

Reviewer#1.

The authors are most grateful for your comments. We have followed your suggestions and revised the manuscript accordingly in many places. Please, find our responses below.

This manuscript considers two-dimensional simulations of internal solitary waves propagating into a region with modelled ice cover. The ice cover is modelled as not moving, and as represented by a piecewise constant value (or perhaps as smoothed) of a drag parameter (which is varied to some degree). The former is sensible, while the latter is perhaps a necessary choice for the model employed. The manuscript is interesting, and the figures provide useful information. The text needs a thorough reading for technical English (if necessary I can provide a list of suggestions when the scientific review is completed). I feel that a version of this manuscript can appear in NPG, but there are some necessary changes/improvements. I enumerate these below, but as an overall comment I would say the results need to discuss the new results in terms of existing literature and the second set of experiments needs a more complete analysis and discussion.

I note that for many of the Yes/No questions the journal asks, the manuscript falls between a strict Yes or NO.

Answer: We have revised and expanded the discussion of the results of the second series of simulations by adding a comparison with known laboratory experiments and numerical calculations (see responses to comments 3) and 7))

1) Self-citation: Proofread to ensure that when a topic is introduced, e.g. shoaling of elevation, the references provided are more than just those of the authors (in particular for numerical studies). This is not just an issue of a longer bibliography. There are quite a few papers I would consider relevant listed, but they tend to appear as lists in the Introduction, and the opportunity to discuss the context of the numerical simulations in terms of these papers is missed.

Answer: We expand our discussion of the results with comparisons with published works.

L. 103 “These results demonstrate a weak effect of free surface on ISW dynamics in considered cases which made it possible in this problem to replace the conditions on the free surface with conditions on the rigid lid. Note that in laboratory experiments (Carr et al., 2008; Luzzatto-Fegiz and Helfrich, 2014) the influence of a free surface on the stability of waves with a trapped core was shown. This effect has been interpreted as the influence of surfactants essential in laboratory-scale processes, however, these Marangoni effects have a negligible impact on the interior of full-scale oceanic waves (Luzzatto-Fegiz and Helfrich, 2014)”

See also the answers to comment 7.

We have added literature:

McPhee, M. G., G. A. Maykut, and J. H. Morison: Dynamics and thermodynamics of the ice/upper ocean system in the marginal ice zone of the Greenland Sea, *J. Geophys. Res.*, 92, 7017-7031, 1987.

Carr, M., Fructus, D., Grue, J., Jensen, A. & Davies, P. A.: Convectively induced shear

instability in large amplitude internal solitary waves. *Phys. Fluids* 20 (12), 126601, 2008.

Luzzatto-Fegiz P and Helfrich K: Laboratory experiments and simulations for solitary waves with trapped cores *J. Fluid Mech.* 757 354-380, 2014.

2) *Details of the numerical model. The basic idea of the top boundary condition is introduced, but right now an interested reader could not reproduce the results on their own. What needs to be done to implement the conditions? What complications result (e.g. in the pressure problem)? Can the drag coefficients really just be discontinuous?*

Answer: The text has been added to clarify numerical model details.

L. 86 “The pressure zero gradient boundary condition was imposed on all boundaries. At the corner of the underwater step, this condition is violated. However, numerical experiments for different resolutions have shown that this problem does not occur at simulated fields of velocity and density.”

L. 111 “The quasi-z-level coordinate system (Maderich et al., 2012) was used to describe this step-like ice layer.”

3) *The model resolution deserves comment. What can one expect to see/resolve (certainly I believe wave fissioning is accurately represented); what do we miss (I think the details of the high shear region near the ice cannot be accurately represented). The Carr et al corrugation paper gives details of the interaction with a no slip boundary layer, and hence provides an easy contrast.*

Answer: We applied the Reynolds averaged equations system closed by a simple subgrid model of turbulence to this real ocean scale problem. The standard boundary conditions for turbulent shear stress under rough ice surfaces were applied. Of course, this model cannot resolve small-scale structures in the turbulent flow. However, it can describe processes of ISW transformation and breaking. The comparison with the simulation of ISW transformation on a single ice keel (Zhang et al., 2022) showed very similar wave evolution despite differences in the parameterization of turbulent mixing (see answer to comment 7). The direct comparison of our simulations with laboratory and numerical experiments by Carr et al. (2010) is difficult since this work considers laboratory-scale processes and, in addition, the incident wave has a length much greater than the length of the corrugated bottom relief elements. Nevertheless, a certain similarity of flow processes was observed in both cases (see answer to comment 7).

4) *I'd prefer “smoothed step” to “step”. It would also help to state for the reader how many points there are across the changing part of the tang-based step.*

Answer: The first series of experiments was performed using a quasi-z-coordinate which allowed us to describe a step-like ice layer without any smoothing. The text has been added:

L. 111 “The quasi-z-level coordinate system (Maderich et al., 2012) was used to describe this step-like ice layer.”

In the second series of experiments, the sigma coordinate system was used to accurately describe flow around the keels placed under a relatively thin ice layer.

L. 118 In the second series of experiments (see Table 2), 12 runs (K1-K12) were performed using a sigma-system of coordinates, which allowed for accurately describing flow around the keel.

5) It would be good to indicate the integration region for the energetics calculations on the appropriate panel of Fig 4. Similarly, the spacing of the equations in the system 6 could be improved (perhaps this is due to them lying at the bottom of the page, and they will likely move in a final version of the manuscript).

Answer: Thank you for the suggestion. We have made changes in the fig 4 and corresponding caption.

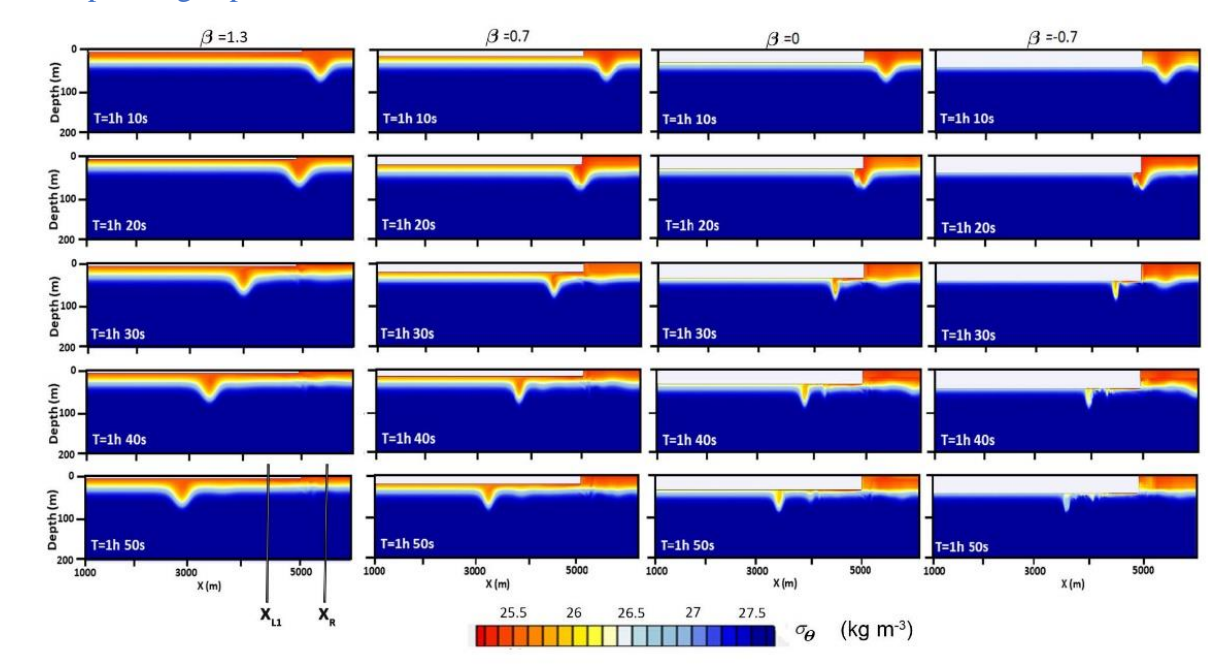


Figure 4. The snapshots of the density field for incident ISW waves with an amplitude of 15 m passing under the ice cover with different drafts. The integration region for the energetics calculations between X_{L1} and X_R is shown.

6) Presumably the g in equation (8) is a reduced gravity? Otherwise I cannot see how a supercritical regime is reached.

Answer: Thank you! We corrected g to g' .

$$L. 178 \quad Fr^2 = \frac{U_1^2}{g'h_1(x)} + \frac{U_2^2}{g'h_2(x)}$$

where U_1 and U_2 are the layer-averaged velocities in each layer, $g' = g\Delta\rho / \rho_0$, where g is the gravity acceleration, $\Delta\rho$ and ρ_0 are the density difference between upper and lower layers and undisturbed density of fluid, respectively, g is the gravity acceleration.

7) I found the second set of experiments, the ice keels, to be a bit tougher to digest, likely due to its brevity. I have questions about the way the boundary layer is parametrized (wouldn't there be more drag over the downstream slope where Carr at all predicted "local hydraulic")

phenomena?). The discussion of Fig 7 seems incomplete (what are the curves shown, what are the details of what is presumably a fitting process?).

Answer: We extended the discussion of the second series of experiments including Fig. 7 and compared the results of our simulations with two limiting cases: interaction ISW with a single keel (Zhang et al., 2022), and ISW propagation over a corrugated bottom when the bottom element length was much less than the ISW wavelength (Carr et al. 2010). The text has been added accordingly:

L. 212 “In the limiting case of the interaction of ISW with a single keel (Zhang et al., 2022b), the maximum energy dissipation was about 25% which is somewhat less than in our calculations, but we need to keep in mind the differences in the calculation parameters and turbulence parametrization. Zhang et al. (2022b) used constant eddy coefficients whereas in our study the model of turbulence was used where eddy coefficients vary in space and time.”

L. 215 “To characterize the dependence of ΔE_{tot} on keel height and distance between keels we introduced parameter $\mu = (h_{ice} + h_k) / L_k$. As seen in Fig. 7 this dependence can be approximated by logarithmic curves. The energy loss ΔE_{tot} increases with the decrease of distance between keels or an increase of keel height. The level of ΔE_{tot} is highest for β values near zero. As seen in Fig. 5a, this range of β corresponds to the regime of strong interaction (III). Energy loss in this regime is maximal, both in the case of the ridged underside of the ice and in the case of smooth ice surfaces with the same parameter β . When β values increase, the dependence of energy loss on the μ and distance between the keels decreases. For $\beta=0.8$ is on the boundary between regime (II) moderate and (I) weak interaction distance between the keels is no longer significant.”

L. 225 “In another limiting case ISW of elevation propagates over a corrugated bottom when the bottom element length was much less than the ISW wavelength (Carr et al., 2010) a comparison with ISW propagated under an ensemble of ice keels of horizontal scales greater than ISW wave length was not straightforward. In addition, Reynolds equations with turbulent closure describe real-scale processes in the ocean, in contrast to laboratory scales in (Carr et al., 2010). Unlike (Carr et al., 2010) we cannot describe in detail the instant spatial-temporal dynamics of high shear region near the ice. However, the Fig. 6b showed wave-induced currents over the keels, their interaction with the apex of the keels and a sequence of lee vortices formed as a result of such interaction (see Fig. 6b T=1h 35 m, T=1h 41m). Similarly to (Carr et al., 2010) the vortices developed after the main wave passed over the keel (see Fig. 6b at T=1h 44 m, T=1h 45m) resulting in deformation of the overlying pycnocline and, in some instances, significant vertical mixing.”

8) *I agree with the comments in line 220. At the same time, I think there have been simulations in the literature of related heat-salt phenomena. It least to point to these as a start of relevant studies.*

Answer: The text was added accordingly:

L. 254 “The next step could be an explicit representation of heat and salt fluxes between the ice cover due to the ISW interaction with the ridged ice, e.g. following flux parametrization by McPhee et al., (1987).”

