

1 **Climatic Characteristics of the Jianghuai Cyclone and Its Linkage**
2 **with Precipitation During the Meiyu Period from 1961 to 2020**

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8 **Abstract.** This study examines the climatic characteristics of 202 Jianghuai cyclones
9 and their linkage with precipitation during the Meiyu period from 1961 to 2020. The
10 results show that cyclones mainly originate from eastern and western Hubei Province.
11 Additionally, we explore the statistical characteristics of intensity, radius, and their
12 positive correlation. In studying the decadal variation of cyclones, we find a similar
13 evolution between the cyclones and Meiyu precipitation. Therefore, we further
14 investigate the correlation between the Jianghuai cyclones and the precipitation during
15 the Meiyu period. There is a positive correlation coefficient of 0.77 between them.
16 Notably, the percentage of precipitation affected by cyclone activities can reach up to
17 47%. The anomalous increase in precipitation caused by cyclones north of 27°N can
18 reach a maximum of 7 mm/day. When a cyclone exists, a significant negative
19 geopotential height anomaly at the 500 hPa level over Mongolia can be traced back to
20 day -4. The abnormally enhanced WPSH (Western Pacific subtropical High),
21 southwesterly low-level jet and negative geopotential height are the dominant factors
22 causing abnormal precipitation during Jianghuai cyclones. Before and after the cyclone
23 develops, water vapor flux and divergence from low latitudes increase abnormally,
24 providing sufficient water vapor conditions for the generation of cyclone precipitation.

25 **1. Introduction**

26 Meiyu is a special rainy season resulting from the progression of the East Asian
27 summer monsoon. The East Asian monsoon originates in the South China Sea in mid-
28 May and then advanced northward, forming rain bands in South China, the Jianghuai
29 region, the Korean Peninsula and Japan (Ding et al., 2004,2007; Qian et al., 2000). This
30 special rainy season is called Meiyu in China, called Changma in South Korea and Baiu
31 in Japan (Ninomiya et al., 1987; Oh et al., 1997; Saito. 1995). The Meiyu front is one
32 of the important weather systems affecting summer precipitation in the middle and
33 lower reaches of the Yangtze River (Pang et al., 2013; Tao et al., 1979; Wang et al.,
34 2014; Zhou et al., 2022). From mid-June to early July, the east of Yichang in Hubei
35 Province experiences continuous rain and short periods of sunshine. These conditions
36 are accompanied by heavy rainfall, strong wind and other weather phenomena during
37 the Meiyu period (Ding. 1992; Zhao et al., 2021; Zhou et al., 2017). In China, the mean
38 annual precipitation during the Meiyu period in the Jianghuai River Basin can reach
39 300 mm, accounting for 30%-40% of the mean annual total precipitation, and even up
40 to 500 mm or more in extreme Meiyu period (Liu et al., 2020). Historically, most
41 summer flood disasters are caused by precipitation anomalies during the Meiyu period.
42 Some scholars have studied and analyzed the representative floods of 1996, 1998, 2016
43 and 2020 (Bao et al., 2021; Su et al., 2021; Zhao et al., 2018; Zhong et al., 2023). These
44 floods, caused by the Meiyu front, had adverse effects on people's safety, lives and
45 property (Yan et al., 2021).

46 Scholars in China have categorized rainstorms caused by Meiyu fronts into three
47 types (Zhang et al., 2004). The first type is the β mesoscale convective rainstorm on the
48 Meiyu front. This type of rainstorm has a range of less than 300 km, with strong
49 intensity and a rapid formation process (He et al., 2007). It is difficult to forecast more
50 than 12 hours and can only be detected using radar for proximity forecasting (Zhang et
51 al., 2002). The second type is the persistent rainstorm located in front of the upper
52 trough of low pressure in the western part of the Meiyu front. It is characterized by a
53 long duration of approximately 5 days but appears less frequently, mainly in western

54 Hubei and western Hunan and Sichuan (Cai et al., 2021; Wu et al., 2020a;). The last
55 type is the rainstorm caused by the Jianghuai cyclone located east of the origin of the
56 Meiyu. The Jianghuai cyclones are affected by the thermal conditions of the sea and
57 land and likely occur in the eastern part of the Meiyu front (Wang et al., 2016). The
58 positive vorticity advection in front of the high-altitude trough and the warm advection
59 in front of the front promote the eastward movement and development of the cyclone
60 (Shen et al., 2019; Zhang et al., 2016). During the development of the cyclone, the lower
61 levels are dominated by the southwest warm and humid airflow, and the upper levels
62 are mainly affected by dry and cold air (Zhao et al., 2008). This type of rainstorm has a
63 large range, high intensity and long duration of precipitation (Wang et al., 2012; Xu et
64 al., 2011).

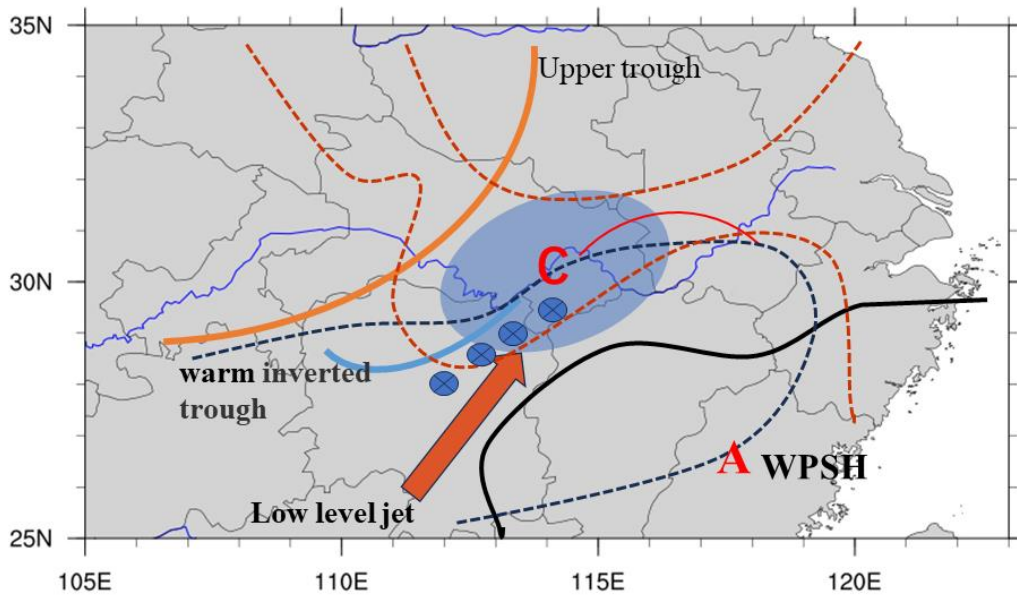
65 Scholars' studies on Jianghuai cyclones during the Meiyu period were initially
66 based on individual case analysis. Xu et al. (2013) studied a cyclone process in 2011
67 and found that the cyclone process lasted up to 36 h. The cyclone rainstorm is
68 distributed on the south side of the cyclone the heavy precipitation mainly occurred in
69 the lower reaches of the Yangtze River. Wu et al. (2020b) studied two different cyclone
70 rainstorm processes and found that rainfall is directly proportional to cyclone intensity.
71 There is a strong convergence center of water vapor flux during cyclone development.
72 Zhou et al. (2020) found that a tornado is generated from the cyclone occlusion stage
73 on July 22. The tornado is under the influence of a strong and fast Jianghuai cyclone
74 and produce heavy precipitation accompanied by thunderstorms.

75 With the improvement of cyclone identification methods and reanalysis data
76 reconstruction, statistical studies of cyclones have been further developed (Simmonds
77 et al., 1999; 2000; Wernli et al., 2006). Yang et al. (2010) modeled the rainstorm process
78 in the lower reaches of the Yangtze River from 1998 to 2005. They found that cyclones
79 account for 62.5% of the rainstorm cases, and more than 70% of the cyclones could
80 develop and produce rainstorms. The Jianghuai cyclone, locate in the lower reaches of
81 the Yangtze River generally exist in the lower troposphere at 700 hPa. The horizontal
82 scale is within 400 km, and the life period on land is generally less than 48 hours. Wang

83 et al. (2015) found that the number of cyclones is lower and their intensity weaker in
84 the 1980s and 1990s. In the early 2000s, cyclones become more frequent, and their
85 intensity increased. After 2010, there is a decreasing trend again. Zhang et al. (2018)
86 divided 60 cases of extreme precipitation in the middle reaches of the Yangtze River
87 from 2008 to 2015 into five types. Among them, extreme precipitation of the Jianghuai
88 cyclone type account for 30%. The stable and maintained Western Pacific subtropical
89 high (WPSH) system is one of the important reasons for the strong precipitation
90 produced by cyclones. Because of the weak cold air force, the intensity of the Jianghuai
91 cyclone is weaker than that in spring (Zhou et al., 2017). The daily analysis of the
92 Jianghuai cyclones during the Meiyu period is easily overlooked. All these studies
93 indicate that the Jianghuai cyclone is an important weather system causing heavy
94 rainfall during the Meiyu period in the middle and lower reaches of the Yangtze River
95 (Wu et al., 2021; Zhang et al., 2018; Zhu et al., 1998).

96 Research on the climatic characteristics and precipitation effects of Jianghuai
97 cyclones during the Meiyu period over the past 60 years has not yielded clear results.
98 In this study, the relative vorticity method is used to objectively identify and track
99 cyclones based on reanalysis data provided by ERA5. We study the climatological
100 characteristics of Jianghuai cyclones during this period are studied, and analyze the
101 correlation between Jianghuai cyclone activity and precipitation. This study provides a
102 reference for the long-term and short-term forecasting of precipitation during the Meiyu
103 period.

104 The remainder of this paper is organized as follows: Section 2 presents the dataset
105 and analytical methods. In Section 3, we show the climatology composite of the cyclone
106 tracks, genesis locations, intensity, lifetime, and other characteristics. We also study the
107 relationship between cyclonic activity frequency and precipitation during the Meiyu
108 period using geopotential height and water vapor flux anomalies. Section 4 offers the
109 main discussion and findings of this study.



110

111 Fig.1 Schematic diagram of the main weather system and the structure of temperature
 112 and pressure field in the middle and low levels of the Jianghuai cyclone. (Red dotted
 113 line: isotherm; Solid black line: contour line; Blue dot: precipitation area; Solid orange
 114 line: 500 hPa upper-level trough; Red arrow: low level jet; Black dotted line: warm
 115 inverted trough; Solid red line: warm shear; Solid blue line: cold shear; Letter C:
 116 cyclone; Letter A: WSPH.)

117 2. Data and methods

118 2.1 Data

119 The time span of all the data is 60 years from 1961 to 2020, and the study area is
 120 located at 108°E to 123°E longitude and 27°N to 34°N latitude. We use the ERA5
 121 relative vorticity hourly data (850 hPa) provided by the European Centre for Medium
 122 Range Weather Forecasts (ECMWF) for Jianghuai cyclone identification and tracking.
 123 This data has a spatial resolution of $0.25^\circ \times 0.25^\circ$ and a temporal resolution of every 6
 124 hours, with each step defined as a 6 hours period. The geopotential height, wind field,
 125 and specific humidity are daily values processed from ERA5 hourly data with a spatial
 126 resolution of $0.25^\circ \times 0.25^\circ$ (Hersbach et al., 2018). The geopotential height and wind
 127 field data include pressure levels approximately at 500 hPa and 850 hPa. The specific

128 humidity data include pressure levels approximately at 850 hPa. The precipitation data
129 are from the CN05.1 grid point observation dataset compiled by the National
130 Meteorological Information Center of China Meteorological Administration with a
131 resolution of $0.25^{\circ} \times 0.25^{\circ}$ (Wu et al., 2013; Xu et al., 2009).

132 The Meiyu intensity index is used to characterize the strength of Meiyu, and the
133 data is from the National Climate Center of China. The calculation area of Meiyu
134 intensity index is defined in the government standard GB/T 33671-2017. The Meiyu
135 intensity index is defined as:

$$136 \quad M = \frac{L}{L_0} + \frac{0.5(R/L)}{R_0/L_0} + \frac{R}{R_0} - 2.5$$

137 M is the Meiyu intensity index. L is the length of the Meiyu in a given year (unit:
138 day) and L_0 means the average length of the Meiyu over the years (units: day). R is the
139 total precipitation of Jianghuai River basin during Meiyu in a given year, and R_0 is the
140 average total precipitation of Jianghuai River basin during Meiyu over the years. The
141 average period is from 1961 to the current year. For example, L_0 and R_0 values for 2000
142 are the averages from 1961 to 2000. Where M between -0.375 and 0.375, China
143 Meteorological Administration defines this year as the normal. Where M between 0.375
144 and 1.25, this year is defined as a little strong. Where M greater than or equal to 1.25,
145 this year is defined as strong. Where M between -1.25 and -0.375, this year is defined
146 as a little weak. Where M less than or equal to -1.25, this year is defined as weak.

147 **2.2 Methods**

148 Scholars propose various methods for identifying extratropical cyclones. The
149 vorticity tracking method suggested by Hodges (1994, 1995) is used in this paper for
150 objective identification and tracking of cyclones. This method primarily uses the
151 relative vorticity field at the 850 hPa to determine feature points of the cyclone, which
152 are then matched within a given time span to correspond with the position and track of
153 the cyclone. In addition to Hodges' relative vorticity method, other scholars have also
154 proposed different methods for identifying cyclones. Lu (2017) improved the
155 extratropical cyclone identification and tracking method involving the nine-point

156 pressure minimum. Jiang et al. (2020) proposed an algorithm for identifying
157 extratropical cyclones on the basis of gridded data. This algorithm is named the eight-
158 section slope detection method.

159 Among them, the most commonly used cyclone tracking methods are the mean
160 sea level pressure method (SLP) and the 850 hPa relative vorticity method. Mailier et.al
161 (2006) and Zhang et.al (2012) studied the tracks of individual cyclones using these two
162 methods. Both studies found that the 850 hPa relative vorticity method can identify and
163 detect cyclone center earlier than the SLP method (Mailier et al., 2006). The reason for
164 this result is that SLP is easily influenced by topography and large-scale background
165 circulation shear vorticity (Hodges, 1994). Therefore, based on this advantage of the
166 relative vorticity method, we select the 850 hPa relative vorticity tracking method. This
167 tracking method can detect low vortex systems earlier and track cyclones for a longer
168 period of time with better stability. Even When closed pressure levels are not visible on
169 satellite maps, the vorticity tracking method can still continue to track cyclone,
170 improving the accuracy of cyclone track data.

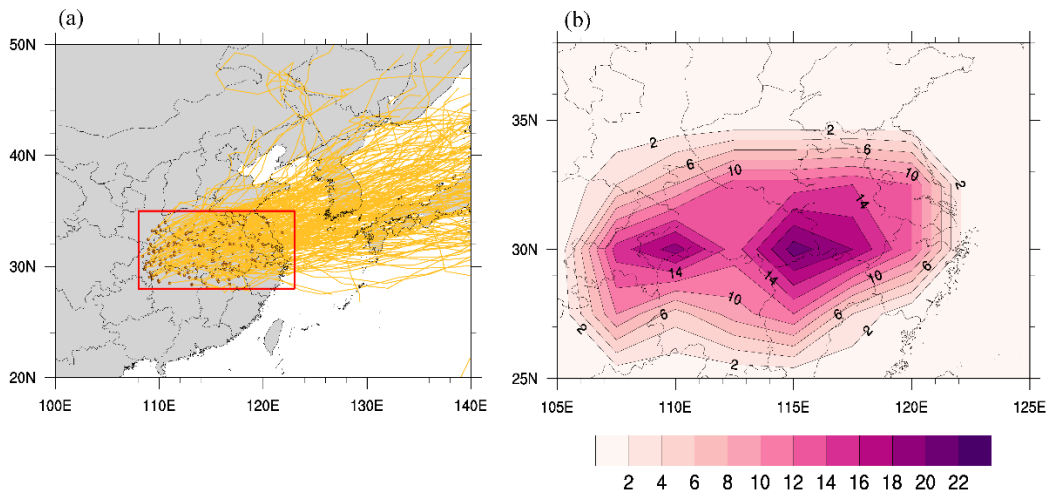
171 **3. Results**

172 **3.1 Climatic characteristics of the Jianghuai cyclone during the Meiyu period**

173 A total of 202 Jianghuai cyclones exist during the Meiyu period from 1961 to 2020.
174 The range of cyclone genesis locations defined by the Jiangsu Meteorological
175 Administration (2017) and the characteristics of the relative vorticity tracking method
176 are used. We adjust the genesis location and remove the cyclones that are generated at
177 sea and have no effect on land precipitation (108°E-123°E, 28°N-35°N). To explore the
178 tracks and genesis location of Jianghuai cyclone, we conduct a statistical analysis
179 (Fig.2). The brown dots represent the genesis locations, the first place meeting the
180 criterion, of the Jianghuai cyclone. The yellow lines indicate the tracks. As shown in
181 the figure, most of these cyclones develop in the Jiangnan Plain and southern Anhui
182 Province before moving eastward towards the Yellow Sea coast. Some cyclones move
183 northward through Shandong Province and reach the Bohai Sea. These tracks are

184 influenced by upper-level guide airflow at 500~700 hPa. They are also affected by the
185 WPSH and its southwest warm and moist air on its edge (Wei et al., 2013).

186 The frequency of cyclone occurrence, as shown in Figure 2b, refers to the total
187 number of cyclones during the Meiyu period from 1961 to 2020. The genesis locations
188 of cyclones are mainly located in the middle and lower reaches of the Yangtze River
189 and the Huaihe River basin, exhibiting an east–west band distribution (Wang et al.,
190 2015; Wu et al., 2021). In particular, the region encompassing Western and Eastern
191 Hubei Province experiences a higher frequency of cyclone occurrence refers to the total
192 number of cyclones during this period. Research has found a close relationship between
193 the genesis locations of cyclones and topography (Xu 2021; Zhang et al., 2012).



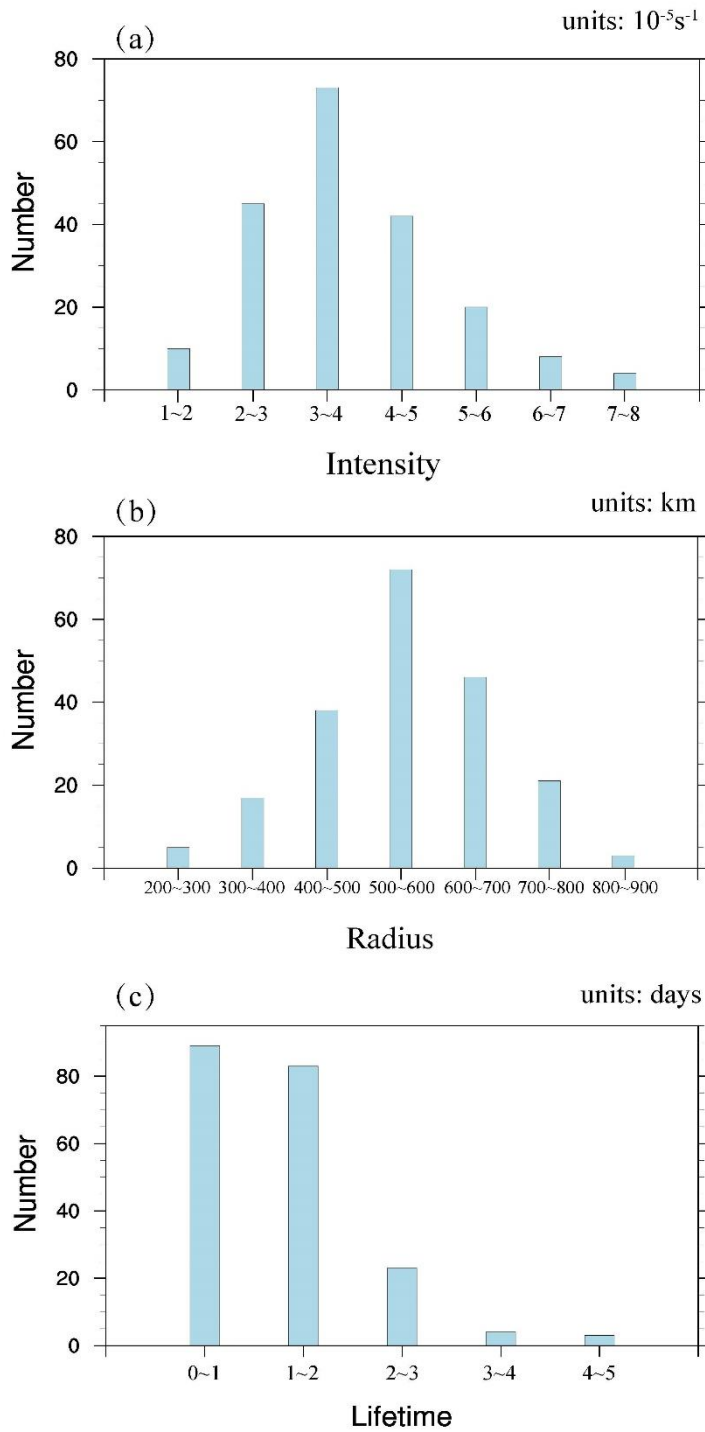
194 Fig 2. Distribution of the cyclone genesis locations, tracks (a) and the frequency of
195 genesis locations refers to the total number of cyclones (b) during the Meiyu period
196 from 1961 to 2020.

197 The climatological characteristics of Jianghuai cyclones over 60 years are
198 examined, with a focus on the intensity, radius, and lifetime of cyclones on land. The
199 intensity of the Jianghuai cyclone is defined as the relative vorticity intensity at the 850
200 hPa cyclone center. Lifetime is defined as the duration during which cyclones affect
201 precipitation on land. Figure 3a illustrates that among the 202 selected cyclones, the
202 intensity of their centers mainly ranges from $1.5 \times 10^{-5} \text{ s}^{-1}$ to $7.3 \times 10^{-5} \text{ s}^{-1}$. The number of
203 cyclones in the range of $3 \times 10^{-5} \text{ s}^{-1}$ to $4 \times 10^{-5} \text{ s}^{-1}$ has the largest proportion, accounting

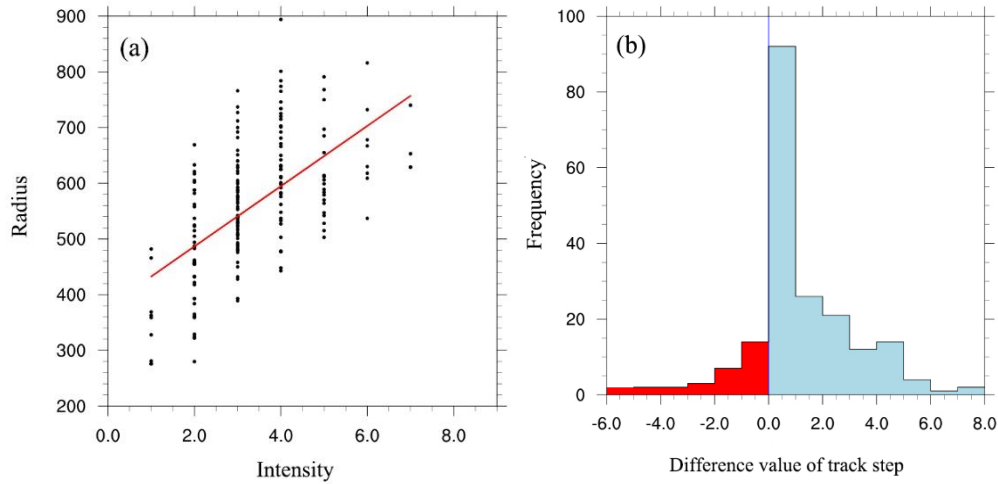
204 for 36% of the total number of cyclones. The total number of cyclones with intensities
205 ranging from $2 \times 10^{-5} \text{ s}^{-1}$ to $5 \times 10^{-5} \text{ s}^{-1}$ amounts to 180, constituting 89% of the total. The
206 relationship between the radius of cyclones and the number of cyclones is shown in
207 Figure 3b. The majority of cyclones have an average radius ranging from 300 to 800
208 km, constituting 96% of the total count. Among these, the highest proportion lies within
209 the range of 500 to 600 km, accounting for approximately 35%. Figure 3c shows the
210 correlation between the lifetime and the number of cyclones. Most cyclones affect
211 precipitation on land for a period of 1-3 days. The number of cyclones' lifetimes that
212 affect precipitation on land within 2 days is 186, accounting for 92% of the total number.

213 The intensity of cyclone is one of the factors affecting its precipitation and impact
214 range during the Meiyu period (Zhao et al., 2010). Figure 4a shows a positive
215 correlation between the maximum intensity and the maximum radius of cyclone
216 development. Therefore, most strong cyclones have a larger horizontal scale than weak
217 cyclones, resulting in greater precipitation and a larger precipitation range. From the
218 distribution of difference values in track step between the maximum intensity and
219 radius of the cyclone shown in Figure 4b, it can be observed that 45% of all cyclones
220 reach both their maximum intensity and horizontal scale simultaneously. Among the
221 remaining Jianghuai cyclones, more of them reach maximum intensity first and
222 continue to develop to the maximum horizontal scale.

223

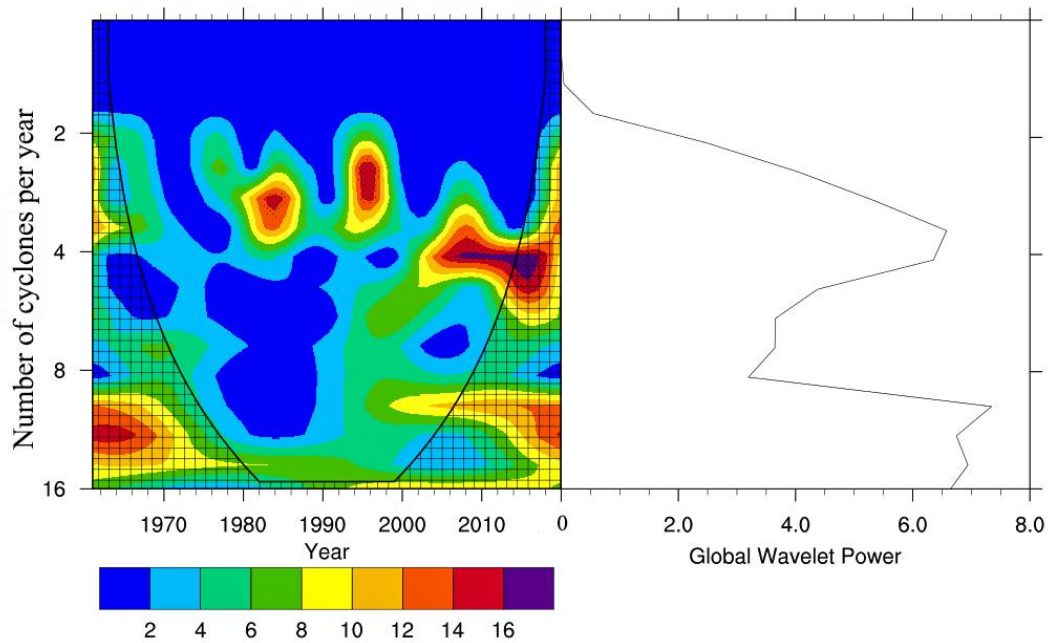


224 Fig 3. Distributions of the number of selected cyclones versus their (a) intensities (units:
 225 10^{-5}s^{-1}), (b) radii (units: km), and (c) lifetimes (units: days).



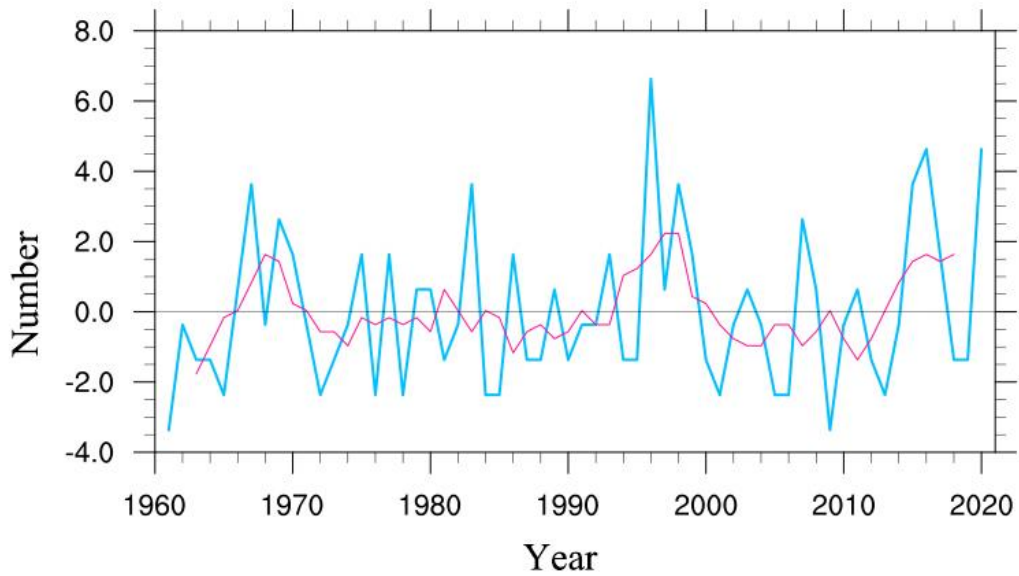
226 Fig 4. Correlation between maximum intensity (units: 10^{-5} s^{-1}) and maximum radius
 227 (units: km) (a) and their difference value of track step during the development of the
 228 Jianghuai cyclone in the Meiyu period (b).

229 The frequency of Jianghuai cyclone occurrence refers to the total number of
 230 cyclones characterized by quasi-periodic variation (Figure 5). The shaded area in the
 231 figure indicates that the 95% confidence interval, according to Student's t-test, is passed.
 232 A strong quasi-periodic variation of 2-4 years is observed for 1980-1990 and 1990-
 233 2000. After 2000, there is an approximate change period of 4-5 years in cyclones. This
 234 change period corresponds to the period of abnormal change in Meiyu. Chen et al.
 235 (2019) pointed out that a quasi-periodic change of 3~4 years is the main component of
 236 abnormal changes in Meiyu. This component of quasi-periodic variable is mainly
 237 influenced by the Indian Ocean dipole (IOD) (Liang et al., 2018). During the positive
 238 phase of the IOD, strong warming of the Indian Ocean triggers a strong Indian monsoon,
 239 leading to a strengthening WPSH and increased precipitation in southern China. The
 240 southwesterly low-level jet, which are enhanced by the positive IOD, also provide water
 241 vapor and warm advection to generate favorable conditions for the development of the
 242 Jianghuai cyclone.



243 Fig 5. Periodic wavelet analysis diagram of Jianghuai cyclones during the Meiyu period
 244 from 1961 to 2020 (units: year) (shadow indicates passing the 95% confidence interval
 245 according to Student's t-test).

246 Jianghuai cyclones are characterized not only by quasi-periodic variation but also
 247 by significant decadal variation. Figure 6 shows the anomaly in activity frequency and
 248 the 5-year sliding average of cyclones during the Meiyu period from 1961 to 2020. The
 249 frequency of cyclone activity is the highest in 1996 and the lowest in 1961 and 2009.
 250 In the long term, there are positive anomalies in cyclone activity frequency in the
 251 middle and lower reaches of the Yangtze River during 1965-1970, 1990-2000, and after
 252 2000, while negative anomalies occurred during 1970-1990 and 2000-2010. The
 253 decadal variation evolution of Jianghuai cyclones is similar to the decadal variation of
 254 precipitation during the Meiyu period (Chen et al., 2019). The decadal variation of
 255 precipitation during the Meiyu period with positive anomaly in 1965-1970, 1995-2000
 256 and 2010-after, and negative anomaly in 1970-1980, 1985-1995 and 2000-2010.



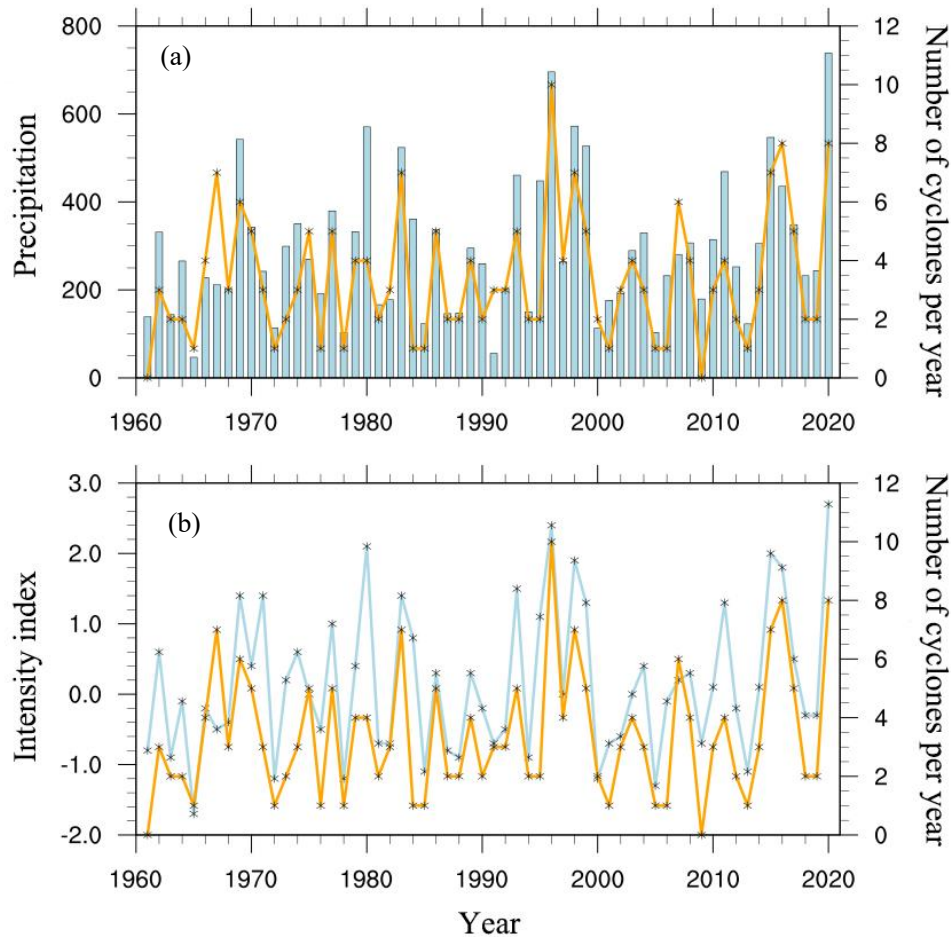
257 Fig 6. Frequency anomaly and 5-year sliding average of cyclones. The blue line shows
 258 the anomalies in the number of cyclones, and the pink line shows the 5-year sliding
 259 average of the anomalies.

260 **3.2 Linkage between cyclone activity and concurrent rainfall in the middle and**
 261 **lower reaches of the Yangtze River.**

262 The Jianghuai cyclones are mainly active in the middle and lower reaches of the
 263 Yangtze River (Huang et al., 2019; Li et al., 2002). Due to the strengthening westward
 264 extension of the WPSH during the Meiyu period, the Jianghuai cyclones are partially
 265 restricted from reaching the sea (Qin et al., 2015; Wu et al., 2020b). They generate
 266 rainstorms in both the middle and lower reaches of the Yangtze River and coastal areas.
 267 A significant portion of the precipitation during the Meiyu period comes from cyclone
 268 precipitation (Zhang et al., 2018). The intensity of Meiyu is typically measured by its
 269 intensity index, which is influenced not only by precipitation but also by the number of
 270 rainy days throughout this period. Both factors collectively determine that year's Meiyu
 271 intensity.

272 The time-series plots of the number of cyclones relate to precipitation and the
 273 intensity index during the Meiyu period from 1961 to 2020 are presented in Figure 7a
 274 and 7b. We find that there is a positive correlation coefficient of 0.77 between the

275 number of cyclones and precipitation in the Meiyu period, passing the student's t-test at
 276 a confidence interval of 99%. Additionally, we observe a positive correlation between
 277 the number of cyclones and the Meiyu intensity index, with a correlation coefficient of
 278 0.76 also passing the student's t-test at a confidence interval of 99%.



279 Fig 7. (a) Changes in precipitation (blue bar chart) (unit: mm/day) and the number of
 280 cyclones (orange line); (b) intensity index (blue line) and the number of cyclones
 281 (orange line) in the Meiyu period from 1961 to 2020.

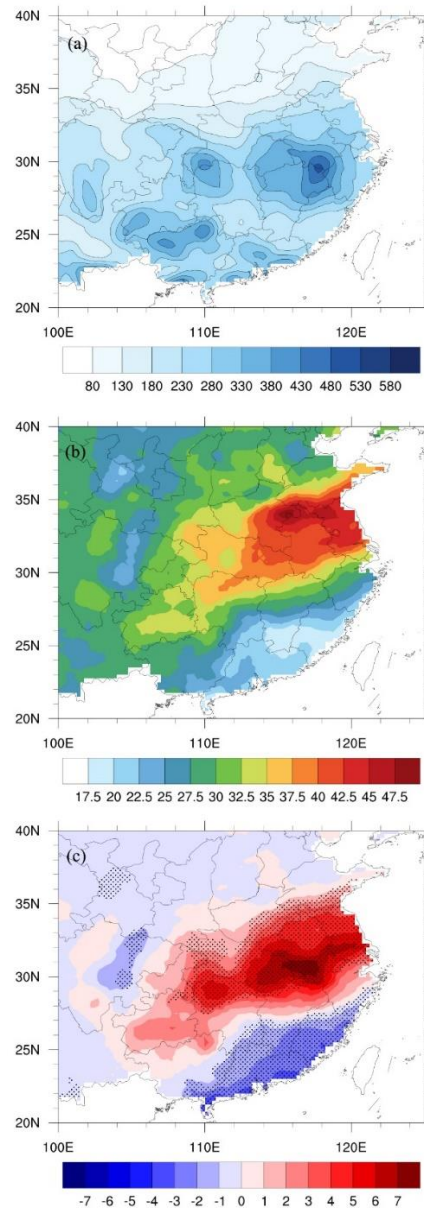
282 Figure 8a shows the spatial distribution of mean annual total precipitation during
 283 the Meiyu period from 1961 to 2020. The areas with high precipitation values in the
 284 middle and lower reaches of the Yangtze River are mainly located in Dabie Mountains
 285 of Anhui Province, northern Jiangxi Province, eastern Hubei Province and western
 286 Hubei Province. In southern Anhui, the maximum average annual precipitation during

287 the Meiyu period can even exceed 480 mm. The occurrence of high precipitation areas
288 during the Meiyu period is closely related to the region topography (Wu et al., 2023).

289 If precipitation and Jianghuai cyclone activity occur on the same day during the
290 Meiyu period, we define that day as a Jianghuai cyclone precipitation day. Figure 8b
291 shows the spatial distribution of the proportion of cyclone precipitation relative to total
292 precipitation during the Meiyu period. As depicted in the figure, the middle and lower
293 reaches of the Yangtze River are primarily affected by cyclone precipitation. The
294 Huaihe River basin in northern Anhui Province is particularly impacted, with cyclone
295 precipitation accounting for over 47% of total precipitation during the Meiyu period. In
296 other regions, cyclone precipitation accounts for more than 35% of total rainfall. In
297 general, the degree of cyclone-influenced precipitation in the middle and lower reaches
298 of the Yangtze River shows an east–west band distribution and a gradual decrease from
299 coastal to inland areas. This suggests that both the concentration of high-value areas
300 and band distribution characteristics are associated with northeastward and eastward
301 tracks followed by Jianghuai cyclones. The capacity for producing precipitation
302 gradually increases as these cyclones develop.

303 Figure 8c shows the spatial distribution of the daily mean precipitation anomaly
304 associated with the Jianghuai cyclone. The shaded region indicates that the 95%
305 confidence interval is passed according to Student's t-test. The mentioned anomalies
306 refer to the difference value between the daily mean values of meteorological elements
307 during selected Jianghuai cyclones and their corresponding values over a sixty-year
308 period. During active periods of the Jianghuai cyclone, abnormal increases in
309 precipitation are observed in the middle and lower reaches of the Yangtze River
310 eastward of 108°E, while Fujian, Guangdong, and other regions experience abnormal
311 decreases. Notably, the precipitation anomalies exceed 7 mm/day in southern Anhui,
312 eastern Hubei, and northern Jiangxi. The high-value regions for precipitation anomalies
313 are consistent with the high-value areas of cyclone occurrence frequency. This suggests
314 a connection between the spatial distribution of precipitation anomalies and cyclone
315 genesis locations. Furthermore, this phenomenon of increasing and decreasing

316 precipitation anomalies is bounded by approximately 27°N and distributed north-south
317 in dipole form.



318
319 Fig 8. (a) Mean annual total precipitation during the Meiyu period from 1961 to 2020
320 (units: mm/year); (b) proportion of Jianghuai cyclone precipitation relative to total
321 precipitation during the Meiyu period (units: %); (c) daily mean precipitation anomaly
322 of the Jianghuai cyclone during the Meiyu period (units: mm/day) (shadow indicates
323 passing the 95% confidence interval according to Student's t-test).

324 The evolution of composite geopotential height and horizontal wind anomalies for

325 two different levels of Jianghuai cyclones from day -4 to +2 during the Meiyu period is
326 shown in Figure 9. Composite geopotential height anomalies are significant at the 95%
327 confidence level based on Student's t-test. Vectors are plotted if wind anomalies are
328 significant at the 95% confidence level based on Student's t-test in at least one direction.

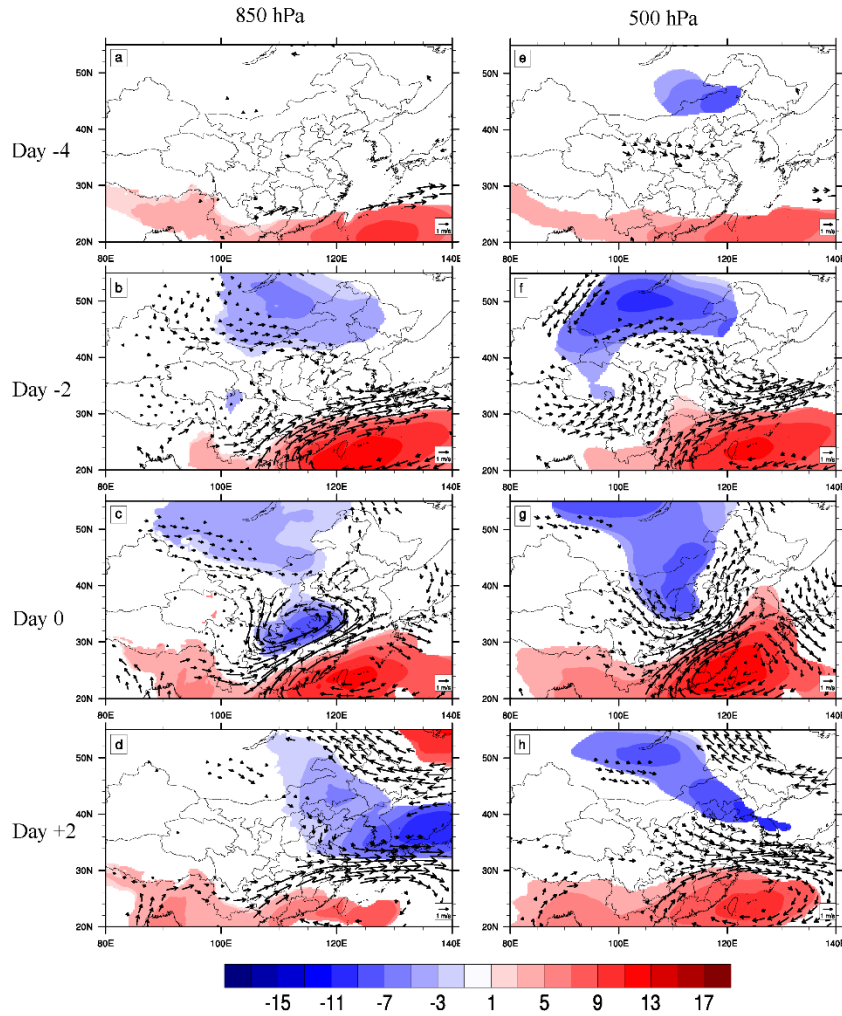
329 Day 0 is the day on which the cyclone first appears in the specified area. On day
330 0, most areas of the middle and lower Yangtze River in the lower and middle
331 troposphere (850 hPa) are covered by significant negative geopotential height
332 anomalies with peak magnitudes greater than -11 gpm. There is a significant positive
333 geopotential height anomaly with a peak magnitude of over 13 gpm on the southeast
334 side of the negative geopotential height anomaly. These anomalies form meridional
335 dipole structures in the geopotential height field of the middle and lower troposphere.
336 The southwest wind anomaly is significant in the middle and lower reaches of the
337 Yangtze River. Southern Anhui Province and northern Jiangxi Province are between the
338 positive geopotential height anomaly and negative geopotential height anomaly. The
339 existence of these anomalies indicates enhanced southwest winds and strengthening of
340 WPSH. The negative geopotential height anomalies at 500 hPa on day 0 are mainly
341 observed in Mongolia, Shanxi and other regions. Strong southwest wind anomalies
342 exist between these positive and negative geopotential height anomalies. The negative
343 geopotential height anomalies exceed -7 gpm in the Mongolian region.

344 The negative geopotential height anomalies on all the isobaric surfaces can be
345 traced back to Mongolia, Inner Mongolia and parts of Northeast China on day -2.
346 Negative geopotential height anomalies at 500 hPa can be traced back to day -4. On day
347 -4, significant southwestern wind anomalies exist in southwestern Hunan at 850 hPa.
348 Significant northwest wind anomalies exist in the Yellow River basin of China at 500
349 hPa. On day -2, the negative geopotential height anomalies in Mongolia, Inner
350 Mongolia and some northeastern areas are enhanced. The positive geopotential height
351 anomalies of the WPSH are enhanced and extend northward to the southern part of the
352 middle and lower reaches of the Yangtze River. There are significant southwest wind
353 anomalies in the south of the middle and lower reaches of the Yangtze River, while

354 there are significant northwest wind anomalies at 500 hPa in northern Anhui Province
355 and Jiangsu Province. The negative geopotential height anomalies move eastward with
356 the formation and development of Jianghuai cyclones. On day +2, anomalous southwest
357 winds combined with northwest winds mainly affect the lower reaches of Yangtze River
358 region. The positive geopotential height anomaly of the WPSH is weakened.

359 Therefore, the abnormal precipitation caused by the Jianghuai cyclone mainly
360 originates from abnormal southwest winds and the strengthening of the WPSH. The
361 enhanced southwesterly low-level jet provides sufficient warm and moist air for the
362 formation of cyclones and promotes the eastward migration of cyclones after formation.
363 Liu et al. (2020) and Zhao et al. (2021) studied the causes of the super strong Meiyu
364 year in 2020, mentioned that the WPSH is unusually strong and westward accompanied
365 by an abnormal increase in precipitation. Liu et al. (2020) found that the enhanced
366 southwesterly low-level jet stream is conducive to the development of vertical
367 movement in the middle and low levels, providing necessary dynamic conditions for
368 sustained precipitation during the Meiyu in 2020.

369 Cold air activity is one of the important factors for the formation of heavy
370 precipitation, as it promotes the convergence and uplift of low level necessary for heavy
371 precipitation (Liu et al., 2020). The enhanced negative geopotential anomaly over
372 Mongolia brings cold and dry air through the westerly jet, contributing to cyclone
373 development. The increased frequency of cyclones over the Yangtze River and Huaihe
374 River results in abnormal increase in precipitation in the middle and lower reaches of
375 the Yangtze River during the Meiyu period. However, due to the strengthening of the
376 WPSH, the southern part of China is controlled by the abnormal positive geopotential
377 height, and the precipitation decreases. Zhao et al. (2021) also found that when the
378 WPSH enhanced, there was a decrease in precipitation in South China.



379 Fig 9. Evolution of composite geopotential height anomalies (shading; units: gpm) and
 380 horizontal wind anomalies (units: m/s) on the 850 hPa and 500 hPa isobaric surfaces
 381 for day -4 (a, e), day -2 (b, f), day 0 (c, g) and day +2 (d, h) for the 202 selected
 382 Jianghuai cyclones. Shading indicates that composite geopotential height anomalies are
 383 significant at the 95% confidence level based on Student's t-test. Vectors are plotted if
 384 wind anomalies are significant at the 95% confidence level based on Student's t-test in
 385 at least one direction.

386 The climatic distribution of water vapor flux and water vapor flux divergence at
 387 850 hPa during the Meiyu period of the Jianghuai cyclone is shown in Figure 10. The
 388 precipitation process of the Jianghuai cyclone mainly involves water vapor brought by
 389 the southwesterly low-level jet of the summer monsoon in the low-latitude area. During
 390 the development of the Jianghuai cyclone, the middle and lower reaches of the Yangtze

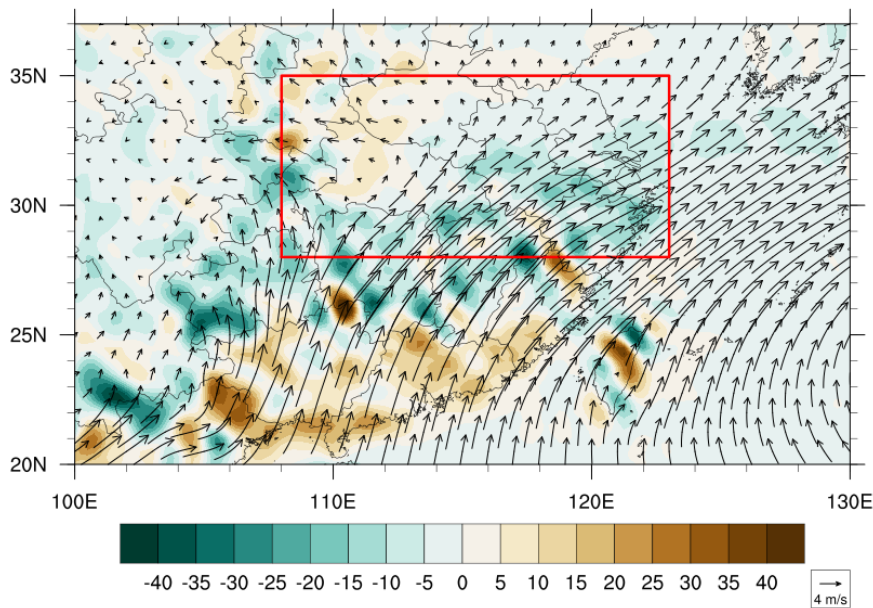
391 River are mostly located in the water vapor convergence area, which is conducive to
392 precipitation generation (Chen et al., 2020).

393 Figure 11 shows the distribution of water vapor flux anomalies and water vapor
394 flux divergence anomalies at the pressure level of 850 hPa during the Jianghuai cyclone
395 from day -1 to day +2. Both the color field and wind vector arrows in the figure passed
396 the 95% significance according to Student's t-test. On day -1, significant water vapor
397 convergence anomaly and southwestward water vapor transport appear in eastern and
398 western Hubei Province. In the middle reaches of the Yangtze River, there is a north-
399 south distribution of water vapor convergence and divergence dipoles. The anomalies
400 of water vapor flux and water vapor flux dispersion are mainly concentrated on day 0.
401 There is significant anomalous water vapor convergence up to $-1 \text{ g} \cdot \text{cm}^{-2} \cdot \text{hPa}^{-1}$ in eastern
402 Hubei Province, Anhui Province and Jiangsu Province on day 0. Anomalous water
403 vapor dispersion exists in the middle and lower reaches of the Yangtze River. On the
404 day +1 of cyclone development, the spatial distribution of water vapor convergence and
405 divergence shifted to an east-west direction. The anomalous transport of water vapor
406 flux also changes to an east-west direction. On day +2, with the eastward movement
407 development of the cyclone, only southern Jiangsu Province and northern Zhejiang
408 Province have abnormal water vapor flux in the eastward direction. At this time,
409 precipitation in that area begins to gradually weaken.

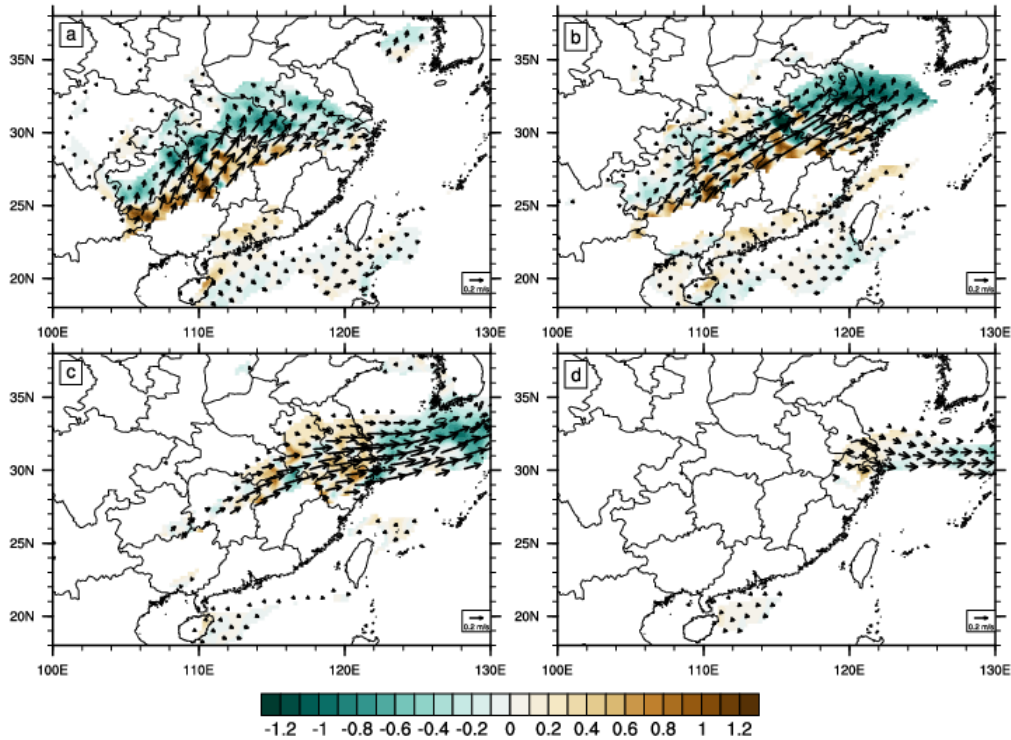
410 From day -1 to day 0, the abnormal water vapor flux and water vapor flux
411 divergence configuration results in the transportation of warm and wet air from the low-
412 latitude area to the middle and lower reaches of the Yangtze River. The negative water
413 vapor flux indicates the occurrence of water vapor convergence, leading to an increase
414 in local water vapor content and ultimately resulting in enhanced precipitation in the
415 region. In a study by Liu et al. (2020) on heavy rainfall in 2020, they found enhanced
416 water vapor transport and repeated occurrences of convergence movement, both
417 contributing to prolonged precipitation duration in the Jianghuai River basin.

418 In contrast, the anomaly of water vapor flux in southern Guangdong and other
419 regions is divergent. This leads to a decrease in local water vapor volume and

420 precipitation within these regions. These results indicate that the variations in water
421 vapor flux and divergence associated with cyclones primarily originate from warm and
422 moist air transported from low latitudes to the middle and lower reaches of the Yangtze
423 River. Therefore, there exists a positive correlation between cyclone activity and
424 precipitation in the middle and lower reaches of the Yangtze River.



425 Fig 10. Distribution of 850 hPa daily mean water vapor flux (unit: $\text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}$) and
426 water vapor flux divergence (unit: $10^{-8}\text{ g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}\cdot\text{s}^{-1}$) of cyclones over the Yangtze
427 and Huaihe rivers during 1961-2020 (color diagram shows water vapor flux divergence,
428 and vector diagram shows water vapor flux).



429 Fig 11. Distribution of the 850 hPa daily mean water vapor flux anomaly (unit: $\text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}$) and water vapor flux divergence anomaly (unit: $10^{-8} \text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}\cdot\text{s}^{-1}$) of
 430 cyclones over the Yangtze and Huaihe rivers during 1961-2020 (color diagram shows
 431 water vapor flux divergence, and vector diagram shows water vapor flux). The colored
 432 region passed the 95% confidence interval according to Student's t-test. If the vapor
 433 flux anomaly is significant at the 95% confidence level for Student's t-test in at least
 434 one direction (zonal or meridian), the vector is plotted.
 435

436 4. Summary and discussion

437 Based on ERA5 reanalysis of sea level pressure data and using the relative
 438 vorticity method to identify and track cyclones, we examine the climatological
 439 characteristics of Jianghuai cyclones. Additionally, we analyze the linkages between
 440 cyclone activity and precipitation in the middle and lower reaches of the Yangtze River
 441 during the Meiyu period.

442 During the Meiyu period, Jianghuai cyclones are mainly generated at the Western
 443 Hubei Province and Eastern Hubei Province. These cyclones develop and move to the

444 sea in the east or northeast direction. There is a positive correlation between the
445 maximum intensity and maximum radius of Jianghuai cyclones. The cyclone
446 occurrence frequency has a 2~4 years quasi-periodic variation for 1980-1990 and 1990-
447 2000, and change to 4~5 years after 2000. The frequency of cyclone activity shows a
448 positive anomaly during 1965-1970, 1990-2000, while it has a negative anomaly during
449 1970-1990 and 2000-2010. Both the quasi-periodic and interdecadal variations of
450 cyclones have a good correspondence well with the precipitation during the Meiyu
451 period.

452 There is a positive correlation between the frequency of cyclone activity and
453 precipitation during the Meiyu period. The frequency of Jianghuai cyclone activity is
454 high in the years with strong Meiyu rainfall and low in the years with weak Meiyu
455 rainfall. The percentage of precipitation affected by Jianghuai cyclone activity in the
456 middle and lower reaches of the Yangtze River can reach up to 47%. The spatial
457 distribution forms an east-west belt, with decreasing influence from the coast to inland.
458 During active periods of Jianghuai cyclones, the precipitation increases abnormally in
459 the middle and lower reaches of the Yangtze River east of 108°E. Precipitation
460 decreases abnormally in Fujian Province and Guangdong Province. The spatial
461 distribution of precipitation anomalies is related to the genesis locations frequency of
462 cyclone, forming positive and negative anomalies distributed north-south as dipoles
463 along approximately 27°N latitude line.

464 The geopotential height anomaly field and the horizontal wind vector anomaly
465 field of the Jianghuai cyclones during the Meiyu period are synthesized and analyzed.
466 There is an enhanced positive geopotential height anomaly of the WPSH during cyclone
467 activity. The negative geopotential height anomaly of Mongolia and the abnormal
468 southwesterly low-level jet are enhanced. All these factors lead to an increasing
469 precipitation in the middle and lower reaches of the Yangtze River. The abnormal
470 leading signal of negative geopotential height in Mongolia can be traced back to day -
471 2 of the cyclone activity, and the signal can be traced back to day -4 at 500 hPa. From
472 day -2 to day 0 of cyclone activity, the abnormal distribution of water vapor flux and

473 water vapor flux divergence cause warm and wet air from low latitudes to be transported
474 to the middle and lower reaches of the Yangtze River. These promote the generation
475 and development of cyclones and increase precipitation in the middle and lower reaches
476 of the Yangtze River.

477 We explore the cyclone characteristics and study the linkage between cyclone
478 activity and precipitation in the Yangtze River region. Spatially, we identified abnormal
479 precipitation patterns by tracing the evolution of geopotential height anomalies and
480 water vapor flux. However, the specific mechanism by which the southwesterly low-
481 level jet affects cyclones during the Meiyu period is not clear enough. Zhang et al. (2018)
482 suggested that the strengthening of the southwesterly low-level jet would lead to the
483 development of α mesoscale low-pressure disturbance near the Meiyu Front and result
484 in extreme precipitation events. Liu et al. (2020) found that the strengthening of the
485 southwesterly low-level caused the southerly meridional strong gradient zone on the
486 north side of the meridional wind maximum center to move northward in the low-level
487 dynamic conditions of the rainstorm process during the Meiyu period. The influence of
488 the southwesterly low-level stream on the development of physical factors promoting
489 the formation of Jianghuai cyclones needs to be considered and analyzed. Zhao et al.
490 (2010) found that Jianghuai cyclones with different intensities have different causes.
491 Therefore, it is also necessary to consider how different intensities of Jianghuai
492 cyclones affect precipitation. These issues require further analysis and research.

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493 **Competing interests**

494 The contact author has declared that none of the authors has any competing interests.

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