

On the potential use of highly oxygenated organic molecules (HOM) as indicators for ozone formation sensitivity

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Abstract. Ozone (O₃), an important and ubiquitous trace gas, protects lives from harm of solar ultraviolet (UV) radiation in the stratosphere but is toxic to living organisms in the troposphere. Additionally, tropospheric O₃ is a key oxidant, and source of other oxidants (e.g., OH and NO₃ radicals) for various volatile organic compounds (VOC). Recently, highly oxygenated organic molecules (HOM) were identified as a new compound group formed from oxidation of many VOC, making up a significant source of secondary organic aerosol (SOA). The pathways forming HOM from VOC involve autoxidation of peroxy radicals (RO₂), formed ubiquitously in many VOC oxidation reactions. The main sink for RO₂ is bimolecular reactions with other radicals, HO₂, NO or other RO₂, and this largely determines the structure of the end products. Organic nitrates form solely from RO₂ + NO reactions while accretion products (“dimers”) solely from RO₂ + RO₂ reactions. The RO₂ + NO reaction also converts NO into NO₂, making it a net source for O₃ through NO₂ photolysis.

There is a highly nonlinear relationship between O₃, NO_x, and VOC. Understanding the O₃ formation sensitivity to changes in VOC and NO_x is crucial for making optimal mitigation policies to control O₃ concentrations. However, determining the specific O₃ formation regimes (either VOC- or NO_x-limited) remains challenging in diverse environmental conditions. In this work we assessed whether HOM measurements can function as a real-time indicator for the O₃ formation sensitivity based on the hypothesis that HOM compositions can describe the relative importance of NO as a terminator for RO₂. Given the fast formation and short lifetimes of the low-volatile HOM (timescale of minutes), they describe the instantaneous chemical regime of the atmosphere. In this work, we conducted a series of monoterpene oxidation experiments in our chamber while varying the concentrations of NO_x and VOC under different NO₂ photolysis rates. We also measured the relative concentrations of HOM of different types (dimers, nitrate-containing monomers, and non-nitrate monomers) and used ratios between these to estimate the O₃ formation sensitivity. We find that for this simple system, the O₃ sensitivity could be described very well based on the HOM measurements. Future work will focus on determining to what extent this approach can be applied in more complex atmospheric environments. Ambient measurements of HOM have become increasingly common during the last decade, and therefore we expect that there already are a large amount groups with available data for testing this approach.

1 Introduction

Ozone (O₃), as a key trace gas in the atmosphere, directly and indirectly affects human lives, and it plays diametrically opposed roles in the troposphere ("bad" ozone) and stratosphere ("good" ozone) (Sandermann, 1996; Staehelin et al., 2001; Seviour, 2022). The formation and depletion of O₃ have been investigated over the past decades (Chapman, 1930; Crutzen, 1970, 1971; Stolarski and Cicerone, 1974; Tiao et al., 1975; Dodge, 1977). Atmospheric O₃ is almost entirely produced through the reaction between atomic oxygen (O³P) and molecular oxygen (O₂) (Wang et al., 2017). In the stratosphere, the O³P source is O₂ photolysis with ultraviolet (UV) wavelengths below 240 nm (Chapman, 1930). Stratospheric O₃, which constitutes approximately 90% of Earth's atmospheric O₃, plays a crucial role in absorbing UV radiation in the UVB band (280–315 nm), protecting organisms on the ground from the harm of UV radiation (Gruijl and Leun, 2000; Seinfeld and Pandis, 2016). Although certain man-made substances, such as chlorofluorocarbons, were found to be responsible for significant depletion of stratospheric O₃, the implementation of the 1987 Montreal Protocol and its subsequent amendments has contributed to the recovery of the stratospheric O₃ layer (Seinfeld and Pandis, 2016; Chipperfield et al., 2017).

In the troposphere, in addition to its role as a greenhouse gas (Ehhalt et al., 2001), O₃ serves as a secondary air pollutant due to its detrimental impacts and indirect emissions (Seinfeld and Pandis, 2016; Nuvolone et al., 2018). Not only is it toxic, but it also participates in chemical reactions that lead to the formation of other harmful molecules (Nuvolone et al., 2018). In contrast to stratospheric O₃ in the troposphere, the source of O³P is NO₂ photolysis at wavelengths less than 420 nm (Madronich et al., 1983). However, the net formation of tropospheric O₃ occurs through chemical reactions involving nitrogen oxides (NO_x=NO+NO₂) and various volatile organic compounds (VOC) in the presence of UV light (Lelieveld and Dentener, 2000). In an ideal "clean" system without any VOC, once O₃ is formed, it readily converts NO back to NO₂ by reacting with NO, resulting in a null cycle as shown below (also illustrated as "NO_x cycle" in Fig. A1):



When VOC are present, they will be oxidized to form peroxy radicals (RO₂) by atmospheric oxidants, such as O₃ and OH (Atkinson and Arey, 2003):



The RO₂ radicals can thus replace O₃ in converting NO into NO₂ (R4a). Some fraction of RO₂ + NO reactions will also lead to the formation of organic nitrates, RONO₂ (R4b) (Atkinson and Arey, 2003):



In summary, the presence of RO₂ radicals, supplied by VOC, can perturb the "NO_x cycle", leading to the net increase of O₃ (Fig. A1). As one of the most characteristic features of the photochemical smog episodes in many cities (Tiao et al., 1975; Tang

et al., 1995; Dickerson et al., 1997), this O₃ formation process (R1–R4a, Fig. A1), known as O₃-NO_x-VOC sensitivity or O₃ formation sensitivity, has been investigated since the last century (Haagen-Smit et al., 1953; Kinoshita, 1982). The Empirical Kinetic Modeling Approach (EKMA) curve, namely O₃ isopleths, was proposed by Dodge (1977) and has been widely used to visually study the O₃ formation sensitivity (Liu and Shi, 2021). The O₃ isopleths reveal a highly nonlinear response of O₃ to its precursors, demonstrating that increasing or decreasing either VOC or NO_x does not consistently exhibit the same behavior regarding their impact on O₃ formation (Meyer Jr et al., 1977; Harris et al., 1982; Sillman et al., 1990). The outcome is dependent on the relative concentrations of VOC (as a surrogate for RO₂) and NO_x, leading to the division of the formation area into NO_x-limited and VOC-limited regimes (Sillman, 1999; Melkonyan and Kuttler, 2012). In the NO_x-limited regime, the concentration of O₃ generally increases with an increase in NO_x, while its response to VOC changes remains relatively small. This is because the supply of RO₂ species from VOC is abundant and R4a is limited by NO (the "NO_x cycle" is saturated in Fig. A1). Conversely, in the VOC-limited regime, an increase in VOC concentration generally leads to an increase in O₃, whereas an increase in NO_x may even result in a decrease in O₃ levels. This is because NO_x is in excess compared to RO₂. Moreover, very high levels of NO_x can directly titrate O₃ (Sillman, 1999) or consume OH radicals (Atkinson et al., 2004), thereby reducing the supply of RO₂ species (R3) and promoting the reaction R2, which results in a decrease in the O₃ concentration. These effects were demonstrated by the amplified O₃ pollution in cities during COVID-19 lockdown, when NO_x emissions dropped dramatically (Sicard et al., 2020).

To mitigate the uncertainties associated with photochemical models and efficiently determine the O₃ formation sensitivity, various photochemical indicators have been utilized since the last century (Wang et al., 2017; Liu and Shi, 2021). For example, the O₃ production efficiency (OPE = $\frac{\Delta O_3}{\Delta NO_z}$) is defined as the number of O₃ molecules produced per molecule of NO_x before the NO_x is oxidized to more stable products, i.e., NO_z species (Trainer et al., 1993; Wang et al., 2017). Smaller values of OPE indicate the inefficiency of the "NO_x cycle" (Fig. A1), suggesting that the supply of RO₂ from VOC becomes the limiting factor. As a result, the photochemical system tends to be VOC-limited. Conversely, when OPE values are higher, the system tends to be NO_x-limited. Additionally, several modifications had been made to the OPE indicator to account for different situations, such as replacing ΔO_3 with ΔO_x (O_x=O₃+NO₂) (Kleinman et al., 2002) and replacing ΔNO_z with ΔNO_y (NO_y=NO_z+NO_x) (Wang et al., 2006). Other chemical species have also been utilized as photochemical indicators, including the ratio of H₂O₂ to HNO₃ (Hammer et al., 2002). High values of $\frac{H_2O_2}{HNO_3}$ indicate high potential for cross-reactions of two HO₂ radicals, which is associated with high $\frac{VOC}{NO_x}$ and thus indicative of the NO_x-limited regime. For a more widespread application, space-based $\frac{HCHO}{NO_2}$ measurements from global O₃ monitoring satellites have been introduced as an indicator (Martin et al., 2004), based on the fact that levels of HCHO and NO₂ in the tropospheric column are closely linked to VOC and NO_x emissions, respectively. However, all these indicators are not inherently linked to the O₃ formation process, and the corresponding threshold values depend on environmental conditions (Liu and Shi, 2021). This makes it challenging to universally apply these indicators.

During the past decade, highly oxygenated organic molecules (HOM) have been recognized as a new group of VOC oxidation products, particularly important for the formation of secondary organic aerosol (SOA) due to their fast formation and low volatilities (Ehn et al., 2014; Bianchi et al., 2019). Aerosols play a significant role in both impacting human health adversely (Kelly and Fussell, 2015) and influencing climate (Boucher et al., 2013). Formed in the atmosphere-mimicking gas

phase and containing six or more oxygen atoms (Bianchi et al., 2019), HOM are produced via RO₂ autoxidation which rapidly
90 increases their oxygen content through intramolecular H atom abstractions followed by O₂ additions (Crounse et al., 2013; Ehn
et al., 2014). Eventually, these highly oxygenated RO₂ will generally be terminated similarly as other RO₂, such as through
bimolecular reactions with NO_x, RO₂ or HO₂ radicals (i.e., R4, R5 and R6).



The O₃ formation precursors, namely NO_x and VOC, thus are intrinsically connected also to HOM formation through RO₂
chemistry. As such, if the daytime HOM distribution is dominated by organic nitrates, it suggests that the majority of RO₂ are
100 being terminated by reactions with NO (R4), thus contributing to O₃ formation. On the other hand, if we observe large amounts
of HOM dimers or non-nitrate monomers, formed from RO₂ + RO₂/HO₂ (R5 and R6), there must be a large fraction of RO₂
that do not contribute to the O₃ formation process, suggesting that increased NO_x would also lead to more O₃. In other words,
ratios of different types of HOM can function as another indicator for determining the sensitivity of O₃ formation. The situation
is complicated by several factors, including challenges in identifying which HOM might have formed from RO₂ termination
105 by RO₂ or HO₂, or by knowing if a HOM is a monomeric product of a larger VOC precursor or a dimeric product from smaller
VOC. Still, if HOM could be used even as a qualitative indicator for O₃ formation, one particular benefit would be that they
would serve as a real-time indicator. This is because both the formation (through autoxidation) and loss (through condensation
onto aerosol particles) of HOM take place on timescales of minutes or less.

In this study, our objective is to assess the viability of using the ratio of HOM dimers or non-nitrate monomers to HOM
110 organic nitrates as an indicator of O₃ formation sensitivity. We conducted a series of experiments in an atmosphere simulation
chamber (Riva et al., 2019a), focusing on the ozonolysis of α-pinene, the most abundantly emitted monoterpene (Pathak et al.,
2007). By varying the concentrations of O₃ formation precursors NO_x and α-pinene, as well as the NO₂ photolysis rate, we
explored the shift between NO_x-limited and VOC-limited regimes in the chemical system. We employed mass spectrometers
and gas monitors to measure HOM products, α-pinene, O₃, and NO_x. We also developed a simple 0-D box model to simulate
115 the concentrations of O₃ and its precursors in the chamber under different conditions. Finally, by analyzing both experimental
and model outcomes, we evaluated the potential of the HOM ratios as indicators of O₃ formation sensitivity (either VOC- or
NO_x-limited) in this system.

2 Methods

2.1 Experiments

120 The experiments were conducted in the COALA chamber, as presented by Riva et al. (2019a). The cuboid chamber is made of
fluorinated ethylene propylene (FEP) and has a volume of 2 m³ with a volume to surface area ratio of 0.2. The chamber was run
in “steady-state mode” (Peräkylä et al., 2020; Krechmer et al., 2020), meaning there was a continuous flow of air and reagents
(O₃, NO₂, and α -pinene) through the chamber. The total flow was around 55 L min⁻¹, giving an average residence time
of approximately 36 minutes ($\tau = \frac{2000 \text{ L}}{55 \text{ L min}^{-1}} \approx 36 \text{ min}$). Each stage, where the experimental conditions remain unchanged,
125 lasted at least 1.5 hours, which is approximately three times the residence time, allowing the chamber to reach a pseudo steady
state, as confirmed by the time series obtained during the experiments (e.g., Fig. 4).

Table 1. Experimental conditions. Each experiment consisted of 3–9 “stages” that corresponded to a specific time period during which the
inputs remain constant. The parameter that was varied included input O₃, α -pinene, or NO_x concentrations, as well as NO₂ photolysis rate
(J_{NO_2}). These variations are indicated in the table by multiple values or ranges in a given cell. Experiment numbers (No.) and number of
total stages per experiment are shown in the first two columns.

Experiment No.	Number of stages	Input			
		J_{NO_2} (s ⁻¹)	O ₃ (ppb)	α -pinene (ppb)	NO _x range (ppb)
1.	7	1.85×10^{-3}	10.5	30/60	0 – 21.5
2.	8	1.85×10^{-3}	15	15/45/60	0 – 21.5
3.	9	1.85×10^{-3}	22.5	10/45/60	0 – 35.2
4.	9	1.15×10^{-3}	15.5	15/45/60	0 – 44.5
5.	9	0.35×10^{-3}	15.5	15/45/60	0 – 44.5
6.	3	1.85×10^{-3}	10/15.5/22.5	60	44.5
7.	8	1.85×10^{-3}	10	30/45/60	0 – 21.5

The details of the conducted experiments are provided in Table 1. The UV LED lights (wavelength \sim 400 nm, manufactured
by LEDlightmake Inc., Shenzhen, China) (Zhao et al., 2023a) used for photolyzing NO₂ were kept on throughout each exper-
iment, while the input concentrations of the precursors NO₂, O₃, and α -pinene were varied across experiments and stages to
130 map out a wide range of different conditions. The photolysis rate was varied by using varying numbers of LED light strips
(1, 3, or 5). Experiments without VOC addition (no. Z1–Z5, Table A1) were used to evaluate the photolysis rates (given in
Table 1) for each number of light strips. We acknowledge that using alkene ozonolysis for this type of study is not ideal as O₃
also reacts with the VOC, making the determination of actual O₃ formation more complicated. The choice of this system was
partly due to our chamber not having the optimal light source for producing OH radicals, thus limiting us to O₃ oxidation, and
135 partly because the HOM spectra from this system has been studied in great detail, making the interpretation of the HOM easier.

All the input reactants, as well as the HOM products, were continuously measured online using instruments described below. The identified HOM species in this study were categorized into three groups: 1) "HOM monomers" (HOM_{Mono}), C_8 – C_{10} compounds without any nitrogen atoms; 2) "HOM organic nitrates" (HOM_{ON}), C_8 – C_{10} compounds with one nitrogen atom; and 3) "HOM dimers" (HOM_{Di}), C_{18} – C_{20} compounds without any nitrogen atoms.

140 2.2 Instrumentation

2.2.1 Mass spectrometers

A nitrate-adduct Chemical Ionization Mass Spectrometer (NO_3 -CIMS, ToFwerk AG/Aerodyne Research, Inc.) was used for online measurements of HOM with high selectivity (Jokinen et al., 2012; Ehn et al., 2014; Riva et al., 2019b). A large sheath flow of 20 L min^{-1} (to minimize wall losses) carries nitric acid (HNO_3) across X-rays, producing nitrate ions (NO_3^-). Then, in an electric field, NO_3^- are guided towards a 10 L min^{-1} sample flow, ionizing targeted HOM molecules by clustering with them (Kürten et al., 2014). Finally, the charged sample molecules are directed through a critical orifice and into an Atmospheric Pressure Interface Time-of-Flight Mass Spectrometer (APi-TOF), where they are detected based on mass-to-charge ratios (m/z) (Junninen et al., 2010). The NO_3 -CIMS was equipped with a standard TOF (HTOF), having a mass resolution of 5000 at m/z 188 Th. The concentrations of HOM were converted from their normalized signals (i.e., the ratio of HOM-containing ions to reagent ions) by multiplying with a calibration factor (C), which takes different efficiencies into account (Jokinen et al., 2012; Bianchi et al., 2019):

$$[\text{HOM}] = C \cdot \frac{\text{HOM}(\text{NO}_3^-)}{\sum_{i=0}^2 (\text{HNO}_3)_i(\text{NO}_3^-)} \quad (1)$$

Calibrating with sulfuric acid (Kürten et al., 2012), we determined C to be $1.56 \times 10^9 \text{ cm}^{-3}$ ($\pm 50\%$) based on a flow-tube model (He et al., 2023). However, in this study, the accuracy of C is less important since the normalized signals of HOM were sufficient for relative comparisons.

A Proton Transfer Reaction Time-of-Flight mass spectrometer (PTR-TOF 8000, Ionicon Analytik GmbH), designed for online measurements of VOC (Jordan et al., 2009), was utilized in our study specifically for the detection of α -pinene. The PTR is an ionization method where water molecules (H_2O) are ionized in a hollow cathode discharge, resulting in the formation of hydronium ions (H_3O^+) (Hansel et al., 1995). Then, as proton donors, H_3O^+ are directed into a drift tube, where trace organic compounds are ionized by proton transfer process with proton affinity as the key parameter (Hansel et al., 1995; Graus et al., 2010). After a differentially pumped ion transfer unit, the charged molecules enter a TOF, where collisions are negligible under a low pressure of $\sim 10^{-6}$ mbar with a high vacuum (Graus et al., 2010). The inlet flow was 1 L min^{-1} with 0.1 L min^{-1} being sampled into the ion drift tube. Further details regarding the calibration and settings of the PTR can be found in Zhao et al. (2023b), and the calibration factor for α -pinene (detected as $\text{C}_{10}\text{H}_{17}^+$ at m/z 137 Th) was $\sim 104 \text{ ppb}^{-1}$ after normalization by primary ion isotope $\text{H}_3^{18}\text{O}^+$ (at 21 Th). The analysis of raw data from the NO_3 -CIMS and PTR-TOF was conducted using the MATLAB-based set of programs called tofTools (version 611) (Junninen et al., 2010).

2.2.2 Gas monitors

The concentrations of NO_x and O_3 were measured by gas monitors. A photometric O_3 analyzer - model 400 (Teledyne API) was used to detect O_3 in the chamber. The amount of O_3 determines how much of a 254 nm UV light signal is absorbed in the sample cell. The absorption difference between the intact sample air and the O_3 -removed air, achieved by a switching valve periodically, enables the determination of the stable O_3 concentrations.

As for NO_x , a NO-NO_2 analyzer - model T200UP (Teledyne API) was utilized. With a high-efficiency photolytic converter, NO_2 is transformed to NO with minimal interference from other gases. Using the chemiluminescence detection principle, NO is measured by reacting with O_3 , yielding light in direct proportion to the amount of NO (Archer et al., 1995). In this way, both sampled NO and total NO_x can be measured, without and with using the photolytic converter, respectively. This enables the determination of the NO_2 concentration in the sample by subtraction.

2.3 Box model

As O_3 was both injected and produced in the chamber, and sinks included reactions with NO and VOC as well as flush out, we constructed a simple 0-D box model (14 reactions, Table A2) to mimic the main reactions and to generate O_3 isopleth diagrams. These isopleths were then used to determine sensitivity regimes for O_3 formation. Reaction rates were adapted from the NIST Chemical Kinetics Database (<https://kinetics.nist.gov/kinetics/index.jsp>). The model did not include closed-shell HOM products and all peroxy radicals were treated as a single term (i.e., " RO_2 "). The box model was first employed to determine NO_2 photolysis rates under different numbers of UV lights in the zero-VOC experiments (Table A1), where only O_3 and NO_2 were used as input species. The high agreement between model and observation (detailed results are shown in section 3.1 and 3.3) showed that the model was adequate for simulating the targeted reactions in our chamber.

3 Results and discussions

3.1 NO₂ photolysis rate determination

Figure 1 illustrates an example of a zero-VOC experiment, while Table A1 provides a comprehensive list of all zero-VOC experiments conducted. Using the box model, NO₂ photolysis rates (J_{NO_2}) were determined by varying the J_{NO_2} parameter in the model until the simulated O₃ and NO_x values agreed with the observations in the zero-VOC experiments (Fig. 1 and A2–A5). With $J_{\text{NO}_2} = 1.85 \times 10^{-3} \text{ s}^{-1}$ with 5 UV lights, the modeled gas concentrations agreed extremely well with measured values (Fig. 1). Similar agreement was observed for different inputs of O₃ and NO₂ (Fig. A2 and A3), indicating the robustness of both the model itself and the fitted J_{NO_2} . Furthermore, employing the same procedure, the photolysis rates with 3 and 1 UV lights were determined to be $1.15 \times 10^{-3} \text{ s}^{-1}$ and $0.35 \times 10^{-3} \text{ s}^{-1}$, respectively (Fig. A4 and A5). The values for J_{NO_2} could also be computed from the observed steady-state and input concentrations of NO_x/O₃ for each condition (details see Table A1).

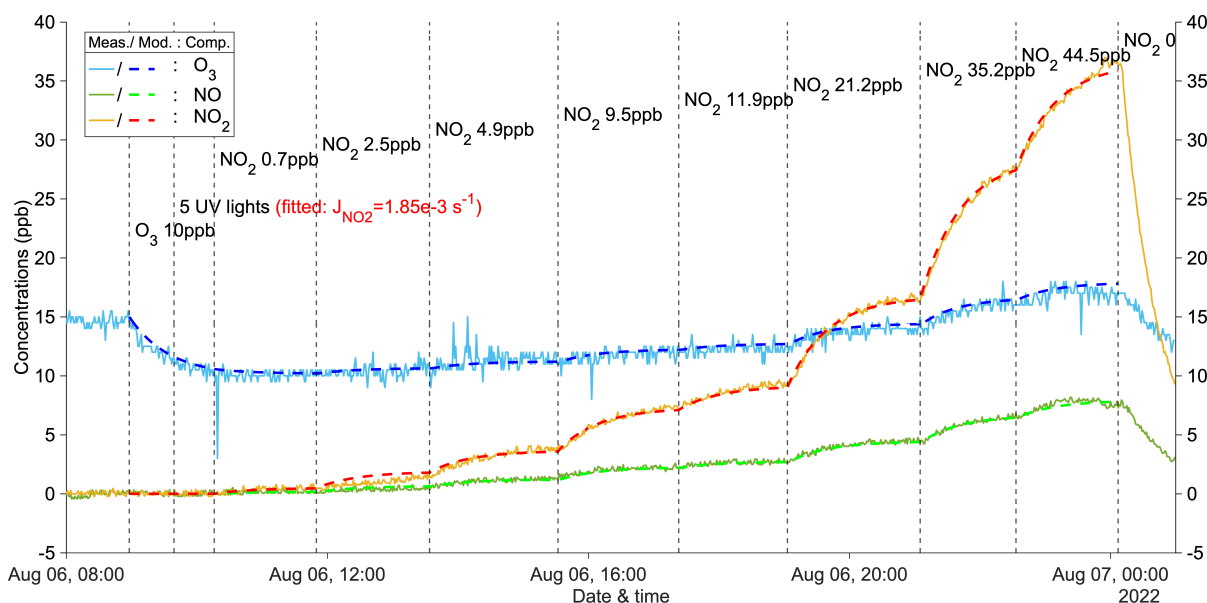


Figure 1. A zero-VOC experiment (Z1) for determining the photolysis rate of NO₂. Measured (abbreviated as Meas.) and modeled (abbreviated as Mod.) concentrations of different compounds (abbreviated as Comp.) are shown in solid and dashed lines respectively. Dashed vertical lines indicate specific time points of operations, with corresponding labels for each operation. Note that the operation labels show input information (input NO₂ = measured NO₂ + NO).

3.2 HOM determination

A steady-state spectrum (experiment no. 2) obtained from the NO₃-CIMS (Fig. 2) illustrates the identified closed-shell HOM products, including HOM_{Mono} (light green), HOM_{ON} (blue), and HOM_{Di} (dark green). Table 2 provides the formulas of all

the selected HOM species. Figure A6 displays the steady-state spectra of all stages, with the corresponding input information
 200 described in Fig. 4a. As expected, with the injection of more NO₂, the signals of HOM_{ON} increased, while those of HOM_{Mono}
 and HOM_{Di} decreased considerably (stage (2) – (5) in Fig. A6). This observation is consistent with the dominance of the RO₂
 + NO reaction over the RO₂ cross-reactions at a few ppb of NO_x (Yan et al., 2016, 2020). After the addition of more α-pinene,
 all signals showed a noticeable rise (stage (6) – (7) in Fig. A6) due to the increased supply of RO₂ species.

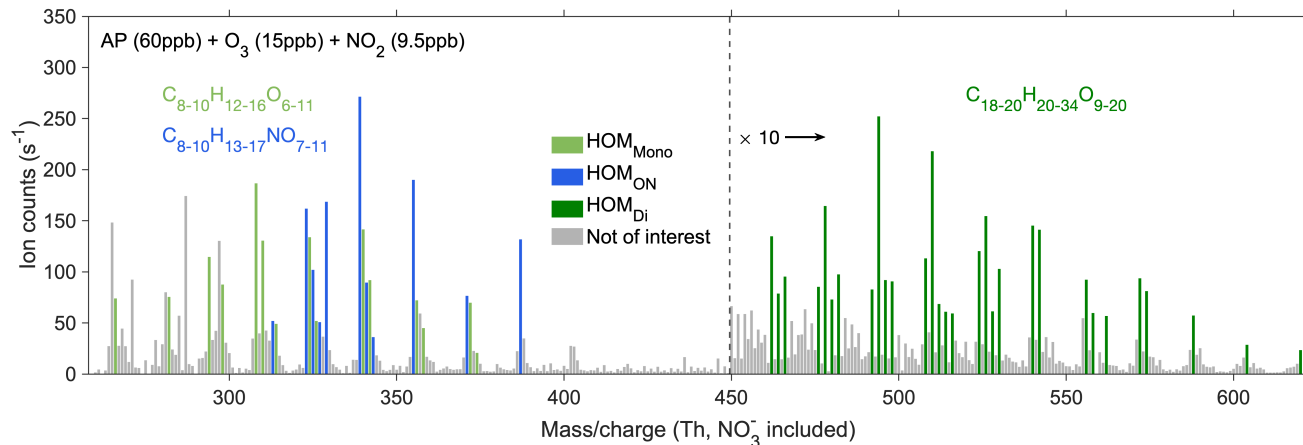


Figure 2. The steady-state spectrum (15 min average) at the stage (7) of experiment no. 2 from the NO₃-CIMS. The spectrum was corrected by subtracting the corresponding background signals. Light green bars show HOM_{Mono}, dark green ones show HOM_{Di}, blue ones show HOM_{ON} while grey ones show peaks not of interest. The peaks not of interest either exhibited relatively low signals, represented radicals, contained too small a number of carbon/oxygen atoms, or had uncertain mass-to-charge ratios. The peaks larger than 450 Th are multiplied by 10.

Table 2. Identified HOM closed-shell species based on the experiment no. 2. The reagent ion NO₃⁻ is excluded.

Monomers	Dimers	
C ₈ H ₁₂ O _x (x=6–9)	C ₁₈ H ₂₂ O _x (x=17–20)	C ₁₉ H ₃₂ O _x (x=10–13)
C ₁₀ H ₁₄ O _x (x=7–11)	C ₁₈ H ₂₄ O _x (x=12,14,16,17)	C ₂₀ H ₂₀ O ₁₅
C ₁₀ H ₁₆ O _x (x=6–11)	C ₁₈ H ₂₆ O _x (x=11–14)	C ₂₀ H ₃₀ O _x (x=10–18)
Organic nitrates	C ₁₈ H ₂₆ O _x (x=11–14)	C ₂₀ H ₃₂ O _x (x=9,11–15)
C ₈ H ₁₃ NO _x (x=8,9)	C ₁₈ H ₃₀ O ₁₃	C ₂₀ H ₃₄ O _x (x=11,12)
C ₉ H ₁₅ NO _x (x=8,9)	C ₁₉ H ₂₆ O _x (x=10–19)	
C ₁₀ H ₁₅ NO _x (x=7–11)	C ₁₉ H ₂₈ O _x (x=9–16)	
C ₁₀ H ₁₇ NO _x (x=7,8)	C ₁₉ H ₃₀ O _x (x=9–13)	

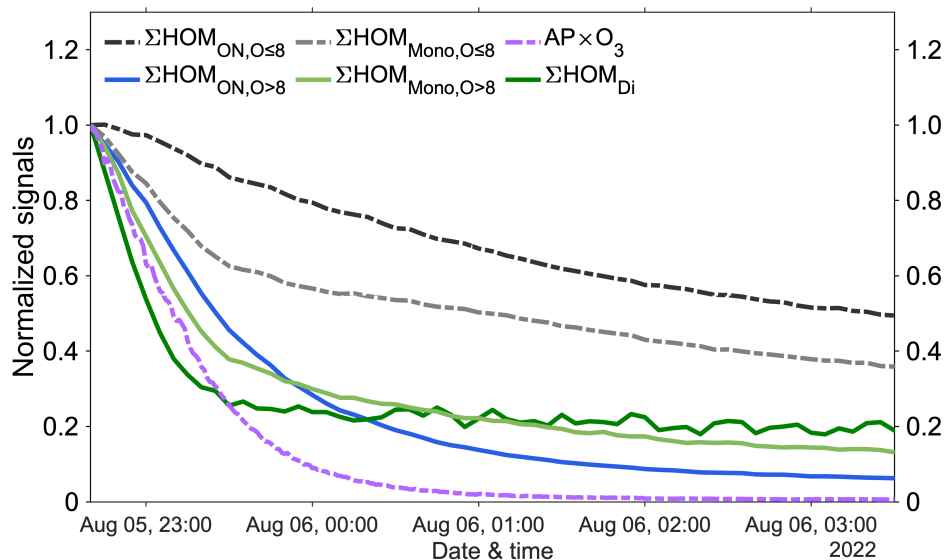


Figure 3. The normalized signal decays after experiment no. 2 ended. All signals were normalized first to primary ions and then to the signals at the moment when experiment no. 2 ended. Dashed lines: $\Sigma\text{HOM}_{\text{ON},\text{O}\leq 8}$ (sum of HOM_{ON} species with less than 9 oxygen atoms, in black), $\Sigma\text{HOM}_{\text{Mono},\text{O}\leq 8}$ (in grey), and $\text{AP}\times\text{O}_3$ (multiplication of α -pinene and O_3 concentrations, in purple). Solid lines: $\Sigma\text{HOM}_{\text{ON},\text{O}>8}$ (in blue), $\Sigma\text{HOM}_{\text{Mono},\text{O}>8}$ (in light green), and $\Sigma\text{HOM}_{\text{Di}}$ (in dark green).

Our experiments showed that HOM organic nitrates with fewer than 9 oxygen atoms ($\text{HOM}_{\text{ON},\text{O}\leq 8}$) exhibited the slowest decay at the end of the experiment (Fig. 3). This can be explained by the evaporation of these semi-volatile $\text{HOM}_{\text{ON},\text{O}\leq 8}$ compounds from chamber walls even after the gas phase production had stopped. Additionally, non-nitrate HOM monomers with fewer than 9 oxygen atoms ($\text{HOM}_{\text{Mono},\text{O}\leq 8}$) also showed an overall slow decay (Fig. 3). As a result, prior to subsequent experiments, the concentration levels of $\text{HOM}_{\text{ON},\text{O}\leq 8}$ and $\text{HOM}_{\text{Mono},\text{O}\leq 8}$ remained high, obscuring actual concentration changes following the addition of NO_x . As such, the ratios of the different compound groups are impacted by memory effects from previous experiments and wall interactions, and would therefore be poor real-time indicators of the O_3 formation sensitivity. In the atmosphere, the long-lived oxygenated VOC (OVOC) could linger for hours to days, and the ratios of e.g., nitrates to non-nitrates would be greatly influenced by various loss terms (photolysis, oxidation, condensation, hydrolysis, etc) that may differ dramatically between different compound groups. In contrast, $\text{HOM}_{\text{ON},\text{O}>8}$, $\text{HOM}_{\text{Mono},\text{O}>8}$ and HOM_{Di} experienced fast decays due to their very low volatilities, corresponding to lifetimes well below one hour when accounting for the fact that HOM formation (roughly estimated by the dashed purple line in Fig. 3) continued while the precursors were being flushed out (the higher normalized background level of HOM_{Di} is due to their lower initial concentrations, but this stable level allowed for accurate background subtraction in each experiment). Therefore, we excluded the less oxygenated and more volatile HOM species, specifically $\text{HOM}_{\text{ON},\text{O}\leq 8}$ and $\text{HOM}_{\text{Mono},\text{O}\leq 8}$ to obtain parameters with short enough lifetimes to truly reflect the ongoing chemistry in real-time. The remaining differences in decay rates in Fig. 3 may reflect slight differences in loss rates

220 also for the highly oxygenated species, but also the yields of the different compounds groups are a function of e.g., reaction rates. For example, dimer formation requires $\text{RO}_2 + \text{RO}_2$ reactions, which are more favored at higher VOC oxidation rates. Consequently, the indicating ratio used in this study is defined as the ratio between the sum of HOM_{Di} or $\text{HOM}_{\text{Mono},\text{O}>8}$ and the sum of $\text{HOM}_{\text{ON},\text{O}>8}$, represented as $\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON},\text{O}>8}}$ (Indicating Ratio 1, abbreviated as IR1) or $\frac{\sum \text{HOM}_{\text{Mono},\text{O}>8}}{\sum \text{HOM}_{\text{ON},\text{O}>8}}$ (IR2). We emphasize that these ratios will be specific to our experiments, as yields for the different groups will vary for different VOC
225 and conditions. Nevertheless, the use of the most oxidized, i.e., the least volatile, HOM is critical in this context, as their loss rates are very similar as they all behave as if they were non-volatile, thus condensing irreversibly to surfaces as the main sink term.

3.3 Indicating ratios

In this section, we detail the conducted experiments. Experiment no. 2, is depicted in Fig. 4, while the other experiments are
230 shown in Fig. A7–A12. It is worth noting that our experiments did not result in notable particle formation, and condensation to walls was always the dominant loss term for HOM. If aerosol formation had been significant, as has been observed in our chamber at higher oxidation rates (Zhao et al., 2023a), HOM would first increase due to fast formation, and then decrease due to condensation sink. Variation of NO_2 and α -pinene input concentrations lead to changes in both indicating ratios (IR1 and IR2) that correspond with changes in O_3 concentrations (Fig. 4), suggesting a possible sensitivity of O_3 formation. From stage
235 (1) to (5), injection of more NO_2 led to increased formation of $\text{HOM}_{\text{ON},\text{O}>8}$, while production of HOM_{Di} and $\text{HOM}_{\text{Mono},\text{O}>8}$ was suppressed. As a result, the indicating ratios decreased. Additionally, the concentration of O_3 increased during this period, but the rate of increase decreased with higher NO_2 inputs. This trend suggests a gradual shift from a NO_x -limited regime to a more VOC-limited regime in the system. For example, during stage (5) with 9.5 ppb NO_2 , the O_3 concentration remained relatively constant. This observation indicates the system may have shifted to the VOC-limited regime. Next, when additional
240 α -pinene was introduced (~ 45 ppb during stage (6)), a significant increase in O_3 concentration was again observed, consistent with the system being in the VOC-limited regime during stage (5). However, after injection of ~ 60 ppb α -pinene during stage (7), the O_3 concentration reached a plateau, indicating that the system had shifted back to the NO_x -limited regime. Moreover, during these two stages, the indicating ratios experienced a substantial increase. In the last stage (8), when additional NO_2 was
245 injected to reach 21.5 ppb, while input α -pinene concentration remained unchanged, the O_3 concentration increased again, and the indicating ratios decreased noticeably, confirming that the previous stage was NO_x -limited. However, it should be noted that the system in this stage might not yet be VOC-limited.

Other experiments with 5 UV lights (Fig. A7, A8, and A11) also exhibited similar time series patterns as described above. One noticeable difference is that higher initial O_3 concentrations resulted in less pronounced increases in O_3 during the first
250 5 stages with varying NO_2 levels. This can be attributed to the reaction $\text{O}_3 + \text{NO}$ (R2) becoming faster, competing with the formation of O_3 from the $\text{RO}_2 + \text{NO}$ reaction (R4a) followed by NO_2 photolysis (R1). As a result, there was a reduced O_3 formation in the presence of higher initial O_3 concentrations due to the scavenging of NO by existing O_3 .

Compared to experiment no. 2 (Fig. 4), experiments with fewer lights (but similar initial concentrations; Fig. A9 and A10) provide insight into the effect of UV light intensities on O_3 formation sensitivity and the consistency of the indicating ratios

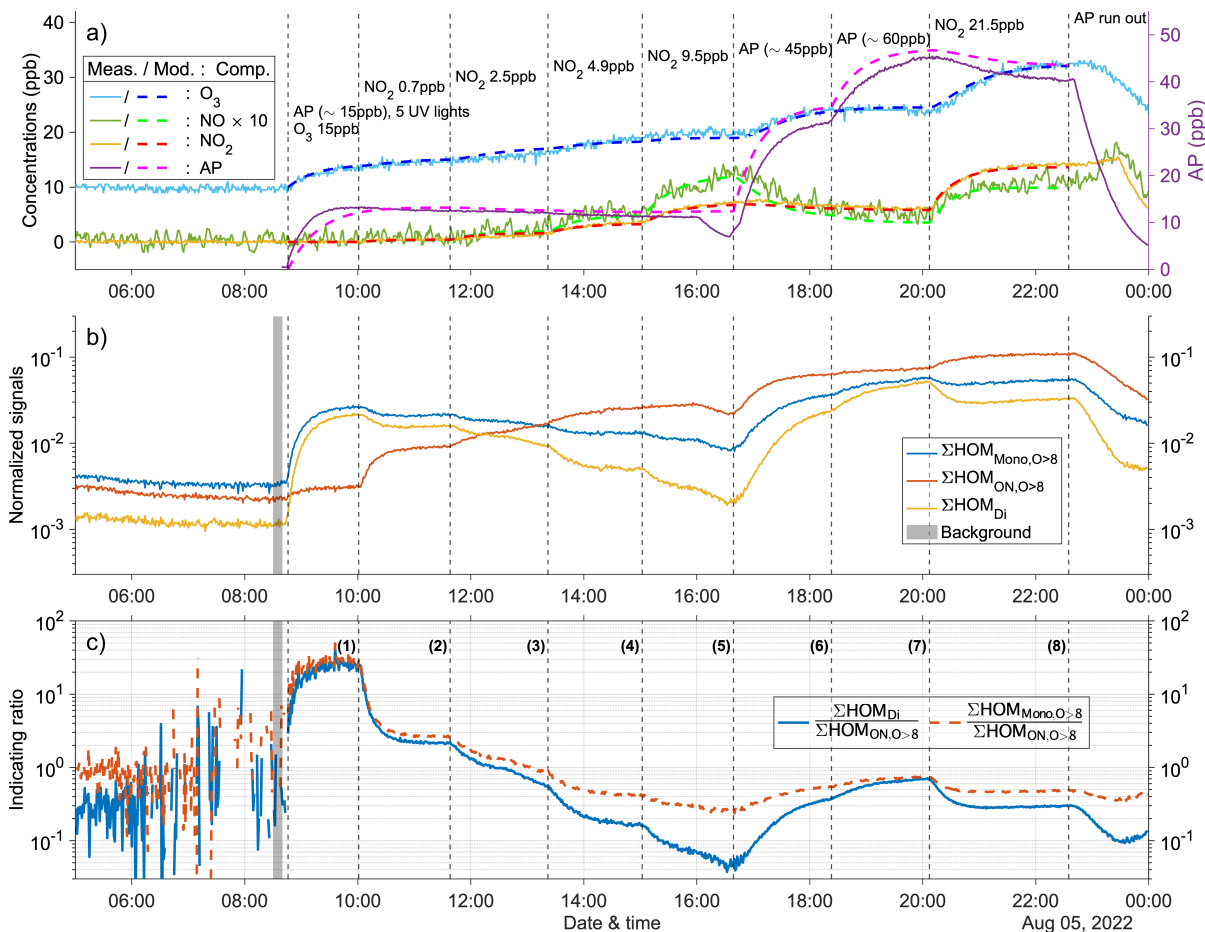


Figure 4. Time series of experiment no. 2 with 15 ppb α -pinene and 15 ppb O₃ as initial inputs. 5 UV lights were on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O₃, NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma\text{HOM}_{\text{Mono},\text{O}>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma\text{HOM}_{\text{ON},\text{O}>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and $\Sigma\text{HOM}_{\text{Di}}$ (sum of HOM dimers); c) IR1 ($\frac{\Sigma\text{HOM}_{\text{Di}}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$) and IR2 ($\frac{\Sigma\text{HOM}_{\text{Mono},\text{O}>8}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a. The bolded number in parentheses in subplot c corresponds to the number of stages. The steady-state mass spectra obtained by the NO₃-CIMS for each stage is shown in Fig. A6.

255 under different light intensities. More lights led to more pronounced increase in O₃ concentrations at the same stages because additional O₃ was produced from NO₂ photolysis. On the other hand, fewer lights resulted in lower NO levels in the system since NO₂ input was the sole source of NO_x. This led to reduced NO₂ formation from RO₂ + NO reaction (R4a), and subsequently less O₃ formation. This aspect is crucial in determining the O₃ sensitivity. In this sense, the presence of fewer lights

implies that higher levels of NO₂ input are required to ensure sufficient NO levels for reaching the VOC-limited regime. This can be confirmed by the stage where ~ 45 ppb α-pinene was injected (Fig. 4, A9, and A10). During this stage, the O₃ concentration showed a significant increase with either 3 or 5 UV lights, indicating that the system had reached the VOC-limited regime in the previous stage. However, when only 1 UV light was used, the O₃ concentration remained relatively constant, suggesting that the system with 1 light did not reach the VOC-limited regime. It is worth mentioning that the indicating ratios exhibited more significant changes when using more lights. This can be attributed to the fact that more lights result in increased production of O₃ and NO, leading to more drastic changes in HOM distributions and thereby influencing the indicating ratios to a greater extent.

Using the box model described in section 2.3, the concentrations of O₃ and its precursors were captured well both qualitatively and quantitatively (e.g., Fig. 4a). In general, simulated concentrations of O₃ during the steady states differed from measured values by at most 10%, while differences for NO_x were even smaller. The largest discrepancy in concentrations (~ 15%) was observed for α-pinene, which can be attributed to the simplifications made in the model. Specifically, the OH concentration will be underestimated if the model did not accurately capture the yields of HO₂, which can be converted into OH via reactions with NO.

Overall, both indicating ratios are promising as indicators of O₃ formation sensitivity. However, in all time series (Fig. 4c, A7c–A12c), IR1 ($\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) exhibited more pronounced changes compared to IR2 ($\frac{\sum \text{HOM}_{\text{Mono}, \text{O} > 8}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) as we shifted the O₃ formation regimes. This highlights that in the well-controlled chamber systems we investigated, IR1 may hold better potential for indicating O₃ formation sensitivity in the absence of other perturbing factors. It can be explained by the fact that the nitrates are solely from RO₂+NO and the dimers solely from RO₂+RO₂. But HOM monomers can be from both of these reactions. On the other hand, in the real atmosphere, IR1 is expected to be much less robust, as discussed in more detail in section 3.5. The intensity of light (i.e., NO₂ photolysis rate) played a crucial role in determining O₃ sensitivity regimes by controlling the amount of NO. Our box model successfully reproduced the measured values of O₃ and its formation precursors, making it reasonable to extend the model to generate O₃ isopleths under chamber conditions beyond those covered in our experiments. With these we can better elucidate how well the HOM ratios can function as O₃ sensitivity indicators in this system.

3.4 Viability of the indicating ratios

In order to validate the indicating ratio, we generated O₃ isopleths by modeling various combinations of input NO₂ and α-pinene concentrations. The indicating ratios obtained from all steady state stages were scattered on the same coordinate system (IR1: Fig. 5–6 and IR2: Fig. A13–A14). The purpose of varying the injection rates of either NO_x or α-pinene within each experiment (e.g., colored curves in Fig. 5) was to shift the system between the VOC-limited and NO_x-limited regimes. Unlike previous studies that used the concentration of VOC as the x-axis (Kinosian, 1982; Chameides et al., 1992), in this study, the x-axis represented the product of the measured α-pinene and O₃ concentrations. This is because the product reflects the potential for RO₂ formation, which reacts directly with NO to contribute to O₃ accumulation (R1–R4a). In typical EKMA plots the oxidant is primarily thought to be OH, the concentration of which is independent of the VOC concentrations, thus the x-axis would be largely equivalent regardless of whether plotting VOC or VOC times oxidant. In our case, the VOC concentration

will directly influence the oxidant (i.e., O_3) concentration through chemical reactions, and therefore we chose to use the current x-axis. The ridge line of the EKMA curves was plotted for each experiment, represented by dotted lines in e.g., Fig. 5b–d. The reason for separating the b–d plots from each other is due to the differences in constant O_3 inputs or NO_2 photolysis rates in each experiment, resulting in distinct O_3 isopleths. For the context of this work, we do not separate a transition regime where O_3 formation is sensitive to both NO_x and VOC, but simply define the VOC- and NO_x -limited regimes based on the ridge line. Farther above the ridge line, the system is more VOC-limited, while farther below the line, the system is more NO_x -limited. It's notable that in the VOC-limited regime, instead of a reduction, the O_3 concentration even increased slowly with the addition of NO_x , primarily due to the photolysis of the input NO_2 .

Comparing experimental and model results, the steady state stages of all experiments with 5 UV lights exhibited a consistent pattern for the indicating ratios, allowing qualitative determination of O_3 formation sensitivity (Fig. 5 and A13). Generally, the farther the steady state point is from the ridge line, the darker the coloring (either blue or green). Specifically, when the color is darker blue, it indicates a smaller value of the indicating ratios, suggesting a higher likelihood of the system being in the VOC-limited regime. Conversely, when the color is darker green, it signifies a higher value of the ratios, indicating a higher likelihood of the system being in the NO_x -limited regime. It is worth noting that experiment no. 7 was essentially a duplicate of experiment no. 1, but with additional stages. When comparing the indicating ratios at stages with the same inputs, as shown in Fig. 5d and A13d, it becomes apparent that the values are highly consistent and closely aligned. Moreover, the background signals of $HOM_{ON,O>8}$, $HOM_{Mono,O>8}$ and HOM_{Di} increased substantially with experiments going on, from $5 \times 10^{-4} - 2 \times 10^{-3}$ (Fig. A7b) to $4 \times 10^{-3} - 2 \times 10^{-2}$ (Fig. A11b). However, the accumulating background did not have a significant impact on the indicating ratios, after the background subtraction (Fig. 5d and A13d). These highlight the remarkable reproducibility of the indicating ratios in our chamber experiments.

To investigate the impact of light intensities, a similar comparison was conducted for experiments with the same initial inputs using 5, 3, and 1 UV lights (Fig. 6 and A14). The observed changes in the pattern of the indicating ratios are in line with those observed in the 5 UV lights experiments described above. The most significant observation is that at lower UV light intensity, the ridge line shifted towards higher NO_x levels at the same potential for RO_2 formation (i.e., the same value of x-axis) (Fig. 6b–d and A14b–d). This finding aligns with the time-series comparison presented in section 3.3, which indicates that at weaker UV intensity, higher input NO_2 is required to generate NO levels sufficient to shift the system towards the VOC-limited regime. When comparing stages with the same conditions except different UV lights, we observed that generally the lower intensity of UV lights corresponded to higher values of the indicating ratios (Fig. 6 and A14), showing a higher likelihood of the system being in the NO_x -limited regime. This finding is in agreement with the shift of the ridge line. The correlation between UV intensities, the indicating ratios, and the position of the ridge line reinforces the relationship between the indicating ratios and O_3 formation sensitivity. It suggests that in addition to the relative changes, the absolute values of the indicating ratios are also informative.

The conclusion can be drawn that the indicating ratios can qualitatively predict the O_3 formation regimes (either VOC- or NO_x -limited). More specifically, based on modeled EKMA curves and steady state HOM ratios (Fig. 5, 6, A13, and A14),

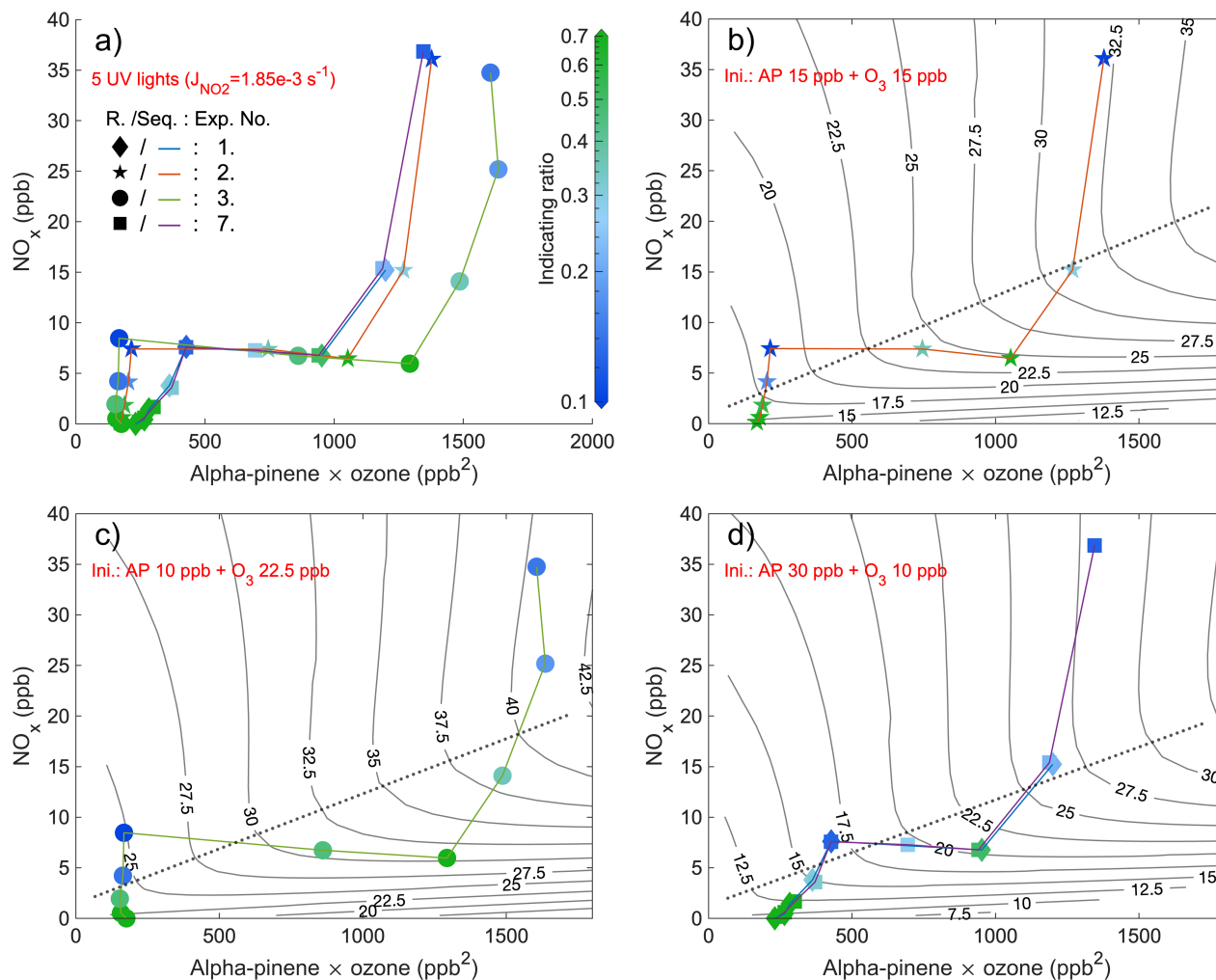


Figure 5. Steady state IR1 ($\frac{\sum \text{HOM}_{\text{D}_i}}{\sum \text{HOM}_{\text{O}_N, \text{O}_{>8}}}$) of experiments from four days with 5 UV lights. X-axis is the multiplication of steady state α -pinene and O_3 concentrations, while y-axis is the steady state NO_x . The scatter points (exp. no. 1 (experiment number 1): diamond; exp. no. 2: star; exp. no. 3: round; exp. no. 7: square) are colored by values of IR1 (abbreviated as R. in the figure), and are connected by curves (exp. no. 1: blue; exp. no. 2: orange; exp. no. 3: green; exp. no. 7: purple) showing the sequence (Seq.) of experimental stages. Panel a combines stages of all four days, and the rest three subplots respectively show the stages of experiments with different initial (Ini.) inputs (exp. no. 1 and 7 are in the same panel d due to same initial inputs). EKMA curves (isopleth of O_3 concentrations in ppb), simulated by the box model, are black solid lines, while dotted lines are corresponding ridge lines.

regardless of the light intensity, IR1/IR2 consistently indicate the VOC-limited regime when below 0.2/0.4, and the NO_x -limited regime when above 0.5/0.7.

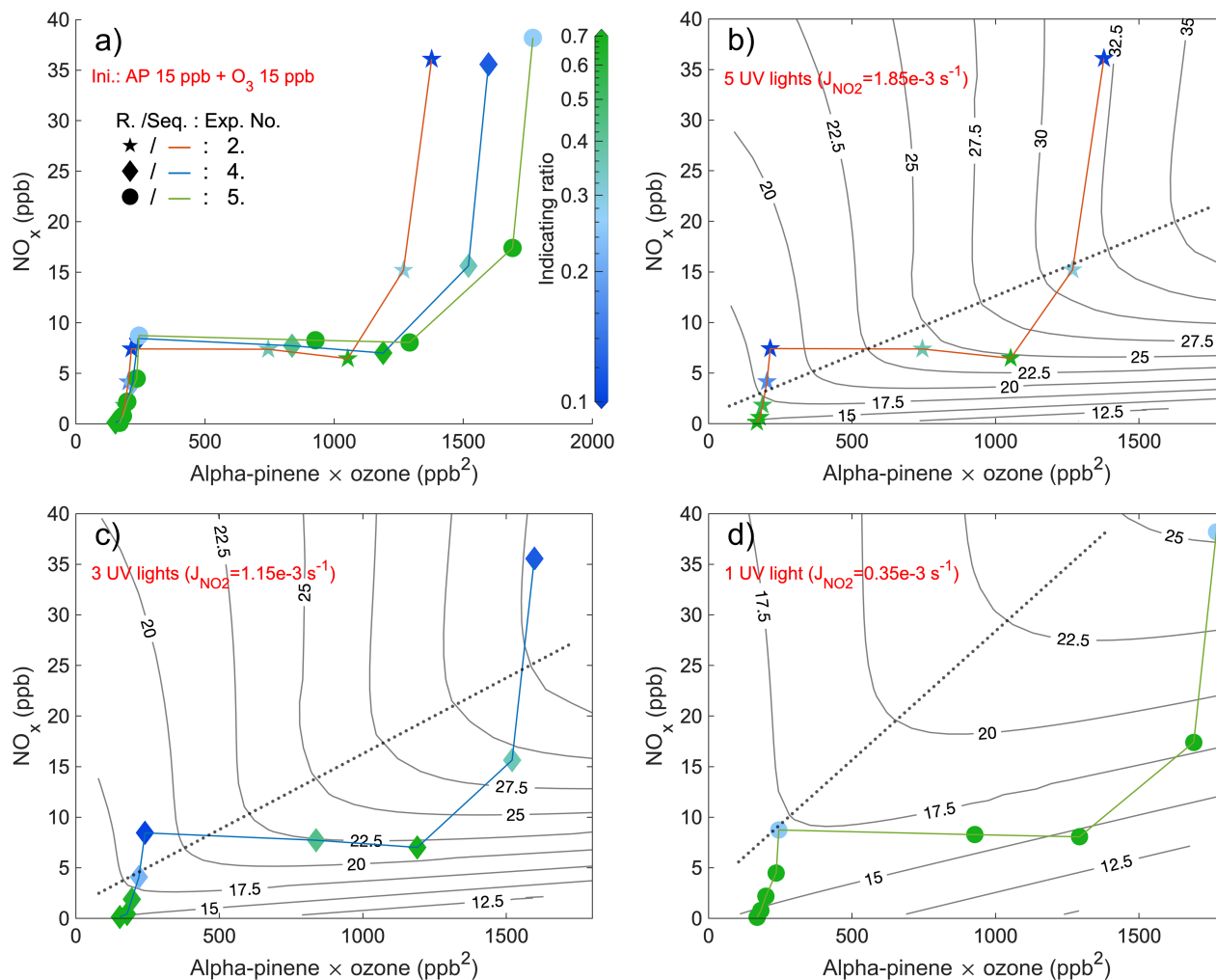


Figure 6. Steady state IR1 ($\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) of experiments from three days with 5, 3, and 1 UV lights, respectively. X-axis is the multiplication of steady state α -pinene and O_3 concentrations, while y-axis is the steady state NO_x . The scatter points (exp. no. 2 (experiment number 2): star; exp. no. 4: diamond; exp. no. 5: round) are colored by values of IR1 (abbreviated as R. in the figure), and are connected by curves (exp. no. 2: orange; exp. no. 4: blue; exp. no. 5: green) showing the sequence (Seq.) of experimental stages. Panel a combines stages of all three days with the same initial (Ini.) inputs, and the rest three subplots respectively show the stages of experiments with different amount of UV lights. EKMA curves by the box model are in black lines and the dotted lines are corresponding ridge lines.

3.5 Implication and further improvements

Our chamber study (section 3.4) confirmed the significant role of the indicating ratios in determining O_3 formation sensitivity qualitatively. However, in the real atmosphere, the conditions will vary significantly more than in our simple system. Most importantly, the amount of different precursor VOC will be vastly greater, and in reactions with OH, the distribution of dif-

ferent types of RO₂ radicals will also be far more complex. As an example, in a rural setting there can be a wide variety of small VOC (C1–C4), isoprene (C5), aromatics (C6–C9) and monoterpenes (C10) that all produce significant amounts of RO₂. Consequently, from these molecules, dimers can form with any carbon number between 2 and 20, while monomers can have carbon numbers between 1 and 10. In addition, both VOC and oxidant concentrations as well as radiation and meteorological patterns will vary over time. This means that it will be difficult to find universal compounds to include in the HOM_{Di}, HOM_{Mono} and HOM_{ON} groups used to calculate the indicating ratios for a given site and a given time. For example, it is likely that there will be few environments where C10-RO₂ from monoterpenes would be efficiently reacting with each other during daytime, meaning that the C20 dimers used here are unlikely to be usable. It remains to be seen whether the HOM_{Mono}-to-HOM_{ON} ratio can be used for C10 compounds in areas with high monoterpene emissions. Nevertheless, conceptually the link between HOM formation pathways and O₃ formation should hold, and it may be possible to determine suitable compound groups for various sites. Our study focused exclusively on α-pinene, but the intrinsic connection between the indicating ratios we proposed and O₃ formation (Fig. A1) is not limited to this specific VOC. Further laboratory and ambient studies are necessary to investigate additional VOC of interest and expand our understanding of the indicating ratios' applicability and generalizability in predicting O₃ formation sensitivity under various atmospheric conditions. In addition, at the extremes, ranging from clear formation of dimeric species to complete lack of dimeric species with abundant organic HOM nitrates, this can be considered a strong qualitative indicator being in a NO_x- or VOC-limited regime, respectively.

4 Conclusions

Both O₃ and HOM are of significant interest, given their impacts from small-scale personal health to large-scale global climate. Due to the intrinsic connection between the formation mechanisms of O₃ and HOM, we suggest new indicators, denoted as $\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$ and $\frac{\sum \text{HOM}_{\text{Mono}, \text{O} > 8}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$, for determining O₃ formation regimes based on the distribution of HOM. One main improvement of using HOM-based indicating ratios compared to those suggested earlier would be the short lifetimes of HOM, which means that these new indicators would be real-time indicators of the formation regime. To assess the viability of the indicating ratios, a series of chamber experiments were carried out using a NO₃-CIMS, a PTR-TOF, and O₃/NO_x monitors. As expected, an increase in NO_x inputs led to an increase in HOM_{ON} and a decrease in HOM_{Di}, HOM_{Mono} and the indicating ratios. Conversely, an increase in α-pinene resulted in a rise in the indicating ratios. Furthermore, when adding enough of one of the O₃ formation precursors (either NO_x or α-pinene), the rate of increase in O₃ concentration slowed down or even stopped. This indicates that the system was shifting to or had already reached the other limited regime.

With a box model, which closely reproduced the measured concentrations of O₃ and its precursors, O₃ isopleths were obtained for different concentrations of NO_x and α-pinene. After drawing ridge lines of the isopleths, it was observed that the indicating ratios provide a qualitative prediction of the O₃ formation regimes: lower values of the ratios indicate a greater likelihood of the system being located in the VOC-limited regime, and vice versa. With less intense UV light (λ ≈ 400 nm), a higher amount of NO₂ was required to shift the system towards the VOC-limited regime. This can be attributed to a decrease in the formation of NO from NO₂ photolysis. Nevertheless, the absolute values of the indicating ratios exhibited a consistent

365 behavior across different intensities of UV light, suggesting that these absolute values are highly valuable for analyzing O₃ formation sensitivity.

The main objective of this study was to evaluate the concept of using HOM distributions in indicating O₃ formation sensitivity. Based on our outcomes, we can conclude that the ratio of HOM dimers or non-nitrate monomers to HOM organic nitrates (i.e., $\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$ or $\frac{\sum \text{HOM}_{\text{Mono}, \text{O} > 8}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) has the capability to indicate O₃ formation regimes (either VOC- or NO_x-limited),
370 in this simple system of monoterpene ozonolysis. An indicating ratio of this kind would aid in better control of O₃ pollution and have the potential to be incorporated as a useful parameter in global models for analyzing O₃ formation sensitivity under diverse environmental conditions.

Nevertheless, future studies will need to assess whether this approach is feasible to be applied in real-world conditions where the chemistry is far more complex. The variability of VOC precursors alone will greatly perturb the ideal situation
375 observed in our chamber. Still, we posit that environments with high monoterpene emissions will also produce abundant C₁₀-RO₂ concentrations, and the comparison of the highly oxygenated monomeric termination products (i.e., nitrates vs non-nitrates) can provide an indication of the relative RO₂ termination pathways. Further studies, both ambient observation as well as chamber experiments involving multiple VOC and oxidants, will be necessary to determine the potential of HOM-based indicators for O₃ formation.

380 *Code availability.* Code is available upon request from the corresponding authors.

Data availability. Data is available upon request from the corresponding authors.

Appendix A: Additional figures and tables

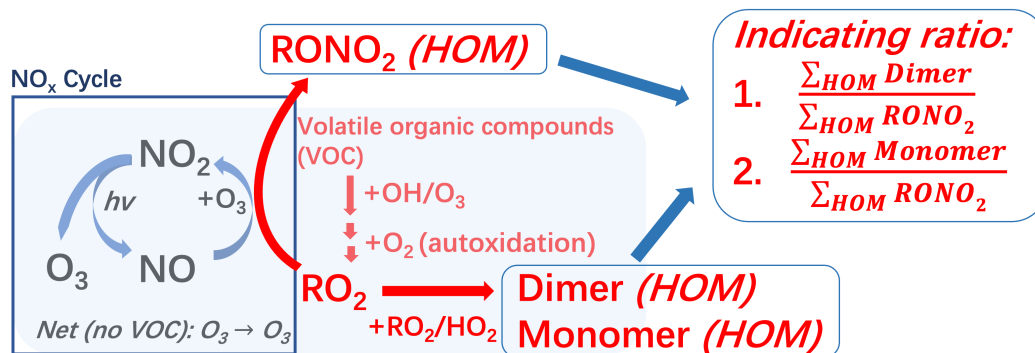


Figure A1. A sketch of the connection between HOM and O₃ formation. Based on the formation connection, two indicating ratios between the HOM species are defined. "RONO₂ (HOM)": nitrate-containing HOM monomers, "Dimer (HOM)": non-nitrate HOM dimers, "Monomer (HOM)": non-nitrate HOM monomers.

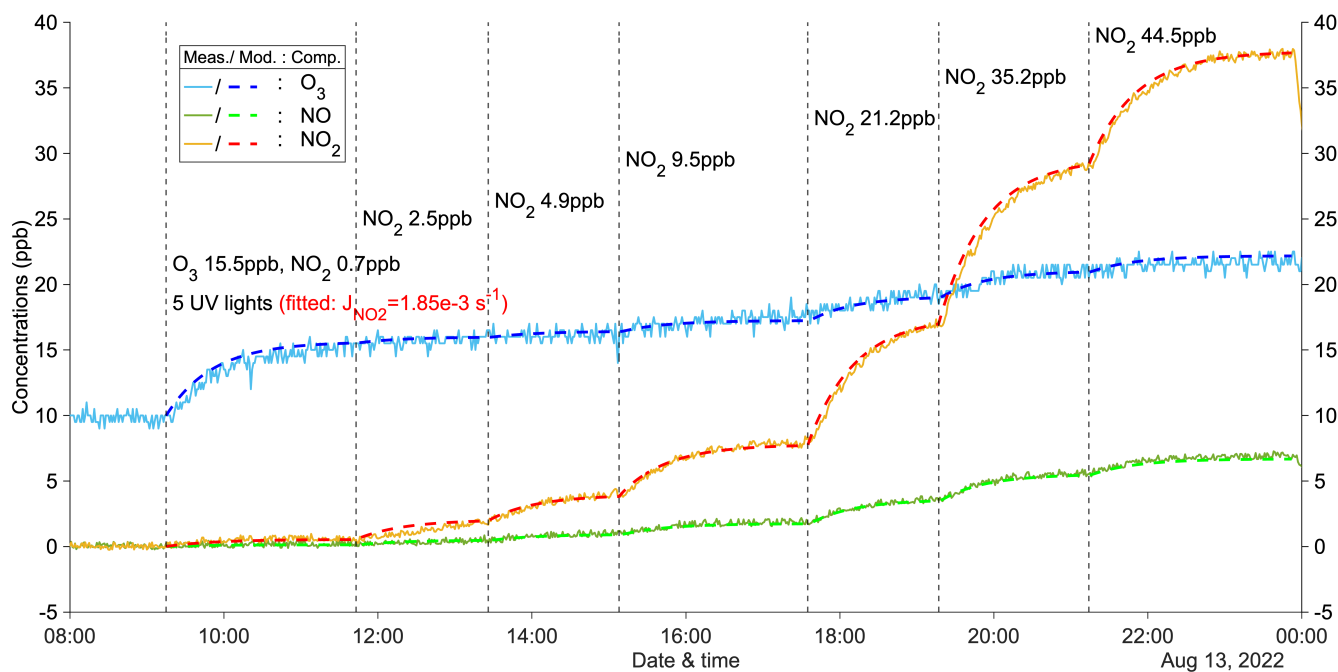


Figure A2. A zero-VOC experiment (Z2) for determining the photolysis rate of NO₂. Measured (abbreviated as Meas.) and modeled (abbreviated as Mod.) concentrations of different compounds (abbreviated as Comp.) are shown in solid and dashed lines respectively. Dashed vertical lines indicate specific time points of operations, with corresponding labels for each operation.

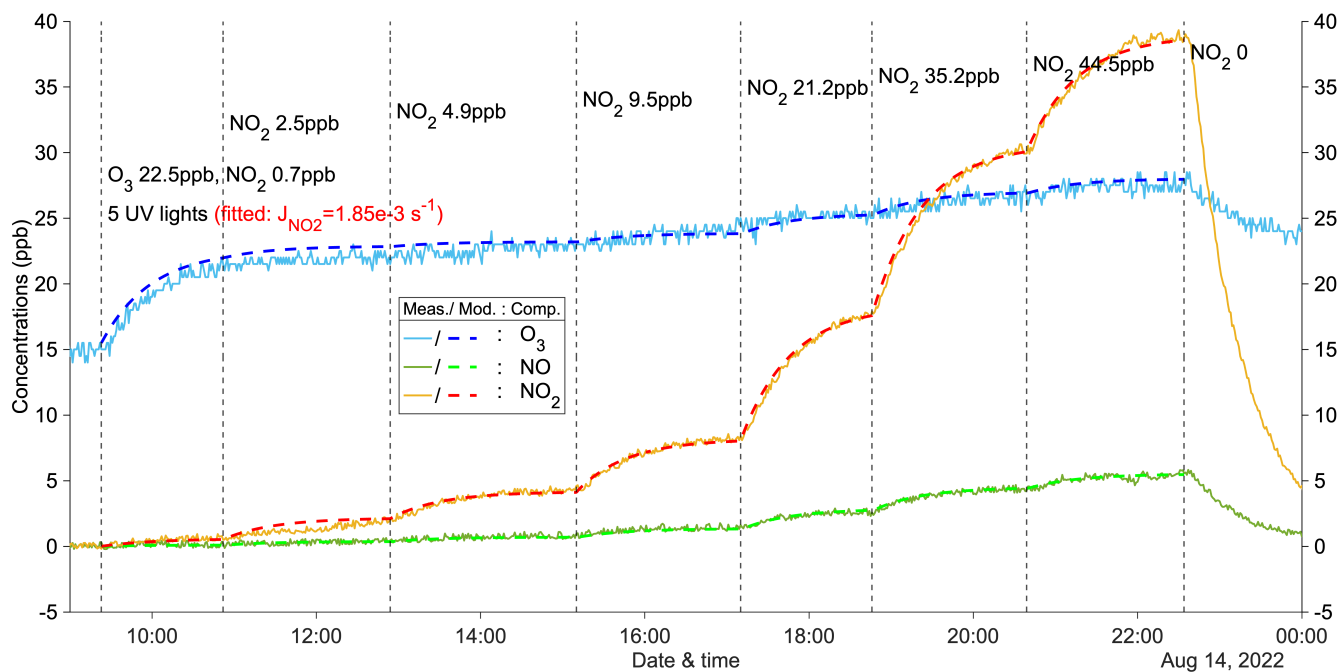


Figure A3. A zero-VOC experiment (Z3) for determining the photolysis rate of NO₂. Measured (abbreviated as Meas.) and modeled (abbreviated as Mod.) concentrations of different compounds (abbreviated as Comp.) are shown in solid and dashed lines respectively. Dashed vertical lines indicate specific time points of operations, with corresponding labels for each operation.

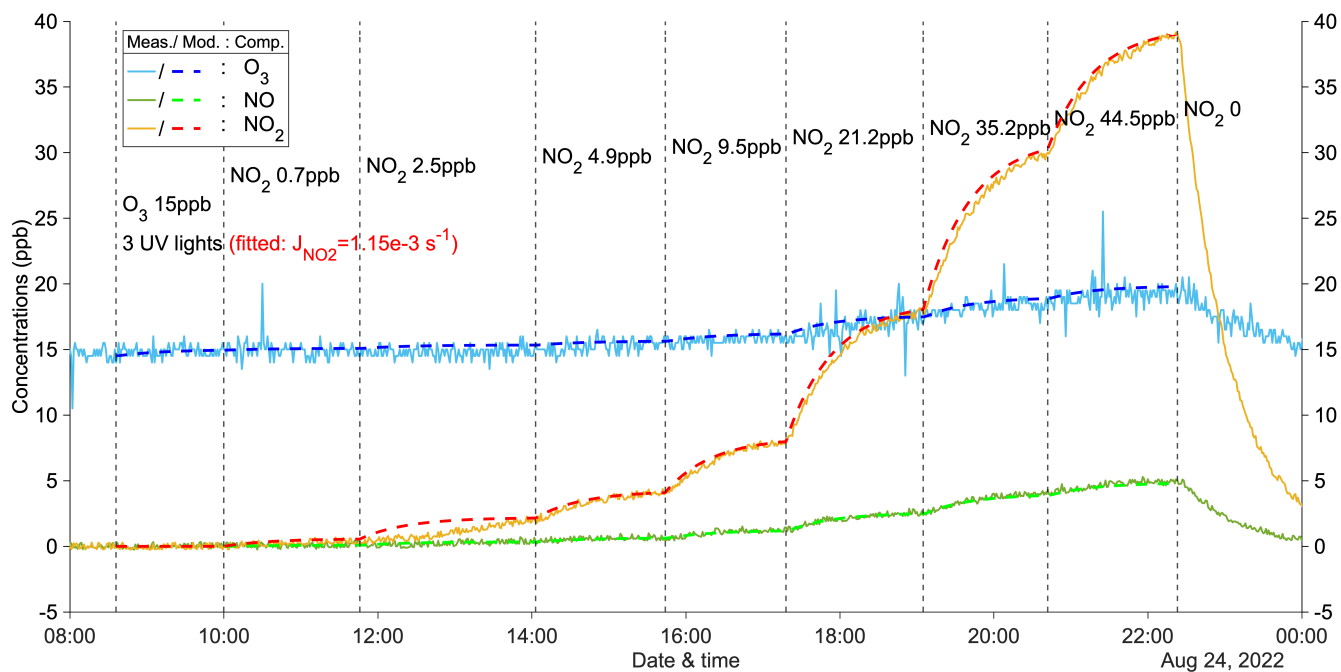


Figure A4. A zero-VOC experiment (Z4) for determining the photolysis rate of NO₂. Measured (abbreviated as Meas.) and modeled (abbreviated as Mod.) concentrations of different compounds (abbreviated as Comp.) are shown in solid and dashed lines respectively. Dashed vertical lines indicate specific time points of operations, with corresponding labels for each operation.

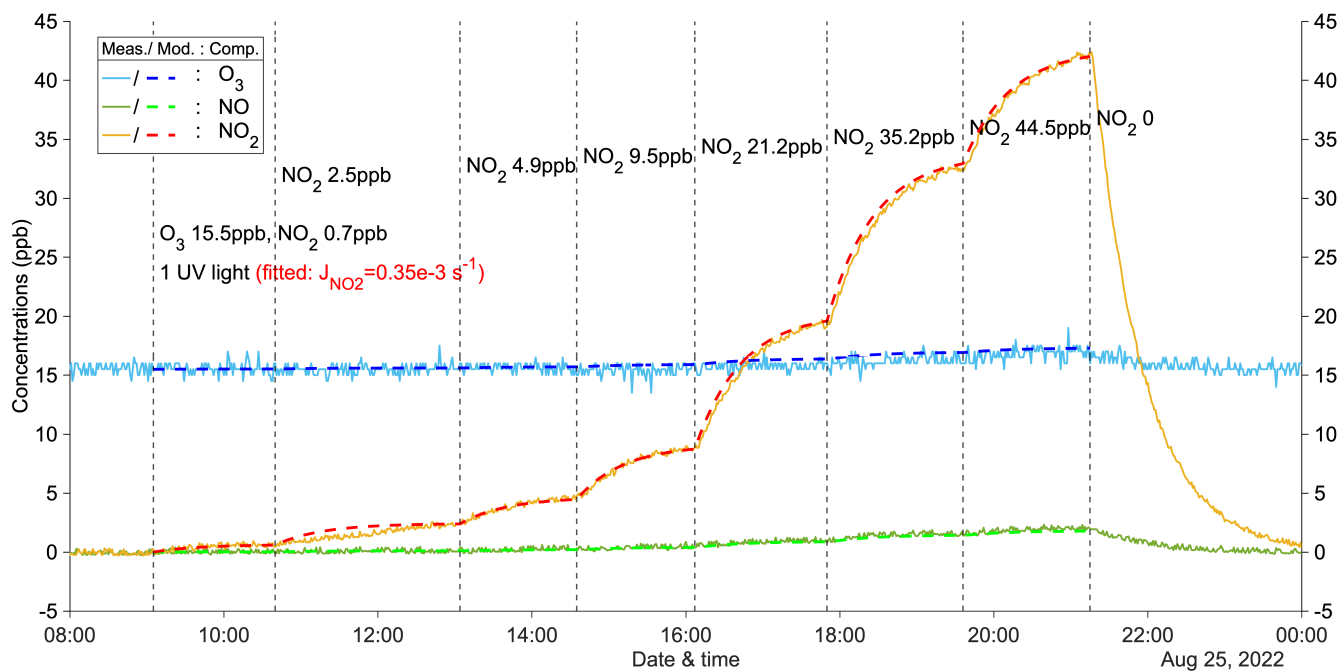


Figure A5. A zero-VOC experiment (Z5) for determining the photolysis rate of NO₂. Measured (abbreviated as Meas.) and modeled (abbreviated as Mod.) concentrations of different compounds (abbreviated as Comp.) are shown in solid and dashed lines respectively. Dashed vertical lines indicate specific time points of operations, with corresponding labels for each operation.

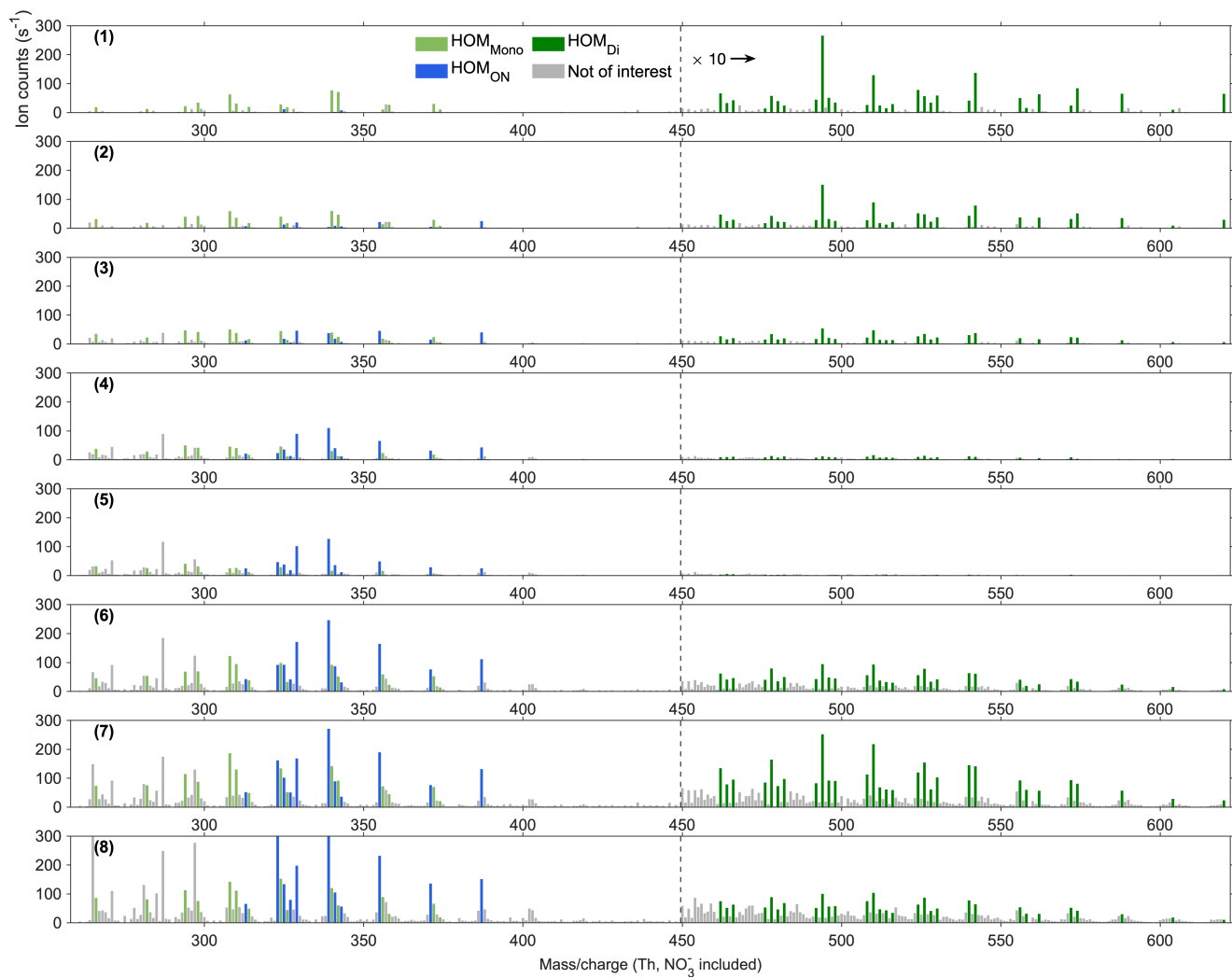


Figure A6. Steady-state spectra (15 min average) at experiment no. 2 from the NO_3 -CIMS. All spectra were corrected by subtracting the corresponding background signals. The number at each row show the order of the stage, consistent with the time series in Fig. 4. Light green bars show HOM_{Mono} , dark green ones show HOM_{Di} , blue ones show HOM_{ON} while grey ones show peaks not of interest. The peaks larger than 450 Th are multiplied by 10.

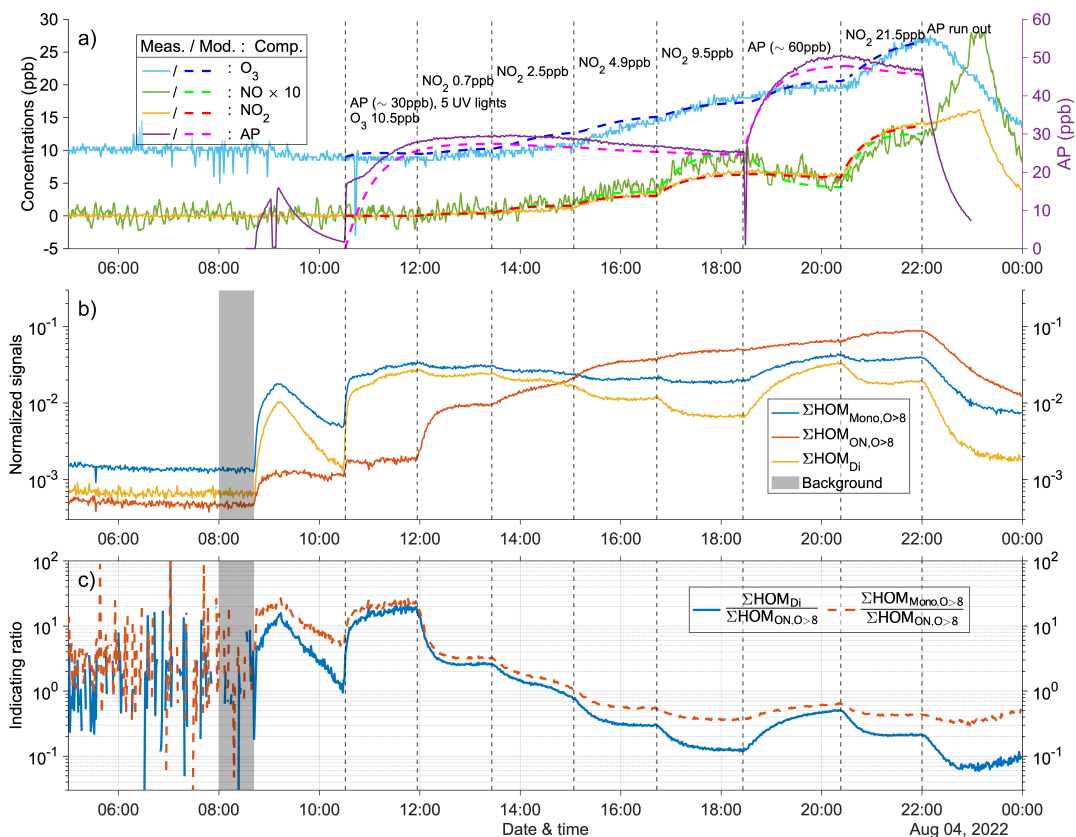


Figure A7. Time series of experiment no. 1 with 30 ppb α -pinene and 10.5 ppb O₃ as initial inputs. 5 UV lights were on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O₃, NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\sum \text{HOM}_{\text{Mono}, \text{O}>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\sum \text{HOM}_{\text{ON}, \text{O}>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and $\sum \text{HOM}_{\text{Di}}$ (sum of HOM dimers); c) IR1 ($\frac{\sum \text{HOM}_{\text{Di}}}{\sum \text{HOM}_{\text{ON}, \text{O}>8}}$) and IR2 ($\frac{\sum \text{HOM}_{\text{Mono}, \text{O}>8}}{\sum \text{HOM}_{\text{ON}, \text{O}>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

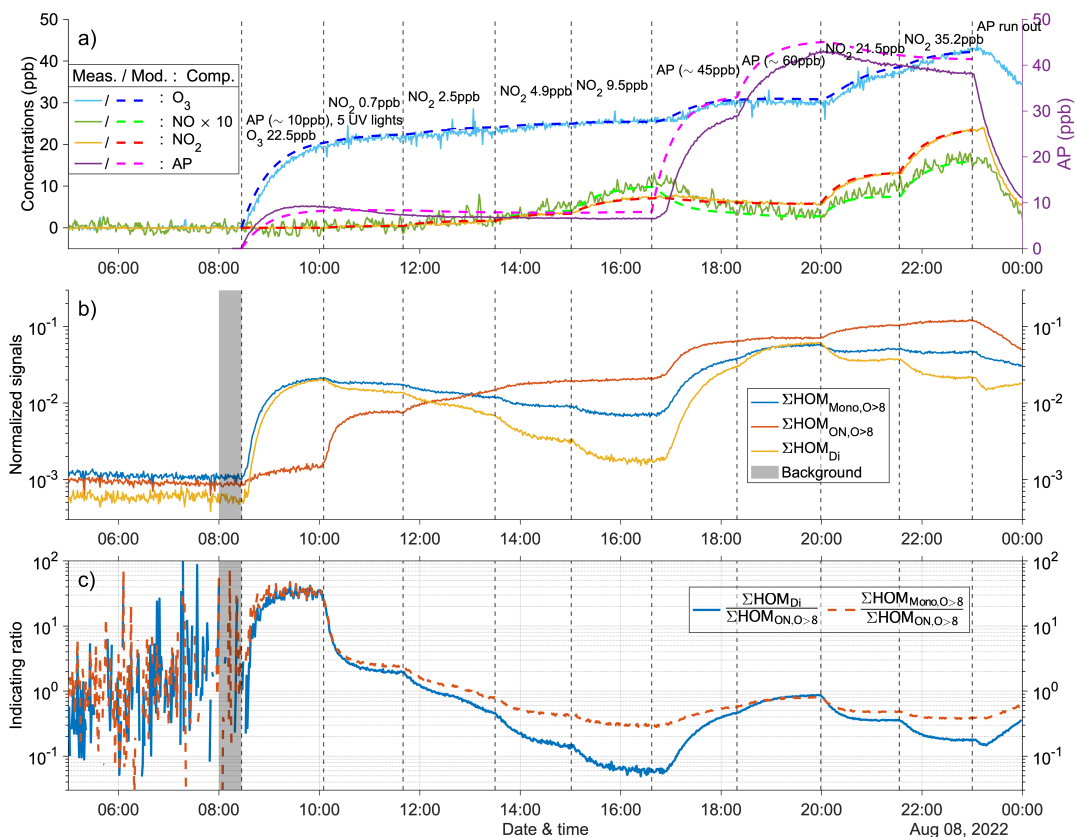


Figure A8. Time series of experiment no. 3 with 10 ppb α -pinene and 22.5 ppb O₃ as initial inputs. 5 UV lights were on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O₃, NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma\text{HOM}_{\text{Mono},\text{O}>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma\text{HOM}_{\text{ON},\text{O}>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and $\Sigma\text{HOM}_{\text{Di}}$ (sum of HOM dimers); c) IR1 ($\frac{\Sigma\text{HOM}_{\text{Di}}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$) and IR2 ($\frac{\Sigma\text{HOM}_{\text{Mono},\text{O}>8}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

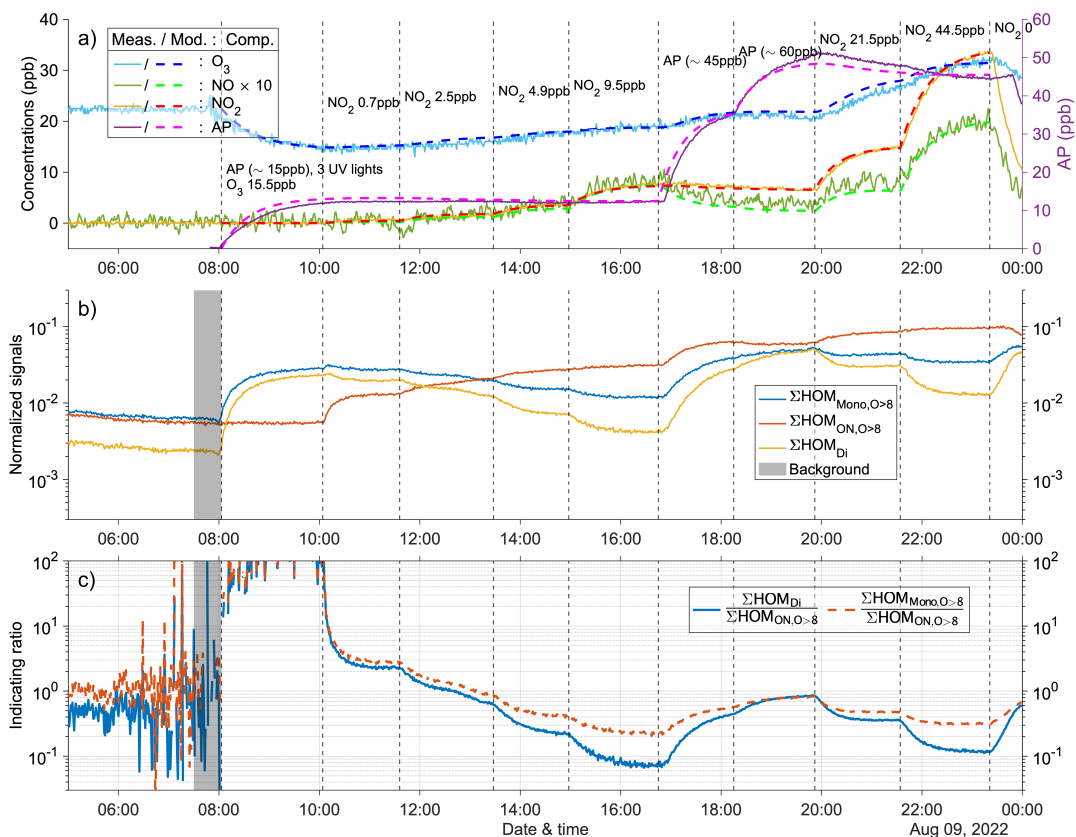


Figure A9. Time series of experiment no. 4 with 15 ppb α -pinene and 15.5 ppb O_3 as initial inputs. 3 UV lights were on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O_3 , NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma HOM_{Mono, O>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma HOM_{ON, O>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and ΣHOM_{Di} (sum of HOM dimers); c) IR1 ($\frac{\Sigma HOM_{Di}}{\Sigma HOM_{ON, O>8}}$) and IR2 ($\frac{\Sigma HOM_{Mono, O>8}}{\Sigma HOM_{ON, O>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

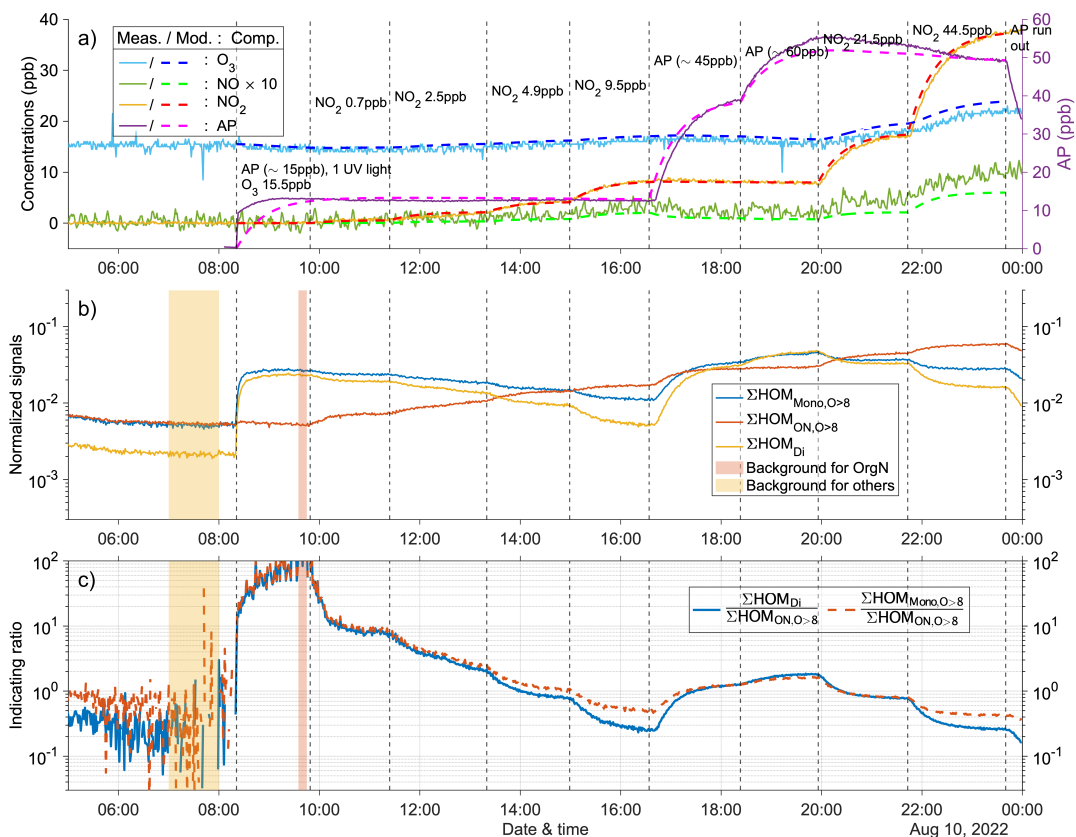


Figure A10. Time series of experiment no. 5 with 15 ppb α -pinene and 15.5 ppb O_3 as initial inputs. 1 UV light was on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O_3 , NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma HOM_{Mono,O>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma HOM_{ON,O>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and ΣHOM_{Di} (sum of HOM dimers); c) $IR1$ ($\frac{\Sigma HOM_{Di}}{\Sigma HOM_{ON,O>8}}$) and $IR2$ ($\frac{\Sigma HOM_{Mono,O>8}}{\Sigma HOM_{ON,O>8}}$). The yellow and red shaded area represent the time periods selected for background subtraction of dimers and organic nitrates, before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

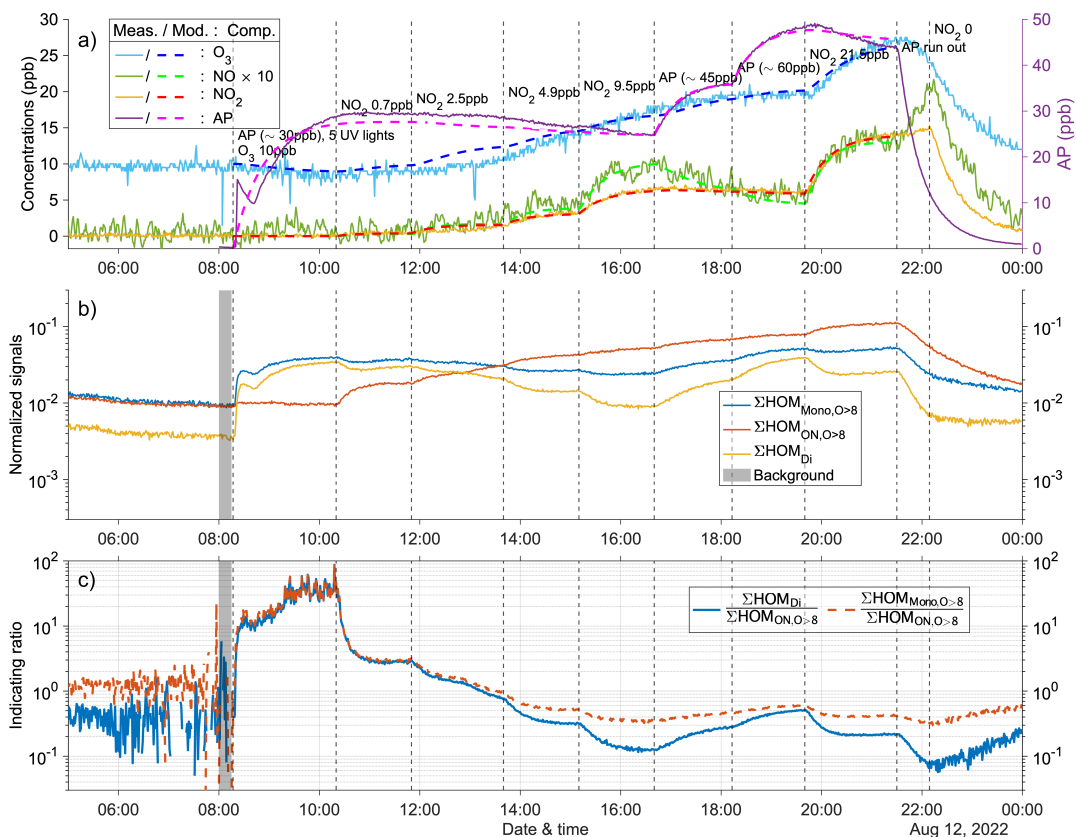


Figure A11. Time series of experiment no. 7 with 30 ppb α -pinene and 10 ppb O_3 as initial inputs. 5 UV lights was on during all stages. Three subplots show the time series of different compounds: a) measured (abbreviated as Meas., in solid lines) and modeled (abbreviated as Mod., in dashed lines) concentrations of O_3 , NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma HOM_{Mono,O>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma HOM_{ON,O>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and ΣHOM_{Di} (sum of HOM dimers); c) IR1 ($\frac{\Sigma HOM_{Di}}{\Sigma HOM_{ON,O>8}}$) and IR2 ($\frac{\Sigma HOM_{Mono,O>8}}{\Sigma HOM_{ON,O>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

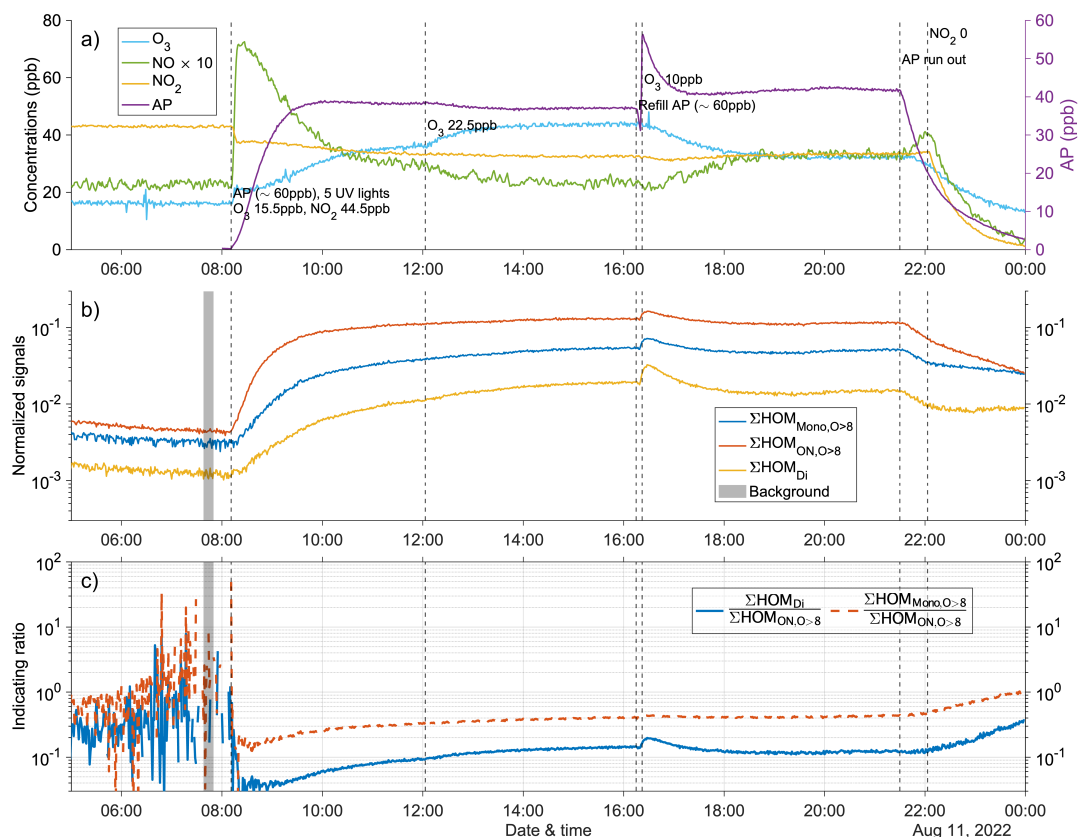


Figure A12. Time series of experiment no. 6, which was meant for collecting more data at the highest NO₂ input (at ~ 44.5 ppb). 5 UV lights was on during all stages. Three subplots show the time series of different compounds: a) measured concentrations of O₃, NO_x (both shown by left y-axis; the NO concentration is multiplied by 10), and AP (i.e., α -pinene, shown by right y-axis); b) normalized signals of $\Sigma\text{HOM}_{\text{Mono},\text{O}>8}$ (sum of non-nitrate HOM monomers with more than 8 oxygen atoms), $\Sigma\text{HOM}_{\text{ON},\text{O}>8}$ (sum of HOM organic nitrates with more than 8 oxygen atoms), and $\Sigma\text{HOM}_{\text{Di}}$ (sum of HOM dimers); c) IR1 ($\frac{\Sigma\text{HOM}_{\text{Di}}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$) and IR2 ($\frac{\Sigma\text{HOM}_{\text{Mono},\text{O}>8}}{\Sigma\text{HOM}_{\text{ON},\text{O}>8}}$). The grey shaded area represents the time period selected for background subtraction before calculating the ratio. Dashed vertical lines indicate specific time points of operations, with the corresponding labels for each operation in the subplot a.

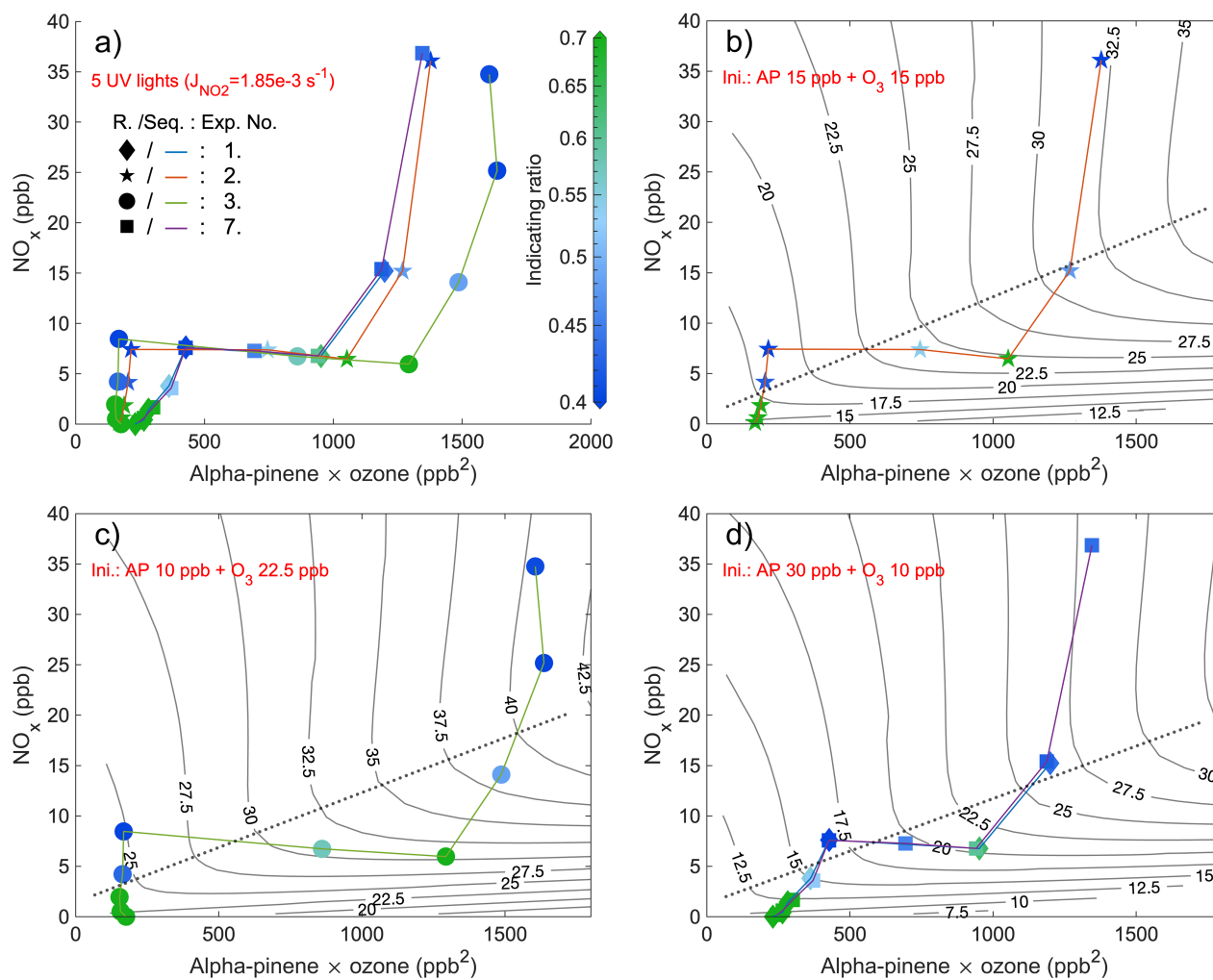


Figure A13. Steady state IR2 ($\frac{\sum \text{HOM}_{\text{Mono}, \text{O} > 8}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) of experiments from four days with 5 UV lights. X-axis is the multiplication of steady state α -pinene and O₃ concentrations, while y-axis is the steady state NO_x . The scatter points (exp. no. 1 (experiment number 1): diamond; exp. no. 2: star; exp. no. 3: round; exp. no. 7: square) are colored by values of IR2 (abbreviated as R. in the figure), and are connected by curves (exp. no. 1: blue; exp. no. 2: orange; exp. no. 3: green; exp. no. 7: purple) showing the sequence (Seq.) of experimental stages. Panel a combines stages of all four days, and the rest three subplots respectively show the stages of experiments with different initial (Ini.) inputs (exp. no. 1 and 7 are in the same panel d due to same initial inputs). EKMA curves (isopleth of O₃ concentrations in ppb), simulated by the box model, are black solid lines, while dotted lines are corresponding ridge lines.

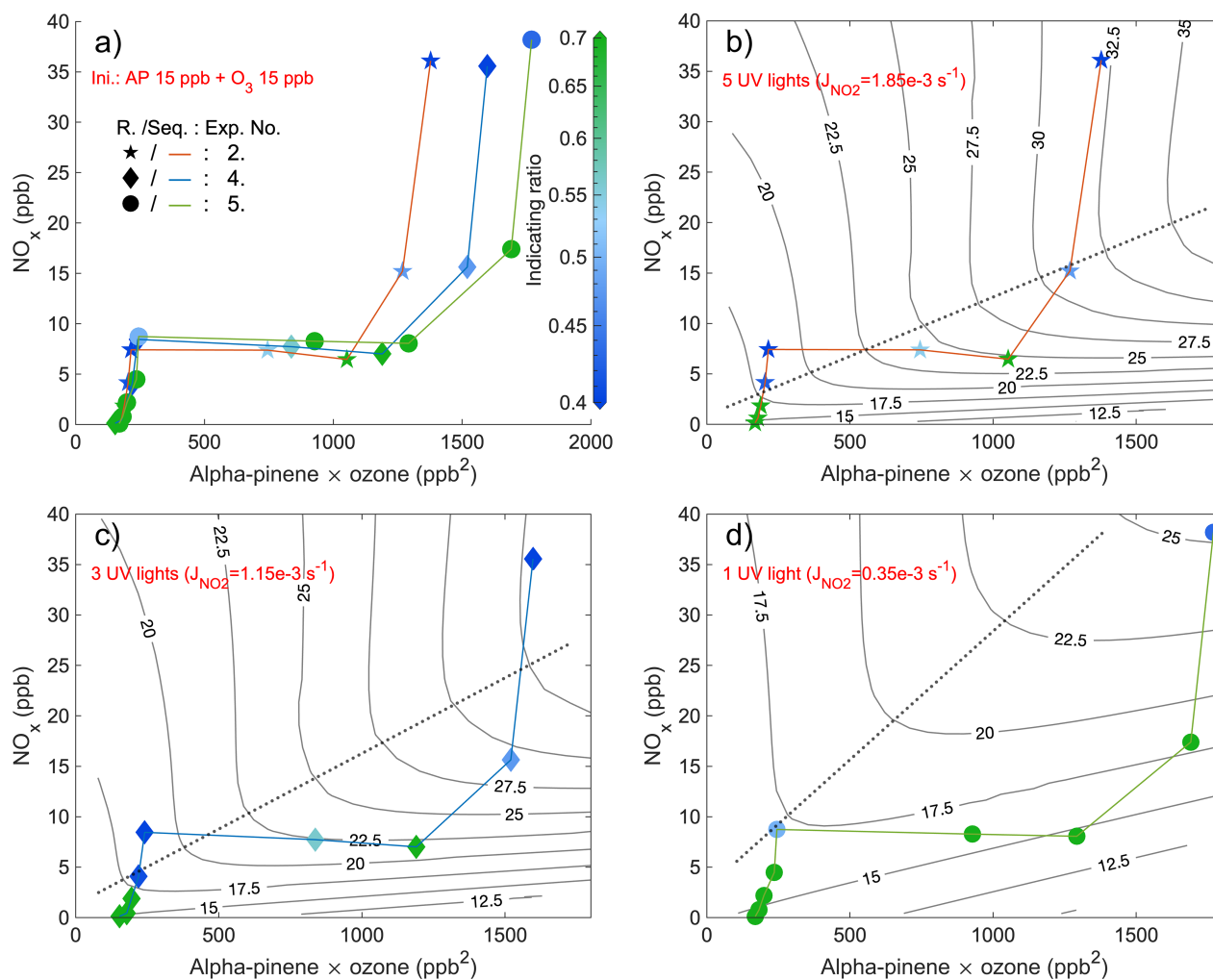


Figure A14. Steady state IR2 ($\frac{\sum \text{HOM}_{\text{Mono}, \text{O} > 8}}{\sum \text{HOM}_{\text{ON}, \text{O} > 8}}$) of experiments from three days (with 5, 3, and 1 UV lights, respectively). X-axis is the multiplication of steady state α -pinene and O₃ concentrations, while y-axis is the steady state NO_x. The scatter points (exp. no. 2 (experiment number 2): star; exp. no. 4: diamond; exp. no. 5: round) are colored by values of IR2 (abbreviated as R. in the figure), and are connected by curves (exp. no. 2: orange; exp. no. 4: blue; exp. no. 5: green) showing the sequence (Seq.) of experimental stages. Panel a combines stages of all three days with the same initial (Ini.) inputs, and the rest three subplots respectively show the stages of experiments with different amount of UV lights. EKMA curves by the box model are in black lines and the dotted lines are corresponding ridge lines.

Table A1. Zero-VOC experiment conditions. Experiment numbers (No.) and number of total stages are shown in the first two columns. Input information include the number and NO_2 photolysis rates (J_{NO_2}) of UV lights ($\lambda \approx 400$ nm) and concentrations of O_3 , α -pinene, and NO_x .

Experiment No.	Number of stages	Input			
		Number of lights	J_{NO_2} (s^{-1})	O_3 (ppb)	NO_x range (ppb)
Z1./Z2./Z3.	9/7/7	5	1.85×10^{-3}	10/15.5/22.5	0.7 – 44.5
Z4.	8	3	1.15×10^{-3}	15	0.7 – 44.5
Z5.	7	1	0.35×10^{-3}	15.5	0.7 – 44.5

From steady-state (ss in subscript) balance of the O_3 concentration, we can write the following expression:

$$\frac{d[\text{O}_3]}{dt} = J_{\text{NO}_2}[\text{NO}_2]_{ss} + \frac{[\text{O}_3]_{input} - [\text{O}_3]_{ss}}{\tau} - k_{\text{O}_3, \text{NO}}[\text{O}_3]_{ss}[\text{NO}]_{ss} = 0$$

where $\tau = \frac{2000 \text{ L}}{55 \text{ L min}^{-1}}$ is the residence time. O_3 has NO_2 photolysis ($J_{\text{NO}_2}[\text{NO}_2]_{ss}$) and its input ($\frac{[\text{O}_3]_{input}}{\tau}$) as sources, with reaction to NO ($-k_{\text{O}_3, \text{NO}}[\text{O}_3]_{ss}[\text{NO}]_{ss}$) and flush-out ($-\frac{[\text{O}_3]_{ss}}{\tau}$) as sinks. We can solve the equation to get NO_2 photolysis rate:

$$J_{\text{NO}_2} = \frac{k_{\text{O}_3, \text{NO}}[\text{O}_3]_{ss}[\text{NO}]_{ss} + \frac{[\text{O}_3]_{ss} - [\text{O}_3]_{input}}{\tau}}{[\text{NO}_2]_{ss}}$$

This expression can be used for each steady state to estimate J_{NO_2} in the corresponding experiment.

Table A2. Reactions and their reaction rate coefficients used for the box model. Note that RO₂ represent all kinds of peroxy radicals, thus there are huge uncertainties regarding reaction rates. This model is only meant for simulating concentrations of O₃ and its precursors.

Reactions	Reaction rate coefficients ^a
1. $NO_2 + hv \xrightarrow{O_2} NO + O_3$	$0.35/1.15/1.85 \times 10^{-3}$
2. $O_3 + NO \rightarrow NO_2 + O_2$	1.8×10^{-14}
3. $\alpha\text{-pinene} + O_3 \xrightarrow{O_2} RO_2 + OH$	8.7×10^{-17}
4. $\alpha\text{-pinene} + OH \xrightarrow{O_2} RO_2$	5.5×10^{-11}
5. $RO_2 + NO \rightarrow RO + NO_2$	1×10^{-11}
6. $RO_2 + RO_2 \rightarrow 2RO + O_2$	8×10^{-13}
7. $RO + O_2 \rightarrow HO_2$	Instantaneous
8. $NO + OH \rightarrow HNO_2$	4×10^{-11}
9. $NO + OH \xrightarrow{O_2} HO_2 + NO_2$	1.2×10^{-13}
10. $NO_2 + OH \rightarrow HNO_3$	4×10^{-11}
11. $HO_2 + NO \rightarrow OH + NO_2$	1×10^{-11}
12. $HO_2 + NO_2 \rightarrow HO_2NO_2$	3×10^{-14}
13. $HO_2 + HO_2/RO_2 \rightarrow \text{peroxides}$	2×10^{-11}
14. $RO_2 \rightarrow \text{wall}$ ^b	2.5×10^{-3}

^a Except the NO₂ photolysis and wall loss rates (s⁻¹), all the other reaction rates (cm³·s⁻¹) are adapted from NIST (National Institute of Standards and Technology) Chemical Kinetics Database (more information see <https://kinetics.nist.gov/kinetics/index.jsp>)

^b In our model, we only considered RO₂ with a wall loss lifetime of 400 s (Peräkylä et al., 2020).

Author contributions. ME, JZ, and JYZ designed the study. JZ, JYZ, YL, and VM conducted the experiments. JYZ analyzed the data and developed the model. ME, JZ, and DR supported the data analysis.

385 *Competing interests.* Douglas Worsnop works for Aerodyne Research, Inc.

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390 References

- Archer, S. L., Shultz, P. J., Warren, J. B., Hampl, V., and DeMaster, E. G.: Preparation of standards and measurement of nitric oxide, nitroxy, and related oxidation products, *Methods*, 7, 21–34, <https://doi.org/10.1006/meth.1995.1004>, 1995.
- Atkinson, R. and Arey, J.: Atmospheric degradation of volatile organic compounds, *Chemical Reviews*, 103, 4605–4638, <https://doi.org/10.1021/cr0206420>, 2003.
- 395 Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M. E., Rossi, M. J., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume I - gas phase reactions of O_x, HO_x, NO_x and SO_x species, *Atmospheric Chemistry and Physics*, 4, 1461–1738, <https://doi.org/10.5194/acp-4-1461-2004>, 2004.
- Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly oxygenated
400 organic molecules (HOM) from gas-phase autoxidation involving peroxy radicals: A key contributor to atmospheric aerosol, *Chemical Reviews*, 119, 3472–3509, <https://doi.org/10.1021/acs.chemrev.8b00395>, 2019.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M., Kondo, Y., Liao, H., and Lohmann, U.: Clouds and aerosols, in: *Climate change 2013: The physical science basis. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change*, pp. 571–657, Cambridge University Press, 2013.
- 405 Chameides, W. L., Fehsenfeld, F., Rodgers, M. O., Cardelino, C., Martinez, J., Parrish, D., Lonneman, W., Lawson, D. R., Rasmussen, R. A., Zimmerman, P., Greenberg, J., Middleton, P., and Wang, T.: Ozone precursor relationships in the ambient atmosphere, *Journal of Geophysical Research: Atmospheres*, 97, 6037–6055, <https://doi.org/10.1029/91JD03014>, 1992.
- Chapman, S.: A theory of upper-atmospheric ozone, *Memoirs of the Royal Meteorological Society*, 3, 103–125, <https://cir.nii.ac.jp/crid/1572261549196038528>, 1930.
- 410 Chipperfield, M. P., Bekki, S., Dhomse, S., Harris, N. R. P., Hassler, B., Hossaini, R., Steinbrecht, W., Thiéblemont, R., and Weber, M.: Detecting recovery of the stratospheric ozone layer, *Nature*, 549, 211–218, <https://doi.org/10.1038/nature23681>, 2017.
- Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of organic compounds in the atmosphere, *The Journal of Physical Chemistry Letters*, 4, 3513–3520, <https://doi.org/10.1021/jz4019207>, 2013.
- Crutzen, P. J.: The influence of nitrogen oxides on the atmospheric ozone content, *Quarterly Journal of the Royal Meteorological Society*,
415 96, 320–325, <https://doi.org/10.1002/qj.49709640815>, 1970.
- Crutzen, P. J.: Ozone production rates in an oxygen-hydrogen-nitrogen oxide atmosphere, *Journal of Geophysical Research*, 76, 7311–7327, <https://doi.org/10.1029/JC076i030p07311>, 1971.
- Dickerson, R. R., Kondragunta, S., Stenchikov, G., Civerolo, K. L., Doddridge, B. G., and Holben, B. N.: The impact of aerosols on solar ultraviolet radiation and photochemical smog, *Science*, 278, 827–830, <https://doi.org/10.1126/science.278.5339.827>, 1997.
- 420 Dodge, M. C.: Combined use of modeling techniques and smog chamber data to derive ozone-precursor relationships, in: *International conference on photochemical oxidant pollution and its control: Proceedings*, vol. 2, pp. 881–889, US Environmental Protection Agency, 1977.
- Ehhalt, D., Prather, M., Dentener, F., Derwent, R., Dlugokencky, E., Holland, E., Isaksen, I., Katima, J., Kirchhoff, V., Matson, P., Midgley, P., Wang, M., Berntsen, T., Bey, I., Brasseur, G., Buja, L., Collins, W. J. J., Daniel, J., Demore, W. B. B., Derek, N., Dickerson, R., Etheridge, D., Feichter, J., Fraser, P., Friedl, R., Fuglestvedt, J., Gauss, M., Grenfell, L., Grüber, A., Harris, N., Hauglustaine, D., Horowitz, L.,
425 Jackman, C., Jacob, D., Jaeglé, L., Jain, A., Kanakidou, M., Karlsdottir, S., Ko, M., Kurylo, M., Lawrence, M., Logan, J. a. A., Manning,

- M., Mauzerall, D., McConnell, J., Mickley, L., Montzka, S., Müller, J. F. F., Olivier, J., Pickering, K., Pitari, G., Roelofs, G. J. J., Rogers, H., Rognerud, B., Smith, S., Solomon, S., Staehelin, J., Steele, P., Stevenson, D., Sundet, J., Thompson, A., Weele, M. v., Kuhlmann, R. v., Wang, Y., Weisenstein, D., Wigley, T., Wild, O., Wuebbles, D., Yantosca, R., Joos, F., and McFarland, M.: Atmospheric chemistry and greenhouse gases, <https://hal.science/hal-03333922>, 2001.
- 430 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, *Nature*, 506, 476–479, <https://doi.org/10.1038/nature13032>, 2014.
- 435 Graus, M., Müller, M., and Hansel, A.: High resolution PTR-TOF: Quantification and formula confirmation of VOC in real time, *Journal of the American Society for Mass Spectrometry*, 21, 1037–1044, <https://doi.org/10.1016/j.jasms.2010.02.006>, 2010.
- Gruijl, F. d. and Leun, J.: Environment and health: 3. Ozone depletion and ultraviolet radiation, *CMAJ*, 163, 851–855, <https://www.cmaj.ca/content/163/7/851>, 2000.
- 440 Haagen-Smit, A. J., Bradley, C. E., and Fox, M. M.: Ozone formation in photochemical oxidation of organic substances, *Industrial & Engineering Chemistry*, 45, 2086–2089, <https://doi.org/10.1021/ie50525a044>, 1953.
- Hammer, M.-U., Vogel, B., and Vogel, H.: Findings on H₂O₂/HNO₃ as an indicator of ozone sensitivity in Baden-Württemberg, Berlin-Brandenburg, and the Po valley based on numerical simulations, *Journal of Geophysical Research Atmospheres*, 107, <https://doi.org/10.1029/2000JD000211>, 2002.
- 445 Hansel, A., Jordan, A., Holzinger, R., Prazeller, P., Vogel, W., and Lindinger, W.: Proton transfer reaction mass spectrometry: on-line trace gas analysis at the ppb level, *International Journal of Mass Spectrometry and Ion Processes*, 149–150, 609–619, [https://doi.org/10.1016/0168-1176\(95\)04294-U](https://doi.org/10.1016/0168-1176(95)04294-U), 1995.
- Harris, G. W., Carter, W. P. L., Winer, A. M., Pitts, J. N., Platt, U., and Perner, D.: Observations of nitrous acid in the Los Angeles atmosphere and implications for predictions of ozone-precursor relationships, *Environmental Science & Technology*, 16, 414–419, <https://doi.org/10.1021/es00101a009>, 1982.
- 450 He, X.-C., Shen, J., Iyer, S., Juuti, P., Zhang, J., Koirala, M., Kytökari, M. M., Worsnop, D. R., Rissanen, M., Kulmala, M., Maier, N. M., Mikkilä, J., Sipilä, M., and Kangasluoma, J.: Characterisation of gaseous iodine species detection using the multi-scheme chemical ionisation inlet 2 with bromide and nitrate chemical ionisation methods, *Atmospheric Measurement Techniques*, 16, 4461–4487, <https://doi.org/10.5194/amt-16-4461-2023>, 2023.
- 455 Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin, R. L. I., Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-API-TOF, *Atmospheric Chemistry and Physics*, 12, 4117–4125, <https://doi.org/10.5194/acp-12-4117-2012>, 2012.
- Jordan, A., Haidacher, S., Hanel, G., Hartungen, E., Märk, L., Seehauser, H., Schottkowsky, R., Sulzer, P., and Märk, T. D.: A high resolution and high sensitivity proton-transfer-reaction time-of-flight mass spectrometer (PTR-TOF-MS), *International Journal of Mass Spectrometry*, 286, 122–128, <https://doi.org/10.1016/j.ijms.2009.07.005>, 2009.
- 460 Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer, K., Kulmala, M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition, *Atmospheric Measurement Techniques*, 3, 1039–1053, <https://doi.org/10.5194/amt-3-1039-2010>, 2010.

- Kelly, F. J. and Fussell, J. C.: Air pollution and public health: emerging hazards and improved understanding of risk, *Environmental Geochemistry and Health*, 37, 631–649, <https://doi.org/10.1007/s10653-015-9720-1>, 2015.
- 465 Kinoshian, J. R.: Ozone-precursor relationships from EKMA diagrams, *Environmental Science & Technology*, 16, 880–883, <https://doi.org/10.1021/es00106a011>, 1982.
- Kleinman, L. I., Daum, P. H., Lee, Y.-N., Nunnermacker, L. J., Springston, S. R., Weinstein-Lloyd, J., and Rudolph, J.: Ozone production efficiency in an urban area, *Journal of Geophysical Research: Atmospheres*, 107, ACH 23–1, <https://doi.org/10.1029/2002JD002529>,
470 2002.
- Krechmer, J. E., Day, D. A., and Jimenez, J. L.: Always lost but never forgotten: Gas-phase wall losses are important in all teflon environmental chambers, *Environmental Science & Technology*, 54, 12 890–12 897, <https://doi.org/10.1021/acs.est.0c03381>, 2020.
- Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a chemical ionization mass spectrometer for the measurement of gaseous sulfuric acid, *The Journal of Physical Chemistry A*, 116, 6375–6386, <https://doi.org/10.1021/jp212123n>, 2012.
- 475 Kürten, A., Jokinen, T., Simon, M., Sipilä, M., Sarnela, N., Junninen, H., Adamov, A., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., Dommen, J., Donahue, N. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hakala, J., Hansel, A., Heinritzi, M., Hutterli, M., Kangasluoma, J., Kirkby, J., Laaksonen, A., Lehtipalo, K., Leiminger, M., Makhmutov, V., Mathot, S., Onnela, A., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L., Schobesberger, S., Seinfeld, J. H., Steiner, G., Tomé, A., Tröstl, J., Winkler, P. M., Williamson, C., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K. S., Kulmala, M., Worsnop, D. R., and Curtius, J.: Neutral molecular cluster formation of sulfuric acid–dimethylamine observed in real time under atmospheric conditions, *Proceedings of the National Academy of Sciences*, 111, 15 019–15 024, <https://doi.org/10.1073/pnas.1404853111>, 2014.
- 480 Lelieveld, J. and Dentener, F.: What controls tropospheric ozone?, *Journal of Geophysical Research: Atmospheres*, 105, <https://doi.org/10.1029/1999JD901011>, 2000.
- Liu, C. and Shi, K.: A review on methodology in O₃-NO_x-VOC sensitivity study, *Environmental Pollution*, 291, 118 249, <https://doi.org/10.1016/j.envpol.2021.118249>, 2021.
- 485 Madronich, S., Hastie, D. R., Ridley, B. A., and Schiff, H. I.: Measurement of the photodissociation coefficient of NO₂ in the atmosphere. I - Method and surface measurements, *Journal of Atmospheric Chemistry*, 1, 3–25, <https://doi.org/10.1007/BF00113977>, 1983.
- Martin, R., Fiore, A., and Donkelaar, A.: Space-based diagnosis of surface ozone sensitivity to anthropogenic emissions, *Geophys. Res. Lett.*, 31, <https://doi.org/10.1029/2004GL019416>, 2004.
- 490 Melkonyan, A. and Kuttler, W.: Long-term analysis of NO, NO₂ and O₃ concentrations in North Rhine-Westphalia, Germany, *Atmospheric Environment*, 60, 316–326, <https://doi.org/10.1016/j.atmosenv.2012.06.048>, 2012.
- Meyer Jr, E. L., Summerhays, J. P., and Freas, W.: Uses, limitations, and technical basis of procedures for quantifying relationships between photochemical oxidants and precursors, in: EPA-450/2-77-021a, Environmental Protection Agency Research Triangle Park, NC, 1977.
- Nuvolone, D., Petri, D., and Voller, F.: The effects of ozone on human health, *Environmental Science and Pollution Research*, 25, 8074–8088, <https://doi.org/10.1007/s11356-017-9239-3>, 2018.
- 495 Pathak, R. K., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of α -pinene at atmospherically relevant concentrations: Temperature dependence of aerosol mass fractions (yields), *Journal of Geophysical Research: Atmospheres*, 112, 2007.
- Peräkylä, O., Riva, M., Heikkinen, L., Quéléver, L., Roldin, P., and Ehn, M.: Experimental investigation into the volatilities of highly oxygenated organic molecules (HOMs), *Atmospheric Chemistry and Physics*, 20, 649–669, <https://doi.org/10.5194/acp-20-649-2020>,
500 2020.

- Riva, M., Heikkinen, L., Bell, D. M., Peräkylä, O., Zha, Q., Schallhart, S., Rissanen, M. P., Imre, D., Petäjä, T., Thornton, J. A., Zelenyuk, A., and Ehn, M.: Chemical transformations in monoterpene-derived organic aerosol enhanced by inorganic composition, *npj Climate and Atmospheric Science*, 2, 1–9, <https://doi.org/10.1038/s41612-018-0058-0>, 2019a.
- Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala, M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical ionization techniques for detecting gaseous oxygenated organic species, *Atmospheric Measurement Techniques*, 12, 2403–2421, <https://doi.org/10.5194/amt-12-2403-2019>, 2019b.
- Sandermann, H.: Ozone and plant health, *Annual Review of Phytopathology*, 34, 347–366, <https://doi.org/10.1146/annurev.phyto.34.1.347>, 1996.
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric chemistry and physics: From air pollution to climate change*, John Wiley & Sons, 2016.
- Seviour, W. J. M.: Good ozone, bad ozone and the Southern Ocean, *Nature Climate Change*, 12, 316–317, <https://doi.org/10.1038/s41558-022-01322-8>, 2022.
- Sicard, P., De Marco, A., Agathokleous, E., Feng, Z., Xu, X., Paoletti, E., Rodriguez, J. J. D., and Calatayud, V.: Amplified ozone pollution in cities during the COVID-19 lockdown, *Science of The Total Environment*, 735, 139 542, <https://doi.org/10.1016/j.scitotenv.2020.139542>, 2020.
- Sillman, S.: The relation between ozone, NO_x and hydrocarbons in urban and polluted rural environments, *Atmospheric Environment*, 33, 1821–1845, [https://doi.org/10.1016/S1352-2310\(98\)00345-8](https://doi.org/10.1016/S1352-2310(98)00345-8), 1999.
- Sillman, S., Logan, J. A., and Wofsy, S. C.: The sensitivity of ozone to nitrogen oxides and hydrocarbons in regional ozone episodes, *Journal of Geophysical Research: Atmospheres*, 95, 1837–1851, <https://doi.org/10.1029/JD095iD02p01837>, 1990.
- Stahelin, J., Harris, N. R. P., Appenzeller, C., and Eberhard, J.: Ozone trends: A review, *Reviews of Geophysics*, 39, 231–290, <https://doi.org/10.1029/1999RG000059>, 2001.
- Stolarski, R. S. and Cicerone, R. J.: Stratospheric chlorine: a possible sink for ozone, *Canadian Journal of Chemistry*, 52, 1610–1615, <https://doi.org/10.1139/v74-233>, 1974.
- Tang, X., Li, J., and Chen, D.: Summertime photochemical pollution in Beijing, *LOCAL AIR POLLUTION IN FAST DEVELOPING COUNTRIES*, 1410, 1465, 1995.
- Tiao, G. C., Box, G. E. P., and Hamming, W. J.: Analysis of Los Angeles photochemical smog data: a statistical overview, *Journal of the Air Pollution Control Association*, 25, 260–268, <https://doi.org/10.1080/00022470.1975.10470082>, 1975.
- Trainer, M., Parrish, D. D., Buhr, M. P., Norton, R. B., Fehsenfeld, F. C., Anlauf, K. G., Bottenheim, J. W., Tang, Y. Z., Wiebe, H. A., Roberts, J. M., Tanner, R. L., Newman, L., Bowersox, V. C., Meagher, J. F., Olszyna, K. J., Rodgers, M. O., Wang, T., Berresheim, H., Demerjian, K. L., and Roychowdhury, U. K.: Correlation of ozone with NO_y in photochemically aged air, *Journal of Geophysical Research: Atmospheres*, 98, 2917–2925, <https://doi.org/10.1029/92JD01910>, 1993.
- Wang, T., Ding, A., Gao, J., and Wu, W. S.: Strong ozone production in urban plumes from Beijing, China, *Geophysical Research Letters*, 33, <https://doi.org/10.1029/2006GL027689>, 2006.
- Wang, T., Xue, L., Brimblecombe, P., Lam, Y. F., Li, L., and Zhang, L.: Ozone pollution in China: A review of concentrations, meteorological influences, chemical precursors, and effects, *Science of The Total Environment*, 575, 1582–1596, <https://doi.org/10.1016/j.scitotenv.2016.10.081>, 2017.
- Yan, C., Nie, W., Äijälä, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H., Jokinen, T., Sarnela, N., Häme, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prévôt, A. S. H., Petäjä, T., Kulmala, M., Sipilä, M., Worsnop, D. R., and Ehn, M.: Source char-

- acterization of highly oxidized multifunctional compounds in a boreal forest environment using positive matrix factorization, *Atmospheric Chemistry and Physics*, 16, 12715–12731, <https://doi.org/10.5194/acp-16-12715-2016>, 2016.
- 540 Yan, C., Nie, W., Vogel, A. L., Dada, L., Lehtipalo, K., Stolzenburg, D., Wagner, R., Rissanen, M. P., Xiao, M., Ahonen, L., Fischer, L., Rose, C., Bianchi, F., Gordon, H., Simon, M., Heinritzi, M., Garmash, O., Roldin, P., Dias, A., Ye, P., Hofbauer, V., Amorim, A., Bauer, P. S., Bergen, A., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Buchholz, A., Mazon, S. B., Canagaratna, M. R., Chen, X., Ding, A., Dommen, J., Draper, D. C., Duplissy, J., Frege, C., Heyn, C., Guida, R., Hakala, J., Heikkinen, L., Hoyle, C. R., Jokinen, T., Kangasluoma, J., Kirkby, J., Kontkanen, J., Kürten, A., Lawler, M. J., Mai, H., Mathot, S., Mauldin, R. L., Molteni, U., Nichman, L.,
- 545 Nieminen, T., Nowak, J., Ojdanic, A., Onnela, A., Pajunoja, A., Petäjä, T., Piel, F., Quéléver, L. L. J., Sarnela, N., Schallhart, S., Sengupta, K., Sipilä, M., Tomé, A., Tröstl, J., Väisänen, O., Wagner, A. C., Ylisirniö, A., Zha, Q., Baltensperger, U., Carslaw, K. S., Curtius, J., Flagan, R. C., Hansel, A., Riipinen, I., Smith, J. N., Virtanen, A., Winkler, P. M., Donahue, N. M., Kerminen, V.-M., Kulmala, M., Ehn, M., and Worsnop, D. R.: Size-dependent influence of NO_x on the growth rates of organic aerosol particles, *Science Advances*, 6, eaay4945, <https://doi.org/10.1126/sciadv.aay4945>, 2020.
- 550 Zhao, J., Häkkinen, E., Graeffe, F., Krechmer, J. E., Canagaratna, M. R., Worsnop, D. R., Kangasluoma, J., and Ehn, M.: A combined gas- and particle-phase analysis of highly oxygenated organic molecules (HOMs) from α -pinene ozonolysis, *Atmospheric Chemistry and Physics*, 23, 3707–3730, <https://doi.org/10.5194/acp-23-3707-2023>, 2023a.
- Zhao, J., Mickwitz, V., Luo, Y., Häkkinen, E., Graeffe, F., Zhang, J., Timonen, H., Canagaratna, M., Krechmer, J. E., Zhang, Q., Kulmala, M., Kangasluoma, J., Worsnop, D., and Ehn, M.: Characterization of the vaporization inlet for aerosols (VIA) for online measurements of
- 555 particulate highly oxygenated organic molecules (HOMs), *EGUsphere*, pp. 1–21, <https://doi.org/10.5194/egusphere-2023-1146>, 2023b.