, although no significant differences were detected between treatments (Table S2).

482

Among all the microcosms, microphytoplankton consistently accounted for the largest proportion of biovolume. From the
 perspective of biovolume proportion, the mineral addition mainly influenced the microphytoplankton and nanoeukaryotes.
 The control had similar phytoplankton biovolume distribution as the treatments from day 1 to day 15, but after that the

486 proportion of microphytoplankton biovolume decreased to a level significantly lower than the treatments. In the control 487 treatment, the proportion of nanoeukaryotes' biovolume increased as the proportion of microphytoplankton decreased. The 488 biovolume of picoeukaryotes, cyanobacteria and cryptophytes increased during the phytoplankton bloom and then 489 decreased drastically after the bloom. There were no significant differences in biovolume proportion observed for 490 picoeukaryotes, cyanobacteria and cryptophytes between the treatments and the control.

491



492

**493** Fig. 6. The photosynthetic performance of the phytoplankton community. (a)  $F_v/F_m$ , the maximum quantum yield of photosynthesis II. **494** (b)  $\alpha$ , the initial slope of the rapid light curves. (c) ETR<sub>max</sub> is the maximum electron transport rate, the maximum potential photosynthetic **495** rate. (d)  $E_k$  is light-saturation parameter, Unit:  $\mu$ mol photons m<sup>-2</sup> s<sup>-1</sup>.

496

497 The temporal development of  $F_v/F_m$ ,  $\alpha$ , ETR<sub>max</sub>, and  $E_k$  is illustrated in Fig. 6. The  $F_v/F_m$  values of the phytoplankton community were approximately  $0.42 \pm 0.01$  and increased to levels > 0.5 during the peak of the phytoplankton bloom on 498 499 day 4 (Fig. 6a). Following the bloom,  $F_v/F_m$  values dropped below 0.3 in the control. However, the decline in  $F_v/F_m$  after 500 the bloom was less pronounced in the two mineral addition treatments with the olivine treatment maintaining higher  $F_v/F_m$ 501 values than the slag treatment (P-smooths <0.05). At the end of the experiment,  $F_v/F_m$  was  $0.22 \pm 0.04$  in the control, 0.35 502  $\pm$  0.01 in the slag treatment, and 0.42  $\pm$  0.02 in the olivine treatment. The temporal development of  $\alpha$  aligned with the 503 patterns observed for  $F_v/F_m$  (compare Fig. 6a and 6b). The maximum values of  $ETR_{max}$  were observed on day 4 in the 504 control and the slag treatment, while in the olivine treatment, it occurred on day 5 (Fig. 6c). Subsequently, ETR<sub>max</sub> 505 continuously decreased until day 10 and then stabilized until the end of the experiment. However, ETR<sub>max</sub> exhibited a 506 subsequent increase in the mineral treatments around day 12. The ETR<sub>max</sub> values were higher in the mineral treatments 507 compared to the control group (P-means <0.001, Table S2). The parameter  $E_k$  decreased from 246 ± 17 µmol photons m<sup>-2</sup> s<sup>-1</sup> on day 1 to  $121 \pm 7 \mu$ mol photons m<sup>-2</sup> s<sup>-1</sup> on day 10, and then it increased again to approximately 200  $\mu$ mol photons m<sup>-</sup> 508 509  $^{2}$  s<sup>-1</sup> by the end of the experiment (Fig. 6d). The change in E<sub>k</sub> did not exhibit significant differences between the treatments 510 and the control (both P-means and P-smooths >0.05).

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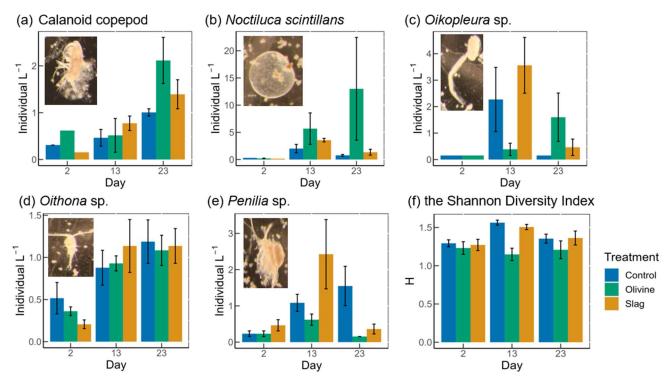


Fig. 7. The dominant zooplankton abundance and community diversity from different treatments. Abundance of dominant zooplankton
in microcosms: (a) calanoid copepod; (b) *Noctiluca scintillans*; (c) *Oikopleura* sp.; (d) *Oithona* sp.; (e) *Penilia* sp.; and (f) the Shannon
diversity index (H) of different treatments and the control. Error bars represent the standard error calculated from three microcosm
replicates. Photographs of each zooplankton group are shown on the corresponding graphs.

519 Thirteen zooplankton taxonomic groups were identified in the microcosms. The dominant taxa were the appendicularian 520 Oikopleura sp., the cyclopoid copepod Oithona sp., the cladoceran Penilia sp., the heterotrophic dinoflagellate Noctiluca 521 scintillans and several calanoid copepods including Acartia sp., Paracalanus sp. and Gladioferens sp. The larvae and eggs 522 of Oikopleura, Penilia and copepod were also observed under the microscope. In general, higher zooplankton numbers were observed after the bloom on day 13 (Fig. 7). The abundance of calanoid copepods and Oithona sp. increased after 523 524 day 2 (Fig. 7a, d), and there was no significant difference between treatments and the control (p-values >0.05, Table S4). 525 The abundance of N. scintillans increased significantly more in the olivine treatment than in the control and the slag 526 treatment, with highest abundance of  $13 \pm 9$  individual L<sup>-1</sup> observed in the olivine treatment on the last day (Fig. 7b). The 527 abundance of Oikopleura in the control and the slag treatment was higher than the olivine treatment on day 13 but was 528 higher in the olivine treatment on day 22 (Fig. 7c). A higher abundance of Penilia sp. was found in the slag treatment on 529 day 13 and in the control on day 23 (Fig. 7e). Due to the patchy distribution of zooplankton, these data have large standard 530 errors and only the differences in the numbers of N. scintillans in the olivine treatment were statistically significantly 531 different from the slag treatment and the control (p-value <0.05, Table S4).

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Considering the control and slag treatment, the Shannon Diversity Index (H) increased from day 2 to day 13 and declined on day 23, while in the olivine treatment, H was lower on day 13 than on day 2 and day 23 (Fig. 7f). The GLMs revealed that the olivine treatment had significantly lower H on day 13 than the control and the slag treatment (p-values <0.001). There were no significant differences in H between the control and the slag treatment (Table S4). The addition of olivine decreased the zooplankton community's diversity. This is mainly driven by distinct trends observed in the abundance of *Oikopleura* sp., *Penilia* sp., and *N. scintillans* (Fig. 7).

## 540 4. Discussion

#### 541 4.1 CO<sub>2</sub> removal potential of slag and olivine

The slag powder created significantly higher CO<sub>2</sub> removal potential than the olivine powder over the course of the study. 542 543 Ca(OH)<sub>2</sub> and CaO in slag and Mg<sub>2</sub>SiO<sub>4</sub> in olivine are likely to be the main functional minerals driving the measured alkalinity enhancement. Total alkalinity increased by 361 µmol kg<sup>-1</sup> in the slag treatment while it increased by only 29 544 µmol kg<sup>-1</sup> in the olivine treatment, equivalent to a potential increase in marine inorganic carbon by 14.7 and 0.9% within 545 3 weeks of their application. When normalizing these alkalinity increases to the same material weight, 1 g of slag would 546 547 release 9626 µmol TA while 1 g of olivine would release 16 µmol TA. Thus, over 3 weeks of experimental incubation, slag 548 is ~600-fold more efficient in releasing alkalinity for particles of this size class (please note that particle size spectra of 549 olivine and slag were similar but not identical; Fig. S1). We can also use these values to make a rough estimate of how 550 much CO<sub>2</sub> these two minerals could potentially sequester. One mole of alkalinity from olivine and slag can sequester 551 approximately 0.85 mole of CO2. Thus, one tonne of slag and olivine powder as used here could sequester 360 and 0.6 kg, 552 respectively, within 3 weeks. It is likely that optimization of particle size and application method may lead to higher 553 efficiencies. Nevertheless, the slag showed potential as an OAE source mineral, even when applied as relatively coarse 554 powder in this experiment.

555

## 556 4.2 Environmental implications of slag and olivine additions

557 The amount of olivine and slag powder added to the treatments differed significantly (100 g of olivine powder were added 558 while only 2 g of slag powder were added to the 53 L microcosms). Our rationale for these different mass additions was to 559 yield somewhat similar amounts of detectable alkalinity enhancement in the dissolved phase, since we already knew from 560 tests before the experiment that slag elevates alkalinity faster than olivine. However, olivine was less efficient in releasing alkalinity than we had anticipated so that even a 50-fold higher addition of olivine (in mass) did not compensate for this 561 562 difference. As such, our experiments are associated with an "apples and oranges issue" in that our perturbation with 563 minerals and associated OAE differs. We noteargue that an adjusted addition of minerals depending on the alkalinity 564 enhancement rate would be consistent with what OAE practitioners may do under real-world conditions. Presumably, OAE 565 deployments may have to adjust the amounts of minerals to detect alkalinity enhancement in the dissolved phase for 566 verification purposes. Nevertheless, to account for the "apples and oranges issue", the following discussion mainly relates 567 the observed environmental effects with the alkalinity enhancement achieved over the course of the study.

# 568 4.2.1. OAE effects on phytoplankton physiology and community

Previous research has hypothesised that OAE-induced changes in seawater carbonate chemistry could delay phytoplankton bloom formation due to reductions in seawater  $pCO_2$  in the aftermath of an OAE deployment (Bach et al., 2019). The buildup of chlorophyll *a* concentration as observed here was indistinguishable between treatments and the control, suggesting no effect of slag- or olivine-based OAE on phytoplankton bloom dynamics under these experimental settings. A lack of bloom delay due to carbonate chemistry is unsurprising for the olivine treatment where the release of alkalinity was small 574 (29 µmol kg<sup>-1</sup> alkalinity release), but somewhat more surprising in the slag treatment where alkalinity was quite rapidly 575 increased by 361 µmol kg<sup>-1</sup>. However, the release was still lower than in a very similar study by Ferderer et al., (2022) 576 where alkalinity was increased by 500 µmol kg<sup>-1</sup> using sodium hydroxide and even there they did not observe a bloom 577 delay. Based on this very limited evidence, it seems that bloom delays do not occur consistently under OAE within the 578 alkalinity ranges tested in this studywithin this alkalinity range.

580 The nutrient data show that the phytoplankton community was most likely N-limited after day 4 so that the release of 581 Si(OH)<sub>4</sub> from olivine and Si(OH)<sub>4</sub> and PO<sub>4</sub><sup>3-</sup> from slag did not stimulate a further increase in chlorophyll-a concentration in the treatments. The development of BSi concentrations is indicative of the prevalence of diatoms in the microcosms but 582 583 differences between treatments and the control were small. The release of Si(OH)<sub>4</sub> through olivine and slag will most likely benefit diatoms but this fertilization effect did not manifest in this specific experiment because N was limiting diatom 584 585 growth. However, when new N is supplied then diatoms will likely take a bigger share of the limiting N pool when olivine 586 or slag are used for OAE, as has been shown in Si(OH)4 manipulation experiments in and outside the context of OAE 587 research (Egge and Jacobsen, 1997; Ferderer et al., 2023). As such, diatoms are likely to benefit from olivine and slag 588 applications. In the case of slag, the release of  $PO_4^{3-}$  will likely be another driver that affects plankton productivity and 589 community composition. As for Si(OH)<sub>4</sub>, however, the effect of additional  $PO_4^{3-}$  did likely not materialise in this 590 experiment because  $PO_4^{3-}$  was not limiting over the course of the study. However, in ecosystems where  $PO_4^{3-}$  is a limiting 591 resource, the application of slag could enhance productivity with associated benefits for higher trophic levels. In contrast, 592 excessive applications of slag and concomitant PO43- release could also pose a risk of eutrophication. Future studies may need to investigate what the most sustainable dose of OAE via olivine and/or slag applications could be and the suitable 593 594 regions for application.

The flow cytometry results further revealed the change in phytoplankton community composition. Both the olivine and slag treatments sustained higher microphytoplankton abundances after the peak of the phytoplankton bloom. This trend is consistent with higher  $F_v/F_m$  values in the treatments than in the control so that it is tempting to assume that photophysiological fitness gain measured with the FRRf led to higher competitiveness of microphytoplankton in the community. Indeed, calculations of the contribution of different phytoplankton groups to total biovolume based on flow cytometry indicate that microphytoplankton were predominantly contributing to the phytoplankton community biovolume so that the responses measured by the FRRf were probably to a large extent driven by this group.

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Apart from the increased microphytoplankton abundance, for the slag treatment, other phytoplankton groups distinguished with flow cytometry did not deviate considerably from the control. The olivine addition, however, triggered more pronounced shifts in the phytoplankton community. In particular, the nanoeukaryotes (roughly between 2-20 μm), picoeukaryotes and the cryptophytes showed relatively higher abundance during the peak of the phytoplankton bloom, and the abundance of cyanobacteria was higher after the bloom. We speculate that this shift following olivine treatment may be attributable to a top-down effect from the decrease in zooplankton grazing effects in microcosms, which will be discussed in section 4.2.2.

611

612 The measurement of photophysiological parameters revealed that the phytoplankton had generally better photosynthetic613 performance in the slag and olivine treatments than in the control, especially after the phytoplankton bloom. During the

614 first 5 days, the changes in phytoplankton photosynthetic performance were indistinguishable between the control and the 615 slag treatment, while the values of  $\alpha$ , ETR<sub>max</sub> and F<sub>v</sub>/F<sub>m</sub> were lower in olivine treatment. At this time all microcosms had 616 similar health because of the relatively high  $NO_x^-$  concentrations and Fe supply (around 500 nmol L<sup>-1</sup>), but the suspended 617 particles in the olivine treatment may have led to artifacts in the measuring of photophysiology by FRRf. Scattering and/or 618 absorption of light by suspended olivine particles is the most parsimonious explanation for the simultaneous depression in 619  $\alpha$ , ETR<sub>max</sub> and F<sub>v</sub>/F<sub>m</sub>. After day 5, the F<sub>v</sub>/F<sub>m</sub>,  $\alpha$  and ETR<sub>max</sub> values decreased significantly faster in the control than in the 620 treatments, and to values lower than the initial condition. A decrease of  $F_v/F_m$  is commonly associated with physiological 621 stress, such as nutrient limitation, and high light stress (Bhagooli, et al., 2021), with Fe limitation causing a more 622 pronounced decline in  $F_v/F_m$  than nitrogen limitation (Gorbunov, et al., 2021). The ETR<sub>max</sub>, which represents the maximum 623 electron transport rate, has also been shown to be negatively affected when phytoplankton experience nitrogen or Fe 624 limitation (Kolber et al., 1994; Gorbunov & Falkowski 2021). Furthermore, the change in photosynthesis performance 625 after day 10 was suspected to be driven by the microphytoplankton because the decrease of F<sub>v</sub>/F<sub>m</sub>, α, and ETR<sub>max</sub> in the 626 control was coupled with the decrease in microphytoplankton abundance while the other phytoplankton groups were in low abundance as in the mineral addition treatments, and the microphytoplankton contributed significantly (75 %) to 627 628 community biovolume. All microcosms were similarly  $NO_x^-$  limited from day 5 onward (Fig. 3) so that N-limitation is 629 unlikely to explain different trends in photophysiological parameters between the control and OAE treatments. Trace metals, 630 especially Fe, released through slag and olivine additions could potentially explain these differences.

632 Several of the trace metals released from slag and olivine are required for photosynthesis. For example, Fe is required for 633 many proteins functioning in photosynthesis, such as cytochromes, ferredoxin, and superoxide dismutase (SOD) (Twining 634 and Baines, 2013), and the addition of Fe can stimulate the growth of phytoplankton (Sunda and Huntsman, 1997) and 635 increase  $F_v/F_m$  (Behrenfeld et al., 2006). The dissolved and particulate Fe concentrations were higher in mineral addition 636 treatments than in the control indicating potentially more Fe available to sustain phytoplankton photosynthesis. While this 637 explanation is intriguing for the observed trends in photophysiology, it remains unclear why such strong differences 638 occurred between mineral addition and control treatments despite dissolved Fe concentrations of ~500 nmol L<sup>-1</sup> at the end 639 of the experiment in the control. In Fe-limited ocean regions, dissolved Fe is at least two orders of magnitude lower, and the enhancement of Fe to  $\sim 1.5$  nmol L<sup>-1</sup> can induce major phytoplankton blooms and relieve photophysiological stress (De 640 641 Baar et al., 2005). It is possible that these coastal phytoplankton species have higher Fe requirements than those from the open ocean where Fe is limiting (Strzepek and Harrison, 2004). We speculate that when Fe was consumed during the 642 643 phytoplankton bloom, bioavailable Fe was much lower in the control, and may have been insufficient to meet the cellular 644 requirements of coastal phytoplankton. Our findings therefore suggest that Fe perturbations mayis not only be relevant for 645 lower Fe open ocean regions but could also be relevant for coastal ocean locations.

646

631

Alternatively, the addition of Mn, Ni and other trace metals from mineral addition may have benefited photosynthesis. Manganese is required for the water-splitting reaction of photosystem II (Armstrong, 2008), and both Mn and Ni are common bioactive trace metals for SODs in marine phytoplankton. The noxious superoxide anion radical  $(O_2^{-})$  generated from aerobic respiration and oxygenic photosynthesis could be harmful to phytoplankton physiology, and SOD removes  $O_2^{-}$ , thus improving photosynthesis (Wafar et al., 1995; Wolfe-Simon et al., 2005). This is consistent with our photosynthetic measurements. Interestingly, although the amounts and types of trace metals released from the slag and

olivine powders were different, they led to relatively similar  $F_v/F_m$  values with only slightly higher  $F_v/F_m$  in the olivine

654 than the slag treatment from days 10-21. Over this time, these trace metal additions could have fertilized different 655 phytoplankton species (Pausch et al., 2019; Balaguer et al., 2022; Guo et al., 2022) possibly because different 656 phytoplankton could have different trace metal requirements, such as for SOD. For example, cyanobacteria have NiSOD, 657 diatoms have MnSOD, dinoflagellates have both FeSOD and MnSOD (Wolfe-Simon et al., 2005). Another explanation is 658 that phytoplankton in the control were limited by bicarbonate while the treatments had sufficient bicarbonate from added 659 minerals. However, we were unable to determine the species-level changes in the phytoplankton community, and hence 660 whether these trace metals, individually or combined, could account for the observed phytoplankton community 661 photosynthetic performance.

662

# 663 4.2.2. OAE impacts on the zooplankton community

664 Slag-based OAE did not significantly influence the zooplankton community composition while olivine-based OAE induced 665 some statistically significant effects, including a lower Shannon diversity. The increase in *N. scintillans* abundance and the 666 decrease in *Penilia* sp. and *Oikopleura* sp. in the olivine treatment indicate that the zooplankton response to OAE can vary 667 among different zooplankton types.

668

669 The observed lower abundance of Oikopleura sp. on day 13 in the olivine treatment may indicate a temporary suppression 670 or a slower growth rate of this zooplankton species in response to the olivine addition. This could be attributed to the 671 potential effects of olivine on the availability of essential nutrients or changes in the physicochemical environment of the 672 water. However, the subsequent increase in *Oikopleura* sp. abundance by day 22 suggests that the growth of this species 673 may have recovered or accelerated in the olivine treatment, leading to a higher abundance compared to the slag treatment 674 and the control on day 22. As discussed in section 4.2.1, reduced Oikopleura sp. abundance was unlikely due to reduced 675 food availability since phytoplankton within the preferred edible size spectrum, such as cyanobacteria and nanoeukaryotes, 676 were even more abundant in the olivine treatment. Instead, we hypothesize it to be an effect of the suspended olivine 677 particles that occurred for approximately the first 5 days of the study that were so plentiful that they turned the enclosed 678 seawater milky and may have clogged the mucous feeding mesh of Oikopleura sp. (Lombard et al., 2011).

679

680 The abundance of *Penilia* sp. and *Oikopleura* sp. was lower in the olivine treatment than the other two groups throughout 681 the experiment while the abundance of N. scintillans was consistently higher. The second bloom of cyanobacteria in olivine 682 is likely potentially to be the results of decreased predators, like Penilia sp. and Oikopleura sp., although the changes in 683 their abundance were not statistically significant between treatments and the control. We cannot provide a particularly 684 convincing hypothesis about what specifically drove these differences in these zooplankton species, al-though it is tempting 685 to speculate that suspended particles present in the olivine treatment at the beginning may have played a role also for those 686 organisms since this was the only apparent systematic difference to the control and slag treatment. The proliferation of N. 687 scintillans can be problematic since heterotrophic dinoflagellate blooms can regulate phytoplankton communities, cause 688 toxicity to aquatic fish, and create a hypoxic sub-surface zone (Baliarsingh et al., 2016; Zhang et al., 2020; Al-Azri et al., 2007), although a bloom of N. scintillans in southeast Australia only induced ichthyotoxicity when the cell concentration 689 reached 2,000,000 cells L<sup>-1</sup> (Hallegraeff et al., 2019). For comparison, we observed a maximum of 32 cells L<sup>-1</sup> in one 690 691 microcosm replicate of the olivine treatment.

693 In comparison to olivine, steel slag seemed to have less potential to affect zooplankton community composition. The 694 abundance of all groups of phytoplankton, apart from microphytoplankton after day 10, was similar in the slag treatment 695 and the control through the experiment. This is probably because the amount of slag powder added in the treatment was 696 much less than the olivine powder resulting in fewer physical particle perturbations to zooplankton. In addition, the 697 chemistry perturbations such as enhanced alkalinity concentration and various dissolved trace metals, especially Mn, from 698 the slag powder did not seem to have a notable direct influence on zooplankton abundance over the three-week period. 699 Even though we did not observe drastic changes in zooplankton abundance during the experiment, considering there was 700 higher microphytoplankton abundance in the slag treatment after day 10, slag powder may benefit some zooplankton 701 especially those who feed on large phytoplankton on a longer time scale.

702

# 703 4.2.3. Dissolved trace metal accumulation in seawater and its environmental implications

704 The addition of olivine and slag as OAE source minerals released trace metals into the seawater, predominantly Al, Fe, Ni, 705 and Cu (olivine) as well as Al, Fe, and Mn (slag). The maximum measured concentrations for dissolved Al, Fe, Ni, Cu, and 706 Mn were 1093, 253, 77, 27, and 810 nmol L<sup>-1</sup>, respectively. The threshold values for drinking water with health or aesthetic 707 considerations by the Australian Drinking Water Guidelines for Al, Fe, Ni, Cu, and Mn are 7400, 5360, 340, 15600, and 708 1800 nmol L<sup>-1</sup>, respectively (NRMMC, 2022). All dissolved trace metal concentrations measured herein are well below 709 these health and aesthetic threshold values. In natural freshwater sources, the concentrations of Al, Fe, Ni, Cu and Mn are 710 generally less than 44000, 71400, 510, 156, and 25400 nmol L<sup>-1</sup> (NRMMC, 2022). Although these natural water data were 711 primarily derived from rivers and streams, they serve as valuable references for evaluating trace metal release in our 712 experiment. Thus, mineral additions to the microcosms as simulated here did not increase thresholds for any of the 713 measured trace metals beyond those that are considered safe for drinking water quality, and they were within the trace 714 metal concentration range in natural water. However, while these guidelines on drinking water provide a good starting point 715 on how to quantify what OAE perturbation could be considered "safe" and "unsafe" with regards to trace metals, it must 716 be recognized that seawater is not drinking water and that critical thresholds may be different in the latter.

717

718 The release of trace metals from OAE materials is considered to have relatively strong effects on biology, particularly in 719 the open ocean where trace metals usually occur in lower concentrations. For example, oceanic Al, Fe, Ni, and Mn 720 concentrations are about 2, 0.5, 8, and 0.3 nmol L<sup>-1</sup> (Bruland and Lohan, 2003; Sohrin and Bruland, 2011). Previous 721 research on OAE-associated trace metal impacts on individual phytoplankton species grown in laboratory environments 722 has shown that concentration thresholds beyond which trace metal induces negative effects on fitness likely differ between 723 species (Guo et al., 2022; Hutchins et al., 2023; Xin et al., 2023). Indeed, our experiment with plankton communities 724 provides further support that several components of the planktonic food web are affected by OAE. However, our experiment 725 does not allow determining whether observed effects were primarily invoked by carbonate chemistry, macronutrient (P and 726 Si), or trace metal perturbations. Thus, dedicated experiments isolating the impact of these factors on plankton will be 727 required in the future.

## 4.2.4. Particulate trace metal accumulation in seawater and its environmental implications

The Derwent Estuary (where we collected our plankton communities) was highly metal polluted due to industrial practice
 (Macleod and Coughanowr, 2019). Both our dissolved and particulate trace metal data indicated high background metal

concentrations, especially for Fe and Zn. Furthermore, the metal:POC ratios found here are higher than reported for open ocean studies or lab cultures. For example, the Fe:POC can vary from 2-136 µmol mol<sup>-1</sup> depending on the cultured phytoplankton species and the environmental dissolved Fe concentration (Kulkarni et al., 2006; Sunda and Huntsman, 1995; King et al., 2012; Boyd et al., 2015). In our results the Fe:POC values ranged from 1200 to 39 000 µmol mol<sup>-1</sup>, which may be due to the particulate trace metal richness of the Derwent Estuary (control) and/or the addition of lithogenic particles (slag and olivine treatment). The presence of abiotic particulate metal sources creates challenges to quantify metal quotas and then to evaluate metal accumulation effects on biological organisms.

738

Our study reveals that the added minerals enriched the particulate trace metal pools to various degrees. Consistent with the dissolved trace metal data, the slag treatment was enriched with particulate Fe and Mn while the olivine treatment was enriched with particulate Fe and Ni. The enhanced particulate Ni and Mn concentrations were higher than before mineral additions and the control levels. This is in line with previous research which indicates a positive correlation between particulate and dissolved trace metal concentrations (Gaulier et al., 2019).

744

Based on the amounts released through OAE as simulated herein, it appears that Ni and Mn have the highest potential to cause toxicity in certain marine organisms (Jakimska et al., 2011). These trace metals have the potential to accumulate in marine organisms over time (bioaccumulation effects), and their increased concentrations in the food chain can lead to adverse effects on the health and well-being of organisms at higher trophic levels (biomagnification effects). One crucial next step will be to investigate whether the enhanced dissolved/particulate trace metal will affect higher trophic levels to estimate the environmental risks of OAE on other marine organisms.

751

## 752 5 Conclusions

753 Our study aimed to assess the environmental impacts of two ground OAE minerals, olivine and steel slag, on coastal 754 plankton communities. Both minerals released alkalinity, leading to an elevation in pH<sub>T</sub>. However, the addition of steel 755 slag exhibited significantly higher efficiency in elevating alkalinity compared to olivine.

756

Approximately 1.9 g L<sup>-1</sup> of olivine powder were added in the olivine treatments, leading to a 29  $\mu$ mol kg<sup>-1</sup> increase in alkalinity and increased concentrations of Si(OH)<sub>4</sub> and trace metals (Fe and Ni). Compared to this relatively modest increase of alkalinity and associated CO<sub>2</sub> removal potential, the impacts on the plankton community appeared to be relatively pronounced. Thus, although our experiment ran for only 3 weeks, and olivine powder may slowly release more alkalinity, the short-term response monitored here suggests that the immediate climatic benefit is relatively small compared to a relatively pronounced environmental effect.

763

Only 0.038 g L<sup>-1</sup> of slag were added to the treatment but this led to an alkalinity enhancement of 361 µmol kg<sup>-1</sup> and the increased concentrations of macronutrients (P and Si) and trace metals (Mn and Fe) additions as well as changes in carbonate chemistry. Although limited environmental impacts were observed from the slag treatment in our experiment, some aspects require further study. For example, the pronounced release of P could cause eutrophication and the relatively rapid increase in pH may be a detrimental aspect if organisms cannot acclimate fast enough. Furthermore, it is essential to

- consider that the composition of steel slag can vary depending on the source factory (Wang et al., 2011; Proctor et al.,
  2000), which may affect the efficiency of carbon removal and change the trace metal perturbation. Nevertheless, just based
  on our experiment, the comparison between the immediate climatic benefit and environmental effect appears to be more
  favourable for slag than olivine.
- 773

Based on our findings, it can be concluded that steel slag powder exhibited fewer environmental impacts on plankton communities compared to olivine powder relative to its capacity for alkalinity enhancement. The results highlight the importance of carefully assessing the environmental consequences of using specific OAE minerals, particularly when considering their potential effects on plankton communities.

778

779 Data availability. Data are available in the Institute for Marine and Antarctic Studies (IMAS) data catalogue, University
780 of Tasmania (UTAS) (<u>https://doi.org/10.25959/X6FH-9K15</u>, Guo, J., & Bach, L. (2023).).

781

Author contributions. LTB, RFS, KMS and JAG designed the experiments and JAG carried them out. LTB, RFS and
 KMS supervised the study. ATT analysed the dissolved/particulate trace metal samples. JAG conducted statistical analyses.
 JAG prepared the manuscript with contributions from all authors.

- 786 **Competing interests.** The contact author has declared that none of the authors has any competing interests.
- 787

785

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796

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