

1 **Climatic characteristics of the Jianghuai cyclone and its linkage with**
2 **precipitation during the Meiyu period from 1961 to 2020**

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8 **Abstract.** This study examines the climatic characteristics of 202 Jianghuai cyclones
9 and their linkage with precipitation during the Meiyu period from 1961 to 2020. The
10 results show that cyclones mainly originate from eastern Hubei Province and south-
11 central Anhui Province, and further explore the statistical characteristics of the strength,
12 radius, and their positive correlation. When studying the decadal variation of cyclones,
13 we find that there is a similar trend between the decadal variation of cyclones and Meiyu
14 precipitation. Therefore, we further investigate the correlation between the Jianghuai
15 cyclones and the precipitation during the Meiyu period. There is a positive correlation
16 coefficient of 0.77 between them. It's worth mentioning that the percentage of
17 precipitation affected by cyclone activities can reach up to 47%. The anomalous
18 increase in precipitation caused by cyclones above 27°N can reach a maximum of 7
19 mm/day. When the cyclone existed, there was a significant altitude anomaly of negative
20 geopotential height can be traced to day -4 at the 500 hPa level over Mongolia. The
21 abnormally enhanced WPSH, southwesterly low-level jet and negative geopotential
22 height are the dominant factors causing abnormal precipitation during Jianghuai
23 cyclones. Before and after the cyclone developed, water vapor flux and divergence from
24 low latitudes abnormally increased. These provide sufficient water vapor conditions for
25 the generation of cyclone precipitation.

26 1. Introduction

27 Meiyu is a special rainy season due to the progress of the East Asian summer
28 monsoon. The East Asian summer monsoon broke out in the South China Sea in mid-
29 May and then advanced northward, forming rain bands in South China, the Jianghuai
30 region, the Korean Peninsula and Japan (Ding et al., 2004,2007; Qian et al., 2000). The
31 name for this special rainy season is called Meiyu in China, while it is called Changma
32 in South Korea and it is called Baiu in Japan (Ninomiya et al., 1987; Oh et al., 1997;
33 Saito. 1995). Meiyu front is one of the important weather systems affecting summer
34 precipitation in the middle and lower reaches of the Yangtze River (Pang et al., 2013;
35 Tao et al., 1979; Wang et al., 2014; Zhou et al., 2022). From mid-June to early July, the
36 east of Yichang, Hubei Province, has continuous rains and short sunshine. These
37 conditions are accompanied by heavy rainfall, strong wind and other weather
38 phenomena in these areas during the Meiyu period (Ding. 1992; Zhao et al., 2021; Zhou
39 et al., 2017). In China, the mean annual precipitation during the Meiyu period in the
40 Jianghuai River Basin can reach 300 mm, accounting for 30%-40% of the mean annual
41 total precipitation, and even up to 500 mm or more in the extreme Meiyu period (Liu et
42 al., 2020). Historically, most of the summer **floods** disasters are caused by precipitation
43 anomalies in the Meiyu period. Some scholars have studied and analyzed the
44 representative floods of 1996, 1998, 2016 and 2020 (Bao et al., 2021; Su et al., 2021;
45 Zhao et al., 2018; Zhong et al., 2023). These floods, caused by the Meiyu front, had
46 adverse effects on people's safety, lives and property (Yan et al., 2021). Scholars in
47 China have divided rainstorms caused by Meiyu fronts into three types (Zhang et al.,
48 2004). The first type is the β mesoscale convective rainstorm on the Meiyu front. This
49 type of rainstorm has a range of less than 300 km with strong intensity and a fast
50 formation process (He et al., 2007). It is difficult to forecast before 12 hours and can be
51 detected only by using radar to make a proximity forecast (Zhang et al., 2002). The
52 second type is the persistent rainstorm located in front of the upper trough of low
53 pressure in the western part of the Meiyu front. It is characterized by a long duration of
54 approximately 5 days but appears less frequently, mainly in western Hubei and western

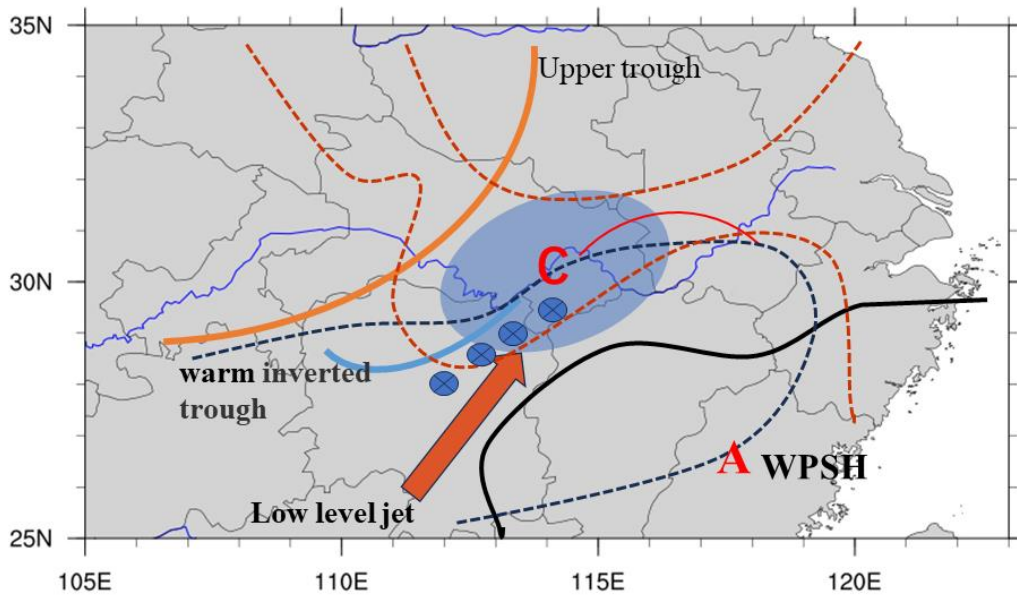
55 Hunan and Sichuan (Cai et al., 2021; Wu et al., 2020a;). The last type is the rainstorm
56 caused by the Jianghuai cyclone located east of the origin of the Meiyu. The Jianghuai
57 cyclones are affected by the thermal conditions of the sea and land and likely occur in
58 the eastern part of the Meiyu front (Wang et al., 2016). The positive vorticity advection
59 in front of the high-altitude trough and the warm advection in front of the front promote
60 the eastward movement and development of the cyclone (Shen et al., 2019; Zhang et
61 al., 2016). During the development of the cyclone, the lower levels are dominated by
62 the southwest warm and humid airflow, and the high levels are mainly affected by dry
63 and cold air (Zhao et al., 2008). This type of rainstorm has a large range, high intensity
64 and long duration of precipitation (Wang et al., 2012; Xu et al., 2011).

65 Scholars' studies on Jianghuai cyclones during the Meiyu period were initially
66 based on individual case analysis. Xu et al. (2013) studied a cyclone process in 2011
67 and found that the cyclone process lasted up to 36 h. The cyclone rainstorm was
68 distributed on the south side of the cyclone. Heavy precipitation during the whole
69 cyclone mainly occurred in the lower reaches of the Yangtze River. Wu et al. (2020b)
70 studied 2 different cyclone rainstorm processes. They found that rainfall is directly
71 proportional to cyclone intensity. There is a strong convergence center of water vapor
72 flux during cyclone development. Zhou et al. (2020) found that a tornado was generated
73 from the cyclone occlusion stage on July 22. The tornado was under the influence of a
74 strong and fast Jianghuai cyclone and produced heavy precipitation accompanied by
75 thunderstorm phenomena. With the improvement of cyclone identification methods and
76 reconstruction of reanalysis data, statistical studies of cyclones have been further
77 developed (Simmonds et al., 1999; 2000; Wernli et al., 2006). Yang et al. (2010)
78 modeled the rainstorm process in the lower reaches of the Yangtze River from 1998 to
79 2005. The cyclones accounted for 62.5% of the rainstorm cases, and more than 70% of
80 the cyclones could develop and produce rainstorms. The Jianghuai cyclone located in
81 the lower reaches of the Yangtze River generally exists in the lower troposphere at 700
82 hPa. The horizontal scale is within 400 km, and the life period on land is generally less
83 than 48 h. Wang et al. (2015) found that the number of cyclones was lower and their

84 intensity was weak in the 1980s and 1990s. In the early 2000s, cyclones were more
85 frequent, and their intensity increased. After 2010, there was again a decreasing trend.
86 Zhang et al. (2018) divided 60 cases of extreme precipitation in the middle reaches of
87 the Yangtze River from 2008 to 2015 into five types. Among them, the extreme
88 precipitation of the Jianghuai cyclone type accounted for 30%. The stable and
89 maintained Western Pacific subtropical high (WPSH) system is one of the important
90 reasons for the strong precipitation produced by cyclones. Because of the weak cold air
91 force, the intensity of the Jianghuai cyclone is weaker than that in spring (Zhou et al.,
92 2017). The daily analysis of the Jianghuai cyclones in the Meiyu period is easy to ignore.
93 All these studies indicate that the Jianghuai cyclone is an important weather system that
94 causes heavy rainfall during the Meiyu period in the middle and lower reaches of the
95 Yangtze River (Wu et al., 2021; Zhang et al., 2018; Zhu et al., 1998).

96 Research on the climatic characteristics and precipitation effects of Jianghuai
97 cyclones during the Meiyu period in the past 60 years has not yielded clear results. In
98 this study, the relative vorticity method is used to objectively identify and track
99 cyclones based on reanalysis data provided by ERA5. The climatological characteristics
100 of the Jianghuai cyclones during this period are studied. We analyze the correlation
101 between Jianghuai cyclone activity and precipitation. This study provides a reference
102 for the long-term and short-term forecasting of precipitation in the Meiyu period.

103 The remainder of the present paper is organized as follows. Section 2 of this paper
104 presents the dataset and analytical methods. In Section 3, we show the climatology
105 composite of the cyclone tracks, genesis locations, intensity, lifetime and so on. There
106 is a positive correlation between the frequency of cyclonic activity and precipitation in
107 the Meiyu period. The relationship between them is studied by means of the
108 geopotential height anomaly and water vapor flux anomaly. Section 4 provides the main
109 discussion and findings of this study.



110

111 Fig.1 Schematic diagram of the main weather system and the structure of temperature
 112 and pressure field in the middle and low levels of the Jianghuai cyclone. (Red dotted
 113 line: isotherm; Solid black line: contour line; Blue dot: precipitation area; Solid orange
 114 line: 500 hPa upper-level trough; Red arrow: low level jet; Black dotted line: warm
 115 inverted trough; Solid red line: warm shear; Solid blue line: cold shear; Letter C:
 116 cyclone; Letter A: WSPH.)

117 2. Data and methods

118 2.1 Data

119 The time span of all the data is 60 years from 1961 to 2020, and the study area is
 120 located at 108°E-123°E, 27°N-34°N. We use the ERA5 relative vorticity hourly data
 121 (850 hPa) released by the European Centre for Medium Range Weather Forecasts
 122 (ECMWF) for Jianghuai cyclone identification and tracking. The spatial resolution of
 123 the data is 0.25°×0.25°, and the temporal resolution is 6 h. Every 6 h was defined as a
 124 step. The data of geopotential height, wind field, and specific humidity are daily data
 125 processed from ERA5 hourly data with a spatial resolution of 0.25°×0.25° (Hersbach
 126 et al., 2018). The geopotential height and wind field data include pressure levels of
 127 approximately 500 hPa, 700 hPa and 850 hPa. The specific humidity data include

128 pressure levels of approximately 500 hPa, 700 hPa and 850 hPa. The precipitation data
129 are from the CN05.1 grid point observation dataset compiled by the National
130 Meteorological Information Center of China Meteorological Administration with a
131 resolution of 0.25°×0.25° (Wu et al., 2013; Xu et al., 2009).

132 We used the Meiyu intensity index to characterize the strength of Meiyu, and data
133 is from the National Climate Center of China. The area for which the Meiyu intensity
134 index is calculated is defined in the article (GB/T 33671-2017). The Meiyu intensity
135 index is defined as:

$$M = \frac{L}{L_0} + \frac{0.5(R/L)}{R_0/L_0} + \frac{R}{R_0} - 2.5$$

137 M is the Meiyu intensity index. L is the length of the Meiyu in a given year (unit:
138 day) and L₀ means the average length of the Meiyu over the years (units: day). R is the
139 total precipitation of Jianghuai River basin during Meiyu in a given year, and R₀ is the
140 average total precipitation of Jianghuai River basin during Meiyu over the years. The
141 average period is from 1961 to the current year. For example, L₀ and R₀ values for 2000
142 are the averages from 1961 to 2000. Where M between -0.375 and 0.375, China
143 Meteorological Administration defines this year as the normal. Where M between 0.375
144 and 1.25, this year is defined as a little strong. Where M greater than or equal to 1.25,
145 this year is defined as strong. Where M between -1.25 and -0.375, this year is defined
146 as a little weak. Where M less than or equal to -1.25, this year is defined as weak.

147 2.2 Methods

148 Scholars have proposed a number of methods to identify extratropical cyclones.
149 The objective identification and tracking method for cyclones used in this paper is the
150 vorticity tracking method proposed by Hodges (1994, 1995). This method mainly uses
151 the relative vorticity field at the 850 hPa to determine the feature points of the cyclone.
152 Feature points are used to correspond to the position of the cyclone and to match the
153 cyclone track within a given time span. In addition to the relative vorticity method of
154 tracking proposed by Hodges, different methods of cyclone identification have also
155 been proposed by other scholars. Lu (2017) improved the extratropical cyclone

156 identification and tracking method involving the nine-point pressure minimum. Jiang
157 et al. (2020) proposed an algorithm for identifying extratropical cyclones on the basis
158 of gridded data. This algorithm is named the eight-section slope detection method.

159 Among them, the most commonly used cyclone tracking methods are the mean
160 sea level pressure method (SLP) and 850 hPa relative vorticity method. Mailier et.al
161 (2006) and Zhang et.al (2012) studied the tracks of individual cyclones in these two
162 methods. Both of them found 850 hPa relative vorticity method can identify and detect
163 cyclone center earlier than the SLP method (Mailier et al., 2006). The reason for this
164 result is that SLP is easily affected by topography and large-scale background
165 circulation shear vorticity (Hodges, 1994). So based on this advantage of the relative
166 vorticity method, we select the 850 hPa relative vorticity tracking method. The relative
167 vorticity tracking method can detect low vortex systems earlier and track cyclones for
168 a longer period of time with better stability. When the closed pressure levels are not
169 visible on the satellite map, the vorticity tracking method can still continue to track the
170 cyclone, improving the accuracy of cyclone track data.

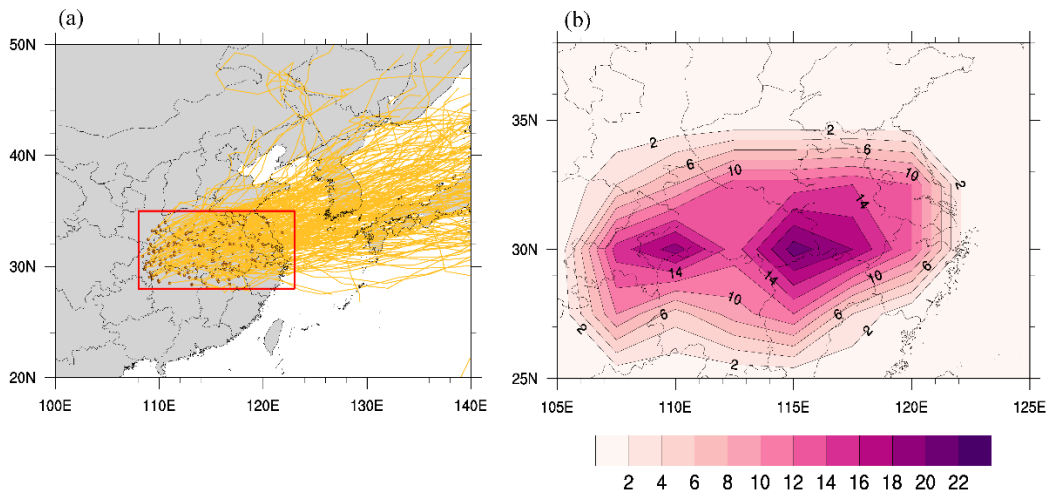
171 **3. Results**

172 **3.1 Climatic characteristics of the Jianghuai cyclone during the Meiyu period**

173 A total of 202 Jianghuai cyclones existed during the Meiyu period from 1961 to
174 2020. The range of cyclone genesis locations defined by the Jiangsu Meteorological
175 Administration (2017) and the characteristics of the relative vorticity tracking method
176 were used. We adjust the genesis location and remove the cyclones that are generated
177 at sea and have no effect on land precipitation (108°E-123°E, 28°N-35°N). To explore
178 the tracks and genesis location of Jianghuai cyclone, we make a statistical of them
179 (Fig.2). The brown dots represent the genesis locations, the first place meeting the
180 criterion, of the Jianghuai cyclone. The yellow lines indicate the tracks of the cyclones.
181 As shown in the figure, most of the cyclones develop in the Jiangnan Plain and southern
182 Anhui Province, then move eastward to the Yellow Sea coast. Some cyclones move
183 northward through Shandong Province and reach the Bohai Sea. These tracks are

184 related to the upper-level guide airflow of 500~700 hPa. They are also affected by the
185 WPSH and its southwest warm and moist air on the edge (Wei et al., 2013).

186 Figure 2b shows the frequency of cyclone occurrence refers to the total number
187 of cyclones during the Meiyu period from 1961 to 2020. The genesis locations of
188 cyclones are mainly located in the middle and lower reaches of the Yangtze River and
189 the Huaihe River basin, with an east–west band distribution (Wang et al., 2015; Wu et
190 al., 2021). The frequency of occurrence refers to the total number of cyclones during
191 the Meiyu period from 1961 to 2020 is higher in the region of the Western Hubei
192 Province and Eastern Hubei Province. Research has found that the genesis locations of
193 cyclones are closely related to the topography (Xu 2021; Zhang et al., 2012).

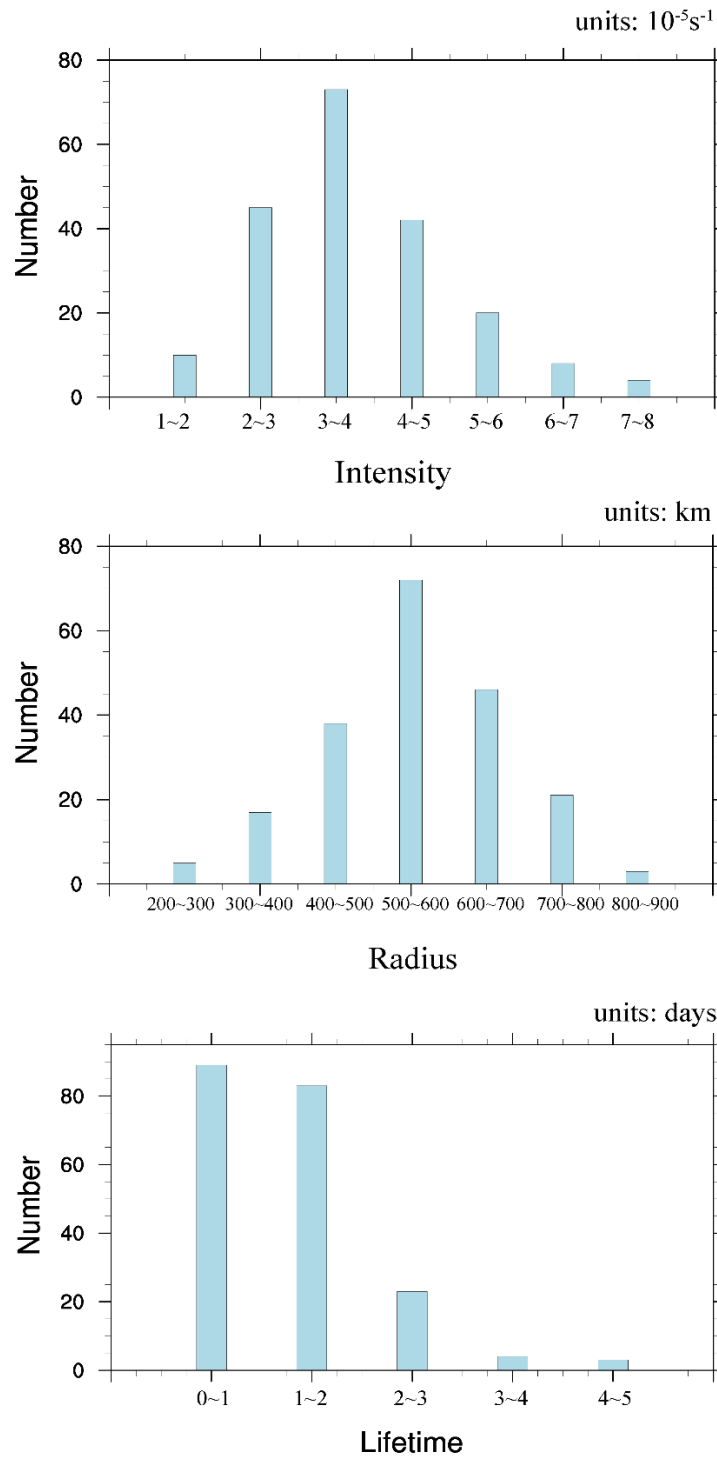


194 Fig 2. Distribution of the cyclone genesis locations, tracks (a) and the frequency of
195 genesis locations refers to the total number of cyclones (b) during the Meiyu period
196 from 1961 to 2020.

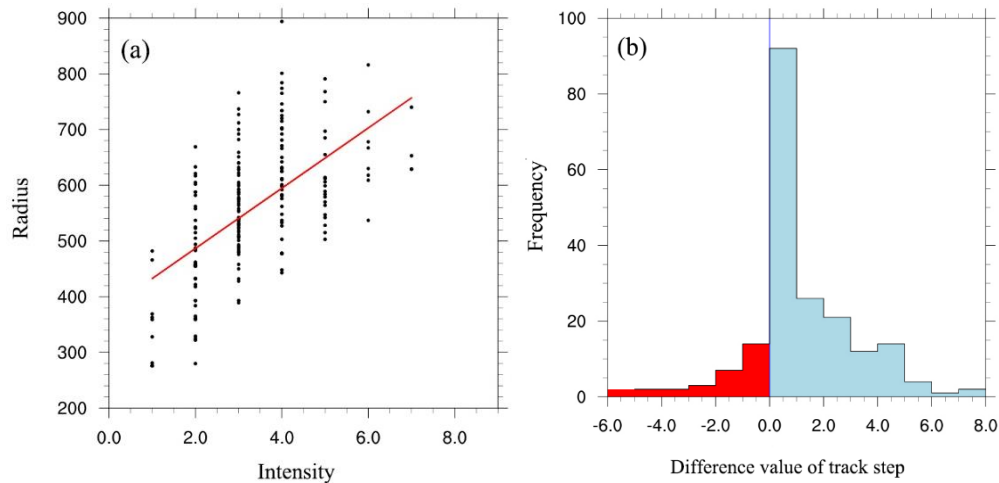
197 To examine the climatological characteristics of Jianghuai cyclones over 60 years,
198 we focus on the intensity, radius, and lifetime of cyclones on land. The intensity of the
199 Jianghuai cyclone is defined as the relative vorticity intensity of the 850 hPa cyclone
200 center. The lifetime is defined as the time of cyclones affecting precipitation on land.
201 Figure 3a shows that among the 202 selected cyclones, the intensity of the cyclone
202 center mainly ranges from $1.5 \times 10^{-5} \text{ s}^{-1}$ to $7.3 \times 10^{-5} \text{ s}^{-1}$. The number of cyclones in the
203 range of $2 \times 10^{-5} \text{ s}^{-1}$ to $3 \times 10^{-5} \text{ s}^{-1}$ has the largest proportion, accounting for 36% of the

204 total number of cyclones. A total of 180 cyclones are in the range of $1.5 \times 10^{-5} \text{ s}^{-1}$ to $5 \times 10^{-5} \text{ s}^{-1}$ in intensity, accounting for 89%. Figure 3b shows the relationship between the
205 radius of cyclones and the number of cyclones. Most of the cyclones have an average
206 radius between 300 and 800 km, accounting for 96% of the total number. The number
207 of cyclones with radii between 500 and 600 km is the largest, accounting for 35%.
208 Figure 3c shows the relationship between the time of cyclones affecting precipitation
209 on land and the number of cyclones. Most of the cyclones affect precipitation on land
210 for 1-3 days, and only one cyclone affects precipitation on land for more than 3 days.
211 The number of cyclones' lifetime that affected precipitation on land within 2 days was
212 186, accounting for 92% of the total number.

214 The intensity of cyclone is one of the factors affecting its precipitation and impact
215 range during the Meiyu period (Zhao et al., 2010). Figure 4a shows a positive
216 correlation between the maximum intensity and the maximum radius of cyclone
217 development. Therefore, the horizontal scale of most strong cyclones is larger than that
218 of weak cyclones, the precipitation is greater, and the precipitation range is larger. From
219 the distribution of difference value of track step between the maximum intensity and
220 the radius of the cyclone shown in Figure 4b, the number of cyclones that reach both at
221 the same time accounts for 45% of the total number of cyclones. Of the remaining
222 Jianghuai cyclones, more reach the maximum intensity first and continue to develop to
223 the maximum horizontal scale.

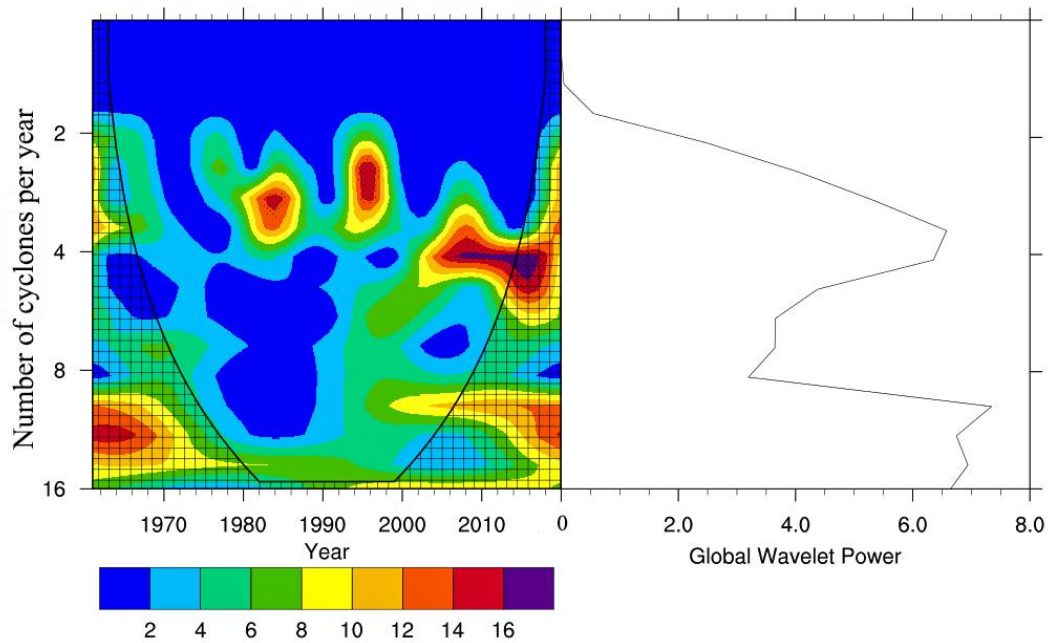


224 Fig 3. Distributions of the number of selected cyclones versus their (a) intensities (units:
 225 10^{-5}s^{-1}), (b) radii (units: km), and (c) lifetimes (units: days).



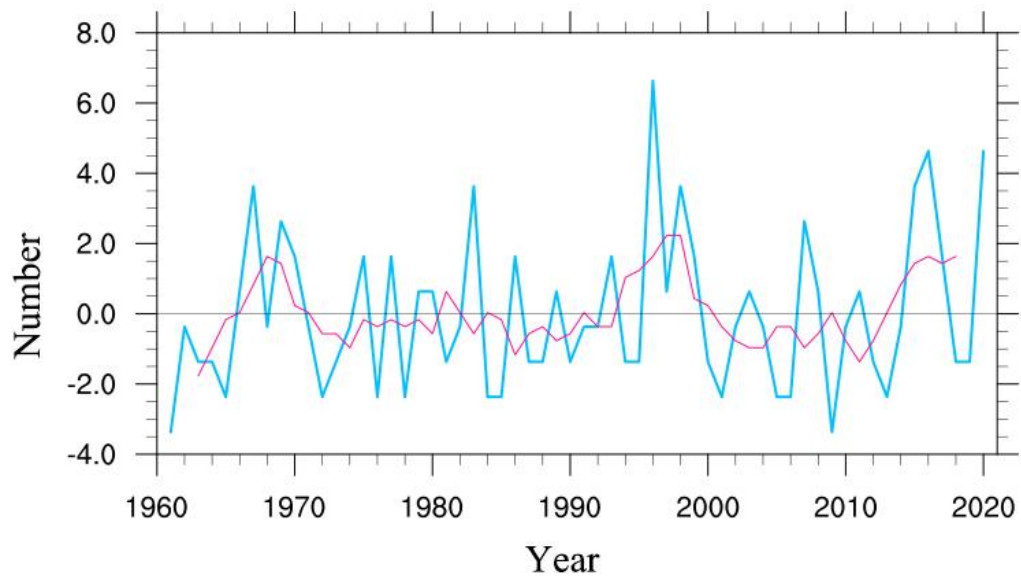
226 Fig 4. Correlation between maximum intensity (units: 10^{-5} s^{-1}) and maximum radius
 227 (units: km) (a) and their difference value of track step during the development of the
 228 Jianghuai cyclone in the Meiyu period (b).

229 The frequency of Jianghuai cyclone occurrence refers to the total number of
 230 cyclones is characterized by quasi-periodic variation (Figure 5). The shaded area in the
 231 figure indicates that the 95% confidence interval according to Student's t-test is passed.
 232 Strong 2~4 years quasiperiodic variation is observed for 1980-1990 and 1990-2000.
 233 After 2000, the quasiperiodic change in cyclones is approximately 4~5 years. This
 234 change period corresponds to the period of abnormal change in Meiyu. Chen et al.
 235 (2019) pointed out that 3~4 years of quasiperiodic change is the main component of
 236 abnormal changes in Meiyu when studying the quasiperiodic change in Meiyu. This
 237 quasi-periodic variation component is mainly influenced by the Indian Ocean dipole
 238 (IOD) (Liang et al., 2018). During the positive phase of the IOD, the strong warming
 239 of the Indian Ocean triggers a strong Indian monsoon. This leads to a strengthening of
 240 the WPSH and an increase in precipitation in southern China. The southwesterly low-
 241 level jet, which are enhanced by the positive IOD, also provide water vapor and warm
 242 advection to generate favorable conditions for the development of the Jianghuai cyclone.



243 Fig 5. Periodic wavelet analysis diagram of Jianghuai cyclones during the Meiyu period
 244 from 1961 to 2020 (units: year) (shadow indicates passing the 95% confidence interval
 245 according to Student's t-test).

246 Jianghuai cyclones are not only characterized by quasi-periodic variation but also
 247 have significant decadal variation. Figure 6 shows the activity frequency anomaly and
 248 5-year sliding average of cyclones during the Meiyu period from 1961 to 2020. The
 249 frequency of cyclone activity was the highest in 1996 and the lowest in 1961 and 2009.
 250 In the long term, the frequency of cyclone activity in the middle and lower reaches of
 251 the Yangtze River with positive anomaly in 1965-1970, 1990-2000, and 2000-after, and
 252 negative anomaly in 1970-1990 and 2000-2010. The decadal variation trend of
 253 Jianghuai cyclones is similar to the decadal variation of precipitation during the Meiyu
 254 period (Chen et al., 2019). The decadal variation of precipitation during the Meiyu
 255 period with positive anomaly in 1965-1970, 1995-2000 and 2010-after, and negative
 256 anomaly in 1970-1980, 1985-1995 and 2000-2010.



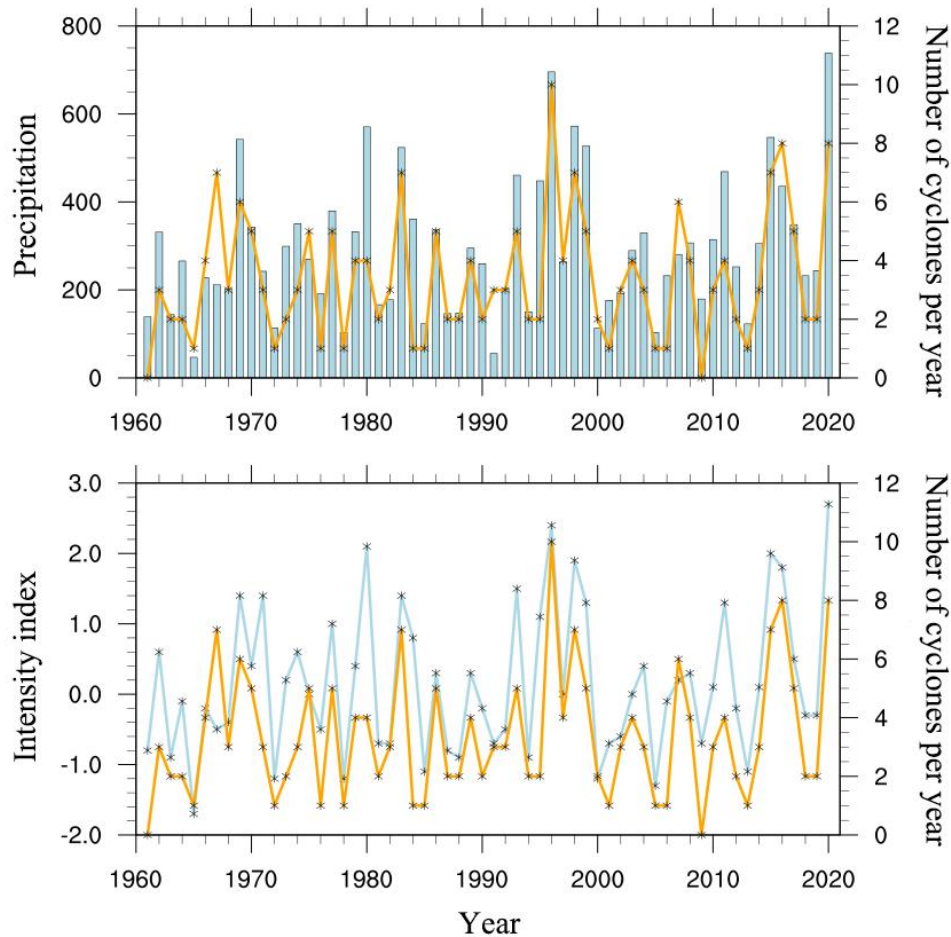
257 Fig 6. Frequency anomaly and 5-year sliding average of cyclones. The blue line shows
 258 the anomalies in the number of cyclones, and the pink line shows the 5-year sliding
 259 average of the anomalies.

260 **3.2 Linkage between cyclone activity and concurrent rainfall in the middle and**
 261 **lower reaches of the Yangtze River.**

262 The Jianghuai cyclones are mainly active in the middle and lower reaches of the
 263 Yangtze River (Huang et al., 2019; Li et al., 2002). Under the influence of the
 264 strengthening westward extension of the WPSH during the Meiyu period, the Jianghuai
 265 cyclones are restricted from entering the sea to some extent (Qin et al., 2015; Wu et al.,
 266 2020b). They form rainstorms in the middle and lower reaches of the Yangtze River and
 267 the coastal areas. A large part of the precipitation in the Meiyu period comes from
 268 cyclone precipitation (Zhang et al., 2018). The intensity of Meiyu is usually expressed
 269 by the Meiyu intensity index. The intensity of precipitation is affected not only by
 270 precipitation but also by the number of precipitation days in the Meiyu period. Both
 271 jointly determine the intensity of Meiyu in that year.

272 The time-series plots of the number of cyclones related to precipitation and the
 273 intensity index during the Meiyu period from 1961 to 2020 are given in Figure 7a and
 274 7b. We found that the number of cyclones has a positive correlation coefficient of 0.77

275 with precipitation in the Meiyu period passing the 99% confidence interval according
 276 to the student's t-test. The number of cyclones was also positively correlated with the
 277 Meiyu intensity index, with a correlation index of 0.76 passing the 99% confidence
 278 interval according to the student's t-test.



279 Fig 7. (a) Changes in precipitation (blue bar chart) (unit: mm/day) and the number of
 280 cyclones (orange line); (b) intensity index (blue line) and the number of cyclones
 281 (orange line) in the Meiyu period from 1961 to 2020.

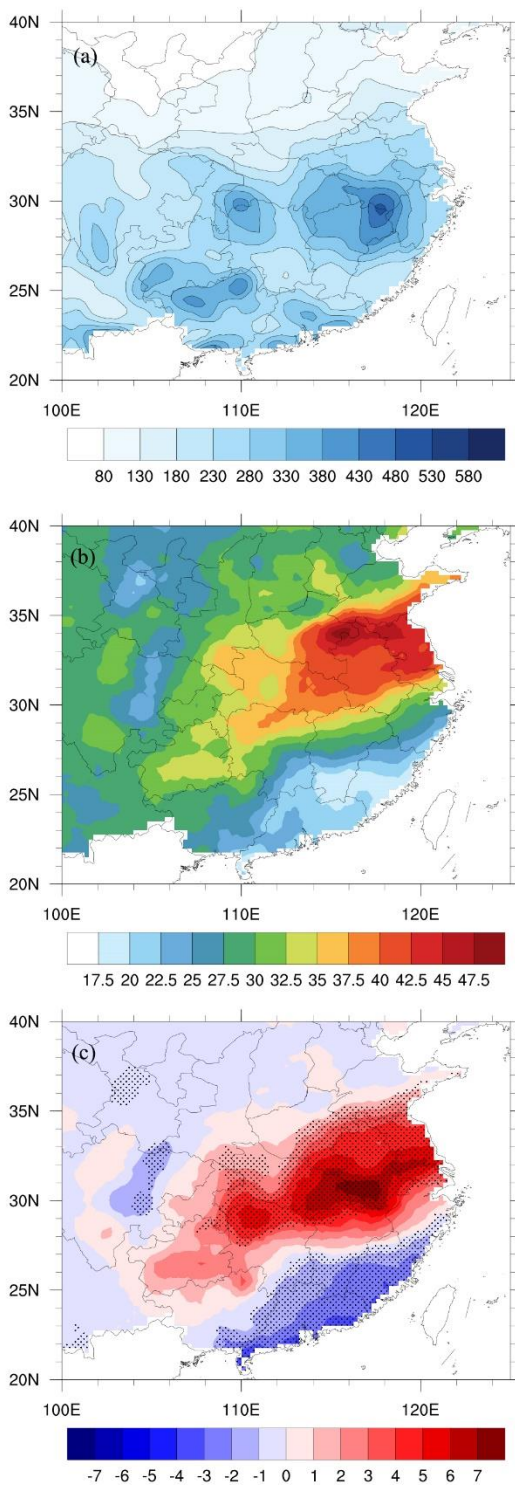
282 Figure 8a shows the spatial distribution of mean annual total precipitation during
 283 the Meiyu period from 1961 to 2020. The areas with large precipitation values in the
 284 middle and lower reaches of the Yangtze River are mainly located in the Dabie
 285 Mountains of Anhui Province, the northern part of Jiangxi Province, the eastern part of
 286 Hubei Province and the western part of Hubei Province. The maximum annual average

287 precipitation during the Meiyu period in southern Anhui can even exceed 480 mm. The
288 occurrence of large precipitation areas during the Meiyu period is closely related to the
289 topography of the region (Wu et al., 2023).

290 If precipitation and Jianghuai cyclone activity existed on the same day during the
291 Meiyu period, we defined that day as a Jianghuai cyclone precipitation day. Figure 8b
292 shows the spatial distribution of the proportion of cyclone precipitation relative to total
293 precipitation during the Meiyu period. As shown in the figure, the main areas affected
294 by cyclone precipitation are the middle and lower reaches of the Yangtze River. The
295 Huaihe River basin in northern Anhui Province is the most affected area. The cyclone
296 precipitation in the Huaihe River basin accounts for more than 47% of the total
297 precipitation during the Meiyu period, while the cyclone-influenced precipitation in
298 other areas accounts for more than 35% of the total precipitation. In general, the degree
299 of cyclone-influenced precipitation in the middle and lower reaches of the Yangtze
300 River shows an east–west band distribution and a gradual decrease from coastal to
301 inland areas. This indicates that the distribution of the large-value area and the
302 characteristics of the band distribution are related to the northeast and eastward tracks
303 of the Jianghuai cyclone. Its precipitation capacity gradually increases with the
304 development of cyclone movement.

305 **Figure 8c shows the spatial distribution of the daily mean precipitation anomaly**
306 **of the Jianghuai cyclone. The shaded part indicates that the 95% confidence interval is**
307 **passed according to Student's t-test. The anomalies mentioned here and later are the**
308 **difference between the daily mean value of meteorological elements of the selected**
309 **Jianghuai cyclones and the corresponding daily mean value of sixty years.** When the
310 Jianghuai cyclone is active, the middle and lower reaches of the Yangtze River to the
311 east of 108°E show an abnormal increase in precipitation. However, Fujian, Guangdong
312 and other places show an abnormal decrease. Among them, the maximum value of
313 abnormally increased precipitation can exceed 7 mm/day in areas such as southern
314 Anhui, eastern Hubei and northern Jiangxi. The large-value areas of precipitation
315 anomalies are consistent with the large-value areas of cyclone occurrence frequency

316 sources. It is inferred that the spatial distribution of precipitation anomalies has a
317 connection with the distribution of cyclone genesis locations. This phenomenon of
318 increasing and decreasing precipitation anomalies is bounded by approximately 27°N
319 and distributed north–south in the form of dipoles.



320 Fig 8. (a) Mean annual total precipitation during the Meiyu period from 1961 to 2020

321 (units: mm/year); (b) proportion of Jianghuai cyclone precipitation relative to total
322 precipitation during the Meiyu period (units: %); (c) daily mean precipitation anomaly
323 of the Jianghuai cyclone during the Meiyu period (units: mm/day) (shadow indicates
324 passing the 95% confidence interval according to Student's t-test).

325 Figure 9 shows the evolution of composite geopotential height and horizontal wind
326 anomalies for two different levels of Jianghuai cyclones from day -4 to +2 during the
327 Meiyu period. Composite geopotential height anomalies are significant at the 95%
328 confidence level based on Student's t-test. Vectors are plotted if wind anomalies are
329 significant at the 95% confidence level based on Student's t-test in at least one direction.

330 Day 0 is the day on which the cyclone first appears in the specified area. Most
331 areas of the lower and middle troposphere (850 hPa) in the middle and lower Yangtze
332 River on day 0 are covered by significant negative geopotential height anomalies with
333 peak magnitudes greater than -11 gpm. There is a significant positive geopotential
334 height anomaly with a peak magnitude of over 13 gpm on the southeast side of the
335 negative geopotential height anomaly. These anomalies form meridional dipole
336 structures in the middle and lower troposphere geopotential height field. The southwest
337 wind anomaly is significant in the middle and lower reaches of the Yangtze River. The
338 south of Anhui Province and the north of Jiangxi Province are between the positive
339 geopotential height anomaly and negative geopotential height anomaly. The existence
340 of these anomalies indicates the enhancement of southwest rapids and the strengthening
341 of the WPSH. The negative geopotential height anomalies at 500 hPa height on day 0
342 are mainly in Mongolia, Shanxi and other places. Strong southwest wind anomalies
343 exist between the positive and negative geopotential height anomalies. The negative
344 geopotential height anomalies in the Mongolian region exceed -7 gpm.

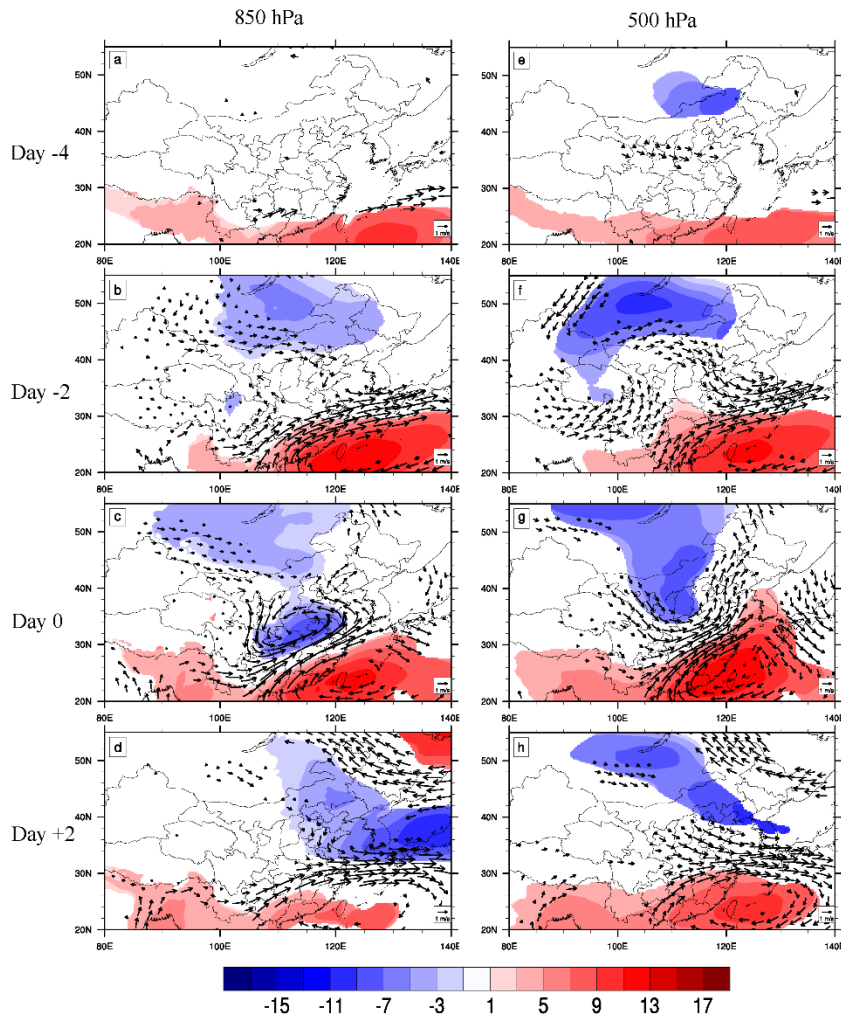
345 The negative geopotential height anomalies on all three isobaric surfaces can be
346 traced back to Mongolia, Inner Mongolia and part of Northeast China on day -2.
347 Negative geopotential height anomalies at 500 hPa can be traced to day -4. On day -4,
348 significant southwestern wind anomalies exist in southwestern Hunan at 850 hPa.

349 Significant northwest wind anomalies exist in the Yellow River basin of China at 500
350 hPa. By day -2, the negative geopotential height anomalies in Mongolia, Inner
351 Mongolia and some northeastern areas are enhanced for all three isobars. The positive
352 geopotential height anomalies of the WPSH are enhanced and extend northward to the
353 southern part of the middle and lower reaches of the Yangtze River. There are significant
354 southwest wind anomalies at the three isobaric surfaces in the south of the middle and
355 lower reaches of the Yangtze River, while there are significant northwest wind
356 anomalies at 500 hPa in the north of Anhui Province and Jiangsu Province. The negative
357 geopotential height anomalies on the three isobaric surfaces move eastward with the
358 formation and development of Jianghuai cyclones. On day +2, the lower reaches of the
359 Yangtze River are mainly affected by the combined action of anomalous southwest
360 winds and northwest winds. The positive geopotential height anomaly of the WPSH is
361 weakened.

362 Therefore, the abnormal precipitation caused by the Jianghuai cyclone mainly
363 comes from the abnormal southwest winds and the strengthening of the WPSH. The
364 enhanced southwesterly low-level jet provides sufficient warm and moist air for the
365 formation of cyclones and promotes the eastward migration of cyclones after formation.
366 Liu et al. (2020) and Zhao et al. (2021) studied the causes of the super strong Meiyu
367 year in 2020, mentioned that the WPSH is unusually strong and westward accompanied
368 by an abnormal increase in precipitation. Liu et al. (2020) found that the enhanced the
369 southwesterly low-level jet stream is conducive to the development of vertical
370 movement in the middle and low levels, which provides the necessary dynamic
371 conditions for the formation of sustained precipitation during the Meiyu in 2020.

372 Cold air activity is one of the important factors for the formation of heavy
373 precipitation, which can promote the convergence and uplift of low level necessary for
374 heavy precipitation (Liu et al., 2020). The enhanced negative geopotential anomaly
375 over Mongolia provides cold and dry air brought by the westerly jet for cyclone
376 development. The increasing frequency of cyclones over the Yangtze River and Huaihe
377 River leads to the abnormal increase in precipitation in the middle and lower reaches of

378 the Yangtze River during the Meiyu period. However, due to the strengthening of the
 379 WPSH, the southern part of China is controlled by the abnormal positive geopotential
 380 height, and the precipitation decreases. Zhao et al. (2021) also found that when the
 381 WPSH enhanced, there was a decrease in precipitation in South China.



382 Fig 9. Evolution of composite geopotential height anomalies (shading; units: gpm) and
 383 horizontal wind anomalies (units: m/s) on the 850 hPa and 500 hPa isobaric surfaces
 384 for day -4 (a, e), day -2 (b, f), day 0 (c, g) and day +2 (d, h) for the 202 selected
 385 Jianghuai cyclones. Shading indicates that composite geopotential height anomalies are
 386 significant at the 95% confidence level based on Student's t-test. Vectors are plotted if
 387 wind anomalies are significant at the 95% confidence level based on Student's t-test in
 388 at least one direction.

389 Figure 10 shows the climatic distribution of water vapor flux and water vapor flux

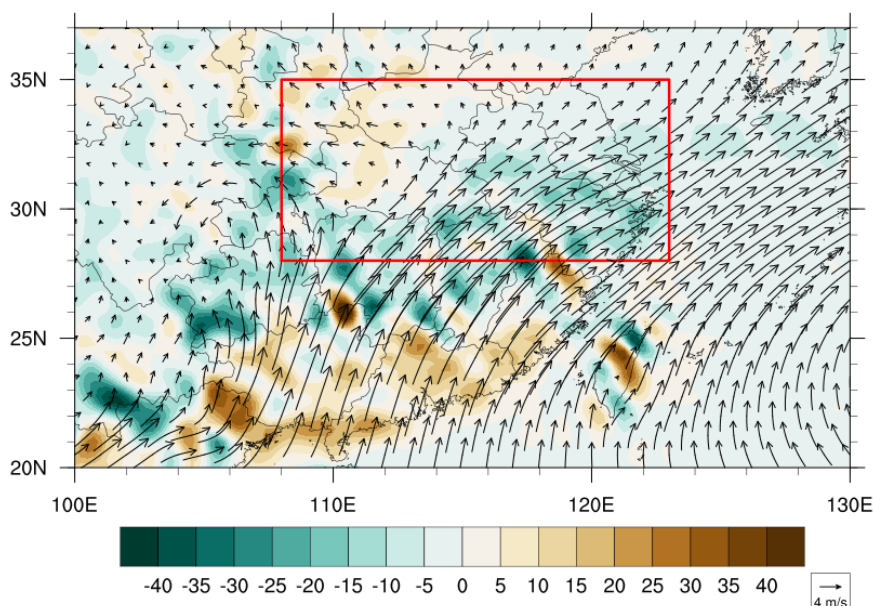
390 divergence at 850 hPa of Jianghuai cyclone during the Meiyu period. The water vapor
391 involved in the precipitation process of the Jianghuai cyclone mainly comes from the
392 water vapor brought by the southwesterly low-level jet of the summer monsoon in the
393 low-latitude area. During Jianghuai cyclone development, the middle and lower reaches
394 of the Yangtze River are mostly in the water vapor convergence area, which is
395 conducive to the generation of precipitation (Chen et al., 2020).

396 Figure 11 shows the distribution of water vapor flux anomalies and water vapor
397 flux divergence anomalies at the pressure level of 850 hPa during the Jianghuai cyclone
398 from day -1 to day +2. The color field and wind vector arrows in the figure both passed
399 the 95% significance according to Student's t-test. On day -1, a significant water vapor
400 convergence anomaly and water vapor transport in the southwest direction appear in
401 eastern and western Hubei Province. In the middle reaches of the Yangtze River, there
402 is a north-south distribution of water vapor convergence and divergence dipoles. The
403 anomalies of water vapor flux and water vapor flux dispersion are mainly concentrated
404 on day 0. There is significant anomalous water vapor convergence up to $-1 \text{ g} \cdot \text{cm}^{-2} \cdot \text{hPa}^{-1}$
405 in eastern Hubei Province, Anhui Province and Jiangsu Province on day 0. Anomalous
406 water vapor dispersion exists in the middle and lower reaches of the Yangtze River. On
407 the day +1 of cyclone development, the spatial distribution of water vapor convergence
408 and divergence shifted to an east-west direction. The anomalous transport of water
409 vapor flux also changes to an east-west direction. On day +2, with the development of
410 the cyclone's eastward movement, only the southern part of Jiangsu Province and the
411 northern part of Zhejiang Province have abnormal water vapor flux in the eastward
412 direction. The precipitation in the area begins to gradually weaken at this time.

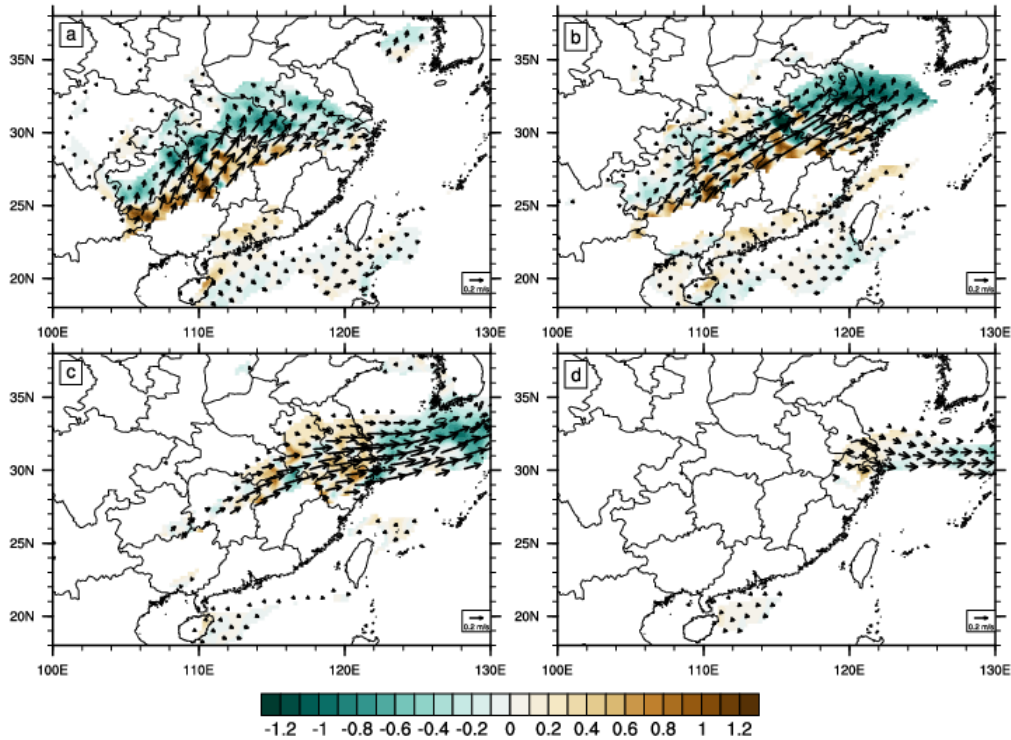
413 From day -1 to day 0, the abnormal water vapor flux and water vapor flux
414 divergence configuration make the warm and wet air in the low-latitude area transport
415 to the middle and lower reaches of the Yangtze River. The abnormal water vapor flux
416 has a negative value, water vapor convergence occurs, local water vapor volume
417 increases, and finally, the precipitation in the region increases. Liu et al. (2020) studied
418 the strong rainfall in 2020, they found that there was an enhanced water vapor transport,

419 and the repeated occurrence of convergence movement. Both of them caused the
420 precipitation time to increase in the Jianghuai River basin.

421 In contrast, the anomaly of water vapor flux in southern Guangdong and other
422 regions is divergent. This leads to a decrease in local water vapor volume and
423 precipitation in this region. These results indicate that the variations in water vapor flux
424 and divergence related to cyclones are mainly from warm and wet air transported from
425 low latitudes to the middle and lower reaches of the Yangtze River. Therefore, there is
426 a positive correlation between cyclone activity and precipitation in the middle and
427 lower reaches of the Yangtze River.



428 Fig 10. Distribution of 850 hPa daily mean water vapor flux (unit: $\text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}$) and
429 water vapor flux divergence (unit: $10^{-8}\text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}\cdot\text{s}^{-1}$) of cyclones over the Yangtze
430 and Huaihe rivers during 1961-2020 (color diagram shows water vapor flux divergence,
431 and vector diagram shows water vapor flux).



432 Fig 11. Distribution of the 850 hPa daily mean water vapor flux anomaly (unit: $\text{g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}$) and water vapor flux divergence anomaly (unit: $10^{-8}\text{ g}\cdot\text{cm}^{-2}\cdot\text{hPa}^{-1}\cdot\text{s}^{-1}$) of
 433 cyclones over the Yangtze and Huaihe rivers during 1961-2020 (color diagram shows
 434 water vapor flux divergence, and vector diagram shows water vapor flux). The colored
 435 region passed the 95% confidence interval according to Student's t-test. If the vapor
 436 flux anomaly is significant at the 95% confidence level for Student's t-test in at least
 437 one direction (zonal or meridian), the vector is plotted.
 438

439 4. Summary and discussion

440 Based on ERA5 reanalysis of sea level pressure data and using the relative
 441 vorticity method to identify and track cyclones, we have examined the impacts of the
 442 climatological characteristics of Jianghuai cyclones. The linkages between cyclone
 443 activity and precipitation in the middle and lower reaches of the Yangtze River during
 444 the Meiyu period are also analyzed.

445 During the Meiyu period, Jianghuai cyclones are mainly generated at the junction
 446 of western Hubei and Chongqing Municipality, eastern Hubei Province, northern

447 Jiangxi Province, central and southern Anhui Province, and Jiangsu and Zhejiang
448 provinces. These cyclones develop and move to the sea in the east or northeast direction.
449 There is a positive correlation between the maximum intensity and maximum radius of
450 Jianghuai cyclones. Its occurrence frequency not only has the characteristics of
451 multicycle variation but also has obvious interdecadal variation, which has a good
452 correspondence with the periodic and interdecadal variation in precipitation in the
453 Meiyu period.

454 There is a positive correlation between the frequency of cyclone activity and
455 precipitation in the Meiyu period. The frequency of Jianghuai cyclone activity is high
456 in the years with strong Meiyu rainfall and low in the years with weak Meiyu rainfall.
457 The percentage of precipitation affected by Jianghuai cyclone activity in the middle and
458 lower reaches of the Yangtze River can reach up to 47%. The spatial distribution is in
459 the shape of an east–west belt, and the degree of influence gradually decreases from the
460 coast to the interior. When the Jianghuai cyclone is active, the precipitation increases
461 abnormally in the middle and lower reaches of the Yangtze River east of 108°E.
462 Precipitation decreases abnormally in Fujian Province and Guangdong Province. The
463 spatial distribution of precipitation anomalies is related to the genesis locations of
464 cyclone frequency, and the positive and negative anomalies are distributed north–south
465 in the form of dipoles based on the latitude line at approximately 27°N as the boundary.

466 The geopotential height anomaly field and the horizontal wind vector anomaly
467 field of the Jianghuai cyclones during the Meiyu period are synthesized and analyzed.
468 There is an enhanced positive geopotential height anomaly of the WPSH during cyclone
469 activity. The negative geopotential altitude anomaly of Mongolia and the abnormal
470 southwesterly low-level jet are enhanced. All of these factors lead to an increase in
471 precipitation in the middle and lower reaches of the Yangtze River. The abnormal
472 leading signal of the negative geopotential height in Mongolia can be traced to day -2
473 of the cyclone activity, and the signal can be traced to day -4 at 500 hPa. From day -2
474 to day 0 of cyclone activity, the abnormal distribution of water vapor flux and water
475 vapor flux divergence cause the warm and wet air at the low latitudes to be transported

476 to the middle and lower reaches of the Yangtze River. They promote the generation and
477 development of cyclones and increase precipitation in the middle and lower reaches of
478 the Yangtze River.

479 We explored the cyclone characteristics and study emphasizes the link between
480 cyclone activity and Yangtze River precipitation. Spatially, abnormal precipitation
481 patterns are identified, tracing the evolution of geopotential height anomalies and water
482 vapor flux. But the specific mechanism by which the southwesterly low-level jet affects
483 cyclones during the Meiyu period is not clear enough. Zhang et al. (2018) suggest that
484 the strengthening of the southwesterly low-level jet will lead to the development of a
485 mesoscale low-pressure disturbance near the Meiyu Front and the occurrence of
486 extreme precipitation. Liu et al. (2020) found that the strengthening of the
487 southwesterly low-level made the southerly meridional strong gradient zone on the
488 north side of the meridional wind maximum center move northward in the low-level
489 dynamic conditions of the rainstorm process during Meiyu. How the southwesterly low-
490 level stream influences the development of physical factors to promote the formation
491 of Jianghuai cyclones remains to be considered and analyzed. Zhao et al. (2010) found
492 that the causes of Jianghuai cyclones with different intensities were different through a
493 case study. Therefore, we think it is also necessary to consider the difference in the
494 influence of different intensities of Jianghuai cyclones on precipitation. These problems
495 need further analysis and research.

496 **Competing interests**

497 The contact author has declared that none of the authors has any competing interests.

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