



# The influences of El Niño–Southern Oscillation on tropospheric ozone in CMIP6 models

Thanh Le<sup>1\*</sup>, Seon-Ho Kim<sup>1</sup>, Jae-Yeong Heo<sup>1</sup> and Deg-Hyo Bae<sup>1\*</sup>

<sup>1</sup>Department of Civil and Environmental Engineering, Sejong University, Seoul 05006, Republic of Korea

5 \*Corresponding author(s): Thanh Le ([levinhthanh.lvt@gmail.com](mailto:levinhthanh.lvt@gmail.com)) and Deg-Hyo Bae ([dhbae@sejong.ac.kr](mailto:dhbae@sejong.ac.kr))

**Abstract.** Ozone in the troposphere is a greenhouse gas and a pollutant, hence, additional understanding of the drivers of tropospheric ozone evolution is essential. The El Niño–Southern Oscillation (ENSO) is a main climate mode and may contribute to the variations of tropospheric ozone. Nevertheless, there is uncertainty regarding the causal influences of ENSO on tropospheric ozone under warming environment. Here, we investigated the links between ENSO and tropospheric ozone

10 using Coupled Modeling Intercomparison Project Phase 6 (CMIP6) data over the period 1850–2014. Our results show that ENSO impacts on tropospheric ozone are primarily found over oceans, while the signature of ENSO over continents is largely nonsignificant. The response of ozone to ENSO may vary depending on specific air pressure levels in the troposphere. These responses are weak in the middle troposphere and are stronger in the upper and lower troposphere. Although there are biases in simulating the signature of ENSO on surface ozone, these signatures in the middle and upper 15 troposphere appear to be more consistent across CMIP6 models. While the response of tropical tropospheric ozone to ENSO is in agreement with previous works, our results suggest that ENSO impacts on tropospheric ozone of the mid-latitude regions over the southern Pacific, Atlantic, and Indian oceans might be more significant than previously understood.

**Keywords:** El Niño–Southern Oscillation; historical simulations; tropospheric ozone; biogeochemistry; CMIP6.

## 1 Introduction

20 Ozone in the troposphere is an important greenhouse gas and pollutant (Archibald et al., 2020; Cooper et al., 2010; Wang et al., 2022). Tropospheric ozone has detrimental effects on human health and ecosystems (Fleming et al., 2018; Franz et al., 2018; Gaudel et al., 2020; Lu et al., 2019; Oliver et al., 2018; Peron et al., 2021; Roberts et al., 2022; Schuberger et al., 2019). Changes in atmospheric ozone may also affect radiative forcing and have effects on climate (Gauss et al., 2006; Myhre et al., 2013).

25 The El Niño–Southern Oscillation (ENSO) is the major mode of climate variability with global impacts (Bjerknes, 1969; Cai et al., 2020; McPhaden et al., 2006) and is expected to affect variations of global tropospheric ozone. ENSO-induced changes in climate and meteorological conditions (Le and Bae, 2022; Lu et al., 2019; Yeh et al., 2018) lead to impacts on ecosystems and the production/removal of ozone in soil, plant, and the water cycle, and the transport of ozone (Ganzeveld et al., 2009; Lin et al., 2019; Neu et al., 2014). In particular, ENSO drives changes in tropospheric and stratospheric



30 circulations which can alter the tropospheric ozone variations (Daskalakis et al., 2022; Domeisen et al., 2019; Koumoutsaris et al., 2008; Lin et al., 2015; Neu et al., 2014; Olsen et al., 2019; Oman et al., 2013; Zeng and Pyle, 2005; Ziemke et al., 2015). In addition, ENSO was revealed to exhibit influences on tropospheric ozone concentrations in many regions by modulating local meteorological conditions (Doherty et al., 2006; Jiang and Li, 2022; Oman et al., 2011; Peiro et al., 2018; Rowlinson et al., 2019; Wie et al., 2021; Xu et al., 2017; Yang et al., 2022; Zhang et al., 2015).

35 Nevertheless, there are uncertainties regarding the causal effects of ENSO on global tropospheric ozone. For instance, while the response of tropospheric ozone to ENSO over the mid-latitude regions remains elusive (Lu et al., 2019; Olsen et al., 2016), further understanding of ENSO impacts on ozone concentrations at multiple air pressure levels in the troposphere is necessary. Despite the high spatial and temporal variability of tropospheric ozone, there are limited observations of past ozone changes at the global scale (Dragani, 2011; Ebojie et al., 2016; Gaudel et al., 2018; Young et al., 2018). Hence, Earth 40 system models remain valuable tools to understand the evolution of tropospheric ozone and the interactions between tropospheric ozone and regional climate (Archibald et al., 2020; Collins et al., 2017; Young et al., 2018). Datasets from Coupled Modeling Intercomparison Project Phase 6 (CMIP6) models provide an important source to better identify the effects of ENSO on global tropospheric ozone.

In the present study, we evaluated the causal impacts of ENSO on tropospheric ozone at the global scale using data from the 45 historical simulations of CMIP6 models. We also discussed the coherency across CMIP6 models in reproducing the connection between ENSO and tropospheric ozone.

## 2 Materials and Methods

### 2.1 Datasets

We used monthly data of mole fraction of ozone in air (i.e., variable ‘o3’) at different air pressure levels (i.e., 1000, 850, 500, and 300 hPa). The CMIP6 models with ozone data available for the historical simulations (Eyring et al., 2016) over the 1850-2014 period are listed in Table S1. The use of various model outputs reduces the uncertainty of the connections between ENSO and tropospheric ozone.

There are biases in simulating tropospheric ozone variations in the models (Griffiths et al., 2021; Turnock et al., 2020; Young et al., 2018), however, CMIP model outputs are still helpful to investigate the effects of ENSO on tropospheric ozone 55 (Archibald et al., 2020; Young et al., 2018).

We employed monthly sea level pressure (SLP) and sea surface temperature (SST) to calculate the time series of the major climate modes (see also section Text S1).

### 2.2 Methods

We assess the possibility of the impacts of ENSO on tropospheric ozone based on the approach employed in recent studies 60 (Le and Bae, 2020, 2022). This method was established using a multivariate predictive model to assess the probability for the



absence of Granger causal effects of ENSO on ozone concentrations. In the computations, we considered the confounding impacts of other major climate modes (i.e., the dipole mode index (DMI; Saji et al., 1999), the Southern Annular Mode (SAM; e.g., Cai et al., 2011), and the North Atlantic Oscillation (NAO; e.g., Hurrell et al., 2003)). Given that the climate changes in the Indian and Atlantic oceans can affect the tropical Pacific (Cai et al., 2019; Ha et al., 2017; Le et al., 2020; Le 65 and Bae, 2019), and modify the connections between ENSO and ozone concentrations, these analyses provide a realistic estimate for the response of ozone concentrations to ENSO. Additional information of the methods employed in this study is described in section Text S1.

### 3 Results

Figure 1 depicts the models' mean map of ozone concentrations at different air pressure levels for the period 1850-2014 of 70 the CMIP6 historical simulations. In the middle and lower troposphere, ozone is higher in the northern hemisphere compared to the southern hemisphere (Figure 1). The agreement between models fluctuates at different air pressure levels. For example, in the upper troposphere (i.e., at 300 hPa pressure level), high consistency across the model is found in the mid-latitude regions, while this consistency is lower in the tropics and polar regions (Figure 1a). In the middle troposphere (i.e., at 500 hPa pressure level), the models' agreement is mainly found in the northern hemisphere (Figure 1b). The simulations 75 of near-surface ozone (i.e., at 850 hPa) are consistent over the tropics and parts of the northern hemisphere (Figure 1c), while the models' agreement is low in reproducing surface ozone (i.e., at 1000 hPa) for most regions (Figure 1d). The standard deviation is normally higher in the tropics compared to other regions (Figure S1).

Figure 2 displays the causal effects of ENSO on global ozone concentrations for the historical period 1850-2014. In Figure 2, we show that the response of ozone to ENSO may vary depending on specific air pressure level. For instance, ENSO impacts 80 on ozone in the upper troposphere (i.e., at 300 hPa pressure level) can be observed over parts of the Pacific Ocean and North America (Figure 2a). In these areas, ENSO is unlikely to exhibit no causal influences on ozone concentrations (i.e.,  $p$ -values were lower than 0.33). The response of ozone to ENSO in the middle troposphere (i.e., at 500 hPa pressure level) is very limited over the tropical Pacific and Atlantic Oceans (Figure 2b). We observe more significant impacts of ENSO on ozone in the lower troposphere compared to the upper and middle troposphere. Specifically, tropical ozone concentrations at near 85 surface (i.e., at 850 hPa) and surface (i.e., at 1000 hPa) levels appear to be sensitive to ENSO (Figures 2c and d). In addition, ENSO impacts on surface ozone can be found over part of the northern North Pacific and mid-latitude regions in the southern hemisphere (Figure 2d). We observe causal influences of ENSO on ozone over Antarctica, however, these influences show low agreement between models (Figures 2c and d). While the signature of ENSO on ozone variations is generally weak over continents of the lower and middle troposphere (Figures 2b-d), this signature is, however, stronger in 90 the upper troposphere over North America (Figure 2a). These results imply that the effects of ENSO on tropospheric ozone over lands are nonsignificant for most regions.



Differences between CMIP6 models in replicating the influences of ENSO on surface (1000 hPa) ozone are shown in Figure 3. Similar results for other pressure levels (300, 500, and 850 hPa) are presented in Figures S2-S4. As revealed in Figure 3, several models (i.e., BCC\_CSM2\_MR, IPSL\_CM6A\_LR, and MPI\_ESM1\_2\_LR) may not reproduce the significant  
95 influences of ENSO on surface ozone over the tropical Pacific and Indian Oceans as described in Figure 2d. The models IPSL\_CM6A\_LR and MPI\_ESM1\_2\_LR may underestimate the response of surface ozone to ENSO over the mid-latitude regions in the southern hemisphere compared to other models. The agreement between models for the impacts of ENSO on surface ozone is low over continents (Figure 2d), partly due to the discrepancy in simulating ozone variability (Figure 1d). While there are biases in simulating the response of surface ozone to ENSO (Figure 2d), these responses in the middle and  
100 upper troposphere appear to be more consistent across models (Figures 2a-c, S2-4).

Springtime surface ozone is more sensitive to ENSO compared to other seasons (Figure 4). In particular, the clear response of springtime surface ozone over the high-latitude north Pacific can be observed (Figure 4a). The impacts of ENSO on surface ozone of other seasons are limited (e.g., over the tropical Pacific and part of southern North America, Figure 4b-d). The results for other air pressure levels (300, 500, and 850 hPa) are shown in Figures S5-S7. We note that the response of  
105 springtime ozone at higher pressure levels is weaker compared to springtime surface ozone for most regions, except for the upper troposphere over east Asia and northwestern North America (Figure S5a). Consistent with the results illustrated in Figure 2b, the impacts of ENSO on seasonal ozone in the middle troposphere (500 hPa) are largely nonsignificant for all seasons (Figure S6).

#### 4 Discussion and conclusions

110 The effects of ENSO on tropospheric ozone over the tropical Pacific (Figures 2-4) show agreement with previous works (Chandra et al., 2007; Peiro et al., 2018). ENSO causes changes in the tropospheric ozone budget over the tropical Pacific by modulating the Walker circulation (Chandra et al., 2007), wind systems (Cai et al., 2021; Le and Bae, 2020; Yeh et al., 2018), and inducing biomass burning (Chandra et al., 2009; Le et al., 2022). Further, significant ENSO impacts on tropical ocean regions described in Figures 2-4 are in agreement with recent works (Olsen et al., 2016; Wespes et al., 2017) using  
115 satellite data.

The significant impacts of ENSO on ozone in the upper troposphere (300 hPa) over the southern and western North America (Figures 2a and S5a) might be associated with the transport of ozone from east Asia (Cooper et al., 2010; Doherty, 2015; Lin et al., 2015). However, these impacts cannot reach the surface levels (Figures 2c-d, S6a, and S7a), consistent with recent work (Lin et al., 2015).

120 We note that the models without the Atmospheric Chemistry module (BCC\_CSM2\_MR, IPSL\_CM6A\_LR, MPI\_ESM1\_2\_LR; See Table S1) provide different results of ENSO impacts compared to the rest models (Figures 3). Hence, improvement of the Atmospheric Chemistry module in the models may provide further understanding of the connection between ENSO and ozone variations.



The weak impacts of ENSO on the mid-level tropospheric ozone (i.e., 500 hPa level, described in Figures 2a and 3) might be  
125 due to the strong exchange between stratospheric ozone and middle to upper tropospheric ozone (Liu et al., 2017; Meul et al., 2018; Neu et al., 2014; Williams et al., 2019).

Several models showed ENSO effects on tropospheric ozone over Antarctica with a low agreement between models (Figures 2-4). These impacts might be associated with the signature of ENSO on stratospheric ozone anomalies over Antarctica (Li et al., 2021; Lin and Qian, 2019).

130 Given that the tropospheric ozone burden and the ozone-induced impacts may increase in some regions in the future (Doherty et al., 2013; Franz and Zaeble, 2021; Gaudel et al., 2020; Griffiths et al., 2021; Verstraeten et al., 2015; Zanis et al., 2022), further analyses of ENSO impacts on tropospheric ozone in future climate projections are necessary.

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## Data Availability

140 The data that support the findings of this study are openly available at the following website: <https://esgf-node.llnl.gov/search/cmip6/>.

## Author contribution

TL designed the study, performed the data analysis, and wrote the manuscript. SHK, JYH and DHB contributed to the interpretation of results and the writing of the manuscript.

## 145 Competing interests

The authors declare that they have no conflict of interest.



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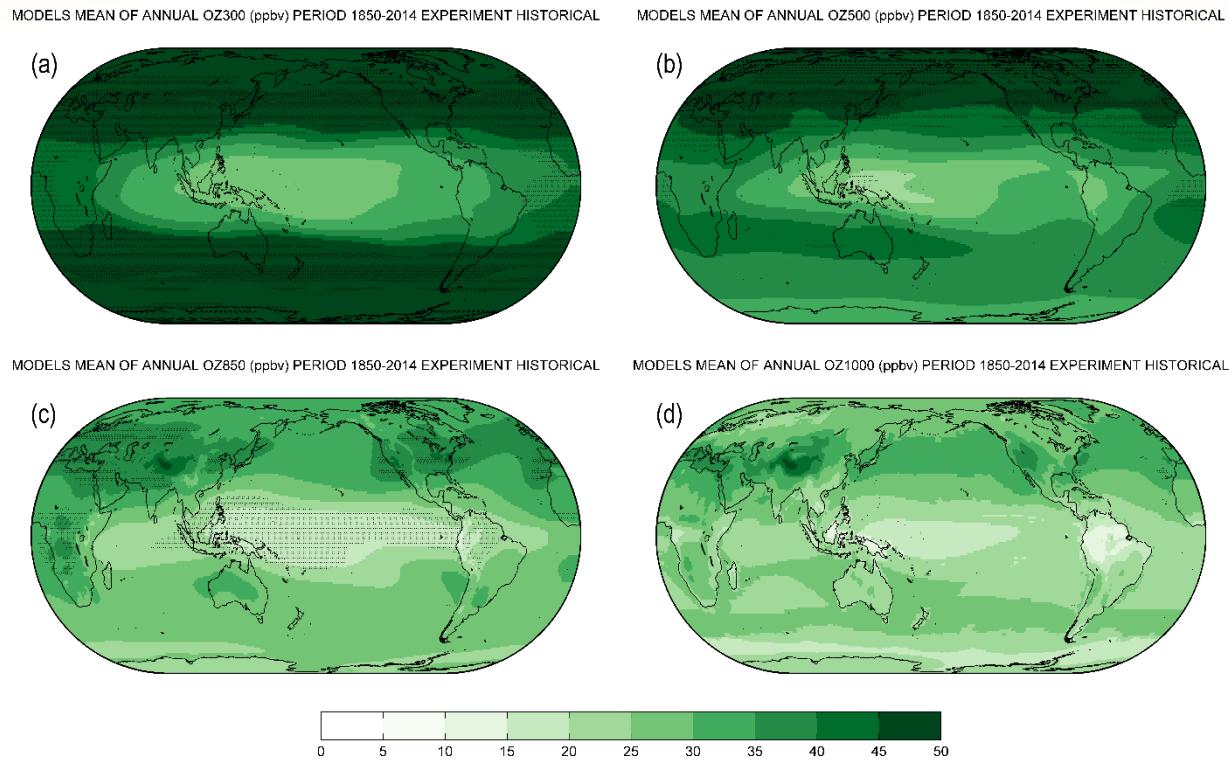


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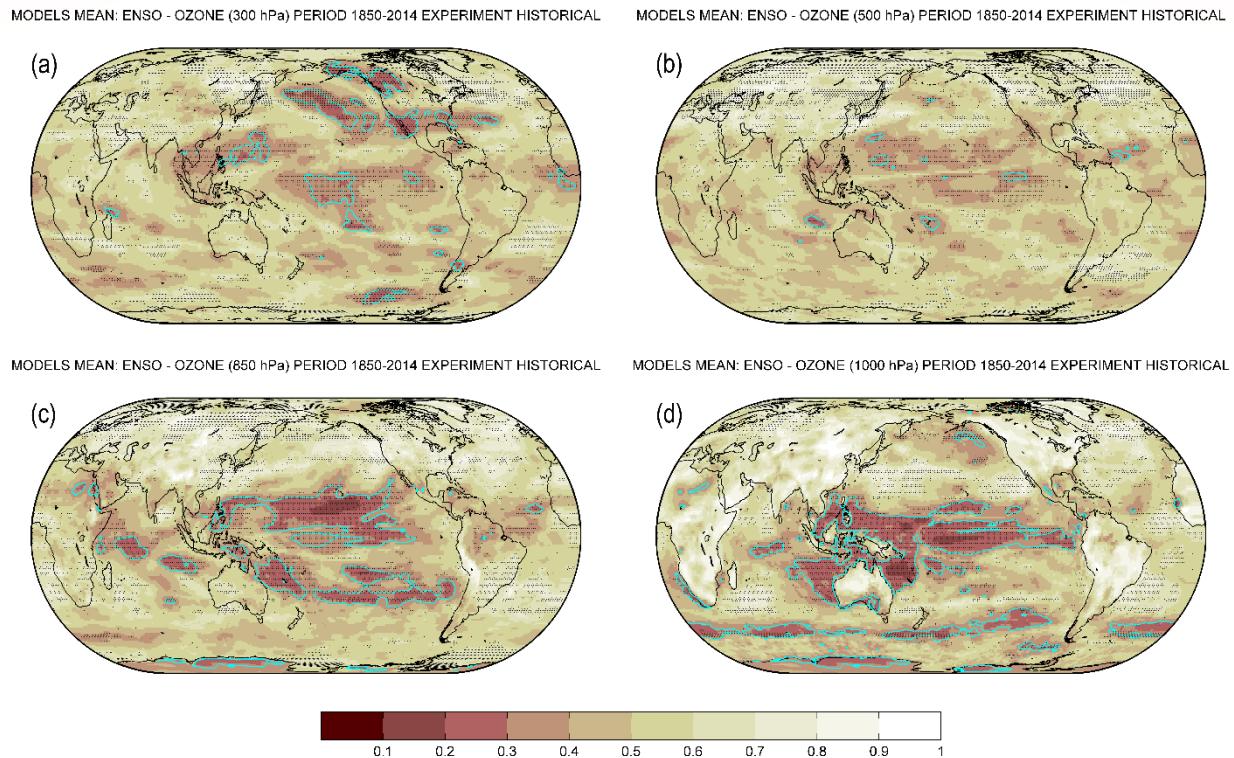
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**Figure 1.** Multi-model mean map of annual mean ozone concentrations (ppbv) for the historical experiment over the 1850-2014 period at 300 hPa (a), 500 hPa (b), 850 hPa (c) and 1000 hPa (d) pressure levels, respectively. Stippling indicates that at least 70% of total models show agreement on the mean ozone concentrations of all models at given grid point. The agreement of an individual model is identified when the difference between the selected model's ozone concentrations and the multi-model mean ozone concentrations is less than one standard deviation of the multi-model mean ozone concentrations.

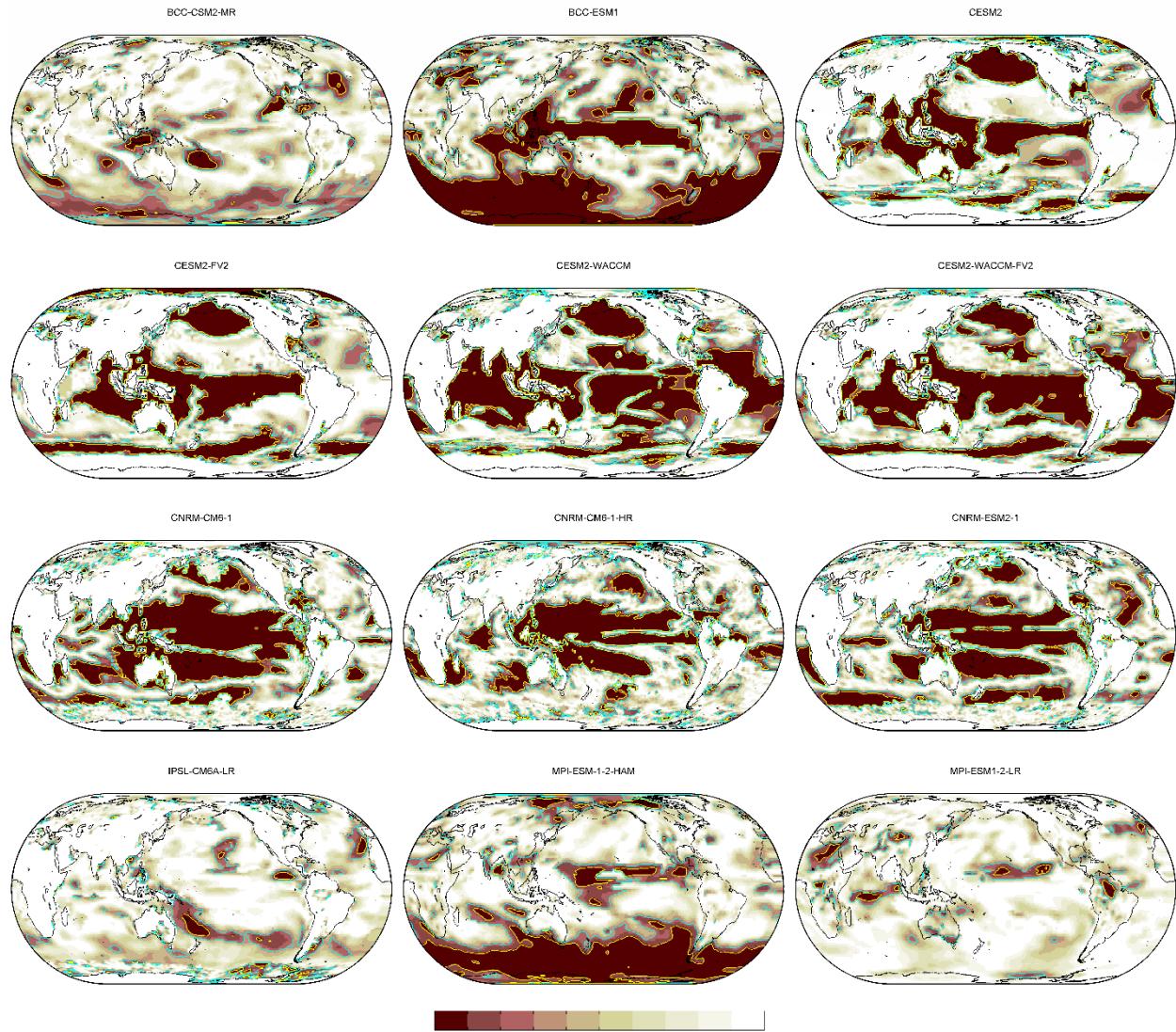
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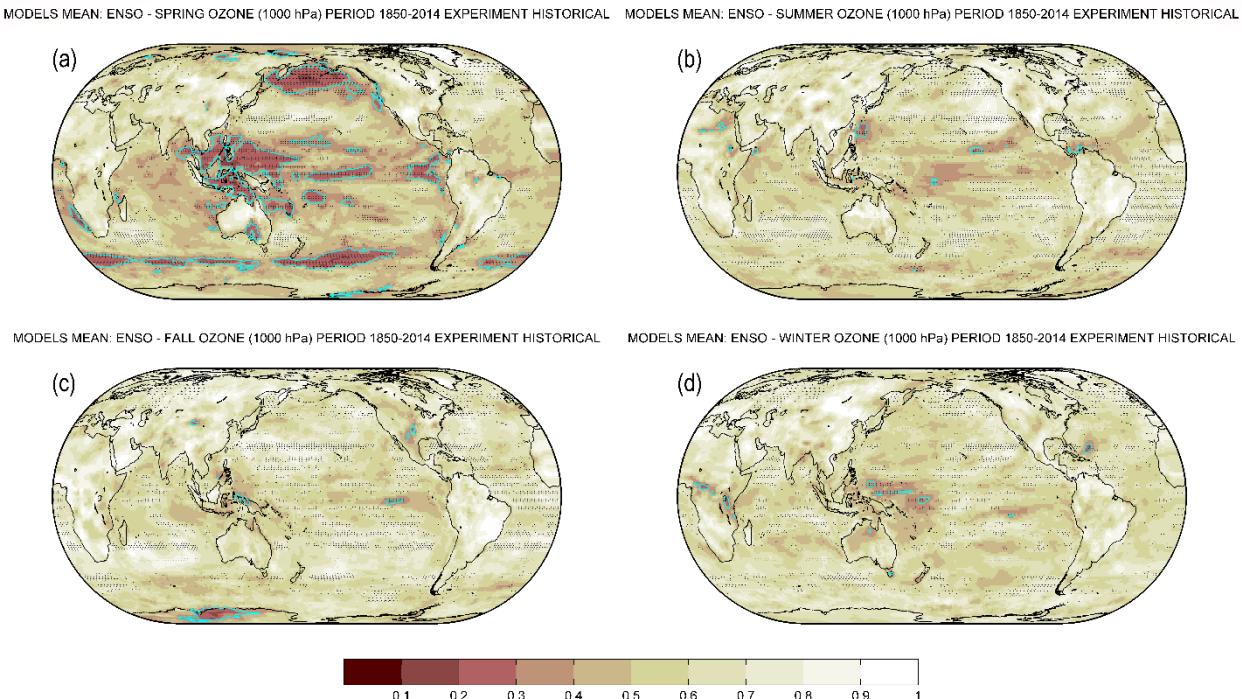
**Figure 2.** Map of multi-model mean probability for the absence of Granger causality from ENSO to annual ozone concentrations for the historical experiment over the 1850–2014 period at 300 hPa (a), 500 hPa (b), 850 hPa (c) and 1000 hPa (d) pressure levels, respectively. Stippling indicates that at least 70% of total models show agreement on the mean probability of all models at given grid point. The agreement of an individual model is identified when the difference between the selected model's probability and the multi-model mean probability is less than one standard deviation of the multi-model mean probability. The cyan contour line denotes  $p$ -value = 0.33. Brown shades indicate low probability of no Granger causality. ENSO = El Niño–Southern Oscillation.

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**Figure 3.** Probability of no Granger causality from ENSO to annual ozone concentrations at 1000 hPa pressure level for the historical experiment over the 1850–2014 period of 12 individual models (see Table S1). The yellow and cyan contour lines denote  $p$ -value = 0.1 and 0.33, respectively. Brown shades imply a low probability of no Granger causality. ENSO: El Niño–Southern Oscillation.



385

**Figure 4.** Multi-model mean probability map of no Granger causality from ENSO in boreal winter [ $D(t)JF(t+1)$ ;  $t$  indicates year  $t$ ] to seasonal mean ozone concentrations at 1000 hPa pressure level over the period 1850-2014. (a) Spring [MAM( $t+1$ )]. (b) Summer [JJA( $t+1$ )]. (c) Fall [SON( $t+1$ )]. (d) Winter [D( $t+1$ )JF( $t+2$ )]. Stippling signifies that at least 70% of total models show agreement on the mean probability of all models at a given grid point. The cyan contour line signifies  $p$ -value = 0.33. Brown shades imply a low probability of no Granger causality. ENSO: El Niño–Southern Oscillation. MAM: March- April-May. JJA: June-July-August. SON: September-October-November. DJF: December-January-February.

390